



1 **No evidence of a human influence on the mineralization of terrestrial dissolved**
2 **organic matter (DOM) in Alpine fluvial networks**

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11 **Abstract**

12 The influence of human activities on the role of inland waters in the global carbon (C)
13 cycle is poorly constrained. In this study, we investigated the impact of human land use on the
14 degradation of terrestrial dissolved organic matter (DOM) and its potential impact on bacterial
15 respiration in ten independent catchments of the Lake Geneva Basin. Sites were selected
16 along a gradient of human disturbance (agriculture and urbanization), and we monitored
17 bacterial respiration in parallel to DOM bioavailability. Fluorescence spectroscopy was used to
18 determine the influence of human land uses on DOM sources and composition as well as the
19 dynamic of degradation or production of the fluorophores identified in our study sites. Higher
20 bacterial respiration measured in agro-urban streams related to a short-term bioreactive pool
21 (0-6 days of incubation) from autochthonous origin, whose the size increased with human
22 disturbance. On the other hand, the degradation of terrestrial DOM was not impacted by
23 human activities and was not found to contribute substantially to aquatic bacterial respiration.
24 Although human land uses controlled DOM sources, composition and bioavailability at the
25 basin scale, our results showed that human activities in the Lake Geneva Basin had a limited
26 impact on the net C exchanges between inland waters and the atmosphere related to DOM
27 mineralization. Considering that greater proportion of autochthonous-DOM signature is a
28 common feature in human-impacted catchments, the implications of our results likely apply
29 more widely. Yet, on a global scale, the influence of human activities remains to be determine
30 given the large diversity of effects of agriculture and urbanization on freshwater DOM
31 depending on the local environmental context.

32 **1. Introduction**

33 Continental surface waters receive more terrestrial carbon (C) than they export toward
34 oceans, leading to the conceptualization of inland waters as active pipes that process, emit
35 and store C during their travel from lands to oceans (Cole et al., 2007). Within this framework,



36 the mineralization of terrestrial dissolved organic matter (DOM) by aquatic heterotrophic
37 bacterial communities is a key process by which terrestrial C returns to the atmosphere through
38 CO₂ emissions (Fasching et al., 2014; Lapierre et al., 2013; Mayorga et al., 2005) and whose
39 flux is significant at the global scale (Battin et al., 2008). The study of DOM transport and
40 transformation along fluvial networks is therefore of primary importance, yet a major gap in
41 knowledge revolves around the impact of human activities on the reactivity and bacterial use
42 (i.e., respiration or allocation to biomass) of terrestrial DOM in inland waters (Creed et al.,
43 2015; Xenopoulos et al., 2021).

44 Agricultural and urban land uses are major catchment features impacting DOM sources
45 and composition in aquatic ecosystems from local (Wilson and Xenopoulos, 2009) to regional
46 scales (Lambert et al., 2017; Williams et al., 2016). Streams draining agricultural landscapes
47 and/or urban catchments are commonly enriched in DOM of low molecular weight compared
48 to streams draining forested catchments whereby DOM is dominated by aromatic, high
49 molecular weight compounds (Lambert et al., 2017; Williams et al., 2010). Greater proportion
50 of lower molecular weight compounds in agro-urban streams can be the consequence of a
51 greater autochthonous algal production and bacterial activity in nutrient-enriched waters (Fuß
52 et al., 2017; Lu et al., 2014; Williams et al., 2010; Wu et al., 2019), reduced hydrological
53 connection with terrestrial sources (Giling et al., 2014; Parr et al., 2015), or transfer of less
54 humified soil organic matter due to agricultural practices (Humbert et al., 2020; Lambert et al.,
55 2017; Landsman-Gerjoi et al., 2020). As lower molecular weight molecules are typically more
56 labile and easily available for uptake to bacterial communities (Berggren et al., 2010; Catalán
57 et al., 2017; Kaplan and Bott, 1989), greater DOM processing can be expected in agro-urban
58 streams (Hosen et al., 2014; Parr et al., 2015). However, the destabilization of a stock of soil
59 organic material built before the conversion of forests or wetlands for agriculture or urban
60 development can lead to the mobilization of large amount of humic and highly aromatic DOM
61 into surface waters (Ekblad and Bastviken, 2019; Graeber et al., 2012; Hu et al., 2016; Petrone
62 et al., 2011). As a consequence, the impact of human land uses on the dynamic of DOM in
63 inland waters may be highly diverse depending on how agriculture and urbanization affect
64 DOM sources, content, and composition as well as external drivers such as inorganic nutrients
65 known to regulate bacterial DOM processing (Guillemette and del Giorgio, 2012; Reche et al.,
66 1998).

67 Different scenarios about the consequences on net C exchanges for surface waters
68 can be envisaged depending on the impact of human land uses on freshwater DOM. First,
69 increasing delivery of colored and aromatic terrestrial DOM can lead to increase CO₂
70 emissions supposing that this terrestrial material gets respired by bacterial communities
71 (Fasching et al., 2014; Lapierre et al., 2013). Second, increase in the export of low molecular
72 weight DOM, either derived from terrestrial sources or produced in-stream, can results in more



73 DOM respired and emitted as CO₂ into the atmosphere (Bodmer et al., 2016; Borges et al.,
74 2018). However, if changes in DOM composition result from an enhancement in aquatic
75 primary production, enhanced respiration of autochthonously-produced DOM shall not lead to
76 higher net CO₂ emissions since the amount of C emitted into the atmosphere would be lower
77 or equivalent to the amount of CO₂ previously fixed by primary aquatic producers. An
78 alternative scenario would be that the release of simple and labile organic compounds derived
79 from autochthonous sources enhances the degradation of terrestrial aromatic DOM by the so-
80 called priming effect (Bianchi, 2011), inducing a final net increase of C emitted to the
81 atmosphere. The priming effect is a process through which labile pool of DOM can enhance
82 (“prime”) the degradation of a more recalcitrant DOM pool based on interactions between
83 microbial communities and/or changes in their functions (Guenet et al., 2010; Kuzyakov et al.,
84 2000) but its occurrence in aquatic ecosystems is highly debated (Attermeyer et al., 2014;
85 Bengtsson et al., 2018; Lambert and Perga, 2019). To determine how human activities may
86 impact the mineralization of terrestrial DOM and infer the consequence on the active role of
87 inland waters into the global C cycle, it is therefore necessary to evaluate the consequences
88 on DOM sources, composition, but also DOM bioavailability and to identify which fraction of
89 the DOM pool fuels respiration. This information is critical to establish the role of human land
90 uses on the linkage between terrestrial and aquatic ecosystems.

91 In this study, we aimed to investigate the impact of human land uses on the role of
92 inland waters as bioreactors with regard to the processing of terrestrial DOM. Water samples
93 were collected at two periods in ten independent catchments selected along a gradient of
94 human pressure (agriculture and urbanization) in the Lake Geneva Basin. Patterns in DOM
95 degradation were investigated based on standardized degradation experiments and by the
96 consumption of specific compounds of low molecular weights. Decreases in dissolved organic
97 carbon (DOC) concentrations and changes in DOM composition (assessed by fluorescence
98 spectroscopy) during incubations were used to unravel the contribution of different fractions to
99 the bulk DOM kinetic degradation as well as to the bacterial respiration measured under field
100 conditions. With this approach combining field work and laboratory experiments, we
101 specifically aimed to identify the origin of DOM contributing to bacterial respiration in human-
102 impacted streams and to evaluate the impact of human activities on the mineralization of
103 terrestrial DOM in Alpine fluvial networks.

104 **2. Material and methods**

105 *2.1 Study sites and sampling*

106 Stream and river waters were collected in ten independent tributaries of the Lake
107 Geneva, the largest lake of Western Europe located at the border between France and
108 Switzerland in the Western Alps (Figure 1). Lake Geneva is situated in the Alpine foreland



109 between the Alps and the Jura mountains and was carved during Quaternary glaciations
110 mostly into the Tertiary Molasse. Drainage areas of streams and rivers ranged from 11 to 5240
111 km², Strahler order from 2 to 7, mean elevation from 614 to 2124 m, and land cover was
112 dominated by forests (39±19%), croplands (30±22%), grasslands (13±14%), and urban areas
113 (11±14%) according to the Swiss Federal Office for the Environment (FOEN). Agriculture is
114 dominated by non-irrigated arable lands and vineyards, and forests by coniferous and broad-
115 leaved forests. Selected streams and rivers drained a mosaic of land cover categories and
116 were classified as agro-urban or forest-grassland streams if the sum of croplands and urban
117 areas extents was higher or lower than 50%, respectively (Table 1).

118 Samples were collected on two occasions, at the end of autumn between the 13th and
119 14th of November 2018 and at the end of winter between the 5th and 7th of March 2019. Water
120 temperatures during field campaigns were 10.1±1.8 °C and 6.6±1.1 °C, respectively, but
121 discharge and precipitation conditions at the time of collection were similar according to some
122 data provided by the FOEN. Water was collected below surface in 2 L acid-washed high-
123 density polyethylene (HDPE) bottles and filtered on site. A known volume of water (between
124 1000 and 1500 mL) was filtered on pre-combusted (450°C for 4 h) Whatman glass fiber filters
125 (GF/F grade, 0.7 µm nominal pore size, 47 mm diameter) for chlorophyll a (Chla)
126 measurements. This filtered water was either stored in 1L acid-washed HDPE bottles for
127 further use in the incubation experiments and the measurements of bacterial metabolism, as
128 described below, or further filtered at 0.2 µm with polyethersulfone (PES) syringe encapsulated
129 filters for dissolved organic carbon (DOC), colored and fluorescent DOM (CDOM and FDOM),
130 soluble reactive phosphorus (SRP), and dissolved inorganic nitrogen (DIN) measurements.
131 Syringe encapsulated filters were first rinsed with ultrapure water (60 mL) in the laboratory and
132 then with 15-20 mL of water from the sampling site before collecting samples. Samples for
133 DOC concentrations were stored in 40 mL acid-washed glass vials with polytetrafluoroethylene
134 (PTFE)-coated septa, and samples for CDOM and FDOM were stored in 40 mL acid-washed
135 amber glass vials with PTFE-coated septa. Samples for SRP, and DIN were stored separately
136 in 50 mL sterile centrifuge tubes. All samples and filters were brought back to the laboratory
137 within 3 hours in cool and dark conditions. Samples for SRP and Chla measurements were
138 frozen at -21°C until analysis. Other samples were stored in a dark chamber at 4-5°C until
139 analysis typically done within the following two weeks.

140 *2.2 Characterization of DOM degradation kinetics*

141 Incubations were prepared once back to the laboratory. Water samples previously
142 filtered on site at 0.7 µm were divided into 250 mL acid-washed glass flasks and incubated for
143 28 days in the dark at 20°C. Different subsets of flasks (in triplicates) were prepared and
144 sacrificed for DOC measurements and DOM characterization every 3-5 days for the first 10



145 days and then every 5-8 days up to day 28. Samples were filtered at 0.2 μm with PES syringe
146 encapsulated filters as described above and stored in a dark chamber at 4-5°C. DOC
147 measurements were done within 48h after collection, and CDOM and FDOM analyses over
148 the following week. Dissolved oxygen depletion during the incubations was avoided by leaving
149 a large headspace within the glass flasks and by a regular (every 3-4 days) renewal of the
150 headspace.

151 Several models can be used to characterize DOM degradation kinetic. Here we applied
152 a first-order exponential decay model in order to derive a decay constant describing the overall
153 dynamic of DOM degradation (Lambert and Perga, 2019; Lu et al., 2013; Shang et al., 2018)
154 as well as the size of the short- and a long-term reactive carbon pools (STRC and LTRC,
155 respectively). Decreasing DOC concentrations during the incubations were modeled using
156 GraphPad Prism 8 software according to the following equation:

$$157 \quad \text{DOC}(t) = \text{DOC}_{\text{cons}} \times e^{-k \cdot t} + \text{DOC}_{\text{residual}}$$

158 where $\text{DOC}(t)$ is the DOC concentration measured at the incubation time t , DOC_{cons} the
159 amount of DOC consumed at the end of the incubation, k the decay constant, and $\text{DOC}_{\text{residual}}$
160 the concentration of the residual pool remaining in solution at the end of the incubation.
161 Biodegradable DOC (BDOC) was calculated as the difference in DOC between the initial and
162 final time. Furthermore, we used the k decay constant from the model to quantify the STRC
163 and LTRC pools: STRC was defined as the amount of DOC consumed within the first six days
164 of incubation and the LTRC as the amount of DOC degraded between days 6 and 28. The
165 separation between STRC and LTRC pools was based on a breakpoint in the degradation
166 curves observed around the 6th day of incubation in almost all experiments. Finally, changes
167 in DOM composition during the incubations were also monitored by fluorescence
168 measurement coupled to parallel factor analysis (PARAFAC) as described below.

169 *2.3 Degradation of low molecular weights compounds*

170 We also determined the consumption of low molecular weight compounds including
171 carbohydrates (CAR), carboxylic and acetic acids (C&AA), and amino acids (AA) using Biolog
172 Ecoplates® (Garland and Mills, 1991; Weber and Legge, 2009, 2010). Ecoplates® are 96-
173 well microplates containing 31 different carbon substrates (in triplicates) plus a tetrazolium dye.
174 The bacterial respiration activity associated with a specific substrate reduces the tetrazolium
175 dye and produces a color measurable at 590 nm in absorbance. The intensity of color
176 development in color can be related to the amount of substrate consumed (e.g., Berggren and
177 del Giorgio 2015). Water from each site (125 μL) filtered at 0.7 μm was added to each well of
178 one Ecoplates® per site, which was then incubated in the dark at 20°C for 3 to 9 days until
179 the average well color development (AWCD) reached an asymptote. The absorbance at 590
180 nm was measured with a TECAN microplate reader one to two times per day. Color



181 development of each carbon substrate was calculated as the blank-corrected absorbance at
182 the time when the AWCD was closest to 0.5 (0.51 ± 0.11 , $n = 60$) following recommendations
183 of Weber and Legge (2010). Carbon substrates were then classified as CAR, C&AA, or AA
184 according to Weber and Legge (2009). The color development for each of these categories
185 was calculated as the mean of the color of the respective substrates normalized by the AWCD.
186 While Ecoplates® also include other organic compounds such as polymers or amines/amides,
187 we focused in this study on the consumption of low molecular weight compounds known to
188 support bacterial respiration (Kaplan and Bott, 1989).

189 *2.4 Bacterial respiration measurements*

190 The 0.7 μm -filtered waters used for incubation experiments were also used to measure
191 bacterial respiration (BR) in stream and river waters, ensuring comparability between water
192 quality and bacterial metabolism. BR was determined from the decrease in dissolved oxygen
193 (DO) in 60 mL borosilicate serum bottles filled with water, sealed with a butyl stopper and
194 crimped with an aluminum cap without headspace (3 serum bottles per site). These vials were
195 equipped with SP-PSt7 oxygen planar sensor spots (PreSens) in order to follow DO
196 consumption after 24h of incubation in the dark at 20°C. Initial (1h after the start of incubation)
197 and final DO was recorded using a PreSens Fibox 4 equipped with a fiber-optic oxygen
198 transmitter. Calibration of the PreSens Fibox 4 (two-point calibration at 0 and 100% oxygen
199 saturation) was performed and verified before measurements. BR data were converted into
200 carbon units using a respiratory quotient of 1.

201 *2.5 Analytical methods*

202 DIN (defined as the sum of nitrate, ammonium and nitrite) was measured by ion
203 chromatography (Metrohm instrument). SRP was determined by spectrophotometry using the
204 ammonium molybdate-potassium antimonyl tartrate method (AFNOR, 2005). DOC
205 concentrations were measured with a total organic carbon analyzer (TOC-L Series,
206 Shimadzu), with a detection limit of 0.01 mg L^{-1} and a precision better than $\pm 5\%$ based on
207 duplicates and standards. Chla concentrations were determined by spectrophotometry after
208 ethanol extraction (90%). Frozen filters were put in 15 mL sterile centrifuge tubes in which 10
209 mL of ethanol (90%) were added. Tubes were vigorously shaken and then installed in an
210 ultrasonic bath at 70°C for 10 min. Tubes were then stored in a dark chamber for a night,
211 centrifuged 5 min at 4000 rpm and then 10 min at 4000 rpm. Chla concentrations were
212 determined from absorbance at 665 nm after correction of sample turbidity concomitantly
213 measured at 750 nm. Absorbance for CDOM was measured with a Lambda 365 UV/vis
214 spectrophotometer (Perkin Elmer) from 200 to 700 nm (1 nm increment) using a 10 cm quartz
215 cuvette. Napierian absorption coefficients were calculated according to



216 $a_{\lambda} = 2.303 \times A_{\lambda}/L$,
217 where a_{λ} is the absorption coefficient (m^{-1}) at wavelength λ , A_{λ} the absorbance at wavelength
218 λ and L the path length of the optical cell in m. Spectral slopes for the intervals 275–295 and
219 350–400 nm were determined from the linear regression of the *log*-transformed a spectra
220 versus wavelength and used to determine the slope ratio (S_R). The slope ratio S_R , calculated
221 as the ratio of $S_{275-295}$ to $S_{350-400}$, is inversely related to the molecular weight distribution of
222 DOM (Helms et al., 2008).

223 FDOM was measured with a Fluorolog-3 spectrometer (Horiba) using a 1 cm quartz
224 cuvette across excitation wavelengths of 270 – 450 nm (5 nm increment) and emission
225 wavelengths of 300 – 500 nm (2 nm increment) in order to build excitation-emission matrices
226 (EEMs). Lamp intensity and instrument calibration were systematically verified before running
227 samples. EEMs were acquired in sample emission to lamp reference mode, and a correction
228 matrix provided by the manufacturer in both excitation and emission dimensions was
229 automatically applied during acquisition. EEMs were then decomposed into individual
230 components using PARAFAC algorithms (Stedmon et al., 2003). Additional samples collected
231 in Lake Geneva and in the Rhône Basin upstream of Lake Geneva were included in the model
232 (total numbers of EEMs > 800). EEMs preprocessing (Raman scattering removal and
233 standardization to Raman units) was performed prior the PARAFAC modeling. Normalization
234 was done using a Milli-Q water sample run the same day as the sample. A eight components
235 PARAFAC model was obtained using the drEEM 0.3.0 Toolbox (Murphy et al., 2013) for
236 MATLAB (MathWorks, Natick, MA, USA). Split-half analysis, random initialization and
237 visualization of residuals EEMs were used to test and validate the model. The positions of
238 maximums peaks of the PARAFAC components were compared to previous studies with the
239 open fluorescence database OpenFluor using the OpenFluor add-on for the open-source
240 chromatography software OpenChrom (Murphy et al., 2014). The maximum fluorescence F_{Max}
241 values of each component for a particular sample provided by the model were summed to
242 calculate the total fluorescence signal F_{Tot} of the sample in Raman units. The relative
243 abundance of any particular PARAFAC component X was then calculated as
244 $\%C_X = F_{\text{Max}}(X)/F_{\text{Tot}}$. Precision of EEMs-PARAFAC, based on replicate measurements ($n = 5$) of
245 different samples, was ± 0.001 R.U. for F_{Max} values of components C1 to C5 and ± 0.003 R.U
246 for F_{Max} values of components C6 to C8, and $\pm 0.2\%$ for $\%C_X$ for components C1 to C5 and
247 $\pm 0.5\%$ for components C6 to C8. Finally, the variations of PARAFAC components during
248 incubations were expressed as $(F_{\text{Max}}(X)_{\text{tf}} - F_{\text{Ma}}(X)_{\text{t0}})/F_{\text{Max}}(X)_{\text{t0}}$ with t_0 and t_f the initial and final
249 values of F_{Max} , respectively. Based on the accuracy of EEM-PARAFAC measurements
250 estimated by replicate measurements (see above), we considered variations to be significant
251 if the median of response ratio was higher than ± 0.05 for components C1 to C5 and ± 0.1 for
252 components C6 to C8.



253 2.6 Statistical analyses

254 Differences in water quality, DOM degradability, BR and consumption of low molecular
255 weight compounds between agro-urban and forest-grassland streams were tested with a Mann
256 Whitney unpaired *t*-test at 0.05 confidence interval level. Differences between the sampling
257 periods were investigated by paired *t*-tests. The level of significance was set to 0.05. The
258 ROUT method implemented in the GraphPad Prism 8 software was used to identify potential
259 outliers. A principal component analysis (PCA) was performed to investigate the importance
260 of human land uses on water quality relative to other geomorphological features (e.g. mean
261 elevation) of the study sites. The data selected for the PCA were DOC, SRP, DIN, and the
262 relative abundance of PARAFAC components collected during the two main campaigns. The
263 PCA was done using the *prcomp* function in the R software, and the *factoextra* package was
264 used to identify the variables that contribute the most to the first two dimensions of the PCA.
265 The sum of cropland and urban areas extents were used as descriptive variables in the PCA
266 biplot.

267 3. Results

268 3.1 Water quality and DOM composition across catchments

269 Periods of sampling were characterized by relatively similar hydro-climatic conditions
270 with low precipitation and water discharge, and no significant difference was found in the
271 overall water quality between the two campaigns (not shown). However, significant differences
272 were observed between sampling sites depending on their dominant land cover (Figures 2 and
273 3). Chl_a (total range of variation between sites and periods from 0.2 to 54.9 $\mu\text{g L}^{-1}$), SRP (from
274 0.7 to 29.3 $\mu\text{g L}^{-1}$), DIN (from 0.3 to 8.8 mg L^{-1}), and DOC (from 0.7 to 5.9 mg L^{-1}) concentrations
275 were higher in agro-urban streams that were also characterized by DOM of lower molecular
276 weight (higher S_R values) and higher BR values (Figure 2).

277 Eight PARAFAC components were identified in our study sites (Table 2, Supplementary
278 Figure S1), all of which having been already described in previous studies (Graeber et al.,
279 2012; Lambert et al., 2017; Massicotte and Frenette, 2011; Stedmon et al., 2011; Stedmon
280 and Markager, 2005; Williams et al., 2016; Yamashita et al., 2010). Components included three
281 humic-like fluorophores (C1, C3, C4), one fulvic-like fluorophore (C2), two microbial protein-
282 like fluorophores (C5 and C6), and the common tryptophane (C7) and tyrosine (C8) protein-
283 like fluorophores. All components exhibited higher F_{Max} values in agro-urban streams, although
284 the most notable increases were observed for protein-like fluorophores (Figure 3).

285 The two first component of the PCA explained 72.4% of the variance (Figure 4). The
286 first principal component (PC1) was related to protein-like components C6 to C8, DOC and
287 SRP (positive scores) and to humic-like components C2 to C4 (negative scores), whereas the
288 second principal component (PC2) was related to higher C5, DOC and DIN concentrations



289 (positive scores) and lower C1 (negative score). The results of the PCA showed that the
290 variability in DOM composition and nutrient loadings was largely driven by land uses (Figure
291 4). Scores along PC1 were positively related to croplands (Pearson $r = 0.49$, $p = 0.0265$), urban
292 areas (Pearson $r = 0.63$, $p = 0.0027$) and negatively to forest (Pearson $r = -0.72$, $p = 0.0003$)
293 and grasslands (Pearson $r = -0.48$, $p = 0.0295$), but not to catchment area (Pearson $r = -0.04$,
294 $p = 0.86$), Strahler order (Pearson $r = 0.13$, $p = 0.57$) or mean elevation (Pearson $r = -0.32$, p
295 $= 0.18$). No relationship was found between PC2 and geomorphological properties of sampling
296 sites, suggesting an in-stream origin for the components C1 and C5. Based on the optical
297 properties of PARAFAC components (Table 2), PC1 represented a shift in the dominant
298 composition of DOM from terrestrial (components C2-C4) to autochthonous (C6-C8)
299 signatures as human disturbance (croplands + urban areas) increases. PC2, however,
300 reflected the in-stream generation of DOM through photodegradation (C1) and bacterial
301 processing of DOM (C5).

302 *3.2 Kinetics of bacterial DOM degradation and consumption of low molecular weight* 303 *compounds*

304 All incubations were successfully modeled by a first order exponential decay model (r^2
305 $= 0.98 \pm 0.02$), and we estimated the decay constants and the amounts of BDOC, STRC and
306 LTRC pools for all experiments (Figure 5). BDOC ranged from 0.2 to 2.3 mg L⁻¹ (mean =
307 1.0 ± 0.6 mg L⁻¹), corresponding to 9.7 to 57.6% of initial DOC (mean = 33.8 ± 11%). Higher
308 amount of BDOC in agro-urban streams was accompanied by higher decay constants (from
309 0.01 to 0.26 d⁻¹, mean = 0.12 ± 0.07 d⁻¹) and greater availability of STRC (from 0.01 to 1.3 mg
310 L⁻¹, mean = 0.5 ± 0.4 mg L⁻¹) but no significant difference was observed regarding the amount
311 of LTRC (from 0.1 to 1.0 mg L⁻¹, mean = 0.5 ± 0.3 mg L⁻¹) across stream categories. Regarding
312 the evolution of the different fractions of DOM during incubations, no significant changes in
313 F_{Max} values were observed for humic-like components during incubations (Figure 6).
314 Component C5, however, exhibited a positive variation in most experiments, showing that this
315 component tended to be produced upon bacterial activity. On the contrary, the other protein-
316 like components C6 – C8 were consumed during incubations. Finally, the consumption of low
317 molecular weights compounds was greater in agro-urban streams for AA and CAR molecules,
318 but no difference was observed regarding the degradation of C&AA (Figure 7).

319 **4. Discussion**

320 The spatial variability in water quality and DOM sources and composition observed in
321 streams and rivers of the Lake Geneva basin echoes numerous previous works illustrating the
322 impact of human activities on freshwater ecosystems (Lambert et al., 2017; Shang et al., 2018;
323 Wilson and Xenopoulos, 2009). Croplands and urbanization resulted in the enrichment in



324 protein-like, low molecular weight DOM that was found to enhance DOM bioavailability as well
325 as bacterial respiration. However, we found no evidence of an impact of human activities on
326 the degradation of terrestrial DOM, neither that the latter contributed to ecosystem respiration.

327 *4.1 Origin and biodegradability of DOM in agro-urban streams*

328 DOC concentrations in agro-urban streams of the Lake Geneva Basin were higher
329 compared to those measured in forest-grassland streams (Figure 2) and these increases were
330 accompanied by a shift in DOM composition toward more autochthonous signatures (Figure
331 4). Effects of agriculture and urbanization on freshwater DOM can vary widely depending on
332 the environmental context and catchment properties (Shang et al., 2018; Stanley et al., 2012),
333 and in our study we attributed this pattern to the combination of enhanced autochthonous
334 production and higher transfer of terrestrial material. Enrichment in nutrients and increased
335 light exposure in agriculture and urban streams can promote primary production (Catford et
336 al., 2007; Stanley et al., 2012; Taylor et al., 2004), and greater algal biomass in our study sites
337 was evidenced by higher Chl_a concentrations (Figure 2) and the subsequent release of
338 protein-like components (C6 – C8) related to algal DOM (Figure 3, Table 2). Although C5 also
339 relates to autochthonous biological production (Stedmon et al., 2011), its accumulation during
340 incubation experiments implies that this component reflected recently DOM produced by
341 bacterial activity (Figure 6). Higher F_{Max} values of C5 in human-impacted catchments thus
342 represented a positive feedback loop where greater primary production enhanced bacterial
343 activity that shaped DOM composition toward a more bacterial signature (see also Williams et
344 al. 2010). Although of less amplitude, higher F_{Max} values of components associated with
345 terrestrial inputs (C2 – C4, Table 2) and/or photobleaching (C1) indicated a more efficient
346 export of terrestrial material in agro-urban streams. As urbanization tends to limit the
347 hydrological connection between terrestrial and aquatic ecosystems (e.g. Hosen et al. 2014),
348 it is likely that this pattern reflected greater erosion of agricultural soils (Celik, 2005; Graeber
349 et al., 2012). While it was not the purpose of this study to investigate deeper the links between
350 human land uses and DOM sources, our results point that agriculture and urbanization can act
351 in concert to disturb both terrestrial and aquatic sources of DOM leading to greater amount of
352 both humic- and protein-like components in agro-urban streams compared to more natural
353 land cover.

354 Along with changes in DOM sources and composition, the bacterial consumption of
355 DOM was strongly impacted by land use (Figure 5). Similarly to previous studies carried out in
356 urban catchments (Hosen et al., 2014; Parr et al., 2015), the amount and decay constants of
357 bioavailable DOM increased in human-influenced streams and rivers. However, both
358 autochthonous and terrestrial DOM pools could have contributed to this pattern in the Lake
359 Geneva Basin. The specific reactivity of fluorophores measured during incubations and the



360 deconvolution of BDOC into short- and long-term reactive pools allowed us to attribute this
361 increase in bioavailability solely to the consumption of organic compounds generated in
362 stream. Algae are a major source of low molecular weight compounds in aquatic ecosystems
363 through exudation and cell lysis (Kaplan and Bott, 1989), and the release of such labile
364 molecules was evidenced by higher consumption of amino acids and carbohydrates in agro-
365 urban streams (Figure 7) and by the efficient degradation of PARAFAC components related to
366 algal DOM (Figure 6). The strong relationships between the amount of BDOC, the decay
367 constants k , and the size of the STRC pool with the initial contribution of protein-like
368 components C6-C8 (Figures 2 and 5, Supplementary Figure S2) provide another evidence that
369 greater DOM bioavailability in agro-urban streams resulted from greater in-stream production.

370 To the inverse, we found no evidence that the bacterial degradation of terrestrial DOM
371 was influenced by human land uses despite higher contents in inorganic nutrient, higher
372 bacterial activity, and freshly produced autochthonous DOM in agro-urban streams. Terrestrial
373 (C2 – C4) and photo-produced (C1) components showed no significant trends during
374 incubations (Figure 6), yet substantial amount of DOC was consumed during the second phase
375 of the incubation as indicated by the similar sizes of the LTRC and STRC pools (Figure 5).
376 LTRC was positively related to DOC concentrations and F_{Max} values of C1 – C4 components
377 but not with protein-like components (data not shown), implying that this C was somehow
378 related to terrestrial inputs. Although the composition of the molecules that formed the LTRC
379 pools and the drivers of its loss remain unclear, our results evidenced that human land uses
380 did not impact the bacterial consumption of this terrestrial material.

381 *4.2 Linking bacterial respiration to DOM origin and implication of human activities on the* 382 *transformation of terrestrial DOM along fluvial networks*

383 The loss of terrestrial DOM through bacterial mineralization as it moves along fluvial
384 networks contributes to CO₂ emissions toward the atmosphere (Lapierre et al., 2013) and
385 human activities are expected to impact the role of inland waters in the global carbon cycle by
386 disturbing DOM sources and composition (Xenopoulos et al., 2021). Although limited to the
387 Lake Geneva Basin, our data do not support this assumption. Higher BR in agro-urban streams
388 was indeed mostly related to the accumulation and mineralization of molecules generated in-
389 stream by aquatic primary producers (Figure 8), although the photodegradation of terrestrial
390 DOM could also have contributed to increase BR through the transformation of complex and
391 aromatic molecules into compounds of lower molecular weight (Bertilsson and Tranvik, 1998)
392 as suggested by the positive relationship between BR and C1. The fact that BR relies mainly
393 on DOM derived from primary production in agro-urban streams implies that most of the C
394 released toward the atmosphere as CO₂ by bacterial respiration corresponded to atmospheric
395 CO₂ previously fixed by aquatic producers and converted into biomass. Considering moreover



396 that algal DOM can fuel BR also downstream, as suggested by the relationships between C6
397 – C8 with BR and STRC (defined as the amount of C consumed after 6 days), our results imply
398 that human activities had only a minor effect on the C budget and exchanges between inland
399 waters and atmosphere in Alpine fluvial networks.

400 Additionally, the consumption of terrestrial DOM did not differ across land use.
401 Moreover, the lack of relationship between LTRC and BR falls in line with experimental studies
402 showing that, in presence of algal derived DOM, terrestrial DOM is preferentially incorporated
403 into bacterial biomass while autochthonous DOM is allocated to respiration (Guillemette et al.,
404 2016). On average, the consumption of terrestrial DOM (i.e., the LTRC pool) at the timescale
405 of our incubation experiments represented $19\pm 8\%$ of initial DOC, an estimation closed to those
406 reported in a medium-size river basin (400 km²) draining temperate wetlands where only 15%
407 of initial DOC was removed along the whole network (Wollheim et al., 2015). Although we
408 cannot exclude that this terrestrial material could be degraded over longer timescale once
409 entered in the lake Geneva (whose the water residence time is about 12 years), our results
410 are consistent with other studies suggesting that inland waters have only a limited ability to
411 degrade terrestrial C delivered to coastal margins (Borges et al., 2015; Coble et al., 2019; del
412 Giorgio and Pace, 2008; Hanley et al., 2013).

413 Finally, despite the large range of size of our sampling sites (Table 1), we found no
414 relationship between Strahler order (ranging from 2 to 7) and DOM composition and reactivity.
415 This observation contrasts with a recent study where stream order (ranging from 1 to 4)
416 correlated negatively with humic-like DOM but positively with protein-like DOM (Shang et al.,
417 2018), a pattern consistent with a general conceptual trend describing DOM transformations
418 along fluvial continuum. Indeed, the control of DOM dynamic along the river continuum is
419 expected to shift from a dominant influence of terrestrial inputs in the headwaters to a dominant
420 influence of in-stream removal and autochthonous production as stream order increases
421 (Creed et al. 2015). In our study, however, human land uses had a major role in controlling
422 DOM sources and reactivity at the basin scale (Figure 4). The only exception was a positive
423 correlation between Strahler order and the relative proportion of C5 (Supplementary Figure
424 S3), indicating that the degradation of autochthonous DOM in agro-urban streams led to a
425 bacterial imprint on the DOM pool that persists along the aquatic continuum.

426 **5. Conclusion**

427 In the Lake Geneva Basin, human land uses were found to alter the terrestrial and
428 aquatic sources of freshwater DOM. Greater autochthonous production of DOM in agro-urban
429 streams led to higher amount of bioavailable DOM, stimulating ecosystem respiration while no
430 influence on the loss of terrestrial DOM was observed. Despite their dominant influence on



431 DOM composition and reactivity at the basin scale, we found that human land uses had a
432 limited effect in terms of net C flux exchanges between inland waters and atmosphere related
433 to DOM mineralization by heterotrophic bacterial communities.

434 Enrichment in protein-like DOM due to greater autochthonous production is a recurrent
435 observation in agricultural and urban catchments, therefore the results of this study are likely
436 not limited to the Lake Geneva Basin. However, caution should be taken when extrapolating
437 the impact of human activities on the role of inland waters on the C cycle. Indeed, the net
438 effects of agriculture and urbanization on freshwater DOM vary widely depending on the
439 environmental context (Stanley et al., 2012), leading to apparent opposite effect on BDOC.
440 Thus, while our results are in line with previous works (Hosen et al., 2014; Parr et al., 2015), it
441 contrasts with studies reporting no influence of human land uses on the bacterial consumption
442 of DOM despite higher contribution of autochthonous DOM (Kadjeski et al., 2020; Lu et al.,
443 2013) or higher DOM degradability in agriculture streams related to larger proportion of soil-
444 derived DOM (Shang et al., 2018). Therefore, additional works on the links between human
445 activities and DOM reactivity and fate are needed in order to fully assess the future of inland
446 waters in the context of the global C cycle.

447

448 **Author Contributions:** T. L. conceived the study with contribution from M.-E. P. T. L., P. P.,
449 and N. E. collected field samples. T. L. made laboratory analysis. T.L. drafted the
450 manuscript which was substantially commented upon and amended by M.-E. P., P. P., and
451 N. E. All co-authors approved the manuscript.

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- 668



669 **Table caption**

670 **Table 1** – Selected properties and dominant classification of sampling sites

671 **Table 2** – Spectral properties (positions of maximum excitation (ex) and emission (em) peaks)

672 of the eight PARARAC components identified in this study, general description and dominant
673 sources based on previous studies.

674



675 **Figure caption**

676 **Figure 1** – Map of the Lake Geneva Basin and the ten independent catchments sampled
677 during this study.

678 **Figure 2** – Boxplots of (A) Chla, (B) SRP, (C) DIN, (D) DOC concentrations and (E) S_R values
679 and (F) BR in agro-urban (grey) and forest-grassland (white) streams. The box represents the
680 first and third quartile, the horizontal line corresponds to the median, the cross corresponds to
681 the average, and the error bars correspond to the maximum and minimum. Mann Whitney
682 unpaired *t*-test were used to test for statistical differences: ns represents not significant, *
683 = $p < 0.05$, ** = $p < 0.01$, *** = $p < 0.001$, **** = $p < 0.0001$.

684 **Figure 3** – Boxplots of F_{Max} values of PARAFAC components in agro-urban (grey) and forest-
685 grassland (white) streams. The box represents the first and third quartile, the horizontal line
686 corresponds to the median, the cross corresponds to the average, and the error bars
687 correspond to the maximum and minimum. Mann Whitney unpaired *t*-test were used to test for
688 statistical differences: ns represents not significant, * = $p < 0.05$, ** = $p < 0.01$, *** = $p < 0.001$, ****
689 = $p < 0.0001$.

690 **Figure 4** – PCA biplot, including loadings plot for the input variables and scores plot for
691 stations. Markers are shaped according to the sampling period and colored according to a
692 gradient of human disturbance (defined as the sum of % croplands and % urban areas,
693 supplementary ordinal variable in the PCA).

694 **Figure 5** – Boxplots of (A) BDOC concentrations, (B) constant decay *k*, (C) STRC and (D)
695 LTRC pools in agro-urban (grey) and forest-grassland (white) streams. The box represents the
696 first and third quartile, the horizontal line corresponds to the median, the cross corresponds to
697 the average, and the error bars correspond to the maximum and minimum. Mann Whitney
698 unpaired *t*-test were used to test for statistical differences: ns represents not significant, *
699 = $p < 0.05$, ** = $p < 0.01$, *** = $p < 0.001$, **** = $p < 0.0001$.

700 **Figure 6** – Response ratio of PARAFAC components during incubation experiments with t_f =
701 day 28 and t_0 = day 0. Grey bars represent threshold of significance above or below which
702 significant production or consumption of component was identified. See text for details.

703 **Figure 7** – Boxplots of (A) AA, (B) CAR, and (C) C&AA consumption in agro-urban (grey) and
704 forest-grassland (white) streams. The box represents the first and third quartile, the horizontal
705 line corresponds to the median, the cross corresponds to the average, and the error bars
706 correspond to the maximum and minimum. Mann Whitney unpaired *t*-test were used to test for
707 statistical differences: ns represents not significant, * = $p < 0.05$, ** = $p < 0.01$, *** = $p < 0.001$, ****
708 = $p < 0.0001$.



709 **Figure 8** – Relationships between BR and (A) DOC concentrations, (B) the sum of initial F_{Max}
710 values of C6, C7, and C8 components, and (C) initial F_{Max} values of C1 component
711



712 Table 1

713

Site	Area (km ²)	Mean elevation (masl)	Strahler order	Forest (%)	Croplands (%)	Urban areas (%)	Classification
1 - La Combe	39	1241	2	73	3	1	Forest-grassland
2 - Le Boiron	11	692	2	62	36	1	Forest-grassland
3 - Le Grand Curbit	14	614	2	24	70	4	Agro-urban
4 - La Venoge	228	696	4	31	60	6	Agro-urban
5 - La Mèbre	21	597	3	18	38	43	Agro-urban
6 - La Sorge	12	595	3	24	65	11	Agro-urban
7 - La Dranse	638		4	33	14	5	Forest-grassland
8 - Le Rhône	5238	2124	7	23	2	3	Forest-grassland
9 - La Veveyse	65	1105	5	45	20	6	Forest-grassland
10 - La Paudèze	16	775	4	33	40	25	Agro-urban

714



715 Table 2

716

Component	Max ex (nm)	Max em (nm)	Description & dominant sources
C1	<270	424	Widespread humic-like fluorophore, terrestrial ^{a,b} and/or photoproducted ^c .
C2	<270 (330, 380)	498	Fulvic-like fluorophore, widespread, terrestrial origin ^{d,e} .
C3	<270 (355)	438	Humic-like fluorophore, widespread, terrestrial origin ^d .
C4	320	402	Low molecular weight humic-like fluorophore, related to agriculture ^{a,b,e} .
C5	300	336	Protein-like fluorophore associated with biological production ^f .
C6	<270	372	Anthropogenic humic-like fluorophore related to algal ^g or bacterial ^f production in urban areas ^{e,g} .
C7	275	332	Tryptophan-like fluorophore, indicative of autochthonous production ^{b,c} .
C8	270	304	Tyrosine-like fluorophore, indicative of autochthonous production ^{b,c} .

^a Stedmon & Markager, 2005; ^b Yamashita et al., 2010; ^c Massicote & Frenette 2011; ^d Graeber et al., 2012; ^e Lambert et al., 2017; ^f Stedmon et al., 2011; ^g Williams et al., 2016.

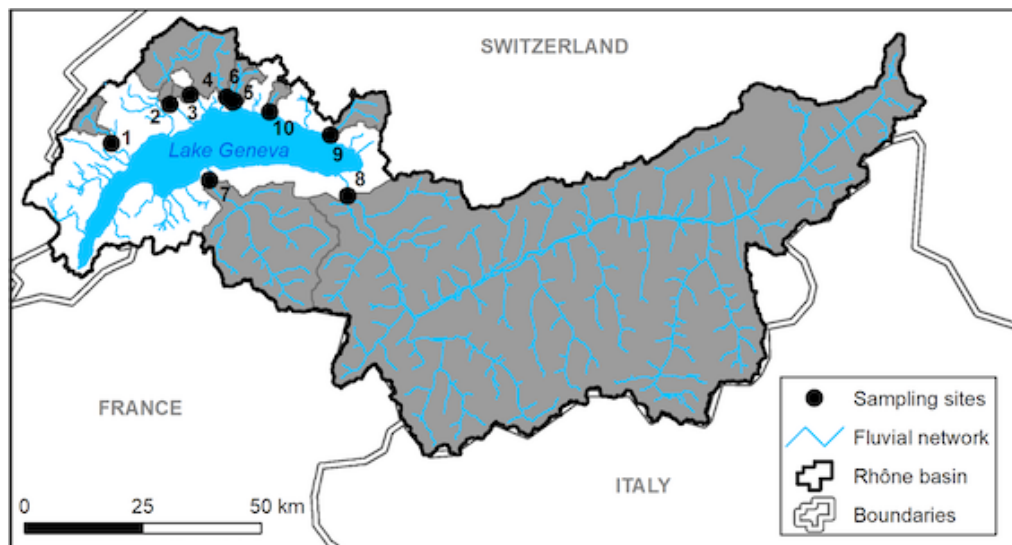
717



718 Figure 1

719

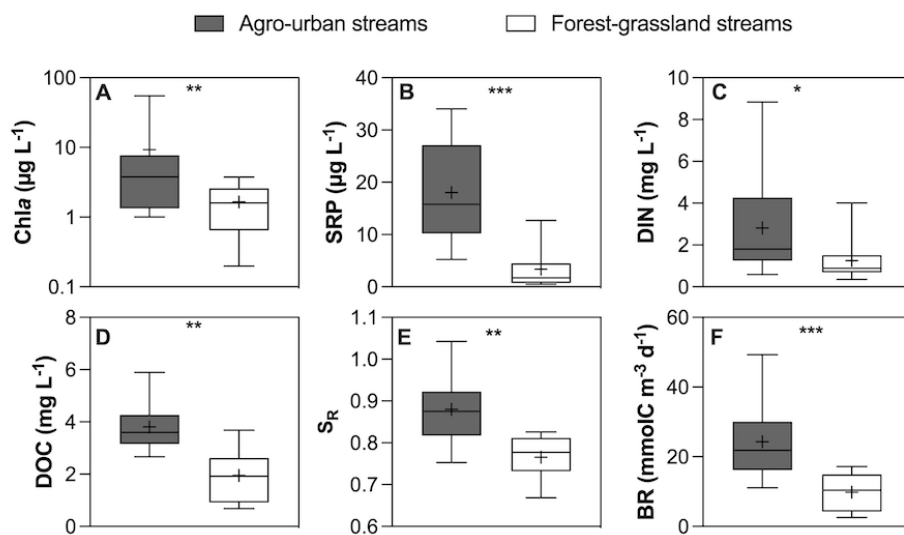
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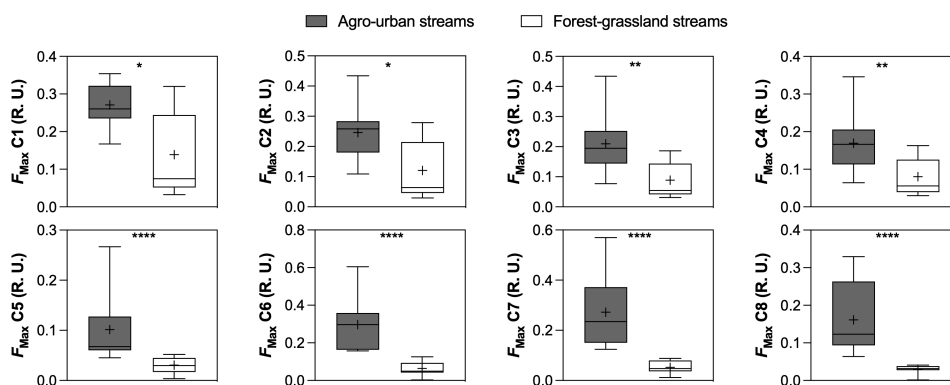
722 Figure 2



723



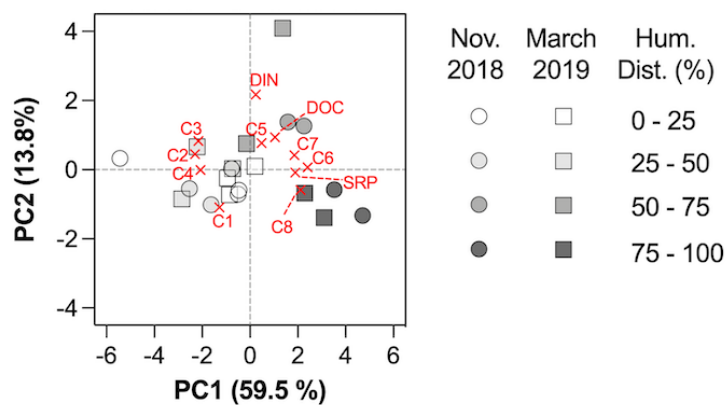
724 Figure 3



725



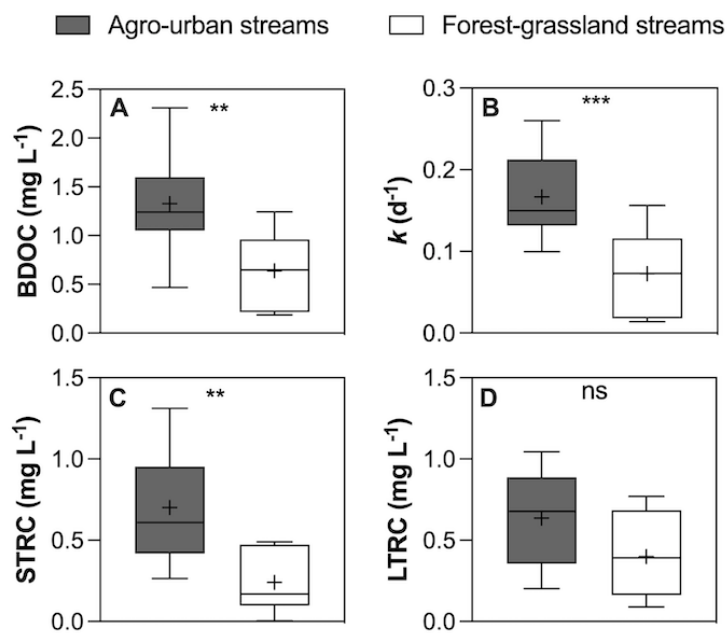
726 Figure 4



727



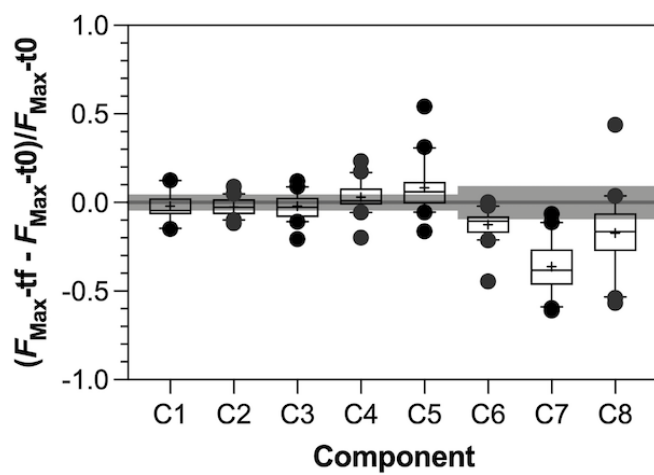
728 Figure 5



729



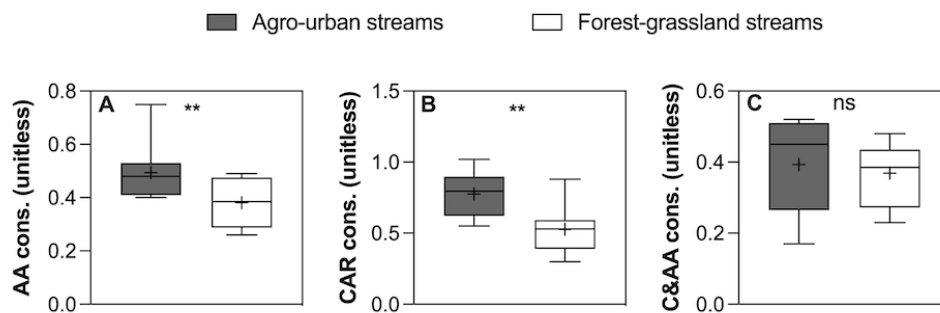
730 Figure 6



731



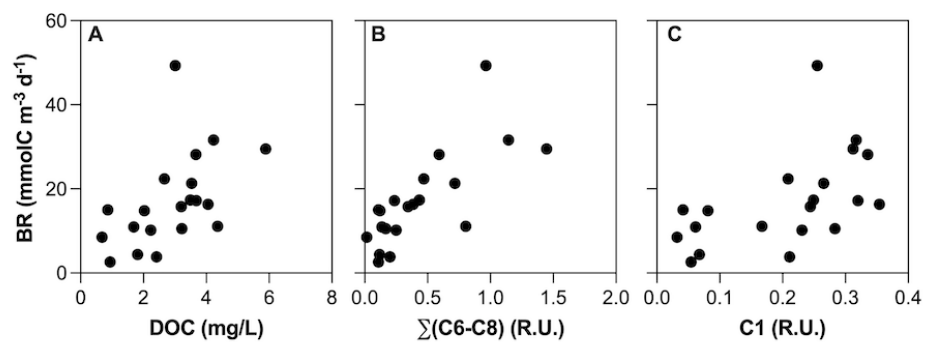
732 Figure 7



733



734 Figure 8



735