



# Calibration of Mg/Ca and Sr/Ca in coastal marine ostracods as proxy of temperature

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Abstract. The Mg/Ca ratio of marine shells has been widely used in temperature paleoreconstructions to understand past marine conditions. Temperature calibrations are known to be species-specific but only available for a few species, despite the large number of known ostracod species. Here, we develop temperature calibrations for two shallow marine ostracods,

- 10 Sinocytheridea impressa and Neomonoceratina delicata, using modern sediment samples. Our results show that adult specimens of these two species are useful as a paleothermometer. We observed significant correlations using the Mg/Ca ratios of both species to the annual (Mg/Ca<sub>S. impressa</sub>=3.7•T-62.7, Mg/Ca<sub>N. delicata</sub>=1.6•T-16.8) and April (Mg/Ca<sub>S. impressa</sub>=2.7•T-37.2, Mg/Ca<sub>N. delicata</sub>=1.6•T-15.7) temperatures. The correlation of temperature to the Mg/Ca ratio of *S. impressa* is more significant, and therefore should be preferred for paleoreconstructions. Re-analysis from satellite data allows us to validate our temperature
- 15 calibration to an extended area around the Pearl River Estuary. Our results show that Mg/Ca of *S. impressa* and *N. delicata* ostracods can be used to reconstruct water temperature at a regional scale, which provides information on the oceanic circulation in coastal areas of the South China Seas. Finally, the Sr/Ca ratios of both species do not correlate with any of the 24 water parameters recorded by the Environmental Protection Department of Hong Kong, including temperature (21.7–24.1 °C), salinity (23.8–33.7 psu), dissolved oxygen (4.3–7.1 mg L<sup>-1</sup>), suspended solids (1.9–35.4 mg L<sup>-1</sup>) and pH (7.7–8.2).

#### 20 1 Introduction

Element/Calcium ratios (E/Ca) of secreted biogenic calcium carbonate by marine organisms, such as foraminifera and corals, have been used as paleo-proxies of environmental marine parameters (Hendy et al., 2002; Lea, 2003; Lea and Boyle, 1989; Linsley et al., 2000; Martin et al., 2002). Mg/Ca in foraminifera and Sr/Ca in corals have been frequently used to reconstruct water temperatures (Beck et al., 1992; Cohen et al., 2001, 2002; Lea, 2003; Rosenthal et al., 2006; Sinclair et al., 1998; Yu

25 and Elderfield, 2008). In addition, Mg/Ca of ostracod shells have also provided valuable information about water temperatures (Chivas et al., 1983, 1986a; De Deckker and Forester, 1988; Dwyer et al., 2002). Several studies have focused their efforts on the development of calibrations for deep water species, such as *Krithe*, which are found in the Pacific, Atlantic, and Arctic oceans (Corrège and De Deckker, 1997; Cronin et al., 1996; Dwyer et al., 1995, 2002; Elmore et al., 2012; Farmer et al., 2012). Other studies have investigated species found in shallow marine environments, such as those from the genus *Loxoconcha* 





- 30 (Cronin et al., 2003, 2005a) and *Cyprideis* (Holmes and De Deckker, 2016; Roberts et al., 2020). However, there is big uncertainty in temperature calibrations to Mg/Ca of marine ostracods due to the low number of calibrations developed in comparison to the number of known species (Holmes and De Deckker, 2012) and the contrasting calibration slopes observed in different species (Yamada et al., 2014). Developing calibrations for new ostracod species can enhance our understanding of the processes controlling ostracod Mg/Ca uptake and broaden its use to other areas of the globe (Holmes and De Deckker,
- 35 2012; Lea, 2003). Mg/Ca and Sr/Ca of ostracod shells have also been used as a proxy of salinity in enclosed water bodies (Chivas et al., 1983) based on two main assumptions: 1) Sr/Ca<sub>water</sub> and Mg/Ca<sub>water</sub> exert a control on Sr/Ca<sub>shell</sub> and Mg/Ca<sub>shell</sub> (Dettman and Dwyer, 2012) and 2) salinity increases simultaneously with the removal of Ca by low-Mg calcite precipitation during evaporation, increasing the water content of Mg relative to Ca (Ito et al., 2003). However, there are several variables that constrain their applicability, such as groundwater inputs, non-equilibrium state of calcite precipitation, among others (Ito
- 40 et al., 2003). In spite of this use in lacustrine systems, Sr/Ca of ostracod shells do not seem to be related to salinity or temperature in shallow marine environments (Dettman and Dwyer, 2012; Ingram, 1998; Roberts et al., 2020). Here, we develop parametric calibrations for Mg/Ca and Sr/Ca of two geochemically unknown species, *Sinocytheridea impressa* and *Neomonoceratina delicata*.

Sinocytheridea impressa and N. delicata are shallow marine ostracods from the superfamily Cytheroidea (Brandão et al.,

- 45 2020), which are mainly distributed in Asian waters (Hong et al., 2019; Tanaka et al., 2011). The abundance of both species has been used as an indicator of sea level transgressions (Chunlian et al., 2013). A high abundance of both species rarely occurs simultaneously as *S. impressa* is commonly found in hypoxic environments with low salinity and high turbidity, while *N. delicata* is common in well-ventilated moderate saline bays (Hong et al., 2019; Irizuki et al., 2005). Paleoenvironmental investigations have focused on the study of assemblages of these two species, but not the elemental composition of their shells
- 50 as indicator of environmental marine variables, such as temperature or salinity. The calibration of E/Ca ratios of their shells with ocean parameters may be used to reconstruct Asian marine environments in different periods and locations, and complement previous studies on sea level transgressions.

Estuaries are ecosystems with high productivity and biodiversity (Day et al., 2013). Water temperature and salinity are important physical properties of these ecosystems, which may control the abundance and type of specimens living in its

- 55 environment. These parameters are controlled by the combined effects of oceanic currents, upwelling waters, river discharge and atmospheric forcings, such as winds and precipitation. The study of water temperature and salinity can help us understand the evolution of marine currents and global atmospheric patterns, which can improve our understanding of glaciation cycles and sea level regressions. For example, Mg/Ca of *Loxoconcha* specimens from Chesapeake Bay has been used as a temperature proxy to evaluate anthropogenic and North Atlantic Oscillation (NAO) impacts on past and present climates (Cronin et al.,
- 60 2005b). In this regard, Hong Kong (HK) and the surrounding waters of the Pearl River Estuary (PRE) are similar. The local waters are affected by multi-annual oceanic and atmospheric patterns such as El Niño-Southern Oscillation (Niu, 2013; Zhang et al., 2013). The development of a new water temperature and salinity proxy will give us a tool to improve our understanding of past marine conditions as well as a better knowledge of oceanic and atmospheric regional patterns.





While estuarine systems have a rich biodiversity, their biological, chemical, and physical properties are highly dynamic (Snedden et al., 2013). Calibration of marine shells with ocean parameters may be challenging under these conditions as the average shell chemistry of several specimens is a result of the calcification of specimens under different marine conditions. Here, we investigate the applicability of ostracod Mg/Ca and Sr/Ca ratios as proxies of temperature in a freshwater-influenced marine system and several factors that control the robustness of the parametric calibrations including: 1) number of shells per site, 2) seasonal variability of ocean parameters, 3) ostracod molting time and 4) the life stage of the ostracods.

### 70 2 Methodology

#### 2.1 Water parameters

## 2.1.1 EPD water samples

Water parameters were obtained from the Environmental Protection Department of Hong Kong (EPD), which records 24 parameters of water quality, including temperature, salinity, pH, and dissolved oxygen around HK (Fig. 1). Monthly

- 75 measurements are performed in 8 water control zones: Tolo Harbour, Southern Waters, Port Shelter, Junk Bay, Deep Bay, North Western, Mirs Bay, Western Buffer, Eastern Buffer and Victoria Harbour (Fig. S1). The records correspond to one or two daily values per month from 1986 to the present date (EPD, 2018). For the calibration, we calculated mean values per month for the last 20 years from the collection time of our sediment samples (i.e. 2012). Extreme values with a probability of exceedance higher than 99% and lower than 1% were removed from each parameter, in order to consider the most probable
- 80 values.

Seawater Mg, Sr and Ca concentrations are available in the Environmental Impact Assessment (EIA) study of the desalination plant in Tseung Kwan O submitted to the EPD in 2013 (Black & Veatch Hong Kong Limited and Water Supplies Department, 2013).









Figure 1 Sampling sites around the Pearl River Estuary (PRE). Ostracods were collected from sediment samples obtained by the Environmental Protection Department of Hong Kong (EPD) and OCEAN-HK. Monthly water quality data is obtained from different sampling locations across Hong Kong (EPD stations). Mean bottom water temperature (BWT) for the year 2011 was obtained from the satellite product of the Copernicus Marine Environment Monitoring Service (CMEMS).

## 90 2.1.2 Copernicus products

Worldwide potential bottom water temperature (BWT), provided by the Copernicus Marine Environment Monitoring Service (CMEMS) of the European Union, is available from 1993 to 2018 at a resolution of 1/12°. Potential BWT is a product calculated from re-analysis, which are intended to be as close as possible to real observations (Drévillon et al., 2018). The product is obtained after the assimilation of satellite observations through real-time marine observations and the modelling of

95 atmospheric and oceanic variables, such as tidal and heat fluxes, among others (CMEMS, 2020). The product has low biases at regional and global scales (<0.4° C and <0.1° C respectively), but higher errors are present in coastal regions due to land cover and river inputs. We calculated the bias and correction factors between the BWT from the CMEMS and EPD in areas where the data overlaps (Fig. S1). These factors were applied to correct the potential BWT in areas where we do not have direct measurements.





#### 100 2.2 Ostracod samples

We investigated *S. impressa* (Brady, 1869; Whatley and Zhao, 1988) and *N. delicata* (Ishizaki and Kato, 1976) from surface sediment samples collected by the EPD in HK in January and July 2012 (Hong et al., 2019; Rodríguez, 2020) and OCEAN-HK in PRE in July 2017. The temperature calibration was developed using only the samples from HK waters (EPD) and it was then validated using ostracods from the PRE sediment samples (Fig. 1). For *S. impressa*, we distinguished adult ostracods by

105 size (Irizuki et al., 2005) and shell morphology. For *N. delicata*, we distinguished adult specimens by size (>450 μm) based on Wang et al. (2018) and Fig. S2.

#### 2.3 Trace element analyses

Elemental concentrations on individual shells were measured by ICP-MS Agilent 7900 in the School of Biological Sciences at The University of Hong Kong. We measured <sup>48</sup>Ca, <sup>43</sup>Ca, <sup>24</sup>Mg, <sup>25</sup>Mg, <sup>86</sup>Sr and <sup>88</sup>Sr. In addition, we measured <sup>27</sup>Al and <sup>56</sup>Fe

- 110 in order to control for potential contamination in our samples. The data was corrected by a blank (2% HNO<sub>3</sub>) measured every third sample and a multi-element standard measured every sixth sample, which was prepared from individual pure elemental solutions (MES1, Mg=28.1 ppb, Sr=29.6 ppb, Al=28.1 ppb, Fe=29.4 ppb, Ca=1830 ppb). A multi-element standard prepared from a pure multi-element solution (MES2, Mg=28.6 ppb, Sr=28.6 ppb, Al=28.6 ppb, Fe = 28.6 ppb and Ca=1861 ppb) and a carbonate standard JCp-1 (Hathorne et al., 2013; Inoue et al., 2004; Okai et al., 2002) were used to check the quality of the
- 115 analysis. Our precision and accuracy improved using <sup>48</sup>Ca and <sup>24</sup>Mg. The accuracy and precision (RSD) of the analysis for JCp-1 and MES2 standards are shown in Table 1 (n=90). The detection limit of the concentrations was initially estimated using the blank as 3σ, resulting in a value lower than 0.3 ppb for Mg, Sr, Al and Fe. Then, the detection limit of the ratios was calculated as 3σ of a solution with a concentration of 0.1 ppb for Mg, Sr, Al and Fe and 990 ppb of Ca (Yu et al., 2005). The precision and accuracy for Al and Fe could not be determined for JCp-1 because of the low ratios within the standard and contamination associated with our analytical procedure.
  - Ostracods shells were sonicated in a methanol bath, rinsed twice with Milli-Q water, bleached with 5% sodium hypochlorite for 12-24 hours and rinsed twice again with Milli-Q water to limit potential contamination affecting the carbonate Mg/Ca and Sr/Ca.





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Ratios	DL	JCp-1 ratio	JCp-1 acc.	JCp-1 RSD (%)	MES2 ratio	MES2 acc.	MES2 RSD
	(mmol mol <sup>-1</sup> )	(mmol mol <sup>-1</sup> )	(%)		(mmol mol <sup>-1</sup> )	(%)	(%)
Mg/Ca	0.7	4.2	0.2	3.7	25.4	-2.7	2.7
Sr/Ca	0.1	8.7	-2.5	2.9	7	-2.5	3.3
Al/Ca	1.4	1.9	>20%*	>20%*	22.9	-2.2	3.0
Fe/Ca	1.0	0.05	Below DL	Below DL	11	-2.2	3.2

Table 1 Detection limit (DL), accuracy (acc.), and precision (RSD) of the standards used to assess the analytical quality of the measurements.

 a) The poor accuracy and precision for Al/Ca ratios in JCp-1 are the result of the low ratios of the standard and external contamination in the analytical procedure.

## 130 3 Results

## 3.1 Hong Kong marine waters

Considering Hong Kong sampling sites, annual mean BWT ranges between  $21.8^{\circ}$  C and  $23.9^{\circ}$  C (Fig. 1). The maximum monthly temperature has been recorded close to the PRE ( $28.6^{\circ}$  C) in August, while the minimum monthly temperature has been recorded eastward Hong Kong Island ( $16.2^{\circ}$  C) in February. Mean annual salinity ranges from 24.7 psu to 33.7 psu. The

135 maximum monthly salinity was measured eastward Hong Kong Island in August (34.3 psu), while the minimum salinity was recorded close to the Pearl River in July (15.6 psu). Annual temperature and salinity inversely correlate considering all EPD stations in Hong Kong waters (R<sup>2</sup>=-0.67). BWT and salinities from the EPD grouped by water control zones are shown in Table 2. A summary of annual mean values of other parameters such as dissolved oxygen and suspended solids can be found in Table S1.

140





Zone	Mg/Ca	Mg/Ca	Sr/Ca	Sr/Ca	Salinity	BWT
	mmol mol <sup>-1</sup>	mmol mol <sup>-1</sup>	mmol mol <sup>-1</sup>	mmol mol <sup>-1</sup>	psu	°C
	(S. impressa)	(N. delicata)	(S. impressa)	(N. delicata)		
Deep Bay	17.8	19.7±4.4	3.3	3.7±0.5	26.5±2.5	23.7±0.3
	(n=1)	(n=9)				
Junk Bay	19.5±4.7	No samples	3.5±0.1	No samples	32.9±0.2	22.5±0.2
	(n=6)					
Mirs Bay	19.8±3.7	18.3±3.3	3.6±0.2	3.4±0.4	33.1±0.5	22.2±0.4
	(n=48)	(n=36)				
Northern	23.6±3.8	19.9±3.2	3.6±0.3	3.5±0.5	30.3±1.1	23.1±0.2
Waters	(n=49)	(n=21)				
Port	18.3±3.2	17.5±3.2	3.6±0.1	3.4±0.3	32.9±0.3	22.5±0.5
Shelter	(n=13)	(n=7)				
Southern	20.9±4.4	17.2±3.9	3.5±0.2	3.2±0.4	32.3±1.0	22.9±0.4
Waters	(n=27)	(n=7)				
Tolo	19.4±2.6	No samples	3.6±0.1	No samples	32.0±0.7	22.8±0.7
Harbour	(n=8)					
Victoria	21.7±4.0	No samples	3.3±0.2	No samples	31.9±0.6	22.9±0.2
Harbour	(n=18)					
HK	21.1±4.2	18.7±3.5	3.5±0.2	3.5±0.4	32.2±1.4	22.7±0.5
Average	(n=170)	(n=80)				

Table 2 Mean and standard deviation (1 sd) of Mg/Ca and Sr/Ca ratios of adult *S. impressa* and *N. delicata*. The number of shells for Mg/Ca and Sr/Ca are the same. Last two columns show mean and standard deviation (1 sd) of annual bottom water salinity and temperature by water control zone of HK.

#### 145 3.2 Ostracods ratios

Mg/Ca mean values of adult *S. impressa* and *N. delicata* ostracods from EPD samples are 21.1 mmol mol<sup>-1</sup> (n=170, min=13.6 mmol mol<sup>-1</sup>, max=34.8 mmol mol<sup>-1</sup>) and 18.7 mmol mol<sup>-1</sup> (n=80, min=10.3 mmol mol<sup>-1</sup>, max=28.1 mmol mol<sup>-1</sup>) respectively. Mg/Ca mean values of adult *S. impressa* and *N. delicata* ostracods from OCEAN-HK are 20.1 mmol mol<sup>-1</sup> (n=51, min=14.4 mmol mol<sup>-1</sup>, max=35.9 mmol mol<sup>-1</sup>) and 19 mmol mol<sup>-1</sup> (n=30, 14.1 mmol mol<sup>-1</sup> to 24.7 mmol mol<sup>-1</sup>) respectively. Mg/Ca

150 mean values in each EPD water control zone are in Table 2.

Sr/Ca mean values of adult *S. impressa* and *N. delicata* ostracods from EPD samples are 3.5 mmol mol<sup>-1</sup> (n=170, min=2.9 mmol mol<sup>-1</sup>, max=4.3 mmol mol<sup>-1</sup>) and 3.5 mmol mol<sup>-1</sup> (n=80, min=2.3 mmol mol<sup>-1</sup>, max=4.7 mmol mol<sup>-1</sup>) respectively, while





for OCEAN-HK samples are 3.6 mmol mol<sup>-1</sup> (n=51, min=2.9 mmol mol<sup>-1</sup>, max=4 mmol mol<sup>-1</sup>) and 3.5 mmol mol<sup>-1</sup> (n=30, min=2.6 mmol mol<sup>-1</sup>, max=4.4 mmol mol<sup>-1</sup>) respectively. Sr/Ca mean values in each EPD water control area are in Table 2.

- Al/Ca mean values of adult *S. impressa* and *N. delicata* from EPD samples are 1 mmol mol<sup>-1</sup> (n=160, min=0 mmol mol<sup>-1</sup>, max=11.7 mmol mol<sup>-1</sup>) and 1.8 (n=76, min=0 mmol mol<sup>-1</sup>, max=5.5 mmol mol<sup>-1</sup>) respectively. Al/Ca mean values of adult *S. impressa* and *N. delicata* from OCEAN-HK samples are 1.2 mmol mol<sup>-1</sup> (n=51, min=0 mmol mol<sup>-1</sup>, max=3.4 mmol mol<sup>-1</sup>) and 1.5 (n=28, min=0 mmol mol<sup>-1</sup>, max=7 mmol mol<sup>-1</sup>) respectively. Mg/Ca and Al/Ca ratios of *S. impressa* and *N. delicata* specimens do not correlate (Fig. S3).
- 160 Fe/Ca mean values of adult *S. impressa* and *N. delicata* from EPD samples are 1.1 mmol mol<sup>-1</sup> (n=160, min=0 mmol mol<sup>-1</sup>, max=11.9 mmol mol<sup>-1</sup>) and 1.5 mmol mol<sup>-1</sup> (n=67, min=0 mmol mol<sup>-1</sup>, max=6.1 mmol mol<sup>-1</sup>) respectively. Fe/Ca mean values of adult *S. impressa* and *N. delicata* from OCEAN-HK samples are 0.6 mmol mol<sup>-1</sup> (n=47, min=0 mmol mol<sup>-1</sup>, max=2.1 mmol mol<sup>-1</sup>) and 1.3 mmol mol<sup>-1</sup> (n=26, min=0 mmol mol<sup>-1</sup>, max=6.5 mmol mol<sup>-1</sup>) respectively.

### 3.3 E/Ca calibrations to water parameters

- 165 Mg/Ca ratios of *S. impressa* significantly correlate to annual water temperature in HK waters ( $R^{2}_{S. impressa} = 0.32$ , ps. *impressa*=0.007), but the Mg/Ca ratios of *N. delicata* ostracods do not correlate if the full data set is used (Table 3). However, the removal of sampling sites with only one shell allow us to produce significant temperature calibrations for both species (Table 3). The highest R<sup>2</sup> at a 95% significance for the temperature calibration to Mg/Ca ratios for *S. impressa* and *N. delicata* was obtained considering a minimum number ( $\eta$ ) of 11 and 3 shells per sampling site respectively (Table 3). The highest R<sup>2</sup> at a
- 170 99.9% significance for S. impressa was obtained considering a minimum number of 7 shells per sampling site, while temperature calibrations for N. delicata did not reach this level of significance. Juvenile ostracods from both species do not correlate with temperature. Juvenile ostracods of S. impressa obtained from samples from the eastern side of HK have similar Mg/Ca ratios to adults, but juveniles close to the PRE have lower Mg/Ca ratios than adults (Fig. 2).





## 175

Minimum number of shells per	S. impressa	N. delicata	
sampling site (η)	R <sup>2</sup> (p-value, sampling sites)	R <sup>2</sup> (p-value, sampling sites)	
1	0.32 (0.007, 22)	0 (0.86, 15)	
2	0.51 (<0.001, 21)	0.37 (0.036, 12)	
3	0.47 (0.001, 20)	0.4 (0.05, 10)	
4	0.57 (<0.001,17)	0.41 (0.063, 9)	
5	0.64 (0.001, 13)	0.41 (0.063, 9)	
6	0.69 (0.001, 12)	0.41 (0.063, 9)	
7	0.7 (0.001, 11)	0.24 (0.408, 5)	
8	0.73 (0.002, 10)	0.16 (0.596, 4)	
9	0.73 (0.002, 10)	0.02 (0.906, 3)	
10	0.76 (0.002, 9)	0.02 (0.906, 3)	
11	0.85 (0.009, 6)	0 (No applicable, 1)	

Table 3 Correlation fit and significance of temperature calibrations using different number of shells per sampling site to calculate the mean value of ostracod Mg/Ca ratios for *S. impressa* and *N. delicata*.

Mg/Ca ratio of *S. impressa* and *N. delicata* also correlates with other water parameters such as volatile suspended solids ( $R_{S. impressa}^2 = 0.60$ , ps. impressa<0.001,  $R_{N. delicata}^2 = 0.49$ , p<sub>N.delicata</sub>=0.037), turbidity ( $R_{S. impressa}^2 = 0.58$ , ps. impressa<0.001,  $R_{N. delicata}^2 = 0.49$ , p<sub>N.delicata</sub>=0.037), turbidity ( $R_{S. impressa}^2 = 0.58$ , ps. impressa<0.001,  $R_{N. delicata}^2 = 0.001$ ,  $R_{N. delicata}^2 = 0.001$ ,  $R_{N. delicata}^2 = 0.001$ ,  $R_{N. delicata}^2 = 0.50$ , p<sub>N.delicata</sub>=0.034), silica ( $R_{S. impressa}^2 = 0.48$ , ps. impressa=0.002,  $R_{N. delicata}^2 = 0.47$ , p<sub>N.delicata</sub>=0.041), salinity ( $R_{S. impressa}^2 = 0.64$ , ps. impressa<0.001,  $R_{N. delicata}^2 = 0.49$ , p<sub>N.delicata</sub>=0.037), nitrite ( $R_{S. impressa}^2 = 0.60$ , ps. impressa<0.001,  $R_{N. delicata}^2 = 0.51$ , p<sub>N.delicata</sub>=0.031) and nitrate ( $R_{S. impressa}^2 = 0.62$ , ps. impressa<0.001,  $R_{N. delicata}^2 = 0.02$ ).







185

Figure 2 Relationships of Mg/Ca and Sr/Ca with annual temperature for adult (filled circles) and A-1 juvenile (open circles) ostracod samples of *S. impressa* ( $\eta$ =7) and *N. delicata* ( $\eta$ =3). Individual samples are shown in light colors. Solid lines show the linear regressions for adult specimens.

The temperature calibration to Mg/Ca is more significant in spring-summer months for both species (Fig. 3), when the ocean 190 temperature increases. The highest correlation occurred in April for *S. impressa* and *N. delicata* ( $\eta_{S. impressa}$ =7,  $R^2_{S. impressa}$ =0.83, ps. impressa<0.001,  $\eta_{N. delicata}$ =3,  $R^2_{N. delicata}$ =0.55, p<sub>N. delicata</sub>=0.015).







Figure 3 (a) Monthly temperature calibrations to Mg/Ca of S. impressa ( $\eta$ =7) and N. delicata ( $\eta$ =3) for adult ostracod. (b) The 195 calibrations for S. impressa and N. delicata for April are Mg/Cas. impressa=2.7•T-37.2 and Mg/Cas. delicata=1.6•T-15.7.

An ANOVA determined that Sr/Ca ratios do not show significant differences between sampling sites for N. delicata (p=0.16). Sr/Ca ratios were significantly different for S. impressa (p=0.02). However this was only the result of samples collected from Victoria Harbour (n=18). After removal of these samples, Sr/Ca ratios of S. impressa were no longer significantly different between sampling sites (p=0.96). Sr/Ca ratios of both species do not correlate with any of the 24 water parameters measured by the EPD.

200

#### **4** Discussion

#### 4.1 Control on Mg/Ca and Sr/Ca ratios of S. impressa and N. delicata

The strongest linear correlation between the 24 parameters (Table S1) measured by the EPD and Mg/Ca ratio was for the annual mean temperature, which suggests that water temperature is the main control of Mg/Ca uptake in adult ostracods of S.

205 impressa and N. delicata (Fig. 2). Correlations with other parameters are also significant but are likely caused by multicollinearity of temperature with other water parameters, such as turbidity, suspended solids, salinity, pH, and nitrite (Fig. S4). The temperature control on Mg/Ca ratio of biogenic material is well-known and has been usually described linearly for ostracods, even though inorganic calcite follows an exponential relationship (Lea, 2003). We observed a linear correlation



210



between water temperature and Mg/Ca ratios of *S. impressa* and *N. delicata* ostracods (Fig. 2), but the significance of the linear calibration is higher for the former species, making it more suitable for temperature reconstructions.

None of the 24 parameters measured by the EPD, including temperature and salinity, exert control on the Sr/Ca of both species. The low Sr/Ca variabilities of *S. impressa* and *N. delicata* suggest that the potential control variables of this ratio, such as water Sr/Ca or vital effects, do not change considerably across different locations within the PRE. A Sr/Ca positive correlation with chlorophyll-a and dissolved oxygen is observed for *S. impressa* specimens, but this relationship is mainly produced by

- 215 specimens with lower Sr/Ca from one sampling site in Victoria Harbour. Dissolved oxygen in this sampling site was particularly low, probably as a result of discharge from the nearby Stonecutters Island sewage treatment plant. The elemental composition of marine waters of HK do not seem to control the ostracod Mg/Ca and Sr/Ca ratios according to our data. Mg/Ca ratios of both species are negatively correlated to salinity (Fig. 4). A potential control of seawater Mg/Ca on ostracod Mg/Ca would be possible if seawater Mg/Ca decreases with salinity. Marine waters have higher Mg and Ca
- 220 concentrations in comparison to freshwaters (Brown et al., 1989; Bruland and Lohan, 2006), and therefore a decrease in Mg concentration cannot occur at high salinity. A higher Ca concentration at a relatively constant Mg concentration is unlikely because previous studies have shown a mostly conservative behavior of Mg and Ca in estuaries and surrounding areas (Millero, 2006; Patra et al., 2012). Moreover, ostracod Sr/Ca is similar at different salinities, supporting the idea that changes in Mg, Sr and Ca water concentrations are not the main control on Mg/Ca and Sr/Ca ostracod ratios. Measurements during 2013 and
- 225 2017 at the desalination plant in Tseung Kwan O (Fig. 1 and S5) show water Mg/Ca and Sr/Ca ranging between 4 and 6 mol mol<sup>-1</sup>, and between 8 to 9 mol mol<sup>-1</sup> respectively during the year. These values are constant even when salinity decreases to 25 psu during the summer, suggesting that the partition coefficient might not be an important control on ostracod Mg/Ca and Sr/Ca ratios.







230 Figure 4 Mg/Ca and Sr/Ca correlation to salinity for *S. impressa* (η=7) and *N. delicata* (η=3) for salinity recorded in April.

The Mg/Ca of A-1 juvenile ostracods of *S. impressa* does not correlate to April temperature (Fig. 2). A significant weak correlation is only observed using June temperature ( $R^2_{S. impressa} = 0.33$ ,  $p_{S. impressa} = 0.031$ ,  $\eta = 2$ ). Juvenile ostracods usually show higher Mg/Ca ratios than adults (Chivas et al., 1986b; Dwyer et al., 2002). However, juvenile specimens of *S. impressa* from sampling sites with salinity lower than 32 psu have lower Mg/Ca ratios than adults (Fig. 4). We hypothesize that environmental

235 factors may increase the bias between ostracod life stages. For example, strong currents can produce preferential post-mortem transport of lighter shells (Boomer et al., 2003), which may foster the mixing of ostracod shells calcified at different temperatures. Our results suggest that juvenile and adult ostracods of *S. impressa* cannot be used indistinctly to reconstruct April temperature supporting previous findings about the incompatible use of Mg/Ca ratios of different ostracod stages (Dwyer et al., 2002).





#### 240 4.2 Factors controlling the robustness of the temperature calibration to Mg/Ca

### 4.2.1 Number of shells per sampling site

The robustness of the annual temperature calibration to Mg/Ca with respect to  $R^2$  and p-values depends on the number of individual shells available to calculate the mean value of each sampling site (Table 3). The removal of sampling sites with a low number of shells increases the  $R^2$  of temperature calibrations in both species, suggesting that a low number of samples

- 245 does not allow us to capture all the natural variability exerted by the temperature on ostracod Mg/Ca ratios. The most significant calibrations were obtained using at least two to four shells for *S. impressa* and at least two to three shells per sampling site for *N. delicata* respectively (Table 3). The margin of error (z•SE, z=critical value, SE= standard error) of confident intervals can be used to provide an estimation of the number of shells needed to obtain a desired level of error. Using temperature calibrations for *Bythocypris* and *Krithe* specimens, Corrège and De Deckker (1997) showed that the use of 4 shells provide an error in the
- 250 temperature prediction lower than 1.4 °C at a 95% confidence. In stratigraphic studies the use of three to five shells is a common practice (Holmes and De Deckker, 2012). Holmes (2008) studied the critical sample size of ostracod specimens to keep a desired error at a 0.99 significance level in Mg/Ca and Sr/Ca from a stratigraphic sequence. The authors concluded that in modern specimens of *Cyprideis Torosa* from shallow marine environments the optimal sample size may vary from 3 to 16 shells. We can use this approach to estimate the number of shells needed per sampling site to keep the error at an acceptable
- 255 level to produce a significant correlation. The use of 1, 2, 3, 4, 5, 6 and 7 shells produce margin of errors of 3.4, 2.4, 2.0, 1.7, 1.5, 1.4 and 1.3 °C at a 95% significance respectively, considering the highest standard deviation observed in our sampling sites ( $\sigma$ =4.7 mmol mol<sup>-1</sup>, n=15) and normally distributed samples (Rodríguez, 2020). Our results suggest that in shallow marine waters the calculation of mean Mg/Ca ratios with at least four individual shells in more than ten sampling sites would likely produce a significant temperature calibration at 99% significance level (p<0.01), accounting for short-term temperature
- variations. In the following analyses, we restricted the minimum number of shells per sampling site to  $\eta=7$  for *S. impressa* and  $\eta=3$  for *N. delicata* as we obtained the highest R<sup>2</sup> at a significance of 99.9% and 95% respectively.

#### 4.2.2 Temporal variability

Most temperature calibrations to Mg/Ca are produced with the annual mean temperature as they are intended to reflect longterm changes of paleoceanographic studies. However, ostracods molt their shells in a short period of time (usually days or

- 265 weeks). Thus, it may be possible to record ocean parameters at a shorter time scale if the correct molting time is known. Our monthly analysis of linear regressions show that the best temperature calibrations to Mg/Ca in terms of high R<sup>2</sup> and significant p-values, were found with the water temperature measured in April for *S. impressa* and *N. delicata* (Fig. 3). This suggests that adult ostracods of both species may inhabit HK waters mainly during spring or early summer, as has been documented for other shallow marine species (Cronin et al., 2005a; Kamiya, 1988). June temperature also correlates with Mg/Ca ratios in both
- 270 species, showing a second peak (Fig. 3). Two periods of calcification has recently been shown for *C. Torosa* (Roberts et al., 2020), which may also occur in *S. impressa* and *N. delicata*. The high correlation between April water temperature and Mg/Ca





ratios indicates the possibility to reconstruct water temperature at a finer temporal scale, which can help to unravel coastal ocean circulation patterns. For example, the interaction of the Hainan, Taiwan and Kuroshio currents (Morton and Wu, 1975) determine the temperature and salinity of Hong Kong waters. *Sinocytheridea impressa* and *N. delicata* may become important

275 tools to determine their interaction by providing information on water temperature during the transition between these currents.

## 4.3 Validation of temperature calibration

April BWTs obtained from the Copernicus product are on average  $-1.5^{\circ}C\pm 1.8$  ( $2\sigma$ ) below the measurements from the EPD (Fig. S1). BWT differences between EPD stations and Copernicus product range from 0 to 3°C, being the largest in stations close to the Pearl River (Fig. S1). We corrected the BWTs obtained from the Copernicus product in all OCEAN-HK sampling locations by  $1.5^{\circ}C$ 

# 280 locations by -1.5°C.

The BWTs estimated from ostracod Mg/Ca ratios in OCEAN-HK sampling locations are in good agreement with BWTs obtained from the Copernicus product for April (Fig. 5 and Table S2). The difference between the estimated temperature by the linear regressions and the potential BWT from the Copernicus product is lower than 3.5% in 7 out of the 8 sites and less than 10% when all sites are considered (Table S2). This suggests that *S. impressa* and *N. delicata* specimens around the PRE

285 follow the regression line developed with ostracods collected from HK waters. Our findings also suggest that calibrations may be done without direct measurements of BWT but using the potential BWT from satellite products. We highlight that the improvement in the quality of BWT products derived from satellite images may facilitate the calibration of multiple species across the world as scientists have global coverage of BWT with a high resolution ( $1/12^{\circ} - 9$  km).







#### 290 Figure 5 Mg/Ca ratios of ostracods from OCEAN-HK over the calibration for S. impressa (Mg/Ca<sub>S. impressa</sub>=2.7•T-37.2) and N. delicata (Mg/Ca<sub>N. delicata</sub>=1.6•T-15.7) performed with samples from the EPD. The BWT from the Copernicus product was used for OCEAN-HK samples.

#### 4.4 Calibrations at a superfamily level

- A unique calibration for ostracods from the same superfamily cannot be performed according to our results. Temperature 295 calibrations to Mg/Ca ratios have been developed for specimens of the same superfamily, including the genus *Loxoconcha* (Cronin et al., 2003), *Cytheropteron* (Yamada et al., 2014) and *Xestoleberis* (Kondo et al., 2005). *Sinocytheridea impressa* and *N. delicata* calcify less Mg/Ca at the same temperature in comparison to other species (Fig. 6). The relative standard errors (RSE) for the calibration of *S. impressa* and *N. delicata* are 1.1 mmol mol<sup>-1</sup> for both species, showing one of the best regression fits of all calibrations for the superfamily *Cytheroidea*. The calibration of *N. delicata* has a similar slope in comparison to
- 300 Xestoleberis specimens (Fig. 6), but S. impressa has the highest slope of all species, which shows that different species from the same superfamily do not share the same calibration slope and intercept. In addition, ostracods from the same superfamily under similar conditions of temperature and salinity may show significant differences in their calibration parameters, such as S. impressa and N. delicata specimens in HK waters or Xestoleberis specimens (Kondo et al., 2005). Factors such as the development of temperature calibrations from multiple species, the number of shells considered per site, variability induced
- 305 after the burial of ostracods, among other reasons, might impoverish the calibration fit. Our study suggests that when speciesspecific calibrations are performed, the relative standard error in the temperature calibration to Mg/Ca may be lower than 1



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mmol mol<sup>-1</sup>. This happens in shells which have not been buried, are oxide-free, and were treated with non-corrosive cleaning reagents such as 5% sodium hypochlorite and methanol (Rodríguez, 2020). Ostracods are temperature sensitive at a very small temperature range and may be useful to understand variation in a localized region. *Sinocytheridea impressa* and *N. delicata* ostracods from other locations can improve the calibration shown in this study.



#### Figure 6 Temperature versus Mg/Ca ratios for specimens of the superfamily Cytheroidea

The lower Mg/Ca in *S. impressa* and *N. delicata* ostracods in comparison to other species from the same superfamily may be the result of other variables such as pH, DIC or the calcite saturation index (Holmes and De Deckker, 2012). *Loxoconcha* from

- 315 Chesapeake Bay (Cronin et al., 2003, 2005a) most likely dwells in the most similar environment as *Sinocytheridea* and *Neomonoceratina* due to its presence in a big estuary in a polyhaline system at similar depths. Measurements of pH and dissolved inorganic carbon (DIC) in HK and Chesapeake Bay waters allow us to compare the saturation index in both basins. A bottom water pH ranging from 7.8 to 8.2 (EPD, 2018) and DIC over 1850  $\mu$ M (Guo et al., 2008; Yuan et al., 2011) indicate that the calcite is oversaturated ( $\Omega \sim 3$ ) in our sampling sites for most of the year. On the other hand, pH and DIC in the bottom
- 320 waters of the Chesapeake lower bay in 2006 ranged from 7.92 to 7.96 and 1717.4 μM to 1865.4 μM respectively, with a calcite saturation index over 2.8 (Brodeur et al., 2019), suggesting that the lower ostracod Mg/Ca ratios in *S. impressa* and *N. delicata* ostracods are not a direct result of different conditions of pH, DIC or calcite saturation index. Species-specific biomineralization types may also play an important role in the temperature response to ostracod Mg/Ca ratios of different species, but our measurements do not allow us to explore this factor.





#### 325 5 Conclusion

The Mg/Ca ratios of *S. impressa* and *N. delicata* ostracods can be used as proxies of water temperature in shallow marine environments. The temperature dependence of *S. impressa* is higher and therefore it is more suitable for temperature reconstructions. This study shows that: 1) the number of shells per sampling site has an important impact on the calibration robustness due to the strong seasonal variability of temperature in estuaries and coastal areas, and therefore we recommend

- 330 the use of four or more shells per sampling site, 2) ostracods can give information on monthly water temperatures 3) the temperature reconstruction based on Mg/Ca ratios of *S. impressa* and *N. delicata* specimens has the potential to give insight to past ocean circulation in coastal areas of the South China Seas and 4) ostracods from the same superfamily show different calibration curves, which does not seem to be controlled by the water Mg/Ca ratios of marine waters, temperature or the carbonate system. A good understanding of ostracod molting time will likely improve calibrations and the identification of the
- 335 calcification temperature for S. impressa and N. delicata ostracods.

## Reference

Beck, J. W., Edwards, R. L., Ito, E., Taylor, F. W., Rougerie, F., Joannot, P. and Henin, C.: Sea-Surface Temperature from Coral Skeletal Strontium/Calcium Ratios, Science, 257(5070), 644–647, 1992.

 Black & Veatch Hong Kong Limited and Water Supplies Department: Desalination Plant at Tseung Kwan O – Feasibility

 340
 Study.
 EIA
 Project
 Profile,
 Hong
 Kong
 SAR.
 [online]
 Available
 from:

 https://www.epd.gov.hk/eia/register/report/eiareport/eia\_2292015/Table of Content\_PDF.htm, 2013.
 Boomer, I., Horne, D. J. and Slipper, I. J.: The use of ostracods in palaeoenvironmental studies, or what can you do with an

Ostracod shell?, Paleontol. Soc. Pap., 9, 2003.

Brady, G. S.: Les fonds de la mer, edited by L. de. Folin and L. Périer., 1869.

Brandão, S. N., Angel, S. N., Karanovic, I., Perrier, V. and Meidla, T.: World Ostracoda Database, Last access 01 July 2020, doi:10.14284/364, 2020.
 Brodeur, J. R., Chen, B., Su, J., Xu, Y. Y., Hussain, N., Scaboo, M. M., Zhang, Y., Testa, J. M. and Cai, W. J.: Chesapeake

bay Inorganic Carbon: Spatial distribution and Seasonal Variability, Front. Mar. Sci., 6, 1–17, doi:10.3389/fmars.2019.00099, 2019.

Brown, J., Colling, A., Park, D., Phillips, J., Rothery, D. and Wright, J.: Seawater: Its composition, properties and behaviour, edited by G. Bearman., 1989.
 Bruland, L. W. and Lohan, M. C.: Controls of Trace Metals in Seawater, in The oceans and marine geochemistry, pp. 23–45., 2006.

Chivas, A. R., De Deckker, P. and Shelley, J. M. G.: Magnesium, strontium, and barium partitioning in nonmarine ostracode

355 shells and their use in paleoenvironmental reconstructions—a preliminary study, in Applications of Ostracoda. Proc. Eighth Int. Symp. Ostracoda. Dep. Geosciences, Univ. Houston., vol. 10, edited by R. Maddocks, pp. 238–249., 1983.



370



Chivas, A. R., De Deckker, P. and Shelley, J. M. G.: Magnesium and strontium in non-marine ostracod shells as indicators of palaeosalinity and palaeotemperature, Hydrobiologia, 143(1), 135–142, doi:10.1007/BF00026656, 1986a.

Chivas, A. R., De Deckker, P. and Shelley, J. M. G.: Magnesium content of non-marine ostracod shells: A new
palaeosalinometer and palaeothermometer, Palaeogeogr. Palaeoclimatol. Palaeoecol., 54(1-4), 43-61, doi:10.1016/0031-0182(86)90117-3, 1986b.

Chunlian, L., Fürsich, F. T., Jie, W., Yixin, D., Tingting, Y. and Jian, Y.: Late Quaternary palaeoenvironmental changes documented by microfaunas and shell stable isotopes in the southern Pearl River Delta plain, J. Palaeogreography, 2(4), 344–361, doi:10.3724/SP.J.1261.2013.00035, 2013.

365 CMEMS: Global ocean physics reanalysis, [online] Available from: https://marine.copernicus.eu/ (Accessed 1 January 2020), 2020.

Cohen, A. L., Layne, G. D., Hart, S. R. and Lobel, P. S.: Implications for the paleotemperature proxy a, Paleoceanography, 16(1), 20–26, 2001.

Cohen, A. L., Owens, K. E., Layne, G. D. and Shimizu, N.: The effect of algal symbionts on the accuracy of Sr/Ca paleotemperatures from coral, Science, 296(5566), 331–333, doi:10.1126/science.1069330, 2002.

Corrège, T. and De Deckker, P.: Faunal and geochemical evidence for changes in intermediate water temperature and salinity in the western Coral Sea (northeast Australia) during the Late Quaternary, Palaeogeogr. Palaeoclimatol. Palaeoecol., 131(3– 4), 183–205, doi:10.1016/S0031-0182(97)00003-5, 1997.

Cronin, T. M., Dwyer, G. S., Baker, P. A., Rodriguez-Lazaro, J. and Briggs, W. M.: Deep-sea ostracode shell chemistry (Mg:Ca

- 375 ratios) and late Quaternary Arctic Ocean history, in Late Quaternary Palaeoceanography of North Atlantic Margins, edited by J. T. Andrews, pp. 117–134, Geological Society, Special Publications., 1996.
  Cronin, T. M., Dwyer, G. S., Kamiya, T., Schwede, S. and Willard, D. A.: Medieval Warm Period, Little Ice Age and 20th century temperature variability from Chesapeake Bay, Glob. Planet. Change, 36(1–2), 17–29, doi:10.1016/S0921-8181(02)00161-3, 2003.
- 380 Cronin, T. M., Kamiya, T., Dwyer, G. S., Belkin, H., Vann, C. D., Schwede, S. and Wagner, R.: Ecology and shell chemistry of Loxoconcha matagordensis, Palaeogeogr. Palaeoclimatol. Palaeoecol., 225(1–4), 14–67, doi:10.1016/j.palaeo.2005.05.022, 2005a.

Cronin, T. M., Thunell, R., Dwyer, G. S., Saenger, C., Mann, M. E., Vann, C. and Seal, I. R.: Multiproxy evidence of Holocene climate variability from estuarine sediments, eastern North America, Paleoceanography, 20(4), doi:10.1029/2005PA001145,

385 2005b.

Day, J., Yanez-Arancibia, A., Kemp, W. M. and Crump, B.: Introduction to Estuarine ecology, in Estuarine ecology, edited by J. Day, B. Crump, W. M. Kemp, and A. Yanez-Arancibia, pp. 1–18, John Wiley and Sons., 2013.

De Deckker, P. and Forester, R. M.: The use of ostracods to reconstruct continental palaeoenvironmental records, in Ostracoda in the Earth Sciences, edited by P. De Deckker, J. P. Colin, and J. P. Peypouquet, pp. 175–199, Elsevier., 1988.

390 Dettman, D. L. and Dwyer, G. S.: The Calibration of Environmental Controls on Elemental Ratios in Ostracod Shell Calcite :





A Critical Assessment, in Developments in Quaternary Sciences, vol. 17, pp. 145–163, Elsevier B.V., 2012. Drévillon, M., Régnier, C., Lellouche, J.-M., Garric, G., Bricaud, C. and Hernandez, O.: Quality Information Document for Global Ocean Reanalysis Products, GLOBAL-REANALYSIS- PHY-001-030. [online] Available from: http://marine.copernicus.eu/services-portfolio/access-to-

- products/?option=com\_csw&view=details&product\_id=GLOBAL\_REANALYSIS\_PHY\_001\_030, 2018.
   Dwyer, G. S., Cronin, T. M., Baker, P. A. and Raymo, E.: North Atlantic deepwater temperature change during late Pliocene and late Quaternary climatic cycles, Science, 270(2), 1347–1351, 1995.
   Dwyer, G. S., Cronin, T. M. and Baker, P. A.: Trace elements in marine ostracodes, in The Ostracoda: Applications in Quaternary Research, vol. 131, edited by J. A. Holmes and A. R. Chivas, pp. 205–225., 2002.
- Elmore, A. C., Sosdian, S., Rosenthal, Y. and Wright, J. D.: A global evaluation of temperature and carbonate ion control on Mg/Ca ratios of ostracoda genus Krithe, Geochem. Geophys. Geosyst., 13(9), 1–20, doi:10.1029/2012GC004073, 2012.
  EPD: Marine water quality in Hong Kong in 2018, Hong Kong SAR. [online] Available from: http://wqrc.epd.gov.hk/pdf/water-quality/annual-report/MarineReport2015eng.pdf, 2018.
  Farmer, J. R., Cronin, T. M. and Dwyer, G. S.: Ostracode Mg/Ca paleothermometry in the North Atlantic and Arctic oceans:
- Evaluation of a carbonate ion effect, Paleoceanography, 27(2), doi:10.1029/2012PA002305, 2012.
  Guo, X., Cai, W. J., Zhai, W., Dai, M., Wang, Y. and Chen, B.: Seasonal variations in the inorganic carbon system in the Pearl River (Zhujiang) estuary, Cont. Shelf Res., 28(12), 1424–1434, doi:10.1016/j.csr.2007.07.011, 2008.
  Hathorne, E. C., Gagnon, A., Felis, T., Adkins, J., Asami, R., Boer, W., Caillon, N., Case, D., Cobb, K. M., Douville, E., Demenocal, P., Eisenhauer, A., Garbe-Schönberg, D., Geibert, W., Goldstein, S., Hughen, K., Inoue, M., Kawahata, H.,
- 410 Kölling, M., Cornec, F. L., Linsley, B. K., McGregor, H. V., Montagna, P., Nurhati, I. S., Quinn, T. M., Raddatz, J., Rebaubier, H., Robinson, L., Sadekov, A., Sherrell, R., Sinclair, D., Tudhope, A. W., Wei, G., Wong, H., Wu, H. C. and You, C. F.: Interlaboratory study for coral Sr/Ca and other element/Ca ratio measurements, Geochem. Geophys. Geosyst., 14(9), 3730– 3750, doi:10.1002/ggge.20230, 2013.

Hendy, E. J., Gagan, M. K., Alibert, C. A., Mcculloch, M. T., Lough, J. M. and Isdale, P. J.: Abrupt Decrease in Tropical
Pacific Sea Surface Salinity at End of Little Ice Age, Science, 295(5559), 1511–1514, 2002.

Holmes, J. A.: Sample-size implications of the trace-element variability of ostracod shells, Geochim. Cosmochim. Acta, 72, 2934–2945, doi:10.1016/j.gca.2008.03.020, 2008.

Holmes, J. A. and De Deckker, P.: The Chemical Composition of Ostracod Shells: Applications in Quaternary Palaeoclimatology, in Developments in Quaternary Science, vol. 17, pp. 131–143, Elsevier B.V., 2012.

Holmes, J. A. and De Deckker, P.: Trace-element and stable-isotope composition of the *Cyprideis torosa* (Crustacea , Ostracoda) shell, J. Micropalaeontology, (2), doi:10.1144/jmpaleo2015-024, 2016.
Hong, Y., Yasuhara, M., Iwatani, H. and Mamo, B.: Baseline for ostracod-based northwestern Pacific and Indo-Pacific shallow-marine paleoenvironmental reconstructions: ecological modeling of species distributions, Biogeosciences, 16(2), 585–604, doi:10.5194/bg-16-585-2019, 2019.





- Ingram, C.: Palaeoecology and geochemistry of Shallow Marine ostracoda from the Sand Hole Formation, inner silver pit, southern North Sea, Quat. Sci. Rev., 17(9–10), 913–929, doi:10.1016/S0277-3791(98)00025-0, 1998.
  Inoue, M., Nohara, M., Okai, T., Suzuki, A. and Kawahata, H.: Concentrations of Trace Elements in Carbonate Reference Materials Coral JCp-1 and Giant Clam JCt-1 by Inductively Coupled Plasma-Mass Spectrometry, Geostand. Geoanalytical Res., 28(3), 411–416, doi:10.1111/j.1751-908X.2004.tb00759.x, 2004.
- 430 Irizuki, T., Matsubara, T. and Matsumoto, H.: Middle Pleistocene Ostracoda from the Takatsukayama member of the Meimi Formation, Hyogo prefecture, western Japan: significance of the occurrence of *Sinocytheridea impressa*, Paleontol. Res., 9(1), 37–54, doi:10.2517/prpsj.9.37, 2005.

Ishizaki, K. and Kato, M.: The basin development of the Diluvium Furuya Mud Basin, Shizuoka prefecture, Japan, based on faunal analysis of fossil ostracodes, Prog. Micropaleontol., 262(77), 1976.

- Ito, E., De Deckker, P. and Eggins, S. M.: Ostracodes and their shell chemistry: Implications for paleohydrologic and paleoclimatologic applications, Paleontol. Soc. Pap., 9, 119–152, doi:10.1017/s1089332600002187, 2003.
  Kamiya, T.: Contrasting population ecology of two species of *Loxoconcha* (Ostracoda, Crustacea) in recent *Zostera* (Eelgrass) beds: Adaptive Differences between Phytal and Bottom-Dwelling species, Micropaleontology, 34(4), 316–331, doi:10.2307/1485600, 1988.
- 440 Kondo, H., Toyofuku, T. and Ikeya, N.: Mg/Ca ratios in the shells of cultured specimens and natural populations of the marine ostracode *Xestoleberis hanaii* (Crustacea), Palaeogeogr. Palaeoclimatol. Palaeoecol., 225(1–4), 3–13, doi:10.1016/j.palaeo.2004.05.026, 2005.

Lea, D. W.: Elemental and isotopic proxies of marine temperatures, Ocean. Mar. geochemistry Treatise Geochemistry, 6, 365–390, doi:http://dx.doi.org/10.1016/B0-08-043751-6/06114-4, 2003.

Lea, D. W. and Boyle, E.: Barium content of benthic foraminifera controlled by bottom-water composition, Nature, 338(6218), 751–753, doi:10.1038/338751a0, 1989.
Linsley, B. K., Wellington, G. M., Schrag, D. P., Matthews, H. E., Owen, T., Meier, R. and Nifio-southem, E.: Decadal Sea

Surface Temperature Variability in the Subtropical South Pacific from 1726 to 1997 A.D., Science, 290(5494), 1145–1148, 2000.

450 Martin, P. A., Lea, D. W., Rosenthal, Y., Shackleton, N. J., Sarnthein, M. and Papenfuss, T.: Quaternary deep sea temperature histories derived from benthic foraminiferal Mg/Ca, Earth Planet. Sci. Lett., 198(1–2), 193–209, doi:10.1016/S0012-821X(02)00472-7, 2002.

Millero, F. J.: Physicochemical controls on seawater, in The oceans and marine geochemistry, edited by H. Elderfield, pp. 1–22, Elsevier., 2006.

455 Morton, B. and Wu, S. S.: The hydrology of the coastal waters of Hong Kong, Environ. Res., 10(3), 319–347, doi:10.1016/0013-9351(75)90029-8, 1975.

Niu, J.: Precipitation in the Pearl River basin, South China: Scaling, regional patterns, and influence of large-scale climate anomalies, Stoch. Environ. Res. Risk Assess., 27(5), 1253–1268, doi:10.1007/s00477-012-0661-2, 2013.





Okai, T., Suzuki, A., Kawahata, H., Terashima, S. and Imai, N.: Preparation of a new Geological Survey of Japan geochemical
reference material: Coral JCp-1, Geostand. Newsl., 26(1), 95–99, doi:10.1111/j.1751-908X.2002.tb00627.x, 2002.
Patra, S., Liu, C. Q., Wang, F. S., Li, S. L. and Wang, B. L.: Behavior of major and minor elements in a temperate river estuary to the coastal sea, Int. J. Environ. Sci. Technol., 9(4), 647–654, doi:10.1007/s13762-012-0097-8, 2012.
Roberts, L. R., Holmes, J. A. and Horne, D. J.: Tracking the seasonal calcification of Cyprideis torosa (Crustacea, Ostracoda)

using Mg/Ca-inferred temperatures, and its implications for palaeotemperature reconstruction, Mar. Micropaleontol.,

- 156(October 2019), 101838, doi:10.1016/j.marmicro.2020.101838, 2020.
   Rodríguez, M.: Trace-elements in ostracods as proxy of environmental parameters in shallow marine waters (in review), Ph.D. thesis. The University of Hong Kong, Hong Kong SAR, China., 2020.
   Rosenthal, Y., Lear, C. H., Oppo, D. W. and Linsley, B. K.: Temperature and carbonate ion effects on Mg/Ca and Sr/Ca ratios in benthic foraminifera: Aragonitic species Hoeglundina elegans, Paleoceanography, doi:10.1029/2005PA001158, 2006.
- Sinclair, D. J., Kinsley, L. P. J. and Mcculloch, M. T.: High resolution analysis of trace elements in corals by laser ablation ICP-MS, Geochim. Cosmochim. Acta, 62(11), 1889–1901, doi:10.1016/S0016-7037(98)00112-4, 1998.
  Snedden, G., Cable, J. and Kjerfve, B.: Estuarine Geomorphology and coastal hydrology, in Estuarine ecology, edited by J. Day, B. Crump, M. Kemp, and A. Yanez-Arancibia, pp. 19–38, John Wiley and Sons., 2013.
  Tanaka, G., Komatsu, T., Saito, Y., Phong, D., Lan, Q., Nguyen, D. P. and Vu, Q. L.: Temporal changes in ostracod
- 475 assemblages during the past 10,000years associated with the evolution of the Red River delta system, northeastern Vietnam, Mar. Micropaleontol., 81(3–4), 77–87, doi:10.1016/j.marmicro.2011.08.001, 2011.
  Wang, H. E., Zhang, H., Cao, M. and Horne, D. J.: Holocene Ostracods from the Hang Hau Formation in Lei Yue Mun, Hong Kong, and their palaeoenvironmental implications, Alcheringa An Australas. J. Palaeontol., 1–11, doi:10.1080/03115518.2018.1511830, 2018.
- Whatley, R. and Zhao, Q.: A revision of Brady's 1869 study of the Ostracoda of Hong Kong, J. Micropalaeontology, 7(1), 21–29, doi:10.1144/jm.7.1.21, 1988.
  Yamada, K., Irizuki, T., Ikehara, K. and Okamura, K.: Calibration of past water temperature in the Sea of Japan based on Mg/Ca of ostracode shells of two shallow marine species in the genus *Cytheropteron*, Palaeogeogr. Palaeoclimatol. Palaeoecol., 410, 244–254, doi:10.1016/j.palaeo.2014.05.042, 2014.
- 485 Yu, J. and Elderfield, H.: Mg/Ca in the benthic foraminifera *Cibicidoides wuellerstorfi* and *Cibicidoides mundulus*: Temperature versus carbonate ion saturation, Earth Planet. Sci. Lett., 276(1–2), 129–139, doi:10.1016/j.epsl.2008.09.015, 2008.

Yu, J., Day, J., Greaves, M. and Elderfield, H.: Determination of multiple element/calcium ratios in foraminiferal calcite by quadrupole ICP-MS, Geochem. Geophys. Geosyst., 6(8), 1–9, doi:10.1029/2005GC000964, 2005.

490 Yuan, X. C., Yin, K., Cai, W. J., Ho, A. Y., Xu, J. and Harrison, P. J.: Influence of seasonal monsoons on net community production and CO<sub>2</sub> in subtropical Hong Kong coastal waters, Biogeosciences, 8(2), 289–300, doi:10.5194/bg-8-289-2011, 2011.





Zhang, Q., Li, J., Singh, V. P., Xu, C. Y. and Deng, J.: Influence of ENSO on precipitation in the East River basin, south China, J. Geophys. Res. Atmos., 118(5), 2207–2219, doi:10.1002/jgrd.50279, 2013.

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