



1 **Carbon storage in phosphorus limited grasslands may**
2 **decline in response to elevated nitrogen deposition: a long-**
3 **term field manipulation and modelling study**

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25 **Abstract**

26 In many temperate ecosystems, nitrogen (N) limits productivity, meaning anthropogenic N
27 deposition can stimulate plant growth and subsequently carbon (C) sequestration. Phosphorus
28 (P) and N-P co-limited grasslands are widespread, yet there is limited understanding of their
29 responses to N deposition, which may transition more ecosystems toward P-limited or N-P co-
30 limited states. Here, we investigate the consequences of enhanced N addition on the C-N-P
31 pools of grasslands in different states of nutrient limitation. We explored the response of a long-
32 term nutrient-manipulation experiment on two N-P co-limited grasslands; an acidic grassland of
33 stronger N-limitation and a calcareous grassland of stronger P-limitation, by combining data with
34 an integrated C-N-P cycling model (N14CP). To explore the role of P-access mechanisms in
35 determining ecosystem state, we allowed P-access to vary, and compared the outputs to plant-
36 soil C-N-P data. Combinations of organic P access and inorganic P availability most closely
37 representing data were used to simulate the grasslands and quantify their temporal response to
38 nutrient manipulation. The model suggested N additions have increased C stocks in the acidic
39 grassland, but decreased them in the calcareous, where N provision exacerbated P-limitation
40 and reduced biomass input to the soil. Furthermore, plant acquisition of organic P may play an
41 important role in reducing P-limitation, as both simulated grasslands increased organic P uptake
42 to meet P demand. We conclude that grasslands of differing limiting nutrients may respond to N
43 deposition in contrasting ways, and stress that as N deposition shifts ecosystems toward P-
44 limitation, a globally important carbon sink risks degradation.

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49 **1. Introduction**

50 Grasslands represent up to a third of terrestrial net primary productivity (NPP) [Hoekstra et al., 2005]
51 and potentially hold over 10% of the total carbon stored within the biosphere [Jones and Donnelly,
52 2004]. The ecosystem services provided by grasslands, such as carbon storage, are highly sensitive to
53 perturbations in their nutrient cycling, including the perturbation of nitrogen (N) inputs from
54 atmospheric deposition [Phoenix et al., 2012].

55 Since the onset of the industrial revolution, human activity has doubled the global cycling of N,
56 meaning that anthropogenic sources of fixed N now surpass natural sources by 7 Tg N yr⁻¹ [Fowler et
57 al., 2013]. Much of this additional N is deposited on terrestrial ecosystems from atmospheric sources.
58 This magnitude of N deposition results in a range of negative impacts on ecosystems (including
59 grasslands) such as reductions in biodiversity [Bobbink et al., 2010; Southon et al., 2013], acidification
60 of soil, and the accumulation of toxic levels of ammonium [Carroll et al., 2003; Horswill et al., 2008;
61 Phoenix et al., 2012]

62 Despite large anthropogenic fluxes of N, most terrestrial ecosystems on temperate post-glacial soils
63 are thought to be N-limited (biomass production is most restricted by N availability) [Vitousek and
64 Howarth, 1991; Du et al., 2020], as weatherable sources of phosphorus (P) remain sufficiently large to
65 meet plant P demand [Vitousek and Farrington, 1997; Menge et al., 2012]. Both empirical and
66 modelling studies have shown that pollutant N, when deposited on N-limited ecosystems, can increase
67 productivity [Tipping et al., 2019] and soil organic carbon (SOC) storage [Tipping et al., 2017], largely
68 as a result of stimulated plant growth. This suggests that while there are negative consequences of N
69 deposition, there may also be benefits from enhanced plant productivity and increases in carbon
70 storage.

71 Whilst most research focuses on N limited ecosystems [LeBauer and Treseder, 2008], a number of
72 studies have highlighted that P limitation and N-P co-limitation are just as prevalent, if not more



73 widespread, than N limitation [Fay et al., 2015; Du et al., 2020; Hou et al., 2020]. In a meta-analysis of
74 grassland nutrient addition experiments spanning five continents, Fay et al. [2015] found that
75 aboveground annual net primary productivity (ANPP) was limited solely by P in 8 sites and co-limited
76 by N and P in 25, compared to only 10 sites showing N limitation alone [Fay et al., 2015]. Similarly, P
77 additions in 652 field experiments increased aboveground plant productivity by an average of 34.9%
78 [Hou et al., 2020], while it is estimated that co-limitation of N and P could constrain up to 39% of the
79 terrestrial surface's productivity [Du et al., 2020].

80 Furthermore, P limitation may be exacerbated by N deposition [Johnson et al., 1999; Phoenix et al.,
81 2004], or become increasingly prevalent as previously N-limited ecosystems transition to N-sufficient
82 states [Goll et al., 2012]. For example, in parts of the Peak District National Park, UK, N deposition has
83 exceeded $3 \text{ g m}^{-2} \text{ yr}^{-1}$, with further experimental additions of $3.5 \text{ g m}^{-2} \text{ yr}^{-1}$ leading to decreases rather
84 than increases in productivity of calcareous grasslands [Carroll et al., 2003], in contrast to previous
85 studies of N deposition enhancement of N-limited productivity [Tipping et al., 2019]. This makes P
86 limitation and N-P co-limitation critical to understand in the context of global carbon and nutrient
87 cycles. By definition, N deposition should impact N-P co-limited and P-limited ecosystems differently
88 to N-limited ones, yet there is little understanding of how N deposition impacts P and N-P co-limited
89 ecosystems.

90 While N deposition may worsen P limitation in some instances, plant strategies for P acquisition, such
91 as changes in root architecture and increased root exudation [Vance et al., 2003] require substantial
92 investments of N, suggesting that in some areas with P depleted soils, N may facilitate enhanced P
93 uptake [Long et al., 2016; Chen et al., 2020]. It has been shown that N deposition can stimulate
94 additional production of extracellular phosphorus-cleaving enzymes by plants [Johnson et al., 1999;
95 Phoenix et al., 2004], thereby increasing plant availability of organic forms of phosphorus in order to
96 help meet plant P demand. This response could be driven by exacerbated P-limitation resulting from
97 N deposition or extra N availability making elevated enzyme production possible.



98 These interdependencies of the C, N and P cycles make understanding an ecosystem's response to
99 perturbations in any one nutrient cycle challenging, particularly when ecosystems are not solely
100 limited in N. This highlights the need for integrated understanding of plant-soil nutrient cycling across
101 the C, N and P cycles, and in ecosystems that are not solely N-limited.

102 Process-based models have a role to play in addressing this, as they allow us to test our mechanistic
103 understanding and decouple the effects of multiple drivers. There has been increasing interest in
104 linking C with N and P cycles in terrestrial ecosystem models [Wang et al., 2010; Achat et al., 2016;
105 Jiang et al., 2019] as the magnitude of the effects that anthropogenic nutrient change can have on
106 biogeochemical cycling are realised [Yuan et al., 2018]. Yet, few modelling studies have explicitly
107 examined the effects of P or N-P co-limitation, likely mirroring the relatively fewer empirical studies
108 of these systems.

109 Combining process-based models with empirical data from long-term nutrient-manipulation
110 experiments offers a valuable opportunity for understanding ecosystem responses to environmental
111 changes that may only manifest after extended periods of time, such as with changes in soil organic
112 C, N and P pools, which typically occur on decadal timescales [Davies et al., 2016a, Janes-Bassett et
113 al., 2020]. Ecological data from these experiments can be used to drive and calibrate process-based
114 models, which in turn can disentangle multiple interacting processes involved in plant-soil nutrient
115 cycling, that otherwise makes interpretation of empirical experiments complex. This allows us to test
116 our assumptions of the key drivers, processes and pathways for carbon and nutrient cycling in
117 grasslands exposed to multiple environmental perturbations.

118 Here, we combine new data from a long-term nutrient manipulation experiment on two contrasting
119 upland grasslands (acidic and calcareous), both N-P co-limited to differing degrees (one more P
120 limited, one more N limited within the co-NP range), with the mechanistic C-N-P plant-soil
121 biogeochemical model; N14CP [Davies et al., 2016b]. We use this model and data to simulate the long-
122 term nutrient manipulation experiment in both grasslands and then use the calibrated model to



123 determine the long-term consequences of differing nutrient limitation on plant and soil C, N and P. To
124 do so, we allow modelled P-access conditions to vary and used the combinations of P-access variables
125 that most closely represented empirical data to simulate the grasslands.

126 Specifically, we aim to first explore how variation in P acquisition parameters, that control access to
127 organic and inorganic sources of P in the model, may help account for differing responses to N and P
128 additions in the empirical data on aboveground biomass carbon and soil C, N and P pools. Secondly,
129 we explore the effects of long-term anthropogenic N deposition at the site and the effects of
130 experimental nutrient additions (N and P) on plant and soil variables of the simulated acidic and
131 calcareous grasslands. This will help improve our understanding of how similarly P-limited or N-P co-
132 limited grasslands may respond to similar conditions. We hypothesise that 1) flexible P-access within
133 the model may help in alleviating P limitation and that 2) grasslands of contrasting nutrient limitation
134 respond to N deposition and nutrient treatment in dissimilar ways, with N deposition exacerbating
135 nutrient limitation in more P-limited grasslands, in turn leading to declining productivity and carbon
136 sequestration.

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146 **2. Methods**

147 **2.1. Field experiment description**

148 Wardlow Hay Cop (henceforth referred to as Wardlow) is a long-term experimental grassland site in
149 the Peak District National Park (UK) [Morecroft et al., 1994]. There are two distinct grassland
150 communities occurring in close proximity; acidic (National vegetation classification U4e) and
151 calcareous (NVC CG2d) semi-natural grasslands (Table S2). Both grasslands share a carboniferous
152 limestone hill but the calcareous grassland sits atop a thin humic ranker [Horswill et al., 2008] and
153 occurs predominantly on the hill brow. In contrast, the acidic grassland occurs in the trough of the
154 hill, allowing the accumulation of wind-blown loess and the formation of a deeper soil profile. As
155 such, the acidic grassland shares the same limestone bedrock but sits atop a palaeo-argillic brown
156 earth soil [Horswill et al., 2008].

157 The biomass in both grasslands show signs of both N and P-limitation, though they differ in the
158 relative strength of limitation by N and P. The acidic grassland is co-limited in N and P, as positive
159 biomass growth responses are observed with additions of both nutrients [Phoenix et al., 2003]. The
160 calcareous grassland, however, is more strongly P-limited, showing increased productivity only with
161 the addition of P [Carroll et al., 2003], though N and P co-limitation has been observed [Phoenix et
162 al., 2003].

163 Nutrients (N and P) have been experimentally added to investigate the effects of elevated N
164 deposition and the influence of P limitation [Morecroft et al., 1994]. Nitrogen treatments simulate
165 additional N deposition to the background level and also act to exacerbate P limitation [Johnson et
166 al., 1999; Phoenix et al., 2004], whereas the P treatment acts to alleviate it. Nutrients are added as
167 solutions of distilled water and applied as fine spray by backpack sprayer, and have been applied
168 monthly since 1995, and since 2017 bi-monthly. Nutrient additions are in the form of NH_4NO_3 for
169 nitrogen and $\text{NaH}_2\text{PO}_4 \cdot \text{H}_2\text{O}$ for phosphorus. Nitrogen is applied at rates of 0 (distilled water control –



170 0N), 3.5 (low nitrogen – LN) and 14 g N m⁻² yr⁻¹ (high nitrogen – HN). The P treatment is applied at a
171 rate of 3.5 g P m⁻² yr⁻¹ (phosphorus – P).

172 Data collected from the Wardlow grasslands for the purpose of this work are; aboveground biomass
173 C, SOC, and total N, which is assumed to be equivalent to modelled SON. This new data is combined
174 with total P data that was collected by Horswill et al. at the site [Horswill et al., 2008]. Summaries of
175 these data are available within the supplementary material (Table S4) and details of their collection
176 and conversion to model-compatible units in supplementary section S1.

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178 **2.2. Summary of model processes**

179 2.2.1. N14CP model summary

180 The N14CP ecosystem model is an integrated C-N-P biogeochemical cycle model that simulates NPP,
181 C, N and P flows and stocks between and within plant biomass and soils, and their associated fluxes
182 to the atmosphere and leachates [Davies et al., 2016b].

183 N14CP was originally developed and tested on 88 northern Europe plot-scale studies, including
184 grasslands, where C, N and P data were available. All but one of the tested ecosystems exhibited N
185 limitation [Davies et al., 2016b]. It has also been extensively and successfully blind-tested against
186 SOC [Tipping et al., 2017] and NPP data from unimproved grassland sites across the UK (~500 and
187 ~300 sites, respectively) [Tipping et al., 2019]. However, the model has not been extensively tested
188 against sites known to exhibit P or N-P co-limitation. A full model description can be found in Davies
189 et al., [2016b], however, a summary of the most relevant features is given here for convenience.

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193 2.2.2. Net primary productivity

194 N14CP simulations run on a quarterly time step and are spun up from the onset of the Holocene
195 (10,000 BP in the model). Plant biomass is simulated in the model as two sets of pools of coarse and
196 fine tissues containing C, N and P. NPP adds to these pools and is calculated on a quarterly basis,
197 with growth occurring in this case in quarters 2 and 3 (spring and summer). The NPP is calculated on
198 the basis of a single limiting factor (i.e. temperature, precipitation, N or P) in accordance with a
199 Liebig's law of the minimum [Davies et al., 2016b]. However, nutrient co-limiting behaviour can
200 occur in the model through increased access to organic P sources in the presence of sufficient N, and
201 by having the rate of N fixation dependent on plant and microbial available P [Davies et al., 2016b].

202 To calculate NPP, first, a regression is used to estimate potential NPP based on climate as in Tipping
203 et al. [2014]. The corresponding plant demand for N and P, driven by plant stoichiometric
204 constraints, to achieve this potential NPP is then calculated and compared with available N and P in
205 the model [Davies et al., 2016b; Tipping et al., 2017].

206

207 2.2.3. Plant available N and P

208 Plant available N is derived from fixation, the decomposition of coarse litter, decomposition of SOM
209 pools and by atmospheric deposition. Plant available P comes from desorption of inorganic P from
210 soil surfaces, the turnover of SOM, and the decomposition of coarse litter [Davies et al., 2016b]. The
211 sorbed inorganic P pool builds over time with inputs of weathered P and sorption of any excess plant
212 available inorganic P, and desorption occurs as a first order process.

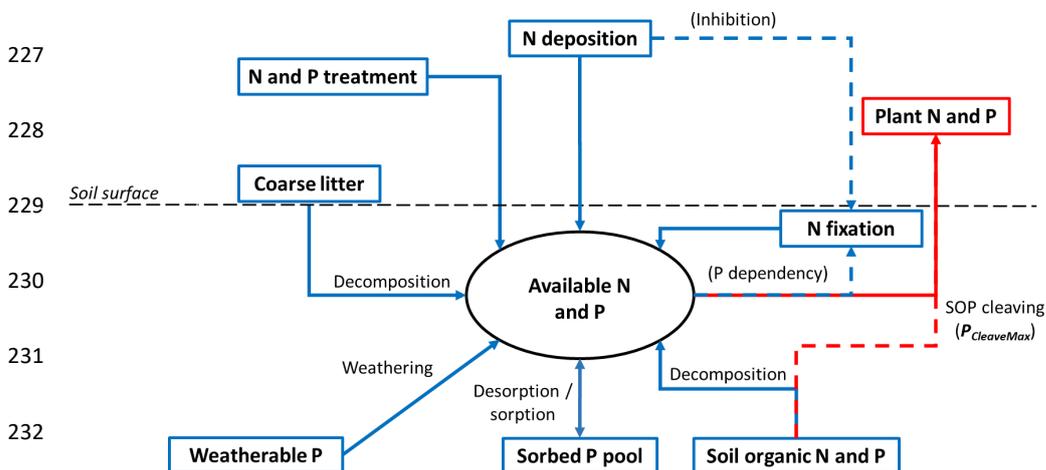
213 Phosphorus enters the plant-soil system by weathering of parent material, the initial value of which
214 ($P_{\text{Weath}0}$ within the model) can be set to a default value, or made site-specific by calibrating this initial
215 condition to soil observational data (as in methods section 2.3.3). In principle, P can be added in
216 small quantities by atmospheric deposition [Ridame and Guieu, 2002] or by local redistribution



217 [Tipping et al., 2014]. For the purpose of this study, P deposition is set to zero as its net contribution
 218 to the total P pool in comparison to weathering is assumed to be minimal. Annual release of
 219 weathered P is determined by a first-order rate constant, which is temperature dependent. Where
 220 the mean temperature falls below 0 °C, it is assumed that no weathering occurs.

221 In the presence of sufficient N and where plant demand for P cannot be met by more accessible P
 222 sources, plants can access P from the soil organic phosphorus (SOP) pool via a cleaving parameter
 223 termed $P_{\text{CleaveMax}}$, which is the maximum quantity of cleavable P within a growing season (g m^{-2}). It is
 224 $P_{\text{CleaveMax}}$ and P_{Weath0} that we allow to vary to account for discrepancies in empirical data.
 225 Contributions of N and P toward the plant available pools are summarised in Figure 1.

226



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Figure 1: An illustration of the plant available N and P pools in the N14CP model. In the model, N can enter the available
 234 pool via atmospheric deposition, biological nitrogen fixation, coarse litter decomposition and decomposition of the soil
 organic matter pools. For P, the two main contemporary sources are the inorganic sorbed pool and from the turnover of soil
 organic matter. The former is derived initially from the weatherable supply of P, defined by its initial condition (P_{Weath0}).
 235 Solid lines indicate input to another pool and a dashed line indicates either a feedback or interaction with another pool.
 These interactions include the downregulation of N fixation by N deposition, the dependency of N fixation on P availability,
 236 and the cleaving of organic P by plants when N is sufficient and other P sources are inaccessible.

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239 Phosphorus access within N14CP is determined by a hierarchal relationship, whereby plants and
240 microbes access the most readily available P sources first and only move onto the next once it has
241 been exhausted. Out of the P sources available to plants (Fig 1), organic P is the least bioavailable
242 within the model hierarchy, hence a depletion in the SOP pool is indicative of severe P stress and low
243 P availability.

244 Plant nutrient demand is defined by Plant Functional Type (PFT), which changes through time in
245 accordance with ecosystem succession, and includes broadleaf woodland, coniferous woodland,
246 shrubs (heather, heather grassland and montane habitats) and herbaceous plants (including neutral,
247 acidic and calcareous grasslands). Each PFT has two stoichiometric end members, allowing the
248 model to represent transitions from N-poor to N-rich species or an enrichment of the fine tissues
249 within a single species (or a combination of both) [Davies et al., 2016b], dependent on available N.
250 This allows a degree of flexibility in plant C:N ratios in response to environmental changes such as N
251 deposition.

252 If the available nutrients cannot meet the calculated plant nutrient demand, the minimum
253 calculated NPP based on either N or P availability is used, giving an estimation of the most limiting
254 nutrient to plant growth. As the limiting nutrient of an ecosystem may not be static through time
255 [Vitousek et al., 2010], and can change in response to external inputs of nutrients such as N
256 deposition [Menge and Field, 2007], by looking at changes in the limiting nutrient, we can better
257 explain model behaviour and its predictions of changes to C, N and P pools.

258

259 2.2.3. Soil processes – organic C, N and P

260 Detailed descriptions of C, N and P inputs, outputs and processes in the soil are explained in Davies
261 et al. [2016b]. Nitrogen enters the system from N fixation and atmospheric N deposition. The former



262 is related to P availability and based on literature values but is downregulated by N deposition
263 (Figure 1).

264 A fraction of plant biomass is converted to litter in each quarterly time step and contributes a
265 proportion of its C, N and P content to SOM, which is sectioned into three pools (fast, slow and
266 passive) depending on turnover rate [Davies et al., 2016b]. Carbon is lost as CO₂ following
267 temperature-dependent decomposition and as dissolved organic carbon. Likewise, N and P are lost
268 via dissolved organic N and P in a proportion consistent with the stoichiometry of each SOM pool.
269 Inorganic N is lost via denitrification and inorganic P can be sorbed by soil surfaces. Both inorganic N
270 and P can be leached in dissolved forms if they are in excess of plant demand.

271

272 **2.3. Simulating the field manipulation experiment with the model**

273 2.3.1. Nutrient applications

274 Nutrient treatments are treated in N14CP as individual plots in the simulations with differing
275 amounts of inorganic N and P applied in line with the field experimental treatments. The N and P
276 nutrient treatments are added to the bioavailable N and P pools of the model on a quarterly basis in
277 line with the model's time-step. While Wardlow nutrient treatments are applied monthly and N14CP
278 quarterly, the annual sum of applied N or P is equivalent, and nutrients are applied during all
279 quarters.

280

281 2.3.2. Input drivers

282 Atmospheric N deposition, climate and PFT history need to be provided as annual time series to
283 drive the model. As the acidic and calcareous sites are co-located, these input timeseries are shared
284 for both grasslands. A summary of the data used for model input and model testing are provided in
285 supplementary Tables S3 and S4 respectively. It is assumed in the model that anthropogenic N



286 deposition was negligible prior to 1800 and the onset of the industrial revolution. After 1800, N
287 deposition is assumed to have increased similarly across Europe [Schopp et al., 2003]. In N14CP, this
288 trend is linearly extrapolated from the first year of data (1880) back to 1800 [Tipping et al., 2012].
289 Data regarding N deposition that is specific to Wardlow was incorporated between the years 2004
290 and 2014 and the Schöpp et al. [2003] anomaly scaled to represent the high N deposition of the site.

291 To provide climate forcing data, daily minimum, mean and maximum temperature and mean
292 precipitation records beginning in 1960 were extracted from the UKPC09 Met office CEDA database
293 (Table S3). The data nearest to Wardlow was calculated by triangulating latitude and longitude data
294 and using Pythagoras' theorem to determine the shortest distance. These data were converted into
295 mean quarterly temperature and precipitation. Prior to this, temperature was assumed to follow
296 trends described in Davies et al. [2016b] and mean quarterly precipitation was derived from Met
297 Office rainfall data between 1960 to 2016 and held constant.

298 To simulate the sites' land use history, PFT was defined on an annual basis using data on Holocene
299 pollen stratigraphy of the White Peak region of Derbyshire [Taylor et al. 1994]. The defined PFT
300 history represents an early colonisation of virgin soil by herbaceous plants following deglaciation
301 (10,000 BP). A succession to broadleaf temperate forest develops and persists until a forest
302 disturbance (but not clearance) by human settlers occurs in 5,190 BP, leaving an open forest mosaic
303 characterised by hazel trees (defined in the model as shrub to distinguish it from forest). This open
304 forest was deforested in 4100 BP to be used as rough grazing pasture for livestock, a practice that
305 continues to this day.

306

307 2.3.3. Model parameters for the acidic and calcareous grasslands

308 The N14CP model has been previously calibrated and tested against a wide range of site data to
309 provide a general parameter set that is applicable to temperate semi-natural ecosystems, without



310 extensive site-specific calibration [Davies et al., 2016b]. The majority of those parameters are used
311 here for both grasslands.

312 However, two parameters relating to P sources and processes were allowed to vary between the
313 sites: the initial condition for the weatherable P pool, P_{Weath0} ; and the rate of plant access to organic
314 P sources, $P_{\text{CleaveMax}}$ (Figure 1).

315 We allowed P_{Weath0} to vary for each grassland as variation in a number of factors including lithology
316 and topography mean that we should expect the flux of weathered P entering the plant-soil system
317 to vary on a site-by-site basis Davies et al. [2016b]. Indeed, we should expect that P_{Weath0} differs
318 between the acid and calcareous grasslands, as despite their proximity, they have differing lithology.
319 Davies et al. [2016b], show that variation in this initial condition considerably helps explain variance
320 in contemporary SOC, SON and SOP stocks between sites. However, it is difficult to set this
321 parameter directly using empirical data, as information on lithology and P release is limited at the
322 site scale.

323 We also allowed $P_{\text{CleaveMax}}$ to vary as this mechanism for P acquisition has been under-explored in
324 previous modelling studies. This is the first time that this model has been knowingly applied to N-P
325 co-limited or P-limited grasslands instead of N-limited sites. Soil organic P has been shown to be an
326 important source of P to plants in P-stressed environments [Balemi and Negisho, 2012]. However,
327 the rates of access to SOP and their controls are relatively poorly understood. We allowed the rate
328 at which P can be cleaved from this pool ($P_{\text{CleaveMax}}$) to vary, to investigate how plant P acquisition
329 might change when more readily accessible P forms become scarcer.

330 We ran a series of simulations systematically varying P_{Weath0} and $P_{\text{CleaveMax}}$ and comparing the results
331 to observations, we simulated the two grasslands and their treatment blocks with a set of a 200
332 parameter combinations. This captured all combinations of 20 values of P_{Weath0} between 50 and 1000
333 g m^{-2} and 10 values of $P_{\text{CleaveMax}}$ between 0 to 1 g m^{-2} per growing season using a \log_{10} spacing to focus
334 on the lower range of $P_{\text{CleaveMax}}$ values. The P_{Weath0} range was set to capture the lower end of P_{Weath0}



335 estimates described in Davies et al. [2016b], which were more likely to be appropriate for these P-
336 poor sites. We explored a range of values for $P_{\text{CleaveMax}}$, from zero where no access to organic sources
337 is allowed, to 1 g m^{-2} per growing season – a rate in the order of magnitude of a fertilizer application.
338 The model outputs were compared to measured aboveground biomass C, SOC, SON (assumed
339 equivalent to total N) and total P (Table S4) for each grassland. We tested how these parameter sets
340 performed by calculating the error between the observations and model outputs of the same
341 variables for each combination of $P_{\text{CleaveMax}}$ and P_{Weath0} . The sum of the absolute errors between
342 modelled and observed plant C and soil C, N and P data were scaled (to account for differing
343 numbers of observations) and summed to provide an F value (Equation 1) as an overall measure of
344 error across multiple observation variables. The parameter combination with the lowest F value that
345 still maintained the grassland's empirical response to nutrient additions for both the acidic and
346 calcareous grasslands (Supplementary section S1.1.), was used within the analysis.

347

$$F = \left(\frac{SAE[C_{SOM}]}{\bar{C}_{SOM,Obs}} \right) / C_n + \left(\frac{SAE[N_{SOM}]}{\bar{N}_{SOM,Obs}} \right) / N_n + \left(\frac{SAE[P_{Total}]}{\bar{P}_{Total,Obs}} \right) / P_n \quad (\text{Equation 1})$$

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356 **3. Results**

357 **3.1. Varying phosphorus source parameters**

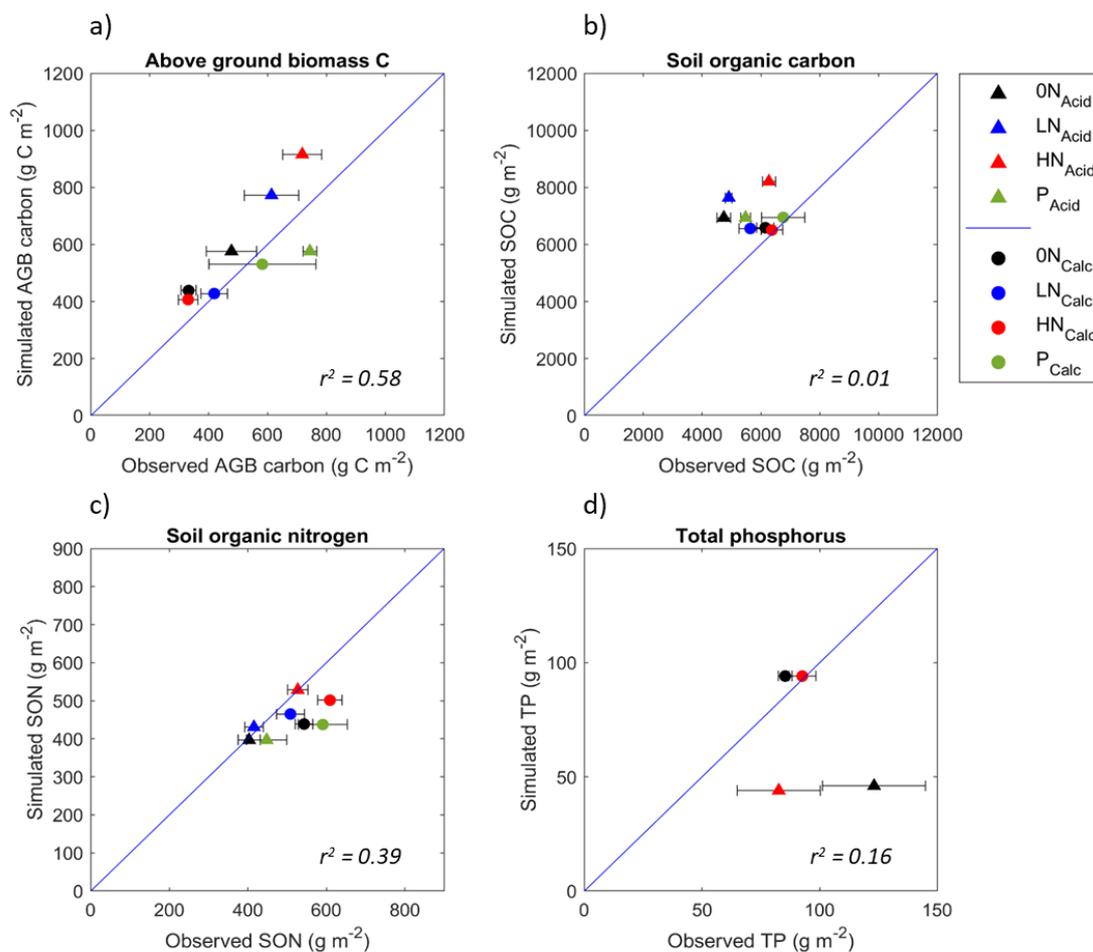
358

359 The model calibration selected parameter values for P_{Weath0} and $P_{\text{CleaveMax}}$ that indicate contrasting
360 use of P sources by the two grasslands, with the acidic grassland capable of acquiring more P from
361 organic sources, having a $P_{\text{CleaveMax}}$ value of $0.3162 \text{ g m}^{-2} \text{ season}^{-1}$ compared to the calcareous, with a
362 value 10 times smaller at $0.0316 \text{ g m}^{-2} \text{ season}^{-1}$. Conversely, inorganic P availability was greater in the
363 calcareous grassland due to the larger weatherable pool of P, P_{Weath0} , at 300 g m^{-2} compared to 150 g
364 m^{-2} in the acidic.

365 The outputs for the calibrated model are shown in Figure 2 against the observations for above-
366 ground biomass C, soil organic C, and N for both the acidic and calcareous grasslands (Fig 2). The
367 model estimates of above ground biomass C are broadly aligned with the observations: capturing
368 variation between the grasslands and treatments ($r^2=0.58$), and on average overestimating the
369 magnitude by 12.9% (SE \pm 11.9) and 12.1% (SE \pm 9.4) for the acidic and calcareous grasslands
370 respectively (Fig 2a). Soil organic C on average was slightly overestimated (7.1% with SE \pm 3.3) for the
371 calcareous grassland (Fig 2b), with a larger average overestimate for the acidic grassland (39.9% with
372 SE \pm 6.8). However, in this latter case the variation between treatments was better captured.

373 Simulated magnitudes of SON are well-aligned with observations for the acidic grassland, with an
374 average error of 2.3% (SE \pm 3.2), whilst the SON at the calcareous grassland was on average
375 underestimated by 17.8% (SE \pm 3.6) (Fig 2c). Finally, the model overestimated total soil P (defined in
376 the model as organic P plus sorbed P) by an average of 6.0% (SE \pm 4.3) for the calcareous but
377 underestimated by 54.7% (SE \pm 8.0) in the acidic grassland, which was the least accurately predicted
378 variable out of those investigated (Fig 2d). Raw data used for Figure 2 are provided in supplementary
379 tables S5 and S6.

380



381 **Figure 2:** A comparison of the observed values of a) aboveground biomass carbon, b) soil organic carbon, c) soil organic nitrogen and d) total soil phosphorus from both grasslands, with simulated values from the model. The blue line represents a 1 to 1 relationship and the closer the data points are to the line, the smaller the discrepancy between observed and modelled data. All data are in grams per metre squared and all treatments for which data were collected are presented. The horizontal error bars represent the standard error of the empirical data means. The r^2 value of regression models fitted to the data are presented to assess closeness to the 1 to 1 line.



382 **3.2. The limiting nutrient through time**

383 The model suggests that the acidic grassland NPP remained N-limited from 1800 through to 2020
384 under most nutrient treatments (Fig 3). Nitrogen deposition increased the potential NPP through
385 time and the grassland moved toward co-limitation in the LN treatment (i.e. the N and P lines were
386 closer) but remained N-limited (Fig 3b). In the HN treatment, the acidic grassland shifted to P
387 limitation as N-limited NPP surpasses P-limited NPP (Fig 3c).

388 The calcareous grassland was also initially N-limited according to the simulation, but was driven
389 through a prolonged (c. 100 year) state of apparent co-limitation until clearly reaching P-limitation in
390 1950, solely as a result of N deposition (Fig 3). In the ON treatment, the grassland remained P-limited
391 but the potential NPP values for N and P are similar, suggesting the grassland is close to co-limitation
392 (Fig 3e). The LN and HN treatment amplified pre-existing P-limitation, lowering the potential NPP of
393 the grasslands (Fig 3f, g). With the addition of P in 1995, P limitation is alleviated, and the ecosystem
394 transitions to a more productive N-limited grassland (Figure 3h).

395 Another way to interpret the extent of nutrient limitation within N14CP with specific reference to P-
396 demand, is to assess the rate of P cleaving through time. These data corroborate the N and P-limited
397 NPP data, showing that in the calcareous grassland, the maximum amount of cleavable P is accessed
398 by plants in the ON, LN and HN treatments from approximately 1900 through to the end of the
399 experimental period in 2020 (Fig S1, Table S14), highlighting its consistent state of P or N-P co-
400 limitation.

401 Conversely, while P is cleaved in the ON control treatment in the acidic grassland, it occurs at
402 approximately one third of the total rate, hence the grassland is not entirely P-limited (Fig S1, Table
403 S10). The LN treatment increases the rate of SOP cleaving and HN causes it to reach its maximum
404 value, confirming the shift to P limitation suggested by the NPP data (Fig S1, Table S10). Soil organic
405 P cleaving does not occur in the P-treated plots of either grassland.

406

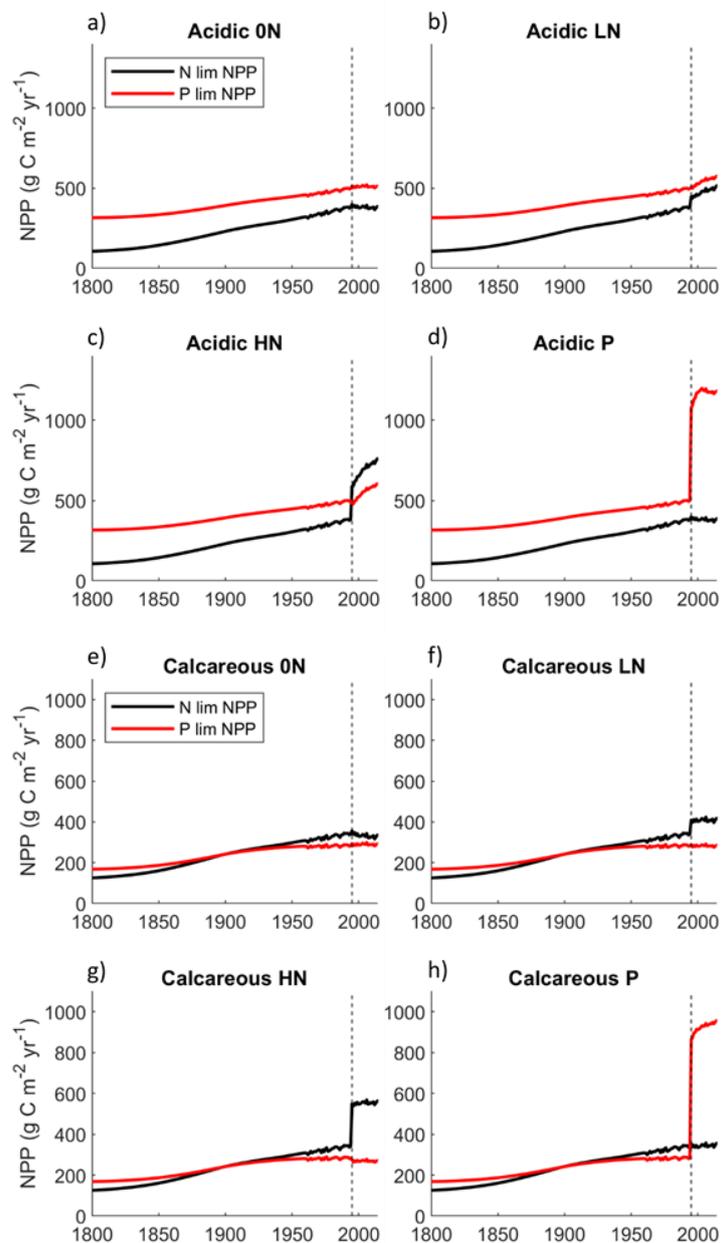


Figure 3: Plots showing the nutrient most limiting productivity for all nutrient treatments in both grasslands. The vertical dashed line is the year of first nutrient addition within the model (1995). The value of the lines represents the maximum amount of productivity attainable given the availability of N and P separately. Due to a Liebig's law of the minimum approach to plant growth, it is the lowest of the two lines that dictates the limiting nutrient of the grassland and represents actual modelled productivity. Where lines share a value, it can be considered in a state of N-P co-limitation.



408 **3.3. Modelled trends and responses to nutrient additions**

409

410 The model allows the temporal trends and responses to nutrient additions to be further explored.

411 Figure 4 provides the temporal responses for the treatments, and Figure 5 a full nutrient budget for

412 the year 2020. Full data for changes in soil C, N and P and plant biomass C pools since 1800 for both

413 grasslands are included in supplementary Table S15. All data used for determining responses of

414 biomass C and soil organic C, N and P pools to experimental nutrient additions are in supplementary

415 Tables S16 (acidic) and S17 (calcareous).

416

417 **3.3.1. Acidic grassland**

418 The modelled time series suggest that in the ON (control) treatment for the acidic grassland,

419 background levels of atmospheric N deposition between the period 1800-2020 resulted in an almost

420 four-fold increase in biomass C, a near-twofold increase in SOC and SON and increased the size of

421 the SOP pool by almost a fifth (Fig 4).

422 Since initiated in 1995, all carbon and nitrogen pools responded positively to N but not P treatments

423 (Fig 5a, c, Tables S7, S8). The LN and HN treatments further increased aboveground biomass C by

424 36.2% and 61.7% (Fig 4a) and increased the size of the total SOC pool by 11.5% and 20.6%

425 respectively (Fig 4c). Similarly, the total SON pool in the acidic grassland increased by 9.7% in the LN

426 treatment and 36.6% in the HN (Fig 4e).

427 Responses of the total SOP pool are in contrast to those of the SOC and SON pools, with LN and HN

428 slightly decreasing SOP by 4.4% and 9.1% respectively, while P addition substantially increased the

429 size of the SOP pool by 76.7% (Fig 4g). Nitrogen treatments facilitated access to SOP from both

430 subsoil and topsoil, increasing plant available P and facilitating its uptake into biomass material (Fig

431 5e, Table S9).

432



433 3.3.2. Calcareous grassland

434 Model simulations for the calcareous grassland also suggest N deposition between 1800 and 2020
435 considerably increased aboveground biomass C, SOC and SON pools (Fig. 4), but to a lesser extent
436 than in the acidic grassland. Soil organic C and SON increased by almost half and biomass C more
437 than doubled. Soil organic P accumulated at a faster rate than in the acidic grassland, increasing by
438 about a third (Fig 4, Table S15).

439 Responses of the aboveground biomass C and SOC pools in the calcareous grassland differ greatly to
440 those of the acidic, declining with N addition and increasing with P addition (Fig 4). This response
441 was ubiquitous to all C pools, with declines in subsoil, topsoil and biomass C (Fig 5b, Table S11).
442 Biomass C declined by 2.4% and 7.3% with LN and HN addition (Fig 4b) and SOC declined by 0.5%
443 and 1.4% with the same treatments (Fig 4d). Phosphorus addition increased biomass C and SOC by
444 22.0% and 6.1% respectively (Fig 4b, d).

445 Nitrogen treatments increased the size of subsoil, topsoil and available N pools, but led to small
446 declines in biomass N (Fig 5d, Table S12). The P treatment slightly reduced subsoil and topsoil SON
447 compared to the control yet increased available N and biomass N, to the extent where biomass N is
448 greater in the P than HN treatment (Fig 5d, Table S12). Total SON increased by 6.4% and 15.0% with
449 LN and HN respectively and declined by 0.2% with P treatment (Fig 4f).

450 The response of the calcareous P pools mirrors that of carbon, with declines in subsoil SOP, topsoil
451 SOP, available P and biomass P with LN and HN addition (Fig 5f, Table S13). The calcareous grassland
452 SOP pool declined by 0.2% with LN and 0.5% with HN addition, with an increase of 20.0% upon
453 addition of P (Fig 4h). The P treatment substantially increased total ecosystem P in the calcareous
454 grassland, particularly in the topsoil sorbed pool (Fig 5f, Table S13).

455

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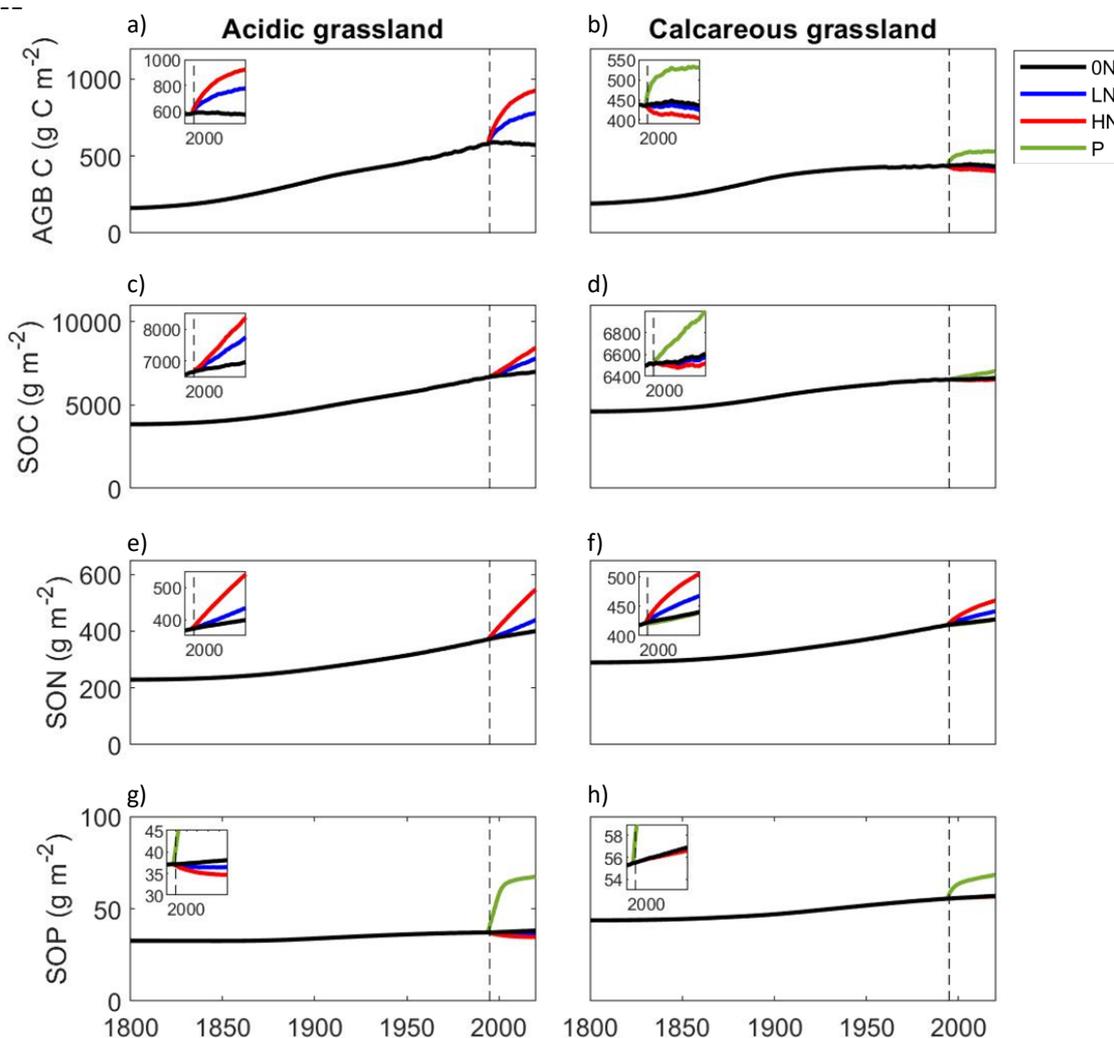


Figure 4: Time series plots of aboveground biomass C, soil organic C, N and P for the acidic (panels a, c, e and g respectively) and calcareous (panels b, d, f and h respectively) grasslands from 1800 to present day. The vertical dashed line represents the year of first nutrient addition (1995) and marks the beginning of experimental nutrient additions. The inset subplots show data from 1990 to 2020 to capture the experimental period (1995–2020) and highlight changes occurring as a result of nutrient additions rather than background N deposition. All nutrient treatments at Wardlow are represented in all panels though not all lines are visible if they do not differ from ON. Both grassland share a y axis.



459

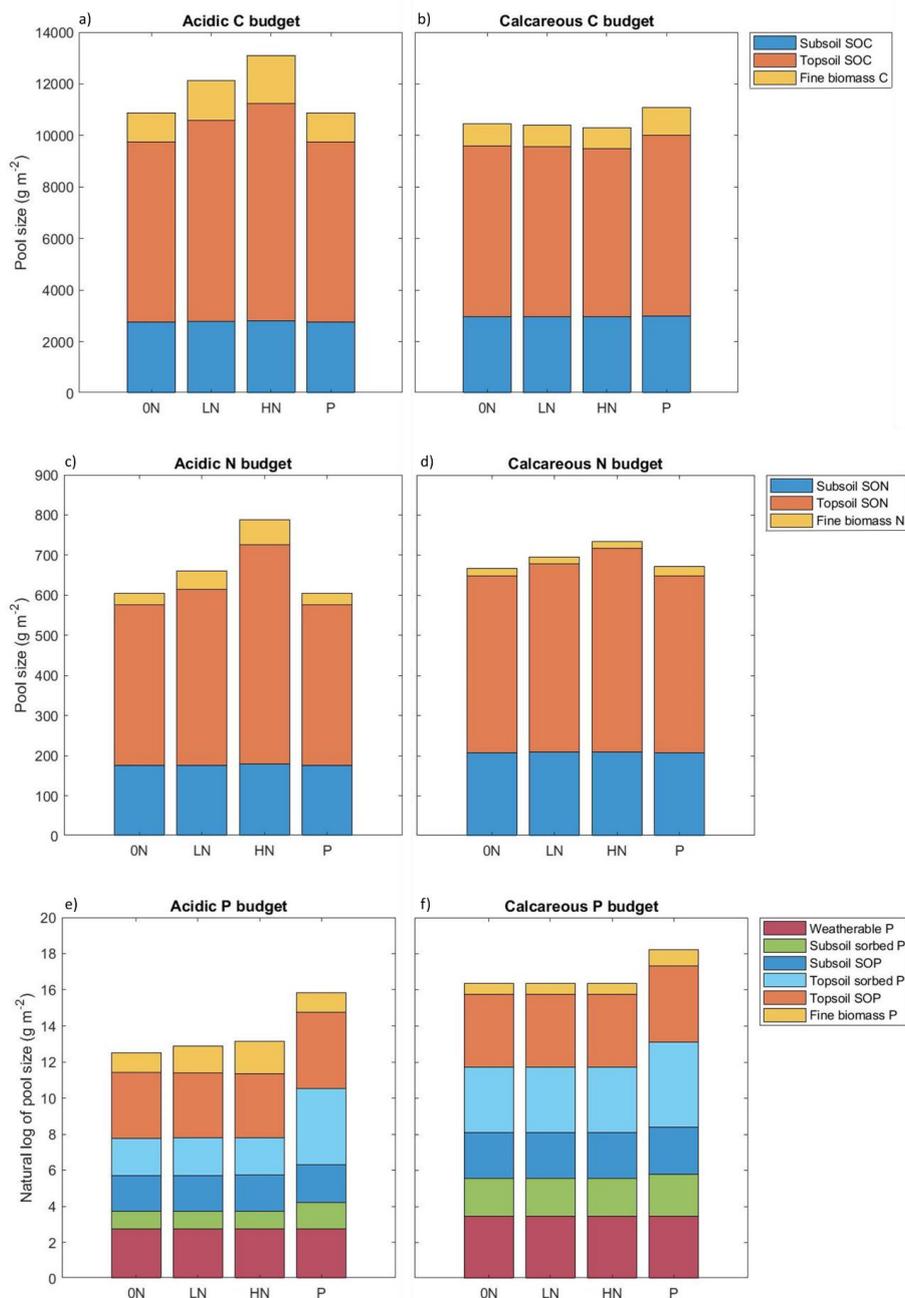


Figure 5: Modelled carbon, nitrogen and phosphorus budgets for the acidic (panels a, c and e) and calcareous (panels b, d, f) grasslands for the year 2020. Modelled sizes of C and N pools are in grams per metre squared, and P pools are presented as \log_n grams per metre squared. Temporary pools such as available N and P and fixed N are not presented here to avoid ‘double counting’ in other pools and wood litter C, N and P are not presented due to their negligible sizes.



460 **4. Discussion**

461 **4.1. Summary of findings**

462 This is the first instance in which N14CP, and to the best of our knowledge; any other integrated C-N-
463 P cycle model, has explicitly modelled N-P co-limited ecosystems and investigated their responses to
464 N deposition and additional nutrient treatments.

465 The model suggests that the acidic grassland was characterised by high access to organic P, with
466 comparatively low inorganic P availability, whereas the calcareous grassland was the opposite, with
467 low organic and high inorganic P availability. The selected combinations of $P_{\text{CleaveMax}}$ and P_{Weath0}
468 resulted in responses to nutrient addition consistent with N limitation in the modelled acidic, and P
469 limitation in the modelled calcareous grassland. This aligned with our empirical understanding of the
470 two real grasslands with co-N-P limitation being more towards either N or P limitation.

471 The modelling highlighted the contrasting impacts of experimental nutrient treatments on these two
472 grasslands, and provided a means for decoupling the effects of deposition and experimental nutrient
473 manipulation. Most notably, the responses of plant biomass C and SOC to N and P addition were in
474 contrast to one another. In the simulations, N addition led to a small decline in biomass and SOC in
475 the calcareous grassland but a substantial increase in the acidic. Nitrogen addition caused SOP to
476 decline in both grasslands as N treatment exacerbated plant P demand, and increasing P limitation in
477 the calcareous grassland.

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483 **4.2. Simulating grassland C, N and P pools by varying plant access to P sources**

484 Although N14CP was not able to replicate a co-limited response for the acidic site, it produced
485 behaviours akin to the most dominant limiting nutrient for both grasslands across multiple variables,
486 with an average discrepancy between observed and modelled data of only 6.6% ($SE \pm 9.1$) and 1.2%
487 ($SE \pm 4.4$) for the acidic and calcareous grasslands respectively across all variables (Table S5). The
488 model's performance suggests that current C-N-P cycle models that employ a Liebig's law of the
489 minimum can provide a broad representation of multiple variables, provided the ecosystem in
490 question's limiting nutrient leans towards N or P limitation.

491 This was achievable in the case of N14CP by varying two P-cycling conditions used by the model,
492 confirming that P acquisition, of both organic and inorganic forms, is a key determinant of
493 contemporary soil carbon and nutrient stocks and flows [Davies et al., 2016b]. In addition, it
494 confirms SOP could be a valuable source of P to plants in P-stressed environments, and we
495 encourage further study aimed at quantifying SOP access by plants [Janes-Bassett et al., 2020;
496 Phoenix et al., 2020].

497 The differences between P access of the two modelled grasslands could reflect the relative
498 availability of different P sources at Wardlow. The acidic grassland forms in a hillside depression
499 where loess has accumulated, distancing the plant community from the limestone beneath. The
500 plant rooting zone of the acidic grassland is not in contact with the bedrock, so roots almost
501 exclusively occur in the presence of organic P sources which can be cleaved and utilised by plants
502 [Caldwell, 2005; Margalef et al., 2017]. Conversely, the calcareous soil rarely exceeds 10 cm depth,
503 and the rooting zone extends to the limestone beneath. This provides plants with greater access to
504 weatherable calcium phosphate [Smits et al., 2012].

505 The rate of organic P access was sufficiently high in the acidic grassland to temporarily overcome P
506 limitation induced by anthropogenic N deposition. Due to its lower $P_{\text{CleaveMax}}$, the calcareous
507 grassland was unable to meet additional P demand driven by N addition, and thus remained P-



508 limited. It should be noted that the model grossly underestimates the acidic TP observations (Fig 2d),
509 as few parameter sets were simultaneously able to simulate the magnitude of the empirical TP pool
510 and the N-limited response of the acidic grassland to nutrient manipulations. Data that distinguishes
511 between organic and inorganic forms of P would help in understanding this discrepancy.

512 N14CP has a number of mechanisms to account for N and P interdependence, meaning that in
513 principle, it is capable of simulating N-P co-limited behaviour. Indeed, we found signs of N-P co-
514 limited behaviour in both grasslands as nutrient treatment altered the limiting nutrient. Available N
515 in the calcareous grassland was marginally greater in the P than ON treatment (but less than LN and
516 HN) (Fig 5d, Table S12), suggesting plants may be using surplus P to acquire N when it becomes
517 limiting. Calcareous biomass N was also highest in the P treatment, though this reflects an absolute
518 increase in N resulting from stimulated growth, and not a substantial acquisition of N from another
519 pool (Fig 5d, Table S12). Similar behaviour was found in the modelled acidic grassland, where LN and
520 HN treatments increased N availability, promoting access to available P (Table S9) and facilitating
521 growth under N addition when it was largely P-limited (Fig 3c).

522 Nitrogen fixation remained unaffected by nutrient treatment in both grasslands (Tables S8, S12).
523 This may be an unintended outcome of another N-P interaction within N14CP, whereby N fixation is
524 downregulated by atmospheric N deposition [Gundale et al., 2013]. However, when N deposition
525 exceeds fixation (as at Wardlow), fixation is essentially nullified (as in Tables S8, S12), meaning
526 deposition becomes the sole source of N to the grassland. This in effect, removes the dependence of
527 N acquisition on P availability, and could make modelling 'true' N-P co-limitation [Harpole et al.,
528 2011] under high levels of N deposition challenging. This could be further explored by allowing N
529 fixation limits in the model to adapt to P nutrient conditions.

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532 **4.3. The limiting nutrient through time**

533 There is some evidence to suggest that modelled transitions of the limiting nutrient may be
534 representative of historical nutrient limitation at Wardlow. Recent (post 1995) strengthening of P
535 limitation (Fig 3g), transition to N limitation in the P-treated calcareous plots (Fig 3h), and transition
536 to P limitation in the acidic HN treatment (Fig 3c), are likely to be accurate representations of the
537 trends in nutrient limitations at the Wardlow grasslands.

538 Strengthening P limitation in both the acidic and calcareous grasslands under increased N input is
539 supported by observations of increased root surface phosphatase enzyme activity in LN and HN
540 treatments [Johnson et al., 1999; Phoenix et al., 2004] that indicate increased P demand.
541 Furthermore, N deposition acidifies soil [Horswill et al., 2008], potentially reducing the availability of
542 mineral P by facilitating the formation of iron and aluminium complexes which act to immobilise P
543 [Kooijman et al., 1998]. Indeed, the model simulated reductions in plant available P for the
544 calcareous grassland in response to the LN and HN treatments (Table S13), further supporting an
545 exacerbated state of P limitation.

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555 **4.4. Modelled trends and responses to nutrient additions**

556 4.4.1. Biomass C and SOC

557 Changes in plant biomass and SOC within N14CP are closely interlinked, due to detrital inputs of
558 biomass being the primary source of SOC accumulation [Davies et al., 2016b]. As such, changes in
559 SOC integrate long term trends in net primary productivity in systems where external nutrients are
560 supplied. In the acidic grassland, biomass in the ON treatment begins to decrease as a result of
561 reducing N deposition following successful legislation to decrease atmospheric N pollution in the UK
562 [Dirnbock et al., 2018]. The provision of additional N in the LN and HN treatments led to large
563 increases in biomass accumulation as the model's limiting nutrient stimulated plant growth. Despite
564 its P-limited condition under the HN treatment (Fig 3c), the acidic grassland continued to accumulate
565 biomass with N addition as the grassland's greater access to topsoil SOP (Table S9) allowed it to
566 acquire sufficient P to stimulate additional growth but not necessarily to alleviate P limitation. This is
567 consistent with the acidic grassland at Wardlow, where N treatment stimulated root surface
568 phosphatases, likely supplying more SOP to plants [D Johnson et al., 1999].

569 In the acidic grassland, LN and HN addition increased SOC almost linearly (Fig 4c). Similar increases in
570 N-limited grassland SOC under N addition have been shown, resulting from significant increases in
571 below-ground carbon input from litter, roots [He et al., 2013] and detrital inputs [Fornara et al.,
572 2013], mechanisms similar to those reported by the model. Similarly, Tipping et al. [2017] used
573 N14CP to show that N deposition onto N-limited UK ecosystems ubiquitously increased SOC storage
574 by an average of 1.2 kgCm^{-2} (c. 10%) between 1750 and 2010 [Tipping et al., 2017].

575 Biomass C and SOC in the calcareous grassland responded positively to P addition, via similar
576 mechanisms to the N-response in the acidic grassland. However, in contrast to the acidic grassland,
577 N addition caused declines in calcareous biomass and SOC, the former of which has been observed
578 at the calcareous grassland at Wardlow [Carroll et al., 2003]. Reductions in calcareous biomass C



579 (and consequently SOC) in the model are a combined result of reductions in bioavailable P (Table
580 S13), occurring via N-driven increases in stoichiometric P demand, in addition to an inability to
581 access sufficient P from the SOP pool (Table S14). Plants therefore cannot meet P demand and new
582 biomass is insufficient to replace senesced plant material, decreasing net biomass C input to the SOC
583 pool.

584 Our results are consistent with findings by Li et al. [2018], who show that N fertilisation of an N-P co-
585 limited grassland reduced SOC stocks by 5-12%, which they attribute to changes in community
586 composition toward a higher proportion of forbs, whose lower tissue C:N increases the
587 decomposability of litter input to the soil, and more rapid microbial degradation of SOC [Li et al.,
588 2018]. However, there is little consensus regarding the fate of SOC under N and P addition in
589 combination or solely P addition [Stiles et al., 2017]. Soil organic C has been found to increase with N
590 and P addition [He et al., 2013], decrease with P [Scott et al., 2015; Luo et al., 2019] and show no net
591 effect on SOM [Fornara et al., 2013].

592 In addition to affecting soil C influx, C efflux can be significantly altered by N deposition. Nutrient
593 fertilisation can lead to decreases in plant tissue C:N and C:P, [Heyburn et al., 2017], increasing the
594 relative availability of nutrients to below-ground microbes and facilitating degradation of SOM [Wild
595 et al., 2014]. Furthermore, N deposition can reduce relative abundances of soil microbes and their
596 enzymes responsible for cellulose and chitin degradation [DeForest et al., 2004; M Waldrop et al.,
597 2004; Tian et al., 2019], slowing SOC decomposition, including in P-limited soils [Tian et al., 2019].
598 Such intricate interactions between soil microbes and N-driven acidification are not detailed within
599 N14CP, therefore, our conclusion that N addition decreases P-limited SOC stocks is attributable to
600 reduced C input rather than increased C output.

601

602



603 4.4.2. SON and SOP

604 Accumulation of SON is similar for both model grasslands, though it accumulates faster under N
605 limitation, as microbial N mineralisation within the model may occur more rapidly to meet the
606 increased demand for N. Much of the additional N bypasses immobilisation processes and rapidly
607 accumulates. Consistent with this, both grasslands at Wardlow accumulated large quantities of
608 simulated N deposition (up to 89% in the calcareous and 38% in the acidic grassland) [Phoenix et al.,
609 2003]. The differences in rates of modelled versus empirical grassland SON accumulation are likely
610 due to edaphic, topographical, and hydrological differences between the empirical grasslands that
611 the model cannot replicate [Phoenix et al., 2003], or a representation of an initial versus long-term
612 response.

613 Nitrogen addition increases plant demand for P and can shift ecosystems toward a state of P
614 limitation or increase the severity of limitation where it already exists [Menge and Field, 2007; An et
615 al., 2011; Goll et al., 2012]. Consistent with this, both simulated grasslands saw SOP decline with LN
616 and HN treatment, worsening P limitation in the calcareous grassland, and depleting the SOP pool in
617 the acidic. While SOP declined in both grasslands, the responses of available and biomass P to
618 nutrient treatments differed markedly between the grasslands. Due to the higher rate of $P_{\text{CleaveMax}}$ in
619 the acidic grassland, more P accumulated in the plant-available pool and hence P does not become
620 the limiting factor under N treatments (Table S9). Conversely, available and biomass P decline under
621 LN and HN addition in the calcareous grassland (Table S13), highlighting how calcareous $P_{\text{CleaveMax}}$
622 capability is insufficient to meet increased P demand.

623



624 **5. Conclusions**

625 We have shown that by varying two P-acquisition parameters within N14CP, we can account for
626 contrasting responses of two N-P co-limited grasslands to long-term nutrient manipulation with
627 reasonable accuracy. This suggests that current measures to account for co-limitation within the
628 model are to some extent sufficient and widely applicable, at least to N-P co-limited ecosystems that
629 are close to N or P limitation. Flexible organic P access allowed the modelled acidic grassland to
630 acquire sufficient P to match the available N from chronic deposition and prevent ‘anthropogenic P
631 limitation’. However, the model suggests that this is an unsustainable strategy, as the SOP pool
632 rapidly degrades, and if N additions are sustained, P limitation becomes likely. Conversely in the
633 calcareous grassland, which was less able to access organic P, additional N provision exacerbated
634 pre-existing P limitation.

635 We further show that anthropogenic N deposition since the onset of the industrial revolution has
636 had a substantial impact on the C, N and P pools of both the acidic and calcareous grasslands, to the
637 extent where almost half of contemporary soil carbon and nitrogen in the model could be from, or
638 caused by, N deposition. Experimental N and P addition had contrasting impacts on the simulated
639 grasslands. In the acidic grassland, N treatment stimulated plant access to soil organic P pools,
640 promoting plant growth and soil carbon sequestration. However, in the calcareous grassland, further
641 N addition simultaneously increased plant P demand and reduced its availability, decreasing plant
642 carbon input to the soil and leading to degradation of soil carbon. Our work therefore suggests that
643 as N deposition shifts more ecosystems toward a state of P limitation or strengthens it where it
644 already occurs [Goll et al., 2012], we may see reductions in sequestration to the point where
645 declines in grassland SOC stocks - one of our largest and most labile carbon pools – may occur.

646



647 *Data availability:* Data archiving is underway with the NERC's Environmental Information Data
648 Centre (EIDC) and a DOI will be available once this process is complete. All data to be archived is
649 present in the supplementary information for review purposes.

650

651 *Author contributions:*

652 CRT: Conceptualisation, data curation, formal analysis, investigation, methodology, project
653 administration, software, validation, visualisation, writing – original draft preparation, writing –
654 review and editing

655 VJB: Conceptualisation, formal analysis, investigation, methodology, supervision, software, writing –
656 review and editing

657 GKP: Conceptualisation, methodology, funding acquisition, project administration, resources,
658 supervision, writing – review and editing

659 BK: Investigation, methodology, supervision, writing – review and editing

660 IPH: Funding acquisition, methodology, resources, supervision, writing – review and editing

661 JD: Conceptualisation, formal analysis, investigation, resources, methodology, supervision, project
662 administration, software, writing – review and editing

663

664 *Competing interests:* The authors declare that they have no competing interests.

665

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673

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681



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