



1 Spring distribution of shelled pteropods across the Mediterranean 2 Sea

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11 12 Abstract

13
14 Shelled pteropods represent an excellent sentinel for indicating exposure to ocean acidification
15 (OA). Here, for the first time, we characterise spring pteropod distribution throughout the
16 Mediterranean Sea, a region that has been identified as a climate change hot-spot. The presence
17 of a west-east natural biogeochemical gradient makes this region a natural laboratory to
18 investigate how the variability in environmental parameters may affect pteropod distribution.
19 Results show that pteropod abundance is significantly higher in the eastern Mediterranean Sea
20 where there is a higher aragonite saturation state (Ω_{ar}), showing that distribution is positively
21 correlated with Ω_{ar} . We also observed a resilience of pteropods to higher temperatures and low
22 nutrient conditions, including phosphorous limitation. The higher abundance of pteropods in
23 ultra-oligotrophic conditions (eastern Mediterranean Sea) suggests that this organism can play
24 an important role as the prime calcifying zooplankton within specific oligotrophic regions.

25 26 1. Introduction

27
28 Thecosome pteropods are holoplanktic mollusks found in all major world oceans (Bednaršek
29 et al., 2012). This organism plays an important role in the biogeochemical cycle (Manno et al.,
30 2019) and in ocean food webs as bacterivores, (predominantly) herbivores, and as prey for
31 higher trophic levels (Conley et al., 2018). Pteropods are very susceptible to changes in
32 carbonate saturation state (Ω) due to their aragonite shell, which is a comparatively highly
33 soluble form of calcium carbonate (Mucci et al., 1989), and they therefore represent an
34 excellent sentinel for indicating the impacts of ocean acidification (OA) (Bednaršek et al.,
35 2016; Manno et al., 2017).



36

37 In the Mediterranean, pteropod species diversity, abundance and distribution are currently
38 poorly documented. The studies that have incorporated information on pteropods in the
39 Mediterranean Sea focus mostly on whole zooplankton communities, with data collected via
40 different methods (i.e. net mesh sizes, collection depth) and in different regions (i.e. Ligurian,
41 Tyrrhenian, Adriatic, Balearic, Ionian, Cretan and Levantine Seas) (Andersen, 2002; Andersen
42 et al., 2004; Batistić et al., 2004; Fernández de Puellas et al., 2007; Mazzocchi et al., 1997).
43 There are only two studies within the Mediterranean that focus solely on pteropod ecology
44 (Howes et al., 2015; Manno et al., 2019) and both are limited to very restricted geographical
45 regions. To our knowledge, there is no study on pteropod abundance and distribution across
46 the Mediterranean Sea covering the whole basin and relatively large biogeochemical gradients.
47 This is limiting our knowledge of this important calcifying zooplankton.

48

49 The Mediterranean Sea has been identified as a climate change hot-spot (Giorgi, 2006) that is
50 particularly responsive and vulnerable to ecosystem changes (Lazzari et al., 2014). This region
51 is undergoing rapid changes as a result of climatic and non-climatic forcing (Cramer et al.,
52 2018) and is experiencing an increase in temperature that is exceeding global trends with a
53 current (as of 2018) annual mean temperature of 1.4°C above late-nineteenth-century level
54 (Cramer et al., 2018). Sea temperatures are expected to rise by 1.5-2°C by the end of this
55 century, at a rate faster than the global average (Lazzari et al., 2014). Sea surface pH is
56 predicted to decrease in line with the global average (approximately 0.3 to 0.4 units by 2100)
57 (Flecha et al., 2015b; Geri et al., 2014; Kapsenberg et al., 2017) or even exceed the global
58 average decrease (Hassoun et al., 2015). The impacts of climate change on marine systems will
59 be diverse and complex, with predicted disruptions to population dynamics, geographical
60 distributions and ecosystem functioning, as well as losses in biodiversity and species richness
61 (Bulling et al., 2010; Harley et al., 2006; Lacoue-Labarthe et al., 2016). It is essential to
62 improve our knowledge of key communities that are likely vulnerable to these changes in order
63 to determine how organisms and communities will respond to ocean conditions under climate
64 change.

65

66 The Mediterranean Sea has distinct biogeochemical regions that cross natural environmental
67 gradients, with the shallow Strait of Sicily splitting the Mediterranean into east and west
68 (Rohling et al., 2009). The eastern Mediterranean is characterised by higher temperatures and
69 salinities than the western basin, which consists of Atlantic water entering from the Gibraltar



70 Strait that are modified moving eastward (Rohling et al., 2009). Surface-water circulation is
 71 mainly driven by thermohaline forcing as well as wind stress (Robinson and Golnaraghi, 1994).
 72 Using data collected from the MedSeA cruise (2013), the average Ω_{ar} (saturation state of
 73 aragonite) in the top 200m of the water column gradually increases from approximately 2.7 in
 74 the Atlantic to approximately 3.6 in the Eastern Mediterranean, meaning these waters are
 75 saturated with respect to aragonite. Phosphate and nitrate have higher concentrations in the
 76 west of the Mediterranean, with a sharp decrease moving to the east of Mediterranean, which
 77 is typified by phosphorus limitation and low concentrations of nitrate (oligotrophic system)
 78 (Krom et al., 1991). These stark changes in marine environmental parameters from west to east
 79 make the Mediterranean Sea a natural laboratory to investigate how the variability in
 80 environmental factor is affecting species distributions.

81
 82 This study aims to investigate pteropod distribution across large spatial scales, such as the
 83 Mediterranean Sea, which has been identified as a gap within pteropod research, particularly
 84 in relation to understanding how these populations will respond to climate change (Bednaršek
 85 et al., 2016). In this region, the presence of a West-East natural environmental gradient enable
 86 us to investigate the interaction between pteropod distribution and environmental parameters
 87 during the spring season. We also present the relationship between pteropods and another major
 88 group of planktic marine calcifier, foraminifera (single-celled, calcareous zooplankton).
 89 Foraminifera were collected during the same research cruise campaign (Mallo et al., 2017) and
 90 are therefore directly comparable with this study on thecosome pteropods.

91

92 **2. Materials and methods**

93

94 Samples were collected from the Mediterranean Sea during the MedSeA (Mediterranean Sea
 95 Acidification in a Changing Climate) cruise from May 2nd to June 3rd, 2013 (Fig. 1) (Ziveri and
 96 Grelaud, 2015). The research cruise covered the majority of Mediterranean sub-basins, starting
 97 from the Atlantic Ocean crossing the western basin and moving through the Levantine Basin
 98 in the east, and then again starting from the north-eastern Ionian Sea, moving to the northern
 99 Algero-Provençal basin in the east (Fig. 1).

100

101 **2.1 Hydrological and chemical collection analyses**

102



Temperature, salinity, oxygen and fluorescence were obtained from the correspondent conductivity-temperature-depth (CTD) stations using an ITS-90 and an oxygen sensor SBE 43 and considering the upper 200m towing depth. The overall accuracy for temperature is $\pm 0.001^{\circ}\text{C}$ and ± 0.0003 for salinity. To determine the seawater carbonate system, samples for total alkalinity (AT) and dissolved inorganic carbon (DIC) were collected from the top 200m of the water column (~5m, 10m, 20m, 40m, 50m and every 25m thereafter until 200m) (see Goyet et al. (2015)). Methods for the analysis of water chemistry (total alkalinity and dissolved inorganic carbon) have been described in Gemayel et al. (2015; Hassoun et al., 2015). Ocean chemistry data were input into the program CO2sys to calculate pH, $p\text{CO}_2$, aragonite saturation (Ω_{ar}) and $[\text{CO}_3^{2-}]$ using the equilibrium constants of Mehrbach et al. (1973) refitted by Dickson and Millero (1987). Photosynthetically active radiation (PAR) was measured at the beginning of each tow. Surface chlorophyll *a* concentration was obtained from MODIS Aqua L2 satellite data (NASA Goddard Space Flight Centre, 2013; Fig. 1). The nutrient concentrations (phosphate $[\text{PO}_4]$ and nitrate $[\text{NO}_3]$) were obtained using OGS (Italian National Institute of Oceanography and Experimental Geophysics) and analysed with a Bran+Luebbe3 AutoAnalyzer (see Grasshoff et al. 1999) and D'Amario et al. (2017) for a detailed methodology of the nutrient analysis).

2.2 Pteropod sample collection and analyses

Sampling for pteropods was conducted using BONGO nets (mesh size $150\mu\text{m}$, 40cm diameter) from the surface to approximately 200m depth. A flow meter attached to the ring of the net was used to determine the volume filtered through the net. Please refer to Supplementary Table 1 for information pertaining to the date, time, location, environmental parameters and volume of water filtered in the plankton tow for each sampling station. From these tow samples, pteropod abundances were determined for each station. Samples were stored in 500ml polycarbonate bottles and kept at 4°C in the dark. pH was measured in all the samples, at the beginning, during and the end of the storing period to ensure that the state of the pteropod shells were not affected by the preservation technique. Pteropod abundance was determined for each station and species were identified and counted using a Leica z16 APO binocular light microscope. Pteropod abundance within the water column was calculated as individuals per cubic meter (ind. m^{-3}). Pteropods were grouped into four target families: Heliconoididae, Limacinidae, Cavoliniidae and Cressedai; and further into seven target species: *Heliconoides inflata*, *Limacina trochiformis*, *Limacina bulimoides*, *Cavolina inflexa*, *Creseis acicula*,



137 *Creseis concia* and *Styliola subula*. The online plankton portal (www.planktonportal.org) was
 138 used to aid in the identification of pteropods to species level.

139

140 2.3 Statistical methods

141

142 All environmental parameters used in the analyses were averaged from 5-200m depth.
 143 Using the environmental parameters (temperature, salinity, oxygen, fluorescence, NO₃, PO₄,
 144 pH, pCO₂, CO₃²⁻ and Ω_{ar}) a principle component analysis was conducted (PCA, varimax
 145 rotation). After an initial analysis, PAR was removed as it did not significantly contribute to
 146 the variation of environmental parameters. Based on the PCA results, a Kruskal-Wallis Test
 147 was used to determine any significant differences in total and individual species abundance
 148 between western stations (1-7a and 19-22) and eastern stations (9-‘16-18’). A parsimonious
 149 canonical correspondence analysis (CCA) was used to determine the significant environmental
 150 parameters affecting pteropod species composition. Using Factor 1 and Factor 2 values from
 151 the PCA and the environmental parameters for each station, alpha values were obtained to
 152 conduct a binary logistic regression model (BLRM) which predicts the odds of having a low-
 153 density station (classified as stations with >1 ind. m⁻³ pteropod). Stations were binned
 154 according to the total density (13 high density stations, 7 low density stations) and were
 155 modelled against Factor 1 and 2, as well as each environmental variable. Pearson’s correlation
 156 coefficients were calculated to determine if there were any significant relationships between
 157 the environmental parameters and total and individual species abundance. The CCA was
 158 analysed using R version 3.6.0 and all other statistical analyses were performed using IBM
 159 SPSS v23.

160

161 3. Results

162

163 The mean absolute abundance of pteropods collected in the Mediterranean Basin was $1.27 \pm$
 164 $1.62(\text{SD})$ ind. m⁻³ (STable 2). The highest abundance was recorded in the Otranto Strait
 165 toward the southern end of the Adriatic Sea with 5.21 ind. per m⁻³ (STable 2, Fig. 2). There
 166 were no pteropods sampled in the Northern Alguero-Balear region (station 20), and the lowest
 167 mean standing stock was recorded at the Catalan-Balearic Station, off the coast of Spain
 168 (STable 2, Fig. 2). The eastern Mediterranean Sea had higher overall pteropod abundances with
 169 an average abundance of 2.13 ind. m⁻³, which was approximately 5x greater than the average



170 pteropod abundance in the western basin (0.42 ind. m⁻³), where the lowest abundances were
 171 recorded.

172

173 A total of five different pteropod families and 11 species were identified. Limacinidae was the
 174 most abundant family (77.7%), which dominate the eastern part of the Mediterranean basin,
 175 followed by Creseidae (15.4%) and Cavolinidae (6.8%). *Heliconoides inflata* was the most
 176 abundant species in the Mediterranean (29.4%) and recorded at all stations, followed by *L.*
 177 *bulimoides* (23.8%) and *L. trochiformis* (23.2%). The three species in Limacinidae followed a
 178 similar distribution pattern with high abundances in the Ionian Sea and the Antikythera Strait
 179 and lower abundances to the south of the Balearic Sea. *Cavolina inflexa* (6.8%), *C. acicula*
 180 (7.7%) and *Creseis conica* (6.8%) all presented the highest abundance in the Eastern Ionian
 181 Sea while *S. subula* (0.68%) abundance was at a maximum in the east of Levantine basin
 182 (STable 2). Specimens within the target families that were unidentifiable to species level made
 183 up 1.4% of the total abundance.

184

185 3.1 Relationship with environmental parameters

186

187 The Kruskal-Wallis Test confirmed significantly greater total pteropod abundance and
 188 abundance of *H. inflata*, *L. trochiformis*, *L. bulimoides* and *C. inflexa* in eastern stations
 189 (STable 3). The results of the Pearson's Correlation show that total pteropod abundance over
 190 the whole Mediterranean Sea was positively correlated with Ω_{ar} , O₂, pH, salinity and
 191 temperature and negatively correlated with NO₄, PO₃ and *p*CO₂ (STable 3).

192

193 The PCA performed on the environmental parameters reveals separation between stations in
 194 the east and the west of the Mediterranean Basin (Fig. 3). The western basin is characterised
 195 by greater nutrient content, lower temperatures and salinities (Fig. 3). The eastern basin is
 196 characterised by ultra-oligotrophic conditions as well as higher temperatures, salinities and
 197 aragonite saturation (STable 1). Factor 1 and 2 together explain approximately 77.55% of the
 198 variation in the environmental parameters. Factor 1 explains approximately 58.08% of
 199 variation in environmental parameters. This factor exhibited positive loadings for temperature
 200 and O₂ (and to a lesser extent Ω_{ar} and salinity) and a negative loading for the nutrients PO₄ and
 201 NO₃ (and to a lesser extent fluorescence). Factor 2 explains approximately 19.47% of variation
 202 in environmental parameters. This factor is characterised by positive loadings for pH, salinity
 203 and Ω_{ar} , and a negative loading for *p*CO₂. Factor 2 is interpreted as the carbonate system in



the Mediterranean Sea, with a higher pH and Ω_{ar} in the east compared to the west. Abundance scores for total pteropod and individual species counts were overlaid onto the PCA plot (Fig. 3 and SFig. 1). Stations in the eastern Mediterranean generally have a greater total abundance of pteropods and are more positively associated with Factor 1 and 2; namely, greater temperatures, salinities, pH levels, Ω_{ar} and O_2 (Fig. 3). In general, the western stations have a lower total abundance and are more positively associated with pCO_2 , nutrients and greater nutrient availability.

The parsimonious CCA reveals that temperature, pH, Ω_{ar} and salinity affect pteropod community composition in the Mediterranean and that species group by family, with species belonging to Limacinidae showing a positive relationship with pH, Ω_{ar} and salinity, and *H. inflata* showing a strong positive correlation with temperature (Fig. 3). Species in Creseidae also group together yet do not correlate positively with temperature, pH, Ω_{ar} and salinity (Fig. 3), reflecting the lower proportion of total abundance for this family in the east of the Mediterranean. *Cavolina inflexa* does not group with the other families, and the results do not indicate a positive correlation with any environmental factor.

The BLRM predicts (90% CI) that an increase in temperature and, to a much greater extent Ω_{ar} , will increase the odds of having a station with a total abundance >1 ind. m^{-3} . Increasing nitrate will reduce the odds of having a station with a total abundance >1 ind./ m^{-3} (STable 3). These results are likely influenced by the higher abundance of family Limacinidae in the east of the Mediterranean Sea. The BLRM correlates well with the CCA results in that pteropod community composition is driven by a similar suite of environmental factors, namely Ω_{ar} , pH, salinity, and to a lesser extent, temperature.

228

229 **4. Discussion**

230

This study shows that springtime shelled pteropod distribution in the Mediterranean Sea is positively correlated with Ω_{ar} saturation across the whole basin. We also found that the variability in Ω_{ar} influences community composition and that changes to the level of Ω_{ar} is likely to cause changes to pteropod populations. These results are driven by the positive relationship of Ω_{ar} with the most abundant family Limacinidae (76.4% of total abundance) and suggest that overall, pteropod distribution is strongly dependent on Ω_{ar} distribution.

237



238 The energetic costs associated with calcification indicate that calcifying organisms will spend
 239 more energy on building their shells under low Ω_{ar} conditions (Waldbusser et al., 2015). A
 240 recent study found that under reduced Ω_{ar} (control - 2.8; reduced - 2.1 Ω_{ar}), calcification in *H.*
 241 *inflatus* decreased and metabolic rate increased (Moya et al., 2016), revealing that even in
 242 oversaturated conditions, variability in Ω_{ar} can generate stress. Small variations in over
 243 saturated aragonite conditions may incur an energetic cost (however small), which can result
 244 in variations in pteropod geographical distribution. Species distribution often reflects
 245 environments containing their optimal parameters in order to optimise energy expenditure
 246 (Parsons, 1990) and our results show that pteropods have a preference for areas with higher
 247 Ω_{ar} in which they will likely devote less energy to building their shell and be under less
 248 metabolic stress. A study in the Mediterranean at CO_2 vents in the Gulf of Naples investigated
 249 pteropod calcification and abundance along an Ω_{ar} gradient (1.9-2.7) (Manno et al., 2019).
 250 Similar to our study, there was a positive correlation of *Creseis conica* abundance with Ω_{ar} in
 251 oversaturated conditions. The focus of this paper however, was to a very small geographic
 252 region in the Tyrrhenian Sea and involved a non-static system (CO_2 vent), whereas in the
 253 present study, we utilised stronger gradients of multiple environmental parameters over a large
 254 geographic scale. The findings in our study indicate that springtime pteropod distribution in
 255 the Mediterranean is strongly dependent on Ω_{ar} distribution.

256

257 The Ω_{ar} in this study ranged from 2.68 in the west of the Mediterranean to 3.61 in the east (Fig.
 258 4A). It is difficult to determine seasonal or inter-annual variability for Ω_{ar} within the
 259 Mediterranean, as studies on carbonate chemistry are either at a single time point over
 260 latitudinal gradients (Álvarez et al., 2014; Millero et al., 1979; Schneider et al., 2007) or a mid
 261 to long-term data series at a single coastal location (Ingrosso et al., 2016; Kapsenberg et al.,
 262 2017). In the case of the long-term data series by Ingrosso et al. (Ingrosso et al., 2016) in the
 263 Gulf of Trieste in the north Adriatic Sea and Kapsenberg et al. (Kapsenberg et al., 2017) in the
 264 Ligurian Sea, samples for Ω_{ar} analysis were collected from shallow coastal locations (~15m
 265 and ~75m respectively) and as such are not reflective of the open sea environments where
 266 sampling in our study was undertaken. Seasonal variation of Ω_{ar} in open ocean waters such as
 267 the North Atlantic (upper 50m and 100m depths; Ω_{ar} 0.4-0.6) (Kim et al., 2015) is likely more
 268 reflective of the seasonal fluctuations experienced in deeper Mediterranean waters, and as such,
 269 the seasonal variability in Ω_{ar} that pteropods are experiencing is not in the range of the
 270 geographical variability within the Mediterranean Sea (up 200m). As Ω_{ar} is the biggest driver
 271 of pteropod distribution, we suggest that the general population trend in this study will be



272 maintained throughout the year, as the eastern Mediterranean will be consistently higher in Ω_{ar}
 273 and therefore more energetically favourable than the west in terms of calcification.

274

275 We found that to a lesser extent, temperature and NO_3 are also drivers of spring pteropod
 276 distribution within the Mediterranean sea, with increases in temperature and decreases in NO_3
 277 predicted by the BLRM to slightly increase the odds of higher pteropod abundance. In the
 278 Mediterranean, sea surface temperatures vary by about 10°C over the year, with the north-
 279 western Mediterranean have a winter average around 10°C and a summer average of 21°C ,
 280 while the south-eastern Mediterranean winter average is around 15°C and the summer around
 281 26°C (Rohling et al., 2009) with a consistent west to east gradient (Fig. 4C). Most of the
 282 recorded temperatures in this study will be experienced at some point over the course of the
 283 year for all pteropod populations, except for the eastern Mediterranean during summer, which
 284 is warmer than the west year-round (Rohling et al., 2009). In light of the ongoing warming of
 285 the Mediterranean Sea, further studies should address the potential implication connected to
 286 the observed positive relationship between pteropod abundance and temperature.

287

288 For NO_3 , there is a well-known east-west gradient of oligotrophy in the Mediterranean (Fig.
 289 4B) that does not have much annual variability (Pasqueron de Fommervault et al., 2015). In our
 290 study, pteropod abundance was $\sim 5\times$ greater in the ultra-oligotrophic eastern Mediterranean.
 291 The negative correlation of pteropod abundance with favourable nutrient conditions (in
 292 particular nitrate) is unexpected, as pteropod global distribution has been previously correlated
 293 with high productivity and nutrient content (Bednaršek et al., 2012; Burrige et al., 2017). A
 294 study of pteropod distribution along a longitudinal gradient in the Atlantic Ocean (Burrige et
 295 al., 2017) found that areas with the highest biomass corresponded to areas with the highest
 296 concentrations of chlorophyll *a* (Huot et al., 2007). However, because in Burrige's study
 297 chlorophyll *a* was correlated with temperature, it is unclear which of the two variables was the
 298 main driver of pteropod distribution. As only two environmental parameters were investigated,
 299 Burrige's study is limited in its' ability to determine how other biologically relevant
 300 parameters for pteropods affect their distribution, such as Ω_{ar} . In the Atlantic Ocean over a
 301 range spanning more than the 45°N to 45°S in Burrige's study, the upper 200m is
 302 supersaturated with respect to aragonite (Jiang et al., 2015) and there is little variability in Ω_{ar}
 303 (Takahashi et al., 2014). If Ω_{ar} is a primary driver of pteropod distribution, as indicated by this
 304 study, when waters are supersaturated and there is little variability in Ω_{ar} , other factors such
 305 as temperature or nutrients may be secondary driving forces affecting pteropod abundance, as



seen in Burridge's study. As we found that Ω_{ar} is the main driver of pteropod distribution in the Mediterranean, it is likely that the negative correlation with NO_3 is not directly causing a negative effect on pteropod abundance and distribution. Pteropod abundances are higher in the oligotrophic and highly aragonite-saturated eastern basin, resulting in a negative correlation with NO_3 .

We are aware that this study is only a snapshot in time, which limits our ability to predict the distribution of pteropods within the Mediterranean over large temporal scales. However, this study represents the first description of pteropod distribution along the whole Mediterranean basin and their relation with relevant environmental parameters. On the basis of the observed correlation between pteropods and the natural geographical variability of Ω_{ar} , we suggest that future research should focus on long-term, Mediterranean-wide monitoring to detect potential changes of pteropod populations.

4.1 Pteropod and foraminifera interaction

Planktic foraminiferal abundance and distribution presented in Mallo et al. (2017) were compared to the pteropod data from this study. The tow samples from Mallo et al. (2017) were collected during the same cruise and within the same nets as the pteropods of the present study, allowing a direct comparison of the two groups of key planktic calcifiers.

Spring pteropod and foraminiferal abundances present an opposite pattern distribution between the west and east of the Mediterranean (Fig. 5). A clear difference in the composition of pteropod and foraminiferal communities is evident in the CCA (Fig. 4), showing foraminiferal communities are positively related to fluorescence and NO_3 , and pteropod communities positively related to temperature, pH, salinity, O_2 and Ω_{ar} . Pteropod abundance is distinctly greater in the eastern ($\bar{x}=2.13$ ind. m^{-3}) Mediterranean than in the west ($\bar{x}=0.47$ ind. m^{-3}), while foraminifera populations are characterised by higher abundances in the west ($\bar{x}=1.87$ ind. $10\ m^{-3}$) than in the east ($\bar{x}=0.96$ ind. $10\ m^{-3}$) (Fig. 5). There was a significant correlation in abundance between western Mediterranean stations ($r=-.757^*$), but not the eastern stations. Although total abundance is distinctly greater for pteropod populations across the entire Mediterranean, abundance for both organisms follows a similar pattern in the western stations (Fig. 5).



Experimental evidence indicates that planktic foraminifera are vulnerable to OA conditions under which they experience reduced calcification and an increased metabolic rate (Davis et al., 2017; Lombard et al., 2010; Manno et al., 2012). Foraminiferal distribution and abundance appear to be driven more by nutrient levels than carbonate saturation levels, as suggested in Mallo et al. (2017). The skeleton of foraminifera is made of calcite, a less soluble form of calcium carbonate than aragonite (Mucci et al., 1989; Subhas et al., 2018). Both pteropods and foraminifera are sensitive to OA due to their carbonate skeletons, however it is likely that foraminifera are less sensitive to carbonate variability due to the lower solubility of their calcite skeleton compared with aragonite. Foraminifera have been shown to be nutrient limited (Gregg and Casey, 2007; Schiebel et al., 2004), and Mallo et al. (2017) suggested that the lower foraminiferal abundance in the eastern Mediterranean (ultra-oligotrophic sector) results from reduced reproduction due to limiting food resources.

Conversely, nutrients and food availability do not appear to be a limiting factor for pteropods and their abundance is ~5x higher in the oligotrophic sector of the Mediterranean than the west. The majority of pteropods belong to the Thecosome order and have a unique feeding method involving the production of a large mucous web that is suspended in the water column, which passively entraps organic particles and motile organisms, enabling them to filter water at high rates (Conley et al., 2018). To ingest the material collected by the mucous webs, they draw the webs into their mouth via ciliary action, a feeding method that may allow them to overcome low food condition due to their ability to capture and filter through relatively large amounts of organic matter (Hamner et al., 1975). In contrast to foraminifera, pteropods are also able to actively swim (Hamner et al., 1975; Lalli and Gilmer, 1989) and potentially search for more favourable food conditions. Further, the survival of pteropod *H. helicina* (Limacinadae family) does not appear to be affected when exposed to prolonged starvation (7 days) (Busch et al., 2014). We suggest that, in the western Mediterranean, pteropods are more able to adapt to low food availability than foraminifera which may be due to a combination of both their feeding method and their ability to withstand starvation.

Thus the difference in the pattern of distribution between foraminifera and pteropods in the eastern Mediterranean may be due to the unfavourable oligotrophic conditions for foraminifera and that pteropods are more dependent upon higher Ω_{ar} than nutrient levels. In the west, with relatively higher concentrations of nutrients, Ω_{ar} is lower, accounting for the reduced abundance pteropods, with foraminiferal and pteropod abundance following the same pattern,



as a secondary factor of importance for pteropods, namely nutrients, affects their distribution here. On the other hand, in the eastern oligotrophic conditions, pteropod abundance is distinctly greater than foraminifera, as it is more strongly driven by the higher Ω_{ar} saturation, and foraminiferal and pteropod abundance no longer follow the same pattern, as the region is very nutrient poor, negatively affecting foraminiferal abundance.

There have been only a handful of studies that investigate the relationship between pteropod and foraminiferal communities. A multi-decadal study of calcareous holo-zooplankton at two sites off the coast of Southern California and Central California (1951-2008) showed no relationship between foraminifera or pteropod abundances (Ohman et al., 2009), unlike our study. In the Gulf Stream, Sargasso Sea and the Gulf of Mexico, pteropod and foraminifera densities followed a similar pattern, with density decreasing closer to oligotrophic conditions (Casey et al., 1979). In Schiebel et al. (2001), foraminifera and pteropod production in the North Atlantic was positively correlated with chlorophyll-*a*. The correlation of abundance for pteropods and foraminifera with nutrients in Casey et al. (1979) and Schiebel et al. (2001) are similar to the findings for the west of the Mediterranean in our study. These studies did not incorporate the carbonate system and our study is the first to utilise a complete spectrum of environmental parameters to investigate the relationship between pteropods and foraminifera.

5. Conclusions

The relationship of shelled Mediterranean pteropods to the natural variability in Ω_{ar} adds to the growing body of support for using pteropods as a bioindicator of changes in ocean carbonate chemistry (Bednaršek et al., 2019; Manno et al., 2017). Overall, we suggest that the higher abundance of pteropods in the eastern than the western Mediterranean sea is a combination of higher Ω_{ar} that is energetically favourable for calcification and their unique method of feeding and/or ability to withstand starvation, allowing them to fill the ecological role in the east of the Mediterranean as the prime calcifying zooplankton. With the increase in OA, the positive correlation between Ω_{ar} and pteropod distribution stresses the need for regular, long-term monitoring of pteropod populations throughout the entire Mediterranean Sea. This study provides new insight into the ecology of shelled pteropods and foraminifera and highlights that, in specific oligotrophic regions, ecological competition could also play an important role in regulating the pelagic calcifier zooplankton communities.



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409

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417

418 References

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- 420 Álvarez, M., Sanleón-Bartolomé, H., Tanhua, T., Mintrop, L., Luchetta, A., Cantoni, C.,
 421 Schroeder, K. and Civitarese, G.: The CO₂ system in the Mediterranean Sea: A basin wide
 422 perspective, *Ocean Sci.*, 10(1), 69–92, doi:10.5194/os-10-69-2014, 2014.
- 423 Andersen, V.: Zooplankton Community During the Transition from Spring Bloom to
 424 Oligotrophy in the Open NW Mediterranean and effects of wind events. 2. Vertical
 425 distributions and migrations, *J. Plankton Res.*, 23(3), 227–242, doi:10.1093/plankt/23.3.227,
 426 2002.
- 427 Andersen, V., Scotto, M., Picheral, M., François, F., Sardou, J. and Nival, P.: Vertical
 428 distributions of macroplankton and micronekton in the Ligurian and Tyrrhenian seas
 429 (northwestern Mediterranean), *Oceanol. Acta*, 21(5), 655–676, doi:10.1016/s0399-
 430 1784(98)90007-x, 2004.
- 431 Batistić, M., Kršinić, F., Jasprica, N., Carić, M., Viličić, D. and Lučić, D.: Gelatinous
 432 invertebrate zooplankton of the South Adriatic: Species composition and vertical distribution,
 433 *J. Plankton Res.*, 26(4), 459–474, doi:10.1093/plankt/fbh043, 2004.
- 434 Bednaršek, N., Možina, J., Vogt, M., Brien, C. and Tarling, G. A.: The global distribution of
 435 pteropods and their contribution to carbonate and carbon biomass in the modern ocean, *Earth*
 436 *Syst. Sci. Data*, 4(1), 167–186, doi:10.5194/essd-4-167-2012, 2012.
- 437 Bednaršek, N., Harvey, C. J., Kaplan, I. C., Feely, R. A. and Možina, J.: Pteropods on the
 438 edge: Cumulative effects of ocean acidification, warming, and deoxygenation, *Prog.*
 439 *Oceanogr.*, 145, 1–24, doi:10.1016/j.pocean.2016.04.002, 2016.
- 440 Bednaršek, N., Feely, R. A., Howes, E. L., Hunt, B. P. V., Kessouri, F., León, P., Lischka, S.,



- 441 Maas, A. E., McLaughlin, K., Nezlin, N. P., Sutula, M. and Weisberg, S. B.: Systematic
 442 Review and Meta-Analysis Toward Synthesis of Thresholds of Ocean Acidification Impacts
 443 on Calcifying Pteropods and Interactions With Warming, *Front. Mar. Sci.*, 6, 227,
 444 doi:10.3389/fmars.2019.00227, 2019.
- 445 Bulling, M. T., Hicks, N., Murray, L., Paterson, D. M., Raffaelli, D., White, P. C. L. and
 446 Solan, M.: Marine biodiversity–ecosystem functions under uncertain environmental futures,
 447 *Philos. Trans. R. Soc. B Biol. Sci.*, 365(1549), 2107–2116, doi:10.1098/rstb.2010.0022,
 448 2010.
- 449 Burrige, A. K., Goetze, E., Wall-Palmer, D., Le Double, S. L., Huisman, J. and Peijnenburg,
 450 K. T. C. A.: Diversity and abundance of pteropods and heteropods along a latitudinal gradient
 451 across the Atlantic Ocean, *Prog. Oceanogr.*, 158, 213–223,
 452 doi:10.1016/j.pocean.2016.10.001, 2017.
- 453 Busch, D. S., Maher, M., Thibodeau, P. and McElhany, P.: Shell Condition and Survival of
 454 Puget Sound Pteropods Are Impaired by Ocean Acidification Conditions, edited by G. E.
 455 Hofmann, *PLoS One*, 9(8), e105884, doi:10.1371/journal.pone.0105884, 2014.
- 456 Casey, R., Gust, L., Leavesley, A., Williams, D., Reynolds, R., Duis, T. and Spaw, J. M.:
 457 Ecological Niches of Radiolarians, Planktonic Foraminiferans and Pteropods Inferred from
 458 Studies on Living Forms in the Gulf of Mexico and Adjacent Waters, *Gulf Coast Assoc.*
 459 *Geol. Soc. Trans.*, 29, 216–223 [online] Available from:
 460 <http://archives.datapages.com/data/gcags/data/029/029001/0216.htm> (Accessed 22 March
 461 2019), 1979.
- 462 Conley, K. R., Lombard, F. and Sutherland, K. R.: Mammoth grazers on the ocean’s
 463 minuteness: A review of selective feeding using mucous meshes, *Proc. R. Soc. B Biol. Sci.*,
 464 285(1878), doi:10.1098/rspb.2018.0056, 2018.
- 465 Cramer, W., Guiot, J., Fader, M., Garrabou, J., Gattuso, J.-P., Iglesias, A., Lange, M. A.,
 466 Lionello, P., Llasat, M. C., Paz, S., Peñuelas, J., Snoussi, M., Toreti, A., Tsimplis, M. N. and
 467 Xoplaki, E.: Climate change and interconnected risks to sustainable development in the
 468 Mediterranean, *Nat. Clim. Chang.*, 8(11), 972–980, doi:10.1038/s41558-018-0299-2, 2018.
- 469 D’Amario, B., Ziveri, P., Grelaud, M., Oviedo, A. M. and Kralj, M.: Hydrology and
 470 geochemistry in the Mediterranean Sea during the MedSeA and Meteor M84/3 cruises (May
 471 2013, April 2011), *Suppl. to D’Amario, B al. Coccolithophore haploid diploid Distrib.*
 472 *patterns Mediterr. Sea can a haplo-diploid life cycle be advantageous under Clim. Chang. J.*
 473 *Plankt. Res.* 1-14, <https://doi.org/10.10.1594/PANGAEA.875923>, 2017.
- 474 Davis, C. V., Rivest, E. B., Hill, T. M., Gaylord, B., Russell, A. D. and Sanford, E.: Ocean



- 475 acidification compromises a planktic calcifier with implications for global carbon cycling,
 476 Sci. Rep., 7(2225), 1–8, doi:10.1038/s41598-017-01530-9, 2017.
- 477 Dickson, A. G. and Millero, F. J.: A comparison of the equilibrium constants for the
 478 dissociation of carbonic acid in seawater media, Deep Sea Res. Part A, Oceanogr. Res. Pap.,
 479 34(10), 1733–1743, doi:10.1016/0198-0149(87)90021-5, 1987.
- 480 Fernández de Puellas, M. L., Alemany, F. and Jansá, J.: Zooplankton time-series in the
 481 Balearic Sea (Western Mediterranean): Variability during the decade 1994–2003, Prog.
 482 Oceanogr., 74(2–3), 329–354, doi:10.1016/j.pocean.2007.04.009, 2007.
- 483 Flecha, S., Pérez, F. F., García-Lafuente, J., Sammartino, S., Ríos, A. F. and Huertas, I. E.:
 484 Trends of pH decrease in the Mediterranean Sea through high frequency observational data:
 485 indication of ocean acidification in the basin, Sci. Rep., 5(1), 16770, doi:10.1038/srep16770,
 486 2015a.
- 487 Flecha, S., Pérez, F. F., García-Lafuente, J., Sammartino, S., Ríos, A. F. and Huertas, I. E.:
 488 Trends of pH decrease in the Mediterranean Sea through high frequency observational data:
 489 Indication of ocean acidification in the basin, Sci. Rep., doi:10.1038/srep16770, 2015b.
- 490 Gemayel, E., Hassoun, A. E. R., Benallal, M. A., Goyet, C., Rivaro, P., Abboud-Abi Saab,
 491 M., Krasakopoulou, E., Touratier, F. and Ziveri, P.: Climatological variations of total
 492 alkalinity and total dissolved inorganic carbon in the Mediterranean Sea surface waters, Earth
 493 Syst. Dyn., 6(2), 789–800, doi:10.5194/esd-6-789-2015, 2015.
- 494 Geri, P., El Yacoubi, S. and Goyet, C.: Forecast of Sea Surface Acidification in the
 495 Northwestern Mediterranean Sea, J. Comput. Environ. Sci., 2014, 1–7,
 496 doi:10.1155/2014/201819, 2014.
- 497 Giorgi, F.: Climate change hot-spots, Geophys. Res. Lett., 33(8), L08707,
 498 doi:10.1029/2006GL025734, 2006.
- 499 Goyet, C., El Rahman Hassoun, A. and Gemayel, E.: Carbonate system during the May 2013
 500 MedSeA cruise, PANGAEA, doi:10.1594/PANGAEA.841933, 2015.
- 501 Grasshoff, K. and Kremling, K.: Methods of seawater analysis, 3rd ed., edited by K.
 502 Grasshoff, M. Ehrhardt, and K. Kremling, Wiley-VCH., 1999.
- 503 Gregg, W. W. and Casey, N. W.: Modeling coccolithophores in the global oceans, Deep Sea
 504 Res. Part II Top. Stud. Oceanogr., 54(5–7), 447–477, doi:10.1016/J.DSR2.2006.12.007,
 505 2007.
- 506 Hamner, W. M., Madin, L. P., Alldrege, A. L., Gilmer, R. W. and Hamner, P. P.: Underwater
 507 observations of gelatinous zooplankton: sampling problems, feeding biology, and behavior.,
 508 Limnol. Oceanogr., 20(6), 907–917, 1975.



- 509 Harley, C. D. G. G., Hughes, A. R., Hultgren, K. M., Miner, B. G., Sorte, C. J. B. B.,
510 Thornber, C. S., Rodriguez, L. F., Tomanek, L. and Williams, S. L.: The impacts of climate
511 change in coastal marine systems, *Ecol. Lett.*, 9(2), 228–241, doi:10.1111/j.1461-
512 0248.2005.00871.x, 2006.
- 513 Hassoun, A. E. R., Gemayel, E., Krasakopoulou, E., Goyet, C., Abboud-Abi Saab, M.,
514 Guglielmi, V., Touratier, F. and Falco, C.: Acidification of the Mediterranean Sea from
515 anthropogenic carbon penetration, *Deep Sea Res. Part I Oceanogr. Res. Pap.*, 102, 1–15,
516 doi:10.1016/J.DSR.2015.04.005, 2015.
- 517 Howes, E. L., Stemmann, L., Assailly, C., Irisson, J. O., Dima, M., Bijma, J. and Gattuso, J.
518 P.: Pteropod time series from the North Western Mediterranean (1967-2003): Impacts of pH
519 and climate variability, *Mar. Ecol. Prog. Ser.*, 531, 193–206, doi:10.3354/meps11322, 2015.
- 520 Huot, Y., Babin, M., Bruyant, F., Grob, C., Twardowski, M. S. and Claustre, H.: Does
521 chlorophyll a provide the best index of phytoplankton biomass for primary productivity
522 studies?, *Eur. Geosci. Union*, 4(2), 707–745 [online] Available from: [www.biogeosciences-](http://www.biogeosciences-discuss.net/4/707/2007/)
523 [discuss.net/4/707/2007/](http://www.biogeosciences-discuss.net/4/707/2007/) (Accessed 9 June 2019), 2007.
- 524 Ingrosso, G., Giani, M., Comici, C., Kralj, M., Piacentino, S., De Vittor, C. and Del Negro,
525 P.: Drivers of the carbonate system seasonal variations in a Mediterranean gulf, *Estuar.*
526 *Coast. Shelf Sci.*, 168, 58–70, doi:10.1016/j.ecss.2015.11.001, 2016.
- 527 Jiang, L. Q., Feely, R. A., Carter, B. R., Greeley, D. J., Gledhill, D. K. and Arzayus, K. M.:
528 Climatological distribution of aragonite saturation state in the global oceans, *Global*
529 *Biogeochem. Cycles*, 29(10), 1656–1673, doi:10.1002/2015GB005198, 2015.
- 530 Kapsenberg, L., Alliouane, S., Gazeau, F., Mousseau, L. and Gattuso, J.-P.: Coastal ocean
531 acidification and increasing total alkalinity in the northwestern Mediterranean Sea, *Ocean*
532 *Sci.*, 13, 411–426, doi:10.5194/os-13-411-2017, 2017.
- 533 Kim, T. W., Park, G. H., Kim, D., Lee, K., Feely, R. A. and Millero, F. J.: Seasonal variations
534 in the aragonite saturation state in the upper open-ocean waters of the North Pacific Ocean,
535 *Geophys. Res. Lett.*, 42(11), 4498–4506, doi:10.1002/2015GL063602, 2015.
- 536 Krom, M. D., Kress, N., Brenner, S. and Gordon, L. I.: Phosphorus limitation of primary
537 productivity in the eastern Mediterranean Sea, *Limnol. Oceanogr.*, 36(3), 424–432,
538 doi:10.4319/lo.1991.36.3.0424, 1991.
- 539 Lacoue-Labarthe, T., Nunes, P. A. L. D., Ziveri, P., Cinar, M., Gazeau, F., Hall-Spencer, J.
540 M., Hilmi, N., Moschella, P., Safa, A., Sauzade, D. and Turley, C.: Impacts of ocean
541 acidification in a warming Mediterranean Sea: An overview, *Reg. Stud. Mar. Sci.*, 5, 1–11
542 [online] Available from:



- 543 <https://www.sciencedirect.com/science/article/pii/S2352485515300116> (Accessed 20 April
 544 2019), 2016.
- 545 Lalli, C. M. and Gilmer, R. W.: Pelagic snails : the biology of holoplanktonic gastropod
 546 mollusks, Stanford University Press. [online] Available from:
 547 https://books.google.es/books/about/Pelagic_Snails.html?id=yIAfwz5cxPMC&redir_esc=y
 548 (Accessed 20 March 2019), 1989.
- 549 Lazzari, P., Mattia, G., Solidoro, C., Salon, S., Crise, A., Zavatarelli, M., Oddo, P. and Vichi,
 550 M.: The impacts of climate change and environmental management policies on the trophic
 551 regimes in the Mediterranean Sea: Scenario analyses, *J. Mar. Syst.*, 135, 137–149,
 552 doi:10.1016/j.jmarsys.2013.06.005, 2014.
- 553 Lombard, F., da Rocha, R. E., Bijma, J. and Gattuso, J.-P.: Effect of carbonate ion
 554 concentration and irradiance on calcification in planktonic foraminifera, *Biogeosciences*,
 555 7(1), 247–255, doi:10.5194/bg-7-247-2010, 2010.
- 556 Mallo, M., Ziveri, P., Mortyn, P. G., Schiebel, R. and Grelaud, M.: Low planktic
 557 foraminiferal diversity and abundance observed in a 2013 West-East Mediterranean Sea
 558 transect, *Biogeosciences Discuss.*, 1–31, doi:10.5194/bg-2016-266, 2017.
- 559 Manno, C., Morata, N. and Bellerby, R.: Effect of ocean acidification and temperature
 560 increase on the planktonic foraminifer *Neogloboquadrina pachyderma* (sinistral), *Polar Biol.*,
 561 35(9), 1311–1319, doi:10.1007/s00300-012-1174-7, 2012.
- 562 Manno, C., Bednaršek, N., Tarling, G. A., Peck, V. L., Comeau, S., Adhikari, D., Bakker, D.
 563 C. E., Bauerfeind, E., Bergan, A. J., Berning, M. I., Buitenhuis, E., Burridge, A. K., Chierici,
 564 M., Flöter, S., Fransson, A., Gardner, J., Howes, E. L., Keul, N., Kimoto, K., Kohnert, P.,
 565 Lawson, G. L., Lischka, S., Maas, A., Mekkes, L., Oakes, R. L., Pebody, C., Peijnenburg, K.
 566 T. C. A., Seifert, M., Skinner, J., Thibodeau, P. S., Wall-Palmer, D. and Ziveri, P.: Shelled
 567 pteropods in peril: Assessing vulnerability in a high CO₂ ocean, *Earth-Science Rev.*, 169,
 568 132–145, doi:10.1016/j.earscirev.2017.04.005, 2017.
- 569 Manno, C., Rumolo, P., Barra, M., d’Albero, S., Basilone, G., Genovese, S., Mazzola, S. and
 570 Bonanno, A.: Condition of pteropod shells near a volcanic CO₂ vent region, *Mar. Environ.*
 571 *Res.*, 143, 39–48, doi:10.1016/j.marenvres.2018.11.003, 2019.
- 572 Mazzocchi, M., Christou, E., Fragopoulou, N. and Siokoufrangou, I.: Mesozooplankton
 573 distribution from Sicily to Cyprus (Eastern Mediterranean) .1. General aspects, *Oceanol.*
 574 *Acta*, 20(3), 521–535, 1997.
- 575 Mehrbach, C., Culbertson, C. H., Hawley, J. E. and Pytkowicz, R. M.: Measurement of the
 576 apparent dissociation constants of carbonic acid in seawater at atmospheric pressure, *Limnol.*



- 577 Oceanogr., 18(6), 897–908, doi:10.4319/lo.1973.18.6.0897, 1973.
- 578 Millero, F. J., Morse, J. and Chen, C.-T.: The carbonate system in the western Mediterranean
 579 sea, Deep Sea Res. Part A. Oceanogr. Res. Pap., 26(12), 1395–1404, doi:10.1016/0198-
 580 0149(79)90007-4, 1979.
- 581 Moya, A., Howes, E. L., Lacoue-Labarthe, T., Forêt, S., Hanna, B., Medina, M., Munday, P.
 582 L., Ong, J.-S., Teyssié, J.-L., Torda, G., Watson, S.-A., Miller, D. J., Bijma, J. and Gattuso,
 583 J.-P.: Near-future pH conditions severely impact calcification, metabolism and the nervous
 584 system in the pteropod *Heliconoides inflatus*, Glob. Chang. Biol., 22(12), 3888–3900,
 585 doi:10.1111/gcb.13350, 2016.
- 586 Mucci, A., Canuel, R. and Zhong, S.: The solubility of calcite and aragonite in sulfate-free
 587 seawater and the seeded growth kinetics and composition of the precipitates at 25°C, Chem.
 588 Geol., 74(3–4), 309–320, doi:10.1016/0009-2541(89)90040-5, 1989.
- 589 Ohman, M. D., Lavaniegos, B. E. and Townsend, A. W.: Multi-decadal variations in
 590 calcareous holozooplankton in the California Current System: Thecosome pteropods,
 591 heteropods, and foraminifera, Geophys. Res. Lett., 36(18), L18608,
 592 doi:10.1029/2009GL039901, 2009.
- 593 Parsons, P. A.: The metabolic cost of multiple environmental stresses: Implications for
 594 climatic change and conservation, Trends Ecol. Evol., 5(9), 315–317, doi:10.1016/0169-
 595 5347(90)90089-V, 1990.
- 596 Pasqueron de Fommervault, O., Mangin, A., Serra, R., Migon, C., Claustre, H., Lavigne, H.,
 597 Ribera, M., Prieur, L., Taillandier, V., Claustre, H., Elo, H., Lavigne, I., Schmechtig, C.,
 598 Poteau, A., Leymarie, E., elie Dufour, A., Besson, F. and Obolensky, G.: Seasonal variability
 599 of nutrient concentrations in the Mediterranean Sea: Contribution of Bio-Argo floats, J.
 600 Geophys. Res. Ocean., 120(12), 8528–8550, doi:10.1002/2015JC011103, 2015.
- 601 Robinson, A. R. and Golnaraghi, M.: The Physical and Dynamical Oceanography of the
 602 Mediterranean Sea, in Ocean Processes in Climate Dynamics: Global and Mediterranean
 603 Examples, pp. 255–306, Springer Netherlands, Dordrecht., 1994.
- 604 Rohling, E. J., Abu-Zied, R., Casford, J. S. L., Hayes, A. and Hoogakker, B.: The Marine
 605 Environment: Present and Past, in The Physical Geography of the Mediterranean, pp. 33–67.,
 606 2009.
- 607 Schiebel, R., Waniek, J., Bork, M. and Hemleben, C.: Planktic foraminiferal production
 608 stimulated by chlorophyll redistribution and entrainment of nutrients, Deep Sea Res. Part I
 609 Oceanogr. Res. Pap., 48(3), 721–740, doi:10.1016/S0967-0637(00)00065-0, 2001.
- 610 Schiebel, R., Zeltner, A., Treppke, U. F., Waniek, J. J., Bollmann, J., Rixen, T. and



611 Hemleben, C.: Distribution of diatoms, coccolithophores and planktic foraminifers along a
 612 trophic gradient during SW monsoon in the Arabian Sea, *Mar. Micropaleontol.*, 51(3–4),
 613 345–371, doi:10.1016/J.MARMICRO.2004.02.001, 2004.
 614 Schneider, A., Wallace, D. W. R. and Körtzinger, A.: Alkalinity of the Mediterranean Sea,
 615 *Geophys. Res. Lett.*, 34(15), 1–5, doi:10.1029/2006GL028842, 2007.
 616 Subhas, A. V., Rollins, N. E., Berelson, W. M., Erez, J., Ziveri, P., Langer, G. and Adkins, J.
 617 F.: The dissolution behavior of biogenic calcites in seawater and a possible role for
 618 magnesium and organic carbon, *Mar. Chem.*, 205, 100–112,
 619 doi:10.1016/j.marchem.2018.08.001, 2018.
 620 Takahashi, T., Sutherland, S. C., Chipman, D. W., Goddard, J. G., Ho, C., Newberger, T.,
 621 Sweeney, C. and Munro, D. R.: Climatological distributions of pH, pCO₂, total CO₂,
 622 alkalinity, and CaCO₃ saturation in the global surface ocean, and temporal changes at
 623 selected locations, *Mar. Chem.*, 164, 95–125, doi:10.1016/j.marchem.2014.06.004, 2014.
 624 Waldbusser, G. G., Hales, B., Langdon, C. J., Haley, B. A., Schrader, P., Brunner, E. L.,
 625 Gray, M. W., Miller, C. A. and Gimenez, I.: Saturation-state sensitivity of marine bivalve
 626 larvae to ocean acidification, *Nat. Clim. Chang.*, 5(3), 273–280, doi:10.1038/nclimate2479,
 627 2015.
 628 Ziveri, P. and Grelaud, M.: Physical oceanography during Ángeles Alvariño cruise
 629 MedSeA2013, PANGAEA., 2015.

Figure Legends

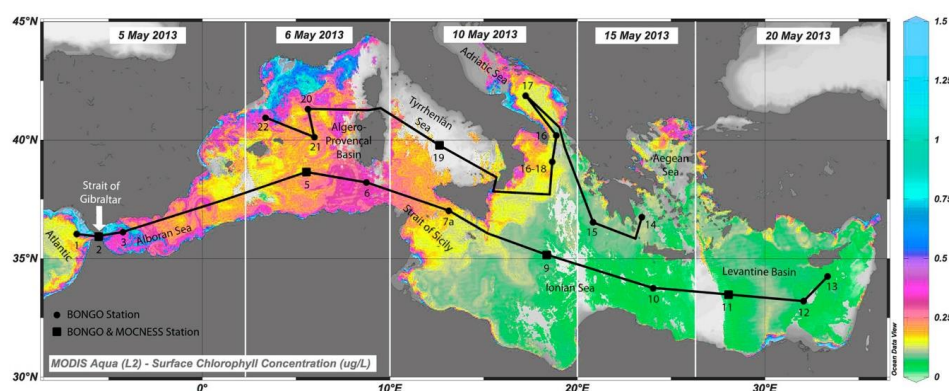




Figure 1. The sampled stations with BONGO nets are indicated by dots. The numbers in the picture represent the station codes: First leg: 1 to 13, second leg: 14 to 22. The colour scale represents the values of surface chlorophyll concentration (in $\mu\text{g/l}$), retrieved from *MODIS Aqua (L2)*, from the closest day to the start of the first transect. Figure made using Ocean Data View. Also in Mallo et al. (2017).

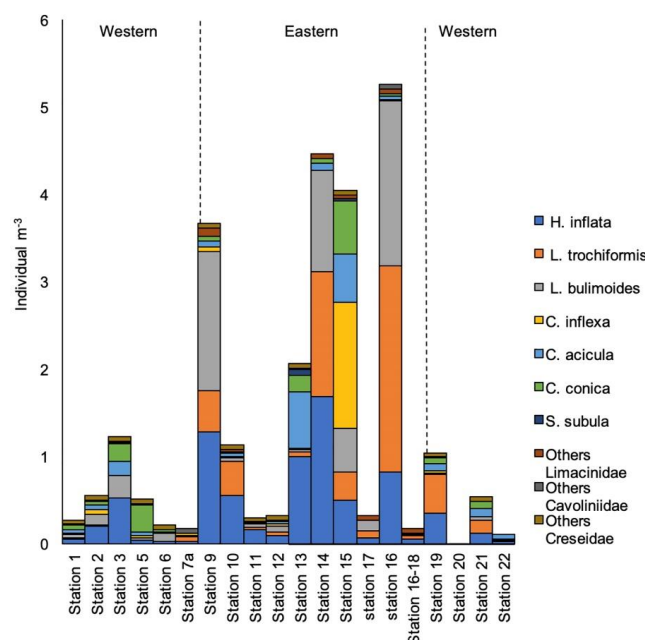


Figure 2. Absolute abundance of planktic pteropods from stations 1-22 on the MedSeA cruise. The category of 'Others' for each family includes specimens that were not a target species in this study or that were unidentifiable to the species level.

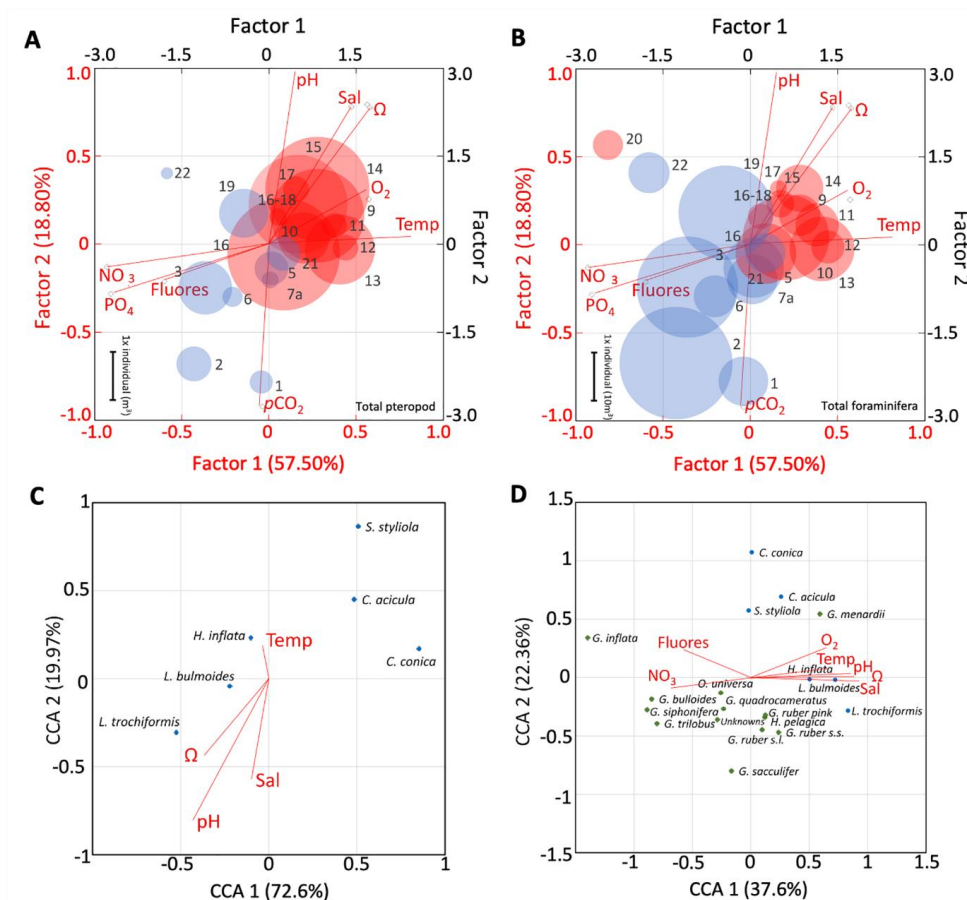


Figure 3. PCA graphs of environmental factors overlaid with absolute abundance values on station scores of **A** pteropods at all stations (ind. per m^{-3}) and **B** foraminifera at all stations (ind. per $10m^{-3}$). The red axes are associated with PCA coordinates and the black axes are associated with the station coordinates. CCA graphs indicating **C** pteropod community composition and **D** pteropod (blue dots) and foraminifera (green dots) community composition. Family Cavoliniidae (*Cavolinia inflata*) is not indicated on the graph as they were not related to any environmental factor or species groupings. Coordinates for *C. inflata* in **C** (1.55, 0.59) and **D** (0.63, 2.33).

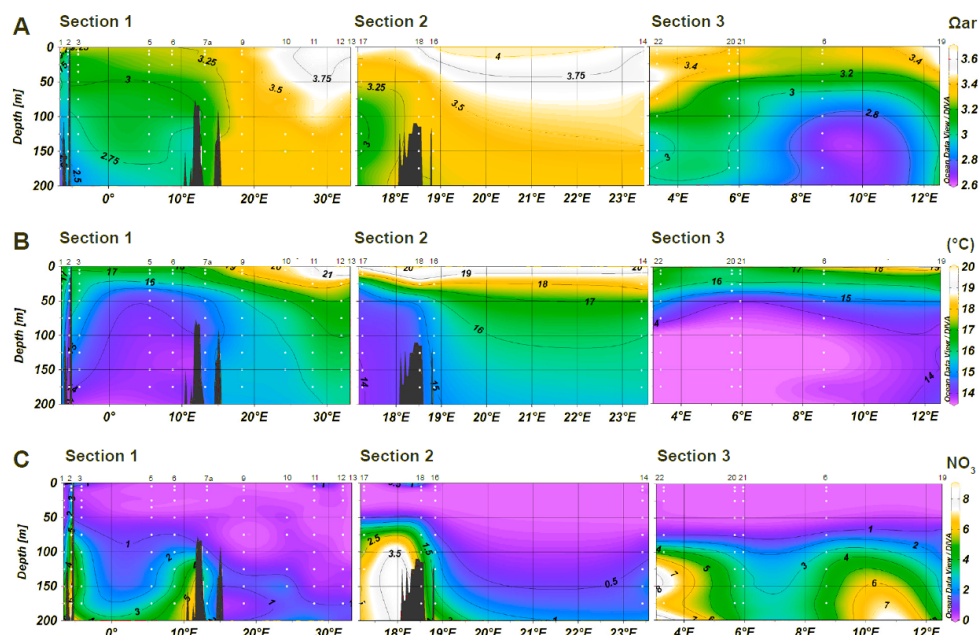
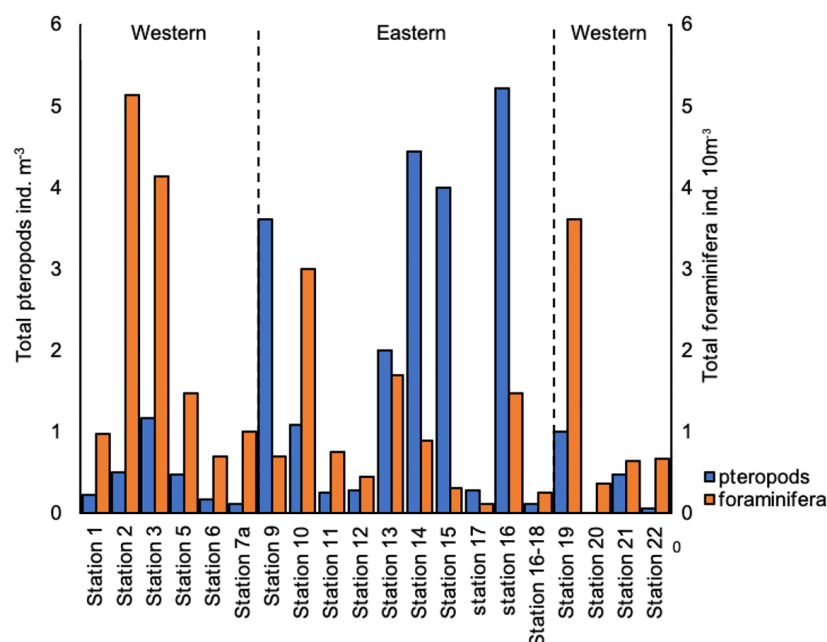


Figure 4. Depth profiles from 0-200 m depth highlighting the gradients of **A** Aragonite saturation (Ω_{ar}) **B** Nitrate (NO_3) and **C** Temperature ($^{\circ}\text{C}$) from the west to the east of the Mediterranean Basin (Section 1), within the Adriatic Sea (Section 2) and in the North-Eastern Mediterranean (Section 3) for factors affecting abundance and the community composition of pteropods. Figures made using Ocean Data View.



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671 **Figure 5.** A comparison of absolute total abundance for pteropods and foraminifera at each
 672 station. Note that the scale of foraminifera abundance is distinctly lower than pteropod
 673 abundance (ind. 10 m⁻³ and ind. m⁻³ respectively), however this graph serves to illustrate
 674 similarities and differences in patterns of abundance.