

1 **Interactions between biogeochemical and management factors explain soil**
2 **organic carbon in Pyrenean grasslands**

3 **Authors**

4 **Antonio Rodríguez^{1,2}, Rosa Maria Canals³, Josefina Plaixats⁴, Elena**

5 **Albanell⁴, Haifa Debouk^{1,2}, Jordi Garcia-Pausas⁵, Leticia San Emeterio^{6,7}, Àngela**

6 **Ribas^{8,9}, Juan José Jimenez¹⁰, M.-Teresa Sebastià^{1,2}**

7 **Afiliations**

8 1: Group GAMES, Department of Horticulture, Botany and Landscaping, School of
9 Agrifood and Forestry Science and Engineering, University of Lleida, Lleida, Spain

10 2: Laboratory of Functional Ecology and Global Change (ECOFUN), Forest Sciences
11 Centre of Catalonia (CTFC), Solsona, Spain

12 3: Grupo Ecología y Medio Ambiente and ISFood Institute, Universidad Pública de
13 Navarra, Campus Arrosadia, Pamplona, Spain.

14 4: Group of Ruminant Research (G2R), Department of Animal and Food Sciences,
15 Universitat Autònoma de Barcelona, 08193, Bellaterra, Spain

16 5: Forest Sciences Centre of Catalonia (CTFC), Solsona, Spain

17 6: Research Institute on Innovation & Sustainable Development in Food Chain
18 (ISFOOD), Universidad Pública de Navarra, 31006 Pamplona, Spain.

19 7: Departamento de Agronomía, Biotecnología y Alimentación, Universidad Pública de
20 Navarra, 31006 Pamplona, Spain.

21 8: Universitat Autònoma de Barcelona, 08193, Bellaterra, Spain.

22 9: Centre for Ecological Research and Forestry Applications (CREAF), 08193, Bellaterra,
23 Spain.

24 10: ARAID/IPE-CSIC, 22700, Huesca, Spain

25 **ORCID**

26 Antonio Rodríguez <http://orcid.org/0000-0002-0536-9902>

27 Elena Albanell <https://orcid.org/0000-0002-6158-7736>

28 Jordi Garcia-Pausas <https://orcid.org/0000-0003-2727-3167>

29 Leticia San Emeterio <http://orcid.org/0000-0002-8063-0402>

30 Àngela Ribas <https://orcid.org/0000-0002-5938-2408>

31 M.-Teresa Sebastià <http://orcid.org/0000-0002-9017-3575>

32

33 **Author contributions**

34 Antonio Rodríguez designed the statistical procedure, carried out the statistical analyses
35 and wrote the original draft.

36 Rosa M Canals was responsible for field monitoring, lab analyses and acquisition of
37 information for the database implementation in the Western Pyrenees (Navarra). She
38 also reviewed the draft.

39 Elena Albanell designed the NIRS study and reviewed the draft.

40 Haifa Debouk sampled and processed some of the data in the PASTUS Database and
41 reviewed the draft.

42 Jordi García-Pausas processed some of the data in the PASTUS Database and
43 reviewed the draft.

44 Josefina Plaixats carried out the chemical analyses of herbage samples for NIR
45 calibration and validation equations and reviewed the draft.

46 Leticia San Emeterio designed methodology and data collection, performed soil and
47 vegetation sampling. She also reviewed the draft.

48 Àngela Ribas sampled and processed some of the data in the PASTUS Database and
49 reviewed the draft.

50 Juan José Jiménez collaborated in the fieldwork and reviewed the draft.

51 M.-Teresa Sebastià contributed to the conception, design and development of the
52 PASTUS database. In addition, she ensured funding and coordinated the projects whose
53 data are included in PASTUS. Finally, she contributed to initial modelling, supervised the
54 development of the paper, read and reviewed the drafts.

55 **Abstract**

56 Grasslands are one of the major sinks of terrestrial soil organic carbon (SOC).
57 Understanding how environmental and management factors drive SOC is challenging
58 because they are scale-dependent, with large-scale drivers affecting SOC both directly
59 and through drivers working at small scales. Here we addressed how regional, landscape
60 and grazing management, soil properties and nutrients, and herbage quality factors
61 affect 20 cm depth SOC stocks in mountain grasslands in the Pyrenees. Taking
62 advantage of the high variety of environmental heterogeneity in the Pyrenees, we built a
63 dataset (n = 128) that comprises a wide range of environmental and management
64 conditions. This was used to understand the relationship between SOC stocks and their
65 drivers considering multiple environments. We found that temperature seasonality
66 (difference between mean summer temperature and mean annual temperature; TSIS)
67 was the most important geophysical driver of SOC in our study, depending on
68 topography and management. TSIS effects on SOC increased in exposed hillsides, slopy
69 areas, and relatively intensively grazed grasslands. Increased TSIS probably favours
70 plant biomass production, particularly at high altitudes, but landscape and grazing
71 management factors regulate the accumulation of this biomass into SOC. Concerning
72 biochemical SOC drivers, we found unexpected interactive effects between grazer type,
73 soil nutrients and herbage quality. Soil N was a crucial SOC driver as expected, but
74 modulated by livestock species and neutral detergent fibre contenting plant biomass;
75 herbage recalcitrance effects varied depending on grazer species. These results
76 highlight the gaps in the knowledge about SOC drivers in grasslands under different
77 environmental and management conditions. They may also serve to generate testable
78 hypotheses in later/future studies directed to climate change mitigation policies.

79 **Keywords**

80 SOC, natural grasslands, grazer type; grazing management, herbage quality; climate
81 change, soil nutrients; topography; temperature seasonality; TSIS

82

83 **Introduction**

84 Soil organic carbon (SOC) is crucial for the functioning of terrestrial ecosystems
85 (Lal, 2004a). SOC enhances soil and water quality and biomass productivity, and
86 has an important role in relation to climate change (Lal, 2004b). Although
87 grasslands have small aboveground biomass compared to other ecosystems,
88 their SOC stocks can be comparable to those in forest ecosystems (Berninger et
89 al., 2015). This is due to their high root biomass and residues, which are a
90 substantial carbon source and can contribute to water retention in soil. This
91 creates favourable conditions for the accumulation of organic matter (Von Haden
92 and Dornbush, 2014; Yang et al., 2018). These attributes, together with the high
93 extent of grassland global cover, make grasslands store around 34% of the
94 terrestrial carbon, mostly in their soils (White et al., 2000).

95 SOC accumulation results from a complex equilibrium between primary
96 production and organic matter decomposition which depends on multiple
97 environmental factors such as climate, soil texture and nutrients, or land
98 management (Jenny, 1941; Schlesinger, 1977). Understanding how these scale-
99 dependent environmental factors drive SOC is challenging because large scale
100 drivers affect also those working at fine spatial scales. This has been described
101 as a hierarchy of controls over SOC (Fig. 1; Manning et al., 2015).

102 Climate is known to be the main SOC driver at broad (global and regional) scales;
103 mean annual precipitation (MAP) and mean temperature (MAT) being the most
104 frequent climate indicators (Wiesmeier et al., 2019). However, climate annual
105 variations represented by seasonality variables are commonly neglected when

106 considering possible SOC drivers in broad-scale models, in spite of being
107 important drivers of plant primary production and enzymatic activity of soil
108 microorganisms (Fernández-Alonso et al., 2018; Garcia-Pausas et al., 2007;
109 Puissant et al., 2018). Other regional and landscape factors like bedrock or
110 topography are also considered to be at the top of the hierarchy because they
111 influence multiple geophysical and biochemical factors affecting SOC, including
112 soil texture and water flow paths (Gray et al., 2015; Hobbey et al., 2015). Next in
113 the hierarchy after regional and landscape factors, are several soil geophysical
114 properties, like pH and texture, which are controlled by climate, bedrock, and
115 which affect SOC through both plant primary production and microbial activity
116 and the capacity to stabilise the SOC (Deng et al., 2016; Xu et al., 2016a).

117 Soil macro and micronutrients are in the next level of the hierarchy, as their
118 abundance is determined by multiple factors, including climate, soil pH, water
119 content or clay content (Hook and Burke, 2000; de Vries et al., 2012). They play
120 an essential role in primary production and herbage quality, and act as resources
121 for microbes to mineralise SOC (Aerts and Chapin, 1999; Vitousek and Howarth,
122 1991). However, these variables are commonly omitted as possible drivers of
123 SOC in the broad-scale studies, especially in those studies focusing on predictive
124 rather than explicative models (Gray et al., 2015; Manning et al., 2015; Zhang et
125 al., 2018). This kind of variables is less frequently available and more difficult to
126 measure than the other indicators used for SOC modelling (Manning et al., 2015).
127 Furthermore, the use of soil nutrients as SOC drivers in linear models can be
128 challenging, as they are often strongly linked to SOC dynamics. This may mask
129 the effect of other drivers acting at larger spatial scales (Bing et al., 2016;
130 Cleveland and Liptzin, 2007; Tipping et al., 2016).

131 Vegetation represents another group of SOC drivers, affected by climate,
132 topography and soil properties and nutrients (Fernández-Martínez et al., 2014;
133 de Vries et al., 2012; Zhu et al., 2019). Plant biomass is the main input of organic
134 carbon into the soil (Shipley and Parent, 1991). However, a not so frequently
135 considered factor is plant litter quality, which can determine decomposition rates
136 and patterns, and hence soil carbon sequestration (Ottoy et al., 2017; Yan et al.,
137 2018, 2019).

138 In addition to these factors, livestock management effects on grassland SOC is
139 a noteworthy issue since they are poorly understood (Wiesmeier et al., 2019). It
140 is known that herbivores can affect SOC through different paths, such as
141 regulating the quantity and quality of organic matter returned to soil (Bardgett and
142 Wardle, 2003), or affecting soil respiration and nutrients by animal trampling or
143 soil microbiota alteration (Lu et al., 2017). Several studies confirmed the
144 interaction between grazing and other SOC drivers at diverse scales (Abdalla et
145 al., 2018; Eze et al., 2018; Lu et al., 2015, 2017; Zhou et al., 2017). Hence,
146 grazing management may be considered a SOC driver with effects at multiple
147 levels of the driver hierarchy (Fig. 1), both affecting other SOC drivers and
148 interacting with them. However, most of the studies investigating grazing effects
149 on SOC focus on grazing intensity, in spite of evidence pointing to a greater role
150 of grazer species in determining vegetation and SOC (Chang et al., 2018;
151 Sebastia et al., 2008).

152 In this study, our goal was to identify the main drivers of SOC stocks and their
153 interactions in Pyrenean mountain grasslands. For this purpose, we considered
154 a wide set of regional, landscape, soil geophysical and biochemical, and herbage
155 quality factors, together with grazing management factors. Mountain grasslands

156 comprise a wide range of all these conditions, which make carbon stocks highly
157 variable (Garcia-Pausas et al., 2007, 2017). For this reason, data analysed here
158 include a wide range of environmental conditions, comparable to studies on SOC
159 developed at continental or even worldwide scales (Table 1). Additionally, we
160 considered an exceptionally broad compilation of drivers (Table 1). To deal with
161 correlations and interactions between SOC drivers, we developed an exhaustive
162 modelling approach based on the controls over function hypothesis (de Vries et
163 al., 2012). To facilitate the formulation of our specific questions to answer in this
164 study, we classified SOC drivers into three main groups (Fig. 1): i) geophysical
165 factors, which include regional and landscape factors and are supposed to be the
166 first sources of variation, ii) biochemical factors, which include soil nutrients and
167 herbage factors and could be conditioned by geophysical factors, and iii) grazing
168 management factors, which could affect SOC through multiple interactions with
169 the rest of the variables at multiple scales. In particular, the specific questions of
170 this study are 1) What are the relative and interaction effects of the geophysical
171 and biochemical SOC controls? 2) How does grazing management regulate the
172 effects of other SOC drivers?

173 **Material & methods**

174 **2.1 Location and sampling design**

175 The set of data used in this study has been extracted from the PASTUS Database
176 (<http://ecofun.ctfc.cat/?p=3538>), which was compiled by the Laboratory of Functional
177 Ecology and Global Change (ECOFUN) of the Forest Sciences Centre of Catalonia
178 (CTFC) and the University of Lleida (UdL). We sourced a wealth of data of 128 grassland
179 patches distributed across the Pyrenees (Fig. S1), and including topographical, climate,
180 soil, herbage and management variables. The elaboration of the PASTUS Database

181 concerning this study is summarised in Fig. S2). The sampled area encompasses a wide
182 variety of temperate and cold-temperate climates, with different precipitation conditions,
183 depending on altitude and geographical location from Mediterranean to Continental and
184 Boreo-Alpine environments (de Lamo & Sebastià, 2006; Rodríguez et al., 2018; Table
185 1). Almost all of the plant species in the grasslands from the PASTUS database are
186 perennial (Sebastià, 2004), and plant diversity is highly heterogeneous as are the
187 environmental conditions (Rodríguez et al., 2018).

188 Sampling in the PASTUS database was designed according to a stratified random
189 scheme, where samples were selected at random within strata. This process was done
190 using the software ArcMap 10 (ESRI, Redlands, CA, USA). The basis for randomization
191 was the map of habitats of Catalonia 1:50000 (Carreras and Diego, 2006) for the Eastern
192 and Central sectors of the Pyrenees, the map of habitats of Madres-Coronat 1:10000
193 (Penin, 1997) for the North-Eastern sector and the land use map of Navarra 1:25000
194 (Gobierno de Navarra, 2003) for the Western sectors. Four variables were initially
195 considered for sampling stratification within each sector: altitude (< 1800 m; 1800-2300
196 m; > 2300 m), slope (0-20°; 20-30°; > 30°), macrotopography (mountain top/southern-
197 facing slope; valley bottom/northern-facing slope) and grazer type (sheep; cattle; mixed).
198 Accordingly, we determined a set of homogeneous grassland patches by crossing the
199 stratification variable layers. Grassland patches were then listed by type and arranged
200 within each list randomly to determine sampling priority. At least one to two replicates of
201 each patch type defined by the stratification variables were sampled.

202 In each sampled grassland patch, a 10 x 10 m² plot was systematically placed in the
203 middle of each homogeneous grassland patch, including a particular plant community.
204 We collected soil and vegetation samples, and assessed environmental variables within
205 each 100 m² plot (see Rodríguez et al., (2018) for additional details about sampling
206 design). Local variables were assessed inside the 100 m² plots. Aboveground biomass
207 was estimated from herbage cut at ground level in four 50 x 50 cm² quadrats placed in a

208 2 x 2 m² subplot inside the 100 m² plot. Herbage from two of the four quadrats were dried
209 and sent to the laboratory for duplicated chemico-bromatological analysis. In addition, in
210 each quadrat, a 20-cm depth soil core was extracted with a 5 x 5 cm probe after herbage
211 was removed. The soil sample in the probe was separated into two soil layers: 0-10 and
212 10-20 cm.

213 **2.2 Regional and landscape environmental drivers**

214 In order to investigate the relationship between SOC and potential environmental drivers,
215 30 independent environmental variables were initially considered (Table S1). These
216 variables were grouped into five sets: regional, landscape, livestock management, soil
217 nutrient stocks, and herbage variables.

218 Regional variables included climate variables and bedrock. Climate variables were
219 determined from Worldclim 2.0 (Fick and Hijmans, 2017). We selected Mean Annual
220 Temperature (MAT), Mean Summer Temperature (MST), Mean Annual Precipitation
221 (MAP) and Mean Summer Precipitation (MSP). The difference between mean annual
222 and mean summer temperature emerged as a relevant explanatory factor of soil organic
223 carbon stocks during previous modelling efforts by one of the co-authors (M-T. Sebastià).
224 Later attempts to improve models by substituting this variable with other temperature
225 indices from climatic databases (Fick and Hijmans, 2017) showed that, for the PASTUS
226 database, this variable provided higher explanatory power than other temperature
227 seasonality indices. Thus, we decided to keep it and here we name it Temperature
228 Seasonality Index of Sebastià (TSIS from now on).

229 Bedrock type was determined in the field and confirmed by the geographical maps
230 mentioned above. Bedrock was categorized into three categories: basic (marls and
231 calcareous rocks), acidic (mostly sandstones and slates) and heterogeneous.

232 Landscape variables included topography and soil type variables. Topography variables
233 included Slope, Aspect, Macrotopography and Microtopography. Slope and Aspect were

234 determined in the field by clinometer and compass respectively. Macrotopography and
235 microtopography were determined visually in the field. Preliminary modelling efforts
236 suggested the reduction of the four macrotopographical positions initially identified in the
237 field into two: Mountain top and south-facing slopes were classified as exposed positions
238 and valley bottoms and north-facing slopes as protected macrotopographical positions.
239 Microtopography included three positions: convexities, concavities and smooth areas.
240 Soil type variables are described in the following.

241 **2.3 Soil physicochemical analysis**

242 To obtain bulk density, we air-dried and weighed the soil samples: we then sieved each
243 sample to 2 mm to separate stones and gravels from the fine earth fraction. The fine
244 fraction was sent to the laboratory for further physicochemical analysis. Standard
245 physicochemical soil analyses were performed in the upper 0-10 cm soil layer of all
246 grasslands. Some analyses were also performed on samples from the 10-20 cm soil
247 layer, including soil organic carbon and total nitrogen. For those variables, we combined
248 0-10 and 10-20 cm values to obtain the whole top 20 cm soil layer.

249 All soil physicochemical analyses were performed on the fine earth, according to
250 standard soil analysis methods. Textural classes were determined by the Bouyoucos
251 method (Bouyoucos, 1936). Soil pH (measured in water), total organic carbon (TOC)
252 total nitrogen (TN), Calcium content (Ca), Extractable phosphorus (P), magnesium (Mg)
253 and potassium (K) were measured on air-dried samples (Schöning et al., 2013; Solly et
254 al., 2014). Soil carbonates were determined using the Bernard calcimeter. Total carbon
255 and nitrogen (N) contents of the fine earth were determined by elemental auto-analyser.
256 The organic C fraction was determined by subtracting inorganic C in the carbonates from
257 the total C. Available phosphorus (P) was extracted by the Olsen method (Olsen, 1954)
258 Ca, Mg and K were extracted by ammonium acetate (Simard, 1993) and measured by
259 flame Atomic Absorption Spectroscopy (AAS) (David, 1960)). Soil organic carbon (SOC)
260 stocks in the upper 20 cm soil layer were then estimated taking into account the organic

261 C concentration in the sample and its bulk density, and subtracting the coarse particle (>
262 2 mm) content, following García-Pausas et al. (2007). Despite recent studies suggesting
263 that fixed mass SOC stocks estimates are preferable to fixed depth methods because
264 they would be more robust to temporal and land use changes in bulk density (Ellert &
265 Bettany 1995), we chose a fixed depth method for measuring SOC stocks. This decision
266 was based on the fact that our work samples came from natural mountain grasslands,
267 where grazing intensity is always low to moderate, and moreover, herbivore presence is
268 seasonal. Therefore, we do not expect important changes in bulk density due to land
269 use. Additionally fixed mass approaches presented the disadvantages of implying more
270 technical difficulties than fixed depth measures, even in the most modern procedures
271 (Haden et al. 2020), and could not deal well with differences in stoniness.

272 **2.4 Herbage chemical and bromatological analysis, and NIRS analysis**

273 All four herbage samples per plot were oven-dried at 60°C to constant weight to
274 determine aboveground biomass and converted into g m⁻². Two out of the four samples
275 were sent to the laboratory for herbage quality analysis. Dried samples were ground to
276 pass a 1 mm stainless steel screen (Cyclotec 1093 Sample mill, Tecator, Sweden) and
277 stored at 4°C until it was needed for use.

278 To develop NIRS prediction models, a random subset of 130 samples was used and
279 analysed in duplicate according to the reference methods mentioned further. Procedures
280 described by AOAC were used to determine dry matter (DM) and ash content or mineral
281 matter (MM). Crude protein (CP) was determined by the Kjeldhal procedure (N x 6.25)
282 using a Kjeltec Auto 1030 Analyser (Tecator, Sweden). Samples were analysed
283 sequentially for neutral detergent fibre (NDF), acid detergent fibre (ADF) and acid
284 detergent lignin (ADL) in accordance with the method described Van Soest et al. (1991)
285 using the Ankom 200 Fibre Analyser incubator (Ankom, USA). The fibre analysis were
286 determined on an ash-free basis and without alpha amylase. We calculated two
287 additional herbage quality indexes often used in the bibliography: NDF/CP and ADL/NH

288 (Stockmann et al., 2013). For each subsample the C and N content (CH and NH)were
289 determined by the Dumas dry combustion method, using an Elemental Analyzer EA1108
290 (Carlo Erba, Milan, Italy).

291 Afterwards, a total of two hundred herbage samples were scanned as described below
292 to collect their NIRS spectra. We estimated chemical and bromatological variables
293 according to the equations derived from the previous calibrations on the initial 130
294 random samples.

295 NIRS data were recorded from 1,100 to 2,500 nm using a FOSS NIRSystem 5000
296 scanning monochromator (Hillerød, Denmark). Separate calibration equations were
297 generated for grassland samples. Reflectance (R) data were collected in duplicate every
298 2 nm. A WinISI III (v. 1.6) software (FOSS, Denmark) was employed for spectra data
299 analysis and development of chemometric models. Prior to calibration, log 1/R spectra
300 were corrected for the effects of scatter using the standard normal variate (SNV), detrend
301 (DT) and multiple scatter correction (MSC) and transformed into first or second derivative
302 using different gap size (nm) and smoothing interval. For each sample, the mean of the
303 spectra from the two lectures were used. Modified partial least square (MPLS) was the
304 regression method used for calibration development and cross validation was
305 undertaken using the standard methodology in the NIRS software program. The
306 performance of the model was determined by the following statistical tools: standard
307 error of calibration (SEC), standard error of cross validation (SECV); coefficient of
308 determination for calibration (R^2) and cross validation (r_{cv}^2); the ratio of performance to
309 deviation (RPD) described as the ratio of standard deviation for the validation samples
310 to the standard error of cross validation ($RPD=SD/SECV$) should ideally be at least three;
311 and the range error ratio ($RER=Range/SECV$) described as the ratio of the range in the
312 reference data to the SECV should be at least 10 (Williams and Sobering, 1996; Williams
313 et al., 2014).

314 **2.5 Livestock management variables**

315 The management variables (grazer type) initially used for sampling stratification were
316 determined from records available in the municipalities of the study area. Once the
317 specific grassland patches to be sampled were determined, we carried out a detailed
318 analysis of the management where the patches were located. To this effect, we carried
319 out detailed surveys among farmers, shepherds and land managers. Sometimes the
320 information collected was modified according to visual records in the field (e.g., cattle
321 and/or cattle dung found in supposedly ungrazed areas). Information from municipalities
322 was usually the most imprecise.

323 We considered two management variables: Grazing intensity and Grazer type. Grazing
324 intensity was determined estimating livestock stocking rates measured as livestock units
325 ha^{-1} (LU ha^{-1}), and treated as a semi-quantitative variable with three categories (1): low
326 (1; lower than 0.2 LU ha^{-1}), medium (2; between $0.2\text{-}0.4 \text{ LU ha}^{-1}$) and high (3; above 0.4
327 LU ha^{-1}). Grazer type was categorised into three main types: sheep, cattle and mixed.
328 Mixed grazing included associations comprising small and big livestock species, mainly
329 sheep and cattle, and more rarely horses. Sheep flocks always included some goats.

330 **2.7 Statistical analyses**

331 We applied two different modelling procedures: Boosted Regression Trees (BRT) and
332 General Linear Models (GLM). BRT is an automatic technique that combines insights
333 from traditional statistical modelling and machine learning traditions (Elith et al., 2008).
334 GLM allowed us to design a hypothesis-based modelling procedure, ensuring that only
335 effects with biological meaning were included; whereas BRT provided information about
336 the data that could be neglected, if only a GLM approach was followed.

337 All the statistical analyses were performed with the software R ver. 3.4.3 (R Core Team,
338 2017), at 95% significance level when appropriate.

339 *Boosted regression trees global model*

340 Including all SOC potential drivers, we fitted a model with BRT to identify the most
341 important variables affecting SOC. BRT uses two algorithms: regression trees and
342 boosting. Regression trees are from the decision tree group of models, and boosting
343 builds and combines a collection of models (Elith et al., 2008). We chose this method
344 because BRT can handle multiple variables better than other techniques as GLM, and
345 can detect automatically curvilinear relationships and interactions, ignoring non-
346 informative ones. We used the gbm and dismo packages (Greenwell et al., 2019;
347 Hijmans et al., 2017), which provide several functions to fit these models.

348 Firstly, we fitted a model with all the drivers (Table S1), configured with 15 folds, a
349 Gaussian distribution of the error, a tree complexity of 5, a learning.rate of 0.005, a
350 bag.fraction of 0.666, and 5 minimum observations by node. Secondly, we reduced the
351 number of drivers by the method described in Elith et al., (2008). We estimated the
352 change in the model's predictive deviance dropping one by one each driver, and re-fitted
353 the model with the set of variables which actually improved model performance (Fig. S3).
354 We checked the relative importance of the drivers and the shape and size of the effects
355 by partial effect plots.

356 *General linear models*

357 We designed and executed a modelling procedure based on general linear models
358 (Legendre and Legendre, 1998) and a hierarchy of controls over function (Díaz et al.,
359 2007; de Vries et al., 2012). We log-transformed SOC using natural logarithm to prevent
360 a breach of the normality assumption by the residuals of the models (Fig. S4). We built
361 two models (Fig. S5), one model based only on geophysical drivers and grazing
362 management (geophysical model), and another model including, in addition to the former
363 drivers, the biochemical drivers: soil nutrients and herbage quality (combined model).
364 With this approach we aimed to avoid ignoring significant effects of the geophysical
365 variables, the original source of variation of SOC stocks according to the hierarchy of

366 controls over function hypothesis (Manning et al., 2015), by masking them with the
367 inclusion of biochemical drivers. We considered that the geophysical factors that
368 potentially affect SOC were regional and landscape (topography and soil type drivers),
369 as they have been widely used in previous studies to model and predict SOC from
370 landscape to continental scales (Manning et al., 2015; Wiesmeier et al., 2019). In
371 addition to soil nutrients and herbage variables, we included again the livestock
372 management variables in the combined model and looked for interactions involving these
373 variables and biochemical drivers of SOC.

374 For model building (Fig. S5A), we added driver groups following a sequential order. For
375 fitting the geophysical model, we started adding regional, landscape and grazing
376 management drivers, and subsequently included soil properties. Afterwards, we
377 sequentially included soil nutrients and herbage drivers to obtain the combined model.
378 We added Management variables from the beginning of the modelling process and re-
379 included the discarded ones in each step to guarantee the detection of interactions
380 between Management variables and the rest of the drivers. Each time we added a set of
381 drivers, we first considered their main effects and some quadratic terms which were
382 found by preliminary analyses with the `scatterplot.matrix` function in the R package `car`
383 (Fox et al., 2018); afterwards we included possible level 2 interactions between all the
384 selected drivers.

385 At every step we selected several candidate terms by a semi-automatic procedure (Fig.
386 S5C) using a genetic algorithm included in the R package `glmulti` (Calcagno, 2015). We
387 used SOC as response variable at the first step, and the residuals of the previous model
388 in the remaining steps (Fig. S5B). This semi-automatic process began by obtaining a
389 best subset of models using the corrected Akaike information criterion (AICc),
390 appropriate when n/k is less than 40, n being the sample size and k the number of
391 parameters in the most complex model (Symonds and Moussalli, 2011). We selected the
392 best model and its equivalents obtained by the genetic algorithm, which were those

393 within 2 Akaike information criterion-corrected (ΔAICc) values of the best model, as a
394 $\Delta\text{AICc} < 2$ indicates that the candidate model is almost as good as the best model
395 (Burnham and Anderson, 2002).

396 For this set of models, we built averaged models using the MUMIn package (Barton,
397 2015). We calculated partial standardized coefficients, obtained by multiplying the
398 unstandardized coefficient in the model by the partial standard deviation of the variable,
399 which is a function of the standard deviation of the variable, the sample size, the number
400 of variables in the model and the variance inflation factor of the variable (Barton, 2015).
401 We selected all the variables with significant effects (alone or in interaction with each
402 other) in the conditional average model, which was preferred over the full average model
403 because we wanted to avoid excessive shrinkage effects at this moment of the modelling
404 procedure (Grueber et al., 2011).

405 Then, we added these terms to the consolidated model, and made a selection through a
406 backward forward procedure. We used several methods to compare and determine the
407 final model, including the AICc, the adjusted determination coefficient R^2 (R_{adj}^2) and
408 model comparison techniques with the “anova()” function in R, using Chi-square tests to
409 test whether the reduction in the residual sum of squares was statistically significant.
410 Once we had the final model we assessed the significance of each term by removing it
411 and performing an F test. For estimating the significance of the main effects we also
412 removed the interaction terms in which they were involved, to avoid transferring the
413 effects of the main terms to the interaction terms (de Vries et al., 2012). We estimated
414 the variance explained by the models through the adjusted determination coefficient R^2
415 (R_{adj}^2).

416 Finally, we estimated the importance of the terms of each model by the lmg method in
417 the relaimpo package (Grömping, 2006), and drew partial effect plots making predictions
418 with the R package emmeans (Lenth et al., 2019).

419 **Results**

420 SOC stocks of the upper 20 cm layer ranged between 2.6 and 23 kg m⁻², with a median
421 and a mean value of 9.1 and 9.6 kg m⁻² respectively. Standard deviation of the mean
422 was 3.15 (n= 125). Minimum, maximum, median and mean values of the continuous
423 drivers are shown in Table S2.

424 **3.1 Relative importance of SOC stocks drivers**

425 The final BRT global model achieved a good goodness of fit, with a cross-validated
426 correlation value of 48% and an explained deviance of 88 %. The most important
427 variables explaining SOC stocks (Fig. 2) were soil N (18%), soil C/N (14%) and clay
428 (13%) although other variables such as aboveground biomass (7%), ADL (6%) or silt
429 (6%) were also relevant for explaining SOC storage. Three important variables in the
430 BRT model, aboveground biomass, silt and soil K, were not selected in the linear models
431 (Tables 2 & 3). Although accounting for a lower importance value than the previous
432 variables (5%), TSIS was the most relevant among the climate drivers considered. TSIS
433 was also noticeably important in both linear models (Fig. S6), especially in the
434 geophysical model, not only as main effect, but in interaction with other variables (lmg:
435 4-10%). According to the combined linear model, soil nutrient and herbage variables
436 were other important SOC stocks drivers (Fig. S7), but many of these effects occurred
437 in interaction with grazer type.

438 **3.2 Geophysical, biochemical and grazing management effects on SOC stocks**

439 The geophysical model (Table 2) explained 34% of the total variance (R^2_{Adj}). Overall,
440 SOC stocks increased with TSIS under certain conditions: exposed hillsides, high slopes
441 and low stocking rates (Fig. 3A, 3B & 3D). On the other hand, clay content had a positive
442 relationship with SOC under low MAP values (Fig. 3C), which turned into negative at
443 high MAP values.

444 Adding nutrient and herbage variables to the previous geophysical model to build the
445 combined model (Table 3) increased the total variance (R^2_{Adj}) up to 61%.
446 Macrotopography and clay effects described by the geophysical model were removed
447 by the new model terms included. SOC increased with C/N (Fig 4A). Soil nitrogen
448 modulated the effects of livestock type and NDF on SOC. Cattle-grazed grasslands
449 stored more SOC than mixed and sheep grazed grasslands under low soil N conditions,
450 whereas the reverse occurred at high soil N levels (Fig. 3B). NDF had negative effects
451 on SOC stocks at high soil N values but had no effect under low soil N levels Fig. 4C).
452 Finally, herbage ADL/NH had positive effects on SOC under mixed and sheep grazing
453 regimes, but there was no effect under cattle management (Fig. 4D).

454 **Discussion**

455 **3.1 Considerations about the modelling procedure**

456 Unsurprisingly, the SOC drivers selected and their main effects in both of the modelling
457 approaches (BRT and GLMs) were highly congruent (Figs. 2-4; S8). Consequently, we
458 preferred to focus on the results from the linear models because this approximation
459 allowed us to build models under a hierarchy of controls over function hypothesis
460 (Manning et al., 2015). Hence, although it is not possible to unequivocally establish the
461 causal links between SOC drivers (Grace, 2006; Grace and Bollen, 2005), with our GLMs
462 procedure we guarantee that the effects of the biochemical variables added in the
463 combined Model on SOC stocks have not been exclusively induced by geophysical
464 drivers (de Vries et al., 2012). If this was the case, soil nutrient and herbage quality
465 drivers wouldn't have entered the combined model as significant terms. This happened
466 with aboveground biomass, which is assumed to be a very important SOC driver, and
467 indeed aboveground biomass was relevant in the BRT model, but in the GLM was
468 substituted by other, more meaningful, variables. In addition, our GLM modelling
469 approach enabled us the selection of biologically meaningful interactions (Manning et

470 al., 2015; de Vries et al., 2012), which cannot be done with a fully automatic approach
471 like BRT. This GLM sequenced modelling procedure, looking for the primary sources of
472 variation, together with the stratified sampling design, is useful as it led us to select a set
473 of lowly correlated drivers for our linear models (Table S5). Furthermore, BRT model
474 provided some valuable information, identifying some relevant SOC drivers which were
475 discarded during the GML modelling, like aboveground biomass, or soil silt and K (Fig.
476 2 and S8). The effects of those drivers were probably masked by the effects of other
477 variables in our linear models (Yang et al., 2009), indicating that these factors were
478 presumably pathways through which other variables drove SOC (de Vries et al., 2012).
479 These variables, identified by BRT and discarded by GLM, should be considered as
480 potential SOC drivers in further studies, particularly when more detailed and difficult to
481 obtain biochemical variables, present in our database, are not available.

482 **3.2 Geophysical, biochemical and grazing management factors driving SOC** 483 **stocks**

484 Considering the difficulties of modelling SOC in a widely heterogeneous mountain
485 environment (Garcia-Pausas et al., 2017), the geophysical Model provided important
486 information about broad-scale and topographic SOC drivers in the Pyrenees. This
487 information could be useful not only for a better understanding of SOC patterns in
488 mountain grasslands, but also for future modelling studies aiming to predict SOC, since
489 geophysical variables are easier and less expensive to acquire and measure compared
490 to biochemical variables (Manning et al., 2015).

491 Most studies on soil carbon usually pinpoint mean temperature and precipitation as the
492 most important climate drivers of SOC (Hobley et al., 2015; Manning et al., 2015;
493 Wiesmeier et al., 2019). Climate regulates large-scale patterns of aboveground net
494 primary production (Chapin et al., 1987). In our study, temperature seasonality (TSIS)
495 was a key driver of SOC, modulated by macrotopography, slope and grazing intensity
496 (Table 2; Fig. 3). The highest variation of TSIS in our database, that is, the broadest

497 temperature seasonality, occurred in cold environments, as compared to mild climates
498 (Fig. S9). In mountain grasslands, cold climates imply a short phenological period of
499 development for plants (Gómez, 2008). Hence, the positive effect of TSIS on SOC could
500 be associated with a higher biomass accumulation in cold locations with more favourable
501 temperatures during summer, this fact reducing geophysical stress for plants and
502 broadening their growth period (Garcia-Pausas et al., 2007; Kikvidze et al., 2005).
503 increase in soil organic matter inputs during summer would overcome an eventual
504 increase of soil organic matter decomposition caused by high temperatures (Sanderman
505 et al., 2003).

506 The interactive effects of TSIS on SOC stocks with macrotopography and slope illustrate
507 the capacity of landscape factors to modulate macroclimate effects on soil (Hook and
508 Burke, 2000). Induced microclimate changes are often the explanation for the effects of
509 topography in SOC (Lozano-García et al., 2016). In our case, SOC stocks increased with
510 temperature seasonality, particularly in exposed locations, including south-facing
511 hillsides and hillside tops (Fig. 3A; Table 2). In protected locations, including shady
512 hillsides and valley bottoms, the hypothesized positive effect of increased TSIS values
513 on plant productivity could be mitigated due to reduced solar radiation, long snow-
514 covered periods and freezing episodes (Garcia-Pausas et al., 2007; López-Moreno et
515 al., 2013). Additionally, differences in SOC between exposed and protected sites may
516 also occur through other mechanisms, for instance the alteration of soil physico-chemical
517 properties (Zhang et al., 2018), or differences in vegetation (Sebastià, 2004). Since we
518 used a hierarchy of controls approach (Manning et al., 2015), these indirect
519 topographical effects on SOC stocks could be behind the exclusion in the linear models
520 of some drivers selected in the BRT model, like silt or pH (Figs. 2 & 3). In addition, SOC
521 stocks decreased with increase of slope, which may be attributed to reduced carbon
522 inputs and increased carbon losses induced by steeper slopes (Yuan et al., 2019 and

523 references therein). However, we found that increased temperature seasonality (TSIS)
524 values partly compensated negative slope effects on SOC.

525 The effect of temperature seasonality on SOC stocks was also modified by grazing
526 management. At low TSIS values, SOC stocks increased under moderate to high grazing
527 pressure; this effect disappeared as TSIS values increased (Fig. 3D). Recent meta-
528 analyses concluded that intensive grazing commonly has decreasing effects on SOC
529 (Abdalla et al., 2018; Eze et al., 2018; Mcsherry and Ritchie, 2013). However, these
530 effects were strongly context-specific, depending on other factors including climate and
531 soil type vegetation (Abdalla et al., 2018; Eze et al., 2018; Mcsherry and Ritchie, 2013).
532 Moreover, moderate grazing intensities can increase SOC inputs by dung deposition,
533 and aboveground and root biomass production (Franzluebbers et al., 2000; Zeng et al.,
534 2015). In our study, grazing intensity was relatively moderate (see methods), therefore
535 in our study increasing stocking rates may increase soil carbon inputs in moderate
536 seasonality locations by enhancing aboveground and belowground productivity.

537 Soil texture also showed interactive effects on SOC stocks with climatic variables. In
538 particular, clay effects on SOC stocks became negative as MAP values increased, (Fig.
539 3C; Table 2). Both MAP and clay content are widely assumed to be positively correlated
540 to SOC (Wiesmeier et al., 2019) but high soil water content caused by high MAP may
541 inhibit decomposition if a shortage of oxygen supply occurs (Xu et al., 2016b).
542 Furthermore, fine texture soils could be waterlogged frequently, leading to inhibition of
543 root growth and soil C allocation belowground (Mcsherry and Ritchie, 2013).

544 **3.2 Geophysical, biochemical and grazing management factors driving SOC** 545 **stocks**

546 Considering the difficulties of modelling SOC in a widely heterogeneous mountain
547 environment (Garcia-Pausas et al., 2017), the geophysical model provided important
548 information about SOC drivers in the Pyrenees. This information could be useful not only

549 for a better understanding of SOC patterns in mountain grasslands, but also for future
550 modelling studies aiming to predict SOC, since geophysical variables are easier and
551 less expensive to acquire and measure compared to biochemical ones (Manning et al.,
552 2015).

553 TSIS was a key driver of SOC with a varying effect depending on macrotopography,
554 slope and grazing intensity (Table 2; Fig. 3). While most of the previous studies
555 addressing soil carbon not included any temperature seasonality variable as potential
556 SOC predictor, usually focusing in mean temperature and precipitation as the most
557 important climate drivers of SOC (Hobley et al., 2015; Manning et al., 2015; Wiesmeier
558 et al., 2019), our models suggest that TSIS and other temperature seasonality indexes
559 should be included in further studies, to provide more evidence of the extent of the effects
560 of temperature seasonality on SOC stocks.

561 This result contrasts with. Climate regulates large-scale patterns of aboveground net
562 primary production (Chapin et al., 1987). In the case of mountain grasslands, cold
563 climates imply a short phenological period of development for plants (Gómez, 2008).
564 Cold sites characterised by low mean temperatures presented a wider spectrum of TSIS
565 values than warm sites, presenting both the lowest and the highest TSIS values (Fig.
566 S9). Hence, the positive effect of TSIS on SOC could be associated with a higher
567 biomass accumulation in cold locations with more favourable temperatures during
568 summer, this fact reducing geophysical stress for plants and broadening their growth
569 period (Garcia-Pausas et al., 2007; Kikvidze et al., 2005). This rise in soil organic matter
570 inputs during summer would overcome an eventual increase of soil organic matter
571 decomposition rates due to high temperatures (Sanderman et al., 2003), which could
572 even be diminished if microbial biomass decreases as a result of soil moisture reduction
573 (Puissant et al., 2018).

574 The interactions of TSIS with macrotopography and slope illustrate the capacity of
575 landscape factors to modulate macroclimate effects on soil (Hook and Burke, 2000).

576 Induced microclimate changes are often the explanation for the effects of topography in
577 SOC (Lozano-García et al., 2016). In our case, SOC stocks increased with temperature
578 seasonality, particularly at mountain-exposed areas (Fig. 3A; Table 2). In protected sites,
579 located in shady slopes and valley bottoms, the hypothesized positive effect of high TSIS
580 values on plant productivity could be mitigated due to lower solar radiation, longer snow-
581 covered periods and freezing episodes (Garcia-Pausas et al., 2007; López-Moreno et
582 al., 2013). Conversely, negative effects of low TSIS values on plant productivity could be
583 compensated thanks to the more humid conditions in protected sites compared to the
584 exposed sites (Garcia-Pausas et al., 2007). Additionally, it is important to take into
585 account that differences in SOC between exposed and protected sites may also occur
586 through other mechanisms, for instance the alteration of soil physico-chemical properties
587 like pH, soil texture or stoniness (Zhang et al., 2018), or differences in vegetation
588 (Sebastià, 2004). Since we used a hierarchy of controls approach (Manning et al., 2015),
589 these topography indirect effects on SOC stocks could be behind the exclusion in the
590 linear models of some drivers selected in the BRT model, like silt or pH (Figs. 2 & 3).

591 In addition, high TSIS values compensated SOC stocks decrease with a greater slope,
592 which may be attributed to reduced carbon inputs and increased carbon losses induced
593 by steeper slopes (Yuan et al., 2019 and references therein). Increases in grazing
594 pressure elevated SOC stocks under low TSIS values (Fig. 3D). This was a surprising
595 result according to recent meta-analyses, which concluded that grazing has commonly
596 decreasing effects on SOC (Abdalla et al., 2018; Eze et al., 2018; Mcsherry and Ritchie,
597 2013). However these effects were strongly context-specific, depending on other factors
598 like climate and soil type vegetation (Abdalla et al., 2018; Eze et al., 2018; Mcsherry and
599 Ritchie, 2013). Moreover, light and medium grazing intensities can increase SOC inputs
600 by dung deposition and promoting aboveground and root biomass production
601 (Franzluebbers et al., 2000; Zeng et al., 2015). Considering that in our natural grasslands
602 all grazing intensities are relatively low (see methods), our medium and high stock rates

603 may increase soil carbon inputs in low seasonality locations by enhancing aboveground
604 and belowground productivity.

605 Interestingly, clay content and precipitation presented interacting effects on SOC (Fig.
606 3C; Table 2). Both MAP and clay content are widely assumed to be positively correlated
607 to SOC (Wiesmeier et al., 2019). High MAP would increase SOC inputs by promoting
608 plant productivity (Author et al., 2000; Hobbey et al., 2015). Clay positive effects are often
609 attributed to a larger contact surface of soil particles (Kennedy et al., 2002), the
610 absorption of negatively charged organic matter, high soil water retention and the
611 exclusion of decomposer organisms due to their low pore size (Krull et al., 2001). In our
612 study, high soil water contents caused by high MAP may inhibit decomposition if a
613 shortage of oxygen supply occurs (Xu et al., 2016b). However, as MAP values increased,
614 clay effect on SOC became negative. To explain low SOC values at high MAP and high
615 clay content, McSherry and Rithchie (2013) hypothesized that finer texture soils could
616 be waterlogged more frequently, leading to inhibition of root growth and soil C allocation
617 belowground.

618 The addition of soil nutrient and herbage variables to our geophysical model implied
619 substitution of terms, including clay content and macrotopography, by newly added
620 variables (Tables 2 & 3). This highlights the importance of indirect effects of these
621 variables on SOC through other small scale drivers (Leifeld et al., 2015; Xu et al., 2016b;
622 Zhu et al., 2019). The combined model was complex and included unfrequently tested
623 effects involving interactions between grazer type, soil nutrients and herbage quality
624 variables (Table 3, Fig 4). Those results must be interpreted cautiously, because they
625 are based on observational data, but can contribute to generate testable hypotheses for
626 later studies about some complex and untested relationships between SOC and its
627 drivers. Interaction experiments concerning soil properties are expensive and rare in the
628 literature (Rillig et al., 2019).

629 For this reason, SOC increased with the C/N ratio (Fig 4A), which may be explained by
630 the difficulty of soil organic matter decomposition by soil microbes, decreasing
631 decomposition rates of SOC with increasing soil C/N (Wanyama et al., 2019; Xu et al.,
632 2016b). A positive relationship between SOC and soil N was also expected, since most
633 of the soil N is in combined form with organic matter (Cambardella and Elliott, 1994).
634 However, in this case, due to the wide range of conditions and the randomized sampling
635 design of the PASTUS database, the raw correlation between soil N and SOC was
636 somehow discrete ($r = 0.297$; $p\text{-value} = 0.001$; $R^2 = 0.088$), in comparison to other studies
637 (i.e. Yan et al., 2020). However, the novelty revealed by our model is that soil N could
638 modulate the effects of certain SOC drivers, including livestock type and herbage NDF.

639 Cattle-grazed grasslands stored more SOC than mixed- and sheep-grazed grasslands,
640 but only under low soil N conditions (Fig. 4B). Chang et al. (2018) found that in a N poor
641 semi-arid grassland, sheep decreased SOC content in comparison to cattle due to
642 vegetation changes caused by their feeding preference for highly palatable forbs
643 (Sebastia et al., 2008), thus promoting less palatable grasses which supported less root
644 biomass. In overall, under low soil N conditions, palatable plants are expected to
645 contribute to SOC inputs through the stimulation of C allocation in forb roots (Ågren and
646 Franklin, 2003; Warembourg et al., 2003) and the increase in the overall plant
647 productivity due to legume atmospheric N fixation (Van Der Heijden et al., 2008).

648 However, these processes could decline under high soil N contents. For instance,
649 legume atmospheric N fixation could be reduced since it requires additional energy in
650 comparison to nitrogen acquisition from the soil (Ibañez et al., 2020; Minchin and Witty,
651 2005). Additionally, sheep selective feeding habits could shift plant leaf traits in the
652 community towards nutrient-conservative leaf traits, which commonly induce fungal-
653 based soil food webs, with slow nutrient-cycling and high SOC storage due to low
654 decomposition rates (Orwin et al., 2010).

655 Additionally, grasslands with mixed grazed regimes stored even more SOC than sheep-
656 grazed grasslands under high soil N conditions (Fig. 4B, Table 3). This result emphasises
657 that mixed livestock assemblages deserve particular attention, because mixed grazing
658 can affect plant composition distinctly from single grazing species regimes, and alter
659 traveling and trampling behaviour of grazing animals (Aldezabal et al., 2019; Chang et
660 al., 2018; Liu et al., 2015).

661 NDF was negatively related to SOC at high soil N values (Fig 4C). NDF proportion
662 represents the amount of structural compounds on litter, and hence is inversely related
663 to non-structural compounds content (Goering and Van Soest, 1970). The latter are the
664 main source of organic matter formation at the early stages of decomposition, and they
665 are incorporated into microbial biomass in a highly efficient way (Cotrufo et al., 2013).
666 However, if microbial necromass was recycled by microbes before its incorporation to
667 mineral-associated organic matter (Córdova et al., 2018), it could be respired and
668 mineralized instead of stored. Thus, our model suggests that incorporation of labile and
669 fast metabolized non-organic compounds to soil organic matter could be a pathway of
670 SOC allocation at high soil N in Pyrenean grasslands.

671 On the other hand, the ADL/NH ratio was positively related to SOC in sheep and mixed
672 grazed grasslands (Fig. 4D). The ADL/NH ratio is a commonly used indicator for the
673 resistance of litter to degradation, particularly at later stages of decomposition (Taylor et
674 al., 1989). Hence, the increase of SOC stocks with ADL/NH should be related to the
675 physical pathway of soil organic matter incorporation, forming coarse particulate organic
676 matter (Cotrufo et al., 2015). Moreover, our model suggests that this pathway would be
677 inhibited under cattle grazing, presumably because of their higher digestive efficiency,
678 and thus less recalcitrant faeces (Wang et al., 2018); and their less selective diet
679 compared to sheep, as the latter would avoid plants with high lignin content, promoting
680 recalcitrant litter (Rosenthal et al., 2012; Sebastià et al., 2008).

681 Our results concerning interactions between grazer type and herbage quality provide
682 some evidence of grazing effects not only through alterations of plant communities that
683 were reported by previous studies in the region (Canals and Sebastià, 2000; Sebastià et
684 al., 2008), but also through interactions with them. Although grazing effects were not the
685 most important factors affecting SOC stocks, this is by far the easiest component to
686 manipulate in order to increase or maintain SOC in soils and face climate change (Komac
687 et al., 2014). Considering our results, we suggest conducting more experiments to
688 investigate grazer type effects on SOC under different soil nutrient conditions, and within
689 plant communities with contrasting herbage quality parameters. Grazing management
690 also has other advantages such as preventing the accumulation of aboveground C, and
691 reducing the risk of forest fires (Nunes and Lourenço, 2017).

692 One key point of our results is that reinforce the idea that grazer type might be at least
693 as important as grazing intensity in regulating grassland ecosystem dynamics (Tóth et
694 al., 2018), and highlight the need for a more thorough research effort in disentangling
695 not only grazing intensity but also grazer type effects on grassland soil organic carbon
696 and nutrient cycling, under different environmental circumstances. The combined model
697 provided some evidence supporting that grazing may affect SOC not only through
698 alterations of plant communities (Canals and Sebastià, 2000; Sebastià et al., 2008), but
699 also through interactions with them. Although grazing effects were not the most important
700 factors affecting SOC stocks, this is by far the easiest component to manipulate in order
701 to increase or maintain SOC in soils and face climate change (Komac et al., 2014).
702 Despite the need of a precise knowledge on the effects of different land uses on
703 ecosystems for climate change mitigation (Lo et al., 2015), studies addressing grazer
704 type effects on SOC are scarce (i.e. Zhou et al., 2017; Chang et al., 2018). Considering
705 our results, we suggest conducting more experiments which investigate grazer type
706 effects on SOC under different soil nutrient conditions, and within plant communities with
707 contrasting herbage quality parameters.

708 **Conclusion**

709 The models presented here show a series of novel broad-scale and local patterns
710 concerning SOC stocks and their geophysical, biochemical and grazing management
711 drivers. Factors driving SOC stocks often interacted in complex ways, within and
712 between spatio-temporal scales. Temperature seasonality (TSIS) was the most critical
713 geophysical factor, affecting SOC through interactions with topographical drivers and
714 grazing intensity. This illustrates how not only climate mean annual conditions should be
715 considered when modelling SOC drivers, but also seasonal patterns. Concerning
716 biochemical factors, we found that the expected positive effect of soil N was modulated
717 by livestock species and herbage NDF; and herbage recalcitrance effects on SOC varied
718 depending on grazer type. Overall, we found a number of interactions highlighting the
719 need to expand knowledge on grassland SOC drivers under different conditions,
720 specially grazing. The latter is the most easily tractable factor affecting SOC. In
721 conclusion, we provided valuable information for further studies dealing with SOC
722 predictions at broad several scales, and laid out the basis to generate new testable
723 hypotheses for future studies, which may be useful for designing improved policies to
724 palliate climate change.

725 **DATA ACCESSIBILITY**

726 Data are not public as the PASTUS database is currently being used for other research
727 projects. Please contact one of us by e-mail for queries concerning the data used in this
728 study.

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1111 **Table captions**

1112 Table 1: Considered factors affecting SOC stocks in some recent studies. V: the study
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1145 **Tables**

1146 Table 1: Considered factors affecting SOC stocks in some recent studies. V: the study
 1147 considers this variable type; -: the study does not consider this variable type.

1148 1: It considers SOC concentrations
 1149 2: It considers total carbon stocks
 1150 3: It considers total carbon stocks and its fractions.

Article	Location	LAT (°)	LONG (°)	MAP (mm)	MAT (°C)	Topography and bedrock	Grazing Management	Soil properties	Soil nutrients	Herbage
Present study	Pyrenees	42.14 – 43.3	-1.22 – 2.26	964 – 1586	1.1 – 9.9	✓	✓	✓	✓	✓
Duarte-guardia et al., 2019	Worldwide	-51.72 – 80.23	-163.95 – 158.25	65 – 5115	-21.2 – 30	✓	-	✓	-	✓**
Abdalla et al., 2018	Worldwide	-45.85 – 51	-114 – 120.7	150 – 1650	0 – 21	-	✓	✓	-	✓
Eze et al., 2018	Worldwide	-44 – 65	-121 – 175	120 – 2000	-4.8 – 26.8	-	✓	✓	✓*	✓**
Peri et al., 2018 ¹	South Patagonia	- 52 – -45	-73.5 – 65.5	139 – 865	4.2 – 11	✓	✓	-	-	✓
Zhang et al., 2018	Northern China	103.5 – 124.16	32.5 – 42.5	500 – 1000	8 – 14	✓	✓	✓	-	-
Zhao et al., 2017	Mongolia	41.95 – 53.93	108.28- 116.2	150 – 400	-1.3 – 2.1	-	✓	✓	-	✓
Zhou et al., 2017 ²	Worldwide	-42.1 – 52.3	-121 – 175	200 – 600	0 – 10	-	✓	-	-	✗
Deng et al., 2016	Eastern China	28.71 – 30.45	120.87 – 122.43	940 – 1720	16.86 – 18.57	✓	-	✓	-	✗
Gray et al., 2015	Eastern Australia	-16.7 – -43.5	-31.8 – -28.7	500 – 2000	10 – 30	✓	✗	✗	-	✓
Lu et al., 2017	Qinghai-Tibetan Plateau	27 – 32	83 – 108	37 – 718	-4.04 – 6.3	-	✓	✗	-	-
Chang et al., 2015 ¹	Tibet	Not Reported	Not Reported	397 – 1910	1.7 – 15.5	✓	-	-	-	✓
Manning et al. 2015 ³	England	50.77– 54.58	-4.43 – 0.87	596 – 3191	6.5 – 10.9	-	✓	✓	-	✓
McSherry & Ritzie 2013	Worldwide	Not reported	Not reported	180 – 950	Not reported	-	✓	✓	-	✓
Garcia-Pausas et al. 2007	Pyrenees	-7 – 2.2	42.5 – 42.9	1416 – 1904	-0.7 – 5	✓	-	✓	-	-

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1152

1153 *Fertilizer effects.

1154 ** Only aboveground and/or belowground biomass index.

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1157 Table 2: Results of the geophysical model for soil organic carbon ($R^2_{Adj} = 0.34$). MAP: mean
 1158 annual precipitation; TSIS: temperature seasonality; Slope: terrain slope; Exposed: Exposed
 1159 position according to Macrotopography; Clay: clay content; Low and medium intensity: Low and
 1160 medium Grazing intensity.

Model term	Estimate	SE	t-value	P-value	
Intercept	-0.525	1.802	-0.291	0.771	
Climate variables					
MAP	0.003	0.001	4.560	<0.001	***
TSIS	-0.098	0.228	-0.429	0.669	
Topography variables					
Slope	-0.339	0.095	-3.569	0.001	***
Exposed	-3.130	0.936	-3.344	0.001	**
Soil type variables					
Clay	0.121	0.027	4.500	<0.001	***
Management variables					
Low intensity	-5.013	1.196	-4.192	<0.001	***
Medium intensity	2.012	1.168	1.722	0.088	
Interactions					
TSIS x Exposed	0.417	0.124	3.358	0.001	**
TSIS x Slope	0.044	0.013	3.452	0.001	***
MAP x Clay	9×10^{-5}	2×10^{-5}	-4.637	<0.001	***
TSIS x Low intensity	0.655	0.159	4.110	<0.001	***
TSIS x Medium intensity	-0.262	0.156	-1.684	0.095	

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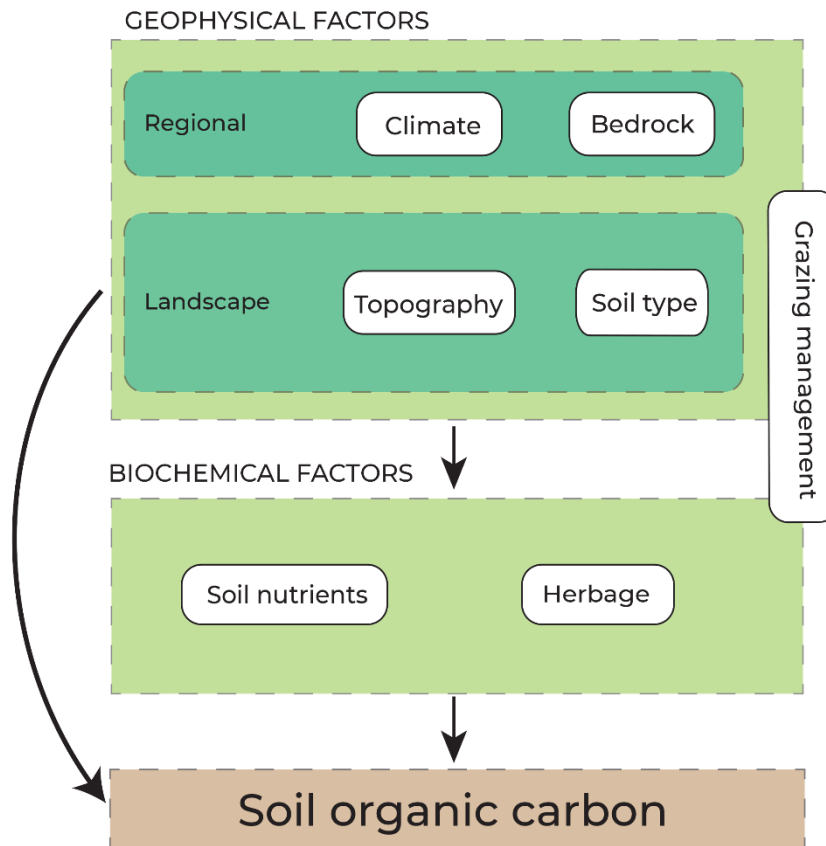
1162

1163 Table 3: Results of the combined model for soil organic carbon ($R^2_{Adj} = 0.61$). MAP: mean
 1164 annual precipitation; TSIS: mean summer temperature minus mean annual temperature; Slope:
 1165 terrain slope; Cattle and Mixed: Cattle and mixed management according to grazing species;
 1166 Low and medium intensity: Low and medium intensity according to Grazing intensity; Soil C/N:
 1167 soil carbon to nitrogen ratio; soil N: soil nitrogen; NDF: neutro-detergent fibre; ADL/NH: acid-
 1168 detergent lignin to nitrogen in the herbage ratio.

Model term	Estimate	SE	t-value	P-value	
Intercept	-0.290	1.458	-0.199	0.843	
Climate variables					
MAP	-0.001	0.000	-2.434	0.017	*
TSIS	-0.004	0.181	-0.022	0.982	
Topography variables					
Slope	-0.225	0.078	-2.868	0.005	**
Management variables					
Cattle	0.487	0.101	4.834	<0.001	***
Mixed	-0.289	0.093	-3.106	0.002	**
Low intensity	-3.249	1.014	-3.204	0.002	**
Medium intensity	1.666	1.073	1.553	0.123	
Soil nutrient variables					
Log(Soil C/N)	0.665	0.076	8.777	<0.001	***
Soil N	3.302	0.617	5.349	<0.001	***
Herbage variables					
NDF	0.014	0.006	2.525	0.013	*
Herbage ADL/NH	0.026	0.009	2.987	0.003	**
Interactions between variable types					
TSIS x Slope	0.030	0.010	2.833	0.006	**
TSIS x Low intensity	0.423	0.136	3.104	0.002	**
TSIS x Medium intensity	-0.214	0.143	-1.495	0.138	
Soil N x Cattle grazing	-0.736	0.168	-4.380	<0.001	***
Soil N x Mixed grazing	0.493	0.175	2.813	0.006	**
Soil N x NDF	-0.039	0.011	-3.505	0.001	***
Cattle x Herbage ADL/NH	-0.030	0.010	-2.872	0.005	**
Mixed x Herbage ADL/NH	0.014	0.011	1.252	0.213	

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1170 **Figures**

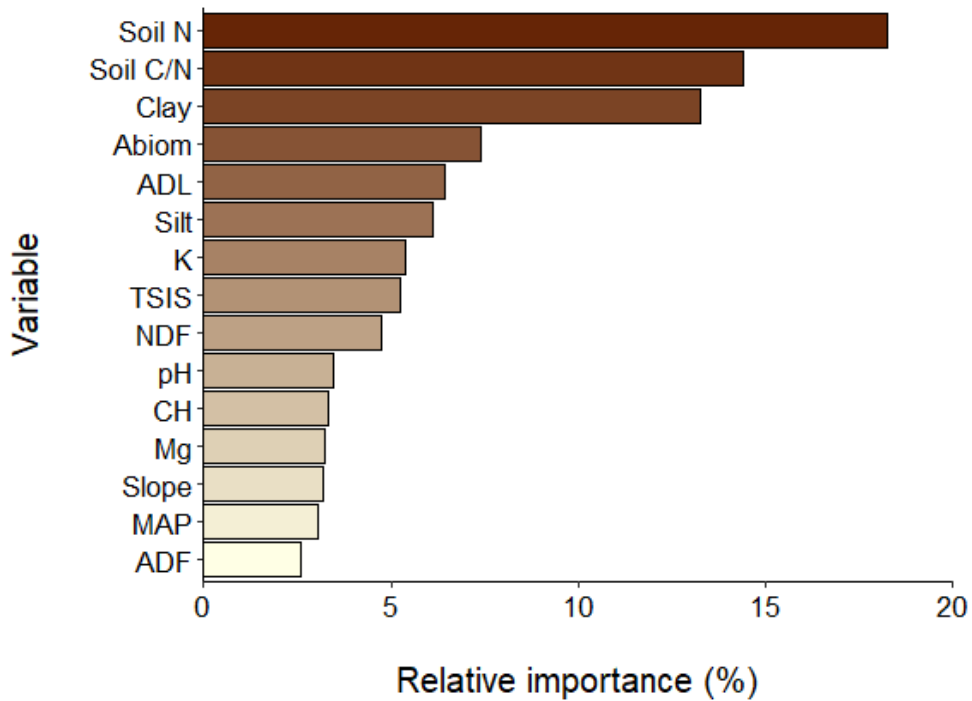


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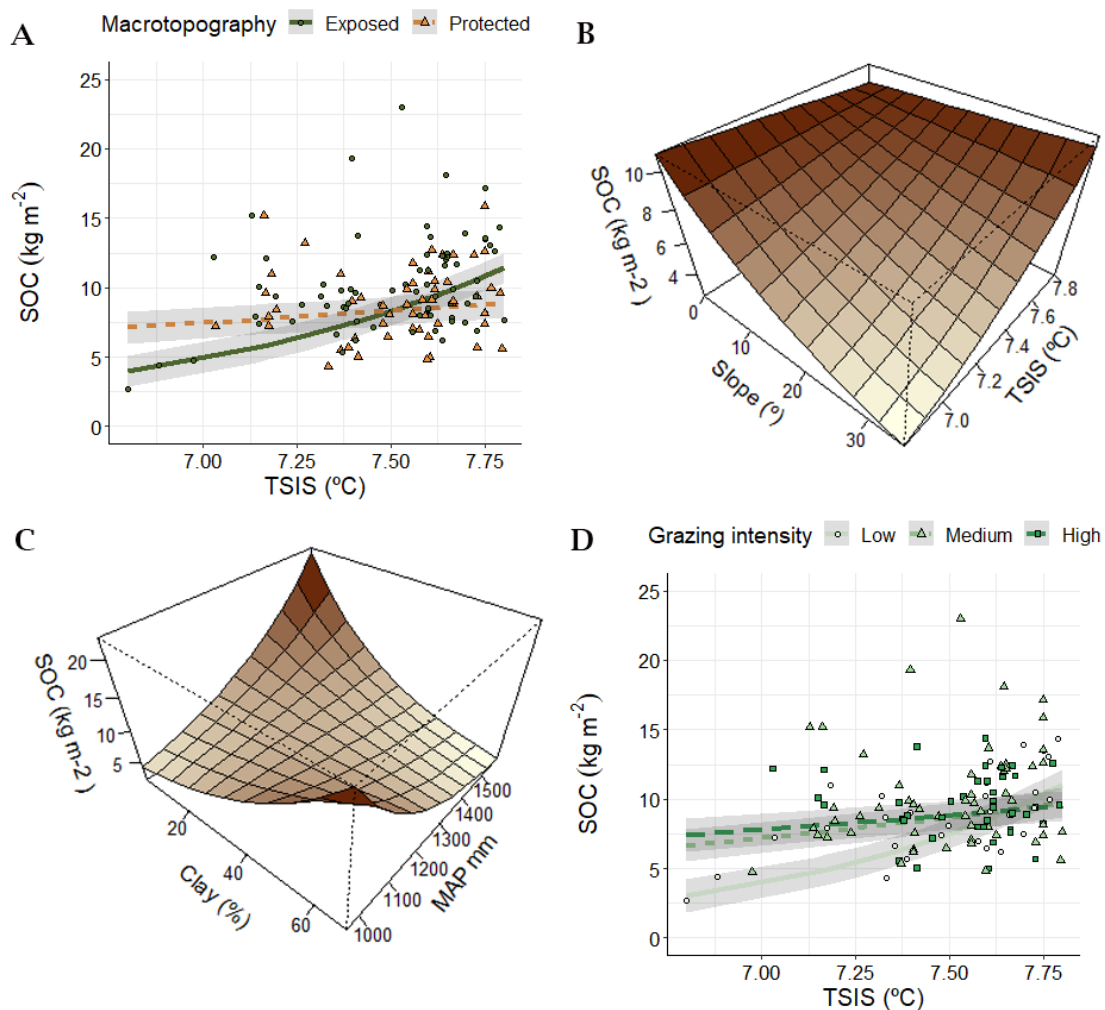
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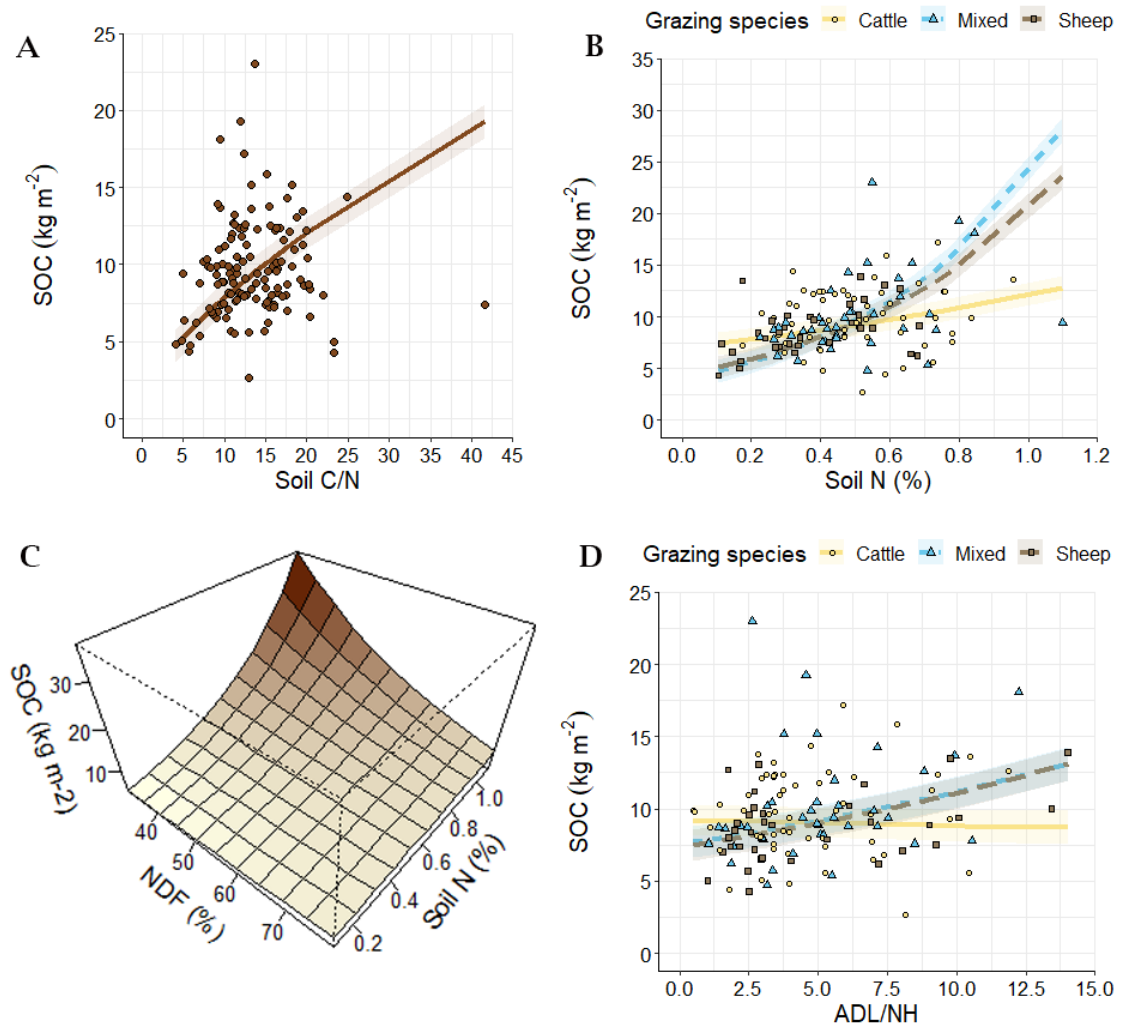
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