

N₂ fixation in the Mediterranean Sea related to the composition of the diazotrophic community, and impact of dust under present and future environmental conditions

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Abstract. N₂ fixation rates were measured in the 0-1000 m layer at 13 stations located in the open western and central Mediterranean Sea (MS) during the PEACETIME cruise (late spring 2017). While the spatial variability of N₂ fixation was not related to Fe, P nor N stocks, the surface composition of the diazotrophic community indicated a strong **longitudinal gradient increasing eastward** for the relative abundance of non-cyanobacterial diazotrophs (NCD) (mainly γ -Proteobacteria) and conversely **eastward decreasing eastward** for UCYN-A (mainly -A1 and -A3) as did N₂ fixation rates. UCYN-A4 and -A3 were identified for the first time in the MS. The westernmost station influenced by Atlantic waters, and characterized by highest stocks of N and P, displayed a patchy distribution of diazotrophic activity with an exceptionally high rate in the euphotic layer of 72.1 nmol N L⁻¹ d⁻¹, which could support up to 19 % of primary production. At this station at 1%PAR depth, UCYN-A4 represented up to 94 % of the diazotrophic community. These *in situ* observations of **greater relative abundance of UCYN-A at stations with higher nutrient concentrations and dominance of NCD at more oligotrophic stations** suggest that nutrient conditions - **even in the nanomolar range -may determine the composition of diazotrophic communities and in turn N₂ fixation rates.** The impact of Saharan dust deposition on N₂ fixation and diazotrophic communities was also investigated, under present and future projected conditions of temperature and pH during

35 short term (3-4 days) experiments at three stations. New nutrients from simulated dust deposition triggered a significant
36 stimulation of N₂ fixation (from 41 % to 565 %). The strongest increase in N₂ fixation was observed at the stations
37 dominated by NCD and did not lead on this short time scale to change in the diazotrophic community composition. Under
38 projected future conditions, N₂ fixation was either **increased** or unchanged; in that later case this was probably due to a too
39 low nutrient bioavailability or an increased grazing pressure. The future warming and acidification likely benefited NCD
40 (*Pseudomonas*) and UCYN-A2 while disadvantaged UCYN-A3 without knowing which effect (alone or in combination) is
41 the driver, especially since we do not know the temperature optima of these species not yet cultivated as well as the effect of
42 acidification.

43 **1. Introduction**

44 The Mediterranean Sea (MS) is considered as one of the most oligotrophic regions of the world's ocean (Krom et al., 2004;
45 Bosc et al., 2004). It is characterized by a longitudinal gradient in nutrient availability, phytoplanktonic biomass and primary
46 production (PP) decreasing eastward (Manca et al., 2004; D'Ortenzio and Ribera d'Alcalà, 2009; Ignatiades et al., 2009;
47 Siokou-Frangou et al., 2010; El Hourany et al., 2019). From May to October, the upper water column is well-stratified
48 (D'Ortenzio et al., 2005), and the sea surface mixed layer (SML) becomes nutrient-depleted leading to low PP (e.g. Lazzari
49 et al., 2012). Most measurements of N₂ fixation during the stratified period have shown low rates ($\leq 0.5 \text{ nmol N L}^{-1} \text{ d}^{-1}$) in
50 surface waters of the open MS (Ibello et al., 2010; Bonnet et al., 2011; Yogeve et al., 2011; Ridame et al., 2011; Rahav et al.,
51 2013a; Benavides et al., 2016) indicating that N₂ fixation represents a minor source of bioavailable nitrogen in the MS
52 (Krom et al., 2010; Bonnet et al., 2011). These low rates are likely related to the extremely low bioavailability in dissolved
53 inorganic phosphorus (DIP) (Rees et al., 2006; Ridame et al., 2011). The high concentrations of dissolved iron (DFe) in the
54 SML due to accumulated atmospheric Fe deposition (Bonnet and Guieu 2006; Tovar-Sánchez et al. 2020; Bressac et al.,
55 2021), suggest that the bioavailability of Fe is not a controlling factor of N₂ fixation (Ridame et al., 2011). Occasionally,
56 high N₂ fixation rates have been reported locally in the northwestern ($17 \text{ nmol N L}^{-1} \text{ d}^{-1}$; Garcia et al., 2006) and eastern MS
57 ($129 \text{ nmol N L}^{-1} \text{ d}^{-1}$; Rees et al., 2006). Usually, the low N₂ fixation rates in the Mediterranean offshore waters are associated
58 with low abundance of diazotrophs, mainly dominated by unicellular organisms (Man-Aharonovich et al., 2007; Yogeve et
59 al., 2011; Le Moal et al., 2011). Unicellular diazotrophs from the photo-heterotrophic group A (UCYN-A, Zehr et al., 1998)
60 largely dominated the cyanobacteria assemblage in the MS (Le Moal et al., 2011), and very low concentrations of
61 filamentous diazotrophic cyanobacteria have only been recorded in the eastern basin (Bar-Zeev et al., 2008; Le Moal et al.,
62 2011; Yogeve et al., 2011). The UCYN-A cluster consist of four sublineages: UCYN-A1, -A2, -A3 and -A4 (Thompson et al.,
63 2014; Farnelid et al., 2016; Turk Kubo et al., 2017; Cornejo-Castillo et al., 2019), of which only UCYN-A1 and -A2 have
64 been previously detected in the MS (Man-Aharonovich et al., 2007; Martinez-Perez et al., 2016; Pierrela Karlusich et al.,
65 2021). Heterotrophic diazotrophs are widely distributed over the offshore surface waters (Le Moal et al., 2011), and the

66 decreasing eastward gradient of surface N₂ fixation rate could be related to a predominance of photo-autotrophic diazotrophs
67 in the western basin and a predominance of heterotrophic diazotrophs in the eastern one (Rahav et al. 2013a).
68 The MS is strongly impacted by periodic dust events, originating from the Sahara, which have been recognized as a
69 significant source of macro- and micronutrients, to the nutrient depleted SML during stratified periods (Guieu and Ridame,
70 2020 and references therein; Mas et al., 2020). Results from Saharan dust seeding experiments during open sea microcosms
71 and coastal mesocosms in the MS, showed stimulation of both PP (Herut et al., 2005; Ternon et al., 2011; Ridame et al.,
72 2014 ; Herut et al., 2016) and heterotrophic bacterial production (BP) (Pulido-Villena et al., 2008, 2014; Lekunberri et al.,
73 2010; Herut et al., 2016). Experimental Saharan dust seeding was also shown to **enhance** N₂ fixation in the western and
74 eastern MS (Ridame et al., 2011; Ternon et al., 2011; Ridame et al., 2013; Rahav et al., 2016a) and to alter the composition
75 of the diazotrophic community (Rahav et al., 2016a), as also shown in the tropical North Atlantic (Langlois et al., 2012).
76 The MS has been identified as one of the primary hot-spots for climate change (Giorgi, 2006). Future sea surface warming
77 and associated increase in stratification (Somot et al., 2008) might reinforce the importance of atmospheric inputs as a source
78 of new nutrients for biological activities **during that season, including** diazotrophic microorganisms **during the stratified**
79 **period**. This fertilizing effect could also be enhanced by the expected decline in pH (Mermex Group, 2011), which could
80 increase the nutrient dust solubility in seawater. Under **nutrients** depleted conditions, predicted elevated temperature and
81 CO₂ concentration favor the growth and N₂ fixation of the filamentous cyanobacteria *Trichodesmium* and of the photo-
82 autotrophic UCYN-B and -C (Webb et al., 2008; Hutchins et al., 2013; Fu et al., 2008, 2014; Eichner et al., 2014; Jiang et
83 al., 2018), whereas effects on UCYN-A and non-cyanobacterial diazotrophs (NCD) are uncertain.
84 In this context, the first objective of this study is to investigate during the season characterized by strong stratification and
85 low productivity, the spatial variability of N₂ fixation rates in relation to **nutrients** availability and diazotrophic **community**
86 composition. The second objective was to study, for the first time, the impact of a realistic Saharan deposition event in the
87 open MS, on N₂ fixation rates and diazotrophic communities composition under present and realistic projected conditions of
88 temperature and pH for 2100.

89

90 **2. Materials and Methods**

91 **2.1 Oceanographic cruise**

92 All data were acquired during the PEACETIME cruise (ProcEss studies at the Air-sEa Interface after dust deposition in the
93 MEditerranean sea) in the western and central MS on board the R/V *Pourquoi Pas ?* from May 10 to June 11, 2017
94 (<http://peacetime-project.org/>) (see the detailed description in Guieu et al., 2020). The cruise track including ten short
95 stations (ST1 to ST10) and three long stations (TYR, ION and FAST) is shown in Fig.1 (coordinates in Table S1). Stations 1
96 and 2 were located in the Liguro-Provencal basin; Stations 5, 6, and TYR, in the Tyrrhenian Sea; Stations 7, 8, and ION in
97 the Ionian Sea; and Stations 3, 4, 9, 10 and FAST in the Algerian basin.

98

99 **2.2 Dust seeding experiments**

100 Experimental dust seedings into six large tanks were conducted at each of the three long stations (TYR, ION and FAST),
101 under present and future conditions of temperature and pH. **Based on previous studies, the location of these stations was**
102 **chosen based on several criteria including because they represent three main bioregions of the MS (Guieu et al., 2020,**
103 **their Fig. S1). They are located along the longitudinal gradient in biological activity, including the activity of**
104 **diazotrophs decreasing eastward (Bonnet et al., 2011; Rahav et al., 2013a).** The experimental setup is fully described in
105 a companion paper (Gazeau et al., 2021a). Briefly, six climate reactors (volume of about 300 L) made in high density
106 polyethylene were placed in a temperature-controlled container, and covered with a lid equipped with LEDs to reproduce
107 natural light cycle. The tanks were filled with unfiltered surface seawater collected at ~5m with a peristaltic pump at the end
108 of the day (T-12h) before the start of the experiments the next morning (T0). Two replicate tanks were amended with
109 mineral Saharan dust (Dust treatments D1 and D2) simulating a high but realistic atmospheric dust deposition of 10 g m^{-2}
110 (Guieu et al., 2010b). Two other tanks were also amended with Saharan dust (same dust flux as in the Dust treatment) under
111 warmer ($\sim +3^\circ \text{ C}$) and more acidic water conditions ($\sim -0.3 \text{ pH unit}$) (Greenhouse treatments G1 and G2). This corresponds
112 to the IPCC projections for 2100 under RCP8.5 (IPCC 2019). Seawater in G1 and G2 was warmed overnight to reach $+3^\circ \text{ C}$
113 and acidified through the addition of CO_2 -saturated $0.2 \text{ }\mu\text{m}$ -filtered seawater ($\sim 1.5 \text{ L}$ in 300 L). The difference in
114 temperature between G (Greenhouse) tanks and other tanks (C, Controls and D, Dust) was $+3^\circ \text{ C}$, $+3.2^\circ \text{ C}$ and $+3.6^\circ \text{ C}$ at
115 TYR, ION and FAST, respectively, and the decrease in pH was -0.31 , -0.29 and -0.33 at TYR, ION and FAST, respectively
116 (Gazeau et al., 2021a). Two tanks were filled with untreated water (Controls C1 and C2). The experiment at TYR and ION
117 lasted three days while the experiment at FAST lasted four days. The sampling session took place every morning at the same
118 time over the duration of the experiments.

119 The fine fraction ($< 20 \text{ }\mu\text{m}$) of a Saharan soil collected in southern Tunisia used in this study, has been previously used for
120 the seeding of mesocosms in the frame of the DUNE project (a DUSt experiment in a low-Nutrient, low-chlorophyll
121 Ecosystem). Briefly, the dust was previously subjected to physico-chemical transformations mimicking the mixing between
122 dust and pollution air masses during atmospheric transport (see details in Desboeufs et al, 2001; Guieu et al., 2010b). This
123 dust contained $0.055 \pm 0.003 \%$ of P, $1.36 \pm 0.09 \%$ of N, and $2.26 \pm 0.03 \%$ of Fe, in weight (Desboeufs et al., 2014). Right
124 before the artificial seeding, the dust was mixed with 2 L of ultrapure water in order to mimic a wet deposition event and
125 sprayed at the surface of the climate reactors D and G. **The succession of operations is fully described in Gazeau et al.,**
126 **(2021a, see their Table 1).**

127

128 **2.3 N₂ fixation and primary production**

129 All materials were acid washed (HCl Suprapur **32%**) following trace metal clean procedures. Before sampling, bottles were
130 rinsed three times with the sampled seawater. For the *in situ* measurements, seawater was sampled using a trace metal clean
131 (TMC) rosette equipped with 24 GO-FLO Bottles (Guieu et al., 2020). At each station, 7 to 9 depths were sampled between
132 surface and 1000 m for N₂ fixation measurements, and 5 depths between surface and $\sim 100 \text{ m}$ for primary production
133 measurements (**one sample par depth**). During the seeding experiments, the six tanks were sampled for simultaneous

134 determination of N₂- and CO₂ net fixation rates before dust seeding (initial time T0) and one day (T1), two days (T2), and
135 three days (T3) after dust addition at TYR and ION Stations. At FAST, the last sampling took place four days (T4) after dust
136 addition.

137 **After collection, 2.3 L of seawater were immediately filtered onto pre-combusted GFF filters to determine natural**
138 **concentrations and isotopic signatures of particulate organic carbon (POC) and particulate nitrogen (PN).** Net N₂
139 fixation rates were determined using the ¹⁵N₂ gas-tracer addition method (Montoya et al., 1996), and net primary production
140 using the ¹³C-tracer addition method (Hama et al., 1983). Immediately after sampling, 1 mL of NaH¹³CO₃ (99 %, Eurisotop)
141 and 2.5 ml of 99 % ¹⁵N₂ (Eurisotop) were introduced to 2.3 L polycarbonate bottles through a butyl septum for simultaneous
142 determination of N₂- and CO₂-fixation. ¹⁵N₂ and ¹³C tracers were added to obtain a ~10 % final enrichment. Then, each bottle
143 was vigorously shaken before incubation for 24 h. The *in situ* samples from the euphotic zone were incubated in on-deck
144 containers with circulating seawater, equipped with blue filters **with different sets of blue neutral density filters (Lee**
145 **Filters)** (percentages of attenuation: 70, 52, 38, 25, 14, 7, 4, 2 and 1 %) to simulate **an irradiance level (% PAR) as close**
146 **as possible to the one corresponding to their depth of origin.** Samples for N₂ fixation determination in the aphotic layer
147 were incubated in the dark in thermostated incubators set at *in situ* temperature. *In situ* ¹³C-PP will not be discussed in this
148 paper as ¹⁴C-PP rates are presented in Maranon et al. (2021) (see details in Fig. S1). The *in situ* ¹³C-PP **and molar C/N**
149 **ratio in the organic particulate matter, measured simultaneously in our samples (see below for details),** were used to
150 estimate the contribution of N₂ fixation to PP.

151 Samples from the dust addition experiments were incubated in two tanks dedicated to incubation: one tank at the same
152 temperature and irradiance as tanks C and D, and another one at the same temperature and irradiance as tanks G. It should be
153 noted that ¹⁴C-PP was also measured during the seedings experiments (Gazeau et al., 2021b).

154 After 24 h incubation, 2.3 L were filtered onto pre-combusted 25 mm GF/F filters, and filters were stored at -25° C. Filters
155 were then dried at 40° C for 48 h before analysis. **POC and PN** as well as ¹⁵N and ¹³C isotopic ratios were quantified using
156 an online continuous flow elemental analyzer (Flash 2000 HT), coupled with an Isotopic Ratio Mass Spectrometer (Delta V
157 Advantage via a conflow IV interface from Thermo Fischer Scientific). **For each sample, POC (in the 0-100m layer) and**
158 **PN (0-1000m) were higher than the analytically determined detection limit of 0.15 μmol for C and 0.11 μmol for N.**
159 **Standard deviations were 0.0007 atom% and 0.0005 atom% for ¹³C and ¹⁵N enrichment, respectively. The atom%**
160 **excess of the dissolved inorganic carbon (DIC) was calculated by using measured DIC concentrations at the**
161 **LOCEAN laboratory (SNAPO-CO₂).** N₂ fixation rates were calculated by isotope mass balance equations as described by
162 Montoya et al. (1996). **For each sample, the ¹³C and ¹⁵N uptake rates were considered as significant when excess**
163 **enrichment of POC and PN was greater than three times the standard deviation obtained on natural samples.**
164 **According to our experimental conditions, the minimum detectable ¹³C and ¹⁵N uptake rates in our samples were 5**
165 **nmol C L⁻¹ d⁻¹ and 0.04 nmol N L⁻¹ d⁻¹, respectively. CO₂ uptake rates were above the detection limit in the upper 0-**
166 **100m, while N₂ fixation was not quantifiable below 300 m depth except at Stations 1 and 10 with rates ~0.05 nmol N**
167 **L⁻¹ d⁻¹ at 500 m depth.** N₂ fixation rates were calculated by isotope mass balanced as described by Montoya et al. (1996).

168 The detection limit for N₂ fixation, calculated from significant enrichment and lowest particulate nitrogen was 0.04 nmol N
169 L⁻¹ d⁻¹. From these measurements, the molar C:N ratio in the organic particulate matter was calculated and used to estimate
170 the contribution of N₂ fixation to primary production. As a rough estimate of the potential impact of bioavailable N input
171 from N₂ fixation on BP, **we used the BP rates presented in companion papers (Gazeau et al., 2021b; Van Wambeke et**
172 **al., 2021), and converted them** in N demand using the molar ratio C/N of **6.8 (Fukuda et al., 1998)**. Trapezoidal method
173 was used to calculate integrated rates over the SML, the euphotic layer (from surface to 1 % photosynthetically available
174 radiation (PAR) depth) and the 0-1000 m water column.

175 It must be noted that N₂ fixation rates measured by the ¹⁵N₂-tracer gas addition method may have been underestimated due to
176 incomplete ¹⁵N₂ gas bubble equilibration (Mohr et al., 2010). However, this potential underestimation is strongly lowered
177 during long incubation (24h).

178 The relative changes (RC, in %) in N₂ fixation in the dust experiments were calculated **as follows** :

$$RC (\%) = 100 \times \frac{(N_2FIXATION_{Tx} - N_2FIXATION_{Control})}{(N_2FIXATION_{Control})}$$

179 with N₂ Fixation_{Tx} the rate in D1, D2, G1 or G2 at Tx, N₂ Fixation_{Control} the mean of the duplicated controls (C1 and C2) at
180 Tx, and Tx the time of the sampling.

181

182 **2.4 Composition of the diazotrophic community**

183 Samples for characterization of the diazotrophic communities were collected during the dust seeding experiments in the six
184 tanks at initial time before seeding (T0) and final time (T3 at TYR and ION, and T4 at FAST). Three liters of water were
185 collected in acid-washed containers from each tank, filtered onto 0.2 μm PES filters (Sterivex) and stored at -80° C until
186 DNA extraction. The composition of the diazotrophic community was also determined at four depths (10, 61, 88 and 200 m)
187 at Station 10. Here, 2 L seawater were collected from the TMC rosette. Immediately after collection, seawater was filtered
188 under low vacuum pressure through a 0.2 μm-Nuclepore membrane and stored at -80° C in cryovials. Nucleic acids were
189 obtained from both filter types using phenol-chloroform extraction followed by purification (NucleoSpin® PlantII kit;
190 Macherey-Nagel). DNA extracts were used as templates for PCR amplification of the nifH gene by nested PCR protocol as
191 fully described in Bigeard et al., (2021, protocol.io). Following polymerase chain reactions, DNA amplicons were purified,
192 and quantified using NanoQuant Plate™ and Tecan Spark® (Tecan Trading AG, Switzerland). Each PCR product was
193 normalized to 30ng/μl in final 50μl and sent to Genotoul (<https://www.genotoul.fr/>, Toulouse, France) for high throughput
194 sequencing using paired-end 2x250bp Illumina MiSeq. All reads were processed using the Quantitative Insight Into
195 Microbial Ecology 2 pipeline (QIIME2 v2020.2, Bolyen et al., 2019). Reads were truncated to 350 bp based on sequencing
196 quality, denoised, merged and chimera-checked using DADA2 (Callahan et al., 2016). A total of 1,029,778 reads were
197 assigned to 635 amplicon sequence variants (ASVs). The table was rarefied by filtering at 1 % relative abundance per sample
198 cut-off that reduced the dataset to 97 ASVs accounting for 98.27 % of all reads. Filtering for homologous genes was done
199 using the NifMAP pipeline (Angel et al., 2018) and translation into amino acids using FrameBot (Wang et al., 2013). This

200 yielded 235 ASVs accounting for 1,022,184 reads (99 %). These remaining ASVs were classified with DIAMOND blastp
201 (Buchfink et al 2015) using a FrameBot translated nifH database (phylum level version; Moynihan, 2020) based on the ARB
202 database from the Zehr Lab (version June 2017; <https://www.jzehrlab.com/nifh>). NifH cluster and subcluster designations
203 were assigned according to Frank et al. (2016). UCYN-A sublineages were assigned by comparison to UCYN-A reference
204 sequences (Farnelid et al., 2016; Turk-Kubo et al., 2017). All sequences associated with this study have been deposited
205 under the BioProject ID: PRJNA693966. Alpha and beta-diversity indices for community composition, were estimated after
206 randomized subsampling. Analyses were run in QIIME 2 and in Primer v.6 software package (Clarke and Warwick, 2001).

207

208 **2.5 Complementary data from PEACETIME companions papers**

209 **Bacterial production-** Heterotrophic bacterial production (BP, *sensus stricto* referring to prokaryotic heterotrophic
210 production) was determined on board using the microcentrifuge method with the ³H- leucine (³H-Leu) incorporation
211 technique to measure protein production (Smith and Azam, 1992). The detailed protocol and the rates of BP are presented in
212 Van Wambeke et al. (2021) for measurements in the water column, and in Gazeau et al. (2021b) for measurements over the
213 course of the dust seeding experiments.

214 **Dissolved Fe-** Dissolved iron (DFe) concentrations (< 0.2 μm) were measured by flow injection analysis with online
215 preconcentration and chemiluminescence detection (FIA-CL). The detection limit was 15 pM (Bressac et al., 2021). DFe
216 concentrations in the water column along the whole transect are presented in Bressac et al. (2021) and for the dust seeding
217 experiments in Roy-Barman et al. (2021).

218 **Dissolved inorganic phosphorus and nitrate-** Concentrations of DIP and nitrate (NO₃⁻) were analyzed immediately after
219 collection on 0.2 μm filtered seawater using a segmented flow analyzer (AAIII HR Seal Analytical) according to Aminot and
220 K  rouel (2007) with respective detection limits of 0.02 μmol L⁻¹ and 0.05 μmol L⁻¹. Samples with concentrations below the
221 limit of detection with standard analysis were analyzed by spectrophotometry using a 2.5 m long waveguide capillary cell
222 (LWCC) for DIP (Pulido-Villena et al., 2010) and a 1 m LWCC for NO₃⁻ (Louis et al., 2015); the limit of detection was 1
223 nM for DIP and 6 nM for NO₃⁻. Samples for determination of NO₃⁻ at nanomolar level were lost from Stations 1 to 4. The
224 dust addition experiments data are detailed in Gazeau et al. (2021a). The water column data are fully discussed in Pulido-
225 Villena et al. (2021) and Van Wambeke et al. (2021).

226

227 **2.6 Statistical analysis**

228 Pearson's correlation coefficient was used to test the statistical linear relationship ($p < 0.05$) between N₂ fixation and other
229 variables (BP, PP, DFe, DIP, NO₃⁻); **it should be noted that the DIN stocks estimated at Stations 1 to 4 (Table S1) were**
230 **excluded from statistical analysis.** In the dust seeding experiments, means at initial time (T0) before dust amendment
231 (average at T0 in C and D treatments; n = 4, see Table 2) were compared using a one-way ANOVA followed by a Tukey
232 means comparison test ($\alpha = 0.05$). When assumptions for ANOVA were not respected, means were compared using a
233 Kruskal–Wallis test and a post hoc Dunn test. To test significant differences ($p < 0.05$) between the slopes of N₂ fixation as a

234 function of time in the C, D and G treatments (n = 8), an Ancova was performed on data presenting a significant linear
235 relationship with time (Pearson's correlation coefficient, $p < 0.05$). Statistical tests were done using XLSTAT and R (**version**
236 **4.1.1 with the stats, tidyverse and FactoMineR packages**).

237

238 **3. Results**

239

240 **3.1 *In situ* N₂ fixation**

241 **3.1.1 Vertical and longitudinal distribution of N₂ fixation**

242 Over the cruise, the water column was well stratified with a shallow SML varying from 7 to 21 m depth (Table S1).
243 Detectable N₂ fixation rates in the 0-1000 m layer ranged from 0.04 to an exceptionally high rate of 72.1 nmol N L⁻¹ d⁻¹ at
244 Station 10 (Fig.2). Vertical N₂ fixation profiles exhibited a similar shape at all stations with maximum values within the
245 euphotic layer and undetectable values below 300 m depth (except at Stations 1 and 10 with rates ~ 0.05 nmol N L⁻¹ d⁻¹ at 500
246 m depth). Within the euphotic layer, all the rates were well above the detection limit (DL = 0.04 nmol N L⁻¹ d⁻¹; minimum *in*
247 *situ* N₂ fixation = 0.22 nmol N L⁻¹ d⁻¹). The highest rates were generally found below the SML and the lowest at the base of
248 the euphotic layer or within the SML (Fig.2). The lowest N₂ fixation rates integrated over the euphotic and aphotic (defined
249 as 1 % PAR depth to 1000 m) layers were found at Station 8, and the highest at Station 10 (Table 1). On average, 59 ± 16 %
250 of N₂ fixation (min 42 % at TYR and ION, max 97 % at Station 10) took place within the euphotic layer (Table 1). The
251 contribution of the SML integrated N₂ fixation to the euphotic layer integrated N₂ fixation was low, on average 17 ± 10 %.

252 **Volumetric surface** (~ 5 m) and euphotic layer integrated N₂ fixation rates exhibited a longitudinal gradient decreasing
253 eastward ($r = -0.59$ and $r = -0.60$, $p < 0.05$, respectively) (Fig.3). Integrated N₂ fixation rates over the SML, aphotic and 0-
254 1000 m layers (**Table 1**) displayed no significant trend with longitude ($p > 0.05$). It should be noted that longitudinal trends
255 with stronger correlations were observed for ¹³C-PP and BP ($r = -0.81$ and $r = -0.82$, $p < 0.05$, respectively, Fig.S2) as well as
256 DIP and NO₃⁻ stocks ($r = -0.68$ and $r = -0.85$, $p < 0.05$, no correlation with DFe stock; data not shown) integrated over the
257 euphotic layer.

258

259 **3.1.2 N₂ fixation and composition of diazotrophs at Station 10**

260 The westernmost Station 10 was in sharp contrast to all other stations with an euphotic integrated N₂ fixation on average 44
261 times higher (Table 1) due to high rates of 2.9 at 37 m and 72.1 nmol N L⁻¹ d⁻¹ at 61 m (i.e. at the deep chlorophyll-a
262 maximum, DCM) (Fig.2). That rate at 61 m was associated with a maximum in PP but not with a maximum in BP. From
263 surface to 200 m depth, the *nifH* community composition was largely dominated by ASVs related to different UCYN-A
264 groups (Fig.4) that represented 86 % at 200 m and up to 99.5 % at the DCM. No UCYN-B and -C or filamentous
265 diazotrophs were detected. The relative abundance of NCD (mainly γ -Proteobacteria *Pseudomonas*) increased with depth (r
266 = 0.96, $p < 0.05$) to reach about 8 % in the mesopelagic layer (200 m). UCYN-A1 and -A4 dominated the total diazotrophic
267 community (from 51 to 99 %). All four UCYN-A had different vertical distributions: the relative abundances of UCYN-A1

268 and -A3 were the highest in surface while UCYN-A4 was dominant at the most productive depths (61 and 88 m). At 61 m
269 depth, where the unusually high rate of N_2 fixation was detected, the community was dominated by both UCYN-A4 (58 %)
270 and UCYN-A1 (41 %).

271

272 **3.1.3 N_2 fixation versus primary production, heterotrophic bacterial production, nutrients**

273 For statistical analysis, **due to the high integrated N_2 fixation rate from Station 10, this rate was** not included in order not
274 to bias the analysis. N_2 fixation rate integrated over the euphotic layer correlated strongly with PP ($r = 0.71$, $p < 0.05$) and BP
275 ($r = 0.76$; $p < 0.05$) (Fig.5). Integrated N_2 fixation over the euphotic layer (and over the SML) was not correlated with the
276 associated DFe, DIP and NO_3^- stocks ($p > 0.05$). It should be noted that DIP and NO_3^- stocks correlated positively with PP
277 and BP ($p < 0.05$) over the euphotic layer (no correlation between DFe stock, and PP, BP).

278

279 **3.2 Response of N_2 fixation and composition of the diazotrophic communities to dust seeding**

280 **3.2.1 Initial characteristics of the tested seawater**

281 N_2 fixation and BP were the highest at FAST while PP was the highest at FAST and ION (Table 2). The N_2 fixation rates
282 were similar at ION and TYR and significantly higher (factor ~ 2.6) at FAST. At TYR and ION, the diazotrophs community
283 was largely dominated by NCD (on average 94.5 % of the total diazotrophic community) whereas at FAST, diazotrophic
284 cyanobacteria, mainly UCYN-A, represented on average 91.4 % of the total diazotrophic community. NO_3^- concentration
285 was the highest at FAST (59 nM) while DIP concentration was the highest at TYR (17 nM) and the lowest at ION (7 nM).
286 The molar NO_3^-/DIP ratio was strongly lower than the Redfield ratio (16/1) indicating a potential N limitation of the
287 phytoplanktonic activity in all experiments. DFe concentrations were all higher than 1.5 nM.

288

289 **3.2.2 Changes in N_2 fixation in response to dust seeding and relationship with changes in primary and heterotrophic 290 bacterial production**

291 All the dust seedings led to a significant stimulation of N_2 fixation relative to the controls under present and future climate
292 conditions (D and G treatments) (Figs.6, **S3**). The reproducibility between the replicated treatments was good at all stations
293 (**mean coefficient of variation (CV%)** < 14 %). The maximum N_2 fixation relative change (RC) was the highest at TYR
294 (+434-503 % in D1 and D2, +478-565 % in G1 and G2) then at ION (+256-173 % in D1 and D2, and +261-217 % in G1 and
295 G2) and finally at FAST (+41-49 % in D1 and D2 and +97-120 % in G1 and G2) (Fig.7). At TYR and FAST, dust addition
296 stimulated N_2 fixation more in the G treatment than in D, whereas at ION the response was similar between the treatments
297 (Fig.**S3**). N_2 fixation measured during the dust seeding experiments correlated strongly with PP at FAST ($r = 0.90$, $p < 0.05$),
298 and with BP at TYR and ION ($r > 0.76$, $p < 0.05$) (Fig. **S4**).

299 **3.2.3 Changes in the diazotrophic composition in response to dust seeding**

300 At TYR and ION, the diazotrophic communities before seeding were largely dominated by NCD (~ 94.5 % of total ASVs,
301 Fig. 8, Table 2). These were mainly γ -proteobacteria related to *Pseudomonas*. Some of these ASVs had low overall **relative**
302 abundance, and therefore did not appear in the top 20 ASVs (Fig.8; Tables S2, S3) but could nevertheless, account for up to
303 16 % in a specific sample. Filamentous cyanobacteria (*Katagnymene*) were also observed at both stations (~ 4.7 % of the
304 total diazotrophs). The community at FAST was initially dominated by UCYN-A phylotypes, mostly represented by UCYN-
305 A1 and -A3 (relative abundance of UCYN-A1 and -A3 **in C and D treatments at T0, n=4** : 34 ± 6 % and 45 ± 2 % of the
306 total diazotrophic composition, respectively, Fig.8). **At TYR and ION, the variability at T0 between replicates was**
307 **higher than at FAST (Fig. S5; C1T0 at TYR was removed due to poor sequencing quality). Also, the diversity**
308 **(Shannon H') was generally higher at TYR and ION at the start of incubations compared to FAST (T0, Fig.S6).** For
309 ION and FAST experiments, *Pseudomonas* related ASVs were more abundant in G treatments at T0 relative to Control and
310 Dust treatments (T0). At the end of the TYR and ION experiments, the community from all treatments appeared to converge
311 (Fig.S5) due to the increase of a few γ -proteobacteria (mainly *Pseudomonas*) that strongly increased in all treatments (Fig.8).
312 At FAST, no difference in the relative abundances of diazotrophs was recorded between D treatment and the controls at T4.
313 **However, when comparing G treatment relative to D at T4, the relative contribution of NCD was higher (82 % in G**
314 **vs. 63 % in D) and the relative abundance of UCYN-A was lower (13 % in G vs. 31 % in D).**

315

316 **4. Discussion**

317 Late spring, at the time of sampling, all the stations were well-stratified and characterized by oligotrophic conditions
318 increasing eastward (Maranon et al., 2021; Fig.8 in Guieu et al., 2020). NO_3^- and DIP concentrations were low in the SML,
319 from 9 to 135 nM for NO_3^- (Van Wambeke et al., 2021) and from 4 to 17 nM for DIP (Pulido-Villena et al., 2021); the
320 highest stocks were measured at the westernmost station (St 10) (Table S1).

321

322 **4.1 General features in N_2 fixation and diazotroph community composition**

323 N_2 fixation rates in the aphotic layer were in the range of those previously measured in the western open MS (Benavides et
324 al., 2016) and accounted, on average, for 41 % of N_2 fixation in the 0-1000 m layer, suggesting that a large part of the total
325 diazotrophic activity was related to heterotrophic NCD in the aphotic layer. N_2 fixation rates in the euphotic layer were of the
326 same order of magnitude (data from St10 excluded) than those previously measured in the open MS in spring and summer
327 (Bonnet et al., 2011; Rahav et al., 2013a). At the tested stations, the surface diazotrophic cyanobacteria were largely
328 dominated by UCYN-A (~ 93 % of the total diazotrophic cyanobacteria, mostly UCYN-A1 and -A3) and the NCD
329 community by γ -proteobacteria (~ 95 % of the total NCD). This is the first time that UCYN-A3 and -A4 are detected in the
330 MS. The photo-autotrophic N_2 fixation was negligible as no UCYN-B and -C were detected and very low abundance of
331 filamentous cyanobacteria was observed.

332

333 **4.2 Longitudinal gradient of N_2 fixation related to the composition of the diazotrophic communities**

334 At Station 10 and FAST, the surface diazotrophic communities were largely dominated by UCYN-A (> 91 %) whereas at
335 TYR and ION they were dominated by NCD (> 94 %) which highlights the predominance of photo-heterotrophic
336 diazotrophy in the western waters of the Algerian Basin and of NCD-supported diazotrophy in the Tyrrhenian and Ionian
337 basins. Surface N₂ fixation exhibited a longitudinal gradient decreasing eastward as previously reported (Bonnet et al., 2011,
338 Rahav et al., 2013a). Strong longitudinal gradients decreasing eastward for the relative abundance of UCYN-A (r = -0.93, p
339 < 0.05) and inversely increasing eastward for NCD were observed (r = 0.89, p < 0.05) (Fig.9). Despite no quantitative
340 abundances of distinct diazotrophs for the studied area (in this and previously published studies), the intensity of the bulk N₂
341 fixation rate was likely related to the overall composition of the diazotrophic communities (here relative abundance of
342 UCYN-A versus NCD). Indeed, surface N₂ fixation rates correlated positively with the relative abundance of UCYN-A
343 (mainly A1 and A3) (r = 0.98, p < 0.05) and negatively with the relative abundance of NCD (r = - 0.99, p < 0.05) (Fig. S7).
344 This could be related, in part, to the variability of the cell-specific N₂ fixation rates that were shown to be higher for UCYN-
345 A relative to NCD (Turk-Kubo et al., 2014; Bentzon-Tilia et al., 2015; Martinez-Perez et al., 2016; Pearl et al., 2018; Mills
346 et al., 2020). Besides, in Atlantic and Pacific Ocean areas when the diazotrophic community is dominated by unicellular
347 organisms, high N₂ fixation rates are mostly associated with a predominance of UCYN-A, and low rates with a
348 predominance of NCD (Turk-Kubo et al., 2014, Martinez-Perez et al., 2016; Moreira-Coello et al., 2017, Fonseca-Batista et
349 al., 2019; Tang et al., 2019).

350

351 4.3 Intriguing Station 10

352 The patchy distribution of the diazotrophic activity at Station 10 was related to an exceptionally high rate at the DCM (72.1
353 nmol N L⁻¹ d⁻¹). **High N₂ fixation rates have previously been observed locally:** 2.4 nmol N L⁻¹ d⁻¹ at the Strait of Gibraltar
354 (Rahav et al., 2013a), ~5 nmol N L⁻¹ d⁻¹ in the Bay of Calvi (Rees et al., 2017), 17 nmol N L⁻¹ d⁻¹ in the northwestern MS
355 (Garcia et al., 2006) and 129 nmol N L⁻¹ d⁻¹ in the eastern MS (Rees et al., 2006). Station 10 was also hydrodynamically
356 "contrasted" compared to the other stations: it was located almost at the centre of an anticyclonic eddy (Guieu et al., 2020),
357 with the core waters (0-200 m) of Atlantic origin (colder, fresher). In such anticyclonic structures, enhanced **exchange** with
358 nutrients rich waters from below take place, **and** combined with lateral mixing, could explain **higher** stocks of NO₃⁻ and DIP
359 in the euphotic layer (Table S1). Nevertheless, the anomaly of N₂ fixation at the DCM was neither associated with anomalies
360 of PP, BP nor NO₃⁻ and DIP concentrations. It only coincided with a minimum in DFe concentration (0.47 nM compared to
361 0.7 to 1.4 nM at the nearby depths, Bressac et al., 2021). ~~which could not be explained solely by the diazotroph uptake.~~
362 **Based on a range of Fe:C (from 7 to 177 μmol:mol) and associated C:N ratios for diazotrophs (*Trichodesmium*,
363 UCYN) from literature (Tuit et al., 2004; Berman-Frank et al., 2007; Jiang et al., 2018), we found that 0.004 nM to
364 0.08 nM of DFe are required to sustain this N₂ fixation rate. Consequently, the minimum in DFe concentration at 61m
365 could not be explained solely by the diazotroph uptake.**

366

367 Despite no correlation between N_2 fixation and the relative abundance of specific diazotrophs ($p > 0.05$) along the profile,
368 the huge heterogeneity in N_2 fixation rate was likely related to the patchy distribution of diazotrophs taxa. Indeed, patchiness
369 seems to be a common feature of unicellular diazotrophs (Robinart et al., 2014; Moreira-Coello et al., 2019). The
370 exceptionally high N_2 fixation rates coincided with the highest relative contributions of UCYN-A and more precisely
371 UCYN-A4. Exceptional N_2 fixation rates at Station 10, impacted by northeast Atlantic surface waters of subtropical origin
372 could thus be related to that incoming waters. Indeed, Fonseca-Batista et al. (2019) reported high N_2 fixation rates (45 and 65
373 $\text{nmol N L}^{-1} \text{d}^{-1}$ with euphotic N_2 fixation rates up to $1533 \mu\text{molN m}^{-2} \text{d}^{-1}$) associated with a predominance of UCYN-A in
374 subtropical Atlantic surface water mass along the Iberian Margin ($\sim 40^\circ \text{N}$ - 11°E). It should be noted that UCYN-A4 was
375 only detected at Station 10, and its relatively high contribution to the **whole diazotrophic community** in the euphotic layer
376 coincided with the highest stocks of P (and N) (**Table S1**). This could reflect higher nutrient requirement(s) of the UCYN-
377 A4 and/or of its eukaryotic partner relative to other sublineages. Another intriguing feature was the high contribution (~ 86
378 %) of UCYN-A in the mesopelagic zone (200 m). As UCYN-A lives in obligate symbiosis with haptophytes from which it
379 receives fixed carbon from photosynthesis (Thompson et al., 2012, 2014), this suggests that this contribution was probably
380 derived from sinking senescing prymnesiophyte-UCYN-A cells, and that the weak N_2 fixation rate at 200m depth is likely
381 only driven by γ -proteobacteria (*Pseudomonas*).

382

383 **4.4 Supply of bioavailable N from diazotrophic activity for fueling primary and heterotrophic bacterial production -** 384 **Relationship with potential controlling factors of N_2 fixation**

385 The relationship established between N_2 fixation, and PP and BP illustrated that in the studied area, N_2 fixation is promoted
386 by UCYN and NCD, and/or could indicate that all processes have the same (co)-limitation. Overall, N_2 fixation was a poor
387 contributor to PP (1.0 ± 0.3 %, **Fig. S8**), as previously shown in the MS (Bonnet et al., 2011; Yogev et al., 2011; Rahav et al.
388 2013a) and BP (**7 ± 1 %**, **Fig. S8**) except at Station 10 where N_2 fixation could support up to 19 % of PP (**Fig. S8**) and
389 supply the entire bioavailable N requirements for heterotrophic prokaryotes (**199 % of BP**). As expected, our results suggest
390 no control of N_2 fixation by DFe and NO_3^- , as previously shown through nutrient additions in microcoms (Rees et al., 2006;
391 Ridame et al., 2011, Rahav et al., 2016b). No correlation was observed between N_2 fixation and DIP which may highlight the
392 spatial variability of the controlling factor of diazotrophs as DIP was shown to control N_2 fixation in the western basin, but
393 not in the Ionian basin (Ridame et al., 2011). Moreover, DIP concentration does not reflect the rapid turnover of P in the
394 open MS and thus could be a poor indicator of DIP availability (Pulido-Villena et al., 2021).

395

396 **4.5 Diazotrophic activity and composition in response to dust addition under present climate conditions**

397 **General features** – In all experiments, simulated wet dust deposition under present climate conditions triggered a significant
398 (41 to 503 %) and rapid (24-48 h) stimulation of N_2 fixation relative to the controls. Despite this strong increase, N_2 fixation
399 rates remained low ($< 0.7 \text{ nmol N L}^{-1} \text{d}^{-1}$) as well as their contribution to PP (< 7 %) and BP (< 5 %) as observed *in situ*

400 (Sect.4.4). All of these results are consistent with those found after dust seeding in mesocosms in a coastal site in the
401 northwestern MS (Ridame et al., 2013) and in the open Cretan Sea (Rahav et al., 2016a).

402 **Temporal Changes in the composition of the diazotrophic community-** Dust addition under present climate conditions did
403 not impact the diazotrophic communities composition. At TYR and ION, the large increase in N_2 fixation recorded after dust
404 addition might be attributed to NCD (mainly γ -proteobacteria), as suggested by the positive correlation between N_2 fixation
405 and BP (Fig S4). At FAST, the community shifted from a large dominance of UCYN-A towards a dominance of NCD both
406 in the dust treatments and unamended controls due to the increase in a few fast growing γ -proteobacteria (mainly
407 *Pseudomonas*). This shift could be attributed to a bottle effect imposed by the tanks which can favor fast growing
408 heterotrophic bacteria (Sherr et al. 1999; Calvo-Diaz et al., 2011). Nevertheless, the increased N_2 fixation after dust seeding
409 at FAST cannot be explained by the shift in composition of the diazotrophic communities because the rates remained quite
410 stable in the controls all along the experiment. Rather, the abundances of diazotrophs have **likely** increased due to dust input,
411 and UCYN-A in association with prymnesiophytes could still be responsible for the majority of the enhanced N_2 fixation as
412 N_2 fixation correlated strongly with PP (Fig. S4).

413 **Variability of the N_2 fixation response among stations** - The highest stimulation of N_2 fixation to dust addition was
414 observed at TYR (mean $RC_D = 321$ %) then at ION (mean $RC_D = 161$ %) and finally at FAST (mean $RC_D = 21$ %) (Fig.7).
415 The differences in the intensity of the diazotrophic response were not related to differences in the initial nutrients stocks
416 (Table S1) and in the nutrients input from dust which was quite similar between experiments (Gazeau et al., 2021a). Briefly,
417 dust input led to a strong increase of $11.2 \pm 0.2 \mu M NO_3^-$ few hours after seeding in the three experiments, and the maximum
418 DIP release was slightly higher at FAST (31 nM) than at TYR and ION (23 ± 2 nM) (Gazeau et al., 2021a). As DFe
419 concentration before seeding was high (≥ 1.5 nM, Table 2), the bioavailability of Fe did not appear to drive the response of
420 N_2 fixation (Ridame et al., 2013). Also, we evidenced in this experiment that NO_3^- release from dust did not inhibit N_2
421 fixation rate driven by UCYN-A and NCD. This was expected for UCYN-A as it lacks NO_3^- assimilation pathways (Tripp et
422 al., 2010; Bombar et al., 2014).

423 N_2 fixation was initially more limited at TYR and ION (as evidenced by the lowest initial rates) compared to FAST, thereby
424 explaining the highest stimulation of N_2 fixation **by** dust seeding at these stations. Interestingly, the stimulation of N_2 fixation
425 was higher at TYR than at ION (Fig.7) while these stations presented the same initial rate supported by NCD. One major
426 difference is that PP was not enhanced by dust seeding at TYR while BP increased in both experiments (Gazeau et al.,
427 2021b) suggesting that NCD-supported N_2 fixation was not limited by organic carbon at this station. As N_2 fixation and BP
428 correlated strongly after the dust seeding (Fig. S4), it means that dust-derived DIP could relieve the ambient limitation of
429 both heterotrophic prokaryotes (BP was co-limited by NP, Van Wambeke et al., 2021) and NCD at TYR. This could explain
430 why DIP concentration in the D treatments became again similar to the controls at the end of this experiment (Gazeau et al.,
431 2021a). At ION characterized by the lowest initial DIP concentration, N_2 fixation and PP were likely DIP (co-)limited as
432 shown for BP (Van Wambeke et al., 2021). **Consequently, diazotrophs as well as non diazotrophs (heterotrophic
433 prokaryotes and photoautotrophs) could all uptake the dust-derived DIP reducing then potentially the amount of DIP**

434 available for each cell that could explain the lower stimulation of N₂ fixation relative to TYR. Consequently,
435 heterotrophic prokaryotes, NCD, and photoautotrophs could out compete for dust derived DIP uptake reducing then the
436 amount of DIP per cell explaining the lower stimulation of N₂ fixation relative to TYR.

437 At FAST, initially dominated by UCYN-A, N₂ fixation and PP correlated strongly after the dust seeding (Fig. S4c). This
438 indicated that dust could relieve either directly the ambient nutrient limitation of both N₂ fixation and PP (Fig.S9) or
439 indirectly through first the relief of the PP limitation of the UCYN-A photoautotroph hosts inducing an increase in the
440 production of organic carbon which could be used by UCYN-A to increase its N₂-fixing activity. **Nutrients from dust could
441 also first enhance the UCYNA-supported N₂ fixation, which in turn could relieve the N limitation of the UCYN-A
442 photoautotrophic host, as the initial NO₃/DIP ratio indicates a potential N limitation of the PP (Table 2).'**

443

444 **4.6 Response to dust addition under future relative to present climate conditions**

445 **General features** -At TYR and FAST, N₂ fixation was more stimulated by dust input under future than present climate
446 conditions (mean RC_{G-TYR}= 478 % and mean RC_{G-FAST}= 54 %) whereas at ION the response was similar (Figs.7, S3). These
447 differences between future and present climate conditions were not related to the nutrients supplied from dust (Gazeau et al.,
448 2021a).

449 The purpose of our study was to study the combined effect of warming and acidification, but we can expect on the short time
450 scale of our experiments (< 3-4 days), that NCD and UCYN-A would not be directly affected by ~~ocean acidification and~~ the
451 ~~associated~~ changes in the CO₂ concentration as they do not fix CO₂ (Zehr et al., 2008). Indeed, no impact of acidification (or
452 pCO₂ increase) on N₂ fixation was detected when the diazotrophic communities were dominated by UCYN-A in the North
453 and South Pacific (Law et al., 2012; Böttjer et al. 2014). **Nevertheless, the decrease in pH may indirectly impact UCYN-
454 A through changes affecting its autotrophic host.**

455

456 **TYR and ION** –Under future climate conditions, the composition of the diazotrophic communities did not change after dust
457 input at TYR and ION relative to present conditions. At TYR, the highest N₂ fixation stimulation might be linked to the
458 increase in the NCD abundances and/or in their cell-specific N₂ fixation rates under future climate conditions. Unfortunately,
459 the impact of increased temperature and decreased pH on the cell-specific N₂ fixation rates of NCD is currently unknown.
460 However, some studies suggest a positive relationship between temperature and abundances of NCD: diazotrophic γ -
461 proteobacteria (γ -24774A11) gene copies correlated positively with temperature (from ~20 to 30° C) in surface waters of the
462 western South Pacific Ocean (Moisander et al., 2014), and Messer et al. (2015) suggested a temperature optima for these γ -
463 proteobacteria around 25-26° C in the Australian tropical waters. At ION, the similar stimulation of N₂ fixation by dust
464 under future climate conditions compared to present climate conditions could be explained by a greater mortality of
465 diazotrophs due to a higher grazing pressure and/or a higher viral activity. Indeed, higher bacterial mortality in the G
466 treatment that could be related to a higher grazing pressure has been observed (Dinasquet et al., 2021). Another explanation

467 is that in spite of the DIP supply from the dust, the DIP bioavailability, initially the lowest at ION, was not sufficient to allow
468 an additional N₂ fixation stimulation.

469 **FAST-** Some differences in the composition of the diazotrophic communities were observed between present and future
470 climate conditions at FAST after dust input: the contribution of NCD (mainly *Pseudomonas*) increased and that of UCYN-A
471 decreased. It must be noted that the duration of the experiment was longer at FAST (4 days) relative to TYR and ION (3
472 days) which could explain at least partly differences between stations. A direct response of increased temperature and/or
473 decreased pH can be considered on a very short time scale (12 hours) by comparing the results in the G treatment at T0 (+3°
474 C, -0.3 pH unit) with those in C and D treatments. The increased contribution of *Pseudomonas* in the G treatment at T0
475 (before dust addition) reveals a likely positive effect of temperature on the growth of this NCD **as an increase in the top-
476 down control on the bacterioplankton was observed after dust seeding under future climate conditions (Dinasquet et
477 al., 2021) instead of a decrease in the top-down control on the bacterioplankton which is strongly suspected to increase under
478 future climate conditions.** Interestingly, despite the decrease in the **relative** contribution of UCYN-A **to the total diazotroph
479 community** after dust addition, we observed contrasted responses within the UCYN-A **group** relative to present climate
480 conditions: **the relative abundance of UCYN-A3** strongly decreased (4.6 % in G vs. 25.4 % in D) whereas **the relative
481 abundance of UCYN-A2** was twice as high (7 % in G vs. 3.4 % in D). **Notably, the relative contribution of UCYN-A1
482 did not appear to be impacted during the dust addition experiment.** These respective changes could be explained by the
483 difference in the temperature tolerance between UCYN-A2 and -A3. Temperature is one of the key drivers explaining the
484 distribution of UCYN-A which appeared to dominate in most of the temperate regions with temperature optima around ~20-
485 24° C (Langlois et al., 2008; Moisaner et al., 2010). However, the temperature optima for the different UCYN-A
486 sublineages, in particular for UCYN-A2 and -A3, are poorly known. Interestingly, Henke et al. (2018) observed that the
487 **absolute** UCYN-A2 abundance was positively affected by increasing temperature, within a range of temperature from about
488 21 to 28° C which is in agreement with our results **although only relative abundances were measured in our study.** Based
489 on the strong positive correlation between N₂ fixation and PP after dust addition **(and no correlation between N₂ fixation
490 and BP, Fig. S4)**, and despite the decrease in the relative abundance of UCYN-A3, the increased stimulation of N₂ fixation
491 under future climate conditions could **likely** be sustained by the increase in the relative abundance of UCYN-A2 which is
492 bigger than UCYN-A3 (Cornejo-Castillo et al., 2019) and could consequently have a higher cell-specific N₂ fixation rate.

493

494 **5. Conclusion**

495 In the MS, N₂ fixation is a minor pathway to supply new bioavailable N for sustaining both PP and BP but can locally
496 support up to 20 % of PP and provide all the N requirement for bacterial activity. UCYN-A **might be supporting** extremely
497 high rates of N₂ fixation (72 nmol.L⁻¹.d⁻¹) in the core of an eddy in the Algerian basin influenced by Atlantic waters. The
498 eastward decreasing longitudinal trend of N₂ fixation in the surface waters is likely related to the spatial variability of the
499 composition of the diazotrophic communities, as shown by the eastward increase in the relative abundance of NCD towards
500 more oligotrophic waters while we observed a westward increase in the relative abundance of UCYN-A. This could reflect

501 lower nutrients requirements for NCD relative to UCYN-A. Through the release of new nutrients, simulated wet dust
502 deposition under present and future climate conditions significantly **stimulated** N₂ fixation. The degree of stimulation
503 depended on the metabolic activity of the diazotrophs (degree of limitation) related to the composition of diazotrophic
504 communities, and on the ambient potential nutrient limitations of diazotrophs, including that of the UCYN-A
505 prymnesiophyte host. The strongest increase in N₂ fixation, not accompanied with a change in the composition of the
506 diazotrophic communities, was observed at the stations dominated by NCD (TYR, ION) where the nutrient limitation was
507 the strongest. Under projected future levels of temperature and pH, the dust effect is either exacerbated or unchanged.
508 Knowing that NCD and UCYN-A do not fix CO₂, we suggest that, on the time scale of our experiments (3-4 days), the
509 exacerbated response of N₂ fixation is likely the result of the warming (from about 21° C to 24° C) which may increase the
510 growth of NCD when nutrient availability allows it, and may alter the composition of UCYN-A community. However, to
511 date, the effect of acidification and temperature optima of the different UCYN-A sublineages are poorly known (or
512 unknown) as these UCYN-A **remain** uncultivated.

513 Future changes in climate, desertification and land use practices could induce an increase in dust deposition to the oceans
514 (Tegen et al., 2004; Moulin and Chiapello, 2006; Klingmüller et al., 2016). The predicted future increase in surface
515 temperature, and the resulting stronger stratification **are expected** to expand the surface of LNLC areas reinforcing
516 consequently the role of new nutrient supply from aeolian dust on the N₂ fixation and probably on the structure of the
517 diazotrophic communities.

518

519 **6. Data availability**

520 Guieu, C. et al. (2020). Biogeochemical dataset collected during the PEACETIME cruise. SEANOE.
521 <https://doi.org/10.17882/75747>.

522

523 **7. Author contributions**

524 FG and CG designed the dust seedings experiments. CR, JD, EB, MB, FVW, FG, VT, AT-S and CG participated to the
525 sampling and analysis. CR and EB performed DNA extraction; EB performed library preparation. CR, JD and SH analyzed
526 the data; CR wrote the manuscript with contributions from all authors.

527

528 **8. Competing interests**

529 The authors declare that they have no conflict of interest

530

531 **9. Special issue Statement**

532 This article is part of the special issue “Atmospheric deposition in the low-nutrient–low-chlorophyll (LNLC) ocean: effects
533 on marine life today and in the future (ACP/BG inter-journal SI)”. It is not associated with a conference.

534

535 **10. Financial support**

536 This study is a contribution to the PEACETIME project (<http://peacetime-project.org>), a joint initiative of the MERMEX and
537 ChArMEX components supported by CNRS-INSU, IFREMER, CEA, and Météo-France as part of the programme
538 MISTRALS coordinated by INSU. PEACETIME was endorsed as a process study by GEOTRACES. JD was funded by a
539 Marie Curie Actions-International Outgoing Fellowship (PIOF-GA-2013-629378). SH and LR were funded by grant 6108-
540 00013 from the Danish Council for independent research to LR.

541

542 **11. Acknowledgments**

543 The authors thank the captain and the crew of the RV Pourquoi Pas ? for their professionalism and their work at sea. We
544 warmly acknowledge our second ‘chieffe’ scientist Karine Desboeufs. We gratefully thank Eric Thiebaut and Pierre
545 Kostyrka for their precious advice with statistical tests. We also thank Kahina Djaoudi and Thibaut Wagener for their
546 assistance in sampling the tanks and TMC-rosette, **Magloire Mandeng-Yogo and Fethiye Cetin for IRMS measurements**
547 **at the Alyses plate-form (SU, IRD). The DIC data used in this study were analyzed at the SNAPO-CO₂ service facility**
548 **at LOCEAN laboratory and supported by CNRS-INSU and OSU Ecce-Terra**

549

550

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902 **Table 1:** Integrated N₂ fixation over the surface mixed layer (SML, from surface to the mixed layer depth), from the surface
 903 to the base of the euphotic layer (1% PAR depth), over the aphotic layer (1%PAR depth to 1000 m), and from surface to
 904 1000 m at all the sampled stations. Contribution (in %) of SML integrated N₂ fixation to euphotic layer integrated N₂
 905 fixation, and contribution of euphotic layer integrated N₂ fixation to total (0-1000 m) integrated N₂ fixation.
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	N ₂ Fix _{SML}	N ₂ Fix _{euphotic}	N ₂ Fix _{aphotic}	N ₂ Fix _{0-1000m}	N ₂ Fix _{SML} /N ₂ Fix _{euphotic}	N ₂ Fix _{euphotic} /N ₂ Fix 0-1000m
	μmolN m ⁻² d ⁻¹	μmolN m ⁻² d ⁻¹	μmolN m ⁻² d ⁻¹	μmolN m ⁻² d ⁻¹	%	%
ST01	14.6	42.6	56.5	99.1	34	43
ST02	10.7	36.0	16.0	51.9	30	69
ST03	7.8	58.3	18.1	76.4	13	76
ST04	10.8	46.6	38.5	85.1	23	55
ST05	4.9	46.3	36.1	82.4	10	56
TYR	4.2	38.6	53.0	91.6	11	42
ST06	9.1	34.9	29.8	64.7	26	54
ST07	10.5	43.5	55.4	98.8	24	44
ION	6.2	40.6	56.5	97.1	15	42
ST08	4.3	27.0	12.3	39.3	16	69
ST09	3.4	50.2	43.3	93.5	7	54
FAST	5.9	58.2	35.7	93.8	10	62
ST10	13.7	1908	63.7	1972	1	97
Mean ± std (ST10 excluded)	7.7±3.5	44±9	38±16	81±20	18%±9%	55%±12%
Mean ± std (all stations)	8.2±3.7	187±517	40±17	227±525	17%±10%	59%±16%

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914 **Table 2:** Initial physico-chemical and biological properties of surface seawater before the perturbation in the dust seeding
 915 experiments at TYR, ION and FAST (average at T0 in C and D treatments, n=4 or data at T-12h in the pumped surface
 916 waters, n=1). The relative abundances of diazotrophic cyanobacteria and NCD (non-cyanobacterial diazotroph) are given as
 917 proportion of total *nifH* sequence reads. DIP: dissolved inorganic phosphorus, DFe: dissolved iron. **The C:N ratio**
 918 **corresponds to the ratio in the organic particulate matter from IRMS measurements (> 0.7µm).** Means that did not
 919 differ significantly between the experiments (p>0.05) are labeled with the same letter (in parenthesis).
 920

	TYR	ION	FAST
Day of sampling	05/17/2017	05/25/2017	06/02/2017
Temperature (° C)*	20.6	21.2	21.5
Salinity*	37.96	39.02	37.07
¹³ C-Primary production, mg C m ⁻³ d ⁻¹	1.23±0.64 (A)	2.53±0.40 (B)	2.82±0.55 (B)
N ₂ fixation nmol N L ⁻¹ d ⁻¹	0.19±0.03 (A)	0.21±0.05 (A)	0.51±0.04 (B)
Relative abundance of diazotrophic cyanobacteria (%)	4.7±3.8 (A)	6.2±6.5 (A)	91.4±6.0 (B)
Relative abundance of NCD (%)	95.3±3.9 (A)	93.8±6.5 (A)	8.6±6.0 (B)
Heterotrophic bacterial production ng C. L ⁻¹ h ⁻¹	26.6 ± 7.0 (AB)	25.9 ± 0.9 (A)	36.3 ± 1.2 (B)
C:N (mol/mol)	9.6±0.8 (A)	10.2±0.8 (A)	9.1±0.5 (A)
DIP, nM*	17	7	13
NO ₃ ⁻ , nM*	14	18	59
NO ₃ ⁻ /DIP, mol/mol	0.8	2.6	4.5
DFe, nM [§]	1.5±0.1 (A)	2.6±0.2 (B)	1.8±0.2 (A)

921 * from Gazeau et al., 2021a

922 § from Roy-Barman et al., 2021

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933 **Figures captions**

934 Figure 1: Locations of the ten short (ST1 to ST10) and three long stations (**TYR**, ION and FAST). Stations 1 and 2 were
935 located in the Provencal basin; Stations 5, 6, and TYR, in the Tyrrhenian Sea; Stations 7, 8, and ION in the Ionian Sea; and
936 Stations 3, 4, 9, 10 and FAST in the Algerian basin. Satellite-derived chlorophyll-a concentration (mg m^{-3}) averaged over the
937 entire duration of the PEACETIME cruise. Ocean color data from MODIS/Aqua, NASA.

938
939 Figure 2: Vertical distribution of N_2 fixation (in $\text{nmol N L}^{-1} \text{d}^{-1}$) in the Provencal (a), Tyrrhenian (b) Ionian (c) and Algerian
940 (d) basins and at Station 10 (e). N_2 fixation rates at Station 10 are plotted ~~individually~~ **in log scale because** of the high
941 fluxes. **Rates under detection limit ($<0.04 \text{ nmol N L}^{-1} \text{d}^{-1}$) are symbolized by crosses.**

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943 Figure 3: **Volumetric** surface (**$\sim 5\text{m}$**) (a) and integrated N_2 fixation from surface to euphotic layer depth (b) along the
944 longitudinal PEACETIME transect (Station 10 was excluded). **Integrated N_2 fixation rate from Station 10 was excluded**
945 **from statistical analysis**

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947 Figure 4: Vertical distribution of the 20 most abundant nifH-ASVs at Station 10, collapsed into major taxonomic groups.

948 Figure 5: N_2 fixation rate integrated over the euphotic layer versus ^{13}C -primary production (a) and bacterial production (b);
949 data at Station 10 were removed.

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951 Figure 6: N_2 fixation rate in $\text{nmol N L}^{-1} \text{d}^{-1}$ during the dust seeding experiments performed at the stations TYR (a), ION (b)
952 and FAST (c) in the replicated controls (black dot), dust treatments under present climate conditions (red square, D
953 treatment) and dust treatments under future climate conditions (green triangle, G treatment). Open symbols were not
954 included in the linear regression

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956 Figure 7: Box plots of the relative changes (in %) in N_2 fixation to the rates measured in the controls over the duration of the
957 dust seeding experiments (**T1, T2, T3 or T4**) at TYR, ION, and FAST stations. D means dust treatments under present
958 climate conditions (D treatment) and G dust treatments under future climate conditions (G treatment). The red cross
959 represents the average.

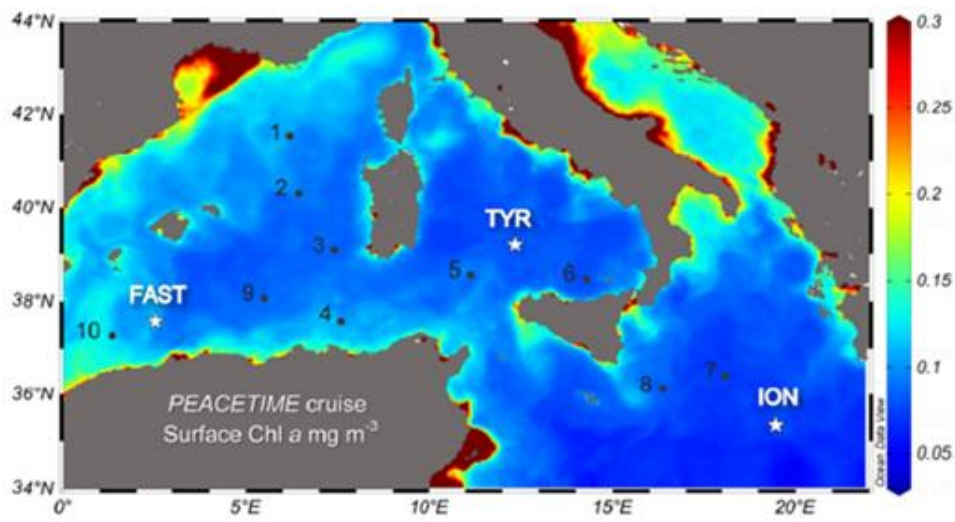
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961 Figure 8: The composition of diazotrophs (based on 20 most abundant ASVs in the tanks) during the dust seeding
962 experiments at the start (T0) and end (T3 at TYR and ION, and T4 at FAST) in each tank, at TYR (Top panel), ION (middle
963 panel) and FAST (bottom panel). C1T0 at TYR was not included due to poor sequencing quality.

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965 Figure 9: Relative abundance of the 20 most abundant nifH-ASVs in surface waters (values at TYR, ION and FAST are
966 based on average of duplicated control and dust treatments at T0).

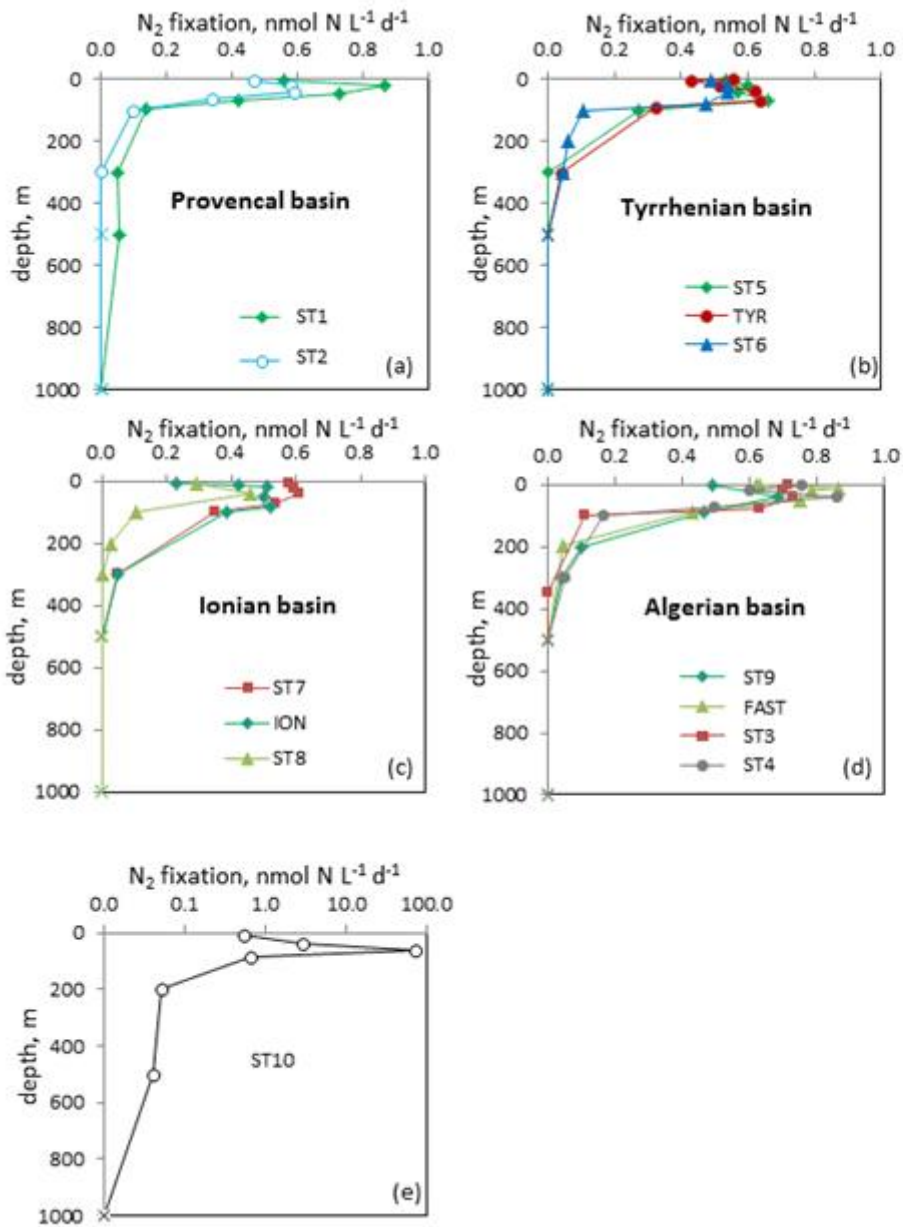
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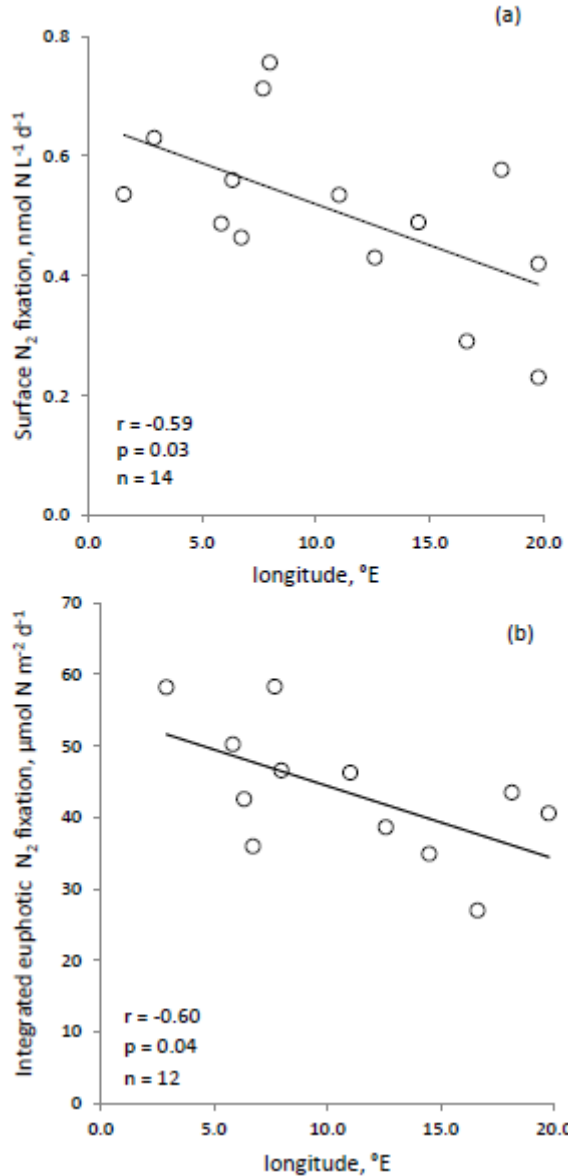
Figure 1 :



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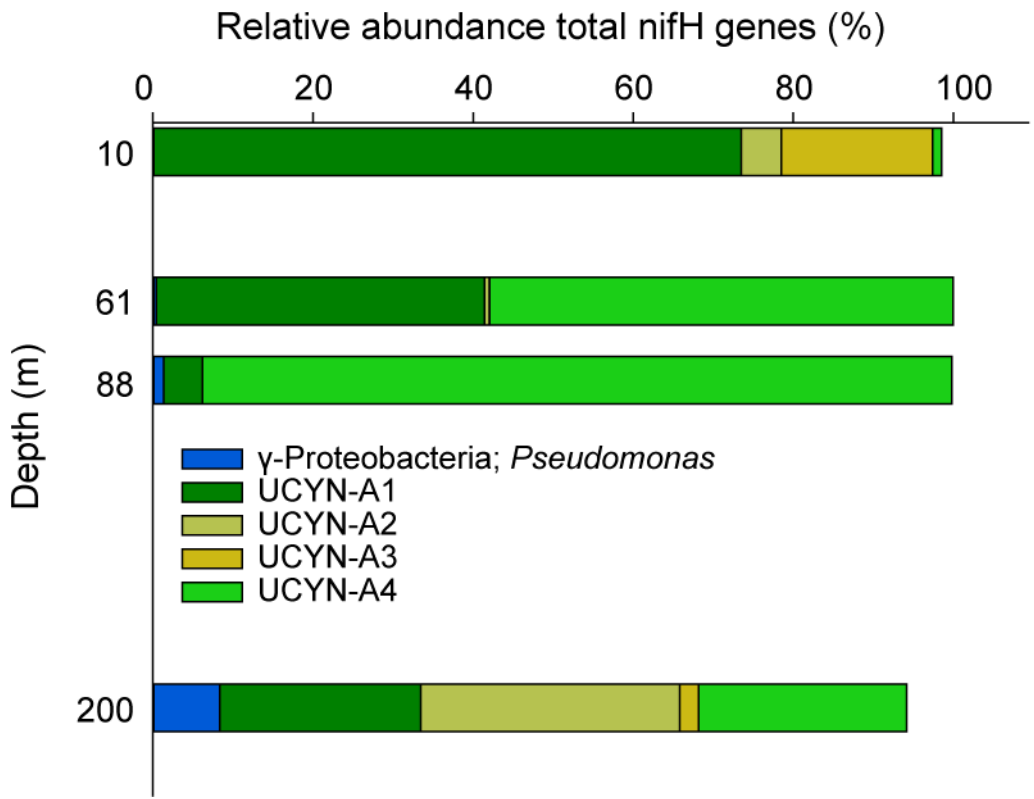
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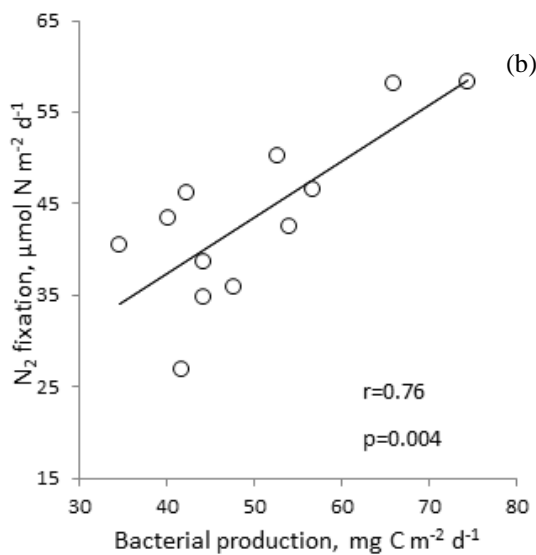
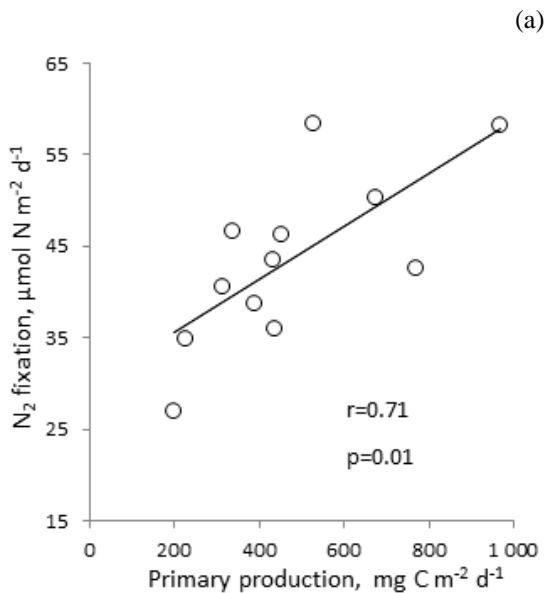
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1014 **Figure 4**

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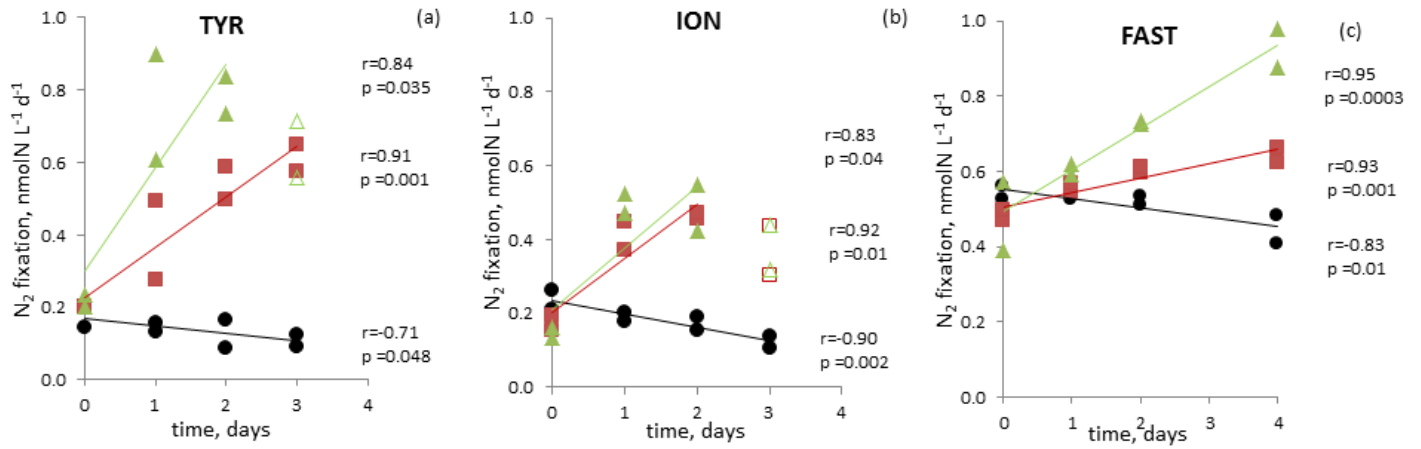
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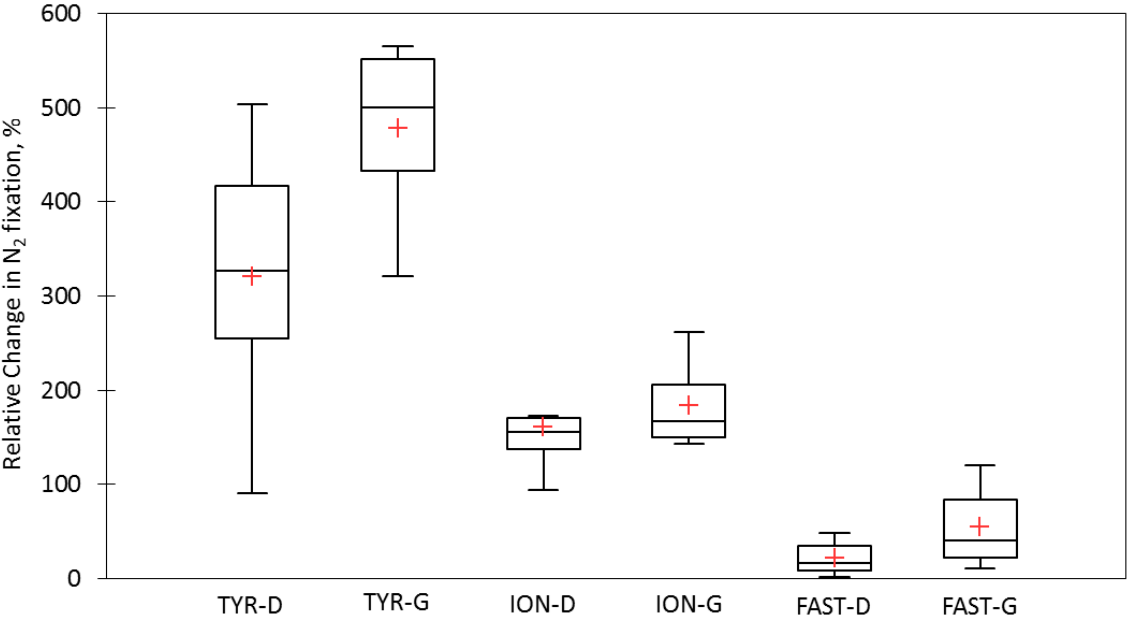
Figure 6



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1066 **Figure 7**

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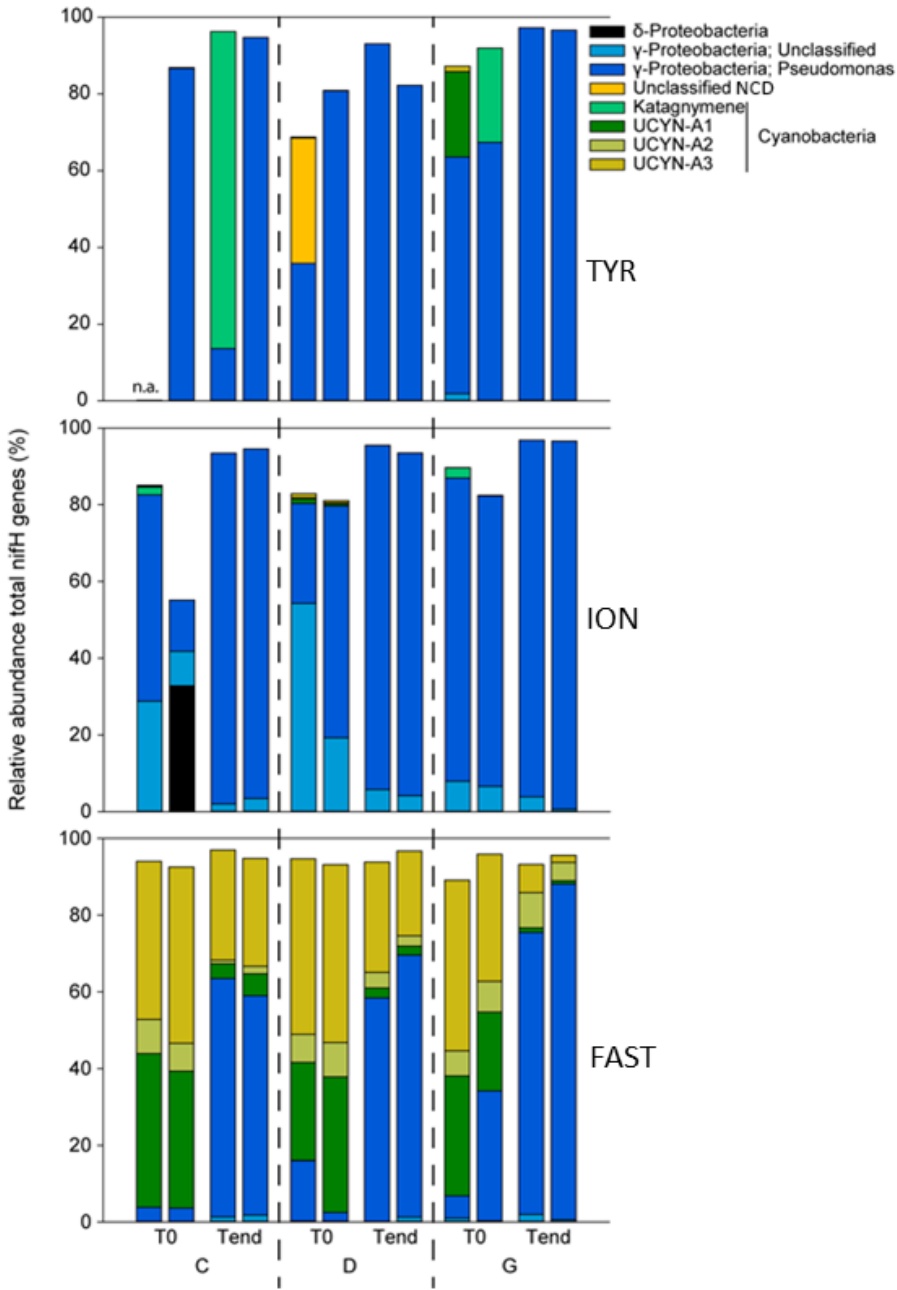
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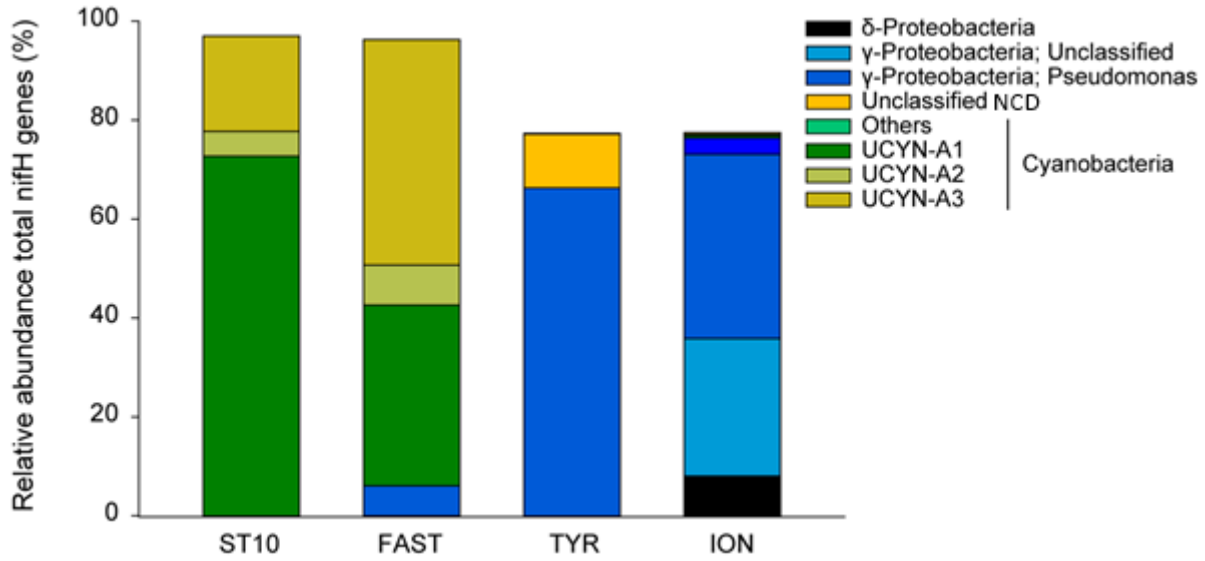
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Figure 9



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