

MS No.: BG-2021-193

Interactive comment on “Effects of soil water content on carbon sink strength in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau”

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November 11, 2021

We appreciate you and reviewers for the valuable and constructive comments. We have carefully considered them all and changed our manuscript accordingly. In the following, we include a point-by-point response to the reviewers #1 and #2, and attached a marked-up manuscript version showing the differences to the initially submitted version. Please note that the line numbers point to the non-marked manuscript.

Interactive comment on “Effects of soil water content on carbon sink strength in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau” by Wei et al.

Received and published for Discussion: 24 September 2021

Response to Comments of Referee #1

Wei et al. reports that “Effects of soil water content on carbon sink strength in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau”. This study investigated the diurnal, seasonal and annual variability of CO₂ fluxes and their drivers in an alpine swamp on the northeastern Tibetan Plateau. This helps to more clearly understand the role of alpine swamp in alpine ecosystem carbon (C) cycling in Tibetan Plateau, because alpine swamp cycling in Tibetan Plateau is less focused at regional scale, compared with alpine steppe and alpine meadow. The text is well-written and clear. Nevertheless, I have reservations about the innovation of scientific questions and the reliability of some results that need to be addressed before the publication of this manuscript.

We thank Reviewer #1 for taking the time to assess our manuscript and for providing general positive comments and main concerns. We believe the comments have helped to improve the manuscript and we carefully considered them. Here specifically we clarify the innovation of our scientific questions addressed in this study. As REF#1 pointed out, alpine swamps in Tibetan Plateau are less focused at regional scale than alpine steppes and alpine meadows. This study highlights among other things the importance of soil water availability regulating carbon sink strength of an alpine swamp, which is characterized by saturated water condition and high soil water content (SWC). The role of soil water has often been neglected or assumed to be less important relative to other factors for carbon (C) cycling. This study provides a four-year field observation dataset to characterize and quantify the importance of soil water controlling the C sink strength of an alpine swamp – one key finding is that a 15% decrease in soil water can induce 25% higher respiration and therefore weaken C sink strength by 20%, and an additional 44% increase of temperature at annual scale can also weaken the C sink strength by about 50% (see answer to comment number 4). These new insights will help us to better understand, model and predict the complex C cycle dynamics in the Tibetan Plateau driven by the almost certain future intensified climate warming.

1. The experimental site is located at Haibei in the northeastern Tibetan Plateau. According to Wei et. al (2021), there are at least six eddy covariance sites at Haibei, including alpine swamp CO₂ fluxes monitoring site. Haibei is the most densely distributed area of eddy covariance sites on the Tibetan Plateau. The strength of CO₂ sink and its diurnal, seasonal and interannual characteristics in alpine swamp at Haibei have been reported in previous publications, such as Zhao et al. (2005) and Zhao et al. (2010), yet it is also the first objective of this study. Thus, the innovation of the objective is not clear to me.

We appreciate REF#1 for sharing this very useful information about other existing eddy covariance sites and related references. We have now cited these previous studies in the revised manuscript see L76 and L77. According to Wei et al. (2021), there are six observational studies about C fluxes at Haibei. However, only three of them focused on alpine swamp meadows (or wetland in Wei et. al (2021)). Among them, one study only focused on a 1-year dataset (Zhang et al., 2008), and the other two characterized the same location (Zhao et al., 2005, 2010). Moreover,

these alpine swamp meadows were reported as a net C source while we our site showed a consistent C sink. The different directions of C exchange suggest that there are still uncertainties in our understanding of C exchange in this alpine swamp meadows, and further insights are obtained from studying multiple years of observations. Therefore, further efforts are still needed to improve our projection of C balance change of this ecosystem under changing climate.

Additionally, as mentioned before, previous studies focusing on C fluxes in alpine swamp meadows did not give enough consideration to the effects of soil water content on C fluxes given their nearly saturated nature. A number of previous studies have shown that temperature is an important driver of ecosystem respiration in similar alpine swamp meadows. For example, in the papers from Zhao et al. (2010) and Zhao et al. (2005), the authors showed that ecosystem respiration follows the exponential variation of soil temperature without considering soil water content. Zhu et al. (2020) also suggested that soil temperature plays the most important role in the change of monthly ecosystem respiration in the alpine wetland at Luanhaizi, northeastern Qinghai-Tibet Plateau. Therefore, in this study, we wanted to characterize and estimate the terrestrial C exchange while considering the potential effects of soil moisture. Meanwhile, it should be noted that Zhao et al. (2010) also noticed that the CO₂ emission rates decrease notably after rain events, and Zhu et al. (2020) confirmed that annual precipitation exhibits significant impact on variation of annual net C uptake. All these existing studies have suggested that C fluxes are related to water availability condition, but few studies have found that soil moisture explicitly affects respiration, and its decrease further reduces net C uptake in alpine swamp meadows, this finding could be an important factor for carbon modelling in the future.

Finally, the addition and further analysis of multiple years of data from new sites is always very important also in a more regional/global context - there is a generalized sparsity of in situ observations where their temporal and spatial coverage is very limited. We believe that any effort and addition to this generalized lack of in situ data will be useful for both flux communities such as FLUXNET and ASIAFLUX and modelling community in general.

2. The main drivers of NEE variation based on the different approaches are contradictory in this study. Net radiation is the leading factor affecting seasonal and annual variability of NEE based on machine learning approach (Fig. 5, Lines 231-250). However, the combined

effect of temperature and soil moisture change is the main factor influencing the annual variation of NEE in Section 4.3 and Table 2. The title of this manuscript only emphasizes the effect of soil moisture. Therefore, there are three different descriptions of the dominant factor of CO₂ sink in this manuscript. The mechanism underlying NEE variation needs to be more rigorously analyzed.

Thank you for this insightful comment. Our first draft didn't present the findings clearly and it was somewhat misleading. First of all, we should acknowledge that the original title did not fully reflect our conclusions, therefore we revised it as follows:

“Radiation, soil water content, and temperature interactions with carbon cycling in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau”.

In the S4.3, we intend to put our results into a broader context by comparing with other surrounding alpine swamp meadow sites to highlight the effects of the complex interactions between temperature and soil water content on carbon fluxes. In this section, we did not intend to explicitly disentangle the most important drivers for the NEE at annual scales among these different sites, given that detailed observations for net radiation are lacking for other sites. Such comparison highlights the importance of SWC (precipitation as a proxy for SWC as explicit SWC is not present in these either) in controlling NEE. This further validates our Random Forest (RF) findings regarding the drivers for the Re in our site (Figure 5). We should keep in mind that NEE is the difference between Re and GPP, environmental variables affecting Re and GPP could affect NEE indirectly (Song et al. 2011). We revealed that the main drivers for the GPP, Re and NEE are remarkably different. Net radiation is a key driver of seasonal and annual NEE and GPP, while soil water content is most important for Re at diurnal, seasonal and annual scales. Therefore, our findings from the RF analyses are not contradictory to the discussions regarding the difference in NEE and the potential influence factors among different sites.

However, your constructive comments point out an interesting scientific issue regarding the divergent drivers for the NEE dynamics at different time scales across different alpine swamp meadow sites. This will be an excellent point to be addressed in our future study.

3. A key conclusion of this study is that ecosystem respiration (Re) increases with decreasing soil water content (SWC). This is based on the comparisons of Re and SWC observations in the late growing seasons of 2014 and 2015 (Section 4.1). However, both SWC and Re in the late growing season are largest in 2017 during the observational period (2014-2017) (Figs. 2 and 4). Thus, the conclusion of “Section 4.1 Low soil moisture is associated with enhanced ecosystem respiration” (Lines 262-312) is likely to be unreliable. All observational years (2014, 2015, 2017 and 2018) are recommended to be considered in the analyses, rather than only two years (2014 and 2015).

Many thanks for this comment - we have not clearly described the comparisons in S4.1 between years in the text. Since C fluxes are affected by plant phenology and climate factors including temperature, soil moisture, and radiation simultaneously (Figure 5), in order to analyze the effects of single factor, ideally, other factors need to be identical or at least close (no significant differences). Based on this theory, we made our comparisons of specific time periods other than all the observation time. We now explicitly implemented the following text in S4.1, L293-298:

“Since C fluxes are affected by plant phenology and climate factors, including temperature, soil moisture, and radiation simultaneously (Fig. 5), to analyze the effect of a single factor, ideally, other factors need to be identical or at least close (no significant differences). Based on this theory and to better understand the underlying mechanisms around how SWC interacts with the C fluxes in the studied alpine swamp meadow ecosystem, we selected a specific group of data for further evaluation other than the entire observation time.”

As described in S2.4, L177-185, we chose explicitly 2014 and 2015 for comparison in S4.1 because there was not a significant change in temperature (<1%) between these two periods, while soil water content decreased significantly more (15.4%) in 2015 (see Table S2). Therefore, within this set of conditions we can compare the influence of soil water content reduction on Re, and its influence further over NEE. Similarly, in S4.2, we also chose 2014 and 2018 for comparison, because the temperature difference between these two periods was greatest (25%) while there was no significant difference in soil water content (0.1%) (see again Table S2), so we could isolate and study the impact of temperature increase over the C fluxes.

Although the soil water content and ecosystem respiration in 2017 were at their highest, the temperature was also higher than in 2015, so we cannot compare it with 2015 to study separately the influence of soil water content change on ecosystem respiration and further net C uptake. To clarify this, we included a more detailed explanation about this comparison in the revision, see S2.4, L177-185:

“To further analyse the effect of soil moisture, radiation, and temperature on C fluxes, we selected two groups of time stamps with a significant difference in SWC but almost identical Ta and Rn (i.e. late growing season of 2014 vs 2015) and a significant difference in Ta but almost identical SWC and Rn (i.e. late growing season of 2014 vs 2018). Additionally, in order to analyse the effect of annual temperature on C fluxes, we selected a group of time stamps with significant differences in Ta but almost identical SWC and Rn (i.e. 2017 vs 2014, and 2018 vs 2014). We made the comparison in each group to exclude the influence of plant phenology, which can influence C fluxes significantly. The magnitude of the differences between C fluxes in the same group were analysed by the independent-sample T-test method.”

Please note also that we included a new column with net radiation to complement the comparison and expand the discussion in order to improve the manuscript's clarity and align better with RF findings, this is including also radiation as suggested by REF#1 indirectly in the earlier point but also explicitly suggested by REF#2 later on.

Table S2. Seasonally aggregated environmental drivers and C fluxes in the late growing season of 2014, 2015, and 2018 and their relative difference between years.

Period	Ta (°C)	Rn (W m ⁻²)	SWC (%)	NEE (g C m ⁻²)	Re (g C m ⁻²)	GPP (g C m ⁻²)
2014 Late GS	6.8±2.6	93.2±49.4	80.7±4.1	-175.6	152.7	328.3
2015 Late GS	6.8±2.5	93.2±43.6	68.3±4.3	-141.6	191.9	333.5
2018 Late GS	8.5±3.4	97.7±47.6	80.8±3.8	-134.3	225.4	359.7
2015 - 2014	0.8%	-0.1%	-15.4%	-19.4%	25.7%	1.6%
2018 - 2014	25%	4.8%	0.1%	-23.5%	47.6%	9.6%

Note: late GS represents late (Aug. - Sep.) growing season.

4. Another key conclusion of this study is that warming leads to higher C losses rather than enhanced C uptake. This is based on the comparisons of Re and GPP observations in the late growing seasons of 2014 and 2018 (Section 4.2). Warming decreases NEE in late growing season but this does not indicate that warming decreases annual NEE. Recently, Wei et al. (2021) found that “plant uptake of CO₂ outpaces losses from permafrost and plant respiration on the Tibetan Plateau” at annual scale based on 32 eddy covariance sites in the Tibetan Plateau. Thus, the authors should more rigorously examine whether warming decreases net C sink of alpine swamp on the Tibetan Plateau.

This is a great point, thanks for bringing this up. In the revised manuscript, we added a new Table S4 to characterize the effect of temperature increase on net C uptake at annual basis and added more discussion accordingly. Based on a new variance analysis, soil water content in 2015 was significantly lower than in 2014, 2017, and 2018 ($p < 0.05$), while there was no significant difference in soil water content between 2014, 2017, and 2018, and there was also no significant difference in net radiation between these four years either. Therefore, we chose 2014, 2017, and 2018, when there was no significant difference in moisture conditions, to study the impact of temperature increase on C fluxes at an annual scale in the alpine swamp meadow. In fact,

according to our data there was a reduction of net C uptake in warmer 2017 and 2018, and an increase of GPP and Re in 2017 and 2018 compared with 2014 (the increase of GPP is relatively lower than Re in warmer 2017 and 2018). Therefore, these extended results indicate that a temperature increase can decrease the net C uptake on an annual scale in this site. We added a new Table S4 in the supplement material:

Table S4. Daily aggregated environmental drivers and C fluxes in 2014, 2017, and 2018.

Year	Ta (°C)	Rn (W m ⁻²)	SWC (%)	NEE (g C m ⁻²)	Re (g C m ⁻²)	GPP (g C m ⁻²)
2014	-0.9 ± 8.2	72.2 ± 52.9	47.8 ± 32.0	-240.3	561.1	801.4
2017	-0.5 ± 8.3	73.4 ± 53.7	50.2 ± 32.8	-117.6	959.3	1076.9
2018	-0.5 ± 9.1	74.7 ± 54.9	49.8 ± 31.6	-113.4	788.4	901.8
2017-2014	44.4%	1.7%	5.0%	-51.1%	71.0%	34.4%
2018-2014	44.4%	3.5%	4.1%	-52.8%	40.5%	12.5%

We further added the following text in Section 4.2, L337-343:

“To evaluate if this finding is also consistent at an annual scale, we further analyzed annual aggregated data. An annual comparison was made between 2014, 2017, and 2018 when SWC were found insignificantly different while temperatures in 2017 and 2018 were 44.4% higher than in 2014 (Table S4). Additionally, this 44.4% increase in Ta in 2017 and 2018 both led to stronger GPP and Re (Table S4). Although both GPP and Re increased, the intensity in Re was greater than GPP, indicating that warmer temperatures have a stronger impact on ecosystem respiration in this site, resulting in an approximately 50% decrease of the net C uptake (Table S4). ”

We thank REF#1 for pointing us towards this nice paper by Wei et al., 2021 “Plant uptake of CO₂ outpaces losses from permafrost and plant respiration on the Tibetan Plateau”. We benefited a lot from this paper. The conclusion from Wei et al. (2021) are based on data of 32 eddy covariance sites during 2002 to 2020, while our study only covers 4-years of year-round observations. Differences in time and space scales may help to explain the differences we found. In fact, in the figure 4 from Wei et al. (2021) there are specific sites from Haibei where RR_{NEP} (Response ratio of NEP to the warming rate) is smaller than 1.0, indicating a negative effect of

warming on the NEP. Potential site-specific differences together with the particularity of water condition in of alpine swamp ecosystem could be the possible reason. We have added more text in the discussion section S4.2, L354-360.

“Wei et al. (2021) also found that the uptake of C by plants will exceed the amount of C release under warmer and wetter climate conditions at an annual scale based on manipulative experiments and model simulations focused on the Tibetan Plateau. Their study is based on a longer-term trend while our study only covers 4-years of year-round observations thus site-specific differences in time and space scales may explain this variability. Nevertheless, such results indicating inconsistent ecosystem responses suggest that there are still large uncertainties regulating the responses of C fluxes to temperature variation and further work is still crucial. ”

Specific comments

1. Line 26, “-168.0 ±-62.5” may should be “-168.0 ± 62.5”.

Typo, thanks for picking it up. Changed accordingly.

2. Lines 75-76, “only a few experiments have been conducted to specifically characterise alpine swamp meadow ecosystem C dynamics.”, but no study focusing on alpine swamp C cycling is mentioned at here. The previous studies focusing on alpine swamp CO₂ fluxes in the Tibetan Plateau are recommended to be mentioned, such as alpine swamp CO₂ fluxes observations at Haibei (Zhao et al., 2005, 2010), Shenzha (Qi et al., 2021), Nam Co (Liu et al., 2020), and Huanhaizi (Zhu et al., 2020).

References

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Zhao, L., Li, Y.N., Zhao, X.Q., Xu, S.X., Tang, Y.H., Yu, G.R., Gu, S., Du, M.Y. and Wang, Q.X., 2005. Comparative study of the net exchange of CO₂ in 3 types of vegetation ecosystems on the Qinghai-Tibetan Plateau. *Chinese Science Bulletin*, 50 (16): 1767-1774. 10.1360/04wd0316.

Zhu, J., Zhang, F., Li, H., He, H., Li, Y., Yang, Y., Zhang, G., Wang, C. and Luo, F., 2020. Seasonal and interannual variations of CO₂ fluxes over 10 years in an alpine wetland on the Qinghai-Tibetan Plateau. *Journal of Geophysical Research-Biogeosciences*, 125 (11). 10.1029/2020jg006011.

Thank you very much for this very useful information concerning previous studies including sites focusing on alpine swamp CO₂ fluxes. Following the recommendation from REF#1 we have cited these previous studies and updated the description in the revised manuscript to further highlight the novelty of this study, S1, L75-78:

“Although many studies concerning ecosystem C dynamics on the QTP have focused on alpine meadow ecosystems (Saito et al., 2009; Zhao et al., 2005, 2010; Zhu et al., 2015b) and alpine swamp meadow ecosystems (Zhao et al., 2005, 2010; Qi et al., 2021; Liu et al., 2020; Zhu et al., 2020), the effect of SWC on the C uptake is still unclear as compared to that of temperature for alpine swamp meadow ecosystems. ”

Interactive comment on Biogeosciences Discuss., <https://doi.org/10.5194/bg-2021-193>

Interactive comment on “Effects of soil water content on carbon sink strength in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau” by Wei et al.

Received and published for Discussion: 10 October 2021

Response to Comments of Referee #2

Wei et al. did an interesting job about carbon balance of an alpine swamp meadow on the Qinghai-Tibet plateau, enriching the dataset of carbon fluxes in alpine wetland. This study highlights that decreasing in soil water content in this wetland ecosystem can stimulate ecosystem respiration significantly and weaken ecosystem carbon sink strength as a result. The writing is good and data processing is reasonable. Some major concerns in data interpretation should be addressed.

We are thankful for the reviewer’s insightful comments that have improved the manuscript. We have carefully considered the reviewer’s remarks and clarified our manuscript accordingly.

Major Comments:

1. In section 3.3, radiation has been identified as the main driver for the variation of GPP and NEE, and air temperature as the second force. However, in section 4.2 and 4.3, only the role of air temperature is discussed through the comparison of different study periods and different studies. This may confuse the reader because the main driver for NEE seems inconsistent in the Result and Discussion section. The authors may want to highlight the role of warming on carbon fluxes, but a statement or discussion on radiation is necessary to make the manuscript with good clarity.

Thanks for your useful comments. We therefore added the following text accordingly at the beginning of the Discussion section, L265-275:

“Since NEE is the difference between Re and GPP, environmental variables affecting Re and GPP could affect NEE indirectly (Song et al., 2011). Radiation affects the magnitude of plant

photosynthesis and controls temperature, one of the key factors related to C fluxes. This suggests that abundant radiation benefits photosynthesis and respiration and thus directly affects the C sink strength in the alpine wetland ecosystem of the Qinghai Lake (Cao et al., 2017). Niu et al. (2017) show that 99% of the interannual variation of NEE in an alpine swamp meadow can be well explained by temperature conditions, precipitation and radiation. The results of this study demonstrate that ecosystem C sequestration is regulated not only by radiation and temperature but also by soil moisture in the alpine swamp meadow site studied herein. Given there is no significant difference in net radiation between the four years we studied, the effects of soil water content and temperature on C fluxes on diurnal, seasonal, and annual scales are therefore discussed in detail below.”

Please note that further implementations including radiation numbers and discussion were made in the following points. But please also note that also the original title has been revised to reflect REFS#2 and #1 points:

“Radiation, soil water content, and temperature interactions with carbon cycling in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau”.

2. In section 4.1, two groups of data (the late growing season of 2014 and 2015) were used to analyze the effects of soil moisture on carbon fluxes. Although the authors declared that air temperature and phenology of these two periods were comparable, I could not find information about radiation of these two periods. As radiation has already been concluded as a main driver for the variation of GPP and NEE in the Result section, it is critical to build the comparison based on comparable radiation, or the results can be pointless. The same issue goes for the comparison between the late growing season of 2014 and 2018 in section 4.2 (L322-324).

We appreciate this comment about the comparisons in S4.1 and S4.2. The REF#2 is right, we did not present the values of net radiation (Rn) in the original comparison because there was no significant difference in Rn between these compared periods. However, following the referees' advice, we have now added radiation number to Table S2 that are addressed and discussed in S4.1 and S4.2.

Table S2. Seasonally aggregated environmental drivers and C fluxes in the late growing season of 2014, 2015, and 2018 and their relative difference between years.

Period	Ta (°C)	Rn (W m ⁻²)	SWC (%)	NEE (g C m ⁻²)	Re (g C m ⁻²)	GPP (g C m ⁻²)
2014 Late GS	6.8±2.6	93.2±49.4	80.7±4.1	-175.6	152.7	328.3
2015 Late GS	6.8±2.5	93.2±43.6	68.3±4.3	-141.6	191.9	333.5
2018 Late GS	8.5±3.4	97.7±47.6	80.8±3.8	-134.3	225.4	359.7
2015 - 2014	0.8%	-0.1%	-15.4%	-19.4%	25.7%	1.6%
2018 - 2014	25%	4.8%	0.1%	-23.5%	47.6%	9.6%

Note: late GS represents late (Aug. - Sep.) growing season.

We further edited the text in S2.4, L177-185 to include radiation not only in this comparison, but also in an additional test looking specifically at annual data following REF#1 comment, number 4 (covered above). REF#1 was not convinced about the conclusion addressed in S4.2 regarding “*Temperature increase leads to higher C losses rather than enhanced C uptake*” - the referee suspected that our claim about warming decreasing NEE in late growing season do not necessarily indicate that warming can decrease annual NEE. Therefore, in order to address this point, we also added a new test to confirm our previous finding (S2.4, L177-185):

“To further analyse the effect of soil moisture, radiation, and temperature on C fluxes, we selected two groups of time stamps with a significant difference in SWC but almost identical Ta and Rn (i.e. late growing season of 2014 vs 2015) and a significant difference in Ta but almost identical SWC and Rn (i.e. late growing season of 2014 vs 2018). Additionally, in order to analyse the effect of annual temperature on C fluxes, we selected a group of time stamps with significant differences in Ta but almost identical SWC and Rn (i.e. 2017 vs 2014, and 2018 vs 2014). We made the comparison in each group to exclude the influence of plant phenology, which can influence C fluxes significantly. The magnitude of the differences between C fluxes in the same group were analysed by the independent-sample T-test method.”

And in S4.2, L337-343:

“To evaluate if this finding is also consistent at an annual scale, we further analyzed annual aggregated data. An annual comparison was made between 2014, 2017, and 2018 when SWC were found insignificantly different while temperatures in 2017 and 2018 were 44.4% higher than in 2014 (Table S4). Additionally, this 44.4% increase in Ta in 2017 and 2018 both led to stronger GPP and Re (Table S4). Although both GPP and Re increased, the intensity in Re was greater than GPP, indicating that warmer temperatures have a stronger impact on ecosystem respiration in this site, resulting in an approximately 50% decrease of the net C uptake (Table S4). ”

3. The comparison between the late growing season of 2014 and 2015 shows that drought in 2015 stimulated Re noticeably (L279-282). However, the differences in GPP between these two periods are not significant (L290), the authors should be more careful to make the statement that high soil water content would suppress GPP (L287-295). Figure 5 shows that the contribution of soil water content to the variation of GPP is small on all the time scales.

The referee is fully right, thanks for pointing to this issue. We agree that our data do not support the potential suppressing effect of SWC on GPP, as since this paragraph is not central to the overall storyline, we have removed the text below as well as the references from the revised manuscript to improve clarity:

“There is evidence that excessive soil water can negatively affect plant physiological and ecological processes by, for example, insufficient supply of metabolic substrates and the production of toxic substances (Jackson and Colmer, 2005), which may reduce the overall plant photosynthetic efficiency (Xu and Zhou, 2011). Although there was no significant difference between the late growing season GPP of 2015 and 2014 in terms of both daily accumulated GPP and diurnal rates of GPP, the decline observed in SWC during September 2015 (the driest month with 65.1% SWC) led to a 11% increase of daily accumulated GPP (Fig. S1; Table S3). The excess of SWC in 2014 caused an inundation of the aboveground plant domain, which also likely contributed to the lower value in GPP by reducing plant photosynthetic efficiency (Cronk et al., 2001; Hirota et al., 2006). Since the increase of GPP

could not offset the increase in Re, September 2015 and the late growing season of 2015 experienced a lower C sink strength.”

4. The authors concluded that warming would weaken carbon sink strength in this alpine swamp meadow ecosystem because it would increase Re more than GPP. The authors also pointed out that other studies have reported opposite results that warming would stimulate GPP more than Re. However, these studies were conducted in different ecosystems, such as Arctic marshlands and Arctic tundra (L334-341). A discussion that focuses on alpine swamp meadow would be more worthwhile for understanding the effects of climate warming on carbon balance of alpine wetland ecosystem.

Thank you very much for comment. Now we have now removed the discussion around different ecosystems and replaced it with the following paragraph in S4.2, L352-360:

“ Liu et al. (2018) concluded that warming has a significant inhibitory effect on GPP and minor effect on Re, resulting in a weaker carbon sequestration capacity of their studied alpine wetland ecosystem. Wei et al. (2021) also found that the uptake of C by plants will exceed the amount of C release under warmer and wetter climate conditions at an annual scale based on manipulative experiments and model simulations focused on the Tibetan Plateau. Their study is based on a longer-term trend while our study only covers 4-years of year-round observations thus site-specific differences in time and space scales may explain this variability. Nevertheless, such results indicating inconsistent ecosystem responses suggest that there are still large uncertainties regulating the responses of C fluxes to temperature variation and further work is still crucial. ”

Specific Comments:

1. The font of ‘CO2’ in the draft should be the same as others.

Corrected accordingly.

2. *Potentilla* in Line 96 and *Kobresia* in L121 should be italic.

Thanks! Corrected accordingly.

3. L174-175: what do you mean by ‘this is one random forest per hour of the day, day of the year and year, respectively’?

We apologize for not been clear. In order to run random forest (RF) diurnally, seasonally and annually, we needed to aggregate the data as revised below in S2.4, L175-176:

“we ran multiple random forests with growing season data binned per hour of the day, day of the year and year, respectively (Table S1). ”

4. Figure 4 and Figure 5 should have tags.

Changed accordingly.

5. L373-379: The stronger C sink strength is first attribute to saturated soil condition rather than lower temperature, but then to higher precipitation and lower temperature.

Thank you very much for pointing this out. As the rest of sites included in Table 2 do not have all available SWC data, in S4.3 we used precipitation as a proxy for soil moisture (see L378-379:

“we examined the temperature and precipitation (as a proxy for SWC) impacts on NEE (Liu et al., 2016).”). This reasoning is consistent with our data; for example, in Figure 2b it is clear that when precipitation is low in 2014 in e.g. August and October compared to other years, SWC decreased significantly.

We further clarified in the text S4.3, L393-398:

“This is likely a result of the inhibiting effects of the nearly-saturated soil condition over soil respiration rather than by the lower temperatures (Sun et al., 2021). Therefore, in permanently or seasonally inundated swamp meadows, high SWC may have triggered lower C loss rates further benefiting C preservation. At our site, the higher C sink strength was likely attributed to higher precipitation (and therefore higher SWC) and lower temperature, which created colder and more humid conditions than other sites (Table 2). ”

6. The data supports the opinion that soil water content is a key control on ecosystem respiration, but soil water content does not affect the overall C sink strength (i.e., NEE) directly.

Agreed, we deleted the words “*and the overall C sink strength*” from the conclusion. Now the sentence was changed to “*Soil moisture, however, has the largest influence over Re variability on diurnal, seasonal, and interannual scales, suggesting that soil water content is a key control on ecosystem respiration.* ”.

Interactive comment on Biogeosciences Discuss., <https://doi.org/10.5194/bg-2021-193>

1 Radiation, soil water content, and temperature
2 interactions with carbon cycling in an alpine swamp
3 meadow of the northeastern Qinghai-Tibet Plateau ~~Effects~~
4 ~~of soil water content on carbon sink strength in an alpine~~
5 ~~swamp meadow of the northeastern Qinghai-Tibet~~
6 ~~Plateau~~

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25 **Abstract.** Predicted intensified climate warming will likely alter the ecosystem net carbon (C) uptake of
26 the Qinghai-Tibet Plateau (QTP). Variations in C sink/source responses to climate warming have been
27 linked to water availability; however, the mechanisms by which net C uptake responds to soil water
28 content in ~~water~~-saturated swamp meadow ecosystems remain unclear. To explore how soil moisture and
29 other environmental drivers modulate net C uptake in the QTP, field measurements were conducted using
30 the eddy covariance technique in 2014, 2015, 2017, and 2018. The alpine swamp meadow presented in
31 this study was a consistent and strong C sink of CO₂ ($-168.0 \pm -62.5 \text{ gC m}^{-2} \text{ y}^{-1}$, average \pm standard
32 deviation) across the entire 4-year study period. A random forest machine-learning analysis suggests that
33 the diurnal, seasonal, and annual variations of net ecosystem exchange (NEE) and gross primary
34 productivity (GPP) were controlled by temperature and ~~net solar~~ radiation. Ecosystem respiration (Re),

35 however, was found mainly regulated by the variability of soil water content (SWC) at different temporal
36 aggregations followed by temperature, the second contributing driver. We further explored how Re is
37 controlled by nearly saturated soil moisture and temperature comparing two different periods featuring
38 almost identical temperatures and significantly differences on SWC and vice versa. Our data suggest
39 that, despite the relatively abundant water supply, periods with a substantial decrease of SWC or increase
40 of temperature produced higher Re and therefore weakened~~lowering~~ the C sink strength. Our results
41 reveal that nearly saturated soil conditions during the warmgrowing seasons can help ~~to~~ -maintain lower
42 ecosystem respiration rates and thus enhance the overall C sequestration capacity in this alpine swamp
43 meadow. We argue that climate warming in the future could change soil hydrological conditions of alpine
44 swamp meadow through degrading permafrost or accelerating thawing-freezing cycling, therefore,
45 having a direct effect on soil respiration and subsequent effect on ecosystem C sink magnitude.~~e.changes~~
46 ~~in soil hydrological conditions induced by a warming climate near permafrost (or seasonal frozen layers)~~
47 ~~may affect the C sink magnitude of wet and cold ecosystems through changes in soil hydrology and the~~
48 ~~subsequent effect on respiration losses-~~

49 1. Introduction

50 Wetlands play a significant role in the global carbon (C) cycle due to ~~the~~a large amount of C stored in
51 their soils. The Qinghai-Tibet Plateau (QTP), with an average altitude of over 4,000 m a.s.l., has
52 approximately 10×10^4 km² of natural wetlands, of which $\sim 50\%$ (4.9×10^4 km²) are alpine swamp
53 meadows. These ecosystems are predominantly located in permafrost areas and are typically soil
54 nutrient-rich and water-logged (Bai et al., 2019; Zhao et al., 2005). Climate change and human
55 disturbance can have profound consequences on permafrost regions (Biskaborn et al., 2019) and
56 significantly impact their hydrological regimes (Lafrenière and Lamoureux, 2019). Hydrological regimes
57 have an important role in controlling wetland functioning (Bohn et al., 2007; Christensen et al., 2003),
58 and the changes of hydrological regimes may put the wetland functioning of the QTP under pressure
59 (Hruby, 1999; Woodward and Wui, 2001; Foti et al., 2013).

60 The QTP is forecasted to be warmer and wetter in the future (Chen et al., 2015; Cheng et al., 2011).
61 Warming may accelerate the microbial breakdown of alpine soil organic C and subsequently increase
62 CO₂ emissions (Zhu et al., 2015a). Warming could also improve C sequestration capacity by enhancing

63 the photosynthetic inputs and growth rates of alpine plants (Fu et al., 2015). Therefore, the potential
64 increase of CO₂ emissions due to warming in alpine regions could be partially offset by enhanced C
65 uptake (Schuur et al., 2009), triggering different net C uptake responses to climate warming. For
66 example, an increase in temperature in the QTP has been associated with net C sinks in the Zoige alpine
67 wetlands (Kang et al., 2014) but also with net C sources in the Damxung alpine swamp meadow (Niu et
68 al., 2017).

69 According to recent studies in QTP alpine grasslands, water conditions such as soil water content (SWC)
70 can be a key factor that changes water-use patterns and ecophysiological characteristics of alpine plants
71 (Wu et al., 2019) and modulate the warming-mediated increase of ecosystem C uptake (Ganjurjav et al.
72 2016; Peng et al. 2014). Ecosystem C processes such as net C uptake and soil respiration may increase
73 with SWC in dry environments, and decrease in water-logged environments (Quan et al., 2019; Taylor
74 et al., 2017). Warming in conjunction with increased precipitation can turn an ecosystem from net source
75 to a sink of C (Zhao et al., 2019), increasing both photosynthesis and respiration rates during warmer and
76 wetter years (López-Blanco et al., 2017, 2018). However, when warming occurs in soils associated with
77 low moisture, soil drought can change ecosystems from C sinks to sources (Ganjurjav et al., 2017).
78 Studies in QTP alpine meadows have indicated that warming significantly stimulates ecosystem net C
79 uptake in wet years but does not affect ecosystem net C uptake in dry years because the positive effects
80 of warming on net C uptake are compensated by the negative effects of lower soil moisture (Peng et al.,
81 2014).

82 Although many studies concerning ecosystem C dynamics on the QTP have focused on ~~the~~ alpine
83 meadow ecosystems (Saito et al., 2009; Zhao et al., 2005, 2010; Zhu et al., 2015b), ~~and only a few~~
84 ~~experiments have been conducted to specifically characterise~~ alpine swamp meadow ecosystems (Zhao
85 et al., 2005, 2010; Qi et al., 2021; Liu et al., 2020; Zhu et al., 2020) C dynamics, the effect of SWC on
86 the C uptake is still unclear as compared to that of temperature for alpine swamp meadow ecosystems.

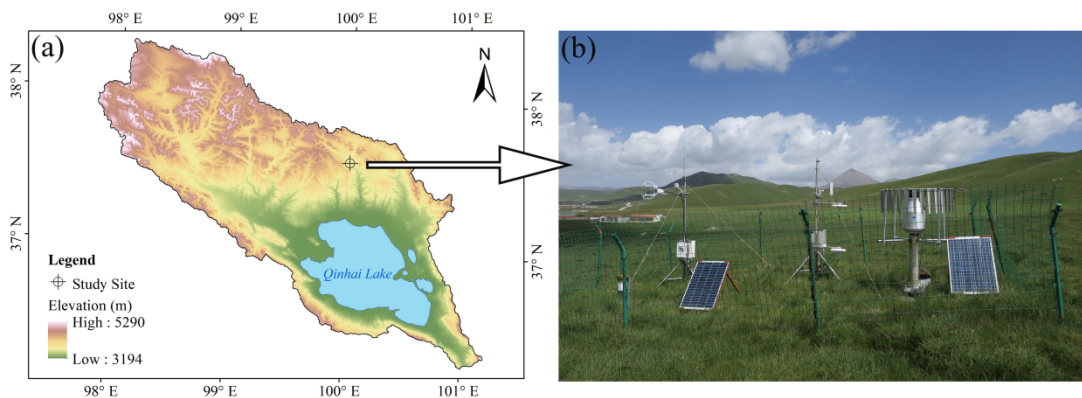
87 Alpine swamp meadow ecosystems are characterised by high SWC, and it remains unclear whether a
88 decrease in SWC would alleviate the stress from saturated water on net C uptake or aggravate drought
89 effects. Given that these ecosystems have high soil moisture compared to typical alpine meadows, the
90 effects of warming and water availability ~~drying~~ on the future C sink strength of the QTP ecosystem
91 remain uncertain ~~unknown~~. These questions and uncertainties require a detailed investigation to

92 understand wetland C source/sink processes and their future C sink strength variations (sign and
93 magnitude). Therefore, the objectives of this study are to (i) quantify the diurnal, seasonal, and annual
94 variations of net ecosystem exchange (NEE), gross primary productivity (GPP), and ecosystem
95 respiration (Re), (ii) identify and quantify the relative importance of different key environmental drivers
96 contributing to the variability observed of NEE, Re, and GPP, and (iii) analyse how these C fluxes
97 respond to soil water availability, ~~key environmental drivers such as~~ temperature, and radiation ~~or soil~~
98 ~~moisture~~ variation in a QTP alpine swamp meadow.

99 2. Materials and methods

100 2.1 Site description

101 The study site (37°35.75' N, 100°00.47' E, 3,571 m a.s.l.) is located in Gangcha County, Qinghai
102 Province, in the north-eastern part of the QTP (Fig. 1(a)). The mean annual temperature and precipitation
103 measured at the Gangcha National Weather Station were 0.1 °C and 389.4 mm between 1982 and 2011,
104 respectively (Zhang et al., 2016). The site area has seasonal permafrost featuring frozen soils between
105 January–March and November–December for a total of about 125–135 days (Zhang et al., 2014). The
106 dominant species of the alpine swamp meadow ecosystem ~~is~~ *Kobresia pygmaea*, accompanied by
107 ~~with~~ *Saussurea pulchra*, *Polygonum viviparum*, and *Potentilla saundersiana*. The average plant height
108 at the experimental site is 7.4 ± 1.5 cm, with a $97 \pm 2\%$ coverage. Our study defined the growing season
109 as the period between June and September. The early (June–July) and late (August–September) growing
110 seasons, as they will be henceforth referred to, correspond with the early growing season and late growing
111 seasons, respectively.



112

113 **Figure 1. (a) Location of the study site in the Qinghai Lake basin in the northeastern part of the QTP. (b)**
 114 **Eddy covariance system measuring water and CO₂ fluxes between the land surface and the atmosphere in the**
 115 **alpine swamp meadow.**

116 2.2 Field measurements

117 An Eddy Covariance (EC) system was installed at the study site (Fig. 1b) to measure the CO₂ fluxes at a
 118 sampling frequency of 10 Hz from 2014 to 2018. Data for 2016 was missing due to equipment
 119 malfunction. The EC system included an open-path CO₂/H₂O infrared gas analyser, which quantified
 120 fluctuations in water vapour and CO₂ fluxes. A 3-D sonic anemometer was ~~also~~ installed at a 2.0 m height
 121 above ground to directly measure horizontal and vertical wind velocity components (u, v, and w). C flux
 122 data were recorded with a data logger (Campbell Scientific Inc.). An automated meteorological station
 123 was installed near the EC station to measure meteorological variables such as air temperature (Ta; °C),
 124 precipitation (P; mm), net radiation (Rn; W m⁻²), wind speed (WS; m s⁻¹), wind direction (WD; °), relative
 125 humidity (RH; %), and vapour pressure deficit (VPD; hPa). The meteorological data were collected at
 126 one-minute intervals and subsequently resampled at 30-minute timesteps to keep pace with the EC data.
 127 More details on the in-situ instrument specifications are summarised in Table 1.

128 **Table 1 Information about the sensors installed in the alpine swamp meadow.**

	Sensor Names	Sensor type	Installation height/depth	Manufacturer
Eddy Covariance	Open-path CO ₂ /H ₂ O infrared gas analyser	EC150	2.0 m	Campbell, US
	Three-dimensional sonic anemometer	CSAT3	2.0 m	Campbell, US
Meteorological observation	Net radiation	NR Lite	1.8 m	Kipp&Zonen, Netherlands
	Wind speed/direction	034B	2m	
	Air temperature/humidity	083E-1-6	0.5 m, 1.5 m	MetOne, US
	Atmosphere pressure	PTB110	In data acquisition box	Vaisala, Finland
	Rain-gauge	7852M-AB	0.7 m	Davis, US

129 Soil water content (SWC; %) were measured at depths of 10, 20, 40, 60, and 100 cm from the soil surface
 130 with EC-H₂O sensors (Decagon Devices, USA) at a 10-minute frequency. The precision of the EC-H₂O
 131 sensors for soil moisture measurements was $\pm 0.03 \text{ m}^3 \text{ m}^{-3}$. As the roots of *Kobresia* meadows are mainly
 132 distributed within the top 20 cm of soil, we focused only on the variation of SWC in the top 20 cm of the
 133 soil.

134 2.3 Eddy Covariance data processing

135 The half-hourly NEE of CO₂ data was calculated using the EddyPro software (version 5.2, LI-COR) from
136 the 10-Hz raw data. During the calculation, three-dimensional rotation was used to correct the data by
137 removing the effects of instrument tilt irregularity on airflow (Wilczak et al., 2001). Webb, Pearman, and
138 Leuning (WPL) (Webb et al., 1980) correction was applied to calculate the averages of CO₂ covariance,
139 rectifying the air density variations induced by heat and water vapour. The half-hourly flux data were
140 quality-checked based on several filtering algorithms, including: (1) the rejection of outliers in sonic
141 temperature, water vapour density, and CO₂ density (Li et al. 2008; Liu et al., 2011), (2) the elimination
142 of data one hour before and after precipitation events (Li et al. 2008; Liu et al., 2011), (3) the removal of
143 negative NEE during the non-growing season (from November to March) (Cao et al., 2017; Qi et al.,
144 2021) attributed to the self-heating effect from EC instruments (Cao et al., 2017), and (4) the exclusion
145 of measurements with weak turbulence conditions at night time. The weak turbulence periods were
146 identified by bootstrapping friction velocity (u^*) thresholds, as described by Papale et al. (2006). This
147 approach effectively divided the data into 4-year and 7-temperature subsets with similar micro-
148 meteorological conditions (except for u^*). The u^* thresholds (5%, 50%, and 95% of bootstrapping) were
149 calculated specifically per year and temperature subset.

150 Based on those different subsets, we gap-filled and partitioned NEE (into GPP and Re) to spread the
151 uncertainty variability that emerged from the different u^* thresholds, similar to López-Blanco et al.
152 (2017). All missing data were marked as -9999 (no data). Negative and positive NEE values represent
153 sink and source of C, respectively. Additionally, a standardised mechanism to fill NEE gaps is needed
154 for adequate data processing (Moffat et al., 2007). Therefore, this study adopted the method described
155 by López-Blanco et al. (2020) using the marginal distribution sampling (MDS) algorithm in the
156 REddyProc gap-filling tool (Reichstein et al., 2016), which was readapted from Reichstein et al. (2005).
157 Finally, NEE was separated into GPP and Re applying the REddyProc partitioning algorithm (Reichstein
158 et al., 2016) for further analyses. This partitioning method is based on the exponential regression of night-
159 time respiration with temperature using the Lloyd-Taylor-Function (Lloyd and Taylor, 1994). Night-time
160 periods were selected via current combined solar radiation and potential radiation thresholds based on
161 the exact solar time, latitude, and longitude. REddyProc estimates the temperature sensitivity from a
162 short-term period, and based on this short-term temperature sensitivity, it estimates the reference

163 temperature in the continuous period of the entire dataset. These estimates were then used to calculate
164 R_e for day-time and night-time, while GPP was estimated based on the difference between NEE and
165 R_e .

166 2.4 Identifying the importance of environmental drivers

167 In order to characterize how environmental conditions impact diurnal, seasonal, and interannual
168 variability of NEE, GPP and R_e at this alpine swamp meadow, we used a novel method based on machine
169 learning.

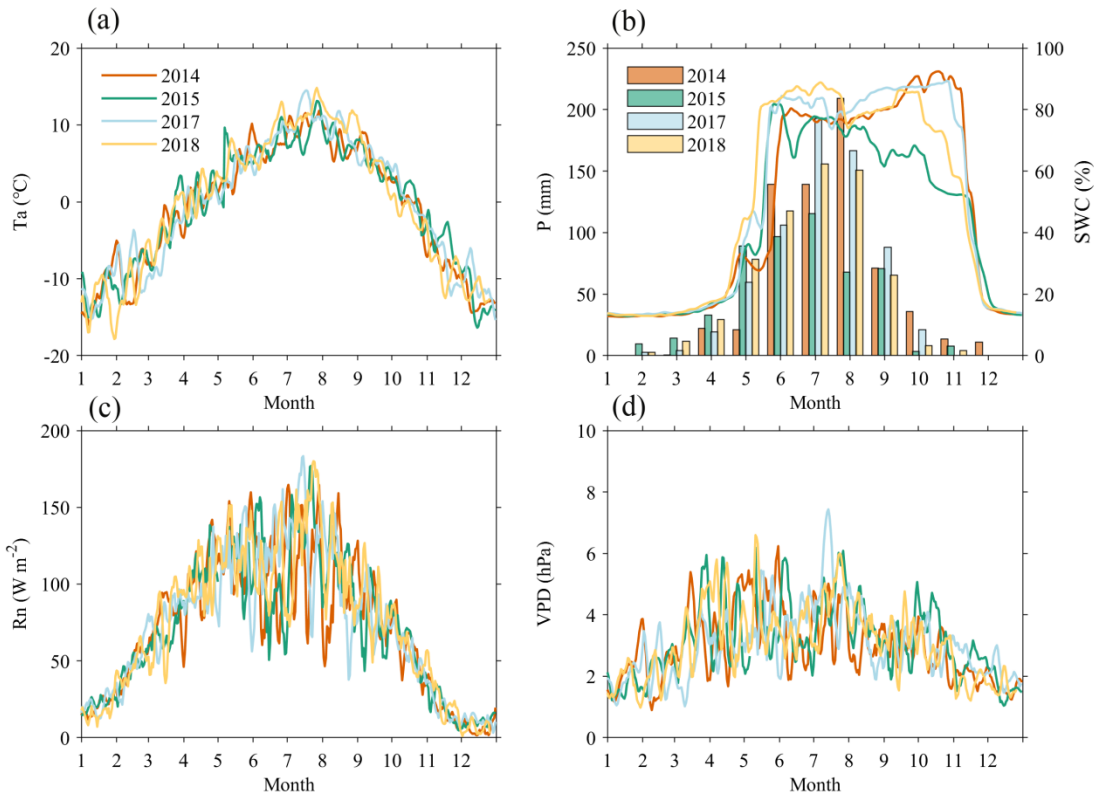
170 Random forest (Breiman, 2001) is a machine learning technique that can be used to quantify and interpret
171 the contribution of environmental drivers (covariates) to the variability of different C fluxes (response
172 variables) by combining multiple individual regression trees. This technique has been increasingly
173 utilized to upscale global C fluxes from eddy covariance data (Zeng et al., 2020) but also to evaluate
174 controls on C cycle processes (Zhang et al., 2017; López-Blanco et al., 2017, 2020). Here, we calculate
175 the relative importance of air temperature (T_a), net radiation (R_n), soil water content (SWC) and vapour
176 pressure deficit (VPD) controlling the C sink strength, photosynthesis, and respiration variability. This
177 random forest algorithm constructs multiple (1000 in this analysis) decision trees during training time
178 with different random subsamples (with replacement) from the same input training dataset. In each
179 cluster classified by random forest, the algorithm generates a multiple linear regression to characterize
180 different C fluxes as a function of environmental drivers (López-Blanco et al., 2017, 2020). This
181 algorithm version (Pedregosa et al., 2011) estimates the relative importance of each covariate between 0
182 and 100%, which correspond to the fraction of decision participating during data clustering. We used the
183 random forest algorithm to evaluate the diurnal, seasonal, and annual patterns of the relative importance
184 of T_a , R_n , SWC, and VPD responsible for the variability of C fluxes. We used data from the June-
185 September period aggregated per hour, and we ran multiple random forests with growing season data
186 binned per hour of the day, day of the year and year, respectively (Table S1). ~~per hour of the day, day of~~
187 ~~the year, and year, respectively (Table S1).~~
188 ~~In order to~~ To further analyse the effect of soil moisture, radiation, and temperature ~~SWC~~ on C fluxes, we
189 selected two groups of time stamps with a significant difference in SWC but almost identical T_a and
190 R_n same T_a (i.e. late growing season of 2014 vs 2015) and a significant difference in T_a but almost

191 identical SWC and Rn (i.e. late growing season of 2014 vs 2018). Additionally, in order to analyse the
192 effect of annual temperature on C fluxes, we selected a group of time stamps with significant differences
193 in Ta but almost identical SWC and Rn (i.e. 2017 vs 2014, and 2018 vs 2014). We made the comparison
194 in each group to exclude the influence of plant phenology, which can influence C fluxes significantly.
195 The magnitude of the differences between C fluxes in the same group~~during the same periods from~~
196 ~~different years~~ were analysed by the independent-sample T-test method.

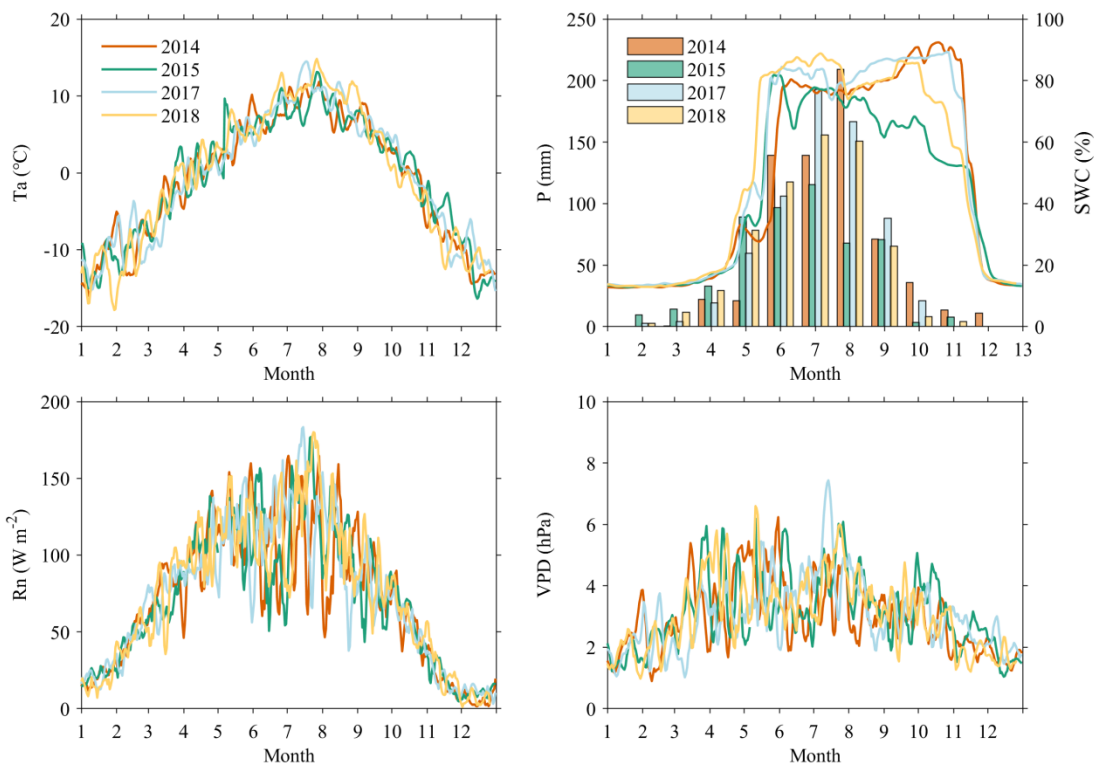
197 3. Results

198 3.1 Meteorological variability

199 Mean daily meteorological variables including Ta, P, SWC, and Rn exhibited evident seasonal variability
200 except for VPD; these variables increased progressively in the early growing seasons, reached their
201 maximum in July and decreased gradually afterwards (Fig. 2(a), (b), (c), (d)). Air temperature during the
202 growing season was 7.7 ± 2.6 , 7.4 ± 2.6 , 8.5 ± 2.9 , and 9.2 ± 3.3 °C in 2014, 2015, 2017, and 2018,
203 respectively while precipitation totalled 662.8, 521.4, 661.2, and 624.3 mm for the same years, falling
204 primarily during the growing season. The precipitation measured during the late growing season of 2015
205 was only half of the amount measured in 2014, 2017, and 2018. The lower precipitation regime led to a
206 marked decline in SWC, making the late growing season of 2015 the driest period among all growing
207 seasons during the study period. The greatest difference in SWC occurred in the late growing season of
208 2014 and 2015, when Ta was the same at 6.8 ± 2.6 and 6.8 ± 2.5 °C, respectively. Compared to 2014,
209 SWC decreased by 15.4% in the late growing season of 2015 (Fig. 2(b); Table S2). Meanwhile, SWC in
210 the late growing season of~~during the same period, the SWC in~~ 2014 and 2018 was almost identical (80.7
211 ± 4.1 and 80.8 ± 3.8 , respectively), but the temperature difference was the largest (25%) compared to
212 any other year (Fig. 2(a); Table S2).



213

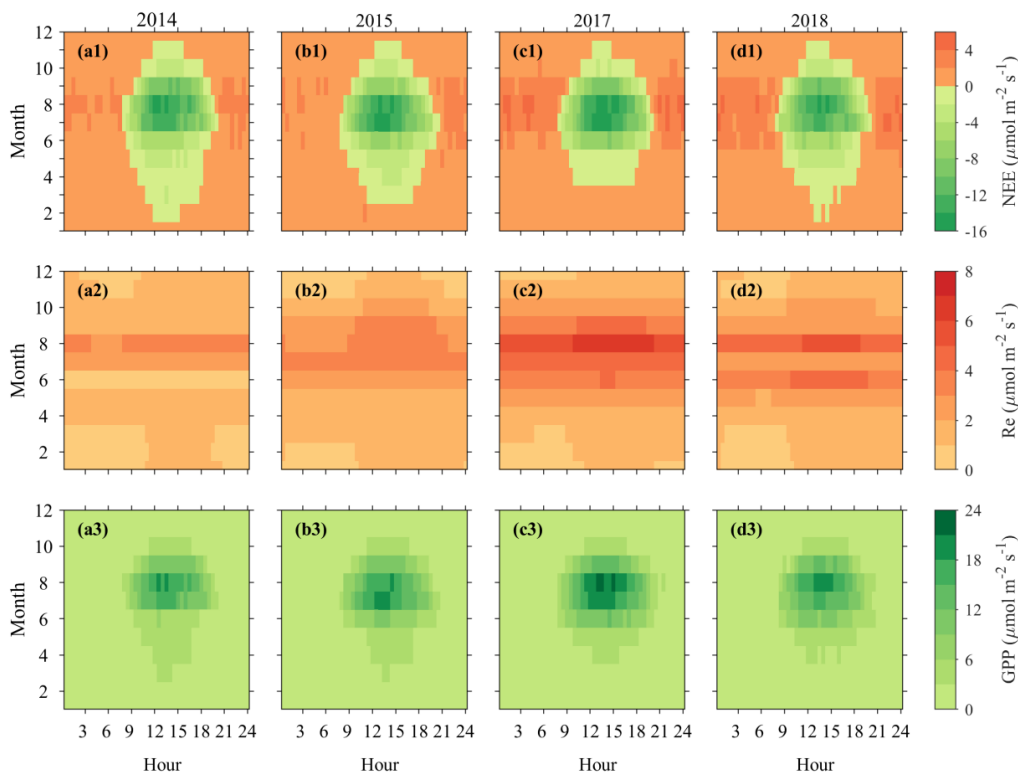


214

215 **Figure 2. Five-day moving average meteorological variables (T_a , P , SWC , R_n , VPD) in the studied swamp**
 216 **alpine meadow.**

217 **3.2 Diurnal, seasonal, and annual variability of CO₂ fluxes**

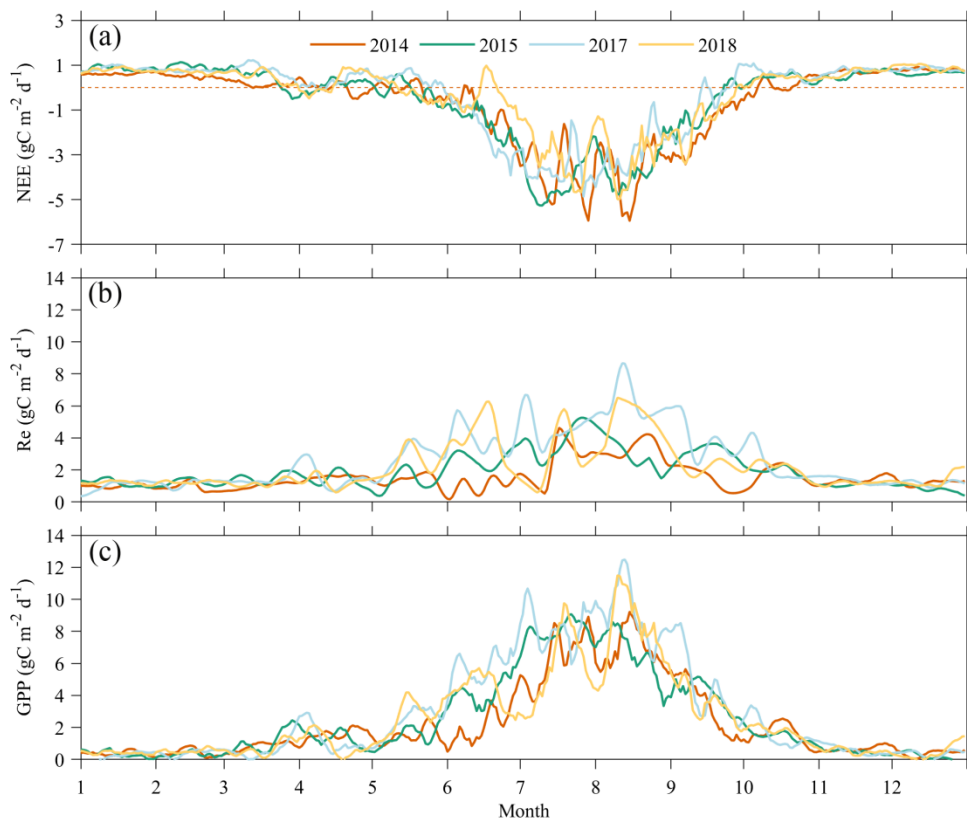
218 During the growing season, NEE (Fig. 3, (a1)–(d1)) and GPP (Fig. 3, (a3)–(d3)) featured a clear peak of
 219 the diurnal variations; both fluxes reached their summit between 12:00 and 14:00 local time. Re,
 220 however, presented a lower daily variability. The rates of NEE, Re, and GPP during the growing season
 221 averaged -2.3 ± 0.3 , 3.2 ± 1.0 , and $5.5 \pm 0.9 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively, for the entire study period (2014,
 222 2015, 2017 and 2018). For the late growing season, the lowest rate of net C uptake was measured in 2015
 223 ($-10.0 \mu\text{mol m}^{-2} \text{s}^{-1}$), whereas 2014 ($-12.4 \mu\text{mol m}^{-2} \text{s}^{-1}$), 2017 ($-12.2 \mu\text{mol m}^{-2} \text{s}^{-1}$), and 2018 ($-12.5 \mu\text{mol}$
 224 $\text{m}^{-2} \text{s}^{-1}$) exhibited more negative NEE values (i.e. stronger net C uptake rate). Between the late growing
 225 season in 2015 and the late growing seasons in 2014, 2017 and 2018, there was a significant difference
 226 in the rates of Re ($p < 0.01$) while no significant difference was found in GPP variability ($p > 0.05$),
 227 suggesting that Re may be the component causing the difference observed in NEE. Specifically, the rates
 228 of Re in the late growing season of 2014 and 2015 were 2.4 ± 0.2 and $3.0 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively,
 229 which indicated that during 2015 drier conditions generated a 25% higher Re compared to 2014. For the
 230 same periods, the rate of Re in warmer 2018 was $3.5 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$, which was also significantly
 231 higher than in 2014 ($2.4 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$).



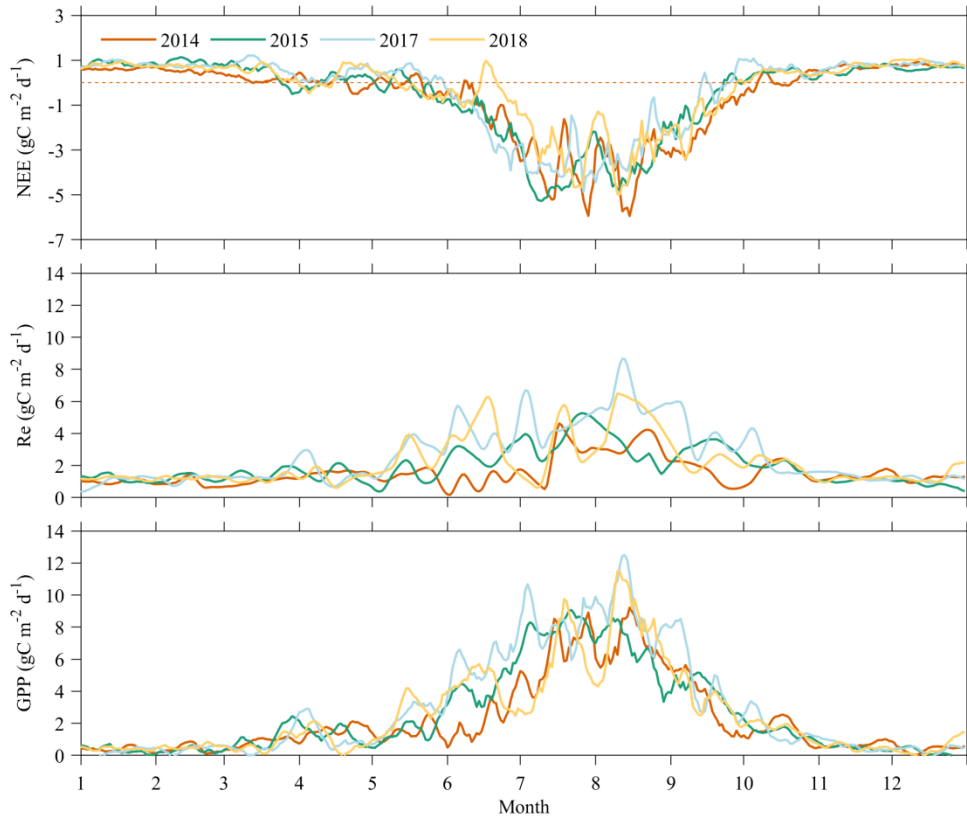
232

233 **Figure 3. Diurnal variability of gap-filled NEE and partitioned Re and GPP in 2014, 2015, 2017 and 2018.**

234 NEE, Re, and GPP also exhibited a strong seasonal variability; the C fluxes gradually increased from
 235 low values in early June to maximum values in the middle of the growing season (late July to early
 236 August on average), followed by a decrease towards the end of the growing season (Fig. 4(a), (b), (c)).
 237 The C sink strength across the growing season was found lowest in 2018, followed by 2017, whereas
 238 2014 and 2015 exhibited relatively higher values (Fig. 4(a)). Notably, the accumulated Re in the late
 239 growing season was significantly higher in 2015 compared to 2014 ($P < 0.05$), when there was no
 240 significant difference in GPP (Fig. 4(b); Table S2). Moreover, the late growing season of 2015 witnessed
 241 the lowest SWC while keeping the same Ta compared to the same period in 2014 (Fig. 2(b); Table S2).
 242 The substantial decline of SWC observed in 2015 appeared to be responsible for the weaker observed C
 243 sink strength. On the other hand, Ta in the late growing season of 2018 was the highest for the same
 244 period among all the years while the SWC remained the same as 2014 (Fig. 2(a)). The late growing
 245 season of 2018 showed overall higher GPP, Re, and lower net C uptake than 2014. Significantly higher
 246 Re in 2018 caused by higher-warmer temperatures eventually led to a decrease of the C sink capacity
 247 (Fig. 4(a), (b), (c); Table S2).



248



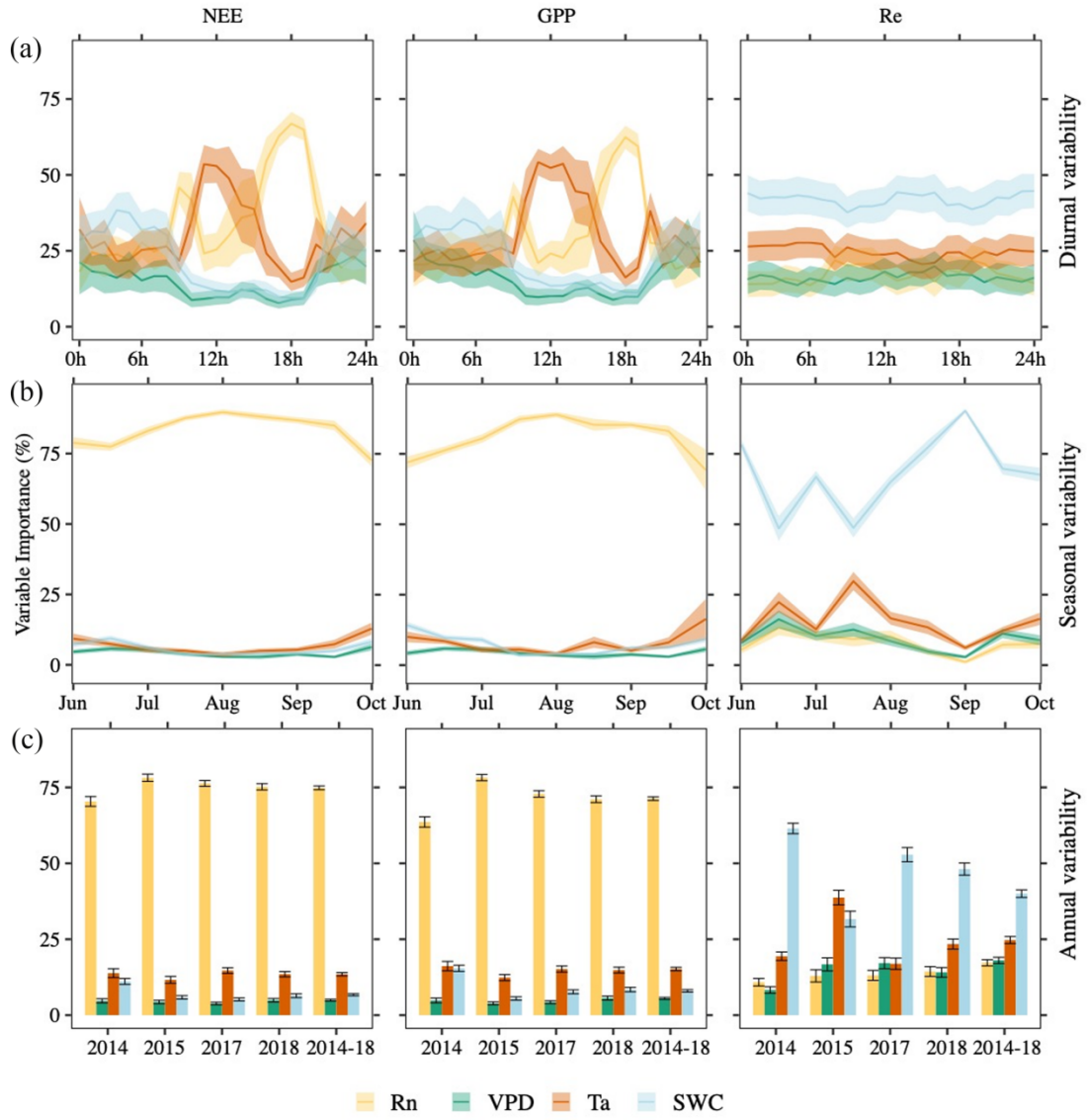
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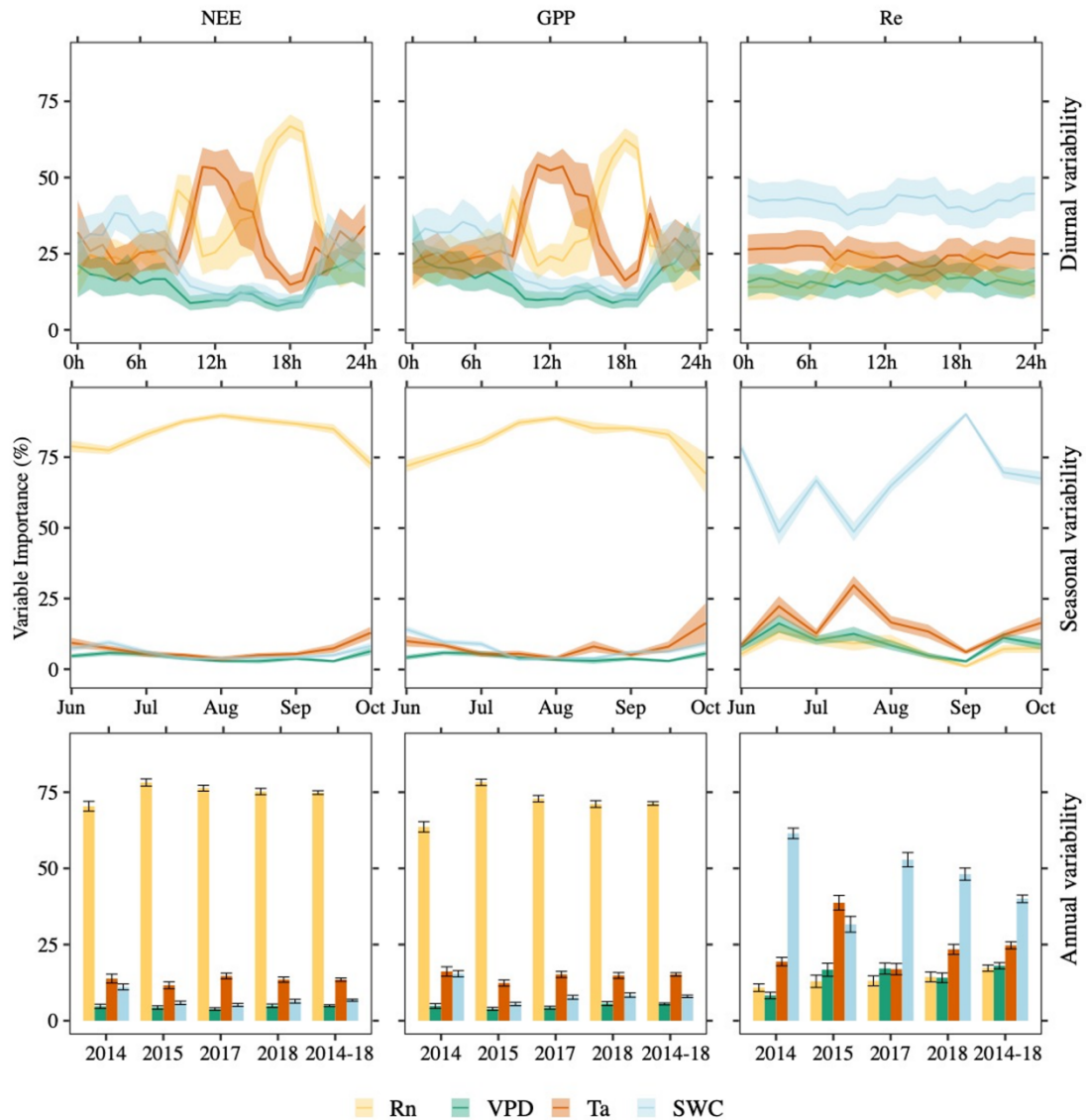
250 **Figure 4. Seasonal variability of daily 5-day moving average daily NEE, Re, and GPP in the swamp alpine**
 251 **meadow.**

252 **3.3 The importance of environmental forcing controlling C fluxes**

253 Our data processed by a machine learning technique suggest that the relative importance of the primary
 254 environmental drivers (Rn, Ta, VPD, and SWC) regulating terrestrial C fluxes varies on different time
 255 scales (i.e., diurnally, seasonally, and interannual scales) in this swamp meadow of QTP (Fig. 5(a), (b),
 256 (c)). The diurnal variability of NEE and GPP was mostly driven by Ta (Fig. 5(a)), especially in the central
 257 hours of the day between 11:00 and 15:00 while Re showed a fairly lower temperature dependence
 258 compared to NEE and GPP (Fig. 5(a)). SWC was ~~a~~ relatively more important than air temperature
 259 controlling the diurnal variability of respiration (Fig. 5(a)). The seasonal variability shaping the terrestrial
 260 C fluxes are regulated not only by meteorological variables but also by plant phenology. To separate the
 261 role of meteorological variables from phenology, we ~~also~~ carried out a random forest analysis every
 262 fortnight and assumed that plant phenology changed little during this time span (Fig. 5(b); seasonal
 263 variability). The analyses based on random forest revealed a distinct seasonal pattern from June to
 264 September, pointing to a marked contribution of net radiation over NEE and GPP (Fig. 5(b)).

265 Interestingly, Re was found mostly regulated by SWC. On an annual scale, the contribution patterns of
266 each environmental driver to the variations of C fluxes are similar to the ones found at the seasonal scale.
267 The interannual variability of NEE and GPP were controlled more clearly by R_nincoming radiation while
268 SWCsoil-moisture revealed a stronger relative importance over Re (Fig. 5(a), (b), (c)). Overall, SWC
269 dynamics seem to be the most important variable explaining the variability observed in the Re data (Fig.
270 5), suggesting that soil moisture plays an essential role on diurnal, seasonal, and interannual basis in this
271 cold swamp meadow ecosystem. Note also that Ta played a secondary role controlling the Re at all
272 assessed time scales (Fig. 5(a), (b), (c)).





274

275 **Figure 5. Contribution to diurnal, seasonal, and annual variation of NEE, GPP, and Re from different**
 276 **environmental drivers (Rn (yellow), Ta (orange), SWC (blue), and VPD (green)). Solid lines (diurnal and**
 277 **seasonal variability) and bars with error bars (annual variability) both illustrate the average \pm standard**
 278 **deviation of the importance across 1000 decision trees. Annual variability refers to the variability of the**
 279 **integrated growing season of 2014, 2015, 2017, and 2018.**

280 4. Discussion

281 Since NEE is the difference between Re and GPP, environmental variables affecting Re and GPP could
 282 affect NEE indirectly (Song et al., 2011). Radiation affects the magnitude of plant photosynthesis and
 283 controls temperature, one of the key factors related to C fluxes. This suggests that abundant radiation
 284 benefits photosynthesis and respiration and thus directly affects the C sink strength in the alpine wetland

285 ecosystem of the Qinghai Lake (Cao et al., 2017). Niu et al. (2017) show that 99% of the interannual
286 variation of NEE in an alpine swamp meadow can be well explained by temperature conditions,
287 precipitation and radiation. The results of this study demonstrate that ecosystem C sequestration is
288 regulated not only by radiation and temperature but also by soil moisture in the alpine swamp meadow
289 site studied herein. Given there is no significant difference in net radiation between the four years we
290 studied, ~~t~~The effects of soil water content and temperature on C fluxes on diurnal, seasonal, and annual
291 scales are therefore discussed in detail below.

292 **4.1 Low soil moisture is associated with enhanced ecosystem respiration**

293 A previous study in alpine swamp meadow ecosystems found that water stress may be the key limiting
294 factor leading to a decline in the photosynthetic rate at noon with a low SWC of 6–21% (Zhang et al.,
295 2018). In this alpine swamp, the soil layer maintained a relatively high~~e~~r SWC due to the frequent
296 precipitation during the growing season. SWC was always greater than 70% during the entire study
297 period (Fig. 2**(b)**). Therefore, microbial activity, and thus heterotrophic respiration were likely
298 suppressed by the anaerobic environment due to saturated soil water condition. At this site, SWC was
299 found to be a more important variable than temperature controlling the variability of ecosystem
300 respiration at different time aggregations (Fig. 5**(a), (b), (c)**). Our results suggest that even under water-
301 saturated conditions, the C dynamics of this alpine swamp meadow are still highly sensitive to changes
302 in soil moisture and could therefore be significantly influenced by future changes in water supply (Li et
303 al., 2015). In fact, previous studies have stressed that soil moisture will likely interact with temperature
304 to affect ecosystem respiration (Han et al., 2013) and therefore modify the overall C sink strength.
305 Moreover, Zhao et al. (2010) also noticed that the CO₂ emission rates decrease notably after rain events,
306 and Zhu et al. (2020) confirmed that annual precipitation exhibits significant impact on variation of
307 annual net C uptake. All these existing studies have suggested that C fluxes are related to water
308 availability condition, but few studies have found that soil moisture explicitly affects respiration.
309 Since C fluxes are affected by plant phenology and climate factors, ~~-~~including temperature, soil
310 moisture, and radiation simultaneously (Fig. 5), to analyze the effects of a single factor, ideally, other
311 factors need to be identical or at least close (no significant differences). Based on this theory and to~~T~~~~e~~
312 better understand the underlying mechanisms around how SWC interacts with the C fluxes in the studied
313 alpine swamp meadow ecosystem, we selected a specific group of data for further evaluation other than

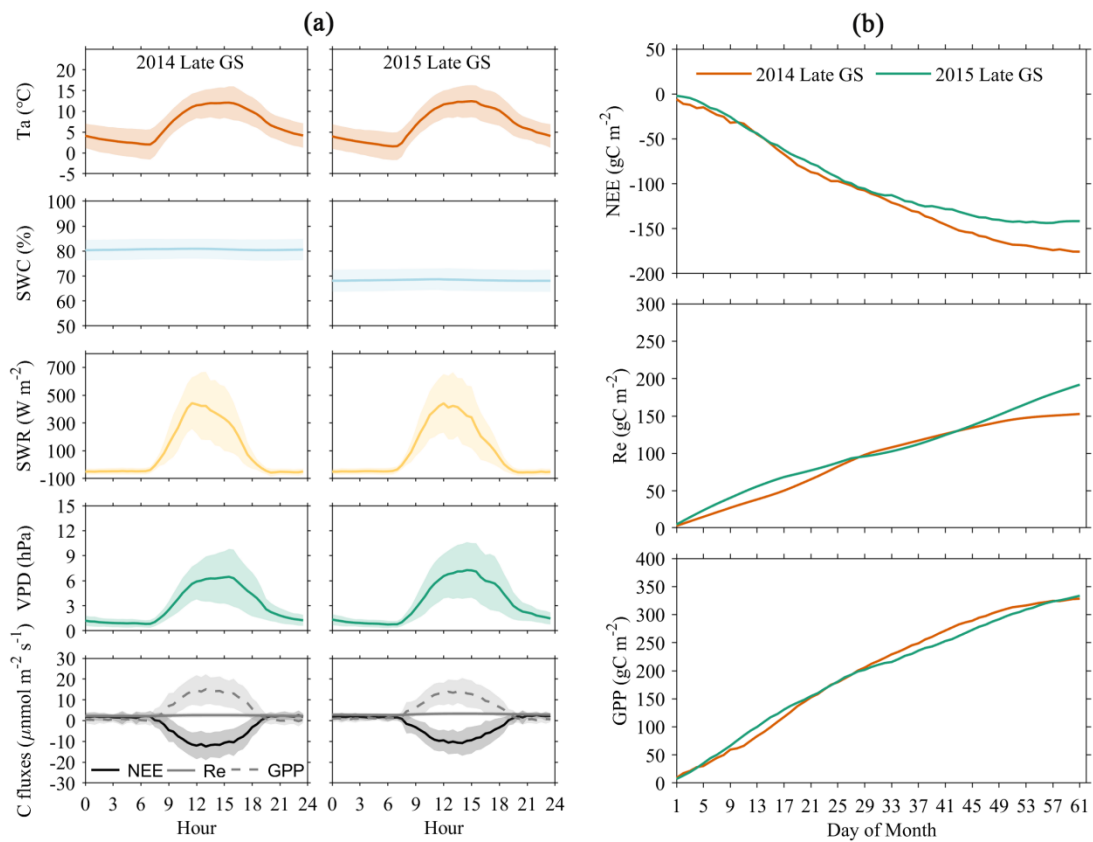
314 the entire observation time. The group contains two late growing season periods, –which have significant
315 differences in SWC but no significant difference in Ta (Fig. 6(a); Table S2).

316 The most significant difference in C fluxes between the late growing season of 2014 and 2015 was
317 observed in Re ($p < 0.05$). Additionally, on both diurnal and seasonal scales, a 15.4% decrease of SWC in
318 the late growing season of 2015 resulted in a 25.7% increase of Re and a 19.4% decrease in net C uptake
319 compared to 2014 (Figs. 6(a), (b); Table S2). This finding suggests that drier conditions likely prevented
320 this alpine swamp meadow from water-logged states, thereby strengthening soil respiration due to
321 improved soil aeration (Wang et al., 2014). According to the literature, the intensification of anaerobic
322 conditions due to water saturated soil can be responsible for weaker respiratory losses (Jansson and
323 Hofmockel, 2020).

~~324 There is evidence that excessive soil water can negatively affect plant physiological and ecological
325 processes by, for example, insufficient supply of metabolic substrates and the production of toxic
326 substances (Jackson and Colmer, 2005), which may reduce the overall plant photosynthetic efficiency
327 (Xu and Zhou, 2011). Although there was no significant difference between the late growing season GPP
328 of 2015 and 2014 in terms of both daily accumulated GPP and diurnal rates of GPP, the decline observed
329 in SWC during September 2015 (the driest month with 65.1% SWC) led to a 11% increase of daily
330 accumulated GPP (Fig. S1; Table S3). The excess of SWC in 2014 caused an inundation of the
331 aboveground plant domain, which also likely contributed to the lower value in GPP by reducing plant
332 photosynthetic efficiency (Cronk et al., 2001; Hirota et al., 2006).~~

~~333 Since the increase of GPP could not offset the increase in Re, September 2015 and the late growing
334 season of 2015 experienced a lower C sink strength.~~ Although the SWC in September 2015 was much
335 greater than the 6–21% range reported by Zhang et al. (2018), our data suggest that a 22.2% reduction in
336 SWC in September 2015 resulted in a 51.6% decline in net C uptake ~~rate~~ compared to September 2014
337 (Fig. S1; Table S3). There is evidence from the literature that the rates of net C uptake in alpine wetlands
338 during the growing season can be lower under drier conditions (Hao et al., 2011), indicating that this
339 alpine swamp meadow ecosystem may be adapted to high levels of SWC (Li et al., 2015), ~~but also that
340 drier conditions may not always favour net C uptake in the cold-adapted alpine swamp meadow
341 ecosystems.~~ Higher SWC may limit the diffusion of oxygen from the atmosphere to the soil, inhibiting
342 the activity of microorganisms and reducing the decomposition rate of soil organic matter, decreasing

343 the nutrients in the soil, and consequently reducing the photosynthetic in alpine wetlands (Chimner and
 344 Cooper, 2003). Our comparisons suggest that drying can weaken the overall C sink strength in this alpine
 345 swamp meadow ecosystem. ~~Therefore, we conclude that the large C sink strength observed in this alpine~~
 346 ~~swamp meadow of the QTP is largely attributed to the inhibiting effects of the nearly saturated soil~~
 347 ~~condition on C release.~~ Wetlands are predicted to experience lower water tables due to permafrost
 348 degradation in the Tibetan Plateau and, therefore, permafrost thaw-induced wetland drying could
 349 enhance the response of C emissions to climate warming (Yu et al., 2020).



350

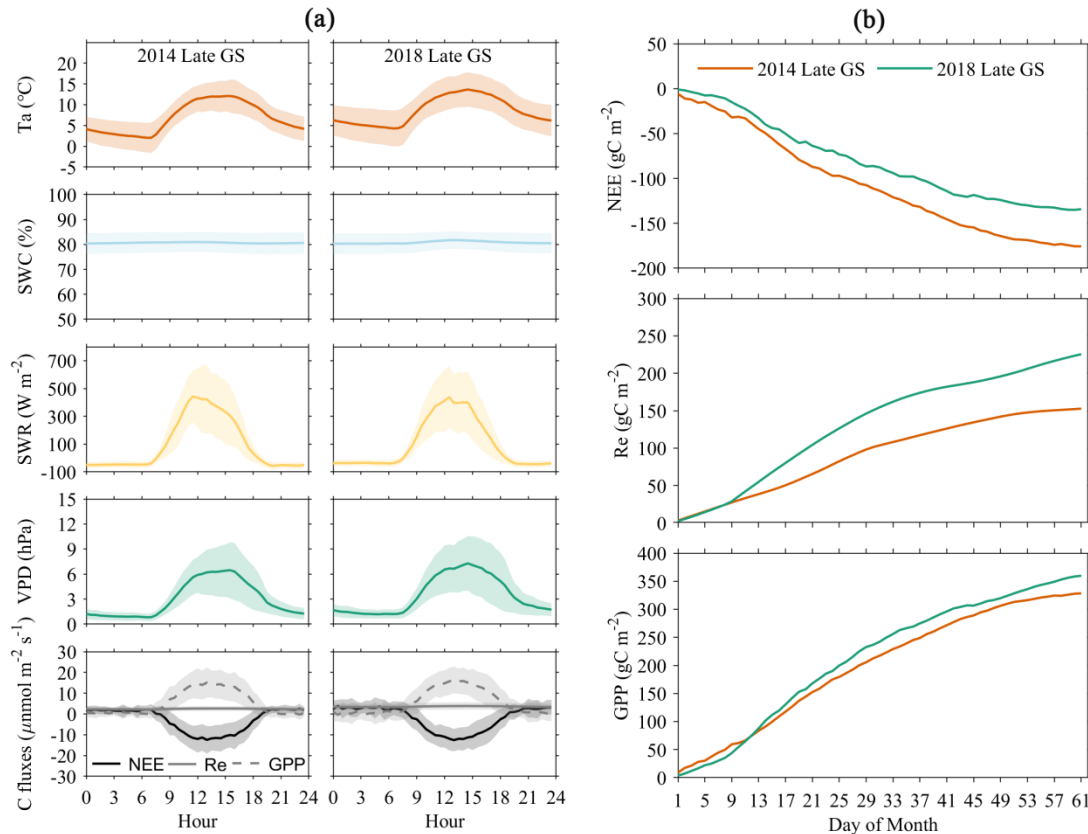
351 **Figure 6. (a) Comparisons of the diurnal variations of environmental drivers (Ta, SWC, Rn, and VPD) and**
 352 **C fluxes (NEE, Re, and GPP) between the late growing season of 2014 and 2015. The shading represents the**
 353 **mean ± standard deviation of the presented variables. (b) Comparisons of the daily accumulated C fluxes**
 354 **(NEE, Re, and GPP) between the late growing season of 2014 and 2015. Note: late GS represents late (Aug.–**
 355 **Sep.) growing season.**

356 4.2 Temperature increase leads to higher C losses rather than enhanced C uptake

357 The important role played by temperature controlling C exchange has been extensively found in alpine
358 marshland ~~not only~~ across the QTP (Qi et al., 2021). For example, Zhao et al. (2010) and Zhao et al.
359 (2005) show that ecosystem respiration follows the exponential variation of soil temperature. Zhu et al.
360 (2020) also suggested that soil temperature plays the most important role in the change of monthly
361 ecosystem respiration in the alpine wetland at Luanhaizi, northeastern Qinghai-Tibet Plateau. ~~but and also~~
362 ~~in other ecosystems such as low and high Arctic tundra (López-Blanco et al., 2017, 2020) and northern~~
363 ~~marshlands (Watson and Byrne, 2009).~~ We therefore explored ~~another~~ ~~an additional~~ comparison between
364 the late growing season of 2014 and 2018 (Fig. 7(a), (b); Table S2) where SWC_c was almost identical at
365 both periods while temperature differed. Compared to 2014, a 25.2% increase in Ta in the late growing
366 season of 2018 led to joint larger GPP and Re fluxes (Fig. 7a, b; Table S2). Although both GPP and Re
367 increased, the intensification in Re was greater than ~~the one found in~~ GPP, indicating that warmer
368 temperatures had a stronger impact on ecosystem respiration, resulting in a decrease of the net C uptake
369 (Fig. 7(b); Table S2). To evaluate if this finding is also consistent at an annual scale, we further analyzed
370 annual aggregated data. An annual comparison was made between ~~the~~ 2014, 2017, and 2018 when
371 SWC were found insignificantly different while temperatures in 2017 and 2018 were 44.4% higher than
372 in 2014 (Table S4). Additionally, this 44.4% increase in Ta in 2017 and 2018 both led to stronger GPP
373 and Re (Table S4). Although both GPP and Re increased, the intensity in Re was greater than GPP,
374 indicating that warmer temperatures have a stronger impact on ecosystem respiration in this site, resulting
375 in an approximately 50% decrease of the net C uptake (Table S4).

376 This comparison suggests that future warming could weaken the overall C sink strength in this alpine
377 swamp meadow ecosystem. Similar climate sensitivities have also been found in recent studies. For
378 example, a study performed by Niu et al. (2017) in an alpine swamp meadow on the central Tibet Plateau
379 suggests Re was more sensitive to increased temperature than GPP. This suggests that global warming
380 may exacerbate future C releases in the alpine wetlands of the QTP (Gao et al., 2019; Niu et al., 2017;
381 Zhu et al., 2020). However, other researchers have also reached different conclusions; for instance, Qi
382 et al. (2021) found that GPP is consistently more sensitive than Re to changes of temperature at daily,
383 seasonal, and annual scales, suggesting that cold conditions can act as strong constraint on C uptake in
384 alpine marshlands. Liu et al. (2018) concluded that warming has a significant inhibitory effect on GPP

385 and minor effect on Re, resulting in a weaker carbon sequestration capacity of their studied alpine
386 wetland ecosystem. Wei et al. (2021) also found that the uptake of C by plants will exceed the amount
387 of C release under warmer and wetter climate conditions at an annual scale based on manipulative
388 experiments and model simulations focused on the Tibetan Plateau. Their study is based on a longer-
389 term trend while our study only covers 4-years of year-round observations thus site-specific differences
390 in time and space scales may explain this variability. Nevertheless, such Additionally, an analysis based
391 ~~on a twenty-year warming experiment manipulated by Sistla et al. (2013) in an Arctic tundra ecosystem~~
392 ~~showed that the insensitive response of net C uptake to warming-induced conditions can be balanced by~~
393 ~~long term variations in vegetation structure and composition together with soil thermodynamics,~~
394 ~~revealing a similar sensitivity to temperature triggered by a compensatory effect between photosynthesis~~
395 ~~and C losses. Similar correlated responses between GPP and Re balancing each other have also been~~
396 ~~found in low and high Arctic tundra sites (López Blanco et al., 2017, 2020), a grassland in northeast~~
397 ~~China (Jiang et al., 2012), and an alpine meadow of the Qinghai Tibet Plateau (Liu et al., 2018).~~
398 ~~Certainly, all these comparison~~ results indicating inconsistent ecosystem responses ~~effects~~ suggest that
399 there are still large uncertainties regulating the responses of C fluxes to temperature variation and further
400 work is still crucial.



401

402 **Figure 7. (a) Comparisons of the diurnal variations of environmental drivers (Ta, SWC, Rn, and VPD) and**
 403 **C fluxes (NEE, Re, and GPP) between the late growing season of 2014 and 2018. The shading represents the**
 404 **mean ± standard deviation of the presented variables. (b) Comparisons of the daily accumulated C fluxes**
 405 **(NEE, Re, and GPP) between the late growing season of 2014 and 2018. Note: late GS represents late (Aug.–**
 406 **Sep.) growing season.**

407 **4.3 Impacts from the eCombined effects of warming temperature and soil moisture changes on C**
 408 **exchange dynamics dynamics**

409 The QTP experienced a higher rates of temperature increase than that of the Northern Hemisphere
 410 average (Zhang et al., 2013). The effects triggered by climate-induced warming over NEE in this area
 411 have been argued to either increase or decrease the net C balance NEE, or even have no effect whatsoever
 412 (Ganjurjav et al., 2018; Li et al., 2020; Wu et al., 2011; Zhu et al., 2017). These inconsistent responses
 413 could be due to water limitations offsetting the C balance or even reversing the effect of elevated
 414 temperatures, which change the decomposition and photosynthetic processes (Wu et al., 2011; Yu et al.,
 415 2013; Zhao et al., 2019). Alpine swamp meadows of the QTP have recently attracted much attention
 416 because they hold 5.9% (~ 1.98 Pg C) of the total grassland soil organic C (~ 33.52 Pg C). Such

417 ecosystems have the highest organic C density ($\sim 50 \text{ kg C m}^{-2}$) and play an important role in the global
418 C cycle (Niu et al., 2017). To test whether the observed SWC effects in this study were representative of
419 other sites in the QTP and put it into a ~~broader~~wider perspective, we examined the temperature and
420 precipitation (as a proxy for SWC)~~warming~~ impacts on NEE_ ~~-in relation to ambient precipitation as a~~
421 ~~proxy for SWC~~ (Liu et al., 2016).

422 Our 4-year dataset revealed that this alpine swamp meadow functioned as a net C sink of -168.0
423 $\pm 62.5 \text{ gC m}^{-2} \text{ y}^{-1}$ at a 3571m asl. According to Wei et. al (2021), there are six observational studies about
424 C fluxes at Haibei, three of which focused on alpine swamp meadows or wetland. These alpine swamp
425 meadows were reported as a net C source while we our site showed a consistent C sink. The different
426 directions of C exchange suggest that there are still uncertainties in our understanding of C exchange in
427 this alpine swamp meadows, and further insights are obtained from studying multiple years of
428 observations. Therefore, further efforts are still needed to improve our projection of C balance change of
429 this ecosystem under changing climate.

430 The NEE observations from this study were within the NEE ranges of previously ~~studies~~ studies in ~~the~~ similar
431 alpine swamp meadows located across the QTP ($-255.5 - 173.2 \text{ gC m}^{-2} \text{ y}^{-1}$) (Table 2). In addition, the
432 NEE estimates of this alpine swamp meadow show a stronger C sink strength than those from alpine
433 meadows ($-161.3 - 85.4 \text{ gC m}^{-2} \text{ y}^{-1}$) (Chai et al., 2017; Wang et al., 2017; Wu et al., 2020), alpine steppes
434 ($-30 - 21.8 \text{ gC m}^{-2} \text{ y}^{-1}$) (Wang et al., 2018; Wang et al., 2020a; Wang et al., 2020b; Wu et al., 2010), and
435 alpine shrublands ($-14 - -67 \text{ gC m}^{-2} \text{ y}^{-1}$) (Zhao et al., 2005, 2006). This is likely a result of the inhibiting
436 effects of the nearly-saturated soil condition over soil respiration rather than by the lower temperatures
437 (Sun et al., 2021). Therefore, in permanently or seasonally inundated swamp meadows, high SWC ~~caused~~
438 ~~by relatively sufficient water supply~~ may have triggered lower C loss rates further benefiting C
439 preservation. At our site, the higher C sink strength was likely attributed to higher precipitation (and
440 therefore higher SWC) and lower temperature, which created colder and more humid conditions than
441 other sites (Table 2).~~The higher C sink strength at our site was likely attributed to higher precipitation~~
442 ~~and lower temperature, which created colder and more humid conditions than other sites (Table 2).~~ It
443 has been demonstrated in the literature that cold and humid conditions favour stronger C sinks in alpine
444 meadow ecosystems (Fu et al., 2009).

445 The interannual comparison of the sites presented in Table 2 shows that under low annual precipitation
 446 conditions (~ 300 mm), the joint effects of warming and reduced precipitation weakened the net C uptake
 447 at Damxung (Niu et al., 2017), and even turned the C sink of Xiaobo Lake wetland into a C source when
 448 comparing 2015 with 2012 and 2013 (Cao et al., 2017; Wu et al., 2018). Under relatively high annual
 449 precipitation (~ 500 mm), the joint effects of warming and increased precipitation ~~also~~ enhanced C
 450 release in Haibei^a ~~Haibei~~ site when comparing 2006 to 2004 and 2005 (Zhao et al., 2010). This indicates
 451 that net C uptake under warming conditions can be weakened even under high annual precipitation rates.

452 **Table 2 Comparison of annual NEE (gC m⁻² y⁻¹) at different sites in the QTP.**

Site	Altitude (m)	Ecosystem	Year	Ta (°C)	P (mm)	Annual NEE	Reference
Haibei ^a (37°35'N, 101°20'E)	3200	AWM	2004	2.3	493.5	101.1	Zhao et al., 2010
			2005	2.2	475.2	44.0	
			2006	3.6	562.4	173.2	
Xiaopo Lake (36°42'N, 100°46'E)	3228	AWM	2015	2.8	304.3	54.6	Wu et al., 2018
			2012	1.2	357.0	-225.6	Cao et al., 2017
			2013	1.2	357.0	-255.5	
Damxung (30°28'N, 91°4'E)	4285	ASM	2009	3.4	208.9	-148.5	Niu et al., 2017
			2011	2.6	393.3	-190.8	
Haibei ^b (37°35'N, 100°00'E)	3571	ASM	2014	-0.9	662.8	-240.3	This study
			2015	-0.6	521.4	-200.1	
			2017	-0.5	661.2	-118.1	
			2018	-0.3	624.3	-113.4	

453 Note: P denotes precipitation, AWM denotes Alpine Wetland Meadow and ASM represents Alpine Swamp Meadow,

454 Haibei^a and Haibei^b denotes different sites.

455 5. Conclusions

456 The alpine swamp meadow from the QTP presented in this study has been found to act as a consistent
 457 and strong sink of CO₂ (-168.0 ± 62.5 g C m⁻² y⁻¹). The results from a novel machine learning technique
 458 revealed that air temperature is the most important variable driving NEE and GPP ~~on~~ a diurnal scale,
 459 while ~~incoming~~ net radiation has a stronger importance controlling the seasonal and interannual

460 variability of the same fluxes. Soil moisture, however, has the largest influence over R_e variability ~~on~~
461 diurnal, seasonal, and interannual scales, suggesting that soil water content is a key control on ecosystem
462 respiration ~~and overall C sink strength~~. In addition, air temperature plays a less important role in
463 regulating the C exchange variability. This study reveals that both drying and warming can suppress net
464 C uptake in water-saturated alpine swamp meadow ecosystems by enhancing ecosystem respiration. The
465 response of net C uptake to climate warming further indicates that the forecasted warming in the QTP
466 will not always increase the net C sink strength. Our results not only highlight the contributions of soil
467 moisture in regulating C sequestration under high water conditions but also support future process-based
468 modelling initiatives focusing on alpine swamp meadow ecosystem C dynamics.

469 **Data availability.** Post-processed data and scripts used in this paper are available from the authors upon
470 request (xyli@bnu.edu.cn).

471 **Author contribution**

472 JQ, XL, LL, TRC, XW, and ELB designed the research. JQ, LL, and ELB processed the data and
473 performed the analyses. YJ and HY helped collect the data. JQ and ELB created the visualization of the
474 outputs. JQ, XL, and ELB prepared the manuscript with contributions (writing, review, and edition) from
475 LL, TRC, and ZJ.

476 **Competing interests.** The authors declare that they have no conflict of interest.

477 **Acknowledgements**

478 The study was financially supported by the second Tibetan Plateau Scientific Expedition and Research
479 Program (STEP 2019QZKK0306), the National Natural Science Foundation of China (NSFC 41730854
480 & 41971029), the Strategic Priority Research Program of Chinese Academy of Sciences
481 (XDA20100102), and projects from the state Key Laboratory of Earth Surface Processes and Resource
482 Ecology. We also gratefully acknowledge the financial support from the China Scholarship Council (No.
483 201906040130) and the Faculty of Technical Science from Aarhus University. Additionally, ELB was
484 funded by the Greenland Research Council, grant number 80.35, financed by the “Danish Program for
485 Arctic Research.

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