- 1 Divergent climate feedbacks on in-winter wheat the growing period and the
- 2 dormancy periods to as affected by sowing date shift of winter wheat in the North
- 3 China Plain
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- 20 Abstracts: Crop phenology exerts measurable impacts on soil surface properties,
- 21 biophysical processes, and climate feedbacks, particularly at local/regional scales. The
- 22 land cover and management changes have strong feedbacks to climate through surface
- 23 biophysical and biochemical processes. Agricultural phenology dynamic exerted

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measurable impacts on land surface properties, biophysical process and climate feedback 24 25 in particular times at local/regional scale. Nevertheless, the response of surface biophysical processes to climate feedbacks as affected by sowing date in winter wheat 26 croplands has been overlooked, especially during winter dormancyBut the responses of 27 28 climate feedback through surface biophysical process to sowing date shift in the winter wheat ecosystem have been overlooked, especially at winter dormancy period. The 29 dynamics of leaf area index (LAI), surface energy balance and canopy temperature (T_c) 30 were simulated by modified SiBcrop model under two sowing date scenarios (Early 31 32 Sowing: EP; Late Sowing: LP) at 10 stations in the North China Plain. The results showed that the SiBcrop with a modified crop phenology scheme better-well simulated 33 the seasonal dynamic of LAI, T_c, phenology, and surface heat fluxes. Earlier sowing date 34 had higher LAI with earlier development than later sowing date. But the response of T_c to 35 sowing date exhibited opposite patterns during the dormancy and active growth periods: 36 EP led to higher T_c (0.05 K) than LP in the dormancy period and lower T_c (-0.2K) in the 37 growth period. The highest difference (0.6 K) between EP and LP happened at the time 38 39 when wheat was sown in EP but wasn't in LP. The higher LAI captured more net 40 radiation with lower surface albedo for warming effect, whist but partitioned more energy into latent heat flux withsurface energy partitioning exerted cooling effect. The 41 climate feedback of sowing date, which was more obvious in winter in the northern areas 42 and in the growing period in the southern areas, was determined by the The relative 43 contributions of albedo-radiative process and partitioning-non-radiative process. 44 determined the climate effect of sowing date shift. The spatial pattern of the climate 45

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- 46 response to sowing date was influence by precipitation and air temperature. _ The study
- 47 <u>climate.</u> <u>-climate effects of the sowing date shift in winter dormancy period are worthy</u>
- 48 Key words: sowing date, canopy temperature, phenology, leaf area index, winter wheat,
- 49 land surface model, North China Plain

1. Introduction

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cover and management changes have strong feedbacks with climate through surface 53 biophysical and biochemical processes (Mahmood et al. 2014). Cropland surface 54 characteristic had been and will continue to be changed through agricultural crop 55 management, such as cropping system (Jeong et al. 2014; Cui et al. 2018), sowing date 56 and phenology shifts (Sacks et al. 2011; Richardson et al. 2013), and cultivars 57 selectionbio geoengineering (Seneviratne et al. 2018), to keep high yield under climate 58 change condition. The changed croplandsurface properties in farmland further generate feedback to regional climate through surface energy partitioning and albedo (a) 60 mechanisms (Cooley et al. 2005; Zhang et al. 2015). It is important to quantify the 61 climate feedback of crop phenology shift _for regional climate prediction and agriculture 62 63 sustainable development. There are evidences that crop phenology has been shifts shifted substantially in the 64 major cultivation areas worldwide (Sacks and Kucharik 2011; Tao et al. 2012; Tao et al. 65 66 2014; Liu et al. 2017). In the North China Plain (NCP), the dates of seedingsowing, 67 dormancy, re-greeninggreen up, anthesis, and maturity in wheat system were changed by 1.5, 1.5, -1.1, -2.7, and -1.4 days/decade (a positive value indicates delay and a negative 68 value indicates advance), respectively (Xiao et al. 2013). The vegetative stages (including 69 periods from dormancy to greenupre-greening, re-greeninggreenup to anthesis) was-were 70 shortened and reproductive stage was prolonged (Xiao et al. 2013). The main 71

contributors including climate change and crop management. Global warming

induced-higher temperature resulted in longer photosynthetic-active period but faster

Land-atmosphere interactions are key components of the climate system. The land

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development rate and shorter growth stages. Crop management, including sowing date 74 75 adjustment and varietal change, reduced the lengths of vegetative stage, but increased the length of reproductive stage (Liu et al. 2010; Liu et al. 2018). The management induced 76 phenology change dynamics is are intended to increase yieldbeneficial for high yielding. 77 The strategies adapting to warmer environment include adopting cultivars with higher 78 accumulated growing degree days (GDD) and later planting. The prolonged grain-filling 79 period of winter wheat benefits the accumulation of carbohydrates organic matter in grain 80 (Reynolds et al. 2012; Liu et al. 2018), and the adjusted sowing date reduces the risks 81 82 such as insect and viral infection, adverse meteorological conditions, and soil water depletion , et al. (Sacks et al. 2010). Model simulation indicated that yield increase of 83 winter wheat was benefitted from cultivars renewal by 12.2-22.6% and fertilization 84 management by 2.1-3.6%; climate change damaged yield by -15.0% for rain-fed type, in 85 the NCP (Xiao et al. 2014). 86 The crop phenology-shifts change affects the seasonal rhythm of crop development 87 and affect theof surface greenness coverage of land surface and energy and water 88 89 exchanges in the boundary layer. For example, maize growth duration prolonged and 90 reached maturity and senesced a couple of weeks later, and the maximum change can reach 47 W m⁻² and -20 W m⁻² for latent heat flux (LH) and sensible heat flux (SH), 91 respectively, when the NDVI is increased by 0.1 in the Agro-IBIS model (Bagley et al. 92

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2015). Earlier planting date and longer grain-filling period increased (decreased) the LH

(SH) by 0.3 (0.2)3 W m⁻², decreased SH by 2.5 W m⁻²-mm/year in June and enhanced the

net radiation (R_n) by 1.2 W m⁻² in October by reducing the interval time from maturity to

harvest in American maize Corn belt (Sacks and Kucharik 2011). The change of surface

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coverage also shows a certainn regional climate feedback. The increased spring surface greenness at farmland, due to the advanced re-greening stage of winter wheat (Xiao et al. 2013; Liu et al. 2017), significantly impacted the patterns of *LH* and *SH* and then the changes of moderate to light rainfall (Zhang et al. 2015). Harvest shifted the key influence factors of the radiative balance and evaporative fraction from leaf area and soil-atmosphere temperature difference to soil moisture in U.S. winter wheat (Bagley et al. 2017), and a shift in radiative forcing with the potential to warm the atmosphere by warming future atmosphere by 1~1.4 °C through decreasing declining *LH*evapotranspiration in the NCP (Cho et al. 2014). So, the influence of phenology on climate feedback through surface biophysical process at local/regional scale is worthy of

further studies (Liu et al. 2017).

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Despite previous studies showed the critical role of crop phenology-dynamic to-in surface energy and water balance, there is an important potential sensitive period that has been ignored in the winter wheat system. During the dormancy period in winter, aboveground canopy of winter wheat remained constant for more than 2 months (Xiao et al. 2013). In view of the close relationships between surface biophysical processes and aboveground canopysurface characteristic (Boisier et al. 2012; Chen et al. 2015; Liu et al. 2017), the length from sowing date to start of dormancy would be the determinant factor to surface biophysical process in winter where winter wheat wildly widely distributed, such as NCP, Pacific Northwest (Wuest 2010) and Southern Great Plains of USA (Bagley et al. 2017), Australia, and numerous countries surrounding the Mediterranean Sea (Mahdi et al. 1998; Schillinger 2011). Compared with climate feedback of other phenology dynamics, such as earlier re-greening stage (Xiao et al. 2013; Zhang et al.

2013), longer reproductive period (Sacks and Kucharik 2011) and inter-cropping period (Cho et al. 2014; Bagley et al. 2017), the climate feedback of sowing date emerges gradually with crop development. Particularly, winter wheat grows faster in early stage and slower as winter approaches, smaller change in sowing date could lead to larger and longer climate feedback in dormancy period, the effects of sowing date on land surface characteristic in dormancy period of winter wheat and other winter crops are relatively indirect and the effects last longer. Recognition of the impacts of sowing date on land surface characteristics and climate feedback would be beneficial to the understanding of human influence on climate change. Therefore, it is necessary to investigate whether dormancy period of winter wheat is sensitive to sowing date. And if so, how sensitivities are surface biophysical process and climate effect?

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2. Data and methods

2.1. Study stations

The NCP, with an area of 4×10^5 km², is the largest winter wheat production region in China, including Hebei, Henan, Shandong, Jiangsu, and Anhui provinces, and Beijing and Tianjin municipalities (Fig.1)._—Summer maize - winter wheat rotation is the main cropping system,—, except Anhui and Jiangsu where winter wheat-rice rotation system is dominated. The satellite data showed a high cropland density above 70% with flat and relatively homogeneous agricultural practices (Liu et al. 2005; Ho et al. 2012). The soil type is classified as sandy loam according to the seven soil textures in the model (Sellers et al. 1996). Two stations with surface fluxes were used for model calibration (Fig.1, blue triangles). Ten randomly distributed stations with complete meteorology and phenology

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with complete meteorology and phenology information were selected for this study (Fig.1, green dots). The natural conditions and agricultural production level of the selected stations are typical to the NCP. The stations maintain good records on both meteorological and winter wheat phenology data since 1981.

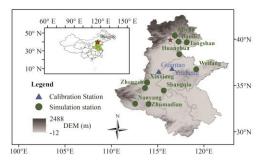


Fig.1 Distribution map of the study area and observation sites.

The 30 m resolution digital elevation model—(DEM), provided by the GlobeLand30 in 2010, and the administrative map were downloaded from the National Catalogue Service For Geographic Information.

2.2 Data

2.2.1 Meteorology

The quality-controlled meteorological data, including air temperature (T_a) , precipitation (P), atmosphere pressure, relative humidity, and wind speed, was obtained from the Chinese Meteorological Administration. Summer monsoon climate dominates the region with an uneven distribution of annual precipitation (Table 1). In the 1980-2012, the average annual P at the selected stations ranged between 550-990 mm, mainly happened in summer. The mean yearly T_a varied between 11-15 °C. In the growing

season of winter wheat (11-12 and 1-6 month), the T_a varied between 7-11 $^{\circ}$ C among stations and P ranged between 170-420 mm, which is consistent with the average climatic conditions in the NCP (A et al. 2016). Climatological mean T_a and accumulated P during the wheat growth period were calculated in the 10 stations and were linearly regressed with the simulated differences between scenarios. Meteorological data was also used to drive the model. Meteorological data was also used to drive the model in the selected 10 stations (Fig.1).

Table 1 Climate conditions of the selected stations in 1980-2012

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														Average
	Station	Jan.	Feb.	Mar.	Apr.	May	Jun.	Jul.	Aug.	Sept.	Oct.	Nov.	Dec.	Wheat
														Season
	Miyun	-5.9	-2.2	4.8	13.5	19.6	24	25.8	24.5	19.3	12	3	-3.7	6.6
	Baodi	-5	-1.4	5.3	13.6	19.5	24	26.1	24.8	19.8	12.7	3.7	-2.7	7.1
	Tangshan	-4.8	-1.3	5.1	13.4	19.4	23.7	25.9	25	20.3	13	4.1	-2.5	7.1
	Huanghua	-3.4	-0.2	5.9	14.1	20.3	25	26.9	25.8	21.2	14.2	5.5	-1.2	8.3
6	Weifang	-2.8	0.1	5.8	13.2	19.2	23.9	26.2	25.2	20.7	14.4	6.4	-0.3	8.2
T_a (°C)	Xinxiang	0	3.3	8.7	15.8	21.2	25.8	27	25.9	21.3	15.3	7.9	1.8	10.6
	Zhengzhou	0.5	3.5	8.7	16	21.5	26	27.1	25.7	21.2	15.5	8.4	2.5	10.9
	Shangqiu	0.1	3.1	8.3	15.1	20.6	25.4	26.9	25.7	21.1	15.3	8.1	2	10.3
	Nanyang	1.6	4.4	9.1	15.8	21.2	25.5	27	26	21.7	16.1	9.4	3.5	11.3
	Zhumadian	1.5	4.2	9	15.7	21.2	25.7	27.2	25.9	21.6	16.3	9.6	3.6	11.3
	Miyun	2.2	4	9.7	19.9	43.1	86.7	180.7	172.6	62.9	25.5	9.4	2.3	177.3
	Baodi	2.7	3.7	9	20.1	36.2	82.4	169.7	142.6	49.7	27.5	10.1	3.6	167.8
n)	Tangshan	3.5	4.1	9.4	22.4	47	83.2	169.7	154.3	50.8	28.2	9.5	3.4	182.5
P (mm)	Huanghua	3.2	5.5	10.1	21.3	42.8	84.2	177.2	111.6	41.5	31	11.9	3.5	182.5
	Weifang	6	10.3	14.9	24.5	45.4	80	136.5	132.1	56.1	32.8	18.8	8.9	208.8
	Xinxiang	4.6	7.1	19.2	25	49.9	65	150	119.5	59.8	32	14.9	5	190.7

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Zhengzhou	9.6	12.4	27.1	30.9	63.6	67.8	146.6	134.7	75.6	40.5	21.1	9.1	241.6
Shangqiu	14.3	16.3	29.3	33	65	85.2	166.8	144.8	68.5	38.2	23.4	12.7	279.2
Nanyang	13.2	15.6	35.2	41.7	78.8	124.5	183.7	131.7	76.3	51.1	30	12.8	351.8
Zhumadian	21.9	24.8	51	50.9	93	128.6	227.7	176.3	98.2	63.9	35.2	18.5	423.9

 T_a means air temperature, and P means precipitation.

2.2.2 Verification data

To verify the applicability of the model, surface flux data was collected from Yucheng and Guantao stations (Fig.1; Table 2). The two stations used the same eddy covariance instruments to measure the surface latent heat flux (LI7500, LI-COR Inc., Lincoln, NE, USA) and sensible heat flux (CSAT-3, Campbell Scientific Inc., Logan, UT, USA), but at different heights (Yucheng:3.3 m; Guantao: 15.6 m). The post-processing software (Yucheng: Eddypro; Guantao: EdiRe) was used to process the raw data such as spike detection, lag correction of H₂O/CO₂ relative to the vertical wind component, sonic virtual temperature correction, coordinating rotation using the planar fit method, corrections for density fluctuation (WPL-correction), and frequency response correction (Liu et al. 2011). The REddyProc was used for gap-filling by method of the look-up table and the mean diurnal variations method (Falge et al. 2001; Wutzler et al. 2018). More details could be referred to (Lei et al. 2010; Liu et al. 2013). Totally 10 complete winter wheat season flux data were used to validate the model (Table 2).

The meteorology conditions were also synchronously measured during flux observation (Table 2). The measurement included T_a , P, atmosphere pressure, relative humidity, wind speed, and sunshine. These data was the inputs of the model. According to the T_a and P, the meteorological conditions were similar between the 10 stations for simulation and the two stations for calibration. More variables were observed at Yucheng

station, such as wheat phenology and leaf area index (LAI) and canopy temperature (T_c). The observed durations of phenology, LAI, and fluxes at Yucheng station were in 2003-2006, 2004-2006, and 2003-2010, respectively.

Table 2 General information about model verification data

		Wheat growing sea	ison		
Station	Period	$T_a\left(\mathbf{K}^{\underline{\circ}}\mathbf{C}\right)$	P (mm)	Measured variables	
Yucheng	2003-2010	282.15 <u>9</u>	226.7	Meteorology, Phenology, LAI, LH, SH, Tc	
Guantao	2008-2010	282.75 <u>9.6</u>	134.4	Meteorology, LH, SH	

 T_a means air temperature, P means precipitation, LAI means leaf area index (m² m⁻²), LH means latent heat flux (W m⁻²), SH means sensible heat flux (W m⁻²). T_c means the simulated canopy temperature ($^{\circ}$ CK).

2.2.3 Phenology of winter wheat

The phenology information was <u>obtained from China agro-meteorological experiment stations</u> manually recorded and available in the period of 1981-2009, except for 2003 at Zhumadian and 1986 and 1988 at Miyun station (Table 3). Phenological statistics showed that the sowing time of winter wheat is generally between DOY (Day Of Year) 270-290 (early and middle October) in the NCP. After sowing, it generally takes about 6-10 days for germination. Winter wheat dormancy stage generally begins in DOY 330-360 (December) and ends in DOY 40-70 (late February and early March), and reaches maturity in DOY 150-160(mid-June). The standard deviation shows that the inter-annual fluctuations of dormant and re-greening period is larger, and harvest period is relatively stable.

For the past 30 years, winter wheat phenology at some stations showed a significant linear trend (Table 4). The sowing and germination periods were significantly delayed in 4 out of 10 stations, and the trend in the dormant and re-greening period was not obvious. Winter wheat matured significantly earlier at five stations. Generally, the autumn and winter phenophases, including sowing, germination and dormancy, are mainly delayed, while spring and summer phenophases, including re-greening and maturity, are primarily advanced. According to the fitting coefficient (a), the duration were was changed by 5.7, 8.1, 4.9, -3.5, and -5.5 in the period of 1981-2009, respectively, for the stages of sowing, germination, dormancy, re-greening and maturity of winter wheat. These results were consistent with our previous studies (Tao et al. 2012; Xiao et al. 2013; Xiao et al. 2015), and indicating that the selected stations were good representation of the NCP.

Table 3 General information on the phenology of winter wheat in the selected stations (unit: DOY)

Station	Period	Sowing	Germination	Dormancy	Re-greening	Maturity	
Miyun	1981-2009	275.52±7.55	284.96±9.03	331.93±6.41	73.59±15.1	168.26±3.46	
Baodi	1981-2009	272.83±4.33	281.55±5.5	335.62±6.86	59.19±42.72	165.97±2.57	
	1981-2009						
Tangshan	(except 2003)	271.86±4.83	279.59±6.04	335.55±6.6	66.62±7.98	169.97±3.23	
Huanghua	1981-2009	274.17±7.83	280.32±7.03	340.38±8.65	62.45±6.56	157.14±3.25	
Weifang	1981-2009	274.1±5.75	284.62±17.93	343.72±7.76	59.59±7.29	160.41±3.42	
Xinxiang	1981-2009	283.59±4.21	291.64±5.14	351.9±10.56	47.55±7.16	152.03±3.3	
Zhengzhou	1981-2009	289.76±5.67	298.45±6.65	360.5±14.08	44.21±7.43	151.34±3.88	

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Shangqiu	1981-2009	287.59±4.07	295.31±4.79	359.21±32.4	47.03±6.43	151.59±2.99
Nanyang	1981-2009	297.21±7.81	306.83±9.03	7.54±14.64	48.22±8.9	149.21±4.99
	1981-2009					
Zhumadian	(except 1986,	289.54±9.33	298.29±11.11	5.46±10.35	49.15±6.84	146.21±4.76
	1988)					

the data was shown in average \pm standard deviation.

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Table 4 Linear trends in winter wheat phenology

Station	Sow	ing	Germi	nation	Dorm	ancy	Re-gre	eening	Matu	ırity	
Station	a	p	a	p	a	p	a	p	a	p	
Miyun	0.62	0.00	0.69	0.00	0.17	0.27	-0.51	0.15	-0.20	0.01	
Baodi	0.31	0.00	0.41	0.00	0.14	0.36	-0.67	0.52	-0.05	0.35	
Tangshan	0.41	0.00	0.51	0.00	0.43	0.00	-0.29	0.11	-0.20	0.00	
Huanghua	0.18	0.31	0.17	0.31	0.38	0.05	-0.07	0.64	-0.13	0.07	7
Weifang	0.20	0.11	0.61	0.13	0.11	0.55	0.14	0.38	-0.12	0.11	
Xinxiang	0.07	0.46	0.12	0.34	0.27	0.26	-0.16	0.33	-0.12	0.10	
Zhengzhou	-0.16	0.21	-0.21	0.17	-0.28	0.41	0.11	0.52	-0.25	0.00	(
Shangqiu	0.03	0.77	0.04	0.68	0.39	0.59	0.10	0.51	-0.07	0.28	
Nanyang	-0.18	0.30	-0.11	0.60	-0.13	0.71	0.12	0.60	-0.38	0.00	(
Zhumadian	0.49	0.02	0.56	0.02	0.21	0.37	0.02	0.89	-0.36	0.00	

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2.3 Methods

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2.3.1 Model calibration and verification

a was the coefficient of linear fitting equation (d/year); p was the significance level;

bolded number means p < 0.05.

The SiBcrop model was selected in this study. SiBcrop is a process-based land surface model adapted from the Simple Biosphere model version 3 (Lokupitiya et al. 2009). The SiB series models (version 1, 2, 3 refers to SiB1, SiB2, SiB3, respectively) are widely adopted land surface models for computing surface energy, water, momentum and CO2 exchange in the boundary layer. The SiBcrop version added the crop-specific submodels simulation of maize, soybean, winter and spring wheats, which eropping system (Lokupitiya et al. 2009). The crop specific submodel was simple and detailed enough in predicting LAI (Lokupitiya et al. 2009). replaces remotely sensed NDVI information by simulated LAI and the fraction of photosynthetically active radiation. The submodel replaces remotely-sensed NDVI information by simulated LAI. SiBcrop simulated fast response processes that vary sub-hourly such as energy, water, carbon and momentum balance of the canopy and soil, as well as the processes that vary daily such as LAI. Surface energy and water fluxes are calculated at each time step on a grid cell basis according to physiologically based formulations of leaf-level photosynthesis, stomatal conductance and respiration (Farquhar et al. 1980; Collatz et al. 1990).

The model was first modified according to the actual situation of winter wheat in the NCP (Chen et al. 2020). The SiBcrop model was originally calibrated in winter wheat – summer fallow system in which the growth time of wheat is relatively abundant (Lokupitiya et al. 2009). However, the NCP is dominated by winter wheat – summer maize system in which the development of wheat is strictly restricted. There are great differences in the varieties, planting date, growth environment and physiological characteristics of winter wheat between the two systems. The modifications including include: (1) the sowing date was postponed to October from original August. (2)

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带格式的:字体:小四 **带格式的:**字体:小四 **带格式的:**字体:小四 tolerance was reduced to 8°C from original 18°C, above which the seven consecutive days for wheat sowing was were counted. (3) The harsh condition of delayed sowing also the daily growth rate, which was modified from 0.07 to 0.03 g m⁻² when GDD was 105-310 °C d. (4) Wheat grows faster when GDD is 769-1074 °C d with maximum dry weight increased from 8 to 12 g and daily rate enlarged from 0.015 to 0.15 g m⁻². (5) Specific leaf area was changed from 0.02 to 0.025 m² g⁻¹ (Najeeb et al. 2016). (6) A subroutine was added to describe the senescence process of canopy when GDD was larger than 1074 °C d according to Tao et al. (Tao et al. 2009). More details could be referred to Chen et al (2020).

After modifications, the simulated biases were within 10 days for wheat emergency and harvest dates, the determination coefficient, root mean square error, and agreement index between simulated and observed LAI were obviously improved from 0.26, 1.89 m² m⁻², and 0.7 to 0.80, 0.99 m² m⁻², and 0.91, respectively. And they were 0.66, 32.37 W m⁻², and 0.84, respectively, for the simulated *LH* (Chen et al. 2020).

2.3.2 Model simulation

Two simulations with different sowing dates were performed to examine the responses of surface biophysical processes to shifts in sowing dates at the selected 10 stations (Fig.1). The planting date was classified into two scenarios: after DOY 265 (early sowing scenario, EP) and after DOY 275 (late sowing scenario, LP). The early and late sowing scenarios were established by artificially limiting the starting time of the sowing date.—The early sowing scenario means that the sowing will not be allowed until DOY 265. Similarly, the late sowing scenario is only allowed after DOY 275. In both scenarios, wheat was sowed at the seventh consecutive days when temperature ranged \$8~25°C,

which means the real sowing date was seven days later. The SiBcrop model was modified to be more cold tolerance (section 2.3.1), which causing the sowing date was less controlled by temperature. The climate variability among stations has less constraint on sowing date. Our previous study showed that the delayed sowing date of winter wheat was mainly caused by the delayed harvest of maize in the NCP (Xiao et al. 2013). The sowing date in the two scenarios is within the climatological average of the region.

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2.3.3. Methods to relate the surface energy balance components with T_c

The Boisier method (Boisier et al. 2012) was adopted to relate the surface energy balance components with T_c . The energy partitioning of a terrestrial surface is expressed

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Where S_d , L_d , and L_u are the downward short-wave radiation, downward long-wave radiation, and upward long-wave radiation, respectively. In order to have a closed surface energy balance, the residual term R was derived explicitly from the other terms in equation (1), and principally accounts for the soil heat flux and canopy storage flux.

The T_c change simulated by model is affected by both radiative (surface albedo effect) and non-radiative processes (surface energy partitioning effect). In order to separate temperature variation caused by the sole change in absorbed short-wave radiation (radiative process), the following equation (Boisier et al. 2012) was used:

$$\Delta T_c = (\varepsilon \sigma)^{-1/4} \left[\left(L_u + \Delta L_u \right)^{1/4} - L_u^{1/4} \right]$$
 (2)

Where ΔT_c is the anomaly of canopy temperature (K). The σ is Stefan-Boltzmann constant (=5.67×10 -8 W m⁻² K⁻⁴). The ε is surface emissivity (= 1). A disturbance in S_d, 298

 L_d , LH, SH or R can be expressed as ΔL_u by fixing non-perturbed terms using equation (1). More details can be found in Boisier et al. (2012).

3 Results

3.1 SiBcrop simulation accuracy

The SiBcrop model had been modified to improve the simulation accuracy of wheat growth and surface fluxes in our previous study (Chen et al. 2020). After modifications, the simulated biases were within 10 days for wheat emergency and harvest dates, the determination coefficient, root mean square error, and agreement index between simulated and observed LAI were obviously improved from 0.26, 1.89 m²-m², and 0.7 to 0.80, 0.99 m²-m², and 0.91, respectively. And they were 0.66, 32.37 W m²-2, and 0.84, respectively, for the simulated *LH* (Chen et al. 2020).

The simulation error for wheat phenology at Yucheng station was within 10 days (Chen et al. 2020). The sowing time under the two sowing scenarios was further compared with observation at the selected 10 stations. The simulated sowing date was stable, generally around DOY278.66 ± 1.15, and DOY 290.34 ± 2.08 for EP and LP scenario, respectively. The observed phenology fluctuated greatly. Wheat was prone to sow later or early generally due to geographical location at some specific stations. In the EP scenario, the stations in the north had a positive difference (delayed sowing date relative to the actual date) compared to the actual phenological period, whereas the stations in the south had a negative difference (advanced sowing date relative to the actual date), because the stations in the north had earlier sowing date than those in the south. In the LP scenario, the stations in the south were relatively close to the actual

phenology, but the stations near the north had a larger positive difference. Overall, the simulation difference of phenology was within 15 days.—The selected scenario covers the actual situation of winter wheat sowing in the NCP. The comparisons improved the representativeness and reliability of the simulation results.

Table 5 The difference between simulated and observed sowing dates under two scenarios at each station

- Control	Scen	ario		
Station	Early sowing	Late sowing		
Miyun	4.19±7.82	17.48±7.55		
Baodi	6.59±4.62	19.59±4.49		
Tangshan	7.41±4.95	20.38±5.47		
Huanghua	4.41±8.02	16.31±7.55		
Weifang	4.34±5.6	15.86±5.55		
Xinxiang	-5.31±4.24	5.59±4.24		
Zhengzhou	-11.41±5.47	-0.38±5.53		
Shangqiu	-9.34±3.84	1.48±3.85		
Nanyang	-19.07±7.87	-8.41±8.08		
Zhumadian	-11.36±8.91	-0.57±9.07		
All	-2.98±10.96	8.7±11.66		

data was show in average \pm standard deviation.

3.2 Seasonal dynamics of LAI and T_c in scenarios

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Wheat LAI curves for the two sowing dates were obviously not overlapped (Fig.2a). The LAI in the EP scenario was larger with earlier development. With the sowing in the LP scenario, LAI difference between the two scenarios gradually narrowed until the spring of the next year when the disparity increased again (Fig.3a). The LAI difference between two scenarios had a valley after the reproductive period. With the approaching of harvest, the difference gradually decreased to 0.

The LAI difference of winter wheat in two scenarios is mainly attributed to the difference in the accumulation of organic matterbiomass. In the EP scenario, earlier sowing means advanced assimilation process and better temperature conditions, more photosynthetic carbon was produced and distributed into leaf. The impact of sowing time on LAI displayed great dissimilarity among stations (Fig.3a). Based on linear regression, the seasonal average of wheat LAI difference between scenarios was highly related with precipitation in the growth period (LAI anomaly = 0.0011 * P - 0.12, $R^2 = 0.59$). The more precipitation, the greater influence of sowing date on growth. The T_a contributed little to the LAI difference between the two scenarios.

According to the T_c difference between scenarios, the following phenologies of winter wheat were relatively important: sowing date, dormancy date, re-greening date and maturity date. Based on the simulation results, the phenological dates used here as follows: EP sowing date, DOY279; LP sowing date, DOY290; dormancy date, DOY334; re-greening date, DOY59; maturity date, DOY170 (Fig.2a). The T_c difference between scenarios was separated into 4 phases: Phase 1, inter-sowing period, when wheat had been sown in the EP but hadn't in the LP; Phase 2: early growing period, from sowing

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354 re-greening date; Phased 4: late growing period, from re-greening date to maturity date (Fig.2b). 355 The most obvious disparity in T_c between two scenarios occurred in the inter-sowing 356 357 period when wheat had been sown in the EP but hadn't in the LP (Fig.2b). The development of early sown winter wheat resulted in higher T_c, with a peak of up to 0.6 K. 358 The growth of wheat in the LP sharply reduced the warming effect in EP, and eventually 359 the EP scenario had lower temperatures (-0.2K) before entering the dormancy period. The 360 361 temperature change process during this period was relatively consistent across the 362 selected stations (Fig.3b). In the late growing period, the EP had lower temperature (-0.1 K) than LP. In particular, LAI difference varied greatly between statioms (Fig. 3a), T_0 363 difference was relatively stable (Fig.3b). 364 Another special period is the dormancy period, when EP had higher T_c than LP with 365 an average of 0.05 K (Fig.3b). With the start of the re-greening period, the EP T_c was 366 gradually lower than LP T_c and dropped to 0 at the harvest time. The T_c dynamics during 367 368 this period was highly heterogeneous among the stations, varying between -0.25~0.25 K. In the dormancy period, the T_c anomaly between scenarios was significantly affected 369 by the T_a in winter (T_c anomaly = -0.023 * T_a + 0.062, R^2 = 0.6, p = 0.005). The lower the 370 T_a , the bigger the T_c difference, which indicating that the influence of sowing date is 371 more important in northern farmland. The linear relationship between P and T_c difference 372 in winter was not obvious. The linear fitting equation between P and T_c anomaly in the 373 growing period: T_c anomaly = -0.0013 * P + 0.057, R^2 = 0.8, p < 0.001. So mMore 374 375 rainfall increased the T_c anomaly in the growing period. The linear fitting equation

date of LP to dormancy date; Phase 3: dormancy period, from dormancy date to

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between T_a and T_c anomaly in the growing period: T_c anomaly = -0.017 * T_a + 0.2, R^2 = 0.53, p = 0.01. Since the T_c anomaly was negative, the higher the T_a , the greater the T_c anomaly. Considering the low temperature and less precipitation at the northern stations, the high temperature and more precipitation at the southern stations, the climate feedback of sowing date shift was more obvious in winter in the northern areas, and in the growing period in the southern areas.

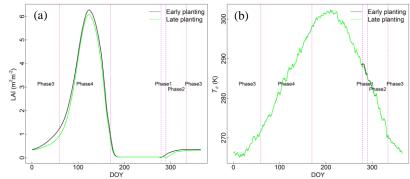


Fig.2 Dynamics of (a) LAI and (b) T_c under two sowing scenarios in winter wheat

growing season

Phase 1: inter-sowing period, when wheat had been sown in the EP but hadn't in the LP; Phase 2: early growing period, from sowing date of LP to dormancy date; Phase 3: dormancy period, from dormancy date to re-greening date; Phased 4: late growing period, from re-greening date to maturity date.

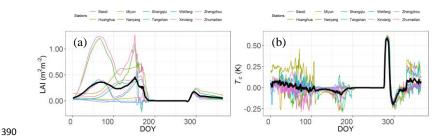


Fig.3 Seasonal differences in (a) LAI and (b) T_c of EP-LP at each station. The average across the stations was shown in bold black line

3.3 Contributions of surface energy balance components to scenario difference in T_c

According to the seasonal dynamics of LAI and T_c , winter wheat growth could not explain the difference in climate effect of sowing time. Specifically, the T_c anomaly between the two scenarios were was reversed between the dormancy (Phase 3December, January, and February) and active growth periods (Phase 2 and Phase 4 other wheat development period with active physiological activity), but with both positive LAI difference (Fig.3). In this section, surface energy balance was used to explain the response of T_c to sowing date.

The flux anomalies of R_n , LH, SH and R were shown in Fig.4a. The EP scenario always maintained higher R_n and LH. Especially winter wheat-covered ground captured more than 10 W m⁻² R_n than bare land. The anomaly of R_n in different sowing dates was maintained within 2 W m⁻². LH generally was covariant with the change in R_n . However, the anomaly of LH in the late growth period was greater than that of R_n , resulting in negative SH, indicating that the EP scenario had stronger LH distribution tendency and

less SH was partitioned. Bigger anomaly of SH was happened in the initial and dormant stages. R anomaly fluctuated obviously only in the initial phase.

The contributions of surface energy balance components to T_c were shown in Fig.4b. Stronger radiation absorption provided more energy for the thermal motion of air and causing positive T_c differences of EP-LP. Correspondingly, higher distribution into LH, SH, and R was conducive to cooling T_c . Therefore, positive LH and SH differences of EP-LP showed negative T_c effects, and negative R difference of EP-LP showed positive T_c effect. The positive T_c anomaly of EP-LP reflected that the radiative process played the major role in the dormancy period. In the active growth time, the cooling effect of LH partitioning dominated the T_c anomaly.

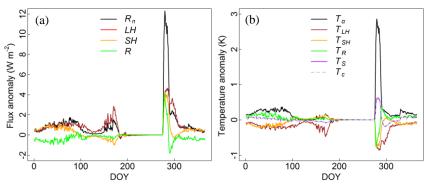


Fig.4 (a) The differences in the surface fluxes between the sowing scenarios of EP and LP, (b) its contributions to T_c anomaly.

 R_n means net radiation, T_a represents the temperature anomaly induced by changes in absorbed solar radiation. T_{LH} represents the temperature anomaly induced by changes in latent flux. T_{SH} represents the temperature anomaly induced by changes in sensible flux. T_R represents the temperature anomaly induced by changes in residual term. T_S represents

the temperature anomaly induced by changes in solar radiation, latent, sensible and residual fluxes.

4 Discussion

4.1 The diverse shift trends in sowing date of winter wheat in the NCP

The spatiotemporal changes of winter wheat phenology had been examined extensively examined in -the NCP. In the period of 1981-2009, the sowing date was on average delayed by 1.5 days/decade, but In the period of 1981-2009, sowing date delayed significantly at 13 station and advanced at 8 stations—out of the 36 agro-meteorological experiment stations were advanced (Xiao et al. 2013). In the NCP, the sowing date were on average delayed by 1.5 day/decade. The diverse trends in sowing date were also existed at the national scale, where 6 stations significantly advanced by up to 9.1 days/decade, and 11 stations significantly delayed by up to 10 days/decade (Tao et al. 2012).

The proper sowing date is key to ensure winter wheat survived through winter andreduce the freezing injury, insect pests and other harmful conditions (Sacks et al. 2010;
Zhang et al. 2012; Newbery et al. 2016). With faster growth in warmer environment, the
sowing date should be postponed to maintain the a proper coverage of winter wheat in
dormancy period. The warming of the NCP is regionally consistent (Shi et al. 2014), and
the diverse change of sowing date will affect the coverage of winter wheat, especially one
fifth stations advanced their sowing date. Earlier sowing may also benefited from the
reduction in freezing damage and the increase in pest diseases caused by higher minimum

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temperature, since more above-ground biomass will not be subject to lethal freezing damage and will resist higher harms from pests and diseases. There are also management practices to counteract the effects of advanced sowing date, such as deep tillage and delayed irrigation, which reduce the development of leaves and stems. <u>Until now, fewer studies had focused on the phenomenon of early sowing date and its underlying causes and countermeasures.</u>

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Although there were literatures reporting that the albedo process in winter is relatively important (Richardson et al. 2013; Lombardozzi et al. 2018), fewer studies directly addressed the influence of different surface characteristics and climate effect through biophysical process in the dormancy period. In the Oklahoma's winter wheat belt, the rapid crop growth during November exhibited a distinct cool anomaly against adjacent regions of dormant grassland. Over the period of December through April, the cool bias was visibly diminished although the greenness difference between grassland and wheat was more distinct (McPherson et al. 2004). The biophysical impacts between maize and perennial grass were simulated using Agro-IBIS model in US corn belt (Bagley et al. 2015). The results showed that much higher LAI of perennial scenario was existed in winter December–February (3 vs 0 m² m⁻²) and in summer June–August (10 vs 4 m² m⁻²). Perennial grass had smaller surface albedo (coupling snow effect) than maize in winter, but showed quite small difference in summer. During winter and summer, the perennial scenario had slightly higher LH than the maize scenario, but the difference in R_n between two scenarios was more than 10W m⁻² in winter (Bagley et al. 2015). The above

studies indicated that the cooling effect of higher LAI was inhibited in winter. The results of this current study indicate that higher LAI in winter has a warming effect, which is different from the conclusion above. The main reason was due to the relative contributions of surface albedo mechanism and surface flux distribution process.

The strong climate feedback in inter-sowing period, when wheat had been sown in the EP but hadn't in the LP, was related to the effect of tillage on maize stubble. The NCP is dominated by summer maize - winter wheat rotation system in which the ground is covered with maize stubble before wheat is sown. The damage of sowing to stubble is conducive to the reduction of albedo since stubble has larger surface reflectivity than soil (O'Brien et al. 2019). The 0.1 increase of surface albedo caused by no-till management, which was also the magnitude of our simulation (Table 6), cooling the hottest summer days by 2 °C or more (Davin et al. 2014). The inter-sowing period is equivalent to no-tillage period, when early sowed wheat absorbed more net radiation with lower albedo by destroying stubble and causing higher temperature (Fig.3b, Fig4a).

Previous studies showed that the increase of vegetation cover caused warming feedback by destroying the high albedo of snow in the case of snow cover (Richardson et al. 2013; Bagley et al. 2015; Lombardozzi et al. 2018). In our simulation, except for the large difference in crop coverage in phase 1, the snow and crop had consistent coverage in other phases (Supplement Table 1), which means albedo difference between two scenarios was not caused by snow. Low soil water content also contributed to the high surface albedo (Seneviratne et al. 2010)(Fig.6b). With the decrease of surface soil moisture, surface albedo increased in winter, which explained why albedo in the winter

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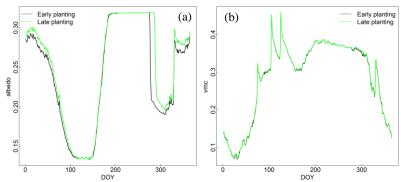


Fig.6 Dynamics of (a) surface albedo and (b) surface soil moisture content under two sowing scenarios in winter wheat growing season

Table 6 The reflectivity of different surface coverings in near-infrared and visible bands in the SiBcrop model

<u>Material</u>	Visible band	Near Infrared band
Green leaf	0.08	<u>0.3</u>

Snow	0.8	0.4
Soil with crop	0.11	0.314
Bare soil	<u>0.33</u>	<u>0.35</u>

4.3 Cooling effect of EP-LP during the growing period

The phenological shifts, such as earlier leaf unfolding, delayed leaf fall, andlengthening of the green-cover season have feedback on climate through biophysical and
biogeochemical processes (Penuelas et al. 2009). Previous studies showed cooling effect
in the photosynthetic active period through surface biophysical mechanism_in
the -cropland (e.g. (Sacks and Kucharik 2011; Zhang et al. 2013; Bohm et al. 2020)).

In the NCP, the increased spring surface greenness at farmland, benefited from advanced re-greening stage of winter wheat (Xiao et al. 2013; Liu et al. 2017), had cooling and wetting effects (Zhang et al. 2013) and suppressed the moderate to light rainfall (Zhang et al. 2015). The analysis found that surface greening increased the partitioning into *LH* and reduced *SH* to cooling surface air and suppression of rainfall (Zhang et al. 2013; Zhang et al. 2015). Distinguished difference between early-covering crops (winter wheat, winter rapeseed, winter barley) and late-covering crops (corn, silage maize, sugar beet) in central Europe caused impacts on simulated surface energy fluxes and temperature in the Noah-MP model, the higher LAI led to an increase in *LH*, decreased in *SH* and eventually surface cooling in May-September (Bohm et al. 2020). The Agro-IBIS model was used to study the impacts on surface energy balance of advanced corn sowing date (10 days): Early sowing means earlier development and

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senescence of LAI, causing stronger disparity of LH than R_n with bigger LAI and probably a slight cooling of T_a in June (Sacks and Kucharik 2011). Similar conclusions were presented based on simulated T_c results with modified SiBcrop.

5 Conclusions

- The dynamics of winter wheat LAI and T_c under two sowing date scenarios were simulated by the SiBcrop model in the NCP, and the T_c disparity between the two scenarios was explained by the surface energy balance. The findings include:
 - (1) Earlier sowing date of winter wheat had higher LAI than later sowing date.
- (2) The T_c disparity between EP and LP is divided into two periods: warming effect in the dormancy period, and cooling effect in the active growth period.
- (3) Surface energy balance can interpret the climate feedback mechanism of sowing date-shift, that is, the dominated role of albedo-radiative process in the dormancy period is surpassed by *LH* partitioning-non-radiative process in the growth period.
- (4) The responses of LAI and T_c to sowing date at station scale were divergent: controlled by T_a in the dormancy period, and influenced by P and T_a in the growth period.

The study had some shortcomings. The single model simulation was highly dependent on the structure and parameterization scheme of the model. The climate feedback was reflected by the canopy temperature. In the SiBcrop model, the spatial distribution of stations was not fully considered in the determination of sowing date, which resulted in too early or too late sowing at some stations.

552 showed that even when land use/cover type remains unchanged, variations in surface 553 Acknowledgments This study is supported by the National Natural Science Foundation of China 554 (41801020). 555 556 References 557 A, D., Xiong, K., Zhao, W., Gong, Z., Jing, R., and Zhang, L.: Temporal trend of climate 558 559 change and mutation analysis of North China Plain during 1960 to 2013. Scientia Geographica Sinica, 36(10), 1555-1564, 2016. 560 Bagley, J. E., Kueppers, L. M., Billesbach, D. P., Williams, I. N., Biraud, S. C., and Torn, 561 M. S.: The influence of land cover on surface energy partitioning and evaporative fraction 562 563 regimes in the US Southern Great Plains. Journal of Geophysical Research-Atmospheres, 122(11), 5793-5807, 2017. 564 Bagley, J. E., Miller, J., and Bernacchi, C. J.: Biophysical impacts of climate - smart 565 agriculture in the Midwest United States. Plant, cell & environment, 38(9), 1913-1930, 566 2015. 567 Bohm, K., Ingwersen, J., Milovac, J., and Streck, T.: Distinguishing between early- and 568 late-covering crops in the land surface model Noah-MP: impact on simulated surface 569 energy fluxes and temperature. Biogeosciences, 17(10), 2791-2805, 2020. 570 Boisier, J. P., de Noblet-Ducoudre, N., Pitman, A. J., Cruz, F. T., Delire, C., van den 571 572 Hurk, B. J. J. M., et al.: Attributing the impacts of land-cover changes in temperate regions on surface temperature and heat fluxes to specific causes: Results from the first 573 LUCID set of simulations. Journal of Geophysical Research-Atmospheres, 117, 2012. 574 Chen, M., Griffis, T. J., Baker, J., Wood, J. D., and Xiao, K.: Simulating crop phenology 575 576 in the Community Land Model and its impact on energy and carbon fluxes. Journal of 577 Geophysical Research-Biogeosciences, 120(2), 310-325, 2015. 578 Chen, Y., Liu, F., Tao, F., Ge, Q., Jiang, M., Wang, M., et al.: Calibration and validation of SiBcrop Model for simulating LAI and surface heat fluxes of winter wheat in the 579 North China Plain. Journal of Integrative Agriculture, 19(9), 2-11, 2020. 580 Cho, M. H., Boo, K. O., Lee, J., Cho, C., and Lim, G. H.: Regional climate response to 581 land surface changes after harvest in the North China Plain under present and possible 582 583 future climate conditions. Journal of Geophysical Research-Atmospheres, 119(8), 584 4507-4520, 2014. Collatz, G. J., Berry, J. A., Farquhar, G. D., and Pierce, J.: The relationship between the 585 Rubisco reaction mechanism and models of photosynthesis*. Plant, Cell & Environment, 586 587 13(3), 219-225, 1990.

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