

Radiation, soil water content, and temperature interactions with carbon cycling in an alpine swamp meadow of the northeastern Qinghai-Tibet Plateau

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Abstract. Predicted intensified climate warming will likely alter the ecosystem net carbon (C) uptake of the Qinghai-Tibet Plateau (QTP). Variations in C sink/source responses to climate warming have been linked to water availability; however, the mechanisms by which net C uptake responds to soil water content in saturated swamp meadow ecosystems remain unclear. To explore how soil moisture and other environmental drivers modulate net C uptake in the QTP, field measurements were conducted using the eddy covariance technique in 2014, 2015, 2017, and 2018. The alpine swamp meadow presented in this study was a consistent and strong C sink of CO₂ (-168.0 ± 62.5 gC m⁻² y⁻¹, average \pm standard deviation) across the entire 4-year study period. A random forest machine-learning analysis suggests that the diurnal, seasonal, and annual variations of net ecosystem exchange (NEE) and gross primary productivity (GPP) were controlled by temperature and net radiation. Ecosystem respiration (Re), however, was found mainly regulated by the variability of soil water content (SWC) at different temporal aggregations followed by temperature, the second contributing driver. We further explored how Re is controlled by nearly saturated soil moisture and temperature comparing two different periods featuring almost identical temperatures and significant differences on SWC and vice versa. Our data suggest that, despite the

35 relatively abundant water supply, periods with a substantial decrease of SWC or increase of temperature
36 produced higher Re and therefore weakened the C sink strength. Our results reveal that nearly saturated
37 soil conditions during the growing seasons can help maintain lower ecosystem respiration rates and thus
38 enhance the overall C sequestration capacity in this alpine swamp meadow. We argue that climate
39 warming in the future could change soil hydrological conditions of alpine swamp meadow through
40 degrading permafrost or accelerating thawing-freezing cycling, therefore, having a direct effect on soil
41 respiration and subsequent effect on ecosystem C sink magnitude.

42 **1. Introduction**

43 Wetlands play a significant role in the global carbon (C) cycle due to a large amount of C stored in their
44 soils. The Qinghai-Tibet Plateau (QTP), with an average altitude of over 4,000 m a.s.l., has
45 approximately 10×10^4 km² of natural wetlands, of which $\sim 50\%$ (4.9×10^4 km²) are alpine swamp
46 meadows. These ecosystems are predominantly located in permafrost areas and are typically soil
47 nutrient-rich and water-logged (Bai et al., 2019; Zhao et al., 2005). Climate change and human
48 disturbance can have profound consequences on permafrost regions (Biskaborn et al., 2019) and
49 significantly impact their hydrological regimes (Lafrenière and Lamoureux, 2019). Hydrological regimes
50 have an important role in controlling wetland functioning (Bohn et al., 2007; Christensen et al., 2003),
51 and the changes of hydrological regimes may put the wetland functioning of the QTP under pressure
52 (Hrubby, 1999; Woodward and Wui, 2001; Foti et al., 2013).

53 The QTP is forecasted to be warmer and wetter in the future (Chen et al., 2015; Cheng et al., 2011).
54 Warming may accelerate the microbial breakdown of alpine soil organic C and subsequently increase
55 CO₂ emissions (Zhu et al., 2015a). Warming could also improve C sequestration capacity by enhancing
56 the photosynthetic inputs and growth rates of alpine plants (Fu et al., 2015). Therefore, the potential
57 increase of CO₂ emissions due to warming in alpine regions could be partially offset by enhanced C
58 uptake (Schuur et al., 2009), triggering different net C uptake responses to climate warming. For
59 example, an increase in temperature in the QTP has been associated with net C sinks in the Zoige alpine
60 wetlands (Kang et al., 2014) but also with net C sources in the Damxung alpine swamp meadow (Niu et
61 al., 2017).

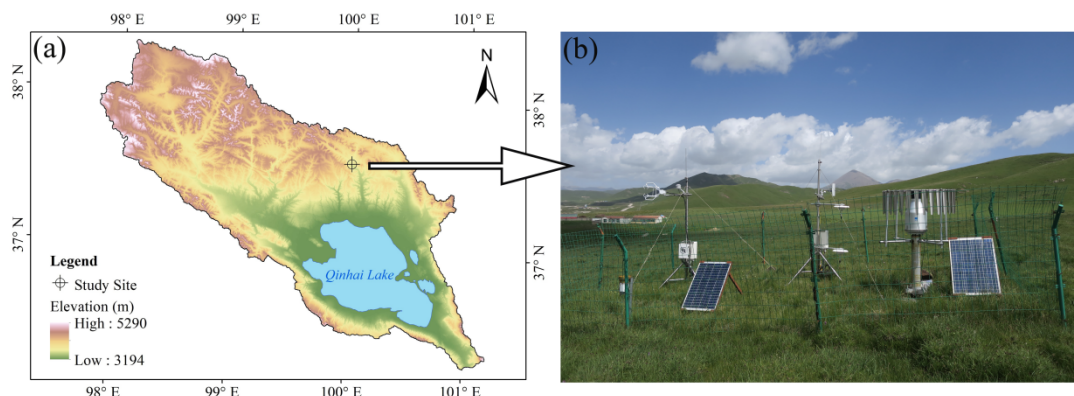
62 According to recent studies in QTP alpine grasslands, water conditions such as soil water content (SWC)
63 can be a key factor that changes water-use patterns and ecophysiological characteristics of alpine plants
64 (Wu et al., 2019) and modulate the warming-mediated increase of ecosystem C uptake (Ganjurjav et al.
65 2016; Peng et al. 2014). Ecosystem C processes such as net C uptake and soil respiration may increase
66 with SWC in dry environments, and decrease in water-logged environments (Quan et al., 2019; Taylor
67 et al., 2017). Warming in conjunction with increased precipitation can turn an ecosystem from net source
68 to a sink of C (Zhao et al., 2019), increasing both photosynthesis and respiration rates during warmer and
69 wetter years (López-Blanco et al., 2017, 2018). However, when warming occurs in soils associated with
70 low moisture, soil drought can change ecosystems from C sinks to sources (Ganjurjav et al., 2017).
71 Studies in QTP alpine meadows have indicated that warming significantly stimulates ecosystem net C
72 uptake in wet years but does not affect ecosystem net C uptake in dry years because the positive effects
73 of warming on net C uptake are compensated by the negative effects of lower soil moisture (Peng et al.,
74 2014).

75 Although many studies concerning ecosystem C dynamics on the QTP have focused on alpine meadow
76 ecosystems (Saito et al., 2009; Zhao et al., 2005, 2010; Zhu et al., 2015b) and alpine swamp meadow
77 ecosystems (Zhao et al., 2005, 2010; Qi et al., 2021; Liu et al., 2020; Zhu et al., 2020), the effect of SWC
78 on the C uptake is still unclear as compared to that of temperature for alpine swamp meadow ecosystems.
79 Alpine swamp meadow ecosystems are characterised by high SWC, and it remains unclear whether a
80 decrease in SWC would alleviate the stress from saturated water on net C uptake or aggravate drought
81 effects. Given that these ecosystems have high soil moisture compared to typical alpine meadows, the
82 effects of warming and water availability on the future C sink strength of the QTP ecosystem remain
83 uncertain. These questions and uncertainties require a detailed investigation to understand wetland C
84 source/sink processes and their future C sink strength variations (sign and magnitude). Therefore, the
85 objectives of this study are to (i) quantify the diurnal, seasonal, and annual variations of net ecosystem
86 exchange (NEE), gross primary productivity (GPP), and ecosystem respiration (Re), (ii) identify and
87 quantify the relative importance of different key environmental drivers contributing to the variability
88 observed of NEE, Re, and GPP, and (iii) analyse how these C fluxes respond to soil water availability,
89 temperature, and radiation variation in a QTP alpine swamp meadow.

90 **2. Materials and methods**

91 **2.1 Site description**

92 The study site (37°35.75' N, 100°00.47' E, 3,571 m a.s.l.) is located in Gangcha County, Qinghai
93 Province, in the north-eastern part of the QTP (Fig. 1(a)). The mean annual temperature and precipitation
94 measured at the Gangcha National Weather Station were 0.1 °C and 389.4 mm between 1982 and 2011,
95 respectively (Zhang et al., 2016). The site area has seasonal permafrost featuring frozen soils between
96 January–March and November–December for a total of about 125–135 days (Zhang et al., 2014). The
97 dominant species of the alpine swamp meadow ecosystem is *Kobresia pygmaea*, accompanied by
98 *Saussurea pulchra*, *Polygonum viviparum*, and *Potentilla saundersiana*. The average plant height at the
99 experimental site is 7.4 ± 1.5 cm, with a $97 \pm 2\%$ coverage. Our study defined the growing season as the
100 period between June and September. The early (June–July) and late (August–September) growing
101 seasons, as they will be henceforth referred to, correspond with the early growing season and late growing
102 seasons, respectively.



103

104 **Figure 1. (a) Location of the study site in the Qinghai Lake basin in the northeastern part of the QTP. (b)**
105 **Eddy covariance system measuring water and CO₂ fluxes between the land surface and the atmosphere in the**
106 **alpine swamp meadow.**

107 **2.2 Field measurements**

108 An Eddy Covariance (EC) system was installed at the study site (Fig. 1b) to measure the CO₂ fluxes at a
109 sampling frequency of 10 Hz from 2014 to 2018. Data for 2016 was missing due to equipment
110 malfunction. The EC system included an open-path CO₂/H₂O infrared gas analyser, which quantified
111 fluctuations in water vapour and CO₂ fluxes. A 3-D sonic anemometer was installed at a 2.0 m height
112 above ground to directly measure horizontal and vertical wind velocity components (u, v, and w). C flux

113 data were recorded with a data logger (Campbell Scientific Inc.). An automated meteorological station
 114 was installed near the EC station to measure meteorological variables such as air temperature (T_a ; °C),
 115 precipitation (P ; mm), net radiation (R_n ; $W\ m^{-2}$), wind speed (WS ; $m\ s^{-1}$), wind direction (WD ; °), relative
 116 humidity (RH ; %), and vapour pressure deficit (VPD ; hPa). The meteorological data were collected at
 117 one-minute intervals and subsequently resampled at 30-minute timesteps to keep pace with the EC data.
 118 More details on the in-situ instrument specifications are summarised in Table 1.

119 **Table 1 Information about the sensors installed in the alpine swamp meadow.**

	Sensor Names	Sensor type	Installation height/depth	Manufacturer
Eddy Covariance	Open-path CO_2/H_2O infrared gas analyser	EC150	2.0 m	Campbell, US
	Three-dimensional sonic anemometer	CSAT3	2.0 m	Campbell, US
Meteorological observation	Net radiation	NR Lite	1.8 m	Kipp&Zonen, Netherlands
	Wind speed/direction	034B	2m	
	Air temperature/humidity	083E-1-6	0.5 m, 1.5 m	MetOne, US
	Atmosphere pressure	PTB110	In data acquisition box	Vaisala, Finland
	Rain-gauge	7852M-AB	0.7 m	Davis, US

120 Soil water content (SWC; %) were measured at depths of 10, 20, 40, 60, and 100 cm from the soil surface
 121 with EC-H₂O sensors (Decagon Devices, USA) at a 10-minute frequency. The precision of the EC-H₂O
 122 sensors for soil moisture measurements was $\pm 0.03\ m^3\ m^{-3}$. As the roots of *Kobresia* meadows are mainly
 123 distributed within the top 20 cm of soil, we focused only on the variation of SWC in the top 20 cm of the
 124 soil.

125 2.3 Eddy Covariance data processing

126 The half-hourly NEE data was calculated using the EddyPro software (version 5.2, LI-COR) from the
 127 10-Hz raw data. During the calculation, three-dimensional rotation was used to correct the data by
 128 removing the effects of instrument tilt irregularity on airflow (Wilczak et al., 2001). Webb, Pearman, and
 129 Leuning (WPL) (Webb et al., 1980) correction was applied to calculate the averages of CO_2 covariance,
 130 rectifying the air density variations induced by heat and water vapour. The half-hourly flux data were
 131 quality-checked based on several filtering algorithms, including: (1) the rejection of outliers in sonic
 132 temperature, water vapour density, and CO_2 density (Li et al. 2008; Liu et al., 2011), (2) the elimination

133 of data one hour before and after precipitation events (Li et al. 2008; Liu et al., 2011), (3) the removal of
134 negative NEE during the non-growing season (from November to March) (Cao et al., 2017; Qi et al.,
135 2021) attributed to the self-heating effect from EC instruments (Cao et al., 2017), and (4) the exclusion
136 of measurements with weak turbulence conditions at night time. The weak turbulence periods were
137 identified by bootstrapping friction velocity (u^*) thresholds, as described by Papale et al. (2006). This
138 approach effectively divided the data into 4-year and 7-temperature subsets with similar micro-
139 meteorological conditions (except for u^*). The u^* thresholds (5%, 50%, and 95% of bootstrapping) were
140 calculated specifically per year and temperature subset.

141 Based on those different subsets, we gap-filled and partitioned NEE (into GPP and Re) to spread the
142 uncertainty variability that emerged from the different u^* thresholds, similar to López-Blanco et al.
143 (2017). All missing data were marked as -9999 (no data). Negative and positive NEE values represent
144 sink and source of C, respectively. Additionally, a standardised mechanism to fill NEE gaps is needed
145 for adequate data processing (Moffat et al., 2007). Therefore, this study adopted the method described
146 by López-Blanco et al. (2020) using the marginal distribution sampling (MDS) algorithm in the
147 REddyProc gap-filling tool (Reichstein et al., 2016), which was readapted from Reichstein et al. (2005).
148 Finally, NEE was separated into GPP and Re applying the REddyProc partitioning algorithm (Reichstein
149 et al., 2016) for further analyses. This partitioning method is based on the exponential regression of night-
150 time respiration with temperature using the Lloyd-Taylor-Function (Lloyd and Taylor, 1994). Night-time
151 periods were selected via current combined solar radiation and potential radiation thresholds based on
152 the exact solar time, latitude, and longitude. REddyProc estimates the temperature sensitivity from a
153 short-term period and based on this short-term temperature sensitivity It estimates the reference
154 temperature in the continuous period of the entire dataset. These estimates were then used to calculate
155 Re for day-time and night-time, while GPP was estimated based on the difference between NEE and Re .

156 **2.4 Identifying the importance of environmental drivers**

157 In order to characterize how environmental conditions impact diurnal, seasonal, and interannual
158 variability of NEE, GPP and Re at this alpine swamp meadow, we used a novel method based on machine
159 learning.

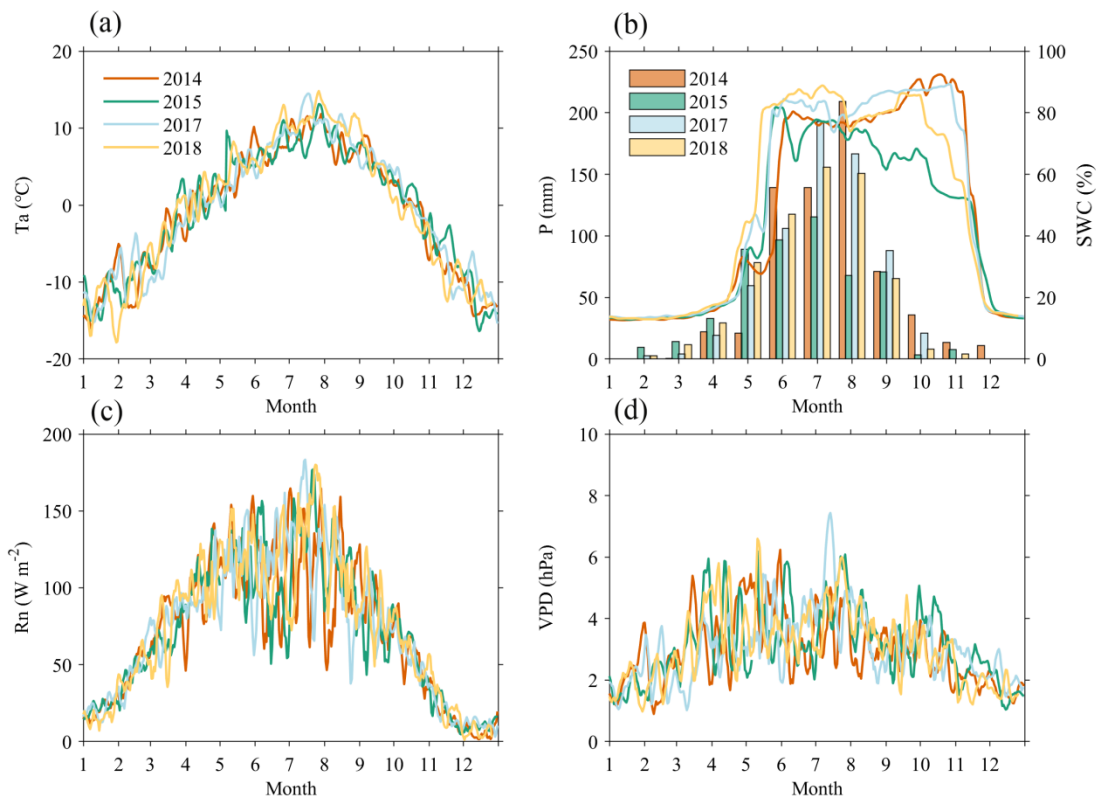
160 Random forest (Breiman, 2001) is a machine learning technique that can be used to quantify and interpret
161 the contribution of environmental drivers (covariates) to the variability of different C fluxes (response
162 variables) by combining multiple individual regression trees. This technique has been increasingly
163 utilized to upscale global C fluxes from eddy covariance data (Zeng et al., 2020) but also to evaluate
164 controls on C cycle processes (Zhang et al., 2017; López-Blanco et al., 2017, 2020). Here, we calculate
165 the relative importance of air temperature (Ta), net radiation (Rn), soil water content (SWC) and vapour
166 pressure deficit (VPD) controlling the C sink strength, photosynthesis, and respiration variability. This
167 random forest algorithm constructs multiple (1000 in this analysis) decision trees during training time
168 with different random subsamples (with replacement) from the same input training dataset. In each
169 cluster classified by random forest, the algorithm generates a multiple linear regression to characterize
170 different C fluxes as a function of environmental drivers (López-Blanco et al., 2017, 2020). This
171 algorithm version (Pedregosa et al., 2011) estimates the relative importance of each covariate between 0
172 and 100%, which correspond to the fraction of decision participating during data clustering. We used the
173 random forest algorithm to evaluate the diurnal, seasonal, and annual patterns of the relative importance
174 of Ta, Rn, SWC, and VPD responsible for the variability of C fluxes. We used data from the June-
175 September period aggregated per hour, and we ran multiple random forests with growing season data
176 binned per hour of the day, day of the year and year, respectively (Table S1).

177 To further analyse the effect of soil moisture, radiation, and temperature on C fluxes, we selected two
178 groups of time stamps with a significant difference in SWC but almost identical Ta and Rn (i.e. late
179 growing season of 2014 vs 2015) and a significant difference in Ta but almost identical SWC and Rn
180 (i.e. late growing season of 2014 vs 2018). Additionally, in order to analyse the effect of annual
181 temperature on C fluxes, we selected a group of time stamps with significant differences in Ta but almost
182 identical SWC and Rn (i.e. 2017 vs 2014, and 2018 vs 2014). We made the comparison in each group to
183 exclude the influence of plant phenology, which can influence C fluxes significantly. The magnitude of
184 the differences between C fluxes in the same group were analysed by the independent-sample T-test
185 method.

186 **3. Results**

187 **3.1 Meteorological variability**

188 Mean daily meteorological variables including Ta, P, SWC, and Rn exhibited evident seasonal variability
189 except for VPD; these variables increased progressively in the early growing seasons, reached their
190 maximum in July and decreased gradually afterwards (Fig. 2(a), (b), (c), (d)). Air temperature during the
191 growing season was 7.7 ± 2.6 , 7.4 ± 2.6 , 8.5 ± 2.9 , and 9.2 ± 3.3 °C in 2014, 2015, 2017, and 2018,
192 respectively while precipitation totalled 662.8, 521.4, 661.2, and 624.3 mm for the same years, falling
193 primarily during the growing season. The precipitation measured during the late growing season of 2015
194 was only half of the amount measured in 2014, 2017, and 2018. The lower precipitation regime led to a
195 marked decline in SWC, making the late growing season of 2015 the driest period among all growing
196 seasons during the study period. The greatest difference in SWC occurred in the late growing season of
197 2014 and 2015, when Ta was the same at 6.8 ± 2.6 and 6.8 ± 2.5 °C, respectively. Compared to 2014,
198 SWC decreased by 15.4% in the late growing season of 2015 (Fig. 2(b); Table S2). Meanwhile, SWC in
199 the late growing season of 2014 and 2018 was almost identical (80.7 ± 4.1 and 80.8 ± 3.8 , respectively),
200 but the temperature difference was the largest (25%) compared to any other year (Fig. 2(a); Table S2).

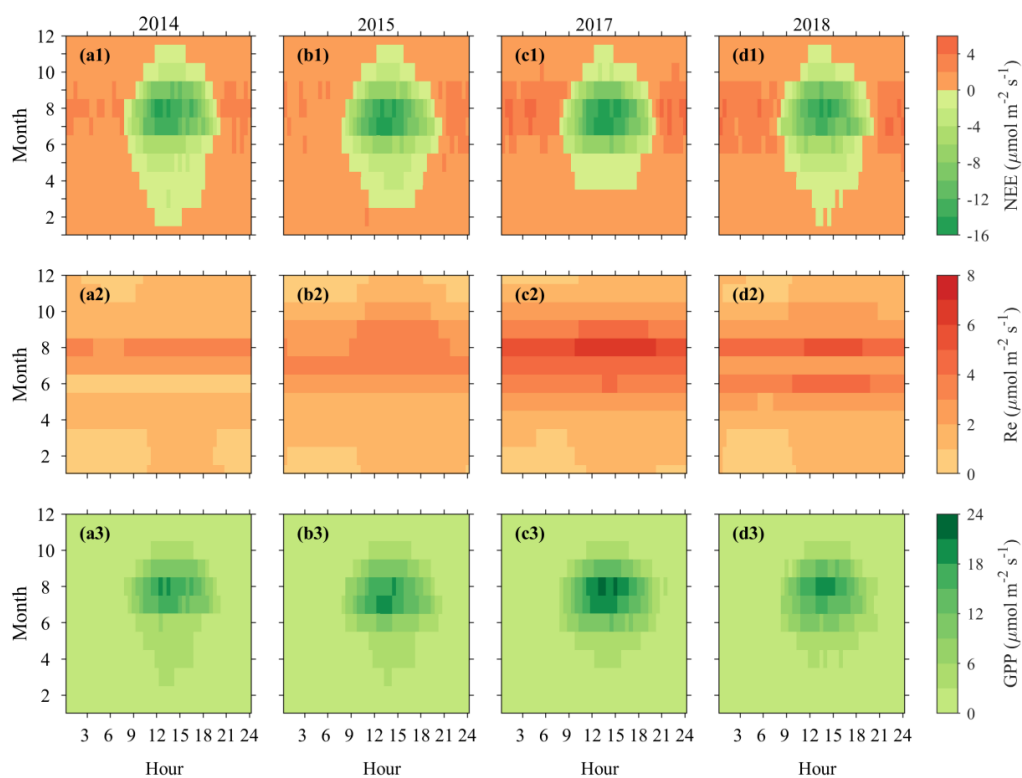


201

202 **Figure 2. Five-day moving average meteorological variables (Ta, P, SWC, Rn, VPD) in the studied swamp**
203 **alpine meadow.**

204 **3.2 Diurnal, seasonal, and annual variability of CO₂ fluxes**

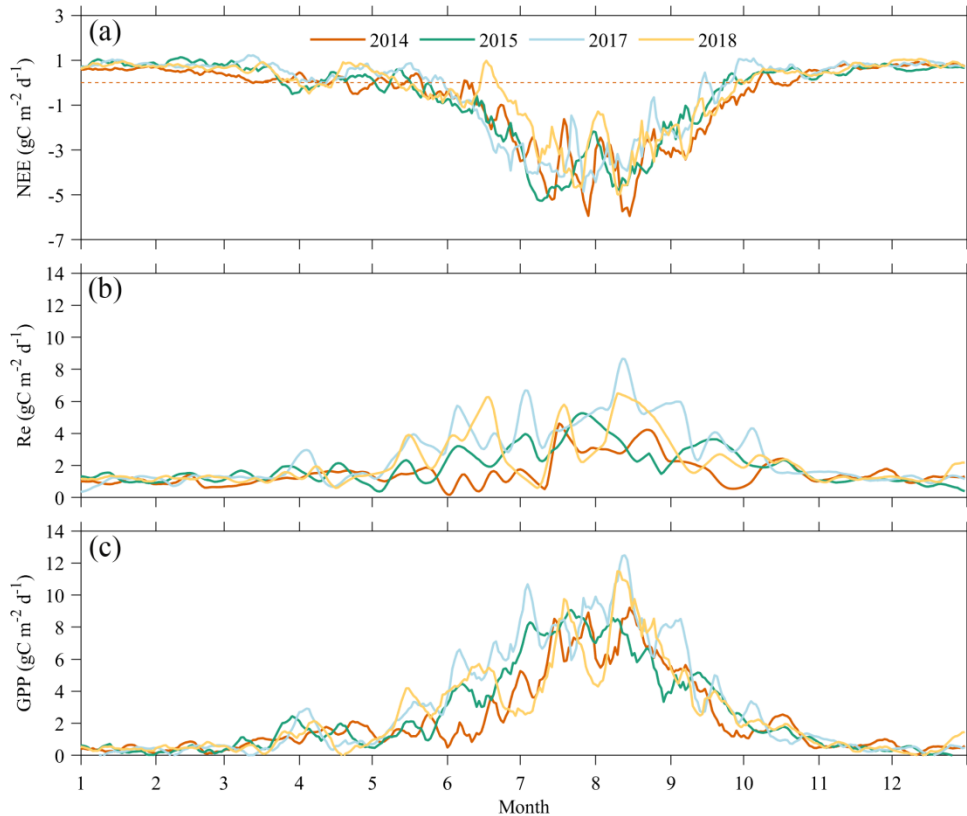
205 During the growing season, NEE (Fig. 3, (a1)–(d1)) and GPP (Fig. 3, (a3)–(d3)) featured a clear peak of
206 the diurnal variations; both fluxes reached their summit between 12:00 and 14:00 local time. Re,
207 however, presented a lower daily variability. The rates of NEE, Re, and GPP during the growing season
208 averaged -2.3 ± 0.3 , 3.2 ± 1.0 , and $5.5 \pm 0.9 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively, for the entire study period (2014,
209 2015, 2017 and 2018). For the late growing season, the lowest rate of net C uptake was measured in 2015
210 ($-10.0 \mu\text{mol m}^{-2} \text{s}^{-1}$), whereas 2014 ($-12.4 \mu\text{mol m}^{-2} \text{s}^{-1}$), 2017 ($-12.2 \mu\text{mol m}^{-2} \text{s}^{-1}$), and 2018 ($-12.5 \mu\text{mol}$
211 $\text{m}^{-2} \text{s}^{-1}$) exhibited more negative NEE values (i.e. stronger net C uptake rate). Between the late growing
212 season in 2015 and the late growing seasons in 2014, 2017 and 2018, there was a significant difference
213 in the rates of Re ($p < 0.01$) while no significant difference was found in GPP variability ($p > 0.05$),
214 suggesting that Re may be the component causing the difference observed in NEE. Specifically, the rates
215 of Re in the late growing season of 2014 and 2015 were 2.4 ± 0.2 and $3.0 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively,
216 which indicated that during 2015 drier conditions generated a 25% higher Re compared to 2014. For the
217 same periods, the rate of Re in warmer 2018 was $3.5 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$, which was also significantly
218 higher than in 2014 ($2.4 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$).



219

220 **Figure 3. Diurnal variability of gap-filled NEE and partitioned Re and GPP in 2014, 2015, 2017 and 2018.**

221 NEE, Re, and GPP also exhibited a strong seasonal variability; the C fluxes gradually increased from
 222 low values in early June to maximum values in the middle of the growing season (late July to early
 223 August on average), followed by a decrease towards the end of the growing season (Fig. 4(a), (b), (c)).
 224 The C sink strength across the growing season was found lowest in 2018, followed by 2017, whereas
 225 2014 and 2015 exhibited relatively higher values (Fig. 4(a)). Notably, the accumulated Re in the late
 226 growing season was significantly higher in 2015 compared to 2014 ($P < 0.05$), when there was no
 227 significant difference in GPP (Fig. 4(b); Table S2). Moreover, the late growing season of 2015 witnessed
 228 the lowest SWC while keeping the same T_a compared to the same period in 2014 (Fig. 2(b); Table S2).
 229 The substantial decline of SWC observed in 2015 appeared to be responsible for the weaker observed C
 230 sink strength. On the other hand, T_a in the late growing season of 2018 was the highest for the same
 231 period among all the years while the SWC remained the same as 2014 (Fig. 2(a)). The late growing
 232 season of 2018 showed overall higher GPP, Re, and lower net C uptake than 2014. Significantly higher
 233 Re in 2018 caused by higher temperatures eventually led to a decrease of the C sink capacity (Fig. 4(a),
 234 (b), (c); Table S2).



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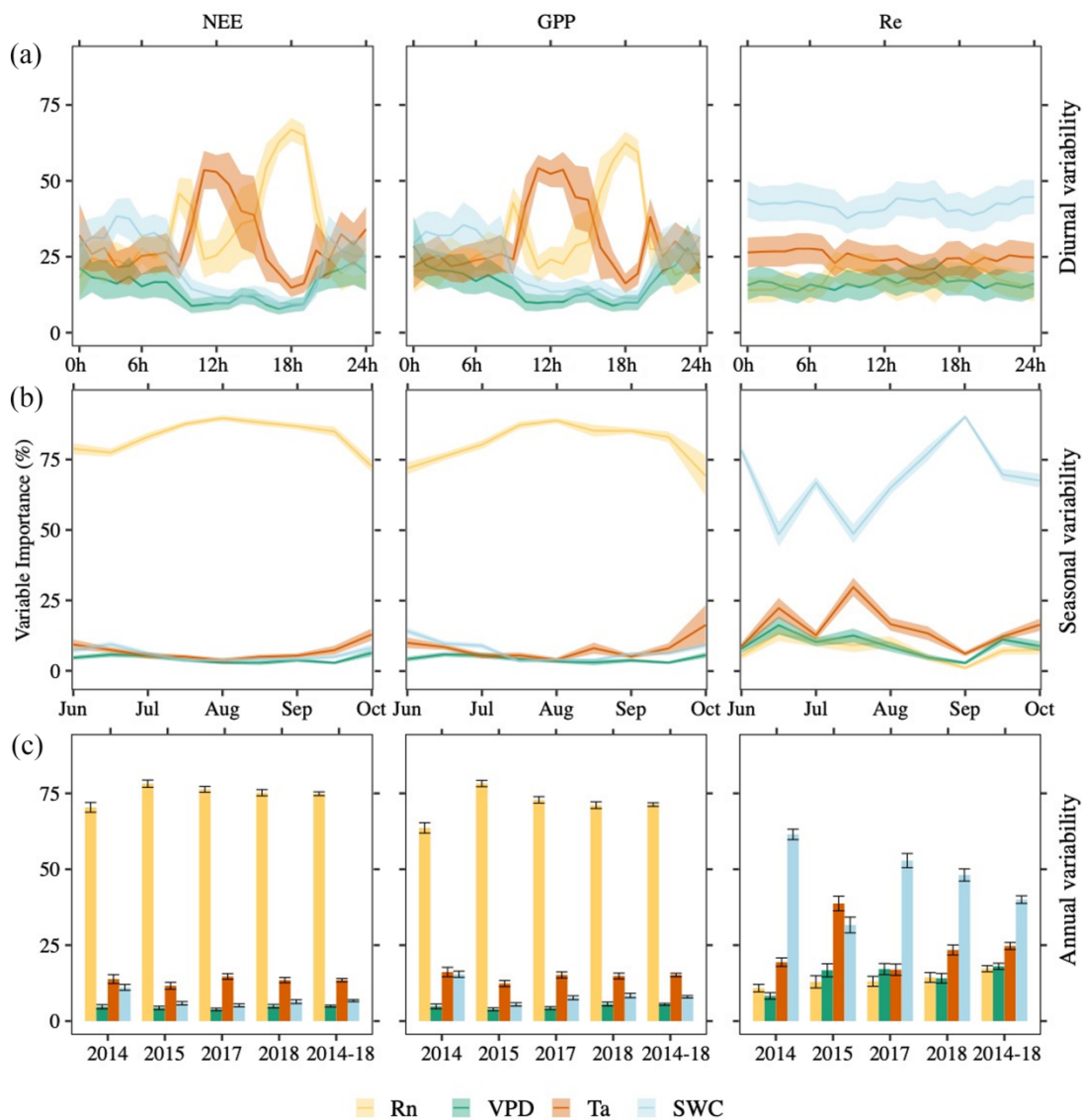
236 **Figure 4. Seasonal variability of daily 5-day moving average daily NEE, Re, and GPP in the swamp alpine**
 237 **meadow.**

238 **3.3 The importance of environmental forcing controlling C fluxes**

239 Our data processed by a machine learning technique suggest that the relative importance of the primary
 240 environmental drivers (R_n , T_a , VPD, and SWC) regulating terrestrial C fluxes varies on different time
 241 scales (i.e., diurnal, seasonal, and interannual scales) in this swamp meadow of QTP (Fig. 5(a), (b), (c)).

242 The diurnal variability of NEE and GPP was mostly driven by T_a (Fig. 5(a)), especially in the central
 243 hours of the day between 11:00 and 15:00 while Re showed a fairly lower temperature dependence
 244 compared to NEE and GPP (Fig. 5(a)). SWC was relatively more important than air temperature
 245 controlling the diurnal variability of respiration (Fig. 5(a)). The seasonal variability shaping the terrestrial
 246 C fluxes are regulated not only by meteorological variables but also by plant phenology. To separate the
 247 role of meteorological variables from phenology, we carried out a random forest analysis every fortnight
 248 and assumed that plant phenology changed little during this time span (Fig. 5(b); seasonal variability).
 249 The analyses based on random forest revealed a distinct seasonal pattern from June to September,
 250 pointing to a marked contribution of net radiation over NEE and GPP (Fig. 5(b)). Interestingly, Re was

251 found mostly regulated by SWC. On an annual scale, the contribution patterns of each environmental
 252 driver to the variations of C fluxes are similar to the ones found at the seasonal scale. The interannual
 253 variability of NEE and GPP were controlled more clearly by Rn while SWC revealed a stronger relative
 254 importance over Re (Fig. 5(a), (b), (c)). Overall, SWC dynamics seem to be the most important variable
 255 explaining the variability observed in the Re data (Fig. 5), suggesting that soil moisture plays an essential
 256 role on diurnal, seasonal, and interannual basis in this cold swamp meadow ecosystem. Note also that Ta
 257 played a secondary role controlling the Re at all assessed time scales (Fig. 5(a), (b), (c)).



258
 259 **Figure 5. Contribution to diurnal, seasonal, and annual variation of NEE, GPP, and Re from different**
 260 **environmental drivers (Rn (yellow), Ta (orange), SWC (blue), and VPD (green)). Solid lines (diurnal and**
 261 **seasonal variability) and bars with error bars (annual variability) both illustrate the average \pm standard**

262 deviation of the importance across 1000 decision trees. Annual variability refers to the variability of the
263 integrated growing season of 2014, 2015, 2017, and 2018.

264 4. Discussion

265 Since NEE is the difference between Re and GPP, environmental variables affecting Re and GPP could
266 affect NEE indirectly (Song et al., 2011). Radiation affects the magnitude of plant photosynthesis and
267 controls temperature, one of the key factors related to C fluxes. This suggests that abundant radiation
268 benefits photosynthesis and respiration and thus directly affects the C sink strength in the alpine wetland
269 ecosystem of the Qinghai Lake (Cao et al., 2017). Niu et al. (2017) show that 99% of the interannual
270 variation of NEE in an alpine swamp meadow can be well explained by temperature conditions,
271 precipitation and radiation. The results of this study demonstrate that ecosystem C sequestration is
272 regulated not only by radiation and temperature but also by soil moisture in the alpine swamp meadow
273 site studied herein. Given there is no significant difference in net radiation between the four years we
274 studied, the effects of soil water content and temperature on C fluxes on diurnal, seasonal, and annual
275 scales are therefore discussed in detail below.

276 4.1 Low soil moisture is associated with enhanced ecosystem respiration

277 A previous study in alpine swamp meadow ecosystems found that water stress may be the key limiting
278 factor leading to a decline in the photosynthetic rate at noon with a low SWC of 6–21% (Zhang et al.,
279 2018). In this alpine swamp, the soil layer maintained a relatively high SWC due to the frequent
280 precipitation during the growing season. SWC was always greater than 70% during the entire study
281 period (Fig. 2(b)). Therefore, microbial activity, and thus heterotrophic respiration were likely
282 suppressed by the anaerobic environment due to saturated soil water condition. At this site, SWC was
283 found to be a more important variable than temperature controlling the variability of ecosystem
284 respiration at different time aggregations (Fig. 5(a), (b), (c)). Our results suggest that even under water-
285 saturated conditions, the C dynamics of this alpine swamp meadow are still highly sensitive to changes
286 in soil moisture and could therefore be significantly influenced by future changes in water supply (Li et
287 al., 2015). In fact, previous studies have stressed that soil moisture will likely interact with temperature
288 to affect ecosystem respiration (Han et al., 2013) and therefore modify the overall C sink strength.
289 Moreover, Zhao et al. (2010) also noticed that the CO₂ emission rates decrease notably after rain events,

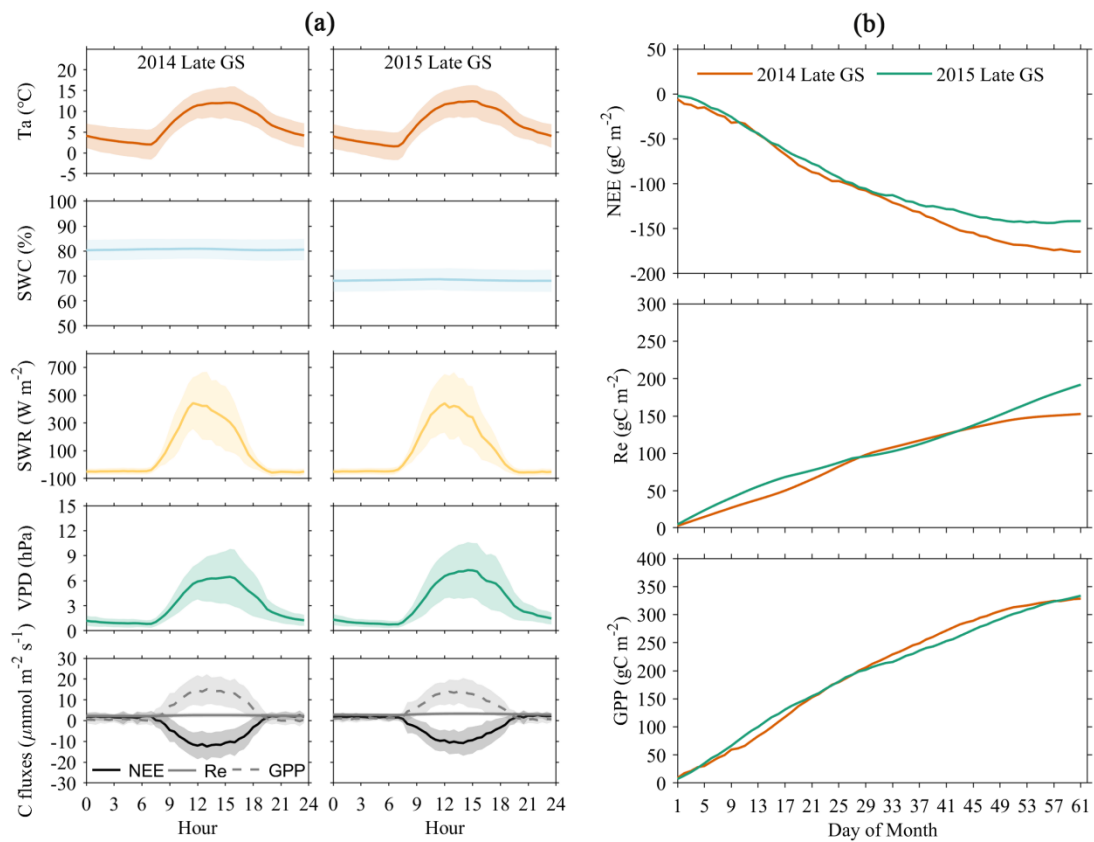
290 and Zhu et al. (2020) confirmed that annual precipitation exhibits significant impact on variation of
291 annual net C uptake. All these existing studies have suggested that C fluxes are related to water
292 availability condition, but few studies have found that soil moisture explicitly affects respiration.

293 Since C fluxes are affected by plant phenology and climate factors, including temperature, soil moisture,
294 and radiation simultaneously (Fig. 5), to analyze the effect of a single factor, ideally, other factors need
295 to be identical or at least close (no significant differences). Based on this theory and to better understand
296 the underlying mechanisms around how SWC interacts with the C fluxes in the studied alpine swamp
297 meadow ecosystem, we selected a specific group of data for further evaluation other than the entire
298 observation time. The group contains two late growing season periods, which have significant differences
299 in SWC but no significant difference in Ta (Fig. 6(a); Table S2).

300 The most significant difference in C fluxes between the late growing season of 2014 and 2015 was
301 observed in Re ($p < 0.05$). Additionally, on both diurnal and seasonal scales, a 15.4% decrease of SWC in
302 the late growing season of 2015 resulted in a 25.7% increase of Re and a 19.4% decrease in net C uptake
303 compared to 2014 (Figs. 6(a), (b); Table S2). This finding suggests that drier conditions likely prevented
304 this alpine swamp meadow from water-logged states, thereby strengthening soil respiration due to
305 improved soil aeration (Wang et al., 2014). According to the literature, the intensification of anaerobic
306 conditions due to water saturated soil can be responsible for weaker respiratory losses (Jansson and
307 Hofmockel, 2020).

308 Although the SWC in September 2015 was much greater than the 6–21% range reported by Zhang et al.
309 (2018), our data suggest that a 22.2% reduction in SWC in September 2015 resulted in a 51.6% decline
310 in net C uptake compared to September 2014 (Fig. S1; Table S3). There is evidence from the literature
311 that the rates of net C uptake in alpine wetlands during the growing season can be lower under drier
312 conditions (Hao et al., 2011), indicating that this alpine swamp meadow ecosystem may be adapted to
313 high levels of SWC (Li et al., 2015). Higher SWC may limit the diffusion of oxygen from the atmosphere
314 to the soil, inhibiting the activity of microorganisms and reducing the decomposition rate of soil organic
315 matter, decreasing the nutrients in the soil, and consequently reducing the photosynthetic in alpine
316 wetlands (Chimner and Cooper, 2003). Our comparisons suggest that drying can weaken the overall C
317 sink strength in this alpine swamp meadow ecosystem. Wetlands are predicted to experience lower water

318 tables due to permafrost degradation in the Tibetan Plateau and, therefore, permafrost thaw-induced
 319 wetland drying could enhance the response of C emissions to climate warming (Yu et al., 2020).



320

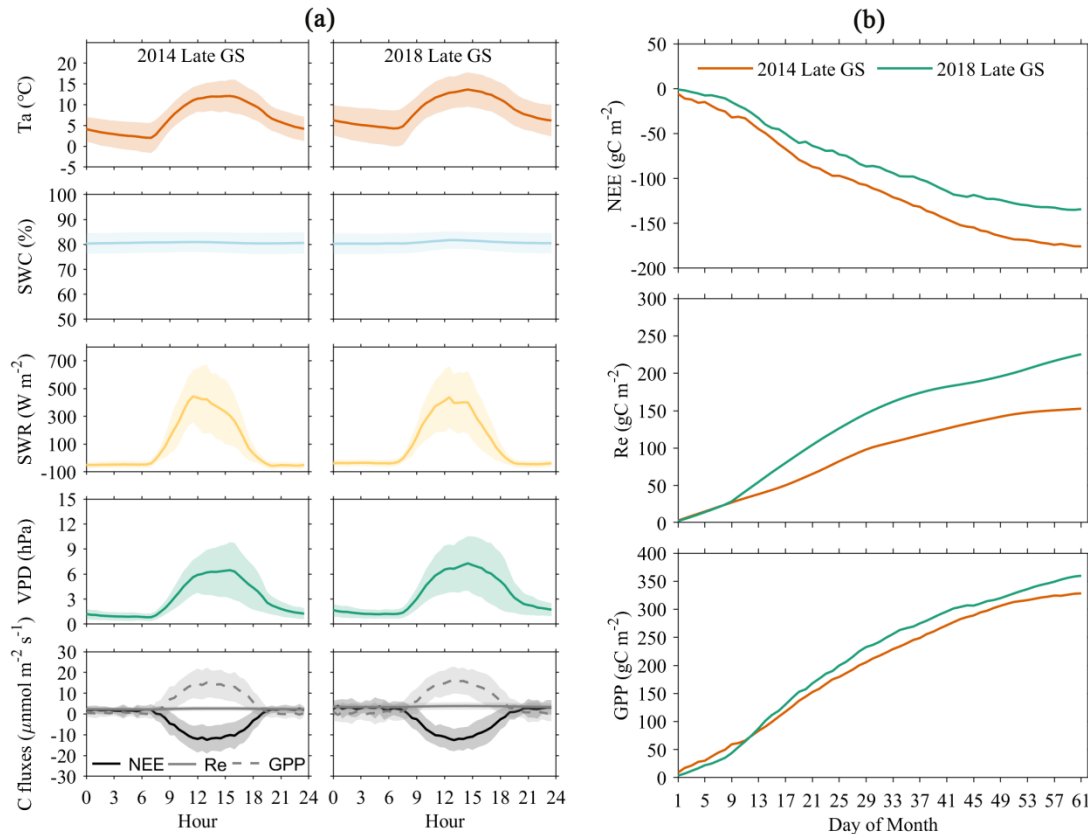
321 **Figure 6. (a) Comparisons of the diurnal variations of environmental drivers (Ta, SWC, Rn, and VPD) and**
 322 **C fluxes (NEE, Re, and GPP) between the late growing season of 2014 and 2015. The shading represents the**
 323 **mean ± standard deviation of the presented variables. (b) Comparisons of the daily accumulated C fluxes**
 324 **(NEE, Re, and GPP) between the late growing season of 2014 and 2015. Note: late GS represents late (Aug.–**
 325 **Sep.) growing season.**

326 4.2 Temperature increase leads to higher C losses rather than enhanced C uptake

327 The important role played by temperature controlling C exchange has been extensively found in alpine
 328 marshland across the QTP (Qi et al., 2021). For example, Zhao et al. (2010) and Zhao et al. (2005) show
 329 that ecosystem respiration follows the exponential variation of soil temperature. Zhu et al. (2020) also
 330 suggested that soil temperature plays the most important role in the change of monthly ecosystem
 331 respiration in the alpine wetland at Luanhaizi, northeastern Qinghai-Tibet Plateau. We therefore explored
 332 another comparison between the late growing season of 2014 and 2018 (Fig. 7(a), (b); Table S2) where

333 SWC was almost identical at both periods while temperature differed. Compared to 2014, a 25% increase
334 in T_a in the late growing season of 2018 led to joint larger GPP and R_e fluxes (Fig. 7a, b; Table S2).
335 Although both GPP and R_e increased, the intensification in R_e was greater than GPP, indicating that
336 warmer temperatures had a stronger impact on ecosystem respiration, resulting in a decrease of the net
337 C uptake (Fig. 7(b); Table S2). To evaluate if this finding is also consistent at an annual scale, we further
338 analyzed annual aggregated data. An annual comparison was made between 2014, 2017, and 2018 when
339 SWC were found insignificantly different while temperatures in 2017 and 2018 were 44.4% higher than
340 in 2014 (Table S4). Additionally, this 44.4% increase in T_a in 2017 and 2018 both led to stronger GPP
341 and R_e (Table S4). Although both GPP and R_e increased, the intensity in R_e was greater than GPP,
342 indicating that warmer temperatures have a stronger impact on ecosystem respiration in this site, resulting
343 in an approximately 50% decrease of the net C uptake (Table S4).

344 This comparison suggests that future warming could weaken the overall C sink strength in this alpine
345 swamp meadow ecosystem. Similar climate sensitivities have also been found in recent studies. For
346 example, a study performed by Niu et al. (2017) in an alpine swamp meadow on the central Tibet Plateau
347 suggests R_e was more sensitive to increased temperature than GPP. This suggests that global warming
348 may exacerbate future C releases in the alpine wetlands of the QTP (Gao et al., 2019; Niu et al., 2017;
349 Zhu et al., 2020). However, other researchers have also reached different conclusions; for instance, Qi
350 et al. (2021) found that GPP is consistently more sensitive than R_e to changes of temperature at daily,
351 seasonal, and annual scales, suggesting that cold condition can act as strong constraint on C uptake in
352 alpine marshlands. Liu et al. (2018) concluded that warming has a significant inhibitory effect on GPP
353 and minor effect on R_e , resulting in a weaker carbon sequestration capacity of their studied alpine
354 wetland ecosystem. Wei et al. (2021) also found that the uptake of C by plants will exceed the amount
355 of C release under warmer and wetter climate conditions at an annual scale based on manipulative
356 experiments and model simulations focused on the Tibetan Plateau. Their study is based on a longer-
357 term trend while our study only covers 4-years of year-round observations thus site-specific differences
358 in time and space scales may explain this variability. Nevertheless, such results indicating inconsistent
359 ecosystem responses suggest that there are still large uncertainties regulating the responses of C fluxes
360 to temperature variation and further work is still crucial.



361

362 **Figure 7. (a) Comparisons of the diurnal variations of environmental drivers (Ta, SWC, Rn, and VPD) and**
 363 **C fluxes (NEE, Re, and GPP) between the late growing season of 2014 and 2018. The shading represents the**
 364 **mean ± standard deviation of the presented variables. (b) Comparisons of the daily accumulated C fluxes**
 365 **(NEE, Re, and GPP) between the late growing season of 2014 and 2018. Note: late GS represents late (Aug.–**
 366 **Sep.) growing season.**

367 4.3 Combined effects of temperature and soil moisture on C exchange dynamics

368 The QTP experienced a higher rate of temperature increase than that of the Northern Hemisphere average
 369 (Zhang et al., 2013). The effects triggered by climate-induced warming over NEE in this area have been
 370 argued to either increase or decrease the net C balance NEE, or even have no effect whatsoever
 371 (Ganjurjav et al., 2018; Li et al., 2020; Wu et al., 2011; Zhu et al., 2017). These inconsistent responses
 372 could be due to water limitations offsetting the C balance or even reversing the effect of elevated
 373 temperatures, which change the decomposition and photosynthetic processes (Wu et al., 2011; Yu et al.,
 374 2013; Zhao et al., 2019). Alpine swamp meadows of the QTP have recently attracted much attention
 375 because they hold 5.9% (~ 1.98 Pg C) of the total grassland soil organic C (~ 33.52 Pg C). Such
 376 ecosystems have the highest organic C density (~ 50 kg C m⁻²) and play an important role in the global

377 C cycle (Niu et al., 2017). To test whether the observed SWC effects in this study were representative of
378 other sites in the QTP and put it into a broader perspective, we examined the temperature and
379 precipitation (as a proxy for SWC) impacts on NEE (Liu et al., 2016).

380 Our 4-year dataset revealed that this alpine swamp meadow functioned as a net C sink of -168.0
381 $\pm 62.5 \text{ gC m}^{-2} \text{ y}^{-1}$ at a 3571m asl. According to Wei et. al (2021), there are six observational studies about
382 C fluxes at Haibei, three of which focused on alpine swamp meadows or wetland. These alpine swamp
383 meadows were reported as a net C source while we our site showed a consistent C sink. The different
384 directions of C exchange suggest that there are still uncertainties in our understanding of C exchange in
385 this alpine swamp meadows, and further insights are obtained from studying multiple years of
386 observations. Therefore, further efforts are still needed to improve our projection of C balance change of
387 this ecosystem under changing climate.

388 The NEE observations from this study were within the NEE ranges of previous studies in similar alpine
389 swamp meadows located across the QTP ($-255.5 - 173.2 \text{ gC m}^{-2} \text{ y}^{-1}$) (Table 2). In addition, the NEE
390 estimates of this alpine swamp meadow show a stronger C sink strength than those from alpine meadows
391 ($-161.3 - 85.4 \text{ gC m}^{-2} \text{ y}^{-1}$) (Chai et al., 2017; Wang et al., 2017; Wu et al., 2020), alpine steppes ($-30 -$
392 $21.8 \text{ gC m}^{-2} \text{ y}^{-1}$) (Wang et al., 2018; Wang et al., 2020a; Wang et al., 2020b; Wu et al., 2010), and alpine
393 shrublands ($-14 - -67 \text{ gC m}^{-2} \text{ y}^{-1}$) (Zhao et al., 2005, 2006). This is likely a result of the inhibiting effects
394 of the nearly-saturated soil condition over soil respiration rather than by the lower temperatures (Sun et
395 al., 2021). Therefore, in permanently or seasonally inundated swamp meadows, high SWC may have
396 triggered lower C loss rates further benefiting C preservation. At our site, the higher C sink strength was
397 likely attributed to higher precipitation (and therefore higher SWC) and lower temperature, which created
398 colder and more humid conditions than other sites (Table 2). It has been demonstrated in the literature
399 that cold and humid conditions favour stronger C sinks in alpine meadow ecosystems (Fu et al., 2009).

400 The interannual comparison of the sites presented in Table 2 shows that under low annual precipitation
401 conditions ($\sim 300 \text{ mm}$), the joint effects of warming and reduced precipitation weakened the net C uptake
402 at Damxung (Niu et al., 2017), and even turned the C sink of Xiaobo Lake wetland into a C source when
403 comparing 2015 with 2012 and 2013 (Cao et al., 2017; Wu et al., 2018). Under relatively high annual
404 precipitation ($\sim 500 \text{ mm}$), the joint effects of warming and increased precipitation enhanced C release in

405 Haibei^a site when comparing 2006 to 2004 and 2005 (Zhao et al., 2010). This indicates that net C uptake
 406 under warming conditions can be weakened even under high annual precipitation rates.

407 **Table 2 Comparison of annual NEE (gC m⁻² y⁻¹) at different sites in the QTP.**

Site	Altitude (m)	Ecosystem	Year	Ta (°C)	P (mm)	Annual NEE	Reference
Haibei ^a (37°35'N, 101°20'E)	3200	AWM	2004	2.3	493.5	101.1	Zhao et al., 2010
			2005	2.2	475.2	44.0	
			2006	3.6	562.4	173.2	
Xiaopo Lake (36°42'N, 100°46'E)	3228	AWM	2015	2.8	304.3	54.6	Wu et al., 2018
			2012	1.2	357.0	-225.6	Cao et al., 2017
			2013	1.2	357.0	-255.5	
Damxung (30°28'N, 91°4'E)	4285	ASM	2009	3.4	208.9	-148.5	Niu et al., 2017
			2011	2.6	393.3	-190.8	
Haibei ^b (37°35'N, 100°00'E)	3571	ASM	2014	-0.9	662.8	-240.3	This study
			2015	-0.6	521.4	-200.1	
			2017	-0.5	661.2	-118.1	
			2018	-0.3	624.3	-113.4	

408 Note: P denotes precipitation, AWM denotes Alpine Wetland Meadow and ASM represents Alpine Swamp Meadow,
 409 Haibei^a and Haibei^b denotes different sites.

410 5. Conclusions

411 The alpine swamp meadow from the QTP presented in this study has been found to act as a consistent
 412 and strong sink of CO₂ (-168.0 ± 62.5 g C m⁻² y⁻¹). The results from a novel machine learning technique
 413 revealed that air temperature is the most important variable driving NEE and GPP on a diurnal scale,
 414 while net radiation has a stronger importance controlling the seasonal and interannual variability of the
 415 same fluxes. Soil moisture, however, has the largest influence over Re variability on diurnal, seasonal,
 416 and interannual scales, suggesting that soil water content is a key control on ecosystem respiration. In
 417 addition, air temperature plays a less important role in regulating the C exchange variability. This study
 418 reveals that both drying and warming can suppress net C uptake in water-saturated alpine swamp meadow
 419 ecosystems by enhancing ecosystem respiration. The response of net C uptake to climate warming further

420 indicates that the forecasted warming in the QTP will not always increase the net C sink strength. Our
421 results not only highlight the contributions of soil moisture in regulating C sequestration under high water
422 conditions but also support future process-based modelling initiatives focusing on alpine swamp meadow
423 ecosystem C dynamics.

424 **Data availability.** Post-processed data and scripts used in this paper are available from the authors upon
425 request (xyli@bnu.edu.cn).

426 **Author contribution**

427 JQ, XL, LL, TRC, XW, and ELB designed the research. JQ, LL, and ELB processed the data and
428 performed the analyses. YJ and HY helped collect the data. JQ and ELB created the visualization of the
429 outputs. JQ, XL, and ELB prepared the manuscript with contributions (writing, review, and edition) from
430 LL, TRC, and ZJ.

431 **Competing interests.** The authors declare that they have no conflict of interest.

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