



1 **Reviews and Syntheses: Spatial and temporal patterns in metabolic**
2 **fluxes inform potential for seagrass to locally mitigate ocean**
3 **acidification**

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21



22 **Abstract:** As global change continues to progress, there is a growing interest in assessing any
23 local levers that could be used to manage the social and ecological impacts of rising CO₂
24 concentrations. While habitat conservation and restoration have been widely recognized for their
25 role in carbon storage and sequestration at a global scale, the potential for managers to use
26 vegetated habitats to mitigate CO₂ concentrations at local scales in marine ecosystems facing the
27 accelerating threat of ocean acidification (OA) has only recently garnered attention. Early studies
28 have shown that submerged aquatic vegetation, such as seagrass beds, can locally draw down
29 CO₂ and raise seawater pH in the water column through photosynthesis, but empirical studies of
30 local OA mitigation are still quite limited. Here, we leverage the extensive body of literature on
31 seagrass community metabolism to highlight key considerations for local OA management
32 through seagrass conservation or restoration. In particular, we synthesize the results from 62
33 studies reporting *in situ* rates of seagrass gross primary productivity, respiration, and/or net
34 community productivity to highlight spatial and temporal variability in carbon fluxes. We
35 illustrate that daytime net community production is positive overall, and similar across seasons
36 and geographies. Full-day net community production rates, which illustrate the potential
37 cumulative effect of seagrass beds on seawater biogeochemistry integrated over day and night,
38 were also positive overall, but were higher in summer months in both tropical and temperate
39 ecosystems. Although our analyses suggest seagrass meadows are generally autotrophic, the
40 modeled effects on seawater pH are relatively small in magnitude. In addition, we illustrate that
41 periods when full-day net community production is highest could be associated with lower
42 nighttime pH and increased diurnal variability in seawater *p*CO₂/pH. Finally, we highlight
43 important areas for future research to inform the next steps for assessing the utility of this
44 approach for management.



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47

48 **1 Introduction**

49 As carbon dioxide (CO₂) emissions continue to rise, there is an intense interest among managers
50 and decision makers in developing local strategies to minimize the social and ecological costs of
51 global change. While habitat restoration or conservation has been recognized for its utility in
52 carbon storage and sequestration at a global scale in both terrestrial and aquatic systems
53 (Canadell and Raupach 2008, Duarte et al. 2010, Agrawal et al. 2011, Mcleod et al. 2011, Nahlik
54 and Fennessy 2016), the potential for using vegetated habitat to mitigate CO₂ concentrations in
55 the surrounding environment at very local scales has received far less attention. In marine
56 ecosystems, however, increasing CO₂ concentrations cause ocean acidification (OA), which
57 poses a threat to species and ecosystems worldwide (Kroeker et al. 2010; 2013). Submerged
58 aquatic vegetation can reduce CO₂ concentrations in the water column through photosynthesis at
59 a local scale (hereafter termed *local OA mitigation*) (Hendriks et al. 2014). Thus, conservation,
60 restoration, and purposeful culturing of submerged aquatic vegetation have emerged as some of
61 the few potential strategies for local OA mitigation for managers, providing possible benefits to
62 other vulnerable species associated with these habitats (Washington State Blue Ribbon Panel on
63 Ocean Acidification 2012, Chan et al. 2016, Nielsen et al. 2018).

64

65 Because both carbon sequestration and local OA mitigation will depend on the productivity of
66 the vegetation, temporal variability in productivity may be especially important for managers
67 considering local OA mitigation strategies. Carbon storage and sequestration in marine



68 ecosystems is largely considered to be a cumulative process, whereas photosynthesis-based, local
69 OA mitigation in the water column will necessarily be more intermittent or transient in time. In
70 particular, variability in potential local OA mitigation can be expected on several different
71 temporal scales. First, daytime and nighttime patterns in photosynthesis and respiration can cause
72 substantial diurnal variability in seawater pH and saturation state (Hendriks et al. 2014, Pacella et
73 al. 2018), such that potential local OA mitigation will vary on hourly timescales with daylight.
74 Moreover, tidal cycles and local hydrodynamics may alter the impact these processes have on
75 ambient water chemistry on hourly timescales as well (Cyronak et al. 2018, Koweek et al. 2018).
76 Local hydrodynamics could significantly influence the time that a water mass experiences
77 chemical alteration by a seagrass meadow, as well as the water depth through which light must
78 penetrate to reach the seagrass canopy and the volume of water that must be modified, and
79 therefore the magnitude of the change. On a longer timescale, seasonal patterns in temperature
80 and light can also cause substantial seasonal variability in the biomass and productivity of
81 submerged aquatic vegetation (Maher and Eyre 2011, Clavier et al. 2014, Ricart et al. 2021).
82 Thus, the degree of any potential local OA mitigation may vary on seasonal or monthly
83 timescales as well (Manzello et al. 2011, Unsworth et al. 2012, Cryonak et al. 2018, Saderne et
84 al. 2019). This potential for temporal variability in photosynthesis-based, local OA mitigation
85 has important implications for how the strategy might be used by managers, particularly with
86 respect to whether potential OA mitigation aligns with windows of vulnerability for sensitive
87 living resources or how variability is integrated through time by important species.
88
89 Seagrasses are productive marine macrophytes that are considered net autotrophic at the global
90 scale (Duarte et al. 2010), suggesting that gross primary production (GPP, defined here as the



91 rate of oxygen production) should exceed respiration (R, defined here as the rate of oxygen
92 consumption, determined using incubations or measurements in the dark) on average. Assessing
93 the utility of seagrass meadows for local OA mitigation, however, requires a better
94 understanding of the spatial and temporal variability in GPP, R, and net community production
95 (NCP, defined here as the GPP-R, or the net carbon flux, when both photosynthesis and
96 respiration are accounted for). For example, diurnal variability in seawater carbonate chemistry
97 could be more pronounced in geographies that support higher daytime GPP and nighttime R
98 (Duarte et al. 2010). In addition, seasonal variability in NCP could be more pronounced in
99 locations where there are larger fluctuations in light and other environmental drivers, such as
100 high latitude or temperate ecosystems.

101

102 Scientists have been quantifying NCP in seagrass meadows for decades (Odum 1956), and
103 insight regarding the potential for temporal and spatial variability in carbon fluxes can be gained
104 using this literature. In particular, understanding whether GPP, R, and NCP vary predictably
105 across time and in different geographies can provide important, first order information about
106 when and where local OA mitigation approaches might be effective, as well as insight into
107 challenges that could arise in particular localities. Although potential local OA mitigation is
108 caused by changes in the dissolved inorganic carbon (DIC) in seawater, which can be influenced
109 by several other important biological and physical factors (Koweek et al. 2018, James et al.
110 2020), the relationship between O₂ fluxes associated with seagrass metabolism and seawater DIC
111 is roughly proportional (i.e., if O₂ production goes up, DIC in seawater will go down). While
112 empirical studies of changes in seawater DIC are currently limited, the comprehensive literature



113 on seagrass community metabolism on O₂ fluxes can provide important spatial and temporal
114 context for managers interested in carbon fluxes and the potential for local OA mitigation.
115
116 Here, we synthesize published studies of seagrass metabolism to characterize the variability in
117 carbon fluxes associated with GPP, R, and NCP across seasons and geographies. In recognition
118 of the substantial temporal diel variability in carbon fluxes associated with daytime NCP and
119 nighttime respiration, as well as the uncertainty in our understanding of how this temporal
120 variability is integrated by vulnerable marine organisms associated with seagrass beds, we focus
121 on both hourly rates of NCP taken during peak daylight hours and full-day NCP. Hourly
122 measurements of NCP collected during peak daylight hours can provide insight into the
123 maximum elevation of seawater pH. Similarly, hourly measurements of respiration provide
124 insight into the potential maximum depression of nighttime pH. In contrast, measurements of
125 NCP taken over longer time periods or that incorporate the full 24 hour cycle (full-day NCP)
126 provide insight into the cumulative effect of seagrass on seawater chemistry. In particular, we
127 tested: (1) If seasonal variability is present in daytime and full-day carbon fluxes, (2) If the
128 temporal variation in carbon fluxes varies among tropical and temperate geographies, and (3)
129 How much of the residual variation in carbon fluxes not accounted for by seasons or geography
130 can be attributed to variation in temperature and aboveground biomass of the seagrass
131 assemblage. To connect the metabolic measurements to seawater chemistry, we model potential
132 changes in bulk seawater pH based on the estimated carbon fluxes given variation in seawater
133 residence time and water depth.

134

135 **2 Methods**



136 We conducted a literature search for *in situ* measurements of seagrass community metabolism
137 using the Web of Science. Search terms included [*seagrass* OR *eelgrass* OR *submerged aquatic*
138 *vegetation*] AND [*metabol** OR *carbon/oxygen fluxes* OR *community prod** *community resp**
139 OR *benthic incubation chambers* OR *primary prod** OR *carbon* chemistry* OR *pH*]. For each
140 paper, we then searched the literature cited for more applicable studies, as well as any papers
141 listed in the Web of Science that cited the study in question. In addition, we searched the datasets
142 used by Duarte et al. (2010) and Unsworth et al. (2012). Studies were included when either O₂ or
143 carbon fluxes of a seagrass-associated community were measured *in situ*. This included studies
144 using a variety of methods, including incubation chambers, eddy correlation techniques, mass
145 balance estimates, and isotope enrichment, among others. Studies were included that were
146 published prior to January 1, 2020.

147

148 Within a single study, regardless of the methods used, each deployment/set of measurements was
149 included as a data point in the synthesis when deployments/measurements were repeated across
150 different locations, months, or species. We collected measurements of GPP, R or NCP from each
151 study using data reported in the text, tables, or graphs using software (Graph Click or Data
152 Thief), or provided by the authors by request. In addition, we recorded information on the
153 photosynthetic quotient (PQ) and respiratory quotient (RQ) values used to convert from O₂ to
154 carbon, as well as other metadata associated with the study (e.g., species, location, temperature,
155 month the study was conducted, etc.). We classified each study as either tropical or temperate
156 based on the designation in the primary study and then classified the metabolic measurements as
157 either (a) hourly rates or (b) daily rates. This classification was defined by the reporting within
158 the studies (i.e., the primary authors either reported hourly or daily rates), but the difference in



159 reporting was ostensibly due to differences in the length of the deployment used to measure
160 metabolism (e.g., <4 hour deployment = an hourly rate, ~12-24 hour deployment = a daily rate).
161 The shorter “hourly” deployments were usually taken during peak daylight hours, which we used
162 to infer the potential for any daytime local OA mitigation. In contrast, we use the daily rates to
163 infer the cumulative, full-day local OA mitigation potential of seagrass. It is important to note
164 that positive daily NCP (used to infer the full-day local OA mitigation potential) can still
165 encompass marked diel or diurnal variability in carbon fluxes that could prove deleterious to
166 seagrass associated species during transient periods of low pH.

167

168 Although no studies included here measured changes in seawater DIC directly, several studies (N
169 = 17) converted metabolism measurements based on oxygen consumption and production to
170 units of carbon. For studies that only reported metabolism in units of oxygen, we converted the
171 reported GPP, R, and NCP measurements to carbon using a PQ or RQ of 1 (Duarte et al. 2010).
172 We then converted all measurements to the same scale: either $mmol\ C/m^2/hour$ or $mmol$
173 $C/m^2/day$. Positive NCP values represent net autotrophy and carbon fluxes *from* the water
174 column *to* the seagrass tissue, and negative NCP values represent net heterotrophy and carbon
175 fluxes *from* the seagrass tissue *to* the water column.

176

177 To assess differences in estimates based on the methods used to measure metabolism, we plotted
178 the carbon fluxes as a function of study type. Based on these plots (Fig. S1), we decided to
179 perform separate analyses for studies that used the “mass balance” approach (Odum 1956) versus
180 other methods (e.g., incubations, eddy correlations). The studies using the mass balance
181 approach often do not differentiate between water column and benthic productivity, and as such,



182 displayed a much higher range and magnitude of responses than those measured by other
183 methods (see *Methodological Analyses in Results* below).

184

185 To assess the drivers of variability in carbon fluxes (hourly and daily rates of GPP, R and NCP),
186 we first assessed the collinearity in the primary drivers of interest: seawater temperature and
187 aboveground biomass for those studies reporting both variables. Because temperature and
188 aboveground biomass are correlated (temperature \times geography: $P = 0.011$, Fig S2), we decided
189 to focus first on the effect of season, using *month* as a predictor variable, which ostensibly
190 encompasses some of the variability in both temperature and aboveground biomass (Figs S3 and
191 S4). Furthermore, the use of month as a predictor variable allowed us to include the maximum
192 number of studies in the analysis, since not all studies reported temperature and aboveground
193 biomass. To standardize months to seasons across the hemispheres, we used the numerical
194 notation for months in the northern hemisphere (i.e., January = 1, etc.). For the southern
195 hemisphere, we subtracted 6 from the numerical notation and used the absolute value. In
196 addition, we tested for differences between seagrass communities in temperate and tropical
197 geographies based on the hypothesis that seagrass meadows in temperate geographies have
198 greater seasonality in light, temperature, and aboveground biomass, and thus, should have a more
199 pronounced seasonal fluctuation (Fig S3 and S4).

200

201 We then tested for effects of temperature and aboveground biomass on the residual variation of
202 the monthly models. Specifically, we first fit mixed-effects models of both hourly and daily rates
203 of GPP, R, and NCP using maximum likelihood with *geography* as a categorical factor (tropical
204 vs. temperate) and a linear and quadratic term for *month* as fixed effects, as well as the two-way



205 interactions between *geography* and each term for *month*. We also included *replicate* nested in
206 *study* as a random effect to account for non-independence arising from the inclusion of repeated
207 measures from the same sites over time and measurements from multiple sites within a single
208 study:

$$\begin{aligned} 209 \quad \text{metabolism} \sim & \text{month} + \text{month}^2 + \text{geography} + \text{month} \times \text{geography} \\ 210 \quad & + \text{month}^2 \times \text{geography} + 1 | \text{study}(\text{replicate}) \end{aligned}$$

211

212 We then used backwards model selection to determine the significance of fixed effects based on
213 likelihood ratio tests. Final models were fit using restricted maximum likelihood to calculate
214 model estimates. We then performed two separate analyses using (1) environmental temperature
215 and (2) aboveground biomass to assess any remaining variability in the residuals from the
216 seasonal models. First, using just the subset of studies that either reported temperature or
217 biomass, we fit the final seasonal models again using restricted maximum likelihood to obtain
218 the conditioned residuals. Then, using these residuals, we fit linear models with *geography* as a
219 categorical factor, plus a linear term for either *temperature* or aboveground *biomass* as well as
220 the interaction between *geography* and *temperature/biomass* as fixed effects. We used
221 backwards model selection, comparing nested models with a series of ANOVAs. Finally, we
222 tested for net autotrophy ($\text{NCP} > 0$) using a one-tailed t-test. All analyses were performed in R
223 (version 3.6.2) (R Core Team 2019) with the following packages as needed: nlme (version
224 3.1.145) (Pinheiro et al. 2020), broom (version 0.7.3) (Robinson et al. 2020), and broom.mixed
225 (version 0.2.6) (Bolker and Robinson 2020).

226



227 To illustrate how water depth and residence time may mediate the effects of the carbon fluxes
228 associated with seagrass communities on bulk seawater pH for potential local OA mitigation, we
229 applied the range of hourly net carbon fluxes (NCP) covered in our synthesis to a simplified,
230 steady state box model developed by Koweeck et al. (2018). We use the hourly rate rather than
231 the full-day rate because we recognize that the effects of seagrass on seawater carbonate
232 chemistry will be intermittent and fluctuate over the daylight hours. We modeled the change in
233 dissolved inorganic carbon as a function of NCP as

$$234 \quad \Delta DIC = \frac{L}{\bar{u}\rho h} \times \text{NCP}$$

235 where L = the box length (m), \bar{u} = mean water velocity (m s^{-1}), ρ = is the seawater density (kg
236 m^3), and h = water depth (m). Because of the familiarity among managers and decision makers
237 with seawater pH, we then converted the delta DIC to pH, assuming a relevant, temperate coastal
238 ocean condition (e.g., total alkalinity = $2300 \mu\text{mol/kg}$, temperature = 15°C , and salinity = 35
239 ppm). We then plotted the change in pH as a function of hourly daytime carbon fluxes (i.e.,
240 hourly NCP) for two different water depths (0.5 and 2m) and three different water residence
241 times (L/\bar{u} = 15 minutes, 60 minutes, and 4 hours) at each water depth. We selected four hours
242 as the maximum duration for the model for two reasons: seagrass beds are rarely extensive
243 enough for water to remain over seagrass for more than a few hours, and longer residence times
244 would tend to overlap with lower-light conditions when the hourly NCP does not apply.

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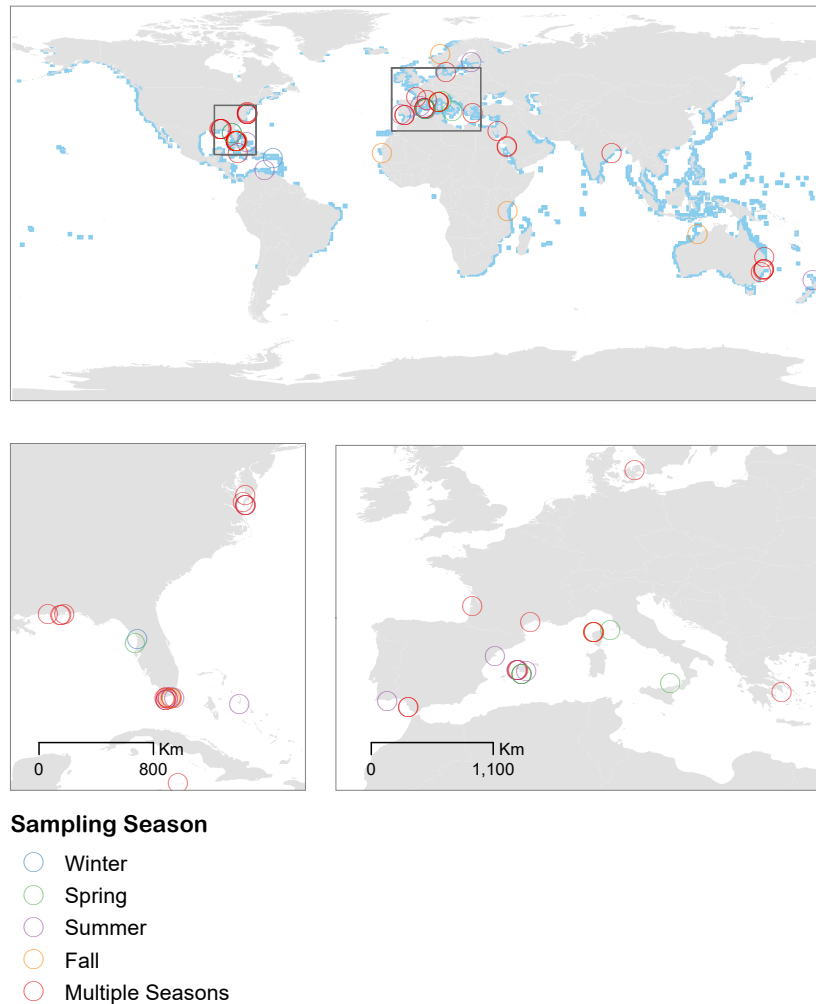
247 **3 Results**

248 *3.1 Description of the database*

249 Using our search criteria, we identified 62 published papers (spanning 1956 to 2020) that
250 reported *in situ* rates of seagrass community metabolism (Table S1). None of the studies directly



251 measured changes in DIC, but 17 of the 62 studies reported metabolic measurements in units of
252 carbon. The complete set of studies spanned temperate and tropical ecosystems. The inclusion of
253 36 temperate studies and an additional 15 tropical studies significantly expanded the scope of
254 inference beyond previous reviews (Duarte et al. 2010; Unsworth et al. 2012). Many studies
255 occurred in the Western Atlantic and Mediterranean (Fig. 1), and there were no studies from the
256 North Pacific. Most studies measured seagrass metabolism during the spring and summer
257 months, while fewer studies measured the metabolism in fall and winter conditions (Fig. 1).
258 Environmental temperature was highest during late summer/early fall months and was higher
259 overall in tropical biomes (Fig S3). Aboveground biomass was highest during summer months
260 and higher in the temperate geographies (Fig S4).
261



262

263 **Figure 1.** Studies included in the analyses span temperate to tropical ecosystems, with many
264 studies occurring in the Western Atlantic and Mediterranean. Most studies measured seagrass
265 metabolism during the spring and summer months, while fewer studies measured the metabolism
266 in fall and winter conditions. Blue pixels represent the distribution of seagrasses from UNEP-
267 WCMC and Short (2018).

268



269 *3.2 Methodological analyses*

270 Our results illustrate greater variability in the ranges of response observed using the “mass
271 balance” method, which extend in magnitude beyond those observed using other methods for
272 measuring both GPP and R (Fig. S1 a-f). This greater variability does not appear to be driven by
273 timing of the measurements as the “mass balance” method produced metabolic measurements of
274 higher variability or magnitude across seasons (Fig S1 g-l). Measurements taken using
275 incubations, eddy correlation methods, and water column measurements of pH using *in situ*
276 sensors and an acoustic doppler velocimeter (or other instrument capable of measuring flow) are
277 generally of similar magnitude and variation.

278

279 *3.3 Spatial and temporal patterns in carbon fluxes*

280 *3.3.1. Daytime carbon fluxes*

281 Measurements of hourly carbon fluxes (N=83 for NCP), typically obtained from shorter duration
282 deployments conducted during peak sunlight hours, reveal differences in seasonal patterns of
283 GPP and respiration. Both GPP and R peak during summer months across both ecosystems (Fig
284 2a-b). Despite higher biomass in temperate systems during summer months (Fig S4), we do not
285 detect a statistical difference in the seasonal patterns among GPP in temperate and tropical
286 ecosystems (Table 1). This result is highly influenced by two studies in tropical geographies
287 (Morgan and Kitting 1984, Herbert and Fourqurean 2008); when these studies are not included,
288 summertime GPP is higher in temperate geographies than in tropical geographies (Fig S5).
289 Similarly, R peaks in summer months in both temperate and tropical ecosystems, and we detect a
290 sharper increase and a higher seasonal peak in R in temperate ecosystems (Fig 2b; Table 1). The
291 seasonal peaks in GPP and R effectively cancel each other out, resulting in no statistically



292 detectable difference in hourly NCP rates across seasons (Fig 2c). Although the net hourly
 293 carbon flux associated with NCP does not vary seasonally, the mean net hourly carbon flux from
 294 the seawater to the seagrass is positive (mean hourly NCP = 5.32 +/- 5.93 SD mmol C/m²/hour),
 295 indicating a net draw down of seawater DIC during peak daylight hours regardless of geography
 296 (one-tailed t-test: $t_{82} = 8.18$, $P < 0.001$). Ninety-three percent of the 83 measurements were net
 297 autotrophic.

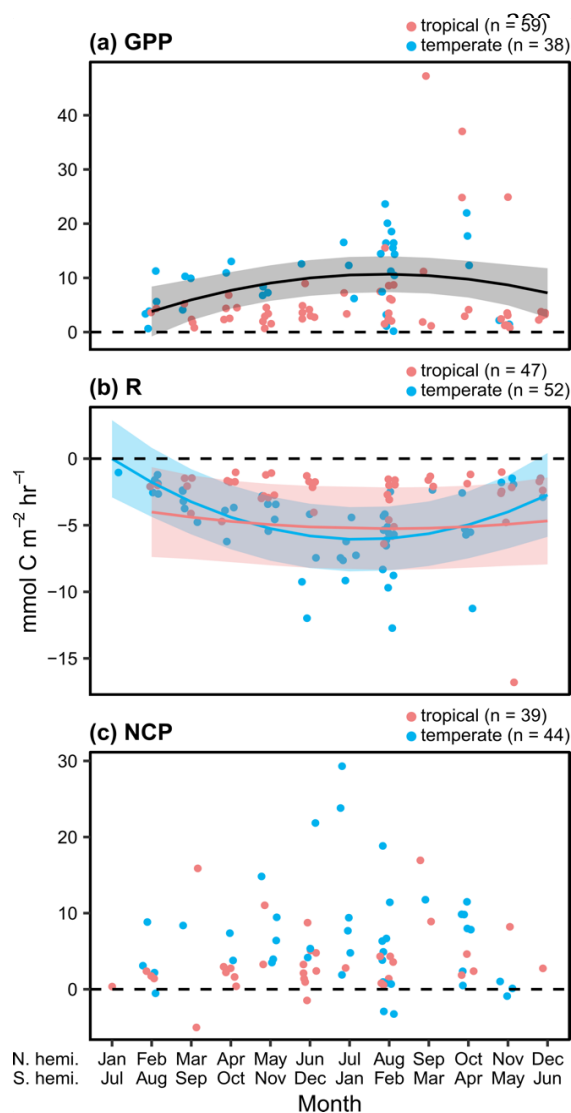


Figure 2. Hourly rates for C fluxes ($\text{mmol}^{-1}\text{m}^{-2}\text{hr}^{-1}$) associated with seagrass communities as a function of time, with GPP measurements taken during peak sunlight hours (~ 10 am to 2 pm local time). Studies performed in temperate versus tropical ecosystems are illustrated by color (blue = temperate, red = tropical), and significance ($p < 0.05$) is denoted by a fit line with a quadratic term and 95%CI. The colors of the lines denote significant differences between tropical and temperate systems in GPP and R, while a black line denotes a significant relationship but no difference between temperate and tropical geographies.



317 **Table 1.** Statistics for linear mixed effects models of hourly rates of carbon fluxes. Terms in grey
 318 were removed from the final model using backward model selection.

319

Response	Outliers removed	Fixed effect	df	L-ratio	p-value
GPP (mmol C/m ² /hr)	0	month			
		month²	5	6.76905	0.0093
		geography	6	0.73356	0.3917
		month × geography	7	0.00066	0.9795
	4	month ² × geography	8	0.32073	0.5712
		month			
		month²	7	13.7499	0.0002
		geography			
		month × geography	8	4.60594	0.0319
		month ² × geography	8	3.5689	0.0589
R (mmol C/m ² /hr)	0	month			
		month ²			
		geography			
		month × geography	8	5.98346	0.0144
		month² × geography	8	5.72857	0.0167
	1	month			
		month ²			
		geography			
		month × geography	8	9.23633	0.0024
		month² × geography	8	8.73371	0.0031
NCP (mmol C/m ² /hr)	0	month	4	0.04176	0.8381
		month ²	5	2.65064	0.1035
		geography	6	0.04703	0.8283
		month × geography	7	3.41018	0.0648
		month ² × geography	8	1.09949	0.2944

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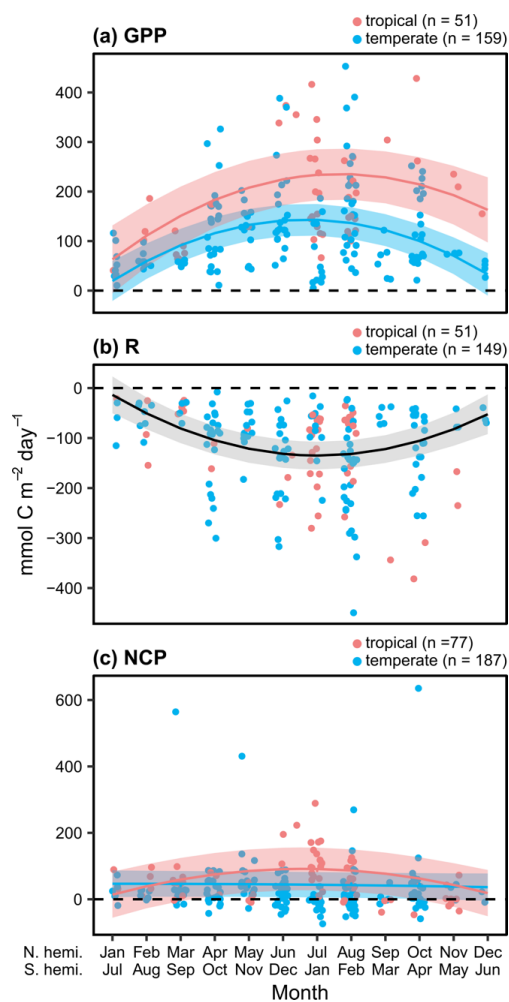
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325 3.3.2 Full-day carbon fluxes

326 We found 164 measurements/deployments that reported full-day NCP using methods that span a
327 wider range of photoperiods and environmental conditions, and thus provide insight into the
328 potential for full-day local OA mitigation. Based on the accompanying daily rates of GPP and R,
329 there is evidence of a seasonal cycle in carbon fluxes to and from the water column associated
330 with seagrass metabolism (Fig. 3a-b). The seasonal fluctuation differed statistically between
331 temperate and tropical geographies, with a sharper slope in the seasonal fluctuation among the
332 tropical studies (Fig 2a). We did not detect a difference in R between geographies. In general, the
333 seasonal fluctuation in GPP exceeds the seasonal fluctuation in respiration in both geographies,
334 resulting in higher daily net carbon flux from the seawater to the seagrass associated with NCP in
335 summer months (Fig. 3c). The seasonal fluctuation in NCP was greater among the tropical
336 studies than the temperate studies (Table 2). The mean NCP for tropical geographies was 62.5
337 (+/- 62.4 SD) mmol C/m²/day, with 84% of the 77 total reported measurements being
338 autotrophic. The mean NCP for temperate geographies was 28.8 (+/- 79.0) mmol C/m²/day, with
339 68% of the 187 total reported measurements being autotrophic. Overall, the seagrass meadows in
340 both geographies were net autotrophic (*one-tailed t-test*: tropical $t_{76}=8.78$, $P<0.001$; temperate
341 $t_{186} = 4.98$, $P<0.001$). Despite these overall trends, there are several individual studies that
342 reported net heterotrophy and net carbon fluxes to the water column, even during summer
343 months.



344

345 **Figure 3.** Daily rates for C fluxes ($\text{mmol}^{-1}\text{m}^{-2}\text{day}^{-1}$) associated with seagrass communities as a
346 function of time. Studies performed in temperate versus tropical ecosystems are illustrated by
347 color (blue = temperate, red = tropical), and significance ($p < 0.05$) is denoted by a fit line with a
348 quadratic term and 95%CI. The colors of the lines denote significant differences between tropical
349 and temperate systems.

350

351



352 **Table 2.** Statistics for linear mixed effects models of daily rates of carbon fluxes. Terms in grey
 353 were removed from the final model using backward model selection.

Response	Fixed effect	df	L-ratio	p-value
GPP (mmol C/m ² /day)	month			
	month²	7	54.5932	<0.0001
	geography			
	month × geography	7	4.28394	0.0385
	month ² × geography	8	0.05199	0.8196
R (mmol C/m ² /day)	month			
	month²	5	40.5399	<0.0001
	geography	6	2.7621	0.0965
	month × geography	7	2.2282	0.1355
	month ² × geography	8	2.04071	0.1531
NCP (mmol C/m ² /day)	month			
	month ²			
	geography			
	month × geography	8	16.9111	<0.0001
	month² × geography	8	15.1501	<0.0001

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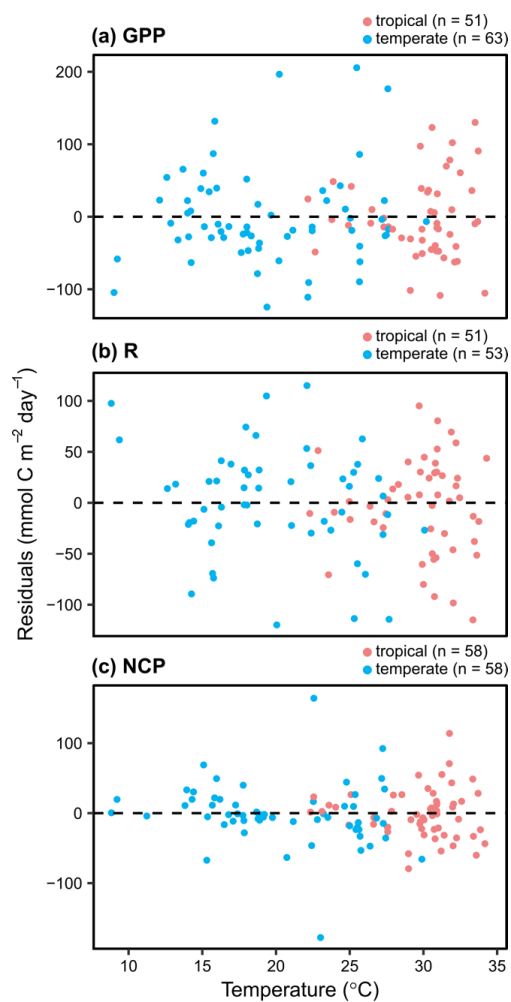
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365 *3.4 Drivers of chemical variability*

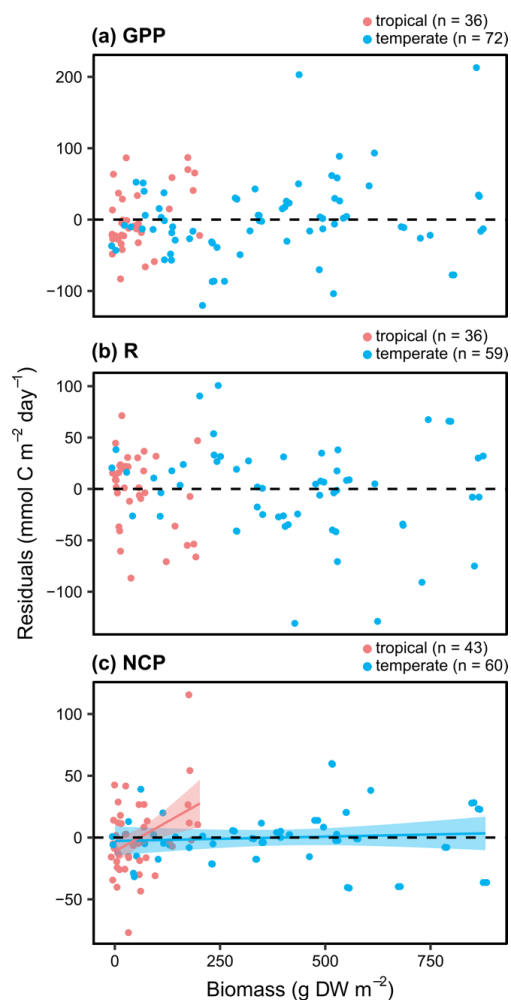
366 Within seasons, there is still marked variation in GPP and respiration (Fig. 2-3). Using the subset
367 of studies that report environmental temperature ($N = 28$), we found that temperature did not
368 explain the residual variability in any metric besides hourly GPP (Fig. 4; Table 3), suggesting the
369 seasonal models may generally account for hypothesized temperature effects. As noted,
370 temperature explained some of the residual variability from the seasonal models of hourly GPP,
371 with the effect differing among tropical and temperate geographies (Fig S6; Hourly GPP
372 *Geography x Temperature*: $F_{42} = 10.83$, $P = 0.001$). Among studies reporting aboveground
373 biomass ($N=23$), biomass explains some of the residual variability in daily NCP, although the
374 effect depends on geography as well (Fig 5; Table 3). Aboveground biomass also explains some
375 of the residual variability in the seasonal models of hourly GPP, respiration, and NCP, and the
376 effect of biomass on hourly GPP also depended on the geography (Fig. S7; Table 3).



377

378 **Figure 4.** Conditioned residuals of the daily rates of C fluxes ($\text{mmol}^{-1}\text{m}^{-2}\text{day}^{-1}$) from a seasonal
379 model as a function of temperature. Studies performed in temperate versus tropical ecosystems
380 are illustrated by color (blue = temperate, red = tropical). None of the relationships are
381 statistically significant.

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383

384 **Figure 5.** Conditioned residuals of the daily rates for seagrass C fluxes ($\text{mmol}^{-1}\text{m}^{-2}\text{day}^{-1}$) from a
385 seasonal model as a function of aboveground biomass measured in the field during metabolism
386 measurements. Studies performed in temperate versus tropical ecosystems are illustrated by color
387 (blue = temperate, red = tropical). Significance is denoted by a fit line and 95% CI.

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391 **Table 3.** Statistics for mixed effects models of the residuals of the hourly metabolic rates as a
 392 function of biomass or temperature and geography and the interactions.

Response	Fixed effect	df _{res}	SS _{res}	SS	F	p-value
GPP (mmol C/m ² /hr)	biomass					
	geography					
	biomass × geography	33	2806.1	-1081.7	20.074	<0.0001
R (mmol C/m ² /hr)	biomass	30	175.9	-20.68	3.864	0.059
	geography	30	161.9	-6.681	1.248	0.273
	biomass × geography	29	155.3	-1.505	0.274	0.605
NCP (mmol C/m ² /hr)	biomass	38	1640.2	-186.1	4.736	0.036
	geography	38	1721.4	-267.4	6.803	0.013
	biomass × geography	37	1454.1	-140.3	3.844	0.058
GPP (mmol C/m ² /hr)	temp					
	geography					
	temp × geography	42	375.4	-71.84	9.704	0.003
R (mmol C/m ² /hr)	temp	61	221.7	-0.930	0.253	0.617
	geography	60	220.8	-0.104	0.028	0.868
	temp × geography	59	220.7	-0.004	0.001	0.976
NCP (mmol C/m ² /hr)	temp	46	1453.0	-41.82	1.334	0.254
	geography	47	2141.0	-688.0	21.78	<0.0001
	temp × geography	45	1411.2	-12.30	0.387	0.537

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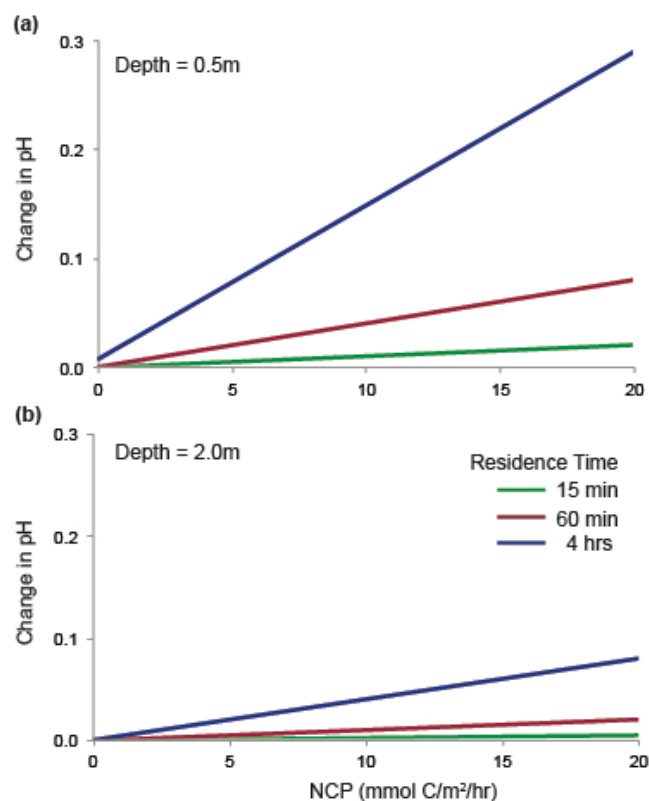
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402 *3.5 Potential OA amelioration*

403 The steady state box model illustrates that the largest potential change in seawater pH occurs
404 when NCP is highest and the water depth and residence time are lowest (Fig. 6). Using the mean
405 hourly NCP from our analysis (~ 5.32 mmol C/m²/hour; Fig. 2C), the potential change in
406 seawater pH in a seagrass meadow that meets the assumptions of the box model (e.g., $\Delta O_2 : \Delta DIC$
407 = 1) in a low flow environment at low tide (i.e., 0.5m water depth) ranges from 0.006 – 0.085 pH
408 units for a residence time from 15 minutes to 4 hours. At the modeled higher tide (i.e., 2m water
409 depth), the potential changes in seawater pH for the same NCP range from 0.001-0.022 pH units
410 for the same residence times (15 minutes to 4 hours).

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416 **Figure 6.** Results from a steady state box model illustrating the change in seawater pH as a
417 function of NCP at (A) 0.5 m water depth and (B) 2 m water depth. The change in seawater pH is
418 illustrated for a series of different residence times (L/u). Both panels are modeled at the same
419 temperature (15°C), total alkalinity ($2300 \mu\text{mol/kg}$), and salinity (35 ppm), with an incoming
420 DIC of $2100 \mu\text{mol/kg}$.

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426 **4 Discussion**

427 *4.1 Spatial and temporal variability in seagrass metabolism*

428 Here, we report that the NCP of seagrass beds during daylight hours is positive and similar
429 across seasons and geographies (Fig 2C). This is ostensibly due to GPP generally exceeding R
430 during daylight and the seasonal fluctuations in hourly rates of GPP and R balancing each other
431 out. If O₂ fluxes translate proportionally to community drawdowns in CO₂, the assumption
432 underlying our box model, our results suggest that the maximum potential local OA mitigation
433 due to seagrass metabolism during daylight hours is similar across time and ecosystems, but
434 small in magnitude.

435

436 We also demonstrate that seagrass beds are generally net autotrophic over the length of the day
437 (based on daily NCP), and the magnitude of this full-day NCP is more pronounced during
438 summer months and in tropical geographies (Fig 3C). However, underlying the summertime
439 peak in full-day NCP is the potential for marked diurnal variability in pH. In particular, the
440 demonstrated summertime peak in hourly respiration rates could drive more pronounced
441 nighttime lows in pH/saturation state during the most pronounced windows for net autotrophy
442 (i.e., summer months with the highest daily NCP). However, diurnal fluctuations in seawater pH
443 associated with seagrass metabolism will also be influenced by hydrodynamics that are not
444 captured in our synthesis (Koweek et al. 2018), and recent field studies of estuarine eelgrass by
445 Ricart et al. (2021) demonstrated that sustained elevations in pH associated with seagrass
446 meadows were not restricted to daylight hours. Regardless, a critical next step for understanding
447 the utility of seagrass beds for climate adaptation is to determine how the demonstrated seasonal



448 variability and the potential diurnal variability in carbonate chemistry are integrated over time by
449 organisms, especially under conditions projected for coastal habitats in the future.

450

451 Given the positive correlations between GPP and temperature, we assume that most of these
452 measurements were taken when the environmental temperatures were not physiologically
453 stressful to the seagrass community. Thus, the positive relationship between environmental
454 temperature and GPP should not necessarily be viewed through the lens of temperature
455 exposures associated with future warming driven by climate change. Continued warming
456 associated with climate change could cause photosynthesis and respiration to decline at stressful
457 temperatures. The probability that the relationship between temperature and metabolic responses
458 will change with future warming is likely to differ geographically based on how close a
459 community is to its thermal limit and the scope for acclimation or adaptation. This may partially
460 explain the differences in the relationships between temperature and the residuals of the seasonal
461 model between geographies, with increasing temperature negatively related to the residuals from
462 tropical geographies and positively related to the residuals from temperate geographies. Many
463 tropical seagrass species are growing close to their photosynthetic and physiological optima (Lee
464 et al. 2007, Koch et al. 2012), and elevated temperatures in these geographies may be
465 detrimental, causing metabolism (in this case hourly GPP) to be less than expected based on
466 seasonal patterns. In contrast, the positive relationship between the residuals from the seasonal
467 model and temperature in temperate geographies indicate that there may be a stimulatory effect
468 before a thermal tolerance threshold is crossed and metabolism decreases.

469



470 The significant relationships between aboveground biomass and the residual variation in hourly
471 GPP and NCP indicate that aboveground biomass also plays an important role in carbon fluxes
472 beyond that which is already captured by any seasonal fluctuations in biomass. Although the
473 relationships between aboveground biomass and the residual variation in hourly GPP are
474 generally what would be expected (higher biomass = higher metabolic rates than expected based
475 on the seasonal model), the negative relationship between aboveground biomass and the residual
476 variation in hourly NCP is somewhat surprising. This relationship suggests that
477 deployments/measurements in seagrass beds with higher aboveground biomass generally had
478 lower hourly NCP than what is predicted by the seasonal model. This negative relationship may
479 be explained by self-shading in dense meadows, or it could be due to other organisms that
480 contribute to daytime respiration (e.g., heterotrophs) that are associated with the higher biomass
481 meadows due to its structural complexity or other habitat features, but are not accounted for in
482 the aboveground biomass measurements. Dedicated experiments may be able to determine the
483 mechanism for these findings; however, the positive relationships between aboveground biomass
484 and the residual variation in daily NCP suggests that, overall, higher aboveground biomass
485 generally increases production relative to respiration.

486

487 *4.2 Implications for local OA mitigation and management*

488 The results of our steady state box model analyses illustrate the potential scope for seagrass NCP
489 to influence seawater pH on an hourly basis (Fig. 6), with the change in pH being proportional to
490 NCP during daylight hours and R during nighttime hours. While the box model is useful in
491 making coarse estimates on what particular NCP values might correspond to in seawater pH, it is
492 important to note that it only represents a first step in translating the seagrass community



493 metabolism estimates to seawater biogeochemistry. This is in part because the ratio of NCP
494 based on carbon fixed and oxygen evolved in seagrass communities is likely to be quite variable.
495 Although the ratio between O₂ produced and carbon fixed by an individual seagrass is generally
496 assumed to be balanced (i.e., 1:1), the other processes that occur in a seagrass meadow, including
497 respiration from organisms living within the seagrass and carbonate production and dissolution,
498 also influence the dissolved inorganic carbon (DIC) concentration in the seawater. Current
499 empirical measurements of NCP_{DIC}:NCP_{O₂} in seagrass meadows range from 0.3 to 6.8 (Ziegler
500 and Benner 1998, Barrón et al. 2006), suggesting the effect of seagrass NCP on seawater pH
501 could be substantially more or less pronounced than illustrated here. Because of this variability
502 in the relationship between O₂ and DIC, care must be taken when interpreting the results from
503 the box model. A better understanding of the NCP_{DIC}:NCP_{O₂} in particular meadows will better
504 inform their potential for local OA mitigation. Finally, the utility of seagrass as a climate
505 mitigation tool will depend on the goal of the management, and in most cases, will require more
506 research. For example, if the goal of management is to prevent negative effects of ocean
507 acidification on oyster growth, then studies that quantify the sensitivity of oyster growth to the
508 variability in pH observed here are still required.

509

510 **5 Conclusions**

511 Few conservation or restoration efforts currently take into account the potential chemical
512 ecosystem services of seagrasses and other submerged aquatic vegetation. Here, we demonstrate
513 that daytime carbon fluxes associated with seagrass metabolism are likely to be similar across
514 seasons and geography, while the full-day carbon fluxes peak during summer months in both
515 tropical and temperate geographies. Integrating across seasons, seagrass meadows are net



516 autotrophic. However, our simplified model results suggest the daytime carbon fluxes reported
517 across the global ocean may translate to small changes in seawater pH. These seasonal patterns
518 largely capture the present-day effects of variability in temperature and aboveground biomass on
519 seagrass metabolism, but likely do not adequately model the effects of future warming as it
520 becomes physiologically stressful.

521

522 These results highlight challenges, as well as gaps in our understanding, that may impede the use
523 of seagrasses for sustained local OA mitigation. In particular, we demonstrate that while peak
524 daytime carbon fluxes are similar across seasons and geographies, nighttime respiration is
525 highest during summer months. Thus, although seagrass beds are generally net autotrophic,
526 nighttime respiration could reduce seawater pH during periods of greatest autotrophy. We
527 provide examples of how water depth and residence time can influence the effect of seagrass on
528 seawater pH, and we demonstrate that the overall magnitude of the effect is likely quite small.

529

530 This work has elucidated several gaps that need to be addressed by the scientific community. For
531 example, certain geographies, such as the North Pacific, are currently underrepresented in our
532 dataset. Thus, continued study of seagrass metabolism and its effects on seawater carbonate
533 chemistry are needed to expand our area of inference. In addition, studies are needed to constrain
534 the relationship between dissolved oxygen fluxes and DIC, and this relationship may be
535 important to elucidate at local scales to truly understand the potential for OA mitigation at a
536 given location. Perhaps most importantly, more information is needed to understand how
537 vulnerable organisms respond to the chemical variability highlighted in our study (Gimenez et al.
538 2018, Lowe et al. 2018), and in particular, how this variability is integrated through time. Despite



539 the considerations of geographic and temporal variability in carbon fluxes illustrated here, we
540 recognize that seagrass conservation and restoration may be important strategies for climate
541 adaptation for numerous other reasons, including carbon sequestration and habitat provisioning.

542

543 **6 Code Availability:** Code is available at

544 <https://github.com/tyekindinger/SeagrassCommunityMetabolism>

545

546 **7 Data Availability:** All data used in this analysis is publicly available via the published studies.

547

548 **8 Author Contributions:** Kroecker: conceptualization, funding acquisition, investigation,
549 methodology, writing – original draft preparation; Kindinger: formal analysis, visualization,
550 writing – review & editing; Hirsh: investigation, visualization, writing – review & editing; Ward:
551 investigation, writing – review & editing; Hill: conceptualization, funding acquisition, writing –
552 review & editing; Jellison, Koweek, Lummis, Rivest, Waldbusser: conceptualization, writing –
553 review & editing; Gaylord: conceptualization, funding acquisition, methodology, writing –
554 review & editing.

555

556 **9 Competing Interests:** The authors declare no competing interests

557

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569

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