Leaf-scale quantification of the effect of photosynthetic gas exchange on $\Delta^{17}O$ of atmospheric CO_2

Getachew Agmuas Adnew[1], Thijs L. Pons[2], Gerbrand Koren[3], Wouter Peters[3,4], Thomas Röckmann[1]

[1]Institute for Marine and Atmospheric research Utrecht (IMAU), Utrecht University, The Netherlands
[2]Institute of Environmental Biology, Utrecht University, The Netherlands
[3]Department of Meteorology and Air Quality, Wageningen University, The Netherlands
[4]Centre for Isotope Research, University of Groningen, The Netherlands

Correspondence to: Getachew Agmuas Adnew (getachewagmuas@gmail.com)

Abstract

Understanding the processes that affect the triple oxygen isotope composition of atmospheric CO_2 during gas exchange can help constrain the interaction and fluxes between the atmosphere and the biosphere. We conducted leaf cuvette experiments under controlled conditions, using three plant species. The experiments were conducted at two different light intensities and using CO_2 with different $\Delta^{17}O$. We directly quantify for the first time the effect of photosynthesis on $\Delta^{17}O$ of atmospheric CO_2. Our results demonstrate the established theory for $\delta^{18}O$ is applicable to $\Delta^{17}O$-CO_2 at leaf-level and we confirm the two key factors determine the effect of photosynthetic gas exchange on the $\Delta^{17}O$ of atmospheric CO_2. The relative difference between $\Delta^{17}O$ of the CO_2 entering the leaf and the CO_2 in equilibrium with leaf water, and the back-diffusion flux of CO_2 from the leaf to the atmosphere, which can be quantified by the c_m/c_a ratio where c_a is the CO_2 mole fraction in the surrounding air and c_m the one at the site of oxygen isotope exchange between CO_2 and H_2O. At low c_m/c_a ratio the discrimination is governed mainly by diffusion into the leaf, and at high c_m/c_a ratio by back-diffusion of CO_2 that has equilibrated with the leaf water. Plants with a higher c_m/c_a ratio modify the $\Delta^{17}O$ of atmospheric CO_2 more strongly than plants with a lower c_m/c_a ratio. Based on the leaf cuvette experiments, the global value for discrimination against $\Delta^{17}O$ of atmospheric CO_2 during photosynthetic gas exchange is estimated to be -0.57 ± 0.14 ‰ using c_m/c_a values of 0.3 and 0.7 for C_4 and C_3 plants, respectively. The main uncertainties in this global estimate arise from variation in c_m/c_a ratios among plants and growth conditions.

1. Introduction

Stable isotope measurements of CO_2 provide important information on the magnitude of the CO_2 fluxes between atmosphere and biosphere, which are the largest components of the global carbon cycle (Farquhar et al., 1989; 1993; Ciais et al., 1997a; 1997b; Flanagan and Ehleringer, 1998; Yakir and Sternberg, 2000; Gillon and Yakir, 2001; Cuntz et al., 2003a; 2003b). A better understanding of the terrestrial carbon cycle is essential for predicting future climate and atmospheric CO_2 mole fractions (Booth et al., 2012). Gross primary productivity (GPP), the total carbon dioxide uptake by vegetation during photosynthesis, can only be determined indirectly and remains poorly constrained (Cuntz, 2011; Welp et al., 2011). For example, Beer et al. (2010) estimated global GPP to be 102-135 PgC yr^{-1} (85 % confidence interval, CI) using machine learning techniques by extrapolating from a database of eddy-covariance measurements of CO_2 fluxes. This estimate has since then been widely used as target for terrestrial vegetation models (Sitch et al., 2015), and replicated based on cross-consistency checks with atmospheric inversions, sun-induced fluorescence (SIF) and global vegetation models (Jung et al., 2020). As an alternative, Welp et al. (2011) estimated global GPP to be 150-175 PgC yr^{-1} using variations in $\delta^{18}O$ of atmospheric CO_2 after the 1997/98 El Nino event; see equation 1 for definition of the δ value.

The concept behind the latter study was that atmospheric CO_2 exchanges oxygen isotopes with leaf and soil water, and this isotope exchange mostly determines the observed variations in $\delta^{18}O$ of CO_2 (Francey and Tans, 1987; Yakir, 1998). Following the 97/98 ENSO event, the anomalous $\delta^{18}O$ signature imposed on tropical leaf and soil waters was transferred to atmospheric CO_2, before slowly disappearing as a function of the lifetime of atmospheric CO_2. This in turn is governed by the land vegetation uptake of CO_2 during photosynthesis, as well as soil invasion of CO_2 (Miller et al., 1999; Wingate et al., 2009). For the photosynthesis term, the equilibration of CO_2 with water is an uncertain parameter in this calculation, partly because the $\delta^{18}O$ of water at the site of isotope exchange in the leaf is not well defined. Importantly, a significant $\delta^{18}O$ variation can occur in leaves due to the preferential evaporation of $H_2{}^{16}O$ relative to $H_2{}^{18}O$ (Gan et al., 2002; Farquhar and Gan, 2003; Gan et al., 2003; Cernusak et al., 2016), which induces a considerable uncertainty in estimating $\delta^{18}O$ of CO_2. Similar considerations for the transfer of the $\delta^{18}O$ signature of precipitation into the soils, and then up through the roots, stems, and leaves make ^{18}O of CO_2 a challenging measurement to interpret (Peylin et al., 1999; Cuntz et al., 2003a; 2003b).

Classical isotope theory posits that oxygen isotope distributions are modified in a mass-dependent way. This means that the $^{17}O/^{16}O$ ratio changes by approximately half of the corresponding change in $^{18}O/^{16}O$ (equation 2), and it applies to the processes involved in gas exchange between atmosphere and plants. However, in 1983 Thiemens and co-workers (Heidenreich and Thiemens, 1983; Thiemens, 1983; Heidenreich and Thiemens, 1986) reported a deviation from mass-dependent isotope fractionation in ozone (O_3) formation called mass-independent isotope fractionation ($\Delta^{17}O$, equation 3). In the stratosphere, the $\Delta^{17}O$ of O_3 is transferred to CO_2 via isotope exchange of CO_2 with $O(^1D)$ produced from O_3 photolysis (Yung et al., 1991; Yung et al., 1997; Shaheen et al., 2007), which results a large $\Delta^{17}O$ in stratospheric CO_2 (Thiemens et al., 1991; 1995;Lyons, 2001; Lämmerzahl et al., 2002; Thiemens, 2006; Kawagucci et al., 2008; Wiegel et al., 2013).

Once $\Delta^{17}O$ has been created in stratospheric CO_2, the only process that modify its signal is isotope exchange with leaf water, soil water and ocean water at the Earth's surface, after CO_2 has re-entered the troposphere (Boering, 2004; Thiemens et al., 2014; Liang and Mahata, 2015; Hofmann et al., 2017). Isotope exchange with leaf water is more efficient relative to ocean water due to the presence of the enzyme carbonic anhydrase (CA), which effectively catalyzes the conversion of CO_2 and H_2O to HCO_3^- and H^+ and vice versa (Francey and Tans, 1987; Friedli et al., 1987; Badger and Price, 1994; Gillon and Yakir, 2001). The isotope exchange in the atmosphere is negligible due to lower liquid water content, lower residence time and the absence of carbonic anhydrase (Mills and Urey, 1940; Miller et al., 1971; Johnson, 1982; Silverman, 1982; Francey and Tans, 1987).

$\Delta^{17}O$ of CO_2 has been suggested as additional independent tracer for constraining global GPP (Hoag et al., 2005; Thiemens et al., 2013; Hofmann et al., 2017; Liang et al., 2017b; Koren et al., 2019), because the processes involved in plant-atmosphere gas exchange are all mass-dependent. Therefore, $\Delta^{17}O$ at the CO_2-H_2O exchange site in the leaf will vary much less than $\delta^{18}O$. Nevertheless, mass-dependent isotope fractionation processes with slightly different three-isotope fractionation slopes are involved, which have been precisely established in the past years. Figure 1 shows how the different processes affect $\Delta^{17}O$ of the H_2O and CO_2 reservoirs involved. The triple isotope slope of oxygen in meteoric waters is taken as reference slope, $\lambda_{Ref}=0.528$ (Meijer and Li, 1998; Barkan and Luz, 2007; Landais et al., 2008; Luz and Barkan, 2010; Uemura et al., 2010) and we assume that soil water is similar to meteoric water. Due to transpiration and diffusion in the leaf, $\Delta^{17}O$ of leaf water gets modified following a humidity dependent three-isotope slope $\theta_{\mathrm{trans}}=0.522-0.008\times h$ (Landais et al., 2006). Exchange of oxygen isotopes between leaf water and CO_2 follows $\lambda_{CO2\text{-}H2O}=0.5229$ (Barkan and Luz, 2012) which determines the $\Delta^{17}O$ of CO_2 inside the leaf at the CO_2-H_2O exchange site. Finally, the $\Delta^{17}O$ of the CO_2 is modified when CO_2 diffuses into and out of the leaf with $\lambda_{diff}=0.509$ (Young et al., 2002).

In the first box model study of Hoag et al. (2005), the small deviations in $\Delta^{17}O$ of CO_2 due to differences in three-isotope slopes were neglected and exchange with water was assumed to reset $\Delta^{17}O$ to 0. Hofmann et al. (2017) included the different isotope effects shown in Figure 1 in their box model. Koren et al. (2019) incorporated all the physico-chemical processes affecting $\Delta^{17}O$ of CO_2 in a 3D atmospheric model and investigated the spatiotemporal variability of $\Delta^{17}O$ and its use as tracer for GPP. Using these and other similar models, numerous measurements of $\Delta^{17}O$ in atmospheric CO_2 from different locations have been performed and used to estimate GPP (Liang et al., 2006; Barkan and Luz, 2012; Thiemens et al., 2014; Liang and Mahata, 2015; Laskar et al., 2016; Hofmann et al., 2017). The three-isotope slopes of the processes involved in the gas exchange (Figure 1) have been precisely determined in idealized experiments. In the advanced models mentioned above it is assumed that when all the pieces are put together they results in a realistic overall modification of $\Delta^{17}O$ of CO_2 in the atmosphere surrounding the leaf. However, this has not been confirmed by measurements previously.

In this study we report the effect of photosynthesis on $\Delta^{17}O$ of CO_2 in the surrounding air at the leaf scale. We measured $\Delta^{17}O$ of CO_2 entering and leaving a leaf cuvette to calculate the isotopic fractionation associated with photosynthesis for three species that are representative for three different biomes. The fast-growing annual herbaceous C_3 species *Helianthus annuus* (sunflower) has a high photosynthetic capacity (A_N) and high stomatal conductance (g_s) and is representative for temperate and tropical crops (Fredeen et al., 1991). The slower growing perennial evergreen C_3 species *Hedera hybernica* (ivy) is representative of forests and other woody vegetation and stress subjected habitats (Pons et al., 2009). The fast-growing, agronomically important crop *Zea mays* (maize) is an herbaceous annual C_4 species with a high A_N and a low g_s, typical for savanna type vegetation (Weijde et al., 2013). The mole fraction of CO_2 at the CO_2-H_2O exchange site (c_m) is an important parameter to determine the effect of photosynthesis on $\Delta^{17}O$ of CO_2. In C_3 plants, the CO_2-H_2O exchange can occur anywhere between the plasma membrane and the chloroplast since the catalyzing enzyme CA has been found in the chloroplast, cytosol, mitochondria and plasma membrane (Fabre et al., 2007; DiMario et al., 2016). For C_4 plants, CA is mainly found in the cytosol and the CO_2-H_2O exchange occurs there (Badger and Price, 1994). In our experiments, sunflower and ivy are used to cover the wide c_m/c_a ratio range among C_3 plants and maize represents the c_m/c_a ratio for the C_4 plants. Using our results from the leaf scale experiments, we estimated the effect of terrestrial vegetation on $\Delta^{17}O$ of CO_2 in the global atmosphere.

1. Theory

1.1. Notation and definition of δ values

Isotopic composition is expressed as the deviation of the heavy to light isotope ratio in a sample relative to a reference ratio and it is denoted as δ, expressed in per mill (‰). In the case of oxygen isotopes, the isotope ratios are $^{18}R = [^{18}O]/[^{16}O]$ and $^{17}R = [^{17}O]/[^{16}O]$ and the reference material is Vienna Standard Mean Ocean Water (VSMOW):

$$\delta^{n}O = \frac{^{n}R_{sample}}{^{n}R_{VSMOW}} - 1,\ n \text{ refers to 17 or 18} \tag{1}$$

For most processes, isotope fractionation depends on mass, and therefore the fractionation against ^{17}O is approximately half of the fractionation against ^{18}O (equation 3).

$$\ln(\delta^{17}O + 1) = \lambda \times \ln(\delta^{18}O + 1) \tag{2}$$

The mass-dependent isotope fractionation factor λ ranges from 0.5 to 0.5305 for different molecules and process (Matsuhisa et al., 1978; Thiemens, 1999; Young et al., 2002; Cao and Liu, 2011). $\Delta^{17}O$ is used to quantify the degree of deviation from equation (2) (see equation 3). Note that $\Delta^{17}O$ changes not only by mass-independent isotope fractionation processes, but also by mass-dependent isotope fractionation processes with a different λ value from the one used in the definition of $\Delta^{17}O$ (Barkan and Luz, 2005; Landais et al., 2006; 2008; Luz and Barkan, 2010; Barkan and Luz, 2011; Pack and Herwartz, 2014).

$$\Delta^{17}O = \ln(\delta^{17}O + 1) - \lambda \times \ln(\delta^{18}O + 1) \quad (3)$$

The choice of λ is in principle arbitrary and in this study, we use $\lambda = 0.528$, which was established for meteoric waters (Meijer and Li, 1998; Landais et al., 2008; Brand et al., 2010; Luz and Barkan, 2010; Barkan and Luz, 2012; Sharp et al., 2018). Equation 3 can be linearized to (Miller, 2002) $\Delta^{17}O = \delta^{17}O - \lambda \times \delta^{18}O$, but this approximation causes an error that increases with $\delta^{18}O$ (Miller, 2002; Bao et al., 2016).

1.2. Discrimination against $\Delta^{17}O$ of CO_2

The overall isotope fractionation associated with the photosynthesis of CO_2 is commonly quantified using the term discrimination as described in (Farquhar and Richards, 1984; Farquhar et al., 1989; Farquhar and Lloyd, 1993). We use the symbol Δ_A for discrimination due to assimilation in this manuscript since the commonly used Δ is already used for the definition of $\Delta^{17}O$ (see equation 3). Δ_A quantifies the enrichment or depletion of carbon and oxygen isotopes of CO_2 in the surrounding atmosphere relative to the CO_2 that is assimilated (Farquhar and Richards, 1984). It can be calculated from the isotopic composition of the CO_2 entering and leaving the leaf cuvette (Evans et al., 1986; Gillon and Yakir, 2000a; Barbour et al., 2016) as:

$$\Delta_A{}^{n}O_{obs} = \frac{{}^{n}R_a}{{}^{n}R_A} - 1 = \frac{\delta^{n}O_a - \delta^{n}O_A}{1+\delta^{n}O_A} = \frac{\zeta\times(\delta^{n}O_a - \delta^{n}O_e)}{1+\delta^{n}O_a - \zeta\times(\delta^{n}O_a - \delta^{n}O_e)} \quad (4)$$

where the indices *e*, *a* and *A* refer to CO_2 entering and leaving the cuvette and being assimilated, respectively. $\zeta = \frac{c_e}{c_e - c_a}$, where c_e and c_a are the mole fractions of CO_2 entering and leaving the cuvette. For quantifying the effect of photosynthesis on $\Delta^{17}O$ in our experiments, the $\Delta_A\Delta^{17}O$ is calculated from $\Delta_A{}^{17}O$ and $\Delta_A{}^{18}O$ using the three-isotope slope $\lambda_{RL} = 0.528$, similar to equation 3. In previous studies slightly, different formulations have been used to define the effect of photosynthesis on $\Delta^{17}O$, and a comparison of the different definitions is provided in the supplementary material (equation S37-S40).

It is important to note that when the logarithmic definition of $\Delta^{17}O$ or $\Delta_A\Delta^{17}O$ is used, values are not additive (Kaiser et al., 2004). In linear calculations, the error gets larger when the relative difference in $\delta^{18}O$ between the two CO_2 gases increases regardless of the $\Delta^{17}O$ of the individual CO_2 gases (Figure S1). Therefore, $\Delta_A\Delta^{17}O$ values have to be calculated from the individual $\Delta_A{}^{17}O$ and $\Delta_A{}^{18}O$ values, and not by linear combinations of the $\Delta^{17}O$ of air entering and leaving a plant chamber.

2. Materials and methods

2.1. Plant material and growth conditions

Sunflower (*Helianthus annuu*s L. cv "sunny") was grown from seeds in 0.6 L pots with potting soil (Primasta, the Netherlands) for about four weeks. All leaves appearing above the first leaf pair were removed to avoid shading. Established juvenile ivy (*Hedera hybernica* L.) plants were pruned and planted in 6 L pots for 6 weeks. Ivy leaves that had developed and matured were used for the experiments. Maize (*Z. mays L*. cv "saccharate") was grown from seed in 1.6 L pots for at least 7 weeks. For maize, the 4th or higher leaf number was used for the experiments when mature. A section of the leaf at about 1/3 from the tip was inserted in the leaf cuvette. They were placed on a sub-irrigation system that provided water during the growth period in a controlled environment growth chamber, air temperature 20ºC, relative humidity 70 % and CO_2 mole fraction of about 400 ppm. The photosynthetic photon flux density (PPFD) was about 300 µmol m^{-2} s^{-1} during a daily photoperiod of 16 hours measured with a PPFD meter (Licor LI-250A, Li-Cor Inc, Nebraska, USA).

2.2. Gas exchange experiments

Gas exchange experiments were performed in an open system where a controlled flow of air enters and leaves the leaf cuvette similar to the setup used by (Pons and Welschen, 2002). A schematic for the gas exchange experimental setup is shown in Figure 2. The leaf cuvette had dimensions of 7 x 7 x 7 cm^3 (lxwxh) and the top part of the cuvette was transparent. The temperature of the leaf was measured with a K type thermocouple. The leaf chamber temperature was controlled by a temperature-controlled water bath kept at 20ºC (Tamson TLC 3, The Netherlands). A halogen lamp (PRADOVIT 253, ERNST LEITZ WETZLAR GMBH, Germany) in a slide projector was used as a light source. Infrared was excluded by reflection from a cold mirror. The light intensity was varied with spectrally neutral filters (PRADOVIT 253, ERNST LEITZ WETZLAR GMBH, Germany).

The CO_2 mole fraction of the incoming and outgoing air was measured with an infrared gas analyzer (IRGA, model LI-6262, LI-COR Inc., Nebraska, USA). The isotopic composition and mole fraction of the incoming and outgoing water vapor were measured with a triple water vapor isotope analyzer (WVIA, model 911-0034, Los Gatos Research, USA). Compressed air (ambient outside air without drying) was passed through soda lime to scrub the CO_2. The CO_2 free air could be humidified depending on the experiment conditions (see Figure 2). The humidity of the inlet air was monitored continuously with a dewpoint meter (HYGRO-M1, General Eastern, Watertown, MA, USA). Pure CO_2 (either normal CO_2 or isotopically enriched CO_2) was mixed with the incoming air to produce a CO_2 mole fraction of 500 ppm. The isotopically enriched CO_2 was prepared by photochemical isotope exchange between CO_2 and O_2 under UV irradiation (Adnew et al., 2019).

An attached leaf or part of it was inserted into the cuvette, the composition of the inlet air was measured, and both IRGA and WVIA were switched to measure the outlet air. Based on the CO_2 mole fraction of the outgoing air the flow rate of the incoming air to the cuvette was adjusted to establish a drawdown of 100 ppm CO_2 due to photosynthesis in the plant chamber. The water vapor content entering the cuvette was adjusted depending on the transpiration rate relative to CO_2 uptake to avoid condensation (Figure 2). The outgoing air was measured continuously until a steady state was reached for CO_2 and H_2O mole fractions and δD and $\delta^{18}O$ of the water vapor. After a steady state was established, the air was directed to the sampling flask while the IGRA and WVIA were switched back to measure the inlet air. The air passed through a $Mg(ClO_4)_2$ dryer before entering the sampling flask.

After sampling, the leaf area inside the cuvette was measured with a LI-3100C area meter (Li-COR, Inc. USA). Immediately afterward, the leaf was placed in a leak tight 9 mL glass vial and kept in a freezer at -20ºC until leaf water extraction.

2.3. Calibration of the Water Vapor Isotope Analyzer (WVIA) and leaf water analysis

The WVIA was calibrated using five water standards provided by IAEA (Wassenaar et al., 2018) for both $\delta^{18}O$ and δD (Figure S2). We did not calibrate the WVIA for $\delta^{17}O$, so the $\delta^{17}O$ data are not used in the quantitative evaluation. The isotopic composition of the water standards ranged from -50.93 to 3.64‰ and -396.98 ‰ to 25.44 ‰ for δD and $\delta^{18}O$, respectively. The detailed characterization and calibration of the water WVIA is provided in the supplementary material (Figure S2 to S4).

Leaf water was extracted by cryogenic vacuum distillation for 4 h at 60ºC following a well-established procedure as shown in Figure S5 (Wang and Yakir, 2000; Landais et al., 2006; West et al., 2006). Details are provided in the supplementary material. The $\delta^{17}O$ and $\delta^{18}O$ of leaf water was determined at the Laboratoire des Sciences du Climat et de l'Environnement laboratory using a fluorination technique as described in (Barkan and Luz, 2005;Landais et al., 2006;2008).

2.4. Carbon dioxide extraction and isotope analysis

CO_2 was extracted from the air samples in a system made from electropolished stainless steel (Supplementary Figure S6). Our system used four commercial traps (MassTech, Bremen, Germany). The first two traps were operated at dry ice temperature (-78ºC) to remove moisture and some organics. The other two traps were operated at liquid nitrogen temperature (-196ºC) to trap CO_2. The flow rate during extraction was 55 mL min^{-1}, controlled by a mass flow controller (Brooks Instruments, Holland). The reproducibility of the extraction system was 0.030 ‰ for $\delta^{18}O$ and 0.007 ‰ for $\delta^{13}C$ determined on 14 extractions (1σ standard deviation, Supplementary Table S1).

The $\Delta^{17}O$ of CO_2 was determined using the CO_2-O_2 exchange method (Mahata et al., 2013; Barkan et al., 2015; Adnew et al., 2019). The CO_2-O_2 exchange system used at Utrecht University is described in (Adnew et al., 2019). In short, equal amounts of CO_2 and O_2 were mixed in a quartz reactor containing a platinum sponge catalyst and heated at 750ºC for 2hrs. After isotope equilibration, the CO_2 was trapped at liquid nitrogen temperature, while the O_2 was collected with 1 pellet of 5Å molecular sieve (1.6 mm, Sigma Aldrich, USA) at liquid nitrogen temperature. The isotopic composition of the isotopically equilibrated O_2 was measured with a Delta^{Plus}XL isotope ratio mass spectrometer in dual inlet mode with reference to a pure O_2 calibration gas that has been assigned values of $\delta^{17}O = 9.254$ ‰ and $\delta^{18}O = 18.542$ ‰ by E. Barkan at the Hebrew University of Jerusalem. The reproducibility of the $\Delta^{17}O$ measurement was better than 0.01 ‰ (Supplementary Table S1).

2.5. Leaf cuvette model

We used a simple leaf cuvette model to evaluate the dependence of $\Delta_A\Delta^{17}O$ on key parameters. In this model, the leaf is partitioned into three different compartments: the intercellular air space, the mesophyll cell, and the chloroplast. In the leaf cuvette model, we used a 100 ppm downdraw of CO_2, similar to the leaf exchange experiments, i.e., the CO_2 mole fraction decreases from 500 ppm in the entering air (c_e) to 400 ppm in the outgoing air (c_o), which is identical to the air surrounding the leaf (c_a) as a result of thorough mixing in the cuvette. The assimilation rate is set to 20.0 µmol $m^{-2}s^{-1}$. The leaf area and flowrate of air are set to 30 cm^2 and 0.7 L min^{-1}, respectively. The isotope composition of leaf water at the site where the H_2O-CO_2 exchange occurs is $\delta^{17}O$ = 5.39 ‰ and $\delta^{18}O$ = 10.648 ‰, which is the mean of the measured $\delta^{17}O$ and $\delta^{18}O$ values of bulk leaf water in our experiments. The leaf water temperature is set to 22°C (similar to the experiment). In the model, the $\delta^{18}O$ of the CO_2 entering the cuvette is set to 30.47 ‰ for all the simulations, as in the normal CO_2 experiments, but the assigned $\Delta^{17}O$ values ranges from -0.5 ‰ to 0.5 ‰ which encompasses both the stratospheric intrusion and combustion components. The corresponding $\delta^{17}O$ of the CO_2 entering the cuvette is calculated from the assigned $\delta^{18}O$ value (30.47 ‰) and $\Delta^{17}O$ values (-0.5 ‰ to 0.5 ‰). For the calculations with this model, we assumed an infinite boundary layer conductance. The leaf cuvette model is illustrated in the supplementary material (Figure S7) and the detailed code and description is available at https://git.wur.nl/leaf_model/D17O.

3. Results

3.1. Gas exchange parameters

Table 1 summarizes the isotopic composition and mole fraction of the CO_2 used in this study for sunflower, ivy and maize. The $\Delta^{17}O$ of CO_2 used in this study varies from -0.215 ‰ to 0.44 ‰ while the $\delta^{18}O$ value is close to 30 ‰ for all the experiments. For all the experiments, the mole fraction of CO_2 entering the leaf (c_a) is 400 ppm whereas the mole fraction of the CO_2 in the intercellular air space (c_i), at the CO_2-H_2O exchange site (c_m) and in the chloroplast (c_c) varies depending on the assimilation rate and metabolism type of the plants. Estimating the mesophyll conductance is described in the companion paper. A detailed description for estimating c_m and c_c is provided in the supplementary material. A list of variables and parameters used in this study are summarized in table 3.

3.2. Discrimination against ^{18}O of CO_2

Figure 3a shows discrimination against ^{18}O associated with photosynthesis ($\Delta_A{}^{18}O$) for sunflower, ivy, and maize as a function of the c_m/c_a ratio. $\Delta_A{}^{18}O$ varies with c_m/c_a, as found in previous studies (Gillon and Yakir, 2000a;Barbour et al., 2016) . For sunflower, we observe $\Delta_A{}^{18}O$ values between 29 ‰ and 64 ‰ for c_m/c_a between 0.54 and 0.86. Ivy shows a relatively little variation of $\Delta_A{}^{18}O$ around a mean of 22 ‰ for c_m/c_a between 0.48 and 0.58. For maize, $\Delta_A{}^{18}O$ is lower than for the C_3 plants measured in this study, with values between 10 ‰ and 20 ‰ for c_m/c_a between 0.15 and 0.37.

For sunflower changing the irradiance from 300 µmol $m^{-2}s^{-1}$ (low light, hereafter LL) to 1200 µmol $m^{-2}s^{-1}$ (high light, hereafter HL) leads to a clear decrease in $\Delta_A{}^{18}O$ (average 22 ‰). For maize, the $\Delta_A{}^{18}O$ change is only 4.4 ‰ on average. For ivy, changing the light intensity does not significantly change the observed $\Delta_A{}^{18}O$. The solid lines in Figure 3a show results of leaf cuvette model calculations, where the dependence of $\Delta_A{}^{18}O$ on c_m/c_a is explored for a set of calculations with otherwise fixed parameters. The model agrees well with the experimental results except for ivy, where the model overestimates the discrimination.

3.3. Discrimination against $\Delta^{17}O$ of CO_2

The discrimination of photosynthesis against $\Delta^{17}O$ of CO_2 ($\Delta_A\Delta^{17}O$) is shown in Figure 3b. $\Delta_A\Delta^{17}O$ is negative for all experiments and it depends strongly on the c_m/c_a ratio and $|\Delta_A\Delta^{17}O|$ increases with c_m/c_a ratio. For instance, for $\Delta^{17}O$ of CO_2 entering the cuvette of -0.215 ‰, $\Delta_A\Delta^{17}O$ is -0.25 ‰ for maize with c_m/c_a ratio of 0.3, -0.3 ‰ for ivy with c_m/c_a ratio of 0.5 ‰ and -0.5 ‰ for sunflower with c_m/c_a ratio of 0.7 (Figure 3b). For sunflower and ivy, $\Delta_A\Delta^{17}O$ is also strongly dependent on the $\Delta^{17}O$ of CO_2 supplied to the cuvette, whereas no significant dependence is found for maize. For an increase in $\Delta^{17}O$ of CO_2 entering the cuvette from -0.215 ‰ to 0.435 ‰, $\Delta_A\Delta^{17}O$ increases from -0.3 ‰ to -0.9 ‰ at c_m/c_a ratio of 0.5 for ivy. For sunflower, an increases $\Delta^{17}O$ of CO_2 entering the cuvette from -0.215 ‰ to 0.31 ‰ increases $\Delta_A\Delta^{17}O$ from -0.8 ‰ to -1.7 ‰ at c_m/c_a ratio of 0.8. The leaf cuvette model results illustrate the shape of the dependence on the c_m/c_a ratio and agree well with the experiments. For the leaf cuvette model, the $\Delta^{17}O$ value of the water is assigned a constant value of -0.122 ‰ (average $\Delta^{17}O$ value for the bulk leaf water).

Figure 4b shows the same values of $\Delta_A\Delta^{17}O$ as a function of the difference between $\Delta^{17}O$ of CO_2 entering the leaf and the calculated $\Delta^{17}O$ of leaf water at the evaporation site where CO_2-H_2O exchange takes place ($\Delta^{17}O_a$ - $\Delta^{17}O_{wes}$), for different c_m/c_a ratios. The leaf cuvette model results (solid lines in Figure 4b) suggest a linear dependence between $\Delta_A\Delta^{17}O$ and ($\Delta^{17}O_a$ - $\Delta^{17}O_{wes}$). The experimental results agree with the hypothesis that $\Delta_A\Delta^{17}O$ is linearly dependent on $\Delta^{17}O_a$ - $\Delta^{17}O_{wes}$ at a certain c_m/c_a ratio. Figure 4a shows the corresponding relation where $\Delta_A\Delta^{17}O$ is divided by $\Delta^{17}O_a$-$\Delta^{17}O_m$. All the values follow the same relationship as function of the c_m/c_a ratio, which can be approximated quite well by an exponential function (equation 5). This function quantifies the dependence of $\Delta_A\Delta^{17}O$ on c_m/c_a, and thus the effect of the diffusion of isotopically exchanged CO_2 back to the atmosphere, which increases with increasing c_m/c_a ratio.

$$\frac{\Delta_A\Delta^{17}O}{\Delta^{17}O_a - \Delta^{17}O_m} = -0.150 \times \exp(3.707 \times c_m/c_a) + 0.028 \qquad (5)$$

Figure 5 a and c show results from the leaf cuvette model that illustrates in more detail how $\Delta^{17}O_e$ and $\Delta^{17}O_{wes}$ affect $\Delta^{17}O_a$ and $\Delta_A\Delta^{17}O$ and their dependence on c_m/c_a. At lower c_m/c_a, only a very small fraction of CO_2 that has undergone isotopic equilibration in the mesophyll diffuses back to the atmosphere, and therefore $\Delta^{17}O_a$ stays close to the incoming $\Delta^{17}O_e$, modified by the fractionation during CO_2 diffusion through the stomata (Figure 5a). Figure 5c confirms that indeed at low c_m/c_a, $\Delta_A\Delta^{17}O$ approaches the fractionation constant expected for diffusion, -0.170 ‰. This diffusional fractionation is independent of the isotopic composition of the CO_2 entering the leaf, and therefore at low c_m/c_a, the $\Delta_A\Delta^{17}O$ curves for the different values of the anomaly of the CO_2 entering the leaf converge. For a high c_m/c_a ratio, the back-diffusion flux of CO_2 that has equilibrated with water becomes the dominant factor, and in this case, the isotopic composition of the outgoing CO_2 converges towards this isotope value, independent of the isotopic composition of the incoming CO_2 (Figure 5a). This can lead to a very wide range of values for the discrimination against $\Delta^{17}O$, because now the effect on $\Delta^{17}O$ of the ambient CO_2 depends strongly on the difference in isotopic composition between incoming CO_2 and CO_2 in isotopic equilibrium with the leaf water.

In the model calculations shown in Figure 5b and d, the isotopic composition of the water was changed from $\Delta^{17}O_{wes}$ = -0.122‰ to 0.300 ‰, whereas all other parameters were kept the same. The value of $\Delta^{17}O_e$ for which $\Delta^{17}O_a$ does not depend on c_m/c_a is shifted accordingly, again being similar to $\Delta^{17}O_m$. At low c_m/c_a $\Delta_A\Delta^{17}O$ converges to the same value as in Figure 5c, confirming the role of diffusion into the stomata as discussed above.

Figure 6 shows how $\delta^{18}O$ and $\Delta^{17}O$ vary in key compartments of the leaf cuvette system that determine the oxygen isotope effects associated with photosynthesis, based on the previously established three-isotope slopes of the various processes (Figure 1). The irrigation water has a $\Delta^{17}O$ value of 0.017. The measured bulk leaf water is 6-16 ‰ enriched in ^{18}O and its $\Delta^{17}O$ value is lower by -0.075 to -0.200 ‰ (mean value -0.121 ‰) than the irrigation water, calculated using a three-isotope slope of λ_{trans} = 0.516 at 80 % humidity (Landais et al., 2006). $\Delta^{17}O$ of leaf water at the evaporation site, calculated from the transpired water, has slightly lower $\Delta^{17}O$, with values between -0.119 ‰ and -0.237 (average -0.184 ‰). Note that the bulk leaf water was not measured for all the experiments. For the experiments where the bulk leaf water is measured, $\Delta^{17}O$ of leaf water at the evaporation site ranges from -0.160 ‰ to -0.231 with an average value of -0.190 ± 0.020 ‰. The calculated isotopic composition of water at the exchange site was thus similar, but slightly lower in $\Delta^{17}O$ than the values measured for bulk leaf water. CO_2 exchanges with the water in the leaf with a well-established fractionation constant (see equation S17, supplementary material) and a three-isotope slope of $\lambda_{CO2\text{-}H2O}$ = 0.5229 (Barkan and Luz, 2012), leading to the lower $\Delta^{17}O$ values of the equilibrated CO_2. In our experiments, the $\Delta^{17}O$ value of CO_2 in equilibrium with leaf water is lower than the $\Delta^{17}O$ value of CO_2 entering the leaf. The $\Delta^{17}O$ of the CO_2 in the intercellular air space is a mixture between two end members, the $\Delta^{17}O$ of the CO_2 entering the leaf and $\Delta^{17}O$ of the CO_2 in equilibrium with leaf water. This explains why the observed values of $\Delta_A\Delta^{17}O$ are negative for the experiments performed in this study.

4. Discussion

4.1. Discrimination against ^{18}O of CO_2

The higher $\Delta_A{}^{18}O_{obs}$ values for sunflower compared to maize and ivy (Figure 3a) are mainly due to a higher back-diffusion flux ($c_m/(c_a-c_m)$). The back-diffusion flux is higher for the C_3 plants sunflower and ivy than for the C_4 plant maize, a consequence of the lower stomatal conductance and higher assimilation rate of C_4 plants (Gillon and Yakir, 2000a; Barbour et al., 2016). In C_4 plants most of the CO_2 entering the stomata is carboxylated by PEPC resulting in a lower CO_2 mixing ratio in the mesophyll which results in a lower back-diffusion flux. The increase of assimilation rate with higher light intensity decreases the c_m/c_a ratio and thus leads to a lower back-diffusion flux, which explains the decreases of $\Delta_A{}^{18}O_{obs}$ for maize and most clearly for sunflower. A similar trend of increase in $\Delta_A{}^{18}O_{obs}$ with an increase in c_m/c_a ratio has been reported in previous studies (Gillon and Yakir, 2000b, a; Osborn et al., 2017). For ivy, $\Delta_A{}^{18}O_{obs}$ and $\Delta_A{}^{17}O_{obs}$ do not decrease with an increase in irradiance, because the change in assimilation rate with irradiance is small. Thus, c_m will not decrease strongly and the effect on the back diffusion is smaller than the variability in $\Delta_A{}^{18}O_{obs}$ of different leaves of the same plant.

In our experiments, photosynthesis causes enrichment $\delta^{18}O$ of atmospheric CO_2 for both C_3 and C_4 plants, i.e. positive values $\Delta_A{}^{18}O$. In principle, $\Delta_A{}^{18}O$ can also be negative if the $\delta^{18}O_m$ are depleted relative to the ambient CO_2. This is in contrast to $\Delta_A{}^{13}C$, which will always be positive since it is determined by the fractionation due to the PEPC and RuBisCO enzyme activity (Figure S8 and S9, supplementary material). In general, in our experiments, the $\Delta_A{}^{18}O_{obs}$ values are about five times larger than $\delta^{18}O_a - \delta^{18}O_e$, the $\delta^{18}O$ difference between CO_2 entering and leaving the cuvette (Figure S10 to S12 supplementary material). This is easy to understand from the definition of Δ_A. Taking $\Delta_A{}^{18}O$ as an example, $\Delta_A{}^{18}O_{obs} = \frac{\zeta(\delta^{18}O_a-\delta^{18}O_e)}{1+\delta^{18}O_a-\zeta(\delta^{18}O_a-\delta^{18}O_e)} \approx \zeta(\delta^{18}O_a - \delta^{18}O_e)$ and in our experiments, $\zeta = c_e / (c_e-c_a) \approx 500 / (500-400) = 5$.

4.2. Discrimination against the $\Delta^{17}O$ of CO_2

The leaf cuvette model includes the isotope fractionations of all the individual processes that have been quantified in dedicated experiments previously (Figure 1). The good agreement of the model results with the measurements (Figure 3a) demonstrates that when all these processes are combined in the quantitative description of a gas exchange experiment, they actually result in a correct quantification of the isotope effects associated with photosynthesis. This has already been demonstrated before for $\Delta_A{}^{18}O_{obs}$ but has now been confirmed for $\Delta_A\Delta^{17}O$.

Unlike ivy and sunflower, maize does not show a significant change in $\Delta_A\Delta^{17}O$ when CO_2 gases with different $\Delta^{17}O$ are supplied to the plant. The C_4 plant maize has a small back-diffusion flux due to its high assimilation rate and low stomatal conductance, leading to a low c_m/c_a ratio. At low c_m/c_a ratios, $\Delta_A\Delta^{17}O$ is expected to be close to the weighted fractionation due to diffusion through boundary layer and stomata. In general, the effect of diffusion on $\Delta^{17}O$ of atmospheric CO_2 can be expressed as follows:

$$\Delta^{17}O_{\mathrm{Modified}} = \Delta^{17}O_a + \left(\lambda_{\mathrm{RL}} - \lambda_{diffusion}\right) \times \ln\alpha_{diffusion} \tag{6}$$

where $\Delta^{17}O_a$ is the $\Delta^{17}O$ of the CO_2 surrounding the leaf, $\Delta^{17}O_{modified}$ is the $\Delta^{17}O$ of the CO_2 modified due to diffusional fractionation and $\lambda_{diffusion}$, λ_{RL} and $\alpha_{diffusion}$ are the oxygen three-isotope relationships during diffusion from the CO_2-H_2O exchange site to the atmosphere, the reference slope used and the fractionation against ^{18}O for CO_2 during diffusion through the stomata. Using the values $\lambda_{RL} = 0.528$, $\lambda_{diffusion} = 0.509$ (Young et al., 2002) and $\alpha_{diffusion}$=0.9912 (Farquhar and Lloyd, 1993), the effect of diffusional fractionation on the $\Delta^{17}O$ of atmospheric CO_2 is -0.168 ‰ regardless of the anomaly of the CO_2 entering the leaf, and the model results confirm this at low c_m/c_a ratio (Figure 5 c and d, inset).

At a high c_m/c_a ratio, $\Delta^{17}O_a$ is dominated by the back-diffusion flux of CO_2 that has equilibrated with water. As a consequence, $\Delta^{17}O_a$ converges to a common value that is independent of the anomaly of the CO_2 entering the cuvette and is determined by the isotopic composition of leaf water. Figure 5 confirms that the end member is equal to the $\Delta^{17}O$ of CO_2 in equilibrium with leaf water, $\Delta^{17}O_m$. In fact, when $\Delta^{17}O_a = \Delta^{17}O_m$, $\Delta^{17}O_a$ does not change with c_m/c_a, indicating that in this case the $\Delta^{17}O$ of the CO_2 diffusing back from the leaf is the same as the $\Delta^{17}O(CO_2)$ entering the leaf.

$\overline{a}_{18}$ is the overall discrimination occurring during the diffusion of $^{12}C^{18}O^{16}O$ from the ambient air surrounding the leaf to the CO_2-H_2O exchange site (see Table 3 for the list of variables). In our study $\overline{a}_{18}$ ranges from 5 ‰ to 7.2 ‰, lower than the literature estimate of 7.4 ‰ (Farquhar et al., 1993). $\overline{a}_{18}$ depends on the ratio of stomatal conductance, which is associated with a strong fractionation of 8.8 ‰, to mesophyll conductance with an associated fractionation of only 0.8 ‰. Therefore, the higher the ratio (g_s/g_{m18}) the lower the $\overline{a}_{18}$ (Table S2, supplementary material). The difference in $\overline{a}_{18}$ of 2.4 ‰ between the literature value of 7.4 ‰ and the lowest $\overline{a}_{18}$ estimate in this study will introduce an error of only 0.046 ‰ in the $\Delta^{17}O$ value (see equation 6). The uncertainty $\overline{a}_{18}$ has lower influence on the $\Delta_A\Delta^{17}O$ of C_3 plants compared to C_4 plants since the diffusional fractionation is less important at the higher c_m/c_a ratio where C_3 plants operate.

4.3. Global average value of $\Delta_A\Delta^{17}O$ and $\Delta^{17}O$ isoflux

We can use the established relationship between $\Delta_A\Delta^{17}O$ and $\Delta^{17}O_a$ - $\Delta^{17}O_{wes}$ for a certain c_m/c_a ratio to provide a bottom-up estimate for the global effect of photosynthesis on $\Delta^{17}O$ in atmospheric CO_2, based on data obtained in real gas exchange experiments. For this, we use results from a recent modeling study, which provides global average values for CO_2 and leaf water ($\Delta^{17}O(CO_2)$ = -0.168 ‰, $\Delta^{17}O(H_2O_{-leaf})$ = -0.067 ‰; (Koren et al., 2019); Figure S13 and 14, supplementary material). The $\Delta^{17}O(CO_2)$ values agree well with the limited amount of available measurements (Table 1).

To extrapolate $\Delta_A\Delta^{17}O$ determined in the leaf scale experiments to the global scale, global average c_m/c_a ratios of 0.7 and 0.3 are used for C_3 and C_4 plants, respectively, similar to previous studies (Hoag et al., 2005;Liang et al., 2017b). From SIBCASA model results we obtained an annual variability of c_i/c_a values with a standard deviation of 0.12 and 0.17 for C_4 and C_3 plants respectively (Figure S15, supplementary material) (Schaefer et al., 2008; Koren et al., 2019). We use this variability as upper limit of the error estimate for c_m/c_a as shown in the light orange and light pink shaded areas in Figure 4b. This error is converted to an error in $\Delta_A\Delta^{17}O$ using the relation with c_m/c_a. Based on the linear dependency of $\Delta_A\Delta^{17}O$

and $\Delta^{17}O_a$-$\Delta^{17}O_{wes}$, we estimate the $\Delta_A\Delta^{17}O$ for tropospheric CO_2 based on the $\Delta^{17}O$ of leaf water and c_m/c_a ratio. In Figure 4b, the dashed black vertical line indicates $\Delta^{17}O_a$-$\Delta^{17}O_{wes}$ obtained from the 3D global model (Koren et al., 2019). The results of the global estimate and parameters used for the extrapolation of leaf scale study to the global scale are summarized in Table 1.

The $\delta^{17}O$ value of atmospheric CO_2 (21.53 ‰) is calculated from the global $\delta^{18}O$ and $\Delta^{17}O$ values (41.5 ‰ and -0.168 ‰, respectively) (Koren et al., 2019). The $\delta^{17}O$ and $\delta^{18}O$ values of global mean leaf water are calculated from the soil water. A global mean $\delta^{18}O$ value of soil water is -8.4 ‰ assuming soil water to be similar to precipitation (Bowen and Revenaugh, 2003; Koren et al., 2019). The $\delta^{17}O$ value of soil water is -4.4 ‰, calculated using equation 7 (Luz and Barkan, 2010).

$$\ln\left(\delta^{17}O_{soil}+1\right) = 0.528\times\ln\left(\delta^{18}O_{soil}+1\right)+0.033 \qquad (7)$$

$\delta^{17}O$ and $\delta^{18}O$ of leaf water are calculated from $\delta^{17}O$ and $\delta^{18}O$ of soil water with fractionation factors of 1.0043 and 1.0084, respectively (Hofmann et al., 2017; Koren et al., 2019). The fractionation factor for $\delta^{17}O$ is calculated using $\alpha^{17} = (\alpha^{18})^{\lambda_{trans}}$ with λ_{trans} =0.516, assuming relative humidity to be 75 % (Landais et al., 2006). The $\delta^{17}O$ and $\delta^{18}O$ values of global mean leaf water are then -0.136 ‰ and -0.131 ‰, respectively. Thus, the difference between global atmospheric CO_2 and leaf water is $\delta^{17}O_{CO2-water}$ = 21.666 ‰ and $\delta^{18}O_{CO2-water}$ = 41.631 ‰. This yields $\Delta^{17}O_{CO2-water}$ = -0.101 ‰, and this value is indicated as dashed black line in Figure 4. The grey shaded area indicates the propagated error using the standard deviation of the relevant parameters in 180 x 360 grid boxes for 12 months of leaf water and 45 x 60 grid boxes for 24 months for CO_2 (Koren et al., 2019). In Figure 4b, the intersection between the dashed black vertical line and the discrimination lines for the representative c_m/c_a ratios of C_3 and C_4 plants corresponds to the $\Delta_A\Delta^{17}O$ value of C_3 and C_4 plants. For C_4 plants (c_m/c_a = 0.3) this yields $\Delta_A\Delta^{17}O$ = -0.3 ‰ (gray dashed line in figure 4b) and for C_3 plants (c_m/c_a = 0.7), $\Delta_A\Delta^{17}O$ = -0.65 ‰ (black dashed line in Figure 4b).

Three main factors contribute to the uncertainty of the extrapolated $\Delta_A\Delta^{17}O$ value. The first is the measurement error, which contributes 0.25 ‰ (standard error for individual experiments). The second factor is the uncertainty in the difference between $\Delta^{17}O$ of atmospheric CO_2 and leaf water, and we use results from the global model to estimate an error. For $\Delta^{17}O$ of atmospheric CO_2, statistics for all 45 x 60 grid boxes for 24 months (2012-2013) show a range of -0.218 ‰ to -0.151 ‰, with a mean of -0.168 ‰ and a standard deviation of 0.013 ‰ (Figure S13, supplementary material). For $\Delta^{17}O$ of the leaf water statistics for all 180 x 360 grid boxes for 12 months show a range of -0.236 ‰ and -0.027 ‰ (Figure S14, supplementary material). The mean is -0.067 ‰ with a standard deviation of 0.041 ‰. From the combined errors we estimate the error in ($\Delta^{17}O_a$ - $\Delta^{17}O_{wes}$) to be 0.043 ‰. The third uncertainty in the extrapolation of $\Delta^{17}O$ comes from the uncertainty in the c_m/c_a ratio. For C_3 and C_4 plants, these errors are indicated by the light orange and light blue shadings in Figure 4b.

Taking these uncertainties into account leads to a mean value of $\Delta_A\Delta^{17}O$ = -0.3±0.18 ‰ for C_4 plants and $\Delta_A\Delta^{17}O$ = -0.65±0.18 ‰ for C_3 plants. The leaf scale discrimination against $\Delta^{17}O$ is then extrapolated to global vegetation using these representative values of $\Delta_A\Delta^{17}O$ and the relative fractions of photosynthesis by C_4 and C_3 plants, respectively as:

$$\Delta_A\Delta^{17}O_{\mathrm{global}} = f_{C4} \times \Delta_A\Delta^{17}O_{C4} + f_{C3} \times \Delta_A\Delta^{17}O_{C3} \quad (8)$$

where f_{C4} and f_{C3} are the photosynthesis weighted global coverage of C_4 and C_3 vegetation. $\Delta_A\Delta^{17}O_{C4}$ and $\Delta_A\Delta^{17}O_{C3}$ quantify the discrimination against $\Delta^{17}O$ by C_4 and C_3 plants, which are calculated using estimated values of c_m/c_a from a model. Using assimilation weighted fractions of 23 % for C_4 and 77 % for C_3 vegetation (Still et al., 2003), the global mean value of $\Delta_A\Delta^{17}O$ obtained from equation 8 is -0.57±0.14 ‰.

Isoflux is the product of isotope composition and gross mass flux of the molecule. In the case of assimilation, the net flux $F_A = F_{AL} - F_{LA}$ is multiplied with the discrimination associated with assimilation (Ciais et al., 1997a). F_{LA} *and* F_{AL} are total CO_2 fluxes from leaf to the atmosphere and from atmosphere to leaf, respectively. The global scale $\Delta^{17}O_A$ isoflux is calculated by multiplying the discrimination with the assimilation flux as:

$$\mathrm{F_A} \times \Delta_A\Delta^{17}O = \mathrm{A} \times (\mathrm{f_{c4}} \times \Delta_A\Delta^{17}O_{c4} + \mathrm{f_{c3}} \times \Delta_A\Delta^{17}O_{c3}) \quad (9)$$

where, A=0.88×GPP is the terrestrial assimilation rate. The factor 0.88 accounts for the fraction of CO_2 released due to autotrophic respiration (Ciais et al., 1997a). The $\Delta_A\Delta^{17}O$ isoflux due to photosynthesis is calculated using a GPP value of 120 $\mathrm{PgCyr^{-1}}$ (Beer et al., 2010) and A=0.88×GPP, resulting in an isoflux of -60±15 ‰ $\mathrm{PgCyr^{-1}}$ globally. This is the first global estimate of $\Delta_A\Delta^{17}O$ based on direct measurements of the discrimination during assimilation. Our value is in good agreement with previous model estimates. Hofmann et al. (2017) estimated an isoflux ranging from -42 to -92 ‰$\mathrm{PgCyr^{-1}}$ (converted to a reference line with λ=0.528) using an average c_m/c_a ratio of 0.7 for both C_4 and C_3 plants and $\Delta^{17}O$ of -0.147 ‰ for atmospheric CO_2. A model-estimated value from (Hoag et al., 2005) is -47 ‰$\mathrm{PgCyr^{-1}}$ (converted to our reference slope of λ=0.528), derived with a more simple model and using $\Delta^{17}O$ of -0.146 ‰ with c_m/c_a ratio of 0.33 and 0.66 for C_4 and C_3 plants, respectively.

The main uncertainty in the extrapolation of $\Delta_A\Delta^{17}O$ from the leaf experiments to the global scale is the uncertainty in the c_m/c_a ratio. The error from the uncertainty in c_m/c_a ratio increases when the relative difference in $\Delta^{17}O$ between CO_2 and leaf water increases (Figure 5b). It is difficult to determine a single representative c_m value for different plants because this value would need to be properly weighted with temperature, irradiance, CO_2 mole fraction and other environmental factors (Flexas et al., 2008; 2012; Shrestha et al., 2019). Recent developments in laser spectroscopy techniques (McManus et al., 2005; Nelson et al., 2008; Tuzson et al., 2008; Kammer et al., 2011) might enable more and easier measurements of c_m/c_a both in the laboratory and under field conditions. This could lead to a better understanding of variations in the c_m/c_a ratio among plant species and, temporally, spatially and environmentally.

5. Conclusions

In order to directly quantify the effect of photosynthetic gas exchange on the $\Delta^{17}O$ of atmospheric CO_2, gas exchange experiments were carried out in leaf cuvettes using two C_3 plants (sunflower and ivy) and one C_4 plant (maize) with isotopically normal and slightly anomalous (^{17}O-enriched) CO_2. Results for ^{18}O agree with results reported in the literature previously. Our results for $\Delta^{17}O$ confirm that the formalism developed by Farquhar and others for $\delta^{18}O$ is also applicable to the evaluation of $\Delta^{17}O$. In particular, our experiments confirm that two parameters determine the effect of photosynthesis on CO_2: 1) the $\Delta^{17}O$ difference between the incoming CO_2 and CO_2 in equilibrium with leaf water and 2) the c_m/c_a ratio, which determines the degree of back-flux of isotopically exchanged CO_2 from the mesophyll to the atmosphere. At low c_m/c_a ratios, $\Delta_A\Delta^{17}O$ is mainly influenced by the diffusional fractionation. Under our experimental conditions, the isotopic effect increased with c_m/c_a, e.g. $\Delta_A\Delta^{17}O$ was -0.3 ‰ and -0.65 ‰ for maize and sunflower with c_m/c_a ratios of 0.3 and 0.7, respectively. However, experiments with mass independently fractionated CO_2 demonstrate that the results depend strongly on the $\Delta^{17}O$ difference between the incoming CO_2 and CO_2 in equilibrium with leaf water. This is supported by calculations with a leaf cuvette model.

$\delta^{18}O$ is largely affected by kinetic and equilibrium processes between CO_2 and leaf water, and also leaf water isotopic inhomogeneity and dynamics. The $\Delta^{17}O$ variation is much smaller compared to $\delta^{18}O$ and is better defined since conventional bio-geo-chemical processes that modify $\delta^{17}O$ and $\delta^{18}O$ follow a well-defined three-isotope fractionation slope. Results from the leaf exchange experiments were upscaled to the global atmosphere using modeled values for $\Delta^{17}O$ of leaf water and CO_2, which results in $\Delta_A\Delta^{17}O$ = -0.57± 0.14 ‰ and a value for the $\Delta^{17}O$ isoflux of -60± 15 ‰ PgCyr^{-1}. This is the first study that provides such an estimate based on direct leaf chamber measurements, and the results agree with previous $\Delta^{17}O$ calculations. The largest contribution to the uncertainty originates from uncertainty in the c_m/c_a ratio and the largest contributions to the isoflux come from C_3 plants, which have both a higher share of the total assimilation and higher discrimination. $\Delta_A\Delta^{17}O$ is less sensitive to c_m/c_a ratios at lower values of c_m/c_a, for instance for C_4 plants, maize.

$\Delta^{17}O$ of tropospheric CO_2 is controlled by photosynthetic gas exchange, respiration, soil invasion, and stratospheric influx. The stratospheric flux is well established and the effect of photosynthetic gas exchange can now be quantified more precisely. To untangle the contribution of each component to the $\Delta^{17}O$ atmospheric CO_2 we recommend measuring the effects of foliage respiration and soil invasion both in the laboratory and at the ecosystem scale.

Code and data availability.

The data used in this study are included in the paper either with figures or tables. The python code for the cuvette model is available at https://git.wur.nl/leaf_model/D17O.

Author contributions.

GAA and TR designed the main idea of the study. GAA and TP designed the leaf cuvette setup. TP monitors plant growth. GAA and TR designed the CO_2 extraction and CO_2-H_2O exchange system. GAA conducted all the measurements. GK provided the leaf cuvette model. WP enabled the work within the ASICA project. All authors discussed the results at different steps of the project. GAA and TR prepared the manuscript with contributions from all the co-authors.

Competing interests.

The authors declare that they have no conflict of interest

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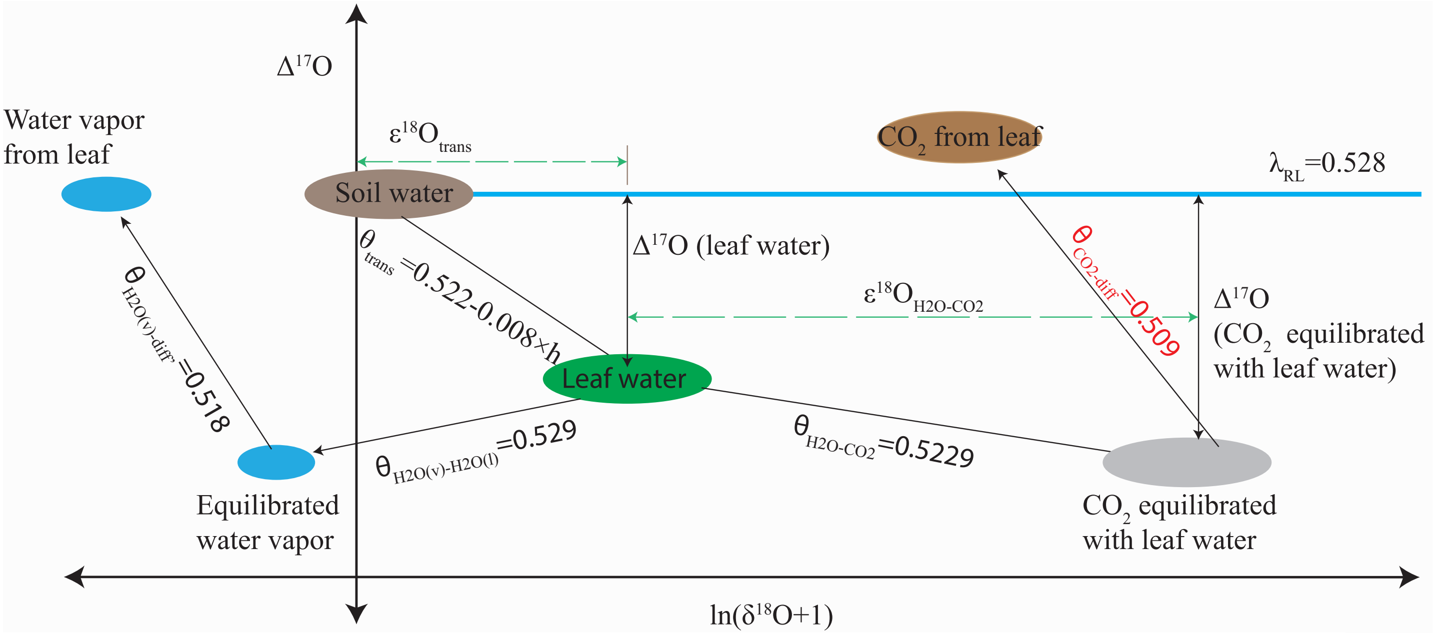

Figure 1 Schematic for mass-dependent isotope fractionation process that affects the Δ^{17}O of the CO_2 and H_2O during the photosynthetic gas exchange (not to scale). The triple oxygen isotope relationship for the individual isotope fractionation processes (both kinetic and equilibrium fractionation) are assigned with θ.

θtrans=0.522-0.008 ×h, where *h* is relative humidity (Landais et al., 2006), in this study the humidity is 75 %, θ_{trans}=0.516. $\theta_{CO2\text{-}H2O}$ (Barkan and Luz, 2012), $\theta_{CO2\text{-}diff}$ (Young et al., 2002), $\theta_{H2O(v)\text{-}H2O(l)}$ (Barkan and Luz, 2005) and $\theta_{H2O(v)\text{-}diff}$ (Barkan and Luz, 2007). Where *v* and *l* for vapor and liquid water, respectively; $\varepsilon^{18}O$ is enrichment or depletion in ^{18}O isotope composition due to the corresponding isotope fractionation process; *diff* and *trans* stand for diffusion and transpiration, respectively.

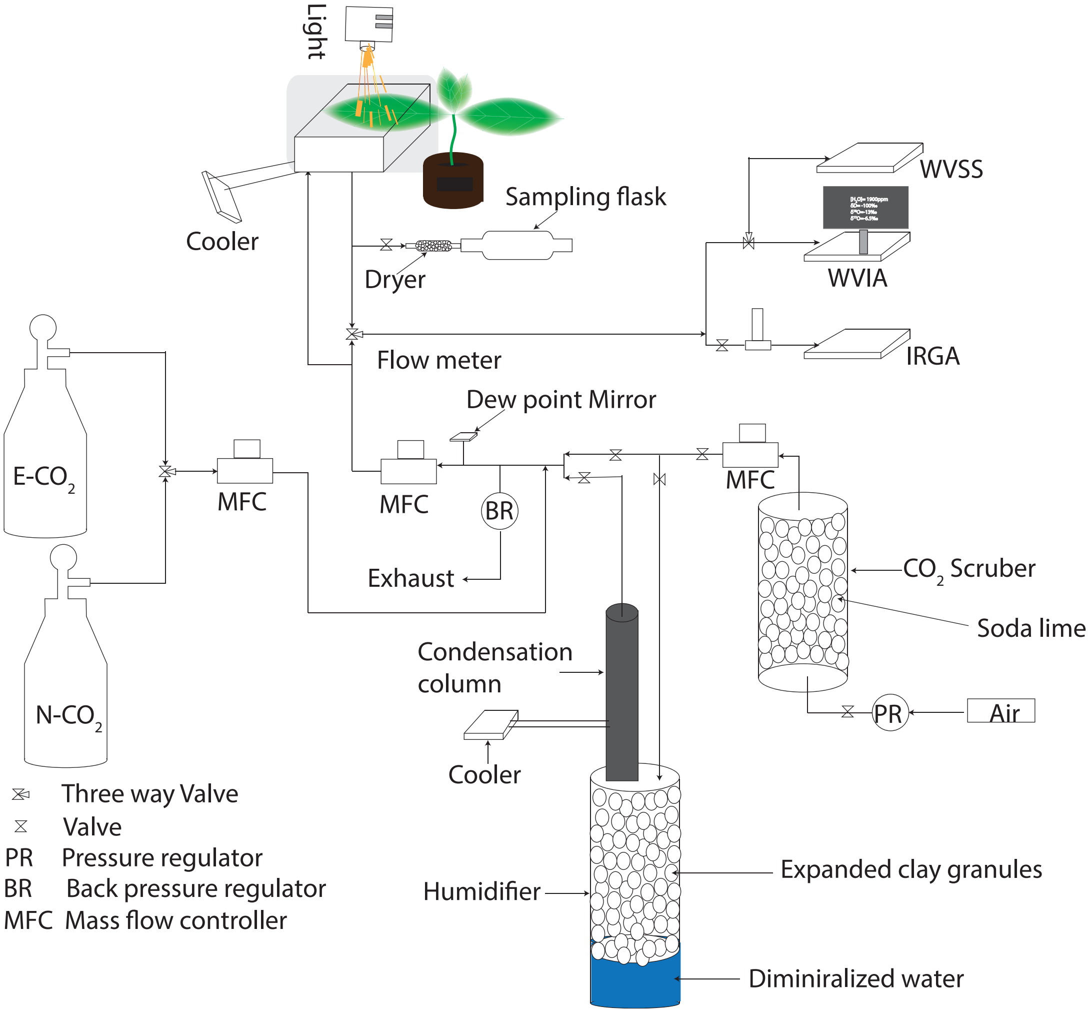

Figure 2 Schematic diagram of the leaf cuvette experimental setup. IRGA stands for the infrared gas analyzer, WVSS is the water vapor standard source, WVIA is the water vapor isotope analyzer, N-CO$_2$ is normal CO$_2$, E-CO$_2$ is ^{17}O-enriched CO$_2$.

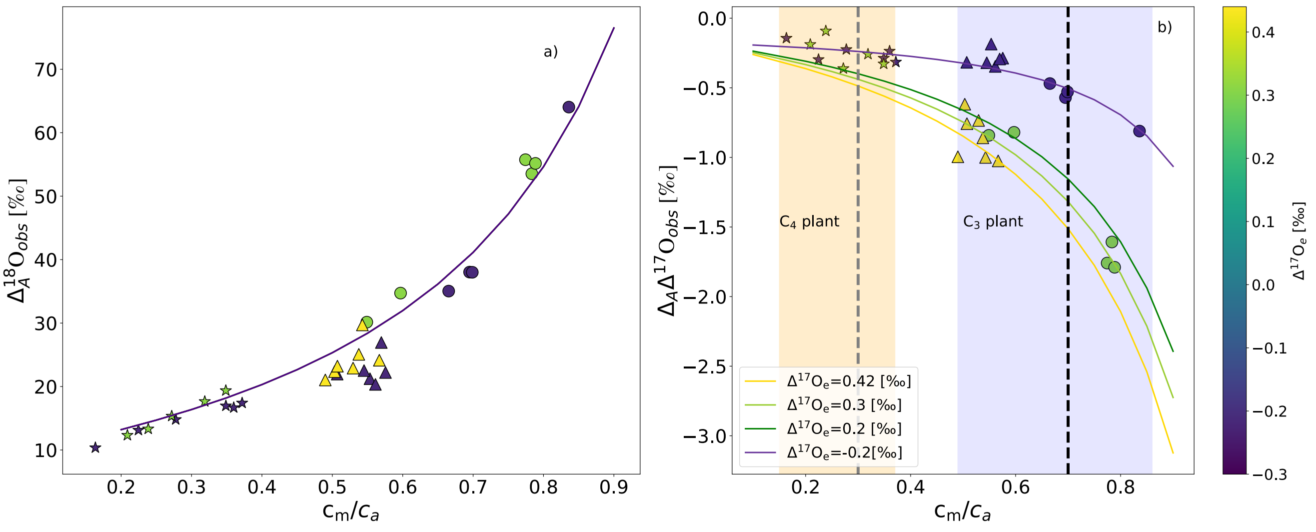

Figure 3 a) $\Delta_A{}^{18}O_{obs}$ during photosynthesis for two C$_3$ plants, sunflower (circles) and ivy (triangles) and C$_4$ plant maize (stars) as a function of c_m/c_a. The solid lines show results from the leaf cuvette model, where $\delta^{18}O$ of the CO$_2$ entering the cuvette is 30.47‰. b) $\Delta_A\Delta^{17}O$ of CO$_2$ as a function of c_m/c_a for isotopically different CO$_2$ gases entering the cuvette (color bar shows $\Delta^{17}O_e$) for sunflower (circles), ivy (triangles) and maize (stars). $\Delta_A\Delta^{17}O$ values calculated using the leaf cuvette model are shown as solid lines in corresponding colors ($\Delta^{17}O_e$ values given in the legend). The shaded areas indicate the c_m/c_a ranges for C$_4$ and C$_3$ plants and the vertical dashed lines indicate the mean c_m/c_a ratio used for extrapolating from the leaf scale to the global scale. Solid line are leaf cuvette model results for the corresponding c_m/c_a ratio.

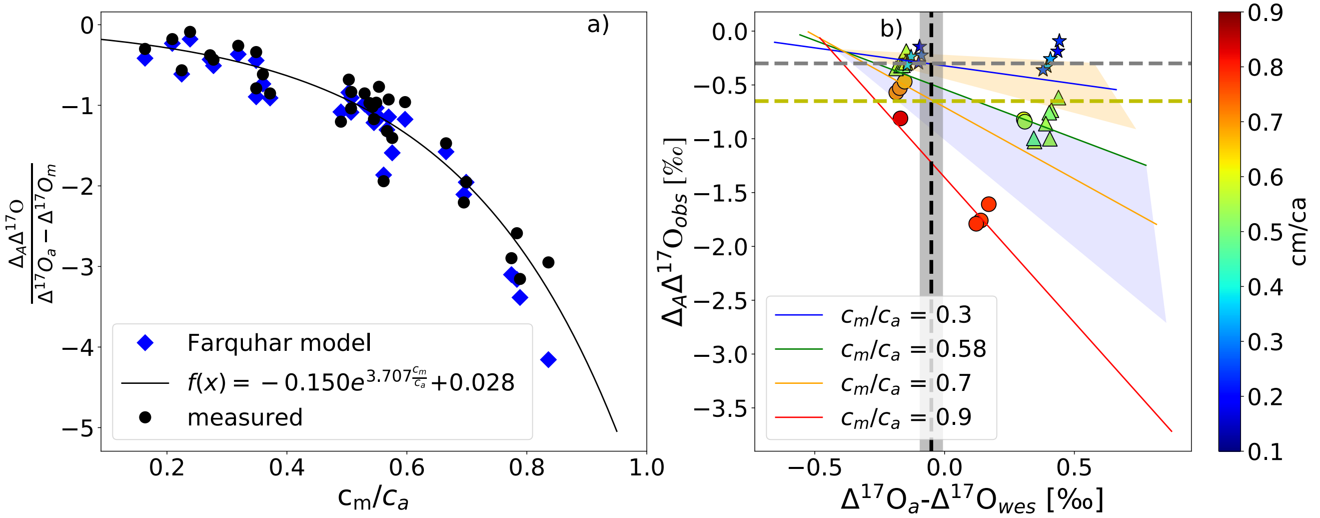

Figure 4 a) Dependency of $\Delta_A\Delta^{17}O$ on the relative difference on the $\Delta^{17}O$ CO_2 entering the leaf and the $\Delta^{17}O$ of CO_2 in equilibrium with leaf water against c_m/c_a ratio. b) dependency of $\Delta_A\Delta^{17}O$ on the difference between the $\Delta^{17}O$ of CO_2 entering the cuvette and the $\Delta^{17}O$ of leaf water at the evaporation site color coded for different c_m/c_a ratios. The solid lines are results of the leaf cuvette model for different c_m/c_a ratios stated in the legend. The dashed vertical black line indicates the difference between the global average $\Delta^{17}O$ value for CO_2 (-0.168 ‰) and leaf water (-0.067 ‰) (Koren et al., 2019). The gray and yellow horizontal dashed lines indicate $\Delta_A\Delta^{17}O$ of C_4 and C_3 plants for c_m/c_a ratio of 0.3 and 0.7, respectively globally.

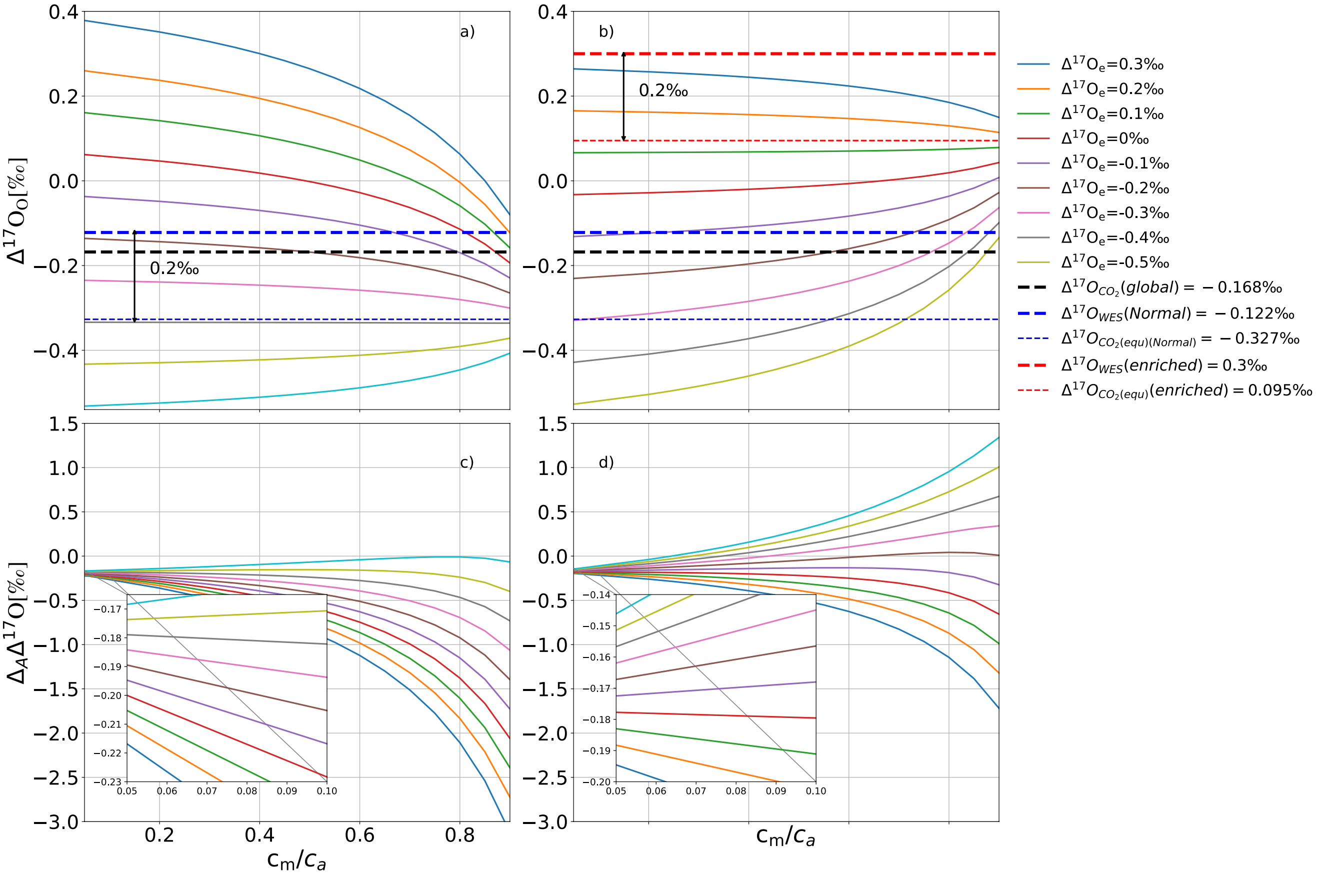

Figure 5 a) and b) $\Delta^{17}O_a$ as a function of c_m/c_a for various values of $\Delta^{17}O_e$ (see legend) for $\Delta^{17}O_{wes}$ = -0.122 ‰ in a) and $\Delta^{17}O_{wes}$ = 0.300 ‰ in b). c) and d) show the corresponding values for $\Delta_A\Delta^{17}O$. $\Delta^{17}O_{global}$ is the global average $\Delta^{17}O$ value for atmospheric CO_2 (Koren et al., 2019). When $\Delta^{17}O$ of CO_2 entering the cuvette is approximately 0.2 ‰ lower than the $\Delta^{17}O$ of leaf water at the CO_2-H_2O exchange site, $\Delta^{17}O$ of the CO_2 leaving the cuvette does not change when c_m/c_a vary.

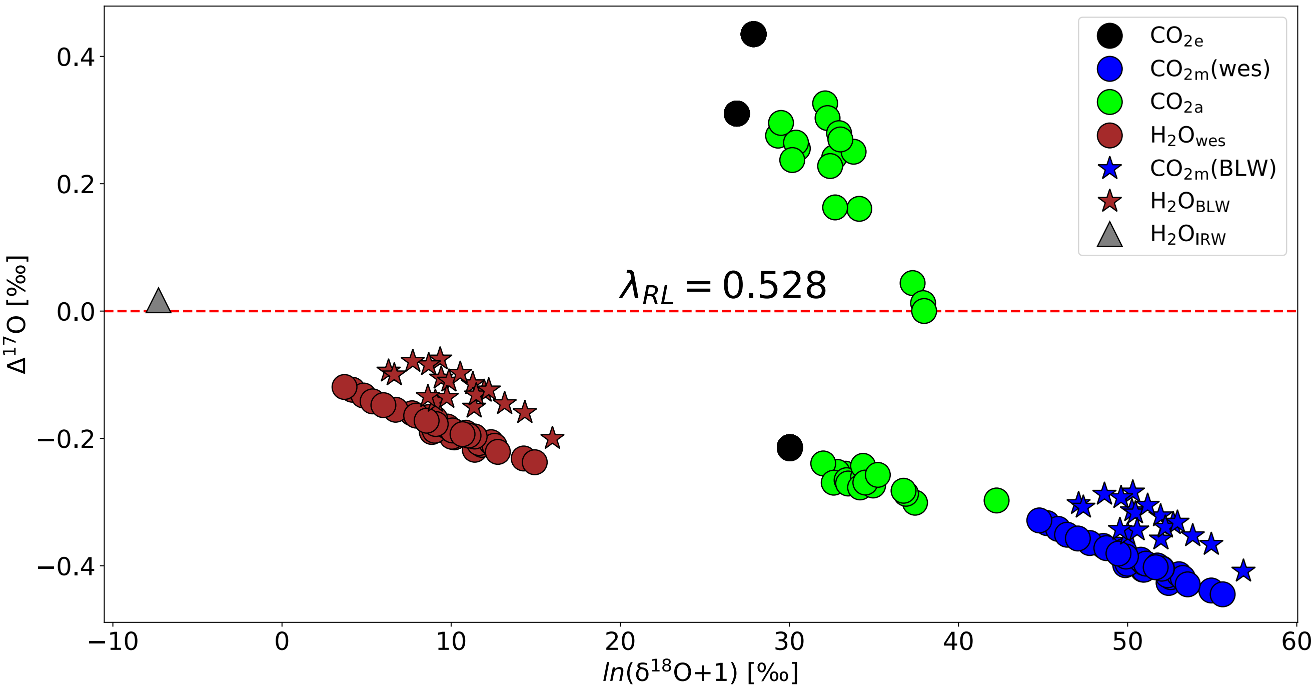

Figure 6 Isotopic composition of various relevant oxygen reservoirs that affect the $\Delta^{17}O$ of atmospheric CO_2 during photosynthesis: irrigation water (grey triangle), calculated leaf water at the evaporation site (brown circles), measured bulk leaf water (brown star), CO_2 entering the cuvette (black circles), CO_2 leaving the leaf cuvette (green circles), CO_2 equilibrated with leaf water at the evaporation site (blue circles), CO_2 equilibrated with bulk leaf water (blue stars). $\Delta^{17}O$ is calculated with λ=0.528.

Table 1: Summary of gas exchange parameters and isotopic composition of maize, sunflower and ivy. Mole fraction at the site of exchange (c_m) is calculated assuming complete isotopic equilibrium with the water at the CO_2-H_2O exchange site. The water at the CO_2-H_2O exchange site is assumed the same as the isotopic composition at the site of evaporation. Number in the parenthesis are the standard deviation of the mean (1σ).

Parameter		Sunflower	Ivy	Maize	Irradiance (μmol m^{-2} s^{-1})
A_n	μmol mol^{-1}m^{-2}s^{-1}	18(0.7)	12(0.7)	17(2)	300
		29(2)	15(2)	32(2)	1200
g_s	mol m^{-2}s^{-1}	0.45(0.14)	0.11(0.02)	0.08(0.01)	300
		0.40(0.04)	0.15(0.03)	0.16(0.02)	1200
$\delta^{18}O_e$	‰	27.26 to 31.80	28.28 to 30.48	27.26 to 30.48	
$\Delta^{17}O_e$	‰	-0.227 to 0.409	-0.215 to 0.435	-0.215 to 0.310	
$\delta^{18}O_a$	‰	33.25 to 43.87	32.64 to 35.86	34.04to 29.764	
$\Delta^{17}O_a$	‰	-0.333 to 0.163	-0.276 to 0.327	-0.270 to 0.296	
$\Delta_A{}^{18}O_{obs}$	‰	57.12(4.70)	22.20(1.32)	17.23(1.32)	300
		34.48(3.25)	24.35(3.09)	12.78(0.83)	1200
$\Delta_A\Delta^{17}O_{obs}$	‰	-2.61 to -0.43	-1.03 to -0.19	-0.36 to -0.09	
$\delta^{18}O_m$	‰	52.02(1.24)	47.17(1.17)	52.62(0.52)	300
		52.62(1.42)	51.09(1.76)	55.15(1.55)	1200
$\Delta^{17}O_m$	‰	-0.41(0.001)	-0.35(0.001)	-0.40(0.01)	300
		-0.41(0.01)	-0.38(0.02)	-0.42(0.02)	1200
c_a	ppm	402 (3)	403 (3)	403 (3)	
c_i	ppm	357(10)	284(0.1)	194(20)	300
		323(10)	301(13)	194(15)	1200
c_c	ppm	277(15)	188(30)		300
		201(42)	163(21)		1200
c_m	ppm	320(10)	220(10)	134(15)	300
		252(27)	214(12)	88 (17)	1200

Table 2: Summary of the parameters used for the extrapolation of leaf scale experiments to the global scale and the results obtained. Bottom part: overview of available Δ^{17}O measurements.

Parameters and values used for global estimation		
Parameter	Value	ref
GPP	120 PgCyr^{-1}	(Beer et al., 2010)
f_{C4}	23 %	(Still et al., 2003)
f_{C3}	77 %	(Still et al., 2003)
c_m/c_a (C$_3$)	0.7	(Hoag et al., 2005)
c_m/c_a (C$_4$)	0.3	(Hoag et al., 2005)
Δ^{17}O leaf water (global mean, modelled)	-0.067±0.04 ‰	(Koren et al., 2019)
Δ^{17}O CO$_2$ (global mean, modelled)	-0.168±0.013 ‰	(Koren et al., 2019)
$\Delta_A\Delta^{17}$O (global mean for C$_4$)	-0.3±0.18 ‰	(Figure 5b, for c_m/c_a ratio of 0.3)
$\Delta_A\Delta^{17}$O (global mean for C$_3$)	-0.65±0.18 ‰	(Figure 5b, for c_m/c_a ratio of 0.7)
$\Delta_A\Delta^{17}$O (global mean for whole vegetation)	-0.57±0.14 ‰	(Equation 13)
$\Delta_A\Delta^{17}$O-isoflux (global mean for C$_4$)	-7.3±4 ‰PgCyr^{-1}	(Equation 14, only for C$_4$)
$\Delta_A\Delta^{17}$O-isoflux (global mean for C$_3$)	-53±15 ‰PgCyr^{-1}	(Equation 14, only for C$_3$)
$\Delta_A\Delta^{17}$O-isoflux (global mean for whole vegetation)	-60±15 ‰PgCyr^{-1}	(equation 14)
$\Delta_A\Delta^{17}$O-isoflux (global mean for whole vegetation)	-47 ‰PgCyr^{-1}	(Hoag et al., 2005)
$\Delta_A\Delta^{17}$O-isoflux (global mean for whole vegetation)	-42 to -92 ‰PgCyr^{-1}	(Hofmann et al., 2017)
Δ^{17}O value of tropospheric CO$_2$		
Δ^{17}O(CO$_2$) for CO$_2$ samples collected in La Jolla-UCSD (California, USA) (1990 to 2000)	-0.173±0.046 ‰	(Thiemens et al., 2014)
Δ^{17}O(CO$_2$) for CO$_2$ samples collected in Israel	0.034±0.010 ‰	(Barkan and Luz, 2012)
Δ^{17}O(CO$_2$) for CO$_2$ samples collected in South china sea (2013-2014)	-0.159±0.084 ‰	(Liang et al., 2017b;Liang et al., 2017a)
Δ^{17}O(CO$_2$) for CO$_2$ samples collected in Taiwan (2012-2015)	-0.150±0.080 ‰	(Liang et al., 2017b;Liang et al., 2017a)
Δ^{17}O(CO$_2$) for CO$_2$ samples collected in California (USA) (2015)	-0.177±0.029 ‰	(Liang et al., 2017b;Liang et al., 2017a)
Δ^{17}O(CO$_2$) for CO$_2$ samples collected in Göttingen (Germany) (2010-2012)	-0.122±0.065 ‰	(Hofmann et al., 2017)

Table 3 List of symbols and variables

Symbol	description	Unit/calculation/value
Gas exchange		
A_n	Rate of CO_2 assimilation	$\frac{u_e}{s}\left(c_e - c_a\left(\frac{1-w_e}{1-w_a}\right)\right)$, mol $m^{-2}s^{-2}$
E	Transpiration rate	$\frac{u_e}{s}\left(\frac{w_a-w_e}{1-w_a}\right)$, mol $m^{-2}s^{-2}$
w_i	Mole fraction of water vapour inside leaf	$\frac{613.65\times e^{\left(\frac{17.502\times T_{leaf}}{240.97+T_{leaf}}\right)}\times 10^{-5}}{P}$, mol mol^{-1}
w_a	Mole fraction of water vapour leaving the cuvette /leaf surrounding	WVIA / IRGA, mol mol^{-1}
w_e	Mole fraction of water vapour entering the cuvette	WVIA/IRGA, mol mol^{-1}
c_e	Mole fraction of CO_2 entering the cuvette	IRGA, mol mol^{-1}
c_a	Mole fraction of CO_2 in the leaf surrounding/ leaving the cuvette	IRGA, mol mol^{-1}
u_e	Flow rate of air entering the cuvette	mol s^{-1}
s	Surface area of the leaf inside the cuvette	m^{-2}
P	Atmospheric pressure	bar
T_{leaf}	Leaf temperature	^{o}C
$g_{s(H2O)}$	Stomatal conductance for water vapour	$\frac{g^t_{H_2O}\times g_{b(H_2O)}}{g_{b(H_2O)}-g^t_{H_2O}}$
$g_{b(H2O)}$	Boundary layer conductance for water vapour	Calibrated for the cuvette we used
g^t_{H2O}	Conductance for water vapor through the boundary layer and stomata	$E\left(\frac{1-\left(\frac{w_i+w_a}{2}\right)}{w_i-w_a}\right)$, mol $m^{-2}s^{-1}$
g_s	Stomatal conductance for CO_2	$\frac{g_{s(H_2O)}}{1.6}$
g_b	Boundary conductance for CO_2	$\frac{g_{b(H_2O)}}{1.37}$
g^t_{CO2}	Conductance for CO_2 through the boundary layer and stomata	$\frac{g_s\times g_b}{g_s+g_b}$
$\boldsymbol{\Gamma^*}$	CO_2 compensation point	45 µmol $m^{-2}s^{-1}$
g_{m13}	CO_2 conductance from intercellular air space to the site of carboxylation calculated using $\Delta_A{}^{13}C$ (for C_3 plants only)	mol $m^{-2}s^{-1}bar^{-1}$
g_{m18}	CO_2 conductance from intercellular air space to CO_2-H_2O exchange site calculated using $\Delta_A{}^{18}O$	mol $m^{-2}s^{-1}bar^{-1}$
g_{m17}	CO_2 conductance from intercellular air space to CO_2-H_2O exchange site calculated using $\Delta_A{}^{17}O$	mol $m^{-2}s^{-1}bar^{-1}$
$g_{m\Delta17}$	CO_2 conductance from intercellular air space to CO_2-H_2O exchange site calculated using $\Delta_A\Delta^{17}O$	mol $m^{-2}s^{-1}bar^{-1}$
c_i	Mole fraction of CO_2 in the intercellular air space	$\frac{\left(g^t_{CO_2}-\frac{E}{2}\right)c_a-A_n}{\left(g^t_{CO_2}+\frac{E}{2}\right)}$ mol mol^{-1}
c_s	Mole fraction of CO_2 at the leaf surface	$c_a - \frac{A_n}{g_b}$
c_m	Mole fraction of CO_2 at the site of CO_2-H_2O exchange	mol mol^{-1}
c_c	Mesophyll conductance to the chloroplast (for C_3 plants)	$c_i - \frac{A_n}{g_{m13}}$ mol mol^{-1}
t^{13}	Ternary correction for $^{13}CO_2$	$\frac{(1+a_{13bs})E}{2g^t_{CO_2}}$
t^{18}	Ternary correction for $C^{18}OO$	$\frac{(1+a_{18bs})E}{2g^t_{CO_2}}$
t^{17}	Ternary correction for $C^{17}OO$	$\frac{(1+a_{17bs})E}{2g^t_{CO_2}}$

R_D	Dark respiration rate	0.8 μmol $m^{-2}s^{-1}$
R_L	Day respiration rate	0.5 × R_D μmol $m^{-2}s^{-1}$
Oxygen and carbon isotope effects		
$\varepsilon^{18}{}_{k}$	Kinetic fractionation of water vapour in air	$\frac{28g_b+19g_s}{g_b+g_s}$, ‰
$\varepsilon^{18}{}_{equ}$	Equilibrium fractionation between liquid and gas phase of water vapor	$2.644 - 3.206(\frac{10^3}{T_{leaf}}) + 1.534(\frac{10^6}{T_{leaf}})$, ‰
a_{13bs}	Weighted fractionation for ^{13}COO as CO_2 diffuses through the boundary layer and stomata	$\frac{(c_s-c_i)a_{13s}+(c_a-c_s)a_{13b}}{c_a-c_i}$ ‰
a_{17bs}	Weighted fractionation for $C^{17}OO$ as CO_2 diffuses through the boundary layer and stomata	$\frac{(c_s-c_i)a_{17s}+(c_a-c_s)a_{17b}}{c_a-c_i}$ ‰
a_{18bs}	Weighted fractionation for $C^{18}OO$ as CO_2 diffuses through the boundary layer and stomata	$\frac{(c_s-c_i)a_{18s}+(c_a-c_s)a_{18b}}{c_a-c_i}$ ‰
a_{13bs}	Weighted fractionation for ^{13}COO as CO_2 diffuses through the boundary layer and stomata	$\frac{(c_s-c_i)a_{13s}+(c_a-c_s)a_{13b}}{c_a-c_i}$, ‰
a_{18bs}	Weighted fractionation for $C^{18}OO$ as CO_2 diffuses through the boundary layer and stomata	$\frac{(c_s-c_i)a_{18s}+(c_a-c_s)a_{18b}}{c_a-c_i}$, ‰
a_{17bs}	Weighted fractionation for $C^{17}OO$ as CO_2 diffuses through the boundary layer and stomata	$\frac{(c_s-c_i)a_{17s}+(c_a-c_s)a_{17b}}{c_a-c_i}$, ‰
$\overline{a}_{17}$	Weighted fractionation of $C^{17}OO$ as it diffuses through the boundary layer, stomata and liquid phase in series	$\frac{(c_i-c_m)a_{17w}+(c_s-c_i)a_{17s}+(c_a-c_s)a_{17b}}{c_a-c_m}$, ‰
$\overline{a}_{18}$	Weighted fractionation of $C^{18}OO$ as it diffuses through the boundary layer, stomata and liquid phase in series	$\frac{(c_i-c_m)a_{18w}+(c_s-c_i)a_{18s}+(c_a-c_s)a_{18b}}{c_a-c_m}$, ‰
a_{13b}	Fractionation in $^{13}CO_2$ as CO_2 diffuses through the boundary layer	2.9‰
a_{13s}	Fractionation in $^{13}CO_2$ as CO_2 diffuses through the stomata	4.4‰
a_m	Fractionation factor for dissolution and diffusion through water	1.8‰
f	Fractionation factor for photorespiration (decarboxylation of glycine)	16‰
e	Fractionation factor for day respiration	$R_D + e^*$, ‰
e^*	Apparent fractionation for day respiration	$\delta^{13}C_a - \Delta_A{}^{13}C - \delta^{13}C_{substrate}$, ‰
b	Fractionation factor for uptake by RubisCO	29‰
α_f	Fractionation due to photorespiration (decarboxylation of glycine)	1+f
α_e	Fractionation due to day respiration	1+e
α_b	Fractionation due to uptake by RubisCO	1+b
a_{17b}	Fractionation of $C^{17}OO$ as CO_2 diffuses through the boundary layer	2.9‰
a_{17s}	Fractionation in $C^{17}OO$ as CO_2 diffuses through stomata	4.4‰
a_{18b}	Fractionation of $C^{18}OO$ as CO_2 diffuses through the boundary layer	5.8‰
a_{18s}	Fractionation in $C^{18}OO$ as CO_2 diffuses through stomata	8.8‰
a_{17w}	Fractionation in $C^{17}OO$ due to diffusion and dissolution in water	0.382‰
a_{18w}	Fractionation in $C^{18}OO$ due to diffusion and dissolution in water	0.8‰
$\varepsilon^{18}{}_{W}$	Equilibrium fractionation of CO_2 and water for $C^{18}OO$	$\frac{17604}{T_{leaf}} - 17.93$, ‰
$\varepsilon^{18}{}_{k}$	kinetic fractionation of water vapor in air	$\frac{28 \times g_b + 19 \times g_s}{g_b + g_s}$
$\varepsilon^{18}{}_{equ}$	equilibrium fractionation between liquid and gas phase water	$2.644 - 3.206 \times (\frac{10^3}{T}) + 1.534 \times (\frac{10^6}{T})$

Isotopic composition		
$\delta^{17}O_A$	δ^{17}O of the assimilated CO_2	$\frac{\delta^{17}O_a-\Delta_A{}^{17}O}{\Delta_A{}^{17}O+1} = \delta^{17}O_a - \frac{c_e}{c_e-c_a}(\delta^{17}O_a - \delta^{17}O_e)$
$\delta^{18}O_A$	δ^{18}O of the assimilated CO_2	$\frac{\delta^{18}O_a-\Delta_A{}^{18}O}{\Delta_A{}^{18}O+1} = \delta^{18}O_a - \frac{c_e}{c_e-c_a}(\delta^{18}O_a - \delta^{18}O_e)$
$\delta^{17}O_{io}$	δ^{17}O of CO_2 in the intercellular air space ignoring ternary correction	$\delta^{17}O_A\left(1-\frac{c_a}{c_i}\right)(1+a17bs)+\frac{c_a}{c_i}\left(\delta^{17}O_a-a17bs\right)+a17bs$, ‰
$\delta^{18}O_{io}$	δ^{18}O of CO_2 in the intercellular air space ignoring ternary correction	$\delta^{18}O_A\left(1-\frac{c_a}{c_i}\right)(1+a18bs)+\frac{c_a}{c_i}\left(\delta^{18}O_a-a18bs\right)+a18bs$, ‰
$\delta^{17}O_i$	δ^{17}O of CO_2 in the intercellular air space	$\frac{\delta^{17}O_{io}+t^{17}\left(\delta^{17}O_A\left(\frac{c_a}{c_i}+1\right)-\delta^{17}O_a\frac{c_a}{c_i}\right)}{1+t^{17}}$, ‰
$\delta^{18}O_i$	δ^{18}O of CO_2 in the intercellular air space	$\frac{\delta^{18}O_{io}+t^{18}\left(\delta^{18}O_A\left(\frac{c_a}{c_i}+1\right)-\delta^{18}O_a\frac{c_a}{c_i}\right)}{1+t^{18}}$, ‰
$\delta^{18}O_{trans}$	δ^{18}O of transpired water vapour	$\left(\frac{w_a}{w_a-w_e}\right)\left(\delta^{18}O_{wa}-\delta^{18}O_{we}\right)+\delta^{18}O_{we}$, ‰
$\delta^{18}O_{wes}$	δ^{18}O of water at the evaporation site	$\delta^{18}O_{wes} = \delta^{18}O_{trans} + \varepsilon^{18}{}_k + \varepsilon^{18}{}_{equ} + \frac{w_a}{w_i} \times (\delta^{18}O_{wa} - \varepsilon^{18}{}_k + \delta^{18}O_{trans})$
$\delta^{17}O_m$	δ^{17}O of CO_2 at the site of CO_2-H_2O exchange	$\left(\delta^{17}O_{wes}+1\right)\times(1+\varepsilon_w^{17})-1$, ‰
$\delta^{18}O_m$	δ^{18}O of CO_2 at the site of CO_2-H_2O exchange	$\left(\delta^{18}O_{wes}+1\right)\times(1+\varepsilon_w^{18})-1$, ‰
$\delta^{13}C_{substrate}$	Isotope (^{13}C) ratio of substrate used for dark respiration	$\frac{\delta^{13}C_a-\Delta_A{}^{13}C}{\Delta_A{}^{13}C+1}$, ‰
$\Delta_A{}^{13}C$	^{13}C-photosynthetic discrimination	$\frac{\zeta(\delta^{13}C_a-\delta^{13}C_e)}{1+\delta^{13}C_a-\zeta(\delta^{13}C_a-\delta^{13}C_e)}$, ‰
$\Delta_A{}^{13}C_{obs}$	^{13}C-photosynthetic discrimination (Farquhar model)	$\left(\frac{1}{1-t}\right)\left[a_{13bs}\frac{c_a-c_i}{c_a}\right]+\left(\frac{1+t}{1-t}\right)\left[a_m\frac{c_i-c_c}{c_a}+b\frac{c_c}{c_a}-\frac{\alpha_b}{\alpha_e}e\frac{R_D}{R_D+A_n}\frac{c_c-\Gamma^*}{c_a}-\frac{\alpha_b}{\alpha_f}f\frac{\Gamma^*}{c_a}\right]$
$\Delta_A{}^{13}C_i$	^{13}C-photosynthetic discrimination (assuming no mesophyll conductance, i.e c_i=c_c)	$\left(\frac{1}{1-t}\right)\left[\bar{a}\frac{c_a-c_i}{c_a}\right]+\left(\frac{1+t}{1-t}\right)\left[b\frac{c_i}{c_a}-\frac{\alpha_b}{\alpha_e}e\frac{R_D}{R_D+A_n}\frac{c_i-\Gamma^*}{c_a}-\frac{\alpha_b}{\alpha_f}f\frac{\Gamma^*}{c_a}\right]$
$\Delta_A{}^{18}O$	^{18}O-photosynthetic discrimination	Equation 3, ‰
$\Delta_A{}^{17}O$	^{17}O-photosynthetic discrimination	$\frac{\zeta(\delta^{17}O_a-\delta^{17}O_e)}{1+\delta^{17}O_a-\zeta(\delta^{17}O_a-\delta^{17}O_e)}$, ‰
$\Delta_A{}^{17}O_{FM}$	Farquhar model for ^{17}O-photosynthetic discrimination	$\frac{\bar{a}_{17}+\frac{c_m}{c_a-c_m}\delta^{17}O_{ma}}{1-\frac{c_m}{a-c_m}\delta^{17}O_{ma}}$ ‰
$\Delta_A{}^{18}O_{FM}$	Farquhar model for ^{18}O-photosynthetic discrimination	$\frac{\bar{a}_{18}+\frac{c_m}{c_a-c_m}\delta^{18}O_{ma}}{1-\frac{c_m}{c_a-c_m}\delta^{18}O_{ma}}$ ‰
$\delta^{17}O_e$	δ^{17}O of CO_2 entering the cuvette	‰
$\delta^{17}O_a$	δ^{17}O of CO_2 leaving the cuvette	‰
$\delta^{18}O_e$	δ^{18}O of CO_2 entering the cuvette	‰
$\delta^{18}O_a$	δ^{18}O of CO_2 leaving the cuvette	‰
$\delta^{17}O_{ma}$	δ^{17}O of CO_2 equilibrated with the leaf water at the evaporating site relative to the CO_2 leaving the cuvette	$\frac{\delta^{17}O_m-\delta^{17}O_a}{1-\delta^{18}O_a}$, ‰
$\delta^{18}O_{ma}$	δ^{18}O of CO_2 equilibrated with the leaf water at the evaporating site relative to the CO_2 leaving the cuvette	$\frac{\delta^{18}O_m-\delta^{18}O_a}{1-\delta^{18}O_a}$, ‰
$\delta^{18}O_{we}$	δ^{18}O of water vapour entering the cuvette	WVIA, ‰
$\delta^{18}O_{wa}$	δ^{18}O of water vapour leaving the cuvette/leaf surrounding	WVIA, ‰