



1 **Interactions between biogeochemical and management factors explain soil**
2 **organic carbon in Pyrenean grasslands**

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29 **Author contributions**

30 Antonio Rodríguez designed the statistical procedure, carried out the statistical analyses
31 and wrote the original draft.

32 Rosa M Canals was responsible from field monitoring, lab analyses and acquisition of
33 information for the data base implementation in the Western Pyrenees (Navarra). She
34 has reviewed the draft.

35 Elena Albanell designed NIRS study and reviewed the draft.

36 Haifa Debouk sampled and processed some of the data in the PASTUS Database and
37 reviewed the draft.

38 Jordi García-Pausas processed some of the data in the PASTUS Database and
39 reviewed the draft.

40 Josefina Plaixats carried out the chemical analysis of herbage samples for NIR
41 calibration and validation equations and reviewed the original draft.

42 Leticia San Emeterio designed methodology and data collection, performed soil and
43 vegetation sampling. She has reviewed the draft.

44 Juan José Jiménez collaborated in the fieldwork and reviewed the draft.

45 M.-Teresa Sebastià contributed to the conception, design and development of the
46 PASTUS database. In addition, she ensured funding and coordinated the projects whose



47 data are included in PASTUS. Finally, she contributed to initial modelling, supervised the
48 development of the paper, and read and reviewed the drafts.

49



50 **Abstract**

51 Grasslands are one of the major sinks of terrestrial soil organic carbon (SOC).
52 Understanding how environmental and management factors drive SOC is challenging
53 because they are scale-dependent, with large scale drivers affecting SOC both directly
54 and through drivers working at detailed spatial scales. Here we addressed how regional,
55 landscape and grazing management, soil properties and nutrients and herbage quality
56 factors affect SOC in mountain grasslands in the Pyrenees. Taking advantage of the high
57 variety of environmental heterogeneity in the Pyrenees, we fit a set of models with
58 explicative purposes using data that comprise a wide range of environmental and
59 management conditions. We found that temperature seasonality (TSIS) was the most
60 important geophysical driver of SOC in our study. TSIS was positively related to SOC
61 but only under certain local conditions: exposed hillsides, steep slopes and relatively
62 highly grazed areas. High TSIS conditions probably are more favourable for plant
63 biomass production, but landscape and grazing management factors buffer the
64 accumulation of this biomass into SOC. Concerning biochemical SOC predictors, we
65 obtained some surprising, interactive effects between grazer type, soil nutrients and
66 herbage quality. Soil N was a crucial factor modulating effects of livestock species and
67 neutral detergent fibre content of plant biomass and herbage recalcitrance effects varied
68 depending on grazer species. These results highlight the gaps in the knowledge about
69 SOC drivers in grassland under different environmental and management conditions,
70 and they may serve to generate testable hypothesis in latter studies directed to climate
71 change mitigation policies.

72 **Keywords**

73 SOC, semi-natural grasslands, grazing management, climate change, soil nutrients

74 **Introduction**



75 Soil organic carbon (SOC) plays key roles in the terrestrial ecosystems (Lal,
76 2004a). SOC enhances soil and water quality and biomass productivity, and has
77 an important role in relation to climate change (Lal, 2004b). Although grasslands
78 have small aboveground biomass compared to other ecosystems, their SOC
79 stocks can be comparable to those in forest ecosystems (Berninger et al., 2015).
80 This is due to their high root biomass and residues, which are a substantial
81 carbon source and can contribute to water retention in soil. This creates
82 favourable conditions for the accumulation of organic matter (Von Haden and
83 Dornbush, 2014; Yang et al., 2018). These attributes, together with the high
84 extent of grassland global cover, make grasslands store around 34% of the
85 terrestrial carbon, mostly in their soils (White et al., 2000).

86 SOC accumulation results from a complex equilibrium between primary
87 production and organic matter decomposition which depends on multiple
88 environmental factors such as climate, soil texture and nutrients or land
89 management (Jenny, 1941; Schlesinger, 1977). Understanding how these
90 environmental factors drive SOC is challenging because they are scale-
91 dependent and are disposed on a hierarchy of controls over SOC, so large scale
92 drivers affect also those working at fine spatial scales (Fig. 1; Manning et al.,
93 2015). Climate is known to be the main SOC driver at broad (global and regional)
94 scales; mean annual precipitation (MAP) and mean temperature (MAT) being
95 the most frequent climate indicators (Wiesmeier et al., 2019). However, climate
96 seasonality variables are be commonly neglected drivers affecting SOC in broad-
97 scale models, in spite of being some important factors for plant primary
98 production or enzymatic activity of soil microorganisms (Fernández-Alonso et al.,
99 2018; Garcia-Pausas et al., 2007; Puissant et al., 2018). Other regional and



100 landscape factors like bedrock or topography are also considered to be at the top
101 of the hierarchy because they influence multiple geophysical and biochemical
102 factors affecting SOC, including soil texture or water flow paths (Gray et al., 2015;
103 Hobbey et al., 2015). Next in the hierarchy after regional and landscape factors,
104 are several soil geophysical properties, like pH and texture, which are controlled
105 by climate, bedrock, and which affect SOC through both plant primary production
106 and microbial activity and the capacity to stabilise the SOC (Deng et al., 2016;
107 Xu et al., 2016a). Soil macro and micronutrients are in the next level of the
108 hierarchy, as their abundance is determined by multiple factors, including climate,
109 soil pH, water content or clay content (Hook and Burke, 2000; de Vries et al.,
110 2012). They play an essential role in primary production and herbage quality, and
111 act as resources for microbes to mineralise SOC (Aerts and Chapin, 1999;
112 Vitousek and Howarth, 1991). However, these variables are commonly omitted
113 in the broad-scale SOC studies, especially if those focus on predictive models
114 instead of explicative ones (Gray et al., 2015; Manning et al., 2015; Zhang et al.,
115 2018). This kind of variables are less frequently available and more difficult to
116 measure than the other indicators used for SOC modelling (Manning et al., 2015).
117 Moreover, the use of soil nutrients as SOC predictors in linear models can be
118 challenging, as they are often so linked to SOC dynamics that their effect can
119 mask the effect of other predictors at higher levels (Bing et al., 2016; Cleveland
120 and Liptzin, 2007; Tipping et al., 2016). Vegetation represents another group of
121 SOC predictors, affected by climate, topography and soil properties and nutrients
122 (Fernández-Martínez et al., 2014; de Vries et al., 2012; Zhu et al., 2019). Plant
123 biomass is the main input of organic carbon into the soil (Shipley and Parent,
124 1991). However, plant litter quality can determine decomposition rates and



125 patterns, and hence soil carbon sequestration (Ottoy et al., 2017; Yan et al., 2018,
126 2019).

127 Apart from these factors, management effects on grassland SOC is a noteworthy
128 issue since they are poorly understood (Wiesmeier et al., 2019). It is known that
129 herbivores can affect SOC through different paths, such as regulating the quantity
130 and quality of organic matter returned to soil (Bardgett and Wardle, 2003), or
131 affecting soil respiration and nutrients by animal trampling or soil microbiota
132 alteration (Lu et al., 2017). However, most of the studies investigating grazing
133 effects on SOC focus on grazing intensity, in spite of evidence pointing to a
134 greater role of grazer species in determining vegetation and SOC (Chang et al.,
135 2018; Sebastia et al., 2008). Moreover, several studies describing interactions of
136 grazing with other SOC predictors at diverse scales have been published (Abdalla
137 et al., 2018; Eze et al., 2018; Lu et al., 2015, 2017; Zhou et al., 2017). Hence,
138 grazing management on grasslands may be considered a unique SOC driver,
139 because it has effects at multiple levels of the driver hierarchy (Fig. 1).

140 In this study, our goal was to identify the main drivers of SOC stocks in semi-
141 natural grasslands of the Pyrenees, assess the interactions between them and
142 describe their relative importance. Mountain grasslands comprise a wide range
143 of climatic, topographic, management and edaphic conditions that make carbon
144 stocks highly variable (Garcia-Pausas et al., 2007, 2017). For this reason data
145 analysed here comprise a wide range of environmental conditions, comparable
146 to studies on SOC developed at continental or even worldwide scales (Table 1).
147 Additionally, we consider an exceptionally broad compilation of predictors (Table
148 1). In particular, the specific questions of this study are 1) how are the effects of
149 the geophysical, widely used predictors located at the top of the hierarchy of



150 controls on SOC? 2) how are the effects of the biochemical, unfrequently used
151 (soil nutrient and herbage), predictors on SOC? 3) Can grazing management
152 regulate the effects of other SOC drivers located at different levels of the
153 hierarchy of controls?

154 **Material & methods**

155 **2.1 Location and sampling design**

156 The set of data used in this study has been extracted from the PASTUS Database
157 (<http://ecofun.ctfc.cat/?p=3538>), which was compiled by the Laboratory of Functional
158 Ecology and Global Change (ECOFUN) of the Forest Sciences Centre of Catalonia
159 (CTFC) and the University of Lleida (UdL). We sourced a wealth of data of 128 grassland
160 patches distributed across the Pyrenees (Fig.S1), and including topographical,
161 climatological, soil, herbage and management variables. The sampled area
162 encompasses a wide variety of temperate and cold-temperate climates, with different
163 precipitation conditions, depending on altitude and geographical location from
164 Mediterranean to Continental and Boreo-Alpine (de Lamo & Sebastià, 2006; Rodríguez
165 et al., 2018; Table 1).

166 Sampling in the PASTUS database was designed according to a stratified random
167 scheme, where samples were selected at random within strata. This process was done
168 using the software ArcMap 10 (ESRI, Redlands, CA, USA). The basis for randomization
169 was the map of habitats of Catalonia 1:50000 (Carreras and Diego, 2006) for the Eastern
170 and Central sectors, the map of habitats of Madres-Coronat 1:10000 (Penin, 1997) for
171 the North-Eastern sector and the land use map of Navarra 1:25000 (Gobierno de
172 Navarra, 2003) for the Western sectors. Four variables were initially considered for
173 sampling stratification within each sector: altitude (< 1800 m; 1800-2300 m; > 2300 m),
174 slope (0-20°; 20-30°; > 30°), macrotopography (mountain top/southern-facing slope;
175 valley bottom/northern-facing slope) and grazing management (sheep grazing; cattle



176 grazing; mixed grazing). Accordingly, we determined a set of homogeneous grassland
177 patches by crossing the stratification variable layers. Grassland patches were then listed
178 by type and arranged within each list randomly to determine sampling priority. At least
179 one to two replicates of each patch type were sampled.

180 In each sampled grassland patch, a 10 x 10 m² plot was systematically placed in the
181 middle of each homogeneous grassland patch, including a particular plant community.
182 Soils and vegetation were sampled inside this 100 m² plot, and environmental variables
183 assessed (see Rodríguez et al., (2018) for additional details about sampling design).
184 Local variables were assessed inside the 100 m² plots. Aboveground biomass was
185 estimated from herbage cut at ground level in four 50 x 50 cm² quadrats placed in a 2 x
186 2 m² subplot inside the 100 m² plot. Herbage from two of the four quadrats were dried
187 and sent to the laboratory for duplicated chemico-bromatological analysis. In addition, in
188 each quadrat, a 20-cm depth soil core was extracted with a 5 x 5 cm probe after herbage
189 was removed. The soil sample in the probe was separated into two soil layers: 0-10 and
190 10-20 cm.

191 **2.2 Regional and landscape environmental drivers**

192 In order to investigate the relationship between soil organic carbon (SOC) and potential
193 environmental drivers, 29 independent environmental variables were initially considered
194 (Table S1). These variables were grouped into five sets: Regional, landscape: livestock
195 management, soil nutrient stocks and herbage variables.

196 Regional variables included climate variables and bedrock. Climate variables were
197 determined from Worldclim 2.0 (Fick and Hijmans, 2017). We selected Mean Annual
198 Temperature (MAT), Mean Annual Precipitation (MAP) and Mean Summer Precipitation
199 (MSP). The difference between mean annual and mean summer temperature emerged
200 as a relevant explanatory factor of soil organic carbon stocks during previous modelling
201 efforts by one of the co-authors (MTS). Latter attempts to improve models by substituting



202 this variable by other temperature indices in climatic databases (Fick and Hijmans, 2017)
203 showed that, for the PASTUS database, this variable provided higher explanatory power
204 than other temperature seasonality indices. Thus, we decided to keep it and here we
205 name it Temperature Seasonality Index of Sebastià (TSIS from now on).

206 Bedrock type was determined in the field and confirmed by the geographical maps
207 mentioned above. Bedrock was categorized into three categories: basic (marls and
208 calcareous rocks), acidic (mostly sandstones and slates) and heterogeneous.

209 Landscape variables included topography and soil type variables. Topography variables
210 included Slope, Aspect, Macrotopography and Microtopography. Slope and Aspect were
211 determined in the field by clinometer and compass respectively. Macrotopography and
212 Microtopography were determined visually in the field. Macrotopography differentiated
213 exposed from protected positions. Mountain top and south-facing slopes were identified
214 as exposed positions and valley bottoms and north-facing slopes as protected positions.
215 Microtopography considered three positions: convexities, concavities and smooth areas.
216 Soil type variables are described in section 2.4.

217 **2.3 Livestock management variables**

218 Regarding livestock management variables, detailed surveys were carried out among
219 farmers, shepherds and land managers. Two management variables were considered:
220 Grazing intensity and Grazer type. Grazing intensity was determined estimating livestock
221 stocking rates measured as livestock units ha⁻¹ (LU ha⁻¹), and treated as a semi-
222 quantitative variable with three categories (Sebastià et al. 2008): low (1; lower than 0.2
223 LU ha⁻¹), medium (2; between 0.2-0.4 LU ha⁻¹) and high (3; above 0.4 LU ha⁻¹). Grazer
224 type was categorized into three main types: sheep grazing, cattle grazing and mixed
225 grazing. Mixed grazing included associations comprising small and big livestock species,
226 mainly sheep and cattle, and more rarely horses. Sheep flocks always included some
227 goats.



228 **2.4 Soil sampling and physicochemical analysis**

229 Soil samples were air-dried and weighted. Each sample was sieved to 2 mm to separate
230 stones and gravels from the fine earth fraction; the fine fraction was sent to the laboratory
231 for physicochemical analysis. Standard physicochemical soil analyses were performed
232 in the upper 0-10 cm soil layer of all grasslands. Some analyses were also performed on
233 samples from the 10-20 cm soil layer, including: soil organic carbon, total nitrogen. For
234 those variables, we later calculated values for the whole top 20 cm soil layer.

235 All soil physicochemical analyses were performed on the fine earth, according to
236 standard soil analysis methods. Textural classes were determined by the Bouyoucos
237 method (Bouyoucos, 1936). Soil pH (measured in water), total organic carbon (TOC)
238 total nitrogen (TN), Calcium content (Ca), Extractable phosphorus (P), magnesium (Mg)
239 and potassium (K) were measured on air dried samples (Schöning et al., 2013; Solly et
240 al., 2014). Soil carbonates were determined using the Bernard calcimeter. Total carbon
241 and nitrogen (N) contents of the fine earth was determined by elemental auto-analyser.
242 The organic C fraction was determined by subtracting inorganic C in the carbonates from
243 the total C. Soil organic carbon (SOC) stocks in the upper 20 cm soil layer were then
244 estimated taking into account the organic C concentration in the sample and its bulk
245 density, and subtracting the coarse particle (> 2 mm) content, following García-Pausas
246 et al. (2007). Available phosphorus (P) was extracted by the Olsen method (Olsen, 1954)
247 Ca, Mg and K were extracted by ammonium acetate (Simard, 1993) and measured by
248 flame Atomic absorption Spectroscopy (AAS) (David, 1960)).

249 **2.5 Herbage chemical and bromatological analysis**

250 A total of two hundred samples were chemical and bromatological analysed by NIRS
251 (near infrared reflectance spectroscopy). All four herbage samples of each plot were
252 oven-dried at 60°C to constant weight. Two of the samples was sent to the laboratory.
253 Dried samples were ground to pass a 1 mm stainless steel screen (Cyclotec 1093



254 Sample mill, Tecator, Sweden) and stored at 4°C until it was needed for use. To develop
255 NIRS equations (see below) subsamples were analysed in duplicate. Procedures
256 described by AOAC were used to determine dry matter (DM) and ash content or mineral
257 matter (MM). Crude protein (CP) was determined by the Kjeldhal procedure (N x 6.25)
258 using a Kjeltec Auto 1030 Analyser (Tecator, Sweden). Samples were analysed
259 sequentially for neutral detergent fibre (NDF), acid detergent fibre (ADF) and acid
260 detergent lignin (ADL) in accordance with the method described Van Soest et al. (1991)
261 using the Ankom 200 Fibre Analyser incubator (Ankom, USA). The fibre analysis were
262 determined on an ash-free basis and without alpha amylase. We calculated two
263 additional herbage quality indexes often used in the bibliography: NDF/CP and ADL/HN
264 (Stockmann et al., 2013). For each subsample the C and N content were determined by
265 the Dumas dry combustion method, using an Elemental Analyzer EA1108 (Carlo Erba,
266 Milan, Italy).

267 **2.6 NIRS analysis**

268 NIRS data were recorded from 1,100 to 2,500 nm using a FOSS NIRSystem 5000
269 scanning monochromator (Hillerød, Denmark). Separate calibration equations
270 were generated for grassland samples. Reflectance (R) data were collected in
271 duplicate every 2 nm. A WinISI III (v. 1.6) software (FOSS, Denmark) was
272 employed for spectra data analysis and development of chemometric models.
273 Prior to calibration, log 1/R spectra were corrected for the effects of scatter using
274 the standard normal variate (SNV), detrend (DT) and multiple scatter correction
275 (MSC) and transformed into first or second derivative using different gap size
276 (nm) and smoothing interval. For each sample, the mean of the spectra from the
277 two lectures were used. Modified partial least square (MPLS) was the regression
278 method used for calibration development and cross validation was undertaken
279 using the standard methodology in the NIRS software program. The performance



280 of the model was determined by the following statistical tools: standard error of
281 calibration (SEC), standard error of cross validation (SECV); coefficient of
282 determination for calibration (R^2) and cross validation (r_{cv}^2); the ratio of
283 performance to deviation (RPD) described as the ratio of standard deviation for
284 the validation samples to the standard error of cross validation ($RPD=SD/SECV$)
285 should ideally be at least three; and the range error ratio ($RER=Range/SECV$)
286 described as the ratio of the range in the reference data to the SECV should be
287 at least 10 (Williams and Sobering, 1996; Williams et al., 2014).

288 **2.7 Statistical analyses**

289 We applied two different modelling procedures, Boosted Regression Trees (BTR) and
290 General Linear Models (GLM). All the statistical analyses were performed with the
291 software R ver. 3.4.3 (R Core Team, 2017), at 95% significance level when appropriate.

292 *Boosted regression trees global model*

293 We fitted a model with BRT to identify the most important variables affecting SOC. BRT
294 uses two algorithms: regression trees are from the classification and regression tree
295 (decision tree) group of models, and boosting builds and combines a collection of models
296 (Elith et al., 2008). We chose this method because BRT can handle multiple variables
297 better than other techniques as GLM, and can detect automatically curvilinear
298 relationships and interactions, ignoring non-informative ones. We used the *gbm* and
299 *dismo* packages (Greenwell et al., 2019; Hijmans et al., 2017), which provide several
300 functions to fit these models.

301 First, we fitted a model with all the predictors (Table S1), configured with 15 folds, a
302 Gaussian distribution of the error, a tree complexity of 5, a learning.rate of 0.005, a
303 bag.fraction of 0.666, and 5 minimum observations by node. Secondly, we reduced the
304 number of predictors by the method described in Elith et al., (2008). We estimated the



305 change in the model's predictive deviance dropping one by one each predictor
306 (supporting information), and re-fitted the model with the set of variables which actually
307 improved model performance Fig. S2). We checked the relative importance of the
308 predictors and the shape and size of the effects by partial effect plots.

309 *General linear models*

310 We designed and executed a modelling procedure based on general linear models
311 (Legendre and Legendre, 1998) and a hierarchy of controls over function (Diaz et al.,
312 2007; de Vries et al., 2012). We log-transformed SOC using natural logarithm to prevent
313 a breach of the normality assumption by the residuals of the models. We built two models
314 (Fig. S4), one model only based on geophysical predictors and grazing management
315 (Geophysical Model), and another model by adding to the former the biochemical
316 predictors: soil nutrients and herbage quality predictors (Combined Model). We
317 considered that the geophysical factors that potentially affect SOC were regional and
318 landscape (topography and soil type predictors), as they have been widely used in
319 previous studies to model and predict SOC from landscape to continental scales
320 (Manning et al., 2015; Wiesmeier et al., 2019). In addition to soil nutrients and herbage
321 variables, we included again the livestock management variables in the Combined Model
322 and looked for interactions involving these variables and biochemical predictors of SOC.

323 For model building (Fig. S4A), we added predictor groups following a sequential order.
324 For fitting the geophysical model, we started adding regional, landscape and grazing
325 management predictors, and subsequently included soil properties. Afterwards, we
326 sequentially included soil nutrients and herbage predictors to obtain the Full Model. We
327 added Management variables from the beginning of the modelling process and re-
328 included the discarded ones in each step to guarantee the detection of interactions
329 between Management variables and the rest of the predictors. Each time we added a
330 set of predictors, we first considered their main effects and some quadratic terms which



331 were found by preliminary analyses with the `scatterplot.matrix` function in the R package
332 `car` (Fox et al., 2018); afterwards we included possible level 2 interactions between all
333 the selected predictors.

334 At every step we selected several candidate terms by a semi-automatic procedure (Fig.
335 S4C) using a genetic algorithm included in the R package `glmulti` (Calcagno, 2015). We
336 used SOC as response variable at the first step, and the residuals of the previous model
337 in the remaining steps (Fig. S4B). This semi-automatic process began by obtaining a
338 best subset of models using the corrected Akaike information criterion (AICc),
339 appropriate when n/k is less than 40, being the sample size and k the number of
340 parameters in the most complex model (Symonds and Moussalli, 2011). We selected the
341 best model and its equivalents obtained by the genetic algorithm, which were those
342 within 2 Akaike information criterion-corrected (ΔAICc) values of the best model, as a
343 $\Delta\text{AICc} < 2$ indicates that the candidate model is almost as good as the best model
344 (Burnham and Anderson, 2002).

345 For this set of models, we built averaged models using the `MUMIn` package (Barton,
346 2015). We calculated partial standardized coefficients, obtained by multiplying the
347 unstandardized coefficient in the model by the partial standard deviation of the variable,
348 which is a function of the standard deviation of the variable, the sample size, the number
349 of variables in the model and the variance inflation factor of the variable (Barton, 2015).
350 We selected all the variables with significant effects (alone or in interaction with each
351 other) in the conditional average model, which was preferred over the full average model
352 because we wanted to avoid excessive shrinkage effects at this moment of the modelling
353 procedure (Grueber et al., 2011).

354 Then, we added these terms to the consolidated model, and made a selection through a
355 backward forward procedure. We used several methods to compare and determine the
356 final model, including the AICc, the adjusted determination coefficient R^2 (R_{adj}^2) and
357 model comparison techniques with the `anova()` function in R, using Chi-square tests to



358 test whether the reduction in the residual sum of squares was statistically significant.
359 Once we had the final model we assessed the significance of each term by removing it
360 and performing an F test. For estimating the significance of the main effects we also
361 removed the interaction terms in which they were involved, to avoid transferring the
362 effects of the main terms to the interaction terms (de Vries et al., 2012). We estimated
363 the variance explained by the models through the adjusted determination coefficient R^2
364 (R_{adj}^2).

365 Finally, we estimated the importance of the terms of each model by the lmg method in
366 the relaimpo package (Grömping, 2006), and drew partial effect plots making predictions
367 with the R package emmeans (Lenth et al., 2019).

368

369



370 **Results**

371 SOC stocks of the upper 20 cm layer ranged between 2.6 and 23 kg m⁻², with a
372 median and a mean value of 9.1 and 9.6 kg m⁻² respectively. Minimum, maximum,
373 median and mean values of the continuous predictors are shown in Table S2.

374 *Relative importance of SOC predictors*

375 The final BRT global model achieved a good goodness of fit, with a cross-
376 validated correlation value of 0.52% and an explained deviance of 88.31%. The
377 most important variables explaining SOC stocks (Fig. 2) were soil N (18.3 %), soil
378 C/N (14.4%) and Clay (13 %) although other variables such as Aboveground
379 biomass (7.3%), ADL (6.4%) or Silt (6.1%) were also relevant for explaining SOC
380 storage. Two of the most important variables in the BRT model, Aboveground
381 biomass and Silt, were not selected in the linear models (Tables 2 & 3). Although
382 accounting for a lower importance value than the previous variables (5%), TSIS
383 was the most relevant selected climate predictor. This variable was also relevant
384 in both linear models (Fig. S5), especially in the Geophysical Model, where TSIS
385 was the most important variable, not only as main effect, but in interaction with
386 other variables (lmg: 4 – 10%). Soil nutrient and herbage variables were also
387 important according to the Combined linear model (Fig. S6), but in this case we
388 identified that many of these effects occurred in interaction between these two
389 predictors with grazer type.

390 *Geophysical effects on SOC stocks*

391 The Geophysical Model (Table 2) explained 34 % of the total variance (R^2_{Adj}).
392 Overall, SOC stocks increased with TSIS under certain conditions: exposed



393 hillsides, high slopes and low stocking rates (Fig. 3A, 3B & 3D). On the other
394 hand, Clay had a positive relationship with SOC under low MAP values (Fig. 3C),
395 which turned into negative at high MAP values (Fig. 6C).

396 Soil nutrient and herbage effects on SOC

397 Adding nutrient and herbage predictors in the previous geophysical model to build
398 the Combined model (Table 3) increased the total variance (R^2_{Adj}) up to 61%.
399 Macrotopography and Clay effects described by the Geophysical model were
400 removed by the new model terms included. SOC increased with C/N (Fig 4A).
401 Soil nitrogen modulated the effects of livestock type and NDF on SOC. Cattle
402 grazed grasslands stored more SOC than mixed and sheep grazed grasslands
403 under low soil N conditions, whereas the opposite occurred at high soil N levels
404 (Fig. 3B). NDF had negative effects on SOC at high soil N values but had no
405 effect under low soil N levels Fig. 4C). Finally, herbage ADL/NH had positive
406 effects on SOC under mixed and sheep grazing regimes, but there was no effect
407 under cattle management (Fig. 4D).

408 Discussion

409 Regional, landscape, management, soil and herbage factors drove SOC stocks
410 in grasslands of the Pyrenees with multiple interactions. The BRT model identified
411 soil N and C/N, texture and herbage variables as the most important predictor
412 groups (Fig. 2), TSIS being the most important climate variable. Both linear
413 models followed a hierarchy of controls over function approach to ensure a
414 unique effect of each driver on SOC. Hence, some variables selected in the BRT
415 model, like aboveground biomass, silt or soil K were not included in these models
416 (Tables 2 & 3). The geophysical model showed how some climate variables (TSIS



417 and MAP) interacted with landscape (macrotopography and slope), soil clay
418 content and grazing intensity (Fig. 3). Whereas, the Combined Model provided
419 information on how herbage quality effects on SOC (NDF and ADL/NH) varied
420 depending on soil N and grazing species, and on how grazer species had
421 different effects depending on soil N content (Fig. 4).

422

423 *Considerations about the modelling procedure*

424 As a regression tree machine learning technique, the BTR model identified a set
425 of SOC predictors (Fig. 2) avoiding some of the linear model disadvantages, like
426 guarding against the elimination of good predictors correlated to others or
427 automatically modelling non-linear effects (Cutler et al., 2007; Elith et al., 2008).
428 Thus, the BRT model included some SOC predictors, like a positive logarithmic-
429 like effect of aboveground biomass or soil K on SOC (Fig. S7), which could be
430 masked by the effects of other variables in our linear models (Yang et al., 2009).
431 However, most of the variables selected and their effects were generally
432 consistent with those explained by the linear models (Fig. 3, 4, S7).
433 Consequently, we preferred to focus on the results from the linear models
434 because our approximation allowed us to build models under a hierarchy of
435 controls over function hypothesis (Manning et al., 2015). Hence, although we
436 could not establish the causal links between SOC predictors (Grace, 2006; Grace
437 and Bollen, 2005), we guaranteed that geophysical drivers included in the first
438 model were not the single common cause of variation of both biotic factors
439 included in the second model and SOC (de Vries et al., 2012). In that case, soil
440 nutrient and herbage quality predictors could not be added to the model as



441 significant terms, as was the case with aboveground biomass. In addition, our
442 modelling approach allowed us to select biologically meaningful interactions
443 (Manning et al., 2015; de Vries et al., 2012), which cannot be done with a fully
444 automatic approach like BRT. Additionally, our sequenced modelling procedure
445 looking for the primary sources of variation, together with the stratified sampling
446 design, lead us to select a set of lowly correlated predictors for our linear models
447 (Table S3).

448

449 *Geophysical predictors driving SOC*

450 Considering the difficulties of modelling SOC in a widely heterogeneous mountain
451 environment (Garcia-Pausas et al., 2017), the Geophysical model provided
452 important information about SOC drivers in the Pyrenees. TSIS was a key
453 predictor of SOC with a varying effect depending on macrotopography, slope and
454 grazing intensity (Table 2). This result contrasts with most of the previous studies
455 addressing soil carbon in mountain grasslands, which usually pinpoint mean
456 temperature and precipitation as the most important climate drivers of SOC
457 (Hobley et al., 2015; Manning et al., 2015; Wiesmeier et al., 2019). Overall, the
458 TSIS effect on SOC was positive under certain conditions. Sites characterised by
459 low mean temperatures presented a wider spectrum of TSIS values than warm
460 sites (Fig. S8). Considering that climate regulates large scale patterns of
461 aboveground net primary production (Chapin et al., 1987), a positive effect of
462 TSIS on SOC could be associated with higher biomass accumulation in cold
463 locations with more favourable temperatures during summer, this fact reducing
464 geophysical stress for plants (Garcia-Pausas et al., 2007; Kikvidze et al., 2005).



465 This plant biomass accumulation during summer would overcome an eventual
466 increase of soil organic matter decomposition rates due to high temperatures
467 (Sanderman et al., 2003), which could even be diminished if microbial biomass
468 decreases as a result of soil moisture reduction (Puissant et al., 2018).

469 The interactions of TSIS with macrotopography and slope illustrate the capacity
470 of landscape factors to modulate macroclimate effects on soil (Hook and Burke,
471 2000). Induced microclimate changes are often the explanation for the effects of
472 topography in SOC (Lozano-García et al., 2016). In our case, SOC stocks
473 increased with temperature seasonality particularly at mountain exposed areas
474 (Fig. 3A; Table 2). In protected sites, located in shady slopes and valley bottoms,
475 the hypothesized positive effect of high TSIS values on productivity could be
476 mitigated due to lower solar radiation, longer snow-covered periods and freezing
477 episodes (Garcia-Pausas et al., 2007; López-Moreno et al., 2013). Conversely,
478 negative effects of low TSIS values on productivity could be compensated thanks
479 to more humid conditions in protected than in exposed sites (Garcia-Pausas et
480 al., 2007). Additionally, it is important to take into account that differences in SOC
481 between exposed and protected sites may also occur through other mechanisms,
482 for instance the alteration of soil physico-chemical properties like pH, soil texture
483 or stoniness (Zhang et al., 2018) or differences in vegetation (Sebastià, 2004).
484 Since we used a hierarchy of controls approach (Manning et al., 2015), these
485 topography indirect effects could be behind the exclusion on the linear models of
486 some predictors selected in the BRT model, like silt or pH (Figs. 2 & 3).

487 In addition, high TSIS values compensated SOC decrease in steep slopes,
488 probably due to reduced carbon inputs and increased carbon losses induced by
489 high soil erosion (Yuan et al., 2019 and references therein). The decrease in



490 SOC stocks under low TSIS values were also compensated by grazing pressure
491 increase (Fig 3D). Recent meta-analyses conclude that grazing has a commonly
492 decreasing, but strongly context-specific effect on SOC, depending on other
493 factors like climate, soil type vegetation or grazing intensity (Abdalla et al., 2018;
494 Eze et al., 2018; Mcsherry and Ritchie, 2013). Particularly, light and medium
495 grazing intensities can increase SOC inputs by dung deposition and promoting
496 aboveground and root biomass production (Franzluebbers et al., 2000; Zeng et
497 al., 2015). Considering that in our semi-natural grasslands all grazing intensities
498 are relatively low (see methods), our medium and high stock rates may increase
499 soil carbon inputs in low seasonality locations by enhancing productivity.

500 Interestingly, clay content and precipitation presented interacting effects on SOC
501 (Fig. 3C; Table 2). Both MAP and clay content are widely assumed to be
502 positively correlated to SOC (Wiesmeier et al., 2019). High MAP would increase
503 SOC inputs by promoting plant productivity (Author et al., 2000; Hobbey et al.,
504 2015). Clay positive effects are often attributed to a larger contact surface of soil
505 particles (Kennedy et al., 2002), the absorption of negatively charged organic
506 matter, high soil water retention and the exclusion of decomposer organisms due
507 to their low pore size (Krull et al., 2001). In our study, high water contents may
508 inhibit decomposition if a shortage of oxygen supply occurs (Xu et al., 2016b).
509 However, as MAP values increased, clay effect on SOC became negative. To
510 explain low SOC values at high MAP and high clay content, McSherry and
511 Ritchie (2013) hypothesized that finer texture soils could be waterlogged more
512 frequently, leading to inhibition of root growth and soil C allocation belowground.

513 *Biochemical predictors driving SOC*



514 Adding soil nutrient and herbage predictors to our modelling procedure implied
515 the substitution of the terms including clay content and macrotopography by the
516 newly added terms (Tables 2 & 3), highlighting the importance of indirect effects
517 of these variables on SOC through other small scale predictors (Leifeld et al.,
518 2015; Xu et al., 2016b; Zhu et al., 2019). In this case, we obtained a complex
519 model with some surprising, less frequently tested effects involving interactions
520 between graze type, soil nutrients and herbage quality variables (Table 3, Fig 4).
521 Although our interpretations have limitations because our models were based on
522 observational data, they can still provide some hints about some of the most
523 complex and unknown relationships between SOC and its drivers. In can also
524 contribute to generate testable hypotheses in latter studies.

525 As expected, SOC increased with the C/N ratio (Fig 4A), which is an indicator of
526 the difficulty of soil organic matter decomposition by soil microbes, decreasing
527 decomposition rates of SOC with increasing soil C/N (Wanyama et al., 2019; Xu
528 et al., 2016b). Conversely, total soil N conditioned livestock type effect on SOC
529 in a surprising way. Cattle grazed grasslands stored more SOC than mixed and
530 sheep grazed ones under low soil N conditions, whereas the opposite occurred
531 at high soil N content (Fig. 4B). Chang et al. (2018) found that in a N poor semi-
532 arid grassland, sheep decreased SOC content in comparison to cattle due to
533 vegetation changes caused by their feeding preference for highly palatable forbs,
534 promoting less palatable grasses which supported less root biomass. A shift
535 towards higher grass biomass with sheep grazing was also found in the Pyrenees
536 (Sebastia et al., 2008). Conversely, in our study mixed grazing increased SOC,
537 probably through effects on soil environment and decomposition processes. Our
538 results suggested that those processes could vary depending on soil conditions.



539 Negative effects of sheep grazing on SOC through their selective feeding could
540 occur mostly in poor N soils (Fig 4B). Under such conditions, palatable plants
541 could produce higher SOC inputs, since plant productivity is more reliant on the
542 ability of fixing atmospheric N of legumes (Van Der Heijden et al., 2008) and the
543 exceptional capacity of forbs to allocate C in roots is especially stimulated (Ågren
544 and Franklin, 2003; Warembourg et al., 2003). However, these processes could
545 be different under different soil N conditions, although the concrete mechanisms
546 are hard to suggest, since livestock type may affect SOC content not only through
547 changes in plant composition, but other differences in certain features of livestock
548 assemblages, like trampling, faeces deposition patterns or effects on plant
549 regrowth, which could promote differences in soil respiration and/or plant
550 productivity (Aldezabal et al., 2019; Chang et al., 2018; Liu et al., 2018), resulting
551 in different SOC levels under different grazers. Grasslands with mixed grazed
552 regimes stored even more SOC than sheep grazed ones under high soil N
553 conditions (Fig. 4B, Table 3). This result emphasises that mixed livestock
554 assemblages deserve particular attention as they can affect plant composition
555 distinctly from single grazing species regimes or alter traveling and trampling
556 behaviours of grazing animals (Chang et al., 2018; Liu et al., 2015).

557 Model terms involving herbage predictors could represent both biochemical and
558 physical pathways of litter incorporation to soil organic matter (Cottrufu 2015). In
559 our model, NDF was negatively related to SOC at high N values (Fig 4C). NDF
560 proportion represents the amount of structural compounds on litter, and hence is
561 inversely related to non-structural compounds content (Goering and Van Soest,
562 1970). The latter are the main source of organic matter formation at the early
563 stages of decomposition, and they are incorporated into microbial biomass in a



564 highly efficient way (Cotrufo et al., 2013). However, if microbial necromass is
565 recycled by microbes before its incorporation to mineral-associated organic
566 matter, it could be respired and mineralized instead of stored (Córdova et al.,
567 2018). Thus, our model suggested that incorporation of labile and fast
568 metabolized non-organic compounds to soil organic matter could be a pathway
569 of SOC allocation at high soil N in Pyrenean grasslands. At low soil N conditions,
570 induced changes in microbial composition or priming effects (De Deyn et al.,
571 2008; Fontaine et al., 2007; Wild et al., 2019; Yan et al., 2018) may disable SOC
572 accumulation through this biochemical pathway.

573

574 On the other hand, the ADL/NH ratio was positively related to SOC in sheep and
575 mixed grazed grasslands (Fig. 4D). The ADL/NH ratio is a commonly used
576 indicator for the resistance of litter to degradation, particularly at later stages of
577 decomposition (Taylor et al., 1989). Hence, the increase of SOC with ADL/NH
578 could be related to the physical pathway of soil organic matter incorporation,
579 forming coarse particulate organic matter (Cotrufo et al., 2015). Moreover, our
580 model suggests that this pathway would be inhibited under cattle grazing,
581 presumably because of their less selective diet and higher digestive efficiency
582 than sheep (Rosenthal et al., 2012; Sebastià et al., 2008). Since lignin content is
583 inversely related to plant palatability (Moore and Jung, 2001), plants with high
584 lignin content will be avoided with greater probability under sheep-based
585 management regimes (Wang et al., 2018), and that would promote differences in
586 recalcitrant litter mineralization rates. Additionally, lower diet selectivity and
587 higher digestive efficiency of cattle compared with sheep, can result into less



588 recalcitrant faeces (Wang et al., 2018), which could explain also SOC differences
589 between grazer types at high ADL/NH conditions.

590 *Implications of livestock effects on SOC*

591 One key point of our results is that they highlight the need for a deeper research
592 effort in disentangling not only grazing intensity but grazer type effects on
593 grassland soil organic carbon and nutrient cycling under different environmental
594 circumstances. Our results concerning interactions between grazer type and
595 herbage quality provide some evidence of grazing effects not only through
596 alterations of plant communities that were reported by previous studies in the
597 region (Canals and Sebastià, 2000; Sebastià et al., 2008), but also through
598 interactions with them. Although grazing effects were not the most important
599 factors affecting SOC stocks, this is by far the easiest component to manipulate
600 in order to increase or maintain SOC in soils and face climate change (Komac et
601 al., 2014). Despite the need of a precise knowledge on the effects of different
602 land uses on ecosystems for climate change mitigation (Lo et al., 2015) studies
603 addressing grazer type effects on SOC are scarce (i.e. Zhou et al., 2017; Chang
604 et al., 2018). Considering our results, we would suggest to carry out more
605 experiments testing the effects of livestock type on SOC under different soil
606 fertility conditions and plant communities with contrasting herbage quality
607 parameters.

608 To conclude, we showed how a combination of regional, landscape,
609 management, soil properties, soil nutrients and herbage factors might drive SOC
610 stocks in the Pyrenees. Among all the regional and landscape scale factors, a
611 seasonality variable, TSIS seemed to be the most decisive, although interacting



612 with some topographical drivers and grazing intensity. To our knowledge, this is
613 the first time these factors were combined together with soil nutrients and
614 herbage quality factors to model SOC. Soil N was a crucial factor modulating the
615 effect of livestock species and NDF, and herbage recalcitrance effect on SOC
616 varied depending on grazer species. Our study highlight the need to expand
617 knowledge about grassland SOC drivers under different conditions, specially
618 grazing, as this is the most easily tractable factor affecting SOC and it has other
619 advantages like preventing the accumulation of aboveground C and reducing the
620 risk of forest fires (Nunes and Lourenço, 2017). We provided the basis to
621 generate new testable hypothesis for latter studies that may be useful to design
622 improved policies to palliate climate change.

623 **DATA ACCESSIBILITY**

624 Data are not public as the PASTUS database is currently being used for other
625 research projects. Please contact one of us by e-mail for queries concerning the
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1045 **Table captions**

1046 Table 1: Considered factors affecting SOC in some recent studies. \checkmark : the study considers this variable type; \times : the
1047 study does not consider this variable type.

1048 Table 2: Results of the geophysical model for soil organic carbon ($R^2_{Adj} = 0.34$).

1049 Table 3: Results of the Combined model for soil organic carbon ($R^2_{Adj} = 0.61$).

1050 **Figure captions**

1051 Figure 1: Conceptual scheme used in this work to relate potential environmental drivers with
1052 SOC. We assume that drivers may affect soil organic carbon (SOC) both directly or hierarchically
1053 through other driver. Interactions between factors from different types could also drive SOC.
1054 Grazing management has a special location as may act through different paths and interact with
1055 factors at different scales.

1056 Figure 2: Relative contributions (%) of predictor variables in the final BRT model obtained. Soil
1057 N: soil nitrogen; Soil C/N: soil carbon to nitrogen ratio, Clay: clay content; Abiom: aboveground
1058 biomass; ADL: acid-detergent lignin; Loam: loam content; K: soil potassium; TSIS: mean
1059 summer temperature minus mean annual temperature; NDF: neutro detergent fibre; pH: soil
1060 pH; CH: carbon in the herbage; Mg: soil magnesium; Slope: terrain slope; MAP: mean annual
1061 precipitation; ADF: acid detergent fibre. See Table S1 for more details about variables

1062 Figure 3. The relationship between SOC and regional and landscape scale factors in the
1063 Geophysical model. In A) solid lines and circles represent exposed hillsides, and dotted lines and
1064 triangles indicate protected hillsides. In D) solid lines and circles indicate low grazing
1065 management intensity, dotted lines and triangles indicate medium grazing management
1066 intensity and dashed lines and squares indicate high grazing management intensity. In A-D line
1067 and plane values are predictions of the model across the corresponding predictors' range



1068 according to estimate marginal means. Grey areas around regression lines indicate standard
1069 errors. In A) and D) points indicate actual values.

1070 Figure 4. The relationship between SOC and biochemical and herbage factors in the Combined
1071 model. In B) and D) solid lines and circle points represent cattle-grazing, dashed lines and
1072 square points indicate sheep-grazing and dotted lines and triangle points indicate mixed-
1073 grazing. In A-D line and plane values are predictions of the model across the corresponding
1074 predictors' range according to estimate marginal means. In A-D line and plane values are
1075 predictions of the model across the corresponding predictors' range according to estimate
1076 marginal means. Grey spectrum indicate 95% confidence intervals. In A) and D) points indicate
1077 actual values.

1078



1080 **Tables**

1081 Table 1: Considered factors affecting SOC in some recent studies. ✓: the study considers this variable type; ✗: the
 1082 study does not consider this variable type.

1083 *Fertilizer effects.

1084 ** Only aboveground and/or belowground biomass index

Article	Location	LAT (°)	LONG (°)	MAP (mm)	MAT (°C)	Topography and bedrock	Grazing Management	Soil properties	Soil nutrients	Herbage
Present study	Pyrenees	42.14 – 43.3	-1.22 – 2.26	964 – 1586	1.1 – 9.9	✓	✓	✓	✓	✓
Duarte-guardia et al., 2019	Worldwide	-51.72 – 80.23	-163.95 – 158.25	65 – 5115	-21.2 – 30	✓	✗	✓	✗	✓**
Abdalla et al., 2018	Worldwide	-45.85 – 51	-114 – 120.7	150 – 1650	0 – 21	✗	✓	✓	✗	✓
Eze et al., 2018	Worldwide	-44 – 65	-121 – 175	120 – 2000	-4.8 – 26.8	✗	✓	✓	✓*	✓**
Peri et al., 2018	South Patagonia	-52 – -45	-73.5 – 65.5	139 – 865	4.2 – 11	✓	✓	✗	✗	✓
Zhang et al., 2018	Northern China	103.5 – 124.16	32.5 – 42.5	500 – 1000	8 – 14	✓	✓	✓	✗	✗
Zhao et al., 2017	Mongolia	41.95 – 53.93	108.28- 116.2	150 – 400	-1.3 – 2.1	✗	✓	✓	✗	✓
Zhou et al., 2017	Worldwide	-42.1 – 52.3	-121 – 175	200 – 600	0 – 10	✗	✓	✗	✗	✗
Deng et al., 2016	Eastern China	28.71 – 30.45	120.87 – 122.43	940 – 1720	16.86 – 18.57	✓	✗	✓	✗	✗
Gray et al., 2015	Eastern Australia	-16.7 – -43.5	-31.8 – -28.7	500 – 2000	10 – 30	✓	✗	✗	✗	✓
Lu et al., 2017	Qinghai-Tibetan Plateau	27 – 32	83 – 108	37 – 718	-4.04 – 6.3	✗	✓	✗	✗	✗
Chang et al., 2015	Tibet	Not Reported	Not Reported	397 – 1910	1.7 – 15.5	✓	✗	✗	✗	✓
Manning et al. 2015	England	50.77–54.58	-4.43 – 0.87	596 – 3191	6.5 – 10.9	✗	✓	✓	✗	✓
McSherry & Ritzie 2013	Worldwide	Not reported	Not reported	180 – 950	Not reported	✗	✓	✓	✗	✓
Garcia-Pausas et al. 2007	Pyrenees	-7 – 2.2	42.5 – 42.9	1416 – 1904	-0.7 – 5	✓	✗	✓	✗	✗



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1086 Table 2: Results of the geophysical model for soil organic carbon ($R^2_{Adj} = 0.34$).

Model term	Estimate	SE	t-value	P-value	
Intercept	-0.525	1.802	-0.291	0.771	
Climate variables					
MAP	0.003	0.001	4.560	<0.001	***
TSIS	-0.098	0.228	-0.429	0.669	
Topography variables					
Slope	-0.339	0.095	-3.569	0.001	***
Exposed	-3.130	0.936	-3.344	0.001	**
Soil type variables					
Clay	0.121	0.027	4.500	<0.001	***
Management variables					
Low intensity	-5.013	1.196	-4.192	<0.001	***
Medium intensity	2.012	1.168	1.722	0.088	
Interactions between variable types					
TSIS x Exposed	0.417	0.124	3.358	0.001	**
TSIS x Slope	0.044	0.013	3.452	0.001	***
MAP x Clay	0.000	0.000	-4.637	<0.001	***
TSIS x Low intensity	0.655	0.159	4.110	<0.001	***
TSIS x Medium intensity	-0.262	0.156	-1.684	0.095	

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1089 Table 3: Results of the Combined model for soil organic carbon ($R^2_{Adj} = 0.61$).

Model term	Estimate	SE	t-value	P-value	
Intercept	-0.290	1.458	-0.199	0.843	
Climate variables					
MAP	-0.001	0.000	-2.434	0.017	*
TSIS	-0.004	0.181	-0.022	0.982	
Topography variables					
Slope	-0.225	0.078	-2.868	0.005	**
Management variables					
Cattle	0.487	0.101	4.834	<0.001	***
Mixed	-0.289	0.093	-3.106	0.002	**
Low intensity	-3.249	1.014	-3.204	0.002	**
Medium intensity	1.666	1.073	1.553	0.123	
Soil nutrient variables					
Log(Soil C/N)	0.665	0.076	8.777	<0.001	***
Soil N	3.302	0.617	5.349	<0.001	***
Herbage variables					
NDF	0.014	0.006	2.525	0.013	*
Herbage ADL/NH	0.026	0.009	2.987	0.003	**
Interactions between variable types					
TSIS x Slope	0.030	0.010	2.833	0.006	**
TSIS x Low intensity	0.423	0.136	3.104	0.002	**
TSIS x Medium intensity	-0.214	0.143	-1.495	0.138	
Soil N x Cattle grazing	-0.736	0.168	-4.380	<0.001	***
Soil N x Mixed grazing	0.493	0.175	2.813	0.006	**
Soil N x NDF	-0.039	0.011	-3.505	0.001	***
Cattle x Herbage ADL/NH	-0.030	0.010	-2.872	0.005	**
Mixed x Herbage ADL/NH	0.014	0.011	1.252	0.213	

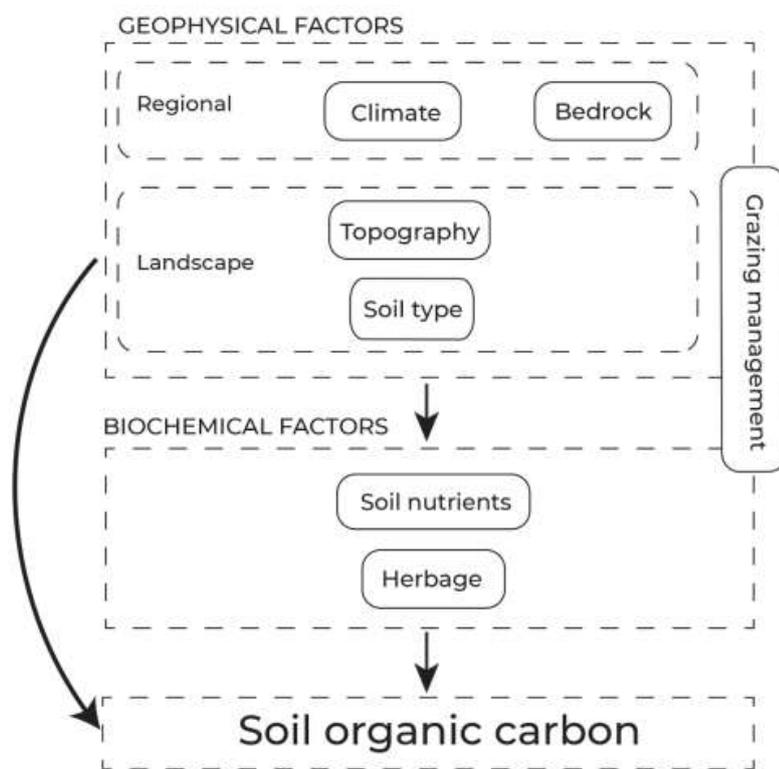
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1093 **Figures**

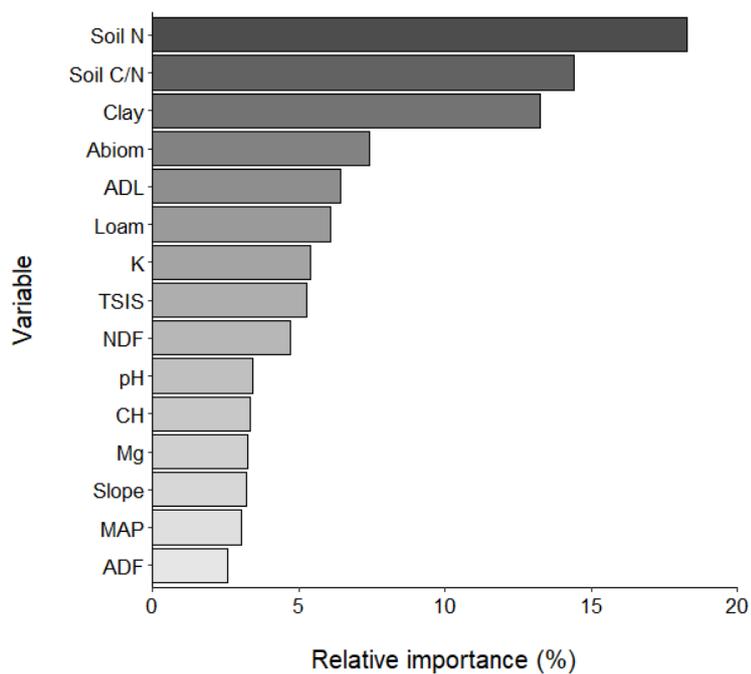


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1095 Figure 1: Conceptual scheme used in this work to relate potential environmental drivers with
1096 SOC. We assume that drivers may affect soil organic carbon (SOC) both directly or hierarchically
1097 through other driver. Interactions between factors from different types could also drive SOC.
1098 Grazing management has a special location as may act through different paths and interact with
1099 factors at different scales.

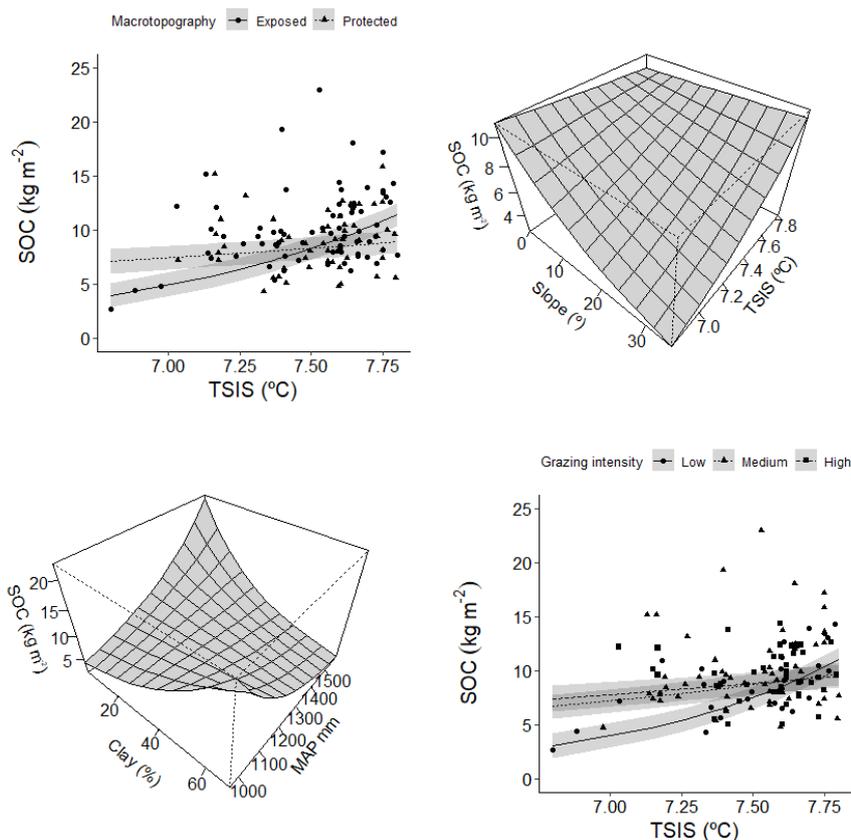
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1102 Figure 2: Relative contributions (%) of predictor variables in the final BRT model obtained. Soil
1103 N: soil nitrogen; Soil C/N: soil carbon to nitrogen ratio, Clay: clay content; Abiom: aboveground
1104 biomass; ADL: acid-detergent lignin; Loam: loam content; K: soil potassium; TSIS: mean
1105 summer temperature minus mean annual temperature; NDF: neutro detergent fibre; pH: soil
1106 pH; CH: carbon in the herbage; Mg: soil magnesium; Slope: terrain slope; MAP: mean annual
1107 precipitation; ADF: acid detergent fibre. See Table S1 for more details about variables.

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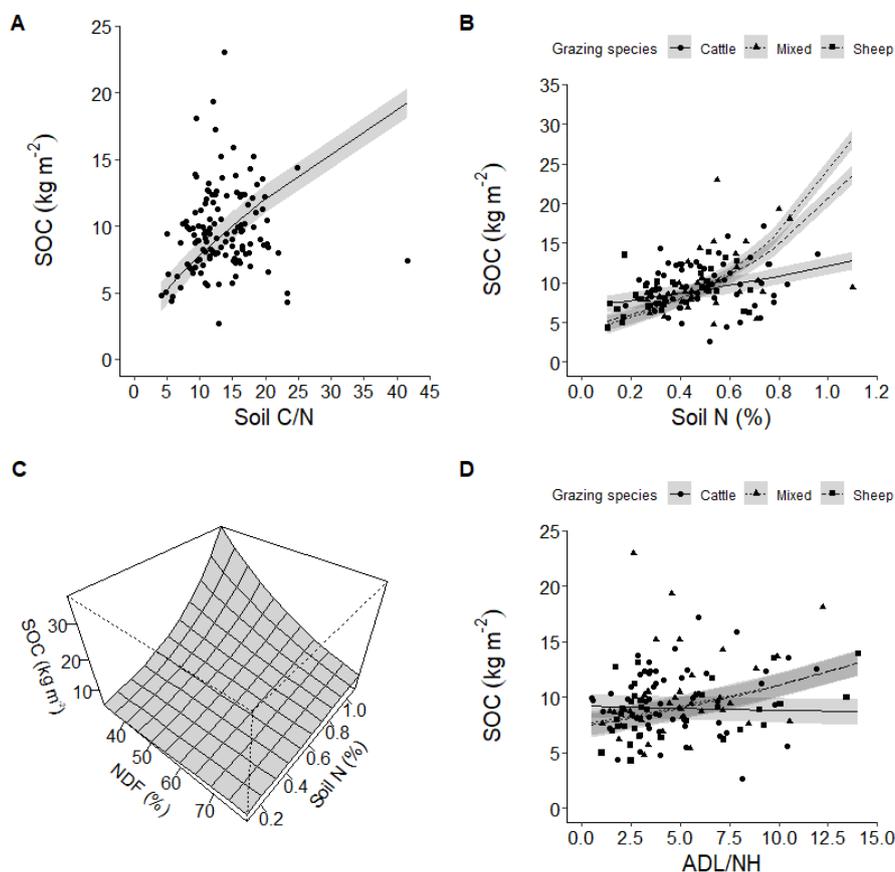


1109 Figure 3. The relationship between SOC and regional and landscape scale factors in the
1110 Geophysical model. In A) solid lines and circles represent exposed hillsides, and dotted lines and
1111 triangles indicate protected hillsides. In D) solid lines and circles indicate low grazing
1112 management intensity, dotted lines and triangles indicate medium grazing management
1113 intensity and dashed lines and squares indicate high grazing management intensity. In A-D line
1114 and plane values are predictions of the model across the corresponding predictors' range
1115 according to estimate marginal means. Grey areas around regression lines indicate standard
1116 errors. In A) and D) points indicate actual values.
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1121 Figure 4. The relationship between SOC and biochemical and herbage factors in the Combined
1122 model. In B) and D) solid lines and circle points represent cattle-grazing, dashed lines and
1123 square points indicate sheep-grazing and dotted lines and triangle points indicate mixed-
1124 grazing. In A-D line and plane values are predictions of the model across the corresponding
1125 predictors' range according to estimate marginal means. In A-D line and plane values are
1126 predictions of the model across the corresponding predictors' range according to estimate
1127 marginal means. Grey spectrum indicate 95% confidence intervals. In A) and D) points indicate
1128 actual values.