



1	Comparing modified substrate induced respiration with selective inhibition
2	(SIRIN) and N <sub>2</sub> O isotope approaches to estimate fungal contribution to
3	denitrification in three arable soils under anoxic conditions
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18	Keywords: selective growth inhibition, $^{15}N$ site preference, fungal denitrification, $C_2H_2$ , isotope
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20	
21	Abstract
22	Pure culture studies provide evidence of the ability of soil fungi to produce nitrous oxide $\left(N_2O\right)$ during
23	denitrification. Soil studies with selective inhibition indicated a possible dominance of fungal
24	compared to bacterial $N_2O$ production in soil, which drew more attention to fungal denitrification.
25	Analyzing the isotopic composition of $N_2O,$ especially the $^{15}\!N$ site preference of $N_2O$ produced
26	$(SP_{N2O})$ , showed that N <sub>2</sub> O of pure bacterial or fungal cultures differed in $SP_{N2O}$ values, which might
27	enable the quantification of fungal $N_2O$ based on the isotopic endmember signatures of $N_2O$ produced
28	by fungi and bacteria.
29	This study aimed to identify the fungal contribution to $N_2 O$ emissions under anaerobic conditions in
30	incubated repacked soil samples by using different approaches to disentangle sources of $N_2 O. \label{eq:N2}$ Three
31	soils were incubated under anaerobic conditions to promote denitrification with four treatments of a
32	modified substrate induced respiration with selective inhibition (SIRIN) approach. While one
33	treatment without microbial inhibition served as a control the other three treatments were amended

with inhibitors to selectively inhibit bacterial, fungal or bacterial and fungal growth. These treatments

were performed in three varieties. In one variety the <sup>15</sup>N tracer technique was used to estimate the 

effect of N<sub>2</sub>O reduction on N<sub>2</sub>O produced, while two other varieties were performed under natural isotopic conditions but with and without acetylene. Three approaches were established to estimate the 

N2O production by a fungal community in soil: i) A modification of the SIRIN approach was used to





calculate N<sub>2</sub>O evolved from selected organism groups, and ii)  $SP_{N2O}$  values from the acetylated treatment were used in the isotope endmember mixing approach (IEM), and iii) the SP/ $\delta^{18}$ O mapping approach (SP/ $\delta^{18}$ O Map) was used to estimate the fungal contribution to N<sub>2</sub>O production and N<sub>2</sub>O reduction under anaerobic conditions from the non-acetylated treatment.

43 The three approaches tested revealed a small fungal contribution to  $N_2O$  fluxes under anaerobic 44 conditions in the soils tested. Quantifying the fungal fraction with modified SIRIN was only possible

45 in one soil and totaled 0.28 $\pm$ 0.09. This was higher than the results obtained by IEM and SP/ $\delta^{18}$ O Map,

 $46 \qquad \text{which accounted zero to } 0.20 \text{ of } N_2 \text{O} \text{ produced to the fungal community.} \\$ 

To our knowledge, this study was the first attempt to quantify the fungal contribution to anaerobic N<sub>2</sub>O production by simultaneous application of three approaches, i.e. modified SIRIN, IEM and SP/ $\delta^{18}$ O Map. While all methods coincided by suggesting a small or missing fungal contribution, further studies under conditions ensuring larger fungal N<sub>2</sub>O fluxes and including alternative inhibitors are needed to better cross-validate the methods.

# 52 **1. Introduction**

The greenhouse gas nitrous oxide (N<sub>2</sub>O) contributes to global warming and to the depletion of the ozone layer in the stratosphere (Crutzen, 1970; IPCC, 2013). The largest anthropogenic N<sub>2</sub>O emissions originate from agricultural soils and are mainly produced during microbial nitrification, nitrifier denitrification and denitrification (Firestone and Davidson, 1989; Bremner, 1997; IPCC, 2013; Wrage-Mönnig et al., 2018). In order to find mitigation strategies for N<sub>2</sub>O emissions from arable soils, it is important to understand N<sub>2</sub>O sources and sinks and thus improve knowledge about the production pathways and the microorganisms involved.

60 For a long time, it was believed that solely bacteria are involved in N<sub>2</sub>O formation during 61 denitrification (Firestone and Davidson, 1989); however, also several fungi are capable of 62 denitrification (Bollag and Tung, 1972; Shoun et al., 1992). Denitrification describes the reduction of 63 nitrate (NO<sub>3</sub>) to dinitrogen (N<sub>2</sub>), with the intermediates nitrite (NO<sub>2</sub>), nitric oxide (NO) and N<sub>2</sub>O (Knowles, 1982). While this entire reaction chain including the ability to reduce  $N_2O$  to  $N_2$  is found 64 65 among bacterial denitrifiers, most fungi lack N<sub>2</sub>O reductase (Nos) (Shoun et al., 1992; Shoun et al., 66 2012; Higgins et al., 2018). Recently, pure culture studies showed that  $N_2O$  from fungal denitrification 67 was often accompanied with N<sub>2</sub>O from abiotic production (Phillips et al., 2016a; Phillips et al., 68 2016b), which may lead to overestimate the importance of fungal N<sub>2</sub>O production. Other studies 69 indicated that only some fungal species (e.g. Fusarium strains) performing respiratory denitrification 70 with substantial amounts of N<sub>2</sub>O production (Higgins et al., 2018; Keuschnig et al., 2020). Even 71 though only a few fungal species were identified to be capable of respiratory denitrification, N<sub>2</sub>O 72 produced by fungi may contribute largely to N2O from denitrification in soil. Firstly, fungi dominate 73 the biomass in soil (up to 96%) compared to bacteria in general and thus fungi could potentially play a 74 dominant role in N<sub>2</sub>O production (Ruzicka et al., 2000; Braker and Conrad, 2011). Thus, a respiratory





75 fungal-to-bacterial (F:B) ratio of 4 is typical for arable soils (Anderson and Domsch, 1975; 76 Blagodatskaya and Anderson, 1998). Secondly, the fact that N<sub>2</sub>O is the major end product of fungal 77 denitrification led to the assumption that the potential activity of fungal N<sub>2</sub>O production in soil may 78 exceed that of bacteria, provided that both microbial groups have the same specific denitrification 79 activity (Shoun et al., 1992; Sutka et al., 2008). Thirdly, co-denitrification was found to often co-occur 80 with fungal denitrification (Shoun and Tanimoto, 1991; Tanimoto et al., 1992). During this fungal 81 pathway, a hybrid N<sub>2</sub>O is formed using one N atom from NO<sub>2</sub><sup>-</sup> and one N atom from compounds like 82 azide or ammonium (NH<sub>4</sub><sup>+</sup>) for N<sub>2</sub>O production (Tanimoto et al., 1992; Shoun et al., 1992; Rohe et al., 2017; Spott et al., 2011). A <sup>15</sup>N tracing approach was used to identify and quantify co-denitrification, 83 84 which contributed about 92% to N<sub>2</sub>O produced in an incubation experiment with a grassland soil under 85 anaerobic conditions (Laughlin and Stevens, 2002). This again stresses the large potential N<sub>2</sub>O 86 production by fungi. However, in pure culture studies not only co-denitrification, but also abiotic N<sub>2</sub>O 87 formation may co-occur with fungal denitrification (Phillips et al., 2016a; Phillips et al., 2016b; Rohe 88 et al., 2017) and pathway differentiation is still challenging.

89 Soil incubation experiments could serve to differentiate between N<sub>2</sub>O produced by fungi and bacteria 90 during denitrification by the application of two antibiotics: streptomycin and cycloheximide, which 91 inhibit bacterial or fungal growth, respectively, by inhibition of the protein biosynthesis. This method 92 is known as substrate induced respiration with selective inhibition (SIRIN) (Anderson and Domsch, 93 1975; Laughlin and Stevens, 2002; Crenshaw et al., 2008; Blagodatskaya et al., 2010; Long et al., 94 2013). A few studies used a modification of this method for N<sub>2</sub>O analysis and found a greater decrease 95 of N<sub>2</sub>O production with fungal than with bacterial growth inhibition (e.g. 89 vs. 23% decrease 96 (Laughlin and Stevens, 2002)), indicating that fungi might dominate N<sub>2</sub>O production (Laughlin and 97 Stevens, 2002; McLain and Martens, 2006; Crenshaw et al., 2008; Blagodatskaya et al., 2010; Long et 98 al., 2013; Chen et al., 2014; Chen et al., 2015).

99 Analysing the isotopic composition of N<sub>2</sub>O might be a promising tool to distinguish between N<sub>2</sub>O 100 from bacterial and fungal denitrification and other pathways. Especially, the isotopomer ratios of N<sub>2</sub>O (i.e.  $N_2O$  molecules with the same bulk <sup>15</sup>N isotopic enrichment but showing different positions of <sup>15</sup>N 101 102 in the linear N<sub>2</sub>O molecule (Ostrom and Ostrom, 2017)) in pure culture studies showed differences in 103 N<sub>2</sub>O of bacterial and fungal denitrification (Sutka et al., 2006; Sutka et al., 2008; Frame and Casciotti, 104 2010; Rohe et al., 2014a; Rohe et al., 2017) and might be suitable for distinguishing between N<sub>2</sub>O produced by bacteria or fungi under denitrifying conditions. Isotopomer ratios of N2O can be 105 106 expressed as <sup>15</sup>N site preference (SP<sub>N2O</sub>), i.e. the difference between  $\delta^{15}$ N of the central and terminal N-107 position of the asymmetric N<sub>2</sub>O molecule (Toyoda and Yoshida, 1999). The  $SP_{N2O}$  values of N<sub>2</sub>O of six 108 pure fungal cultures was between 16 and 37 ‰ (Sutka et al., 2008; Rohe et al., 2014a; Maeda et al., 109 2015; Rohe et al., 2017), whereas several bacteria produced N<sub>2</sub>O with  $SP_{N2O}$  values between -7.5 and 110 +3.5 ‰ during denitrification (Toyoda et al., 2005; Sutka et al., 2006; Rohe et al., 2017). However, the 111  $SP_{N2O}$  value of N<sub>2</sub>O produced by pure bacterial cultures during nitrification is approximately 33 ‰ and





112 interferes with SP<sub>N2O</sub> values of fungal denitrification (Sutka et al., 2006; Sutka et al., 2008; Rohe et al., 113 2014a). This demonstrates the difficulty to use  $SP_{N2O}$  values as an indicator for different organism 114 groups contributing to N<sub>2</sub>O production from soil, where different pathways may co-occur. Although  $SP_{N20}$  values are independent of isotopic signatures of the precursors,  $\delta^{15}N$  and  $\delta^{18}O$  values of 115 produced N<sub>2</sub>O ( $\delta^{15}$ N<sup>bulk</sup><sub>N2O</sub> and  $\delta^{18}$ O<sub>N2O</sub>, respectively) result from the isotopic signature of the 116 117 precursor and isotopic fractionation during N2O production (Toyoda et al., 2005; Frame and Casciotti, 118 2010). Interpretation of  $\delta^{18}O_{N2O}$  values is even more complex, because O exchange during 119 denitrification between water and denitrification intermediates alters the final  $\delta^{18}O_{N2O}$  value (Garber and Hollocher, 1982; Aerssens et al., 1986; Kool et al., 2007; Rohe et al., 2014b; Rohe et al., 2017). 120 121 However, recently fungal and bacterial N<sub>2</sub>O showed different ranges for  $\delta^{18}O_{N2O}$  values and this 122 isotopic signature may also be helpful in differentiation of these pathways (Lewicka-Szczebak et al., 123 2016). Moreover,  $\delta^{15}N^{\text{bulk}}_{N2O}$ ,  $\delta^{18}O_{N2O}$  and  $SP_{N2O}$  values are in the course of denitrification affected by isotopic fractionation due to N<sub>2</sub>O reduction. During N<sub>2</sub>O reduction, the <sup>14</sup>N<sup>16</sup>O bond is preferentially 124 broken compared to <sup>14</sup>N<sup>18</sup>O or <sup>15</sup>N<sup>16</sup>O, resulting in residual N<sub>2</sub>O, that is relatively isotopically enriched 125 126 in <sup>15</sup>N and <sup>18</sup>O and shows larger  $SP_{N2O}$  values compared to  $SP_{N2O}$  values of N<sub>2</sub>O from denitrification 127 without the reduction step (Popp et al., 2002; Ostrom et al., 2007). Quantification of N<sub>2</sub>O reduction to 128 N<sub>2</sub> during denitrification is possible by analyzing <sup>15</sup>N<sub>2</sub> fluxes in <sup>15</sup>N tracing experiments using <sup>15</sup>N 129 enriched substrates (Well et al., 2006; Lewicka-Szczebak et al., 2014). To quantify N2O reduction and 130 the pathways producing N<sub>2</sub>O based on N<sub>2</sub>O isotopocules (i.e. N<sub>2</sub>O with differing number or positions 131 of N or O isotopes (Ostrom and Ostrom, 2017)), the isotope mapping approach was developed using isotope fractionation factors together with  $\delta^{15}N^{\text{bulk}}$  values of N<sub>2</sub>O precursors ( $\delta^{15}N_{NOx}$ ) as well as 132  $\delta^{15}N^{\text{bulk}}_{N20}$  and  $SP_{N20}$  values of N<sub>2</sub>O produced (Toyoda et al., 2011). Recently, this isotope mapping 133 approach was further developed (SP/ $\delta^{18}$ O Map) using  $\delta^{18}O_{N2O}$  and  $SP_{N2O}$  values of N<sub>2</sub>O and  $\delta^{18}O$ 134 135 values of precursors (Lewicka-Szczebak et al., 2014; Lewicka-Szczebak et al., 2017). This approach uses different slopes of N<sub>2</sub>O reduction and mixing lines in the  $\delta^{18}O$  – SP isotope plot and allows for 136 137 differentiation of isotope effects due to N<sub>2</sub>O reduction and admixture of fungal N<sub>2</sub>O.

138 Based on the above cited ranges for the isotopomer endmembers of fungal and bacterial 139 denitrification, and assuming that only fungi and bacteria are responsible for N<sub>2</sub>O production the 140 fraction of fungal N<sub>2</sub>O can be calculated using the isotope endmember mixing approach (IEM) with  $SP_{N2O}$  values of N<sub>2</sub>O produced in soil ( $SP_{prod}$ ), provided N<sub>2</sub>O reduction, which is altering  $SP_{N2O}$  values 141 142 of emitted N<sub>2</sub>O, does not occur (Ostrom et al., 2010; Ostrom and Ostrom, 2011). This can be ensured 143 in laboratory experiments by inhibiting  $N_2O$  reduction to  $N_2$  using acetylene ( $C_2H_2$ ) during anaerobic incubation experiments (Yoshinari and Knowles, 1976; Groffman et al., 2006; Well and Flessa, 2009; 144 145 Nadeem et al., 2013). Hence,  $C_2H_2$  inhibition might be suitable to quantify  $SP_{prod}$  values in soils 146 exhibiting significant N<sub>2</sub>O reduction and would thus allow quantification of fungal N<sub>2</sub>O fluxes based on SP<sub>prod</sub> values. For the SP/δ<sup>18</sup>O Map, the inhibition of N<sub>2</sub>O reduction is not needed. Hence, N<sub>2</sub>O 147 148 reduction can be estimated together with the N<sub>2</sub>O mixing due to application of two isotopic signatures





149 of N<sub>2</sub>O. While it is generally assumed that SP<sub>prod</sub> values of N<sub>2</sub>O produced by fungal pure cultures is 150 transferable to N<sub>2</sub>O produced by fungal soil communities, this has not yet been proven. Until now, 151 studies reporting possible ranges of fungal contributions to  $N_2O$  fluxes from soil were based on  $SP_{prod}$ 152 values of pure cultures (Köster et al., 2013b; Zou et al., 2014; Lewicka-Szczebak et al., 2017; 153 Senbayram et al., 2018; Senbayram et al., 2020; Lewicka-Szczebak et al., 2014), but uncertainty of 154 this approach arose from the fact that the full range of  $SP_{prod}$  values is between 16 and 37% (Sutka et al., 2008; Maeda et al., 2015; Rohe et al., 2017). It would thus be useful to constrain fungal SP<sub>prod</sub> 155 156 values for a specific soil or soil type. 157 So far, the described methods for distinguishing between fungal and bacterial N<sub>2</sub>O emission have not 158 been evaluated and compared in the same soil and their accuracy and possible bias remains unknown. 159 Therefore, this study aims at (i) determining the fungal contribution on N<sub>2</sub>O production by denitrification under anoxic conditions and glucose addition using three arable soils and approaches: 160 modified SIRIN, IEM and the SP/818O Map, (ii) to compare the fungal contribution on N2O production 161 162 determined by these approaches and thus assess factors of potential bias of the methods, and (iii) to 163 estimate the  $SP_{N2O}$  values from a fungal soil community and thus to evaluate the transferability of the 164 pure culture range of the fungal  $SP_{N2O}$  endmember values.

### 165 **2.** Materials and Methods

#### 166 2.1 Soil samples

All experiments were conducted with three arable soils differing in texture to provide different conditions for denitrification. As one soil was sampled at two different time points, we conducted four experiments: Experiment 1 with loamy sand sampled in December 2012, Experiment 2 with sand sampled in January 2013, Experiment 3 with silt loam sampled in December 2012, and Experiment 4 with loamy sand sampled in June 2011.

172 Soil samples of the upper 30 cm were collected in plastic bags aerated via cotton wool stoppers and 173 stored at 6 °C for maximally two months. To get information about the initial soil status, total contents 174 of C and N in soil samples were analyzed by dry combustion of grinded samples (LECO TruSpec, 175 Germany). The soil pH was measured in 0.01 M CaCl<sub>2</sub>. The mineral nitrogen content (Nmin) of soil 176 samples was determined before and after fertilization by extracting NO3- and NH4+ with 0.01 M 177 calcium chloride dihydrate (CaCl<sub>2</sub>  $\cdot$  2 H<sub>2</sub>O) according to ISO 14255 and analyzing NO<sub>3</sub><sup>-</sup> and NH<sub>4</sub><sup>+</sup> concentrations in the extracts with a Continuous-Flow-Analyzer (SKALAR, Germany). The  $\delta^{15}$ N and 178 179  $\delta^{18}$ O values of NO<sub>3</sub><sup>-</sup> and NO<sub>2</sub><sup>-</sup> ( $\delta^{15}$ N<sub>NOx</sub> and  $\delta^{18}$ O<sub>NOx</sub>, respectively) in soil extracts (with 0.01 M calcium 180 chloride dihydrate (CaCl<sub>2</sub> · 2 H<sub>2</sub>O)) were analyzed by the bacterial denitrifier method (Casciotti et al., 181 2002). Respiratory biomass of the three soils was analyzed with substrate induced respiration (SIR) 182 according to Anderson and Domsch (1978) and the respiratory F:B ratio was analyzed with substrate 183 induced respiration with selective inhibition (SIRIN) in summer 2010 by a computer-generated 184 selectivity analysis: "SIR-SBA 4.00" (Heinemeyer, copyright MasCo Analytik, Hildesheim, Germany)





- 185 (Anderson and Domsch, 1975). The scheme of glucose and growth inhibitor combinations is listed
- 186 below in section "Methodological approach". For further characteristics of the soils, see Table 1.
- 187 2.2 Methodological approach
- 188 2.2.1 SIRIN pre-experiment

189 As in most studies applying the SIRIN method on N<sub>2</sub>O emissions (e. g. Laughlin and Stevens, 2002; 190 Chen et al., 2014; Ladan and Jacinthe, 2016), a pre-experiment was conducted, in order to get 191 information about optimal substrate and inhibitor concentrations for substrate induced growth 192 inhibition. The SIR method (Anderson and Domsch, 1978) was used to get information about the 193 amount of respiratory biomass in soil. To this end, we added different concentrations of glucose (0.50, 194 0.75, 1.0, 1.5, 2.0, 3.0, 4.0, 5.0, 6.0 mg g<sup>-1</sup> dry weight (dw) soil) to find the optimal glucose 195 concentration ( $c_{out}$ (glucose)), which is the glucose concentration that causes maximum initial 196 respiration rates (Anderson and Domsch, 1978). Copt(glucose)) was 1.0 mg g<sup>-1</sup> for Experiment 2 (sand) 197 and 1.5 mg g<sup>-1</sup> for Experiments 1, 3 and 4 (loamy sand and silt loam). Glucose served as substrate to initiate microbial growth (Anderson and Domsch, 1975). 198 199 We conducted SIRIN for determining the respiratory F:B ratio according to Anderson and Domsch

200 (1975). Selectivity of the inhibitor combinations of streptomycin (bacterial respiratory inhibitor) and cycloheximide (fungal respiratory inhibitor) were tested with the following concentrations, 0.75, 1.0, 201 202 1.5 mg g<sup>-1</sup> dw, respectively. The optimal concentration for inhibition of fungal respiration was 0.75 mg  $g^{-1}$  dw soil cycloheximide ( $c_{out}$ (cycloheximide)) and for bacterial respiratory inhibition 1.0 mg  $g^{-1}$  dw 203 204 soil streptomycin ( $c_{opt}$ (streptomycin)).

205 Table 1: Soil characteristics of three arable soils from Germany used for incubation experiments (Exp.) 206 (standard deviation in brackets).

Exp. (Year)	Soil texture	Soil type (WRB)	Location	C content [%]	N content [%]	NH4 <sup>+</sup> [mg N L <sup>-1</sup> ]	NO3 <sup>-</sup> [mg N L <sup>-1</sup> ]	pH (CaCl <sub>2</sub> )	δ <sup>15</sup> N <sub>NOx</sub> [‰] <sup>e</sup>	δ <sup>18</sup> O <sub>N</sub> <sup>Ox</sup> [‰] <sup>e</sup>	F:B <sup>f</sup>	Biomass <sup>g</sup> [µg C gdw <sup>-1</sup> soil]
1 (2012) 4 (2011)	Loamy sand	Haplic Luvisol	Braun- schweig <sup>a</sup>	1.43 (<0.01)	0.10 (<0.01)	0.04	1.25	5.67	3.98	-4.82	2.6	234
2 (2013)	Sand	Gleyic Podzol	Wenne- bostel <sup>b</sup>	2.31 (0.04)	0.14 (<0.01)	0.02	0.56	5.54	0.73	-2.68	2.6	161
3 (2013)	Silt loam	Haplic Luvisol	Götting- en <sup>c</sup>	1.62 (0.02)	0.13 (<0.01)	n.d. <sup>d</sup>	2.05	7.38	4.18	2.32	4.9	389

207 <sup>a</sup>Experimental Station of the Friedrich-Löffler Institute, Braunschweig, Germany

208 <sup>b</sup>private agricultural field North of Hannover, water protection area Fuhrberger Feld, Germany

<sup>c</sup>Reinshof Experimental Farm, Georg-August-University, Göttingen, Germany <sup>4</sup>not detectable (i.e. below detection limit of 0.005 mg L<sup>-1</sup> NH<sub>4</sub><sup>+</sup>-N) 209

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<sup>e</sup>Isotopic values of natural soil NO<sub>3</sub> using the denitrifier method (Casciotti et al., 2002). 211

212 <sup>f</sup>Respiratory fungal-to-bacterial (F:B) ratio analyzed by SIRIN method (Anderson and Domsch, 1973,

213 214 1975)

<sup>g</sup>Respiratory biomass analyzed by CO<sub>2</sub> production from SIR method.





215	2.2.2 Soil incubation with selective inhibition to determine $N_2O$ forming processes
216	The experimental design included two factors, (i.) microbial inhibition by fungal and/or bacterial
217	inhibitors and (ii.) activity of $N_2O$ reductase analyzed either by inhibition with $C_2H_2$ or quantification
218	by <sup>15</sup> N tracing. To address factor (i.), the SIRIN method for determination of the respiratory F:B ratio
219	based on CO2 emission was modified to determine N2O production by microbial groups. However, in
220	contrast to previous studies by Laughlin and Stevens (2002), McLain and Martens (2006),
221	Blagodatskaya et al. (2010) and Long et al. (2013), we did not pre-incubate the soil with the growth
222	inhibitors, as this could result in changes of the microbial community (e.g. preferential growth of
223	selected organisms). We intended to disturb microbial communities as little as possible.
224	The soil was sieved (2 mm) and pre-incubated at 22 °C for five to seven days in the dark with cotton
225	wool stoppers to allow respiration and aerobic conditions in soil bags. Four microbial inhibitor
226	treatments (each in triplicate) with $c_{opt}$ (glucose) for each soil were established:
227	A Control, without growth inhibitors
228	B With streptomycin sulfate $(C_{42}H_{84}N_{14}O_{36}S_3)$ to inhibit bacterial growth
229	C With cycloheximide $(C_{15}H_{23}NO_4)$ to inhibit fungal growth
230	D With streptomycin and cycloheximide, to inhibit bacterial and fungal growth
231	To address factor (ii.), all microbial inhibitor treatments were conducted in three $N_2O$ reductase
232	varieties, i.e.: with <sup>15</sup> N-NO <sub>3</sub> fertilizer (variety "traced") to quantify N <sub>2</sub> O reduction to N <sub>2</sub> , with natural
233	abundance $NO_3^-$ and 10 kPa $C_2H_2$ in the headspace (variety "+ $C_2H_2$ ") to block $N_2O$ reductase, and
234	with natural abundance $\mathrm{NO}_3^-$ but without blocking $\mathrm{N}_2\mathrm{O}$ reductase, i.e. no $\mathrm{C}_2\mathrm{H}_2$ added (variety "-
235	$C_2H_2$ "). In total, there were 48 experimental treatments and 144 vessels (four Experiments with four
236	inhibitor treatments (A, B, C, D) and three varieties ( <i>traced</i> , $+C_2H_2$ and $-C_2H_2$ ) each in triplicates).
237	The soil was adjusted to 80% water filled pore space (WFPS) with distilled water and simultaneously
238	fertilized with NO <sub>3</sub> <sup>-</sup> (varieties $-C_2H_2$ and $+C_2H_2$ with 50 mg N kg <sup>-1</sup> KNO <sub>3</sub> in Experiment 1, 2 and 3 and
239	with 60 mg N kg <sup>-1</sup> NaNO <sub>3</sub> in Experiment 4 and <i>traced</i> variety with 50 mg N kg <sup>-1 15</sup> N-KNO <sub>3</sub> in
240	Experiment 1, 2 and 3 and 60 mg N kg <sup>-1 15</sup> N-KNO <sub>3</sub> in Experiment 4 with a $^{15}$ N-labeling of 50 atom%
241	(at%)). For each treatment, we incubated 100 g dw soil in 850 mL preserving jars (J. WECK GmbH u.
242	Co KG, Wehr, Germany) with gas inlet and outlet equipped with three port luer lock plastic stopcocks
243	(Braun, Melsungen, Germany). According to the original SIRIN method (Anderson and Domsch,
244	1973, 1978) and a mixture of $c_{opt}$ (glucose) and carrier material talcum (5 mg talcum g dw <sup>-1</sup> ) was added
245	to soil of treatment A and together with the growth inhibitors to the soil of treatments B, C and D. The
246	soil and additives of each treatment were mixed for 90 seconds with a handheld electric mixer. During
247	packing, the soil density was adjusted to a target soil density of 1.6 g cm <sup>-3</sup> in Experiment 1, 2 and 4
248	and of 1.3 g cm <sup>-3</sup> in Experiment 3. To ultimately achieve denitrifying conditions in all treatments and
249	to avoid catalytic NO decomposition in the $+C_2H_2$ variety (Nadeem et al., 2013), the headspace of the
250	closed jars was flushed with N2 to exchange the headspace 10 times. Directly following, 85 mL of the





251 gas in the headspace in variety  $+C_2H_2$  were exchanged by pure  $C_2H_2$  resulting in 10 kPa  $C_2H_2$  in the 252 headspace. The manual sample collection of 14 mL gas in duplicates with a plastic syringe was

performed after six, eight and ten hours (Experiment 1, 2 and 3) or two, four and eight (Experiment 4)

of incubation time, respectively. The removed gas was replaced by the same amount of N<sub>2</sub>.

#### 255 2.3 Gas analysis

256 Gas samples were analyzed for N<sub>2</sub>O and CO<sub>2</sub> concentrations ( $c(N_2O)$  and  $c(CO_2)$ ) with gas 257 chromatography (GC, Agilent 7890A, Agilent, Böblingen, Germany). The detection limit of N<sub>2</sub>O was 258 0.04 ng N h<sup>-1</sup> with a measurement precision of 1% and for CO<sub>2</sub> the detection limit was 4 ng C h<sup>-1</sup> with 259 a measurement precision of 0.5%. As a control, N<sub>2</sub> and O<sub>2</sub> concentrations in the samples were analyzed 260 with GC to ensure anaerobic conditions during the incubation for N<sub>2</sub>O production from denitrification. 261 The N<sub>2</sub>O isotopic analysis of the gas samples of varieties  $-C_2H_2$  and  $+C_2H_2$  were performed on a pre-262 concentrator (PreCon, Thermo-Finnigan, Bremen, Germany) interfaced with a GC (Trace Gas Ultra, 263 Thermo Scientific, Bremen, Germany) and analyzed by isotope ratio mass spectrometry (IRMS, Delta V, Thermo Fisher Scientific, Bremen, Germany) (Brand, 1995; Toyoda and Yoshida, 1999; Köster et 264 al., 2013b). The analytical precision was 0.1 ‰, 0.2 ‰ and 1.5 ‰ for  $\delta^{15} N^{\text{bulk}}_{\text{N2O}}$ ,  $\delta^{18} O_{\text{N2O}}$  and  $SP_{N2O}$ 265 266 values, respectively.

267 The gas samples of variety traced from Experiment 1, 2, and 3 were analyzed for the 29/28 and 30/28 268 ratios of N<sub>2</sub> according to Lewicka-Szczebak et al. (2013) using a modified GasBench II preparation 269 system coupled to IRMS (MAT 253, Thermo Scientific, Bremen, Germany). The gas samples of 270 variety traced from Experiment 4 were analyzed at the Centre for Stable Isotope Research and 271 Analysis (University of Göttingen, Germany). The N<sub>2</sub> produced was analyzed using an elemental 272 analyzer (Carlo Erba ANA 1500) that was coupled to dual inlet IRMS (Finnigan MAT 251) (Well et 273 al., 1998; Well et al., 2006). Isotopic values of N2O of Experiment 4 (variety traced) were analyzed in 274 the same lab using a pre-concentration unit coupled to IRMS (Precon-DeltaXP, Thermo Scientific, 275 Bremen, Germany) (Well et al., 2006). Isotope ratios were used applying the non-random distribution 276 approach to calculate the fraction of  $N_2$  and  $N_2O$  originating from the <sup>15</sup>N-labelled N pool as well as 277 the <sup>15</sup>N enrichment of that N pool  $(a_p)$  (Bergsma et al., 2001; Spott et al., 2006).

### 278 2.4 Inhibitor effects

For interpretation of  $N_2O$  or  $CO_2$  production, the validity of the experimental results with respect to fungal and bacterial  $N_2O$  fluxes was checked using a flux balance comparing the sum of bacterial and fungal inhibition effects (treatments B and C) to the dual inhibition effect (treatment D):

282 D = A - [(A - B) + (A - C)]

283 With A, B, C and D representing the  $N_2O$  production rates of the last sampling time of treatment A, B,

(Eq. 1)

284 C and D, respectively. Assuming that in the other three treatments (A, B and C) non-inhibitable N<sub>2</sub>O





- 285 production was equal to treatment D, N<sub>2</sub>O produced by bacteria or fungi should show the following
- 286 relation between the four treatments:
- 287 (A D) = (B D) + (C D) (Eq. 2)
- 288 The fungal contribution to  $N_2O$  production during denitrification with microbial inhibition ( $F_{FDmi}$ ) can
- 289 be calculated, when  $N_2O$  production of treatment D is significantly smaller than  $N_2O$  production of
- treatments A, B and C by:
- 291  $F_{FDmi} = \frac{(A-C)}{(A-D)}$  (Eq. 3)
- 292 2.5 Isotope methods
- 293 2.5.1 Isotope endmember mixing approach (IEM)

The fungal fraction ( $F_{FD}$ ) contributing to N<sub>2</sub>O production from denitrification in soil samples was calculated according to the isotope mixing model (IEM) proposed by Ostrom et al. (2010), which was established for calculating the bacterial fraction ( $F_{BD}$ ) of N<sub>2</sub>O production. Assuming that bacteria (*BD*) and fungi (*FD*) are the only microorganisms responsible for denitrification in soil, the <sup>15</sup>N site preference values of produced N<sub>2</sub>O (*SP*<sub>prod</sub>) results from the *SP*<sub>N2O</sub> mixing balance:

$$299 SP_{prod} = F_{FD} * SP_{FD} + F_{BD} * SP_{BD} (Eq. 4)$$

300 where  $F_{FD}$  and  $F_{BD}$  represent the fraction of N<sub>2</sub>O produced by fungi and other N<sub>2</sub>O sources than fungal 301 denitrification, respectively, and  $SP_{BD}$  and  $SP_{BD}$  are the respective  $SP_{N2O}$  endmember values (Ostrom et 302 al., 2010; Ostrom and Ostrom, 2011). This calculation was based on the assumption that the sum of 303  $F_{BD}$  and  $F_{FD}$  equals 1 and that N<sub>2</sub>O reduction to N<sub>2</sub> is negligible. The mean  $SP_{FD}$  value was assumed to 304 be 33.6 % (Sutka et al., 2008; Maeda et al., 2015; Rohe et al., 2014a; Rohe et al., 2017) and the  $SP_{BD}$ 305 value from heterotrophic denitrification was assumed with minimum and maximum values from -7.5 306 to +3.7 ‰ (Yu et al., 2020). For this IEM approach, only results from variety  $+C_2H_2$  could be used to 307 calculate the fungal fraction contributing to N<sub>2</sub>O production ( $F_{FD SP}$ ), as microorganisms of this variety 308 produce  $N_2O$  that is not affected by reduction to  $N_2$ . The  $F_{FD}$  sp contributing to  $N_2O$  production during 309 denitrification was calculated from the measured  $SP_{N2O}$  value from treatment A of variety  $+C_2H_2$  as 310 SP<sub>prod</sub> value (Eq. 4). In case successful inhibition (modified SIRIN approach), Eq. 4 was solved for the 311  $SP_{FD}$  value using  $F_{FD}$ ,  $F_{BD}$ , and  $SP_{prod}$  values of the respective variety.

# 312 2.5.2 SP/ $\delta^{18}$ O isotope mapping approach (SP/ $\delta^{18}$ O Map)

313 The  $F_{FD}$  contributing to N<sub>2</sub>O production from denitrification in soil samples was also estimated with 314 the SP/ $\delta^{18}$ O Map ( $F_{FD\_MAP}$ ) (Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et al., 2020). This 315 method allows for estimation of both: the  $F_{FD}$  and N<sub>2</sub>O product ratio [N<sub>2</sub>O/(N<sub>2</sub>+N<sub>2</sub>O)] (product 316 ratio<sub>Map</sub>). For precise estimations, the  $\delta^{18}$ O values of soil water ( $\delta^{18}O_{H2O}$ ) applied in the experiments 317 are needed and these values were not determined. However, since we have independent information on





the N<sub>2</sub>O product ratio from the *traced* variety, we can calculate the possible  $\delta^{18}O_{H2O}$  values of soil to 318 get the nearest N2O product ratios in natural and <sup>15</sup>N treatments. The fitting of values was performed 319 for mean, minimal und maximal values of SPBD (-1.9, -7.5 and 3.7‰, respectively) and aimed at 320 321 obtaining the minimal difference between product ratio<sub>Map</sub> and measured in traced variety, i.e., the 322 minimal value of (product ratio<sub>15N</sub> - product ratio<sub>Map</sub>)<sup>2</sup> for  $-C_2H_2$  and  $+C_2H_2$  variety (for explanation of 323 the product ratio see next section). This further allows obtaining the possible ranges for  $F_{FD}$  for 324 particular fitted values (Table 4). The calculations with this approach may be performed assuming two 325 different scenarios of the interplay between N<sub>2</sub>O mixing and reduction (Lewicka-Szczebak et al., 326 2017; Lewicka-Szczebak et al., 2020) but for this study both scenarios yield almost identical results 327 (maximal difference of 0.02 in N<sub>2</sub>O product ratio and  $F_{FD}$  was found), due to  $F_{BD}$  near 1. Hence, we 328 only provide the results assuming the reduction of bacterial N<sub>2</sub>O followed by mixing with fungal N<sub>2</sub>O.

329 2.5.3 Product ratio  $[N_2O/(N_2+N_2O)]$  of denitrification

330 The variety *traced* served to assess N<sub>2</sub>O reduction during denitrification in each experiment. The 331 product ratio of denitrification  $[N_2O/(N_2+N_2O)]$  as given by the variety *traced (product ratio<sub>15N</sub>)* was 332 calculated as:

333 
$$product \ ratio_{15N} = \frac{{}^{15}N_{N2O}}{{}^{15}N_{N2} + {}^{15}N_{N2O}}$$
 (Eq. 5)

334 with  ${}^{I_3}N_{N2O}$  and  ${}^{I_3}N_{N2}$  representing N<sub>2</sub>O and N<sub>2</sub> produced in the  ${}^{15}$ N-labeled fertilizer pool. To check 335 the effectiveness of C<sub>2</sub>H<sub>2</sub> to block the N<sub>2</sub>O reduction, *product ratio*<sub>15N</sub> was compared with *product* 336 *ratio*<sub>C2H2</sub>, where the latter can be calculated from N<sub>2</sub>O production rates of varieties  $-C_2H_2$  and  $+C_2H_2$ :

337 product 
$$ratio_{C2H2} = \frac{N_2 O_{-C2H2}}{N_2 O_{+C2H2}}$$
 (Eq. 6)

338 with  $N_2O_{-C2H2}$  and  $N_2O_{+C2H2}$  representing the N<sub>2</sub>O produced in varieties  $-C_2H_2$  and  $+C_2H_2$ , respectively. 339 If *product ratio*<sub>15N</sub> and *product ratio*<sub>C2H2</sub> were in agreement, a complete blockage of N<sub>2</sub>O reduction 340 could be assumed. This enabled to estimate reduction effects on the isotopic signatures of N<sub>2</sub>O by 341 comparing the isotopic values of N<sub>2</sub>O produced without N<sub>2</sub>O reduction effects of variety  $+C_2H_2$  ( $\delta 0$ 342 values) with isotopic values of N<sub>2</sub>O of variety  $-C_2H_2$ .

The information on the product ratio was used as an additional possibility to calculate the  $F_{FD}$  also for variety  $-C_2H_2$ . First, the Rayleigh-type model presented by Lewicka-Szczebak et al. (2017) and Senbayram et al. (2018) for similar closed-system incubations, the <sup>15</sup>N site preference values of produced N<sub>2</sub>O, i.e. without its reduction to N<sub>2</sub>O ( $SP_{prod}$ ), of variety  $-C_2H_2$  was calculated by correcting SP values of emitted N<sub>2</sub>O, i.e. after partial reduction of produced N<sub>2</sub>O ( $SP_{N2O-r}$ ) from variety  $-C_2H_2$ with the net isotope effect of N<sub>2</sub>O reduction ( $\eta r$ ) and the *product ratio*<sub>15N</sub> as follows:  $SP_{prod} = SP_{N2O-r} + \eta r \ln(product ratio_{15N})$  (Eq. 7)

350 According to (Yu et al., 2020) the  $\eta r$  was assumed to be -6‰. Secondly, Eq.4 was used to calculate the

351  $F_{FD}$  by using  $SP_{prod}$  values of variety  $-C_2H_2$  ( $F_{FD SPcalc}$ ) obtained from Eq. 7





# 352 2.6 Sources of N<sub>2</sub>O produced

353 Assuming that denitrification is the only process producing N<sub>2</sub>O in the incubation experiment, the 354 expected <sup>15</sup>N enrichment in N<sub>2</sub>O produced ( ${}^{15}N_{N2O exp}$ ) was given by

(Eq. 8)

355 
$${}^{15}N_{N20\_exp} [at\%] = \frac{(N_{soil} x {}^{15}N_{nat}) + (N_{fert} x {}^{15}N_{fert})}{N^{bulk}}$$

with  $N_{soil}$ ,  $N_{fert}$  and  $N^{bulk}$  describing the amount of N [mg] in unfertilized soil samples, fertilizer and fertilized soil samples, respectively and  ${}^{15}N_{nat}$  and  ${}^{15}N_{fert}$  is standing for  ${}^{15}$ N enrichment under natural conditions (0.3663 at%) and in fertilizer (50 at%), respectively. Comparison of measured  ${}^{15}$ N enrichment in N<sub>2</sub>O and  ${}^{15}N_{N2O\_exp}$  gave information about the contribution of processes other than denitrification to N<sub>2</sub>O production.

### 361 2.7 Statistical Analysis

362 We conducted several three-way analyses of variance (ANOVA) to test significant effects of soil, experimental variety and treatment on N<sub>2</sub>O production, CO<sub>2</sub> production, and  $SP_{N2O}$ ,  $\delta^{15}N^{\text{bulk}}_{N2O}$  and 363 364  $\delta^{18}O_{N2O}$  values. The pairwise comparison with Tukey's HSD test was made to find differences between soils, varieties and treatments influencing N<sub>2</sub>O production, CO<sub>2</sub> production, and isotopic values. 365 366 Significant effects of soils and treatments on product ratio<sub>C2H2</sub> and product ratio<sub>15N</sub> were tested by two-way ANOVA, while differences between soils and treatments influencing the product ratios were 367 368 tested with pairwise comparison with Tukey's HSD test. Effects of varieties  $-C_2H_2$  and traced on N<sub>2</sub>O 369 and CO<sub>2</sub> production were tested by ANOVA. For this ANOVA, the N<sub>2</sub>O production rate had to be 370  $log_{10}$ -transformed to achieve homogeneity of variance and normality. The significance level  $\alpha$  was 0.1 371 for every ANOVA. For some ANOVAs treatments were excluded, when replicates were n < 3. The 372  $N_2O$  or  $CO_2$  production rates of variety  $+C_2H_2$  were followed over three sampling times by regression. For statistical analysis, we used the program R (R Core Team, 2013). Excel Solver tool was used to 373 determine the  $\delta^{18}O_{H2O}$  values in the application of SP/ $\delta^{18}O$  Map calculations. 374

**375 3. Results** 

#### $3.1 N_2O$ production rates

377 N<sub>2</sub>O and CO<sub>2</sub> production rates of all treatments were similar in magnitude in almost all cases and 378 mostly indistinguishable (Table 2). CO<sub>2</sub> production rates were determined to get additionally 379 information about the denitrifying process. N<sub>2</sub>O production rates exhibited increasing trends with 380 ongoing incubation time for every soil with large variations within the treatments. Contrary to that, 381 CO<sub>2</sub> production rates showed decreasing trends (Figure 1, exemplarily shown for data of variety 382  $+C_2H_2$ ). Calculations of inhibitor effects were based on average N<sub>2</sub>O and CO<sub>2</sub> production rates of the 383 entire incubation period, i.e. 10 hours of incubation time for Experiment 1, 2 and 3 and 8 hours for 384 Experiment 4.



393



385 N<sub>2</sub>O and CO<sub>2</sub> production rates of all  $+C_2H_2$  varieties differed significantly among soils (P < 0.001) and  $N_2O$  production rates differed also significantly among treatments (P < 0.001). Largest  $N_2O$ 386 production rate about 5.5 to 6.1 µg N kg<sup>-1</sup>h<sup>-1</sup> was obtained in Experiment 1 and 3, while in Experiment 387 2 and 4 N<sub>2</sub>O production rates were lower (2.6 and 2.7  $\mu$ g N kg<sup>-1</sup>h<sup>-1</sup>, respectively). N<sub>2</sub>O and CO<sub>2</sub> 388 production rates were significantly larger in variety  $+C_2H_2$  than in variety  $-C_2H_2$  of Experiment 1, 3 389 390 and 4 (P = 0.002, P < 0.010 and P < 0.010 for N<sub>2</sub>O production rate and P = 0.027, P < 0.010 and 391 P = 0.008 for CO<sub>2</sub> production rate, respectively) (Table 2), while  $-C_2H_2$  and  $+C_2H_2$  varieties of Experiment 2 did not differ in N<sub>2</sub>O and CO<sub>2</sub> production rates (P = 0.402 and P = 0.288, respectively). 392



401



- 394 Figure 1: Time series of average N<sub>2</sub>O and CO<sub>2</sub> production rates during incubation of variety  $+C_2H_2$  at the 395 three sample collection times of each soil (Experiment 1 - 4) for treatment A without growth inhibitors, B
- with bacterial growth inhibition, C with fungal growth inhibition, and D with bacterial and fungal growth
- 397 inhibition; *P*-values for linear regressions (significance level  $\alpha \le 0.05$ ). For all significant regressions,  $R^2$ -
- 398 values were  $\ge 0.46$  and in the case of non-significance,  $R^2$ -values were  $\le 0.40$ .

and: There was no detectable CO<sub>2</sub> production in Experiment 4 at the first sampling time after 2 hours.
 (Figure is continued on next page)











Figure 1 continued.

405 Without blockage of N<sub>2</sub>O reductase (variety  $-C_2H_2$ ), N<sub>2</sub>O production rates of treatment A varied significantly among experiments with mean values between 1.6 and 3.6  $\mu$ g N kg<sup>-1</sup> h<sup>-1</sup> ( $P \le 0.001$ ) 406 (Table 2). In Experiment 1, N<sub>2</sub>O production rate was significantly larger (2.7  $\mu$ g N kg<sup>-1</sup> h<sup>-1</sup>) than in 407 408 Experiment 4 (1.6  $\mu$ g N kg<sup>-1</sup> h<sup>-1</sup>) (P = 0.028) in variety -C<sub>2</sub>H<sub>2</sub>. The inhibitor application of each variety 409 revealed in most cases that treatment A (without growth inhibitors) produced most N<sub>2</sub>O, followed by 410 either treatment B (bacterial growth inhibitor; more N<sub>2</sub>O compared to treatment C in Experiments 2, 3 411 and 4) or treatments C (fungal growth inhibitor; more N<sub>2</sub>O compared to treatment B in Experiment 1). 412 In varieties  $-C_2H_2$ ,  $+C_2H_2$  and traced varieties, non-inhibitable organisms (treatment D) showed 413 smallest N<sub>2</sub>O production rates in most cases (i. e. except of variety  $-C_2H_2$  of Experiment 1, varieties -414  $C_2H_2$  and traced of Experiment 3 and variety traced of Experiment 4). Microbial inhibitor treatments 415 differed significantly in N<sub>2</sub>O fluxes of variety  $+C_2H_2$  of each experiment (always  $P \le 0.040$ ), while this was not the case for inhibitor treatments of varieties  $-C_2H_2$  and traced of Experiment 4 (P = 0.154416 417 and P = 0.154, respectively). Significant deviations of treatments without (A) or with full inhibition





- 418 (D) were found in the following cases (Table 2): N<sub>2</sub>O production rate of treatment A was significantly
- 419 larger compared to the other three treatments of Experiment 1 ( $+C_2H_2$  and  $-C_2H_2$ ), Experiment 2
- 420  $(+C_2H_2)$  and Experiment 3  $(+C_2H_2)$ ; treatment D was significantly smaller compared to the other three
- 421 treatments in Experiment 2  $(-C_2H_2)$  only and compared to treatments A and C of Experiment 1
- 422  $(+C_2H_2)$ . Comparing varieties  $-C_2H_2$  and *traced*, N<sub>2</sub>O and CO<sub>2</sub> rates did not differ (P = 0.991 for N<sub>2</sub>O
- 423 production rate and P = 0.490 for CO<sub>2</sub> production rate, respectively), confirming that <sup>15</sup>N-labeling did
- $424 \quad \ \ not \ affect \ N_2O \ and \ CO_2 \ processes.$





426	Table 2: Average CO <sub>2</sub> and N <sub>2</sub> O production rates and N <sub>2</sub> O isotopic values of N <sub>2</sub> O of the last sample
427	collection with and without $C_2H_2$ application in the headspace (varieties $-C_2H_2$ and $+C_2H_2$ ) of each soil

428 (Experiment 1 - 4) for treatments A without, B with bacterial, C with fungal, and D with bacterial and

429 fungal growth inhibition, respectively (standard deviation in brackets, n = 3).

Treatment/ variety	mean N <sub>2</sub> O [µg N kg <sup>-1</sup> h <sup>-1</sup> ]	mean CO <sub>2</sub> [μg C kg <sup>-1</sup> h <sup>-1</sup> ]	δ <sup>18</sup> Ο <sub>N2O</sub> [‰]	$\delta^{15} N^{bulk}{}_{N2O}$	<i>SP</i> <sub>N20</sub> [‰]
Experiment 1 (	Loamy sand, winter	2012)			
$A / -C_2H_2$	2.7 (0.4)a	12.3 (1.7)a	13.1 (0.2)a	-21.9 (1.7)a	1.6 (0.8)a
<b>B</b> / - $C_2H_2$	1.8 (0.2)b	12.8 (1.6)a	13.0 (<0.1)*	-24.2 (0.7)*	-1.3 (0.2)*
C / -C <sub>2</sub> H <sub>2</sub>	2.0 (0.1)b	11.2 (0.5)a	14.6 (0.4)a	-20.0 (0.8)a	-1.6 (0.5)a
$D / -C_2 H_2$	2.1 (0.3)b	13.7 (0.4)a	15.2 (0.5)*	-20.2 (1.8)*	-0.3 (0.5)*
A / + $C_2H_2$	5.5 (0.5)a	17.0 (1.0)a	8.5 (0.1)a	-22.1 (0.3)a	-0.4 (0.3)a
$B / + C_2 H_2$	3.5 (0.1)b	16.1 (0.5)a	7.5 (0.1)a	-26.1 (0.2)a	-1.2 (1.0)b
$C / + C_2 H_2$	4.4 (0.2)c	14.0 (0.6)a	9.3 (0.2)a	-22.4 (0.4)a	-0.9 (0.4)b
$D / + C_2 H_2$	3.3 (0.2)b	14.4 (1.4)a	7.8 (0.3)*	-24.2 (0.1)*	-2.3 (0.7)*
Experiment 2 (	Sand, winter 2012)				
$A / -C_2H_2$	3.2 (0.4)a	13.1 (1.0)a	15.5 (1.8)a	-18.9 (2.6)a	-0.9 (2.5)a
$B / -C_2H_2$	2.4 (<0.1)b	12.1 (0.2)a	15.0 (1.3)a	-23.4 (2.5)a	-0.8 (<0.1)a
$C / -C_2 H_2$	2.5 (0.2)b	12.0 (0.5)a	14.3 (0.1)a	-21.8 (0.2)a	-1.8 (0.2)a
$D / -C_2 H_2$	2.0 (0.3)b	11.0 (0.2)a	13.4 (0.3)a	-24.5 (0.1)a	-1.2 (0.3)a
A / $+C_2H_2$	2.7 (0.4)a	12.7 (2.0)a	12.6 (0.3)a	-18.9 (4.6)a	-1.4 (0.3)a
$B / + C_2 H_2$	2.6 (0.2)a	13.4 (0.7)a	12.3 (0.1)a	-24.6 (0.2)b	-2.0 (0.2)a
$C / + C_2 H_2$	2.5 (0.2)a	12.2 (0.5)a	12.7 (0.1)*	-23.3 (0.2)*	-1.7 (0.4)*
$D / + C_2 H_2$	1.9 (0.2)b	11.7 (0.6)a	12.2 (0.3)a	-26.0 (0.1)b	-1.5 (0.9)a
Experiment 3 (	Silt loam, winter 201	13)			
$A / -C_2H_2$	3.6 (0.2)a	12.3 (1.0)a	26.0 (0.5)a	-20.8 (0.5)a	-0.5 (0.4)a
$B / -C_2H_2$	3.3 (0.4)a	11.6 (1.8)a	24.1 (0.2)b	-22.0 (0.2)b	-0.1 (0.4)a
C / -C <sub>2</sub> H <sub>2</sub>	2.8 (0.1)a	10.6 (0.6)a	27.3 (0.1)b	-20.6 (0.3)a	0.6 (0.2)a
$D / -C_2 H_2$	2.9 (0.4)a	11.2 (0.7)a	26.3 (0.3)a	-21.0 (0.1)a	-0.04 (0.18)a
A / + $C_2H_2$	6.1 (0.3)a	13.3 (1.2)a	15.2 (0.1)a	-25.6 (0.8)a	-2.8 (0.2)a
$B / + C_2 H_2$	5.5 (0.3)b	12.4 (0.8)a	14.9 (0.2)a	-26.3 (<0.1)a	-3.5 (0.4)a
$C / + C_2 H_2$	5.2 (0.2)b	11.7 (0.3)a	16.2 (<0.1)*	-25.2 (0.1)*	-4.0 (0.4)*
$D / + C_2 H_2$	5.1 (<0.1)b	13.0 (0.6)a	16.0 (0.1)b	-25.1 (0.1)a	-4.3 (0.5)a
Experiment 4 (	Loamy sand, summe	er 2011)			
$A / -C_2 H_2$	1.8 (0.1)a	24.5 (1.4)a	25.7 (0.3)a	-30.6 (0.2)a	12.1 (1.6)a
$B / -C_2H_2$	1.2 (0.7)a	20.9 (0.2)b	28.0 (5.0)a	-32.3 (0.7)a	7.7 (1.4)b
C / -C <sub>2</sub> H <sub>2</sub>	1.0 (0.05)a	18.4 (1.9)b	29.3 (0.1)a	-30.0 (0.5)a	4.3 (1.0)c
$D / -C_2 H_2$	0.7 (0.6)a	16.3 (1.2)b	28.9 (1.2)a	-31.8 (2.2)a	3.4 (2.0)c
$A / +C_2H_2$	2.6 (0.3)a	20.8 (3.1)a	13.5 (0.5)*	-34.7 (0.1)*	-1.0**
$B / + C_2 H_2$	2.3 (0.2)a	17.9 (2.4)a	14.3 (1.7)a	-33.8 (0.9)a	-4.9 (0.9)a
$C / + C_2 H_2$	1.2 (1.0)a	17.4 (4.2)a	19.0 (7.0)a	-33.1 (2.8)a	-1.7 (2.7)b
$D / + C_2 H_2$	1.6 (0.1)a	15.0 (1.3)a	14.8 (0.5)a	-35.7 (0.2)a	-4.9 (0.7)c

430 Letters denote significant differences (P < 0.1) among treatments and varieties within a soil.

Asterisks indicate that only two samples (\*) or one sample (\*\*) of triplicates were analyzable.





#### 433 3.2 Isotopologues of N<sub>2</sub>O produced in different varieties and treatments

434 3.2.1 Variety  $+C_2H_2$ 

435  $SP_{N2O}$  values of all experiments, and all treatments of variety  $+C_2H_2$  were within a narrow range 436 between -4.9 and -0.4 ‰ (Table 2), and differed only significantly among treatments of Experiment 4 437 (P = 0.002). In general, there were only small differences among treatments:  $SP_{N20}$  values of 438 treatments A in variety  $+C_2H_2$  differed significantly among soils (P < 0.001), with largest SP<sub>N20</sub> values 439 in Experiment 1 (-0.4 ‰) and smallest  $SP_{N20}$  values in Experiment 3 (-2.8 ‰).  $SP_{N20}$  values of 440 treatment D in variety  $+C_2H_2$  of all soils varied between -1.5 and -4.9 ‰, but only SP<sub>N20</sub> values of 441 Experiment 2 differed significantly from  $SP_{N2O}$  values of the other Experiments (P = 0.006). For 442 treatments B of variety  $+C_2H_2$  SP<sub>N20</sub> values differed only significantly between Experiment 1 and 4, 2 and 4, and 1 and 3 (each P = 0.002). SP<sub>N20</sub> values from treatment C in variety  $+C_2H_2$  did not differ 443 significantly (P = 0.600). For every soil we found significantly larger  $\delta^{18}O_{N2O}$ ,  $\delta^{15}N^{\text{bulk}}_{N2O}$  and  $SP_{N2O}$ 444 445 values in variety  $-C_2H_2$  than in variety  $+C_2H_2$  (P < 0.001), except for Experiment 2, where  $\delta^{15}N^{\text{bulk}}_{N2Q}$ values of variety  $-C_2H_2$  were indistinguishable from those of variety  $+C_2H_2$  (P = 0.400). However, 446 only in a few varieties there were significant differences in  $\delta^{18}O_{N2O}$ ,  $\delta^{15}N^{bulk}_{N2O}$  or  $SP_{N2O}$  values 447 448 between treatments with fungal and bacterial inhibition (B and C, respectively) (Table 2). N<sub>2</sub>O 449 reduction blockage in varieties  $+C_2H_2$  was successful in most cases (Experiment 2, 3 and 4). SP<sub>N20</sub> 450 values of this variety are thus assumed to be valid estimates of  $\delta 0$ , i.e.  $SP_{prod}$  values of N<sub>2</sub>O production, 451 and can thus be used for applying the IEM.

# 452 3.2.2 Variety $-C_2H_2$

453  $SP_{N20}$  values of all experiments and inhibitor treatments of variety  $-C_2H_2$  were within a range of -1.8 454 to 12.1 % (Table 2) and did not differ among inhibitor treatments (P = 0.037). SP<sub>N20</sub> values in variety  $-C_2H_2$  of Experiment 4 was particularly large (3.4 - 12.1 ‰) compared to the other experiments (1.6 to 455 456 -1.6 ‰). As already stated above,  $SP_{N2O}$  values of variety  $-C_2H_2$  were significantly larger than  $SP_{N2O}$ 457 values of variety  $+C_2H_2$  (up to 2.4, 1.5, 4.6 and 4.1‰ in Experiment 1, 2, 3 and 4, respectively). 458 Generally, most SP<sub>prod</sub> values of variety  $-C_2H_2$  (Eq. 7) were smaller than SP<sub>N20</sub> values of variety  $-C_2H_2$ 459 but still larger than  $SP_{N2O}$  values of variety  $+C_2H_2$  and are presented in Table S1 (supplementary 460 Material).

### 461 3.2.3 Variety *traced*

462 The <sup>15</sup>N-labeling of N<sub>2</sub>O ( $^{15}N_{N2O}$ ) or N<sub>2</sub> produced ( $^{15}N_{N2}$ ) gave information about the incorporated N 463 from <sup>15</sup>N-labeled NO<sub>3</sub><sup>-</sup> into N<sub>2</sub>O or N<sub>2</sub> as well as about the N<sub>2</sub>O reduction to N<sub>2</sub>. Microorganisms in 464 each treatment used the <sup>15</sup>N-labeled NO<sub>3</sub><sup>-</sup> in variety *traced* (Table 3) and expected <sup>15</sup>N<sub>N2O</sub> depended on 465 the initial N abundance in NO<sub>3</sub><sup>-</sup> of unfertilized soil (Eq. 7). Experiment 4 is the only one showing a





- large discrepancy between measured (about 30 at%) and calculated  ${}^{15}N_{N2O}$  exp (49 at%) in N<sub>2</sub>O, 466
- 467 whereas the other experiments showed close agreement (Table 3).
- 468 3.3 Product ratios of denitrification and efficiency of N2O reductase blockage by C2H2
- 469 Product ratio C2H2 as well as product ratio 15N of Experiment 2 were significantly larger than of the other
- 470 experiments ( $P \le 0.001$ ) (Table 3). Product ratio<sub>15N</sub> of treatment B was significantly larger than of
- 471 treatment C and D of Experiment 4 (P = 0.032), while all other treatments of other soils did not differ.
- 472 Product ratio<sub>C2H2</sub> did not differ significantly among treatments (P = 0.400). In order to test the
- 473 efficiency of blockage of N<sub>2</sub>O reduction by  $C_2H_2$  application, product ratio<sub>C2H2</sub> (Eq. 5) was compared
- 474 with product ratio<sub>15N</sub> (Eq. 6). In Experiment 1, product ratio<sub>C2H2</sub> was by far smaller than product
- 475 ratio<sub>15N</sub>, while both calculated product ratios were in similar ranges in the other three experiments and
- 476 thus a successful blockage of N2O reduction was assumed for those experiments.

477 Table 3: Average CO<sub>2</sub> and N<sub>2</sub>O production rates of the last sample collection after 10 or 8 hours of variety 478 traced, respectively, with <sup>15</sup>N labeling in N<sub>2</sub>O (<sup>15</sup>N-N<sub>2</sub>O) and the calculated product ratio<sub>15N</sub> of variety traced 479 and product ratio<sub>C2H2</sub> calculated from N<sub>2</sub>O production rates of variety  $-C_2H_2$  and  $+C_2H_2$  of each soil 480 (Experiment 1 - 4) for treatments A without, B with bacterial, C with fungal, and D with bacterial and 481 fungal growth inhibition, respectively (standard deviation in brackets, n = 3).

	maan N O	meen CO		15 N	Calc. total	Calc. total
Treatment	$[\mu g N kg^{-1} h^{-1}]$	$[\mu g N kg^{-1} h^{-1}]^*$	<sup>15</sup> N <sub>N20</sub> [at%]	[at%] <sup>a</sup>	ratio <sub>15N</sub> <sup>b*</sup>	ratio <sub>C2H2</sub> c*
Experiment 1 (	Loamy Sand, 2012	2)				
А	2.6 (0.4)	13.1 (1.7)	36.8 (0.1)		0.80 (0.02)	0.48 (0.07)
В	1.5 (0.3)	11.5 (2.4)	36.4 (0.2)	20	0.76 (0.02)	0.48 (0.05)
С	1.9 (1.5)	12.2 (1.1)	36.9 (<0.1)	59	0.72 (0.05)	0.45 (0.04)
D	1.5 (<0.1)	12.5 (0.5)	36.8 (0.1)		0.69 (0.02)	0.54 (0.05)
Experiment 2 (	Sand, 2012)					
А	2.4 (<0.1)	12.9 (0.1)	43.2 (<0.1)		0.94 (0.01)	1.04 (0.10)
В	1.9 (<0.1)	11.6 (0.2)	43.0 (0.1)	4.4	0.94 (0.01)	0.81 (0.04)
С	2.4 (0.1)	12.8 (0.6)	43.2 (0.1)	44	0.95 (0.01)	0.99 (0.09)
D	1.7 (0.1)	12.0 (0.3)	42.7 (0.1)		0.93 (0.01)	0.98 (0.04)
Experiment 3 (	Silt loam, 2013)					
А	2.9 (0.2)	10.4 (0.5)	35.8 (<0.1)		0.62 (<0.01)	0.52 (0.04)
В	3.2 (0.2)	12.0 (0.9)	35.5 (<0.1)	24	0.62 (0.01)	0.59 (0.02)
С	2.2 (0.3)	9.8 (2.0)	35.5 (<0.1)	34	0.59 (0.02)	0.48 (0.04)
D	2.3 (0.1)	9.9 (0.7)	35.3 (<0.1)		0.62 (0.01)	0.51 (0.04)
Experiment 4 (	Loamy Sand, 201	1)				
А	1.6 (0.6)	31.1 (12.5)	31.1**		0.54 (0.05)	0.63 (0.10)
В	1.7 (<0.1)	23.2 (3.0)	26.5**	40	0.59 (0.03)	0.63 (0.17)
С	1.2 (<0.1)	17.9 (0.8)	30.1*	49	0.50 (0.01)	0.62 (0.02)
D	1.2 (<0.1)	17.1 (0.4)	33.5*		0.50 (0.01)	0.53 (0.12)

Asterisks indicate that only two samples (\*) or one sample (\*\*) were analyzed.  ${}^{al5}N_{N2Oexp}$  [at%] was calculated from Eq. 7. 482

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<sup>b</sup>product ratio<sub>15N</sub> =  $[N_2O/(N_2+N_2O)]$  with N<sub>2</sub>O or N<sub>2</sub> production rates from variety traced; see Eq. 5 484

485 <sup>c</sup>*product ratio*<sub>C2H2</sub> =  $[N_2O_{-C2H2}/N_2O_{+C2H2}]$  with N<sub>2</sub>O production rate from varieties  $-C_2H_2$  and  $-C_2H_2$ ; see Eq. 486 6, cf. Table 2





487 3.4 Fungal contribution to N<sub>2</sub>O production from denitrification by microbial inhibitor approach
 488 (modified SIRIN)

489 When calculating  $F_{FDmi}$ , N<sub>2</sub>O production rates of treatment D must be significantly smaller compared 490 to the other treatments and the flux balance according to Eq. 1 and 2 must be consistent. This was only 491 the case in Experiment 2 of variety  $+C_2H_2$ . The calculated  $F_{FDmi}$  (Eq. 3) was  $0.28 \pm 0.90$  (Table 5). The 492 respective flux of fungal N<sub>2</sub>O was  $0.24 \pm 0.08 \ \mu g \ N \ kg^{-1} \ h^{-1}$ . For all other experiments calculation of 493  $F_{FDmi}$  was not possible.

494 3.5 Fungal contribution to N<sub>2</sub>O production from denitrification by the SP endmember mixing
 495 approach (IEM) and SP/δ<sup>18</sup>O isotope mapping approach (SP/δ<sup>18</sup>O Map)

When applying SP/ $\delta^{18}$ O Map, we can assess the plausibility of the determined  $F_{FD}$  values based on the 496  $\delta^{18}O_{H2O}$  values obtained from the fitting ( $\delta^{18}O_{H2O}$  value in Table 4) and the fitting outcome, i.e. the 497 difference between product ratio<sub>15N</sub> and product ratio<sub>MAP</sub> (Diff in Table 4). The most probable  $\delta^{18}O_{H2O}$ 498 499 value for our experiments can be assumed based on the fact that Braunschweig tap water was added to 500 soil and the original soil water also represent the isotope characteristics typical for this region which is 501 about -7.4‰ (long-term mean Braunschweig precipitation water (Stumpp et al., 2014)). Depending on 502 the season and evaporative losses, this value may slightly vary and the most possible range of soil 503 water in our experiments may vary from about -11 to -4‰ as observed in other experiments conducted 504 in our laboratory in similar conditions (Lewicka-Szczebak et al., 2014; Rohe et al., 2014a; Lewicka-505 Szczebak et al., 2017; Rohe et al., 2017). Taking this into account, we can say that for Experiment 1, 506 the fungal contribution must be below 0.02, because to obtain any larger  $F_{FD}$  values unrealistically 507 small  $\delta^{18}O_{H2O}$  values (of -14.9‰) must be fitted (see Table 4). For Experiment 2 both the smaller 508  $F_{FD MAP}$  values of 0.01 and the larger ones up to 0.15 are possible, since they are associated with very 509 realistic  $\delta^{18}O_{H2O}$  values (of -6.3 and -10.1, respectively) and identical *Diff* of 0.04 (Table 4). For 510 Experiment 3 the only plausible fitting can be obtained for the smallest  $SP_{BD}$  values, which are 511 associated with a  $\delta^{18}O_{H2O}$  value of -5.6% (Table 4). Although the *Diff* for this fitting is slightly higher, 512 the other fittings must be rejected due to unrealistic  $\delta^{18}O_{H2O}$  values (of -1.7 and +3.7‰), hence  $F_{FD MAP}$ 513 values must be 0.04-0.09. Similarly, for Experiment 4, the only plausible fitting can be obtained for 514 the smallest  $SP_{BD}$  values, which are associated with a  $\delta^{18}O_{H2O}$  value of -6.8‰ (Table 4) and indicate 515  $F_{FD MAP}$  values from 0.11 to 0.20. Here this fitting also shows clearly the smallest Diff of only 0.01 516 (Table 4). However, except for Experiment 4, where the Diff is smallest for the last fitting, the Diff 517 values for other experiments are very similar for different fittings with the largest values in 518 Experiment 3. A better fit (showing smaller *Diff* values) was not possible with any other  $SP_{BD}$  and 519  $\delta^{18}O_{H2O}$  values. The  $F_{FD}$  sp ranged between 0 and approximately 0.15 (Table 5). The results obtained 520 from SP/ $\delta^{18}$ O Map show  $F_{FD MAP}$  reaching up to 0.14, 0.15, 0.09 and 0.20 for Experiments 1, 2, 3, and 521 4 respectively (Table 4, Table 5). 522





523	Table 4: Summary of the results provided by SP/ $\delta^{18}$ O Map for fraction of fungal denitrification ( $F_{FD MAP}$ )
524	and N <sub>2</sub> O product ratio (product ratio <sub>MAP</sub> ) in the acetylated ( $+C_2H_2$ ) and non-acetylated ( $-C_2H_2$ ) treatments
525	for 3 possible $SP_{N2O}$ values from bacterial denitrification ( $SP_{BD}$ ): mean (-1.9‰), maximal (3.7‰), and
576	(1, 1, 2, 5) TI $(1, 2, 5)$ I $(1, 2, 5)$ $(1, 2, 5$

minimal (-7.5%). The  $\delta^{18}$ O values of soil water ( $\delta^{18}$ O<sub>H2O</sub>) were fitted to get the lowest difference (*Diff*) between product ratio determined with <sup>15</sup>N treatment and SP/ $\delta^{18}$ O Map (*product ratio<sub>15N</sub>* and *product ratio<sub>MAP</sub>*). The most plausible fittings are bolded (see discussion for reasons of this choice). 527 528

Experiment	Variety	product ratio <sub>15N</sub>	<i>SP<sub>BD</sub></i> [‰]	$\delta^{18}O_{\rm H2O}$ [%)	product ratio <sub>MAP</sub>	Diff	F <sub>FD MAP</sub>
1	$-C_2H_2$	0.66	-1.9	-11.2	0.66	0.00	-0.01
	$+C_2H_2$	1	-1.9	-11.2	1.00	0.00	0.02
	$-C_2H_2$	0.66	3.7	-6.1	0.65	0.01	-0.14
	$+C_2H_2$	1	3.7	-6.1	1.00	0.00	-0.16
	$-C_2H_2$	0.66	-7.5	-14.9	0.66	0.00	0.08
	$+C_{2}H_{2}$	1	-7.5	-14.9	1.00	0.00	0.14
2	$-C_2H_2$	0.94	-1.9	-6.3	0.90	0.04	0.01
	$+C_2H_2$	1	-1.9	-6.3	1.04	0.04	0.01
	$-C_2H_2$	0.94	3.7	-1.2	0.90	0.04	-0.16
	$+C_2H_2$	1	3.7	-1.2	1.04	0.04	-0.18
	$-C_2H_2$	0.94	-7.5	-10.1	0.90	0.04	0.13
	$+C_2H_2$	1	-7.5	-10.1	1.04	0.04	0.15
3	$-C_2H_2$	0.61	-1.9	-1.7	0.54	0.07	-0.03
	$+C_{2}H_{2}$	1	-1.9	-1.7	1.04	0.04	-0.05
	$-C_2H_2$	0.61	3.7	3.7	0.54	0.07	-0.14
	$+C_2H_2$	1	3.7	3.7	1.03	0.03	-0.24
	$-C_2H_2$	0.61	-7.5	-5.6	0.53	0.08	0.04
	$+C_2H_2$	1	-7.5	-5.6	1.04	0.04	0.09
4	$-C_2H_2$	0.60	-1.9	-3.3	0.66	0.06	0.15
	$+C_{2}H_{2}$	1	-1.9	-3.3	0.96	0.04	-0.03
	$-C_2H_2$	0.60	3.7	1.5	0.72	0.12	0.08
	$+C_2H_2$	1	3.7	1.5	0.91	0.09	-0.21
	$-C_2H_2$	0.60	-7.5	-6.8	0.61	0.01	0.20
	$+C_2H_2$	1	-7.5	-6.8	0.99	0.01	0.11

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- 550 Table 5: Ranges of the fraction of N<sub>2</sub>O produced by fungi ( $F_{FD}$ ) from four soil experiments using four 551 different approaches: Fungal fraction was calculated using a) the microbial inhibitor approach (modified
- 552 SIRIN) (FFDmil, b) the isotopomer endmember mixing approach (IEM) by SP isotope mixing balance
- 553  $(F_{FD_sSP})$ , c) the IEM by  $SP_{N20}$  isotope mixing balance (IEM) for results from variety  $-C_2H_2$  with reduction
- 554 correction to calculate the  $SP_{N20}$  values ( $F_{FD_SPcalc}$ ), and d) the  $\delta^{18}O/SP$  Map ( $F_{FD_MAP}$ ) with  $\delta^{18}O_{N20}$  and 555  $SP_{N20}$  values from variety  $-C_2H_2$  and variety  $+C_2H_2$ . Negative values by IEM and  $\delta^{18}$ O/SP Map are
- 556 assumed to be zero.

Experiment	$F_{FDmi}^{a}$	$F_{FD\_SP}^{b}$	$F_{FD\_SPcalc}^{c}$	$F_{FD\_MAP}{}^{\mathbf{d}}$
1	n.d.	0-0.15	0-0.19	0-0.02
2	0.19-0.37	0-0.14	0-0.15	0.01-0.15
3	n.d.	0-0.09	0-0.18	0.04-0.09
4	n.d.	0-0.11	0-0.21	0.11-0.20

557 <sup>a</sup>Fungal fraction on N<sub>2</sub>O production calculated Eq. 3.

<sup>b</sup>Fungal fraction on N<sub>2</sub>O production calculated by Eq. 4 for variety  $+C_2H_2$  with assuming  $SP_{N2O}$  values of 558

559 N2O produced by bacteria were 3.7 ‰ (resulting in negative fraction and therefore set to zero) or -7.5 ‰. 560 <sup>c</sup>Eq. 4 to solve for fungal fraction in variety  $-C_2H_2$  with assuming  $SP_{N20}$  values of N<sub>2</sub>O produced by 561 bacteria was 3.7 (resulting in negative fraction and therefore set to zero) or -7.5 ‰ and using reduction 562 correction with  $\eta_r$ =-6 ‰ to calculate SP<sub>prod</sub> values (Senbayram et al., 2018; Yu et al., 2020).

563 <sup>d</sup>Fungal fraction on N<sub>2</sub>O production calculated by SP/ $\delta^{18}$ O Map with assuming most probable SP<sub>N2O</sub> 564 values from bacterial denitrification (according to Table 4)

- 565 n.d.-not determined because of insufficient inhibition.
- 566

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Figure 2: SP/ $\delta^{18}$ O isotope mapping approach (SP/ $\delta^{18}$ O Map) to estimate the contribution of bacteria or 570 fungi to N<sub>2</sub>O produced according to Lewicka-Szczebak et al. (2017) and Lewicka-Szczebak et al. (2020). 571 The isotopic values for natural abundance treatments with acetylene addition ( $+C_2H_2$ , empty symbols) and 572 without acetylene addition ( $-C_2H_2$ , corresponding filled symbols) are shown for four experiments (1-4). 573 The grey rectangles indicate expected ranges of isotopic signatures for heterotrophic bacterial 574 denitrification (BD) and fungal denitrification (FD) (Yu et al. 2020). The black solid line is the mixing line 575 connecting the average expected values for BD and FD, while the red solid line is the mean reduction (for 576 the mean SP values for BD) line and the red dashed line is the minimum reduction line (for the minimal 577  $SP_{N20}$  values for BD).

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## 582 $3.6 SP_{N2O}$ values of N<sub>2</sub>O produced by the fungal soil community

583 Solving Eq. 4 for  $SP_{FD}$  enables to calculate  $SP_{N2O}$  values from the fungal soil community for 584 Experiment 2 (Table 6). Estimates for the ranges of  $F_{FD}$  and  $F_{BD}$  from the results  $(+C_2H_2)$  of the 585 modified SIRIN were obtained ( $F_{FDmi}=0.19-0.37$  and  $F_{BD}=1-F_{FDmi}$  resulted in a range between 0.63 and 0.81, respectively, see section "3.4 Fungal contribution to  $N_2O$  production from denitrification by 586 microbial inhibitor approach (modified SIRIN)"). The  $SP_{prod}$  values of N<sub>2</sub>O ( $SP_{prod} = -1.4$  ‰) of the 587 588 respective treatment A (Table 2, variety  $+C_2H_2$ ) served to calculate  $SP_{N2O}$  values for fungal 589 denitrification for Experiment 2. Assuming -7.5 or 3.7 % for the bacterial SP<sub>N20</sub> endmember values of 590 N<sub>2</sub>O (Toyoda et al., 2005; Sutka et al., 2006; Yu et al., 2020) resulted in SP<sub>FD</sub> values between -10 ‰ 591  $(SP_{BD} = 3.7 \text{ })$  and 25  $(SP_{BD} = -7.5 \text{ })$  (Table 6). The respective  $SP_{FD}$  value for variety  $-C_2H_2$  was 592 in a very similar range between -17 ‰ and 27 ‰ (Table 6) using  $SP_{prod}$  values ( $SP_{prod} = -1.0$  ‰) of the 593 respective treatment A (Table S1).

594

595Table 6:  $SP_{FD}$  values (i.e.  $SP_{N2O}$  values of N2O produced by fungi) by solving Eq. 4 using  $F_{FDmi}$  and  $F_{BD}$ 596from results of modified SIRIN approach and using  $SP_{prod}$  values of varieties  $+C_2H_2$  and  $-C_2H_2$  of597Experiment 2.

Treatment	<i>SP</i> <sub>prod</sub> <b>[‰]</b>	$SP_{BD}[\%]^{a}$	$F_{FDmi}^{b}$	$F_{BD}^{b}$	SP <sub>FD</sub> [%)
		-7.5	0.19	0.81	25
	1.4	3.7	0.19	0.81	-23
$+C_2H_2$	-1.4	-7.5	0.37	0.63	9
		3.7	0.37	0.63	-10
		-7.5	0.19	0.81	27
CII	1.0	3.7	0.19	0.81	-17
-C <sub>2</sub> H <sub>2</sub>	-1.0	-7.5	0.37	0.63	10
		3.7	0.37	0.63	-9

598  $SP_{N20}$  endmember values of bacterial denitrification were taken for calculation (Eq. 4) according to 599 studies with pure cultures (Toyoda et al., 2005; Sutka et al., 2006; Yu et al., 2020).

600 **b**Ranges of  $F_{FDmi}$  and  $F_{BD}$  were calculated using the modified SIRIN approach.

### 601 **4. Discussion**

602 To our knowledge, this was the first attempt to determine  $SP_{N2O}$  values by fungi or bacteria from soil 603 communities using microbial growth inhibitors with a modification of SIRIN and comparing microbial inhibitor and isotopic approaches (IEM and SP/ $\delta^{18}$ O Map) to estimate fungal contribution to N<sub>2</sub>O 604 605 production from denitrification in anoxic incubation. Using IEM revealed that the fungal contribution 606 to N<sub>2</sub>O production was small ( $F_{FD,SP} \le 0.15$  or  $F_{FD,MAP} \le 0.20$ ) in the three soils tested (Table 5). Only 607 one experiment with modified SIRIN allowed the calculation of the fungal fraction producing N<sub>2</sub>O 608 during denitrification ( $F_{FDmi}$  between 0.19 and 0.37 in Experiment 2), which was larger than the  $F_{FD}$  by 609 two isotope approaches ( $\leq 0.20$ ). While the three approaches coincided in showing dominance of 610 bacterial denitrification, the isotopic approaches yielded similar estimates of  $F_{FD}$  and thus did not confirm largest  $F_{FD}$  of Experiment 2. The strict application of the SIRIN method prescribes proof of 611





612 selectivity of the inhibitors (i.e., streptomycin should not inhibit fungi and cycloheximide should not 613 inhibit bacteria). The SIRIN results obtained with respect to N<sub>2</sub>O production by the fungal or bacterial 614 fraction were rather unsatisfactory and led to unsolved questions, which are discussed in the following 615 sections.

616 4.1 Experimental setup

617 Inhibitor effects, expressed by smaller N<sub>2</sub>O production with selective inhibitors (treatments B, C and 618 D) compared to treatments without inhibitors (A), were only minor in the present study. Previous 619 studies found much larger inhibitor effects by pre-incubating the soil with selective inhibitors 620 (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013; Chen et al., 2014). The 621 experimental design of our incubation setup was, however, in agreement with the original SIRIN 622 method for respiration (Anderson and Domsch, 1975, 1978) without soil pre-incubation with selective 623 inhibitors to minimize disturbance of the soil microbial community. Another study performing similar 624 experiments without pre-incubation with inhibitors did not find effectiveness of application of both 625 antibiotics during long-term application (up to 48 h) (Ladan and Jacinthe, 2016). Inhibitor application 626 without pre-incubating with inhibitors was contrary to previous studies focusing on N<sub>2</sub>O production 627 (Laughlin and Stevens, 2002; Blagodatskaya et al., 2010; Long et al., 2013) and we suppose that pre-628 incubation with selective inhibitors changes the F:B ratio compared to the undisturbed soil 629 considerably more than soil incubation without this pre-incubation step. Additionally, although 630 Blagodatskaya et al. (2010) did not find more inhibitor efficiency after a period of 1 to 20 hours of 631 pre-incubation with streptomycin, they found greater inhibitor effects of cycloheximide with pre-632 incubation phases. This could indicate that microbial distribution changed after exposition to this 633 inhibitor. Anderson and Domsch (1975) stated already that CO<sub>2</sub> production of initially active 634 organisms can only be ensured up to six or eight hours of experimental duration and biomass activity 635 is changed by both inhibitors.

636 It has to be noticed that pre-incubation in previous studies was without glucose, while N2O production 637 was analyzed after the addition of glucose as substrate in the present as well as previous studies 638 (Laughlin and Stevens, 2002; McLain and Martens, 2006; Blagodatskaya et al., 2010; Long et al., 639 2013). Glucose initiates the growth of active heterotrophic organisms. Pre-incubation under 640 denitrifying conditions is not needed for microorganisms to produce denitrifying enzymes as pure 641 cultures synthesized enzymes capable of denitrification within two to three hours (USEPA, 1993). We 642 started gas sample collection after two or four hours, when organisms should have produced 643 denitrifying enzymes and microbial growth of initially active organisms should have started. With incubation time production rates of CO<sub>2</sub> decreased, probably because experimental incubation 644 645 conditions provoked unfavorable conditions and physiological changes, e.g. increasing partial pressure 646 within the closed jars.





647 The conventional practice of SIRIN implies determination of copt(glucose), copt(streptomycin) or 648  $c_{opt}$ (cycloheximide) with an "Ultragas 3" CO<sub>2</sub> analyzer (WösthoffCo., Bochum) (Anderson and 649 Domsch, 1973) with continuous gas flow and we used this method to determine optimal 650 concentrations for SIRIN and used these concentrations for the modified SIRIN approach as well. This 651 optimization procedure was not used in other studies (Laughlin and Stevens, 2002; Blagodatskaya et 652 al., 2010; Long et al., 2013). We supposed that optimal concentrations for  $CO_2$  respiration could work 653 as well for denitrification, if both inhibitors are apt to inhibit the denitrification process as well. SIRIN 654 has so far been tested with isolated cultures and soils for microbial growth on agar and CO2 655 production (Anderson and Domsch, 1975, 1973), but information on N<sub>2</sub>O producing processes, 656 especially denitrification, is still lacking and should be investigated in further studies.

#### 657 4.2 Inhibitor effects

658 Even with both growth inhibitors (treatment D) N<sub>2</sub>O production was large in all experiments, i.e., 659 often not significantly smaller than in the other three treatments. Thus, we suppose similar 660 contributions of non-inhibitable organisms in all treatments. Non-inhibitable organisms could be, for 661 example, bacteria or fungi that are not in growth stage or may be not affected by inhibitors. These 662 organisms could be archaea as well, which are also known to be capable of denitrification (Philippot et 663 al., 2007; Hayatsu et al., 2008). It is known, that archaea are not affected by streptomycin or 664 cycloheximide (Seo and DeLaune, 2010). However, effects of archaeal occurrence in soil or secondary 665 effects on fungi or bacteria were not tested in this study. As stated before, Ladan and Jacinthe (2016) 666 did not find effective inhibition of denitrification by either inhibitor for denitrification although 667 streptomycin and cycloheximide are commonly used to inhibit denitrification of selective groups. 668 Thus, similar experiments with different inhibitors, such as the bactericide bronopol and the fungicide 669 captan presented by Ladan and Jacinthe (2016), should be conducted to evaluate inhibition approaches 670 and isotopic endmember approaches.

# 4.3 Is SIRIN without $C_2H_2$ suitable to examine the fungal contribution to $N_2O$ production in soil?

672 In order to determine SP<sub>N2O</sub> values without alteration by partial reduction of N<sub>2</sub>O to N<sub>2</sub>, C<sub>2</sub>H<sub>2</sub> was used 673 to quantitatively block N2O reduction during denitrification. We found the expected effect of C2H2 674 application, i.e. larger N<sub>2</sub>O production rates in variety  $+C_2H_2$  compared to variety  $-C_2H_2$ . Calculated 675 product ratios varied between 0.5 and 0.95 (product ratio15N) in all soils, showing that N2O reduction 676 can have significant effects on measured N<sub>2</sub>O production and isotopic values. The product ratio is 677 controlled by the reaction rate or by the activity of enzymes capable of N<sub>2</sub>O reduction (Nos) in the 678 system. The calculated *product ratio*<sub>C2H2</sub> was within the same range as *product ratio*<sub>I5N</sub> in Experiment 679 2, 3 and 4 (maximal 9% difference), providing the effective blockage of N<sub>2</sub>O reductase in variety 680  $+C_2H_2$ . Only in Experiment 1 product ratio<sub>15N</sub> and product ratio<sub>C2H2</sub> differed by about 34% with larger 681 calculated reduction in the tracer variety, which might be explained by potential incomplete inhibition





682 by the  $C_2H_2$  method. Nadeem et al. (2013) found some artifacts with  $C_2H_2$ , which resulted in smaller 683  $N_2O$  production rates due to NO oxidation accelerated by  $C_2H_2$  application in the presence of very 684 small oxygen (O) amounts ( $\geq 0.19$  mL L<sup>-1</sup>). Moreover incomplete C<sub>2</sub>H<sub>2</sub> diffusion into denitrifying 685 aggregates might also lead to incomplete N<sub>2</sub>O reductase blockage (Groffman et al., 2006). Both 686 potential methodological errors cannot be excluded for Experiment 1. For the other three experiments 687 (2, 3 and 4) it can be supposed that the isotopic signature of N<sub>2</sub>O of variety  $+C_2H_2$  showed isotopic 688 signatures of produced N<sub>2</sub>O without influences of N<sub>2</sub>O reduction. By comparing varieties  $-C_2H_2$  and  $+C_2H_2$ , isotopologue values of all soils (except  $\delta^{15}N^{\text{bulk}}_{N^{2O}}$  values of Experiment 2) of variety  $-C_2H_2$ 689 were significantly larger than those that of variety  $+C_2H_2$ . The enrichment of residual N<sub>2</sub>O in heavy 690 691 isotopes results from the isotope effect associated with N<sub>2</sub>O reduction (Jinuntuya-Nortman et al., 2008; 692 Well and Flessa, 2009; Lewicka-Szczebak et al., 2014). This explains why  $C_2H_2$  application is essential for analyzing N2O produced by different microbial organism groups from soil using solely 693 694 the modified SIRIN approach without additional isotopic approaches.

695 Moreover, when applying SIRIN without quantifying N<sub>2</sub>O reduction, fungal denitrification is 696 potentially overestimated due to the impact of SIRIN inhibitors on N<sub>2</sub>O reduction. It is evident that 697  $N_2O$  fluxes represent net  $N_2O$  production, i. e. the difference between gross  $N_2O$  production by the 698 microbial community and N<sub>2</sub>O reduction, mainly by heterotrophic bacterial denitrifiers (Müller and 699 Clough, 2014). The goal of SIRIN application has been to determine the contribution of fungi and 700 bacteria, respectively, to net N<sub>2</sub>O production. It has been shown that N<sub>2</sub>O released by microorganisms 701 to air filled pore space can be partially consumed by denitrifiers before being emitted (Clough et al., 702 1998). This means that fungal N<sub>2</sub>O can also be subject to reduction by bacterial denitrifiers. 703 Consequently, inhibiting bacterial denitrification by SIRIN would lead to an overestimation of fungal 704 contribution to N<sub>2</sub>O production. Until now, this effect has not been considered in previous SIRIN 705 papers on fungal N<sub>2</sub>O. This effect can only be evaluated by measuring N<sub>2</sub>O reduction in all inhibitor 706 treatments as in our study. If true, the N<sub>2</sub>O reduction with bacterial inhibition should be smaller than 707 that of the treatments without inhibition or with fungal inhibition. Though, with fungal inhibition, N<sub>2</sub>O 708 reduction is also assumed to be smaller than that without inhibition, because N<sub>2</sub>O produced by fungi is 709 missed for bacterial reduction. The product ratio is a measure for the  $N_2O$  reduction to  $N_2$ . However, 710 regarding the product  $ratio_{15N}$ , there was no evidence of different N<sub>2</sub>O reduction effects between the 711 SIRIN treatments. The product ratio<sub>C2H2</sub> also revealed indistinguishable values between SIRIN 712 treatments in Experiment 1 and 4, but it was slightly larger in Experiment 3 with bacterial inhibition 713 compared to the other treatments. However, this effect was very small, which would only cause small 714 overestimation of fungal contribution. The smallest N<sub>2</sub>O reduction was found in Experiment 2 715 (product ratio<sub>C2H2</sub> values near 1), with smallest product ratio<sub>C2H2</sub> with bacterial inhibition (0.81). This 716 could result in an overestimation of bacterial contribution, since with blockage of N2O reduction, gross 717  $N_2O$  production of bacteria is measured. The product ratio<sub>15N</sub> and product ratio<sub>C2H2</sub> were between 0.5 718 and 1 and  $N_2O$  reduction was thus never consuming most of the produced  $N_2O$ . Hence both the  $C_2H_2$ 





719 and Streptomycin effects on SIRIN results were probably low. But the product ratio in soil 720 denitrification exhibits the full range from 0 to 1, meaning that this effect can be quite relevant and 721 must thus be considered in future studies.

722 4.4  $SP_{N2O}$  values of N<sub>2</sub>O produced by microbial communities

723 The  $SP_{N2O}$  values of each soil indicated predominantly bacteria to be responsible for N<sub>2</sub>O production 724 during denitrification, assuming that results of  $SP_{N2O}$  values of denitrification by pure bacterial 725 cultures is transferable to bacteria of soil communities contributing to denitrification. The latter 726 assumption has been confirmed repeatedly in soil incubation studies, where in absence of N<sub>2</sub>O 727 reduction smallest SP<sub>N20</sub> values have been found that were within the range of bacterial pure cultures 728 (Lewicka-Szczebak et al., 2015; Lewicka-Szczebak et al., 2017; Senbayram et al., 2018). Therefore, 729 there was no unequivocal evidence of fungi contributing to N<sub>2</sub>O production during denitrification, 730 although the isotopic approaches revealed a fungal contribution up to 0.20 on  $N_2O$  production during 731 denitrification. The  $SP_{N20}$  values of treatment A within variety  $+C_2H_2$  showed that the signature of 732 produced N<sub>2</sub>O was not affected by reduction effects and might give evidence of the microbial 733 community contributing to N2O production regarding differences in SPN2O values of pure bacterial or 734 fungal culture studies (Sutka et al., 2006; Sutka et al., 2008; Frame and Casciotti, 2010; Rohe et al., 735 2014a). However, variations in  $SP_{N20}$  values of treatments A of variety  $+C_2H_2$  are very small and do 736 not give a clear evidence of any differences in microbial soil community producing N<sub>2</sub>O. Lewicka-737 Szczebak et al. (2014) analyzed  $SP_{N2O}$  values of denitrification with blockage of N<sub>2</sub>O reduction by 738  $C_2H_2$  for the same soils as used in the present study for Experiment 1 and 4 as well as Experiment 3 739 and revealed  $SP_{N2O}$  values between -3.6 and -2.1 ‰, which is similar to the respective  $SP_{N2O}$  values of 740 the present study from -4.9 to -0.4 %. This reinforces the conclusion that bacteria dominate gross N<sub>2</sub>O 741 production under anoxic conditions in both these soils. However, other studies found larger SP<sub>N20</sub> 742 values of produced N<sub>2</sub>O unaffected by the reduction effect of up to +6 ‰ (Köster et al., 2013a) most 743 probably as a result of larger contributions of fungi to N<sub>2</sub>O production. However, those results were 744 obtained in an experimental setup with ambient oxygen concentration, without glucose amendment 745 and without C<sub>2</sub>H<sub>2</sub> inhibition of N<sub>2</sub>O reduction since N<sub>2</sub> gas fluxes were directly measured. It was also 746 discussed before that short-time incubations under static conditions as presented here, may promote 747 bacterial over fungal growth, which may also be transferable to denitrification activity by both 748 organism groups (Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et al., 2014). Additionally to this, 749 the selection of glucose as substrate in the selected concentration may promote bacteria compared to 750 fungi even more (Koranda et al., 2014; Reischke et al., 2014).

# 751 $4.5 \, \delta^{18}O_{N2O}$ values

The analysis of  $\delta^{l8}O_{N2O}$  values can give information about O exchange between water and denitrification intermediates by various microorganisms (Aerssens et al., 1986; Kool et al., 2007; Rohe





et al., 2014b; Rohe et al., 2017). The range of  $\delta^{I8}O_{N2O}$  values in our study for variety  $+C_2H_2$  (7.5 to 19.0 ‰) was quite similar to the range found by Lewicka-Szczebak et al. (2014) for the same soils (4.8 to 16.3 ‰), where almost complete O exchange with soil water was documented. Hence, for this study the O exchange was probably also very high. However, there were no remarkable differences in  $\delta^{I8}O_{N2O}$  values among treatments within one variety and soil and therefore we assume no differences in O exchange among the treatments.

The information on  $\delta^{18}O_{N20}$  values combined with known  $\delta^{18}O_{H20}$  values is also precious information 760 761 for differentiation between N<sub>2</sub>O mixing and reduction processes (Lewicka-Szczebak et al., 2017). 762 However, for this study,  $\delta^{18}O_{H2O}$  values were not analyzed. However, due to parallel *traced* variety 763 experiments, we could determine possible  $\delta^{18}O_{H2O}$  values for the particular SP<sub>N2O</sub> values of bacterial denitrification mixing endmember (Table 4). Since the  $\delta^{18}O_{H2O}$  value for the particular geographic 764 765 region can be assessed based on the known isotopic signatures of meteoric waters (Lewicka-Szczebak 766 et al., 2014; Stumpp et al., 2014; Lewicka-Szczebak et al., 2017; Buchen et al., 2018) the most plausible ranges of  $\delta^{18}O_{H2O}$  values can be used to indicate the plausible ranges of  $F_{FD MAP}$  values. In 767 case of precisely determined  $\delta^{18}O_{H2O}$  values, the calculated  $F_{FD MAP}$  values could be more precise, 768 769 however, here we show that in case of missing  $\delta^{18}O_{H2O}$  values but known product ratio, the SP/ $\delta^{18}O$ 770 Map can also provide information on N<sub>2</sub>O production pathway contributions.

## 771 4.6 Co-denitrification

772 The influence of co-denitrification, which is predominantly associated to fungi (Spott et al., 2011), 773 may have a large impact on N<sub>2</sub>O production, since Laughlin and Stevens (2002) found N<sub>2</sub>O production 774 in their experiment derived to 92% from co-denitrification and only 8% from denitrification. So far, there is no study on  $SP_{N2O}$  values of N<sub>2</sub>O produced by co-denitrification. Co-denitrification could have 775 776 been a contributing process in Experiment 4. When N in N<sub>2</sub>O originates only from <sup>15</sup>N-labeled soil  $NO_3^{-}$ , measured  $\delta^{15}N^{bulk}_{N20}$  values as well as the <sup>15</sup>N enrichment of the labelled N pool producing N<sub>2</sub>O 777 778  $(a_p)$  should show identical <sup>15</sup>N enrichment to the labeled soil NO<sub>3</sub>. During co-denitrification, when 779 one N atom in N<sub>2</sub>O originates from labeled NO3<sup>-</sup> and the other one from another unlabeled and unknown N source, this results in  $a_p$  values and <sup>15</sup>N enrichment of produced N<sub>2</sub>O smaller than the 780 respective enrichment of the NO3<sup>-</sup> pool. The <sup>15</sup>N enrichment of soil NO3<sup>-</sup> was about 60% larger than 781 782 the analyzed <sup>15</sup>N enrichment in N<sub>2</sub>O, leading to the assumption that N<sub>2</sub>O was produced not only by 783 denitrification. We also calculated  $a_p$  values of the other three experiments (data not shown) which 784 coincided with the <sup>15</sup>N enrichment of N<sub>2</sub>O (Table 3). Since  $a_p$  would not be affected by contributions 785 of unlabelled N<sub>2</sub>O we can thus exclude the possibility that this smaller enrichment could be caused by 786 dilution of enriched N<sub>2</sub>O from denitrification by N<sub>2</sub>O production from an unknown N source and thus 787 verified that this was due to formation of hybrid N<sub>2</sub>O, probably via co-denitrification (Spott et al., 788 2011). In the other experiments there was no indication of co-denitrification being relevant for N<sub>2</sub>O 789 production since <sup>15</sup>N enrichments of NO3<sup>-</sup> and N2O coincided. The question arises, why hybrid N2O





790 formation was only found when the loamy sand was sampled in summer (June, Experiment 4) but not 791 when it was sampled during winter (December, Experiment 1). Information on substrates for co-792 denitrification, i.e.  $NO_2^{-1}$  and  $NH_4^{+}$  or certain organic N compounds could have been different due to 793 seasonal effects. Moreover, seasonal impacts on microbial community could have been relevant. Since 794 these possible factors were not assessed in our study and their impact on co-denitrification is still 795 poorly understood, it is currently not possible to give an answer here. Thus, only the  $SP_{N2O}$  values in 796 Experiment 4 might be influenced by co-denitrification. But since  $SP_{N2O}$  values of the acetylated 797 treatments of Experiment 4 coincided with the  $SP_{N2O}$  value range of bacterial denitrification and also 798 with  $SP_{N2O}$  values of the other experiments, our data give no indication that co-denitrification produces 799  $N_2O$  with  $SP_{N2O}$  values differing from bacterial denitrification.

4.7 Calculating the fungal fraction contributing to N<sub>2</sub>O production and SP<sub>FD</sub> values

801 Due to the inefficiency of microbial inhibition regarding N<sub>2</sub>O production in most cases, calculation of 802  $F_{FDMi}$  contributing to N<sub>2</sub>O production was only possible for Experiment 2. Comparing the modified 803 SIRIN with the isotopic approaches revealed that the fungal fraction contribution to N<sub>2</sub>O production 804 was smaller (about 0.28 in modified SIRIN,  $\leq 0.15$  with IEM,  $\leq 0.20$  with SP/ $\delta^{18}$ O Map) than the 805 bacterial fraction. Although we did not obtain a very clear picture of various microorganisms 806 contributing to N<sub>2</sub>O production due to the large uncertainties of the calculated fractions, all approaches 807 coincided by showing dominance of bacterial N<sub>2</sub>O. In contrast to SIRIN, the isotopic approaches 808 yielded similar estimates of  $F_{FD}$  for all experiments.

809 In some soil studies using helium incubations the SP<sub>Prod</sub> values obtained by correction for the 810 reduction effect on  $SP_{N20}$  values showed significantly larger values than  $SP_{N20}$  of bacterial 811 denitrification (Köster et al., 2013a; Lewicka-Szczebak et al., 2017; Lewicka-Szczebak et al., 2014; 812 Senbayram et al., 2018; Senbayram et al., 2020). Therefore, it can be supposed that based on the 813 isotopic approaches various soils may largely differ in the microbial community that contributes to 814 N<sub>2</sub>O from denitrification. The three tested soils seemed to contain a microbial community where fungi 815 have minor contributions to N<sub>2</sub>O emissions from denitrification compared to bacteria. However, this 816 may also be due to the applied experimental setup favoring bacterial denitrification by static and 817 strictly anoxic conditions and due to the choice of glucose as substrate. Senbayram et al. (2018) could 818 show in an incubation experiment with sufficient NO3<sup>-</sup> supply, that fungal contribution to 819 denitrification was larger with straw compared to a control without straw addition.

The fungal  $SP_{FD}$  values (section 3.6 "SP of  $N_2O$  produced by the fungal soil community") by SIRIN were highly variable with values between -23 and +25 ‰, which is smaller than the  $SP_{N2O}$  range of  $N_2O$  known from pure cultures (16 - 37 ‰) (Sutka et al., 2008; Rohe et al., 2014a). Unfortunately, both ranges exhibit a large overlap but also some discrepancy, which precludes a clear conclusion whether or not Experiment 2 yielded valid estimates of fungal  $SP_{N2O}$  values. There may be different





826 fungal fraction contributing to denitrification of the tested soils was only small compared to that of 827 bacteria, SP<sub>N2Q</sub> values were estimated using a large endmember range known from pure culture studies 828 only, and possible SIRIN artefacts may have occurred as discussed above. The isotopic approaches 829 should thus be further investigated with soils, where presumable fungi contribute largely to N<sub>2</sub>O 830 production during (e. g. acid forest soils, or litter-amended arable soils) (Senbayram et al., 2018) and 831 using SIRIN with suitable inhibitors (Ladan and Jacinthe, 2016). The critical question whether the 832 isotopic signatures of fungal N<sub>2</sub>O determined in pure culture studies are transferable to natural soil 833 conditions cannot be fully answered with this study due to large uncertainties associated with the 834 results of the SIRIN method.

### 835 5. Conclusions

836 Selective inhibitor and isotopic approaches coincided in showing dominance of bacterial denitrification. Neither the modified SIRIN approach, nor IEM or SP/ $\delta^{18}$ O Map approaches yielded 837 838 larger contributions of the fungal N<sub>2</sub>O fraction in any experiment. Both selective growth inhibitors of 839 modified SIRIN confirmed the expected effect on N<sub>2</sub>O production only in one out of four experiments, 840 and  $SP_{N2O}$  values of fungal N<sub>2</sub>O calculated from this treatment did not appear to be a valid estimate of 841 this value and need further evaluation. There might be several artefacts in the modified SIRIN, where 842 further studies should focus on, e.g. including the effectiveness of inhibitors, changes in microbial 843 community during pre-incubation with inhibitors and effects of bacterial consumption of N<sub>2</sub>O 844 produced by fungi in the presence of bacterial growth inhibitors. The present study could show that 845 consideration of N<sub>2</sub>O reduction in further studies is inevitably necessary. Further studies should also 846 determine the range of  $SP_{N2O}$  values known from fungal denitrification as well as the effect of specific 847 inhibitors on microbial groups producing N2O and reducing N2O during denitrification.

848 Data availability. Gas emission and isotopic data are available from the authors on request.

849 *Author contribution.* HF, NWM, RW and THA designed the experiment. LR carried out the 850 experiment at Thünen Institute for Climate-Smart Agriculture in Braunschweig. AG, DLS and RW 851 helped with isotopic analysis and DLS performed the  $\delta^{18}$ O/SP Map. LR, RW and DLS prepared the 852 manuscript with contributions from all co-authors.

853 *Competing interests.* The authors declare that they have no conflict of interest.

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