

Radiation, soil water content, and temperature effects on carbon cycling in an alpine swamp meadow of the northeastern Qinghai-Tibetan Plateau

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Abstract. Predicted intensified climate warming will likely alter the ecosystem net carbon (C) uptake of the Qinghai-Tibetan Plateau (QTP). Variations in C sink/source responses to climate warming have been linked to water availability; however, the mechanisms by which net C uptake responds to soil water content in saturated swamp meadow ecosystems remain unclear. To explore how soil moisture and other environmental drivers modulate net C uptake in the QTP, field measurements were conducted using the eddy covariance technique in 2014, 2015, 2017, and 2018. The alpine swamp meadow presented in this study was a persistent and strong C sink of CO₂ (-168.0 ± 62.5 g C m⁻² y⁻¹, average \pm standard deviation) across the entire 4-year study period. A random forest machine-learning analysis suggested that the diurnal and seasonal variations of net ecosystem exchange (NEE) and gross primary productivity (GPP) were regulated by temperature and net radiation. Ecosystem respiration (Re), however, was found mainly regulated by the variability of soil water content (SWC) at different temporal aggregations, followed by temperature, the second contributing driver. We further explored how Re is controlled by nearly saturated soil moisture and temperature comparing two different periods featuring almost identical temperatures and significant differences on SWC and vice versa. Our data suggest that, despite the relatively abundant

35 water supply, periods with a substantial decrease of SWC or increase of temperature produced higher R_e
36 and therefore weakened the C sink strength. Our results reveal that nearly saturated soil conditions during
37 the growing seasons can help maintain lower ecosystem respiration rates and thus enhance the overall C
38 sequestration capacity in this alpine swamp meadow. We argue that soil respiration and subsequent
39 ecosystem C sink magnitude in alpine swamp meadows could likely be affected by future changes in soil
40 hydrological conditions caused by permafrost degradation or accelerated thawing-freezing cycling due
41 to climate warming.

42 **1. Introduction**

43 Wetlands play a significant role in the global carbon (C) cycle due to a large amount of C stored in their
44 soils. The Qinghai-Tibetan Plateau (QTP), with an average altitude of over 4,000 m a.s.l., has
45 approximately 10×10^4 km² of natural wetlands, of which $\sim 50\%$ (4.9×10^4 km²) are alpine swamp
46 meadows. These ecosystems are predominantly located in permafrost areas and are typically soil
47 nutrient-rich and water-logged (Bai et al., 2019; Zhao et al., 2005). Climate change and human
48 disturbance can have profound consequences on permafrost regions (Biskaborn et al., 2019) and
49 significantly impact their hydrological regimes (Lafrenière and Lamoureux, 2019). Hydrological regimes
50 have an important role in controlling wetland functioning (Bohn et al., 2007; Christensen et al., 2003),
51 and the changes of hydrological regimes may put the wetland functioning of the QTP under pressure
52 (Hrubby, 1999; Woodward and Wui, 2001; Foti et al., 2013).

53 The QTP is forecasted to be warmer and wetter in the future (Chen et al., 2015; Cheng et al., 2011).
54 Warming may accelerate the microbial breakdown of alpine soil organic C and subsequently increase
55 CO₂ emissions (Zhu et al., 2015a). Warming could also improve C sequestration capacity by enhancing
56 the photosynthetic inputs and growth rates of alpine plants (Fu et al., 2015). Therefore, the potential
57 increase of CO₂ emissions due to warming in alpine regions could be partially offset by enhanced C
58 uptake (Schuur et al., 2009), triggering different net C uptake responses to climate warming. For
59 example, an increase in temperature in the QTP has been associated with net C sinks in the Zoige alpine
60 wetlands (Kang et al., 2014) but also with net C sources in the Damxung alpine swamp meadow (Niu et
61 al., 2017).

62 According to recent studies in QTP alpine grasslands, water conditions such as soil water content (SWC)
63 can be a key factor that changes water-use patterns and ecophysiological characteristics of alpine plants
64 (Wu et al., 2019) and modulate the warming-mediated increase of ecosystem C uptake (Ganjurjav et al.
65 2016; Peng et al. 2014). Ecosystem C processes such as net C uptake and soil respiration may increase
66 with SWC in dry environments, but decrease with SWC in water-logged environments (Quan et al., 2019;
67 Taylor et al., 2017). Warming in conjunction with increased precipitation can turn an ecosystem from
68 net source to a sink of C (Zhao et al., 2019), increasing both photosynthesis and respiration rates during
69 warmer and wetter years (López-Blanco et al., 2017, 2018). However, when warming occurs in soils
70 associated with low moisture, soil drought can change ecosystems from C sinks to sources (Ganjurjav et
71 al., 2017). Studies in QTP alpine meadows have indicated that warming significantly stimulates
72 ecosystem net C uptake in wet years but does not affect ecosystem net C uptake in dry years, because
73 the positive effects of warming on net C uptake are compensated by the negative effects of lower soil
74 moisture (Peng et al., 2014).

75 However, the existing studies concerning ecosystem C dynamics on the QTP mainly focused on alpine
76 meadows (Saito et al., 2009; Zhao et al., 2005, 2010; Zhu et al., 2015b), only a few analyses have been
77 conducted to specifically characterize C dynamics in alpine swamp meadows (Zhao et al., 2010; Qi et
78 al., 2021; Liu et al., 2020; Zhu et al., 2020). The magnitudes and interannual variations of net ecosystem
79 exchange (NEE) in alpine wetlands from the QTP are proved to be closely related to radiation,
80 precipitation, and temperature (Cao et al., 2017; Niu et al., 2017). Temperature has been identified as an
81 important driver for ecosystem respiration (R_e) in alpine swamp meadows, and Zhao et al. (2005, 2010)
82 found that R_e follows an exponential relationship with soil temperature. Other studies also noticed that
83 rainfall is an important determinant of the interannual C sink/source strength in alpine swamp meadows
84 (Liu et al., 2019; Zhu et al. 2020). For example, CO_2 emissions were reported to decrease notably after
85 rain events (Zhao et al., 2010). Even though alpine swamp meadow ecosystems are characterised by high
86 SWC, the role of SWC on C cycling has often been neglected or assumed to be less important. Compared
87 to other factors, the effects of SWC on the net C uptake in alpine swamp meadows are still unclear.
88 Climate warming and the associated enhanced evapotranspiration and permafrost degradation may
89 change soil hydrology dramatically (Andresen et al., 2020; Zhao et al., 2019). Considering the critical
90 role that SWC played in regulating C uptake and soil respiration of other ecosystem types (Ganjurjav et

91 al. 2016; Peng et al., 2014; Quan et al., 2019; Taylor et al., 2017; Wu et al., 2019), it is important to
92 understand whether the change of SWC would aggravate the saturated water stress or trigger drought
93 effects on net C uptake in the alpine swamp meadow ecosystem under future climate warming.

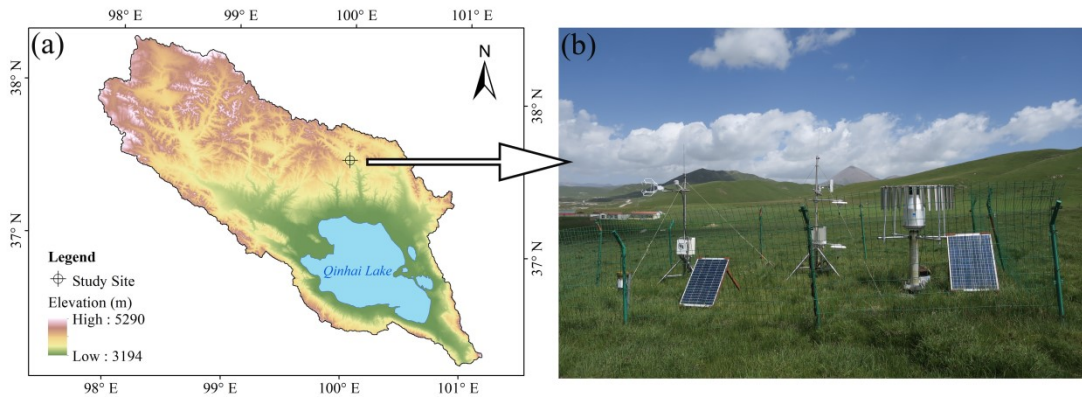
94 These uncertainties require a detailed investigation to understand wetland C source/sink processes and
95 their potential future C sink strength variations (sign and magnitude). In addition, as compared with
96 alpine meadows, there still needs long-term continuous observations for the alpine swamp meadow to
97 investigate C dynamics during dry and wet years. So a four-year field observation dataset is provided in
98 this study to characterize and quantify the importance of SWC in addition to temperature and net
99 radiation on the C sink strength of an alpine swamp meadow. Therefore, the objectives of this study are
100 to (i) quantify the diurnal and seasonal variations of net ecosystem exchange (NEE), gross primary
101 productivity (GPP), and ecosystem respiration (Re), (ii) identify and quantify the relative importance of
102 different key environmental drivers contributing to the variability observed of NEE, Re, and GPP, and
103 (iii) analyse how these C fluxes respond to soil water availability, temperature, and radiation variation in
104 a QTP alpine swamp meadow. This study would provide new insights into better understanding of the
105 complex C cycle dynamics in the Tibetan Plateau driven by the almost certain future intensified climate
106 warming.

107 **2. Materials and methods**

108 **2.1 Site description**

109 The study site (37°35.75' N, 100°00.47' E, 3,571 m a.s.l.) is located in Gangcha County, Qinghai
110 Province, in the north-eastern part of the QTP (Fig. 1(a)). The mean annual temperature and precipitation
111 measured at the Gangcha National Weather Station were 0.1 °C and 389.4 mm between 1982 and 2011,
112 respectively (Zhang et al., 2016). The site area has seasonal permafrost featuring frozen soils between
113 January–March and November–December for a total of about 125–135 days (Zhang et al., 2014). The
114 dominant species of the alpine swamp meadow ecosystem is *Kobresia pygmaea*, accompanied by
115 *Saussurea pulchra*, *Polygonum viviparum*, and *Potentilla saundersiana*. The average plant height at the
116 experimental site is 7.4 ± 1.5 cm, with a $97 \pm 2\%$ coverage. Our study defined the growing season as the
117 period between June and September. The early (June–July) and late (August–September) growing

118 seasons, as they will be henceforth referred to, correspond with the early growing season and late growing
 119 seasons, respectively.



120
 121 **Figure 1. (a) Location of the study site in the Qinghai Lake basin in the northeastern part of the QTP. (b)**
 122 **Eddy covariance system measuring water and CO₂ fluxes between the land surface and the atmosphere in the**
 123 **alpine swamp meadow.**

124 **2.2 Field measurements**

125 An Eddy Covariance (EC) system was installed at the study site (Fig. 1b) to measure the CO₂ fluxes at a
 126 sampling frequency of 10 Hz from 2014 to 2018. Data for 2016 was missing due to equipment
 127 malfunction. The EC system included an open-path CO₂/H₂O infrared gas analyser, which quantified
 128 fluctuations in water vapour and CO₂ fluxes. A 3-D sonic anemometer was installed at a 2.0 m height
 129 above ground to directly measure horizontal and vertical wind velocity components (u, v, and w). C flux
 130 data were recorded with a data logger (Campbell Scientific Inc.). An automated meteorological station
 131 was installed near the EC station to measure meteorological variables such as air temperature (T_a; °C),
 132 precipitation (P; mm), net radiation (R_n; W m⁻²), wind speed (WS; m s⁻¹), wind direction (WD; °), relative
 133 humidity (RH; %), and vapour pressure deficit (VPD; hPa). The meteorological data were collected at
 134 one-minute intervals and subsequently resampled at 30-minute timesteps to keep pace with the EC data.
 135 More details on the in-situ instrument specifications are summarised in Table 1.

136

137 **Table 1 Information about the sensors installed in the alpine swamp meadow.**

	Sensor Names	Sensor type	Installation height/depth	Manufacturer
Eddy Covariance	Open-path CO ₂ /H ₂ O infrared gas analyser	EC150	2.0 m	Campbell, US

	Three-dimensional sonic anemometer	CSAT3	2.0 m	Campbell, US
Meteorological observation	Net radiation	NR Lite	1.8 m	Kipp&Zonen, Netherlands
	Wind speed/direction	034B	2m	
	Air temperature/humidity	083E-1-6	0.5 m, 1.5 m	MetOne, US
	Atmosphere pressure	PTB110	In data acquisition box	Vaisala, Finland
	Rain-gauge	7852M-AB	0.7 m	Davis, US

138 Soil water content (SWC; %) were measured at depths of 10, 20, 40, 60, and 100 cm from the soil surface
139 with EC-H₂O sensors (Decagon Devices, USA) at a 10-minute recording frequency. The precision of the
140 EC-H₂O sensors for soil moisture measurements was $\pm 0.03 \text{ m}^3 \text{ m}^{-3}$. As the roots of *Kobresia* meadows
141 are mainly distributed within the top 20 cm of soil, we focused only on the variation of SWC in the top
142 20 cm of the soil.

143 2.3 Eddy Covariance data processing

144 The half-hourly NEE data was calculated using the EddyPro software (version 5.2, LI-COR) from the
145 10-Hz raw data. During the calculation, three-dimensional rotation was used to correct the data by
146 removing the effects of instrument tilt irregularity on airflow (Wilczak et al., 2001). Webb, Pearman, and
147 Leuning (WPL) (Webb et al., 1980) correction was applied to calculate the averages of CO₂ covariance,
148 rectifying the air density variations induced by heat and water vapour. The half-hourly flux data were
149 quality-checked based on several filtering algorithms, including: (1) the rejection of outliers in sonic
150 temperature, water vapour density, and CO₂ density (Li et al. 2008; Liu et al., 2011), (2) the elimination
151 of data one hour before and after precipitation events (Li et al. 2008; Liu et al., 2011), (3) the removal of
152 negative NEE during the non-growing season (from November to March) (Cao et al., 2017; Qi et al.,
153 2021) attributed to the self-heating effect from EC instruments (Cao et al., 2017), and (4) the exclusion
154 of measurements with weak turbulence conditions at night time. The weak turbulence periods were
155 identified by bootstrapping friction velocity (u^*) thresholds, as described by Papale et al. (2006). This
156 approach effectively divided the data into 4-year and 7-temperature subsets with similar micro-
157 meteorological conditions (except for u^*). The u^* thresholds (5%, 50%, and 95% of bootstrapping) were
158 calculated specifically per year and temperature subset.

159 Based on those different subsets, we gap-filled and partitioned NEE (into GPP and Re) to spread the
160 uncertainty variability that emerged from the different u^* thresholds, similar to López-Blanco et al.
161 (2017). All missing data were marked as -9999 (no data). Negative and positive NEE values represent
162 sink and source of C, respectively. Additionally, a standardised mechanism to fill NEE gaps is needed
163 for adequate data processing (Moffat et al., 2007). Therefore, this study adopted the method described
164 by López-Blanco et al. (2020) using the marginal distribution sampling (MDS) algorithm in the
165 REddyProc gap-filling tool (Reichstein et al., 2016), which was readapted from Reichstein et al. (2005).
166 Finally, NEE was separated into GPP and Re by applying the REddyProc partitioning algorithm
167 (Reichstein et al., 2016) for further analyses. This partitioning method is based on the exponential
168 regression of night-time respiration with temperature using the Lloyd-Taylor-Function (Lloyd and
169 Taylor, 1994). Night-time periods were selected via current combined solar radiation and potential
170 radiation thresholds based on the exact solar time, latitude, and longitude. REddyProc estimates the
171 temperature sensitivity from a short-term period, and based on this short-term temperature sensitivity, it
172 estimates the reference temperature in the continuous period of the entire dataset. These estimates were
173 then used to calculate Re for day-time and night-time, while GPP was estimated based on the difference
174 between NEE and Re.

175 **2.4 Identifying the importance of environmental drivers**

176 In order to characterize how environmental conditions impact diurnal and seasonal variability of NEE,
177 GPP and Re at this alpine swamp meadow, we used a novel method based on machine learning. Random
178 forest (Breiman, 2001) is a machine learning technique that can be used to quantify and interpret the
179 contribution of environmental drivers (covariates) to the variability of different C fluxes (response
180 variables) by combining multiple individual regression trees. This technique has been increasingly
181 utilized to upscale global C fluxes from eddy covariance data (Zeng et al., 2020) but also to evaluate
182 controls on C cycle processes (Zhang et al., 2017; López-Blanco et al., 2017, 2020). Here, we calculate
183 the relative importance of air temperature (T_a), net radiation (R_n), soil water content (SWC) and vapour
184 pressure deficit (VPD) controlling the C sink strength, photosynthesis, and respiration variability. This
185 random forest algorithm constructs multiple (1000 in this analysis) decision trees during training time
186 with different random subsamples (with replacement) from the same input training dataset. In each
187 cluster classified by random forest, the algorithm generates a multiple linear regression to characterize

188 different C fluxes as a function of environmental drivers (López-Blanco et al., 2017, 2020). This
189 algorithm version (Pedregosa et al., 2011) estimates the relative importance of each covariate between 0
190 and 100%, which correspond to the fraction of decision participating during data clustering. We used the
191 random forest algorithm to evaluate the diurnal and seasonal patterns of the relative importance of Ta,
192 Rn, SWC, and VPD responsible for the variability of C fluxes. We used data from the June-September
193 period aggregated per hour, and we ran multiple random forests with growing season data binned per
194 hour of the day, day of the year and yearly, respectively (Table S1).

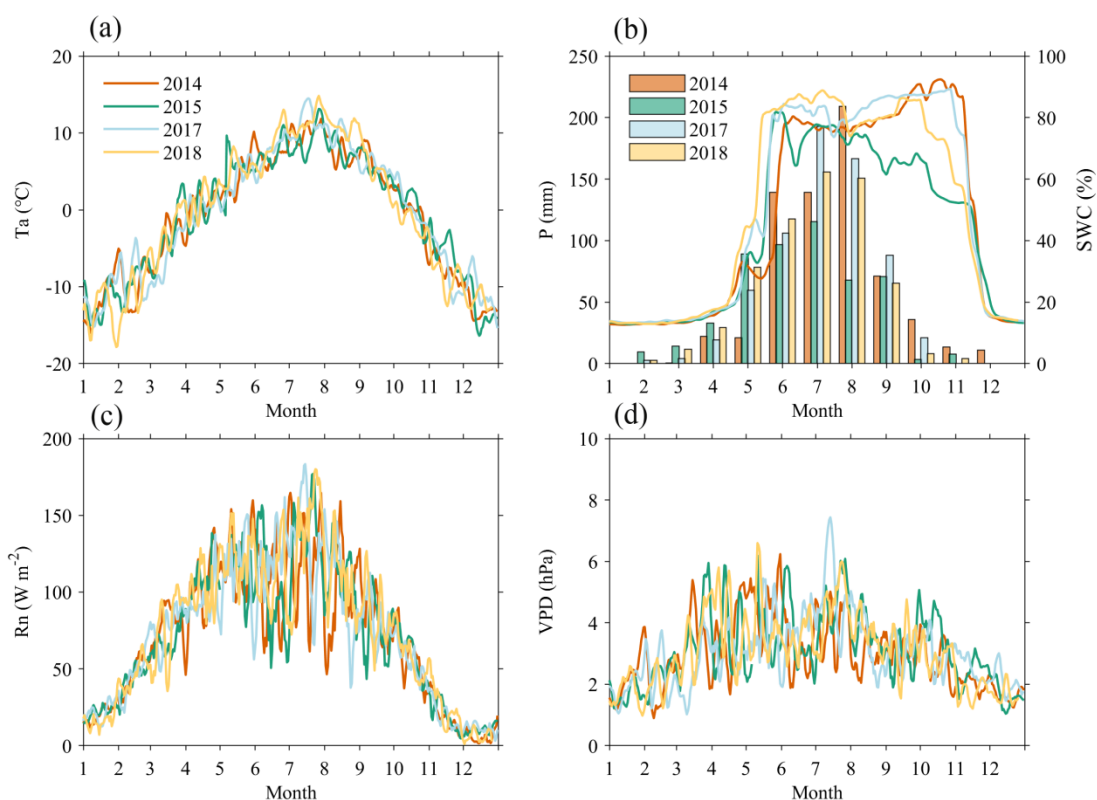
195 Since C fluxes are affected by plant phenology and climate factors, including temperature, soil moisture,
196 and radiation simultaneously, to analyze the effect of a single factor, ideally, other factors need to be
197 identical or at least closed (no significant differences). In each comparison, data of the same period in
198 each year were selected to exclude the influence of plant phenology. To further analyse the effect of soil
199 moisture, radiation, and temperature on C fluxes, we selected a specific group of data for further
200 evaluation other than the entire observation time. The group of data contains two late growing season
201 periods: periods with a significant difference in SWC but almost identical Ta and Rn (i.e., late growing
202 season of 2014 vs 2015) and periods with a significant difference in Ta but almost identical SWC and
203 Rn (i.e., late growing season of 2014 vs 2018). Additionally, in order to analyse the effect of annual
204 temperature on C fluxes, we selected a group of time stamps with significant differences in Ta but almost
205 identical SWC and Rn (i.e., 2017 vs 2014, and 2018 vs 2014). The magnitude of the differences between
206 C fluxes in the same group were analysed by the independent-sample T-test method.

207 **3. Results**

208 **3.1 Meteorological variability**

209 Mean daily meteorological variables (including Ta, P, SWC, and Rn) exhibited evident seasonal
210 variability except for VPD; these variables increased progressively in the early growing seasons, reached
211 their maximum in July and decreased gradually afterwards (Fig. 2(a), (b), (c), (d)). Air temperature
212 during the growing season was 7.7 ± 2.6 , 7.4 ± 2.6 , 8.5 ± 2.9 , and 9.2 ± 3.3 °C in 2014, 2015, 2017, and
213 2018, respectively while precipitation totalled 662.8, 521.4, 661.2, and 624.3 mm for the same years,
214 falling primarily during the growing season. The precipitation measured during the late growing season
215 of 2015 was only half of the amount measured in 2014, 2017, and 2018. The lower precipitation regime

216 led to a marked decline in SWC, making the late growing season of 2015 the driest period among all
 217 growing seasons during the study period. The greatest difference in SWC occurred in the late growing
 218 season of 2014 and 2015, when T_a at the same period were very closed (T_a was 6.8 ± 2.6 and $6.8 \pm 2.5^\circ\text{C}$
 219 in 2014 and 2015, respectively). Compared to 2014, SWC decreased by 15.4% in the late growing season
 220 of 2015 (Fig. 2(b); Table S2). Meanwhile, SWC in the late growing season of 2014 and 2018 was almost
 221 identical (80.7 ± 4.1 and 80.8 ± 3.8 , respectively), but the temperature difference was the largest (25%)
 222 compared to any other years (Fig. 2(a); Table S2).



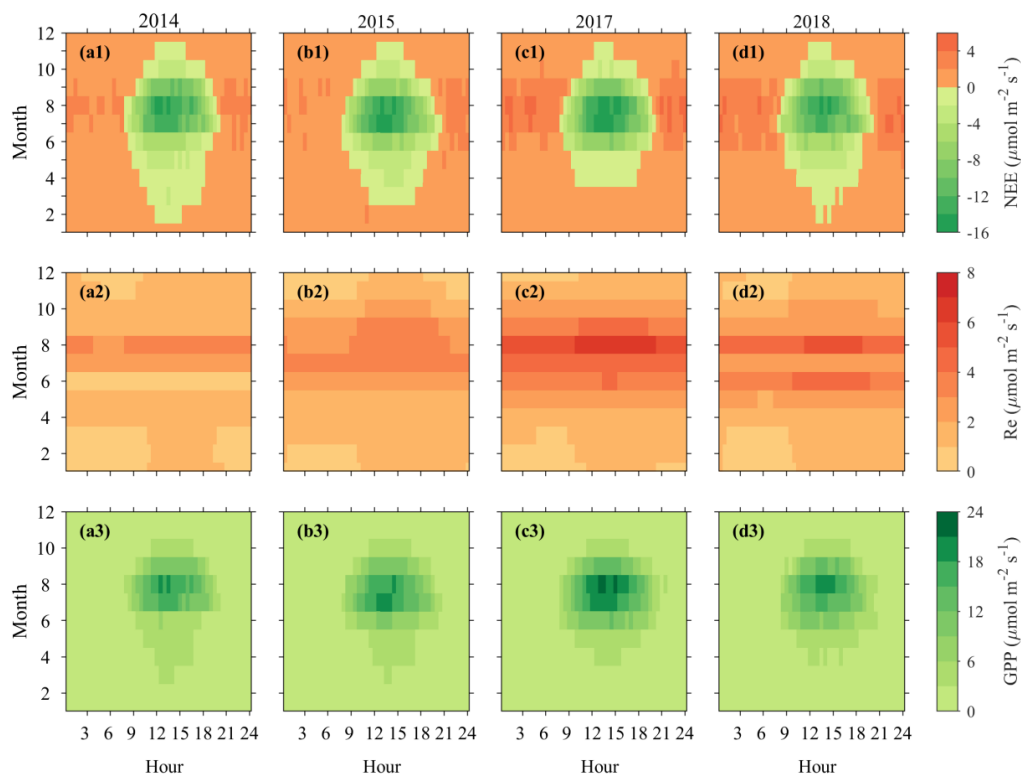
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224 **Figure 2. Five-day moving average meteorological variables (T_a , P, SWC, R_n , and VPD) in the studied swamp**
 225 **alpine meadow.**

226 3.2 Diurnal and seasonal variability of CO_2 fluxes

227 During the growing season, NEE (Fig. 3, (a1)–(d1)) and GPP (Fig. 3, (a3)–(d3)) featured a clear peak of
 228 the diurnal variations; both fluxes reached their summit between 12:00 and 14:00 local time. R_e ,
 229 however, presented a lower daily variability. The rates of NEE, R_e , and GPP during the growing season
 230 averaged -2.3 ± 0.3 , 3.2 ± 1.0 , and $5.5 \pm 0.9 \mu\text{mol m}^{-2} \text{s}^{-1}$, respectively, for the entire study period (2014,
 231 2015, 2017 and 2018). For the late growing seasons, the lowest rate of net C uptake was measured in

232 2015 ($-10.0 \mu\text{mol m}^{-2} \text{s}^{-1}$), whereas 2014 ($-12.4 \mu\text{mol m}^{-2} \text{s}^{-1}$), 2017 ($-12.2 \mu\text{mol m}^{-2} \text{s}^{-1}$), and 2018 (-12.5
 233 $\mu\text{mol m}^{-2} \text{s}^{-1}$) exhibited more negative NEE values (i.e., stronger net C uptake rate). Between the late
 234 growing season in 2015 and the late growing seasons in 2014, 2017 and 2018, there was a significant
 235 difference in the rates of Re ($p < 0.01$) while no significant difference was found in GPP variability
 236 ($p > 0.05$), suggesting that Re may be the component causing the difference observed in NEE. Specifically,
 237 the rates of Re in the late growing season of 2014 and 2015 were 2.4 ± 0.2 and $3.0 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$,
 238 respectively, which indicated that the drier conditions in 2015 generated a 25% higher Re compared to
 239 2014. In addition, the rate of Re in the late growing season was significantly higher in warmer 2018 (3.5
 240 $\pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$) than in 2014 ($2.4 \pm 0.2 \mu\text{mol m}^{-2} \text{s}^{-1}$).

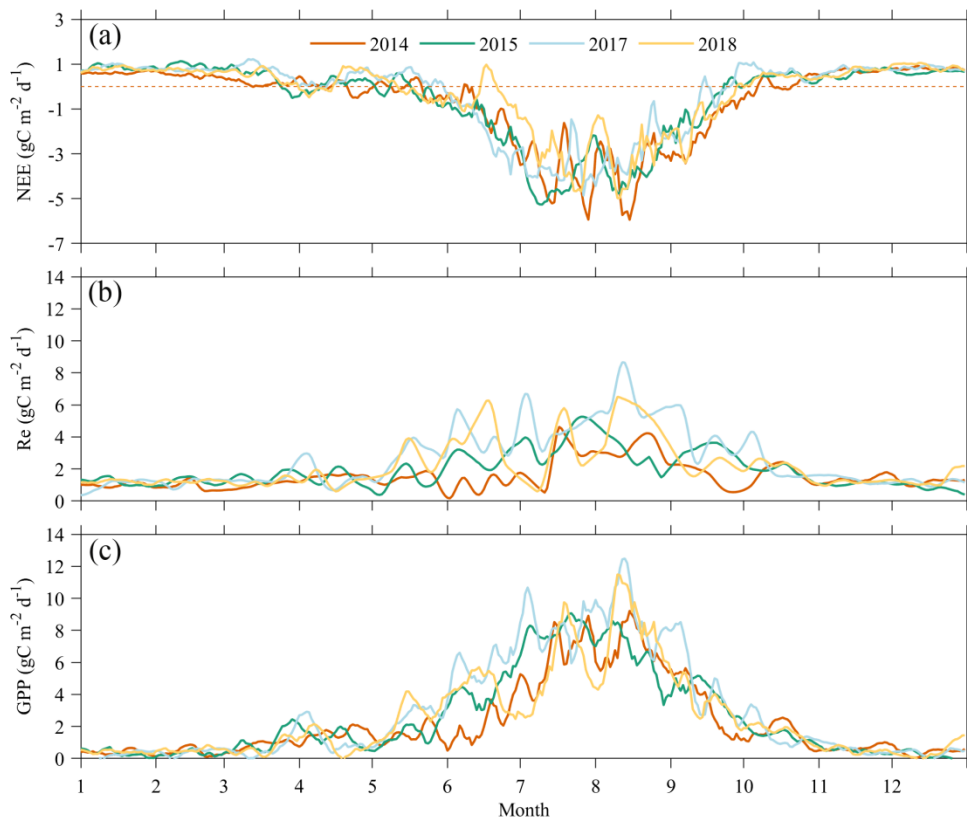


241

242 **Figure 3. Diurnal variability of gap-filled NEE and partitioned Re and GPP in 2014, 2015, 2017 and 2018.**

243 NEE, Re, and GPP also exhibited a strong seasonal variability; the C fluxes gradually increased from
 244 low values in early June to the maximum in the middle of the growing season (late July to early August
 245 on average), followed by a decrease towards the end of the growing season (Fig. 4(a), (b), (c)). The C
 246 sink strength across the growing season was found lowest in 2018, followed by 2017, whereas 2014 and
 247 2015 exhibited relatively higher values (Fig. 4(a)). Notably, the accumulated Re in the late growing

248 season was significantly higher in 2015 compared to 2014 ($P < 0.05$), while there was no significant
 249 difference in GPP (Fig. 4(b); Table S2). Moreover, the late growing season of 2015 witnessed the lowest
 250 SWC while keeping the same T_a compared to the same period in 2014 (Fig. 2(b); Table S2). The
 251 substantial decline of SWC observed in 2015 appeared to be responsible for the weaker observed C sink
 252 strength. On the other hand, T_a in the late growing season of 2018 was the highest for the same period
 253 among all the years while the SWC remained the same as 2014 (Fig. 2(a)). The late growing season of
 254 2018 showed overall higher GPP and R_e , but lower net C uptake than 2014. Significantly higher R_e in
 255 2018 caused by higher temperatures eventually led to a decrease of the C sink capacity (Fig. 4(a), (b),
 256 (c); Table S2).



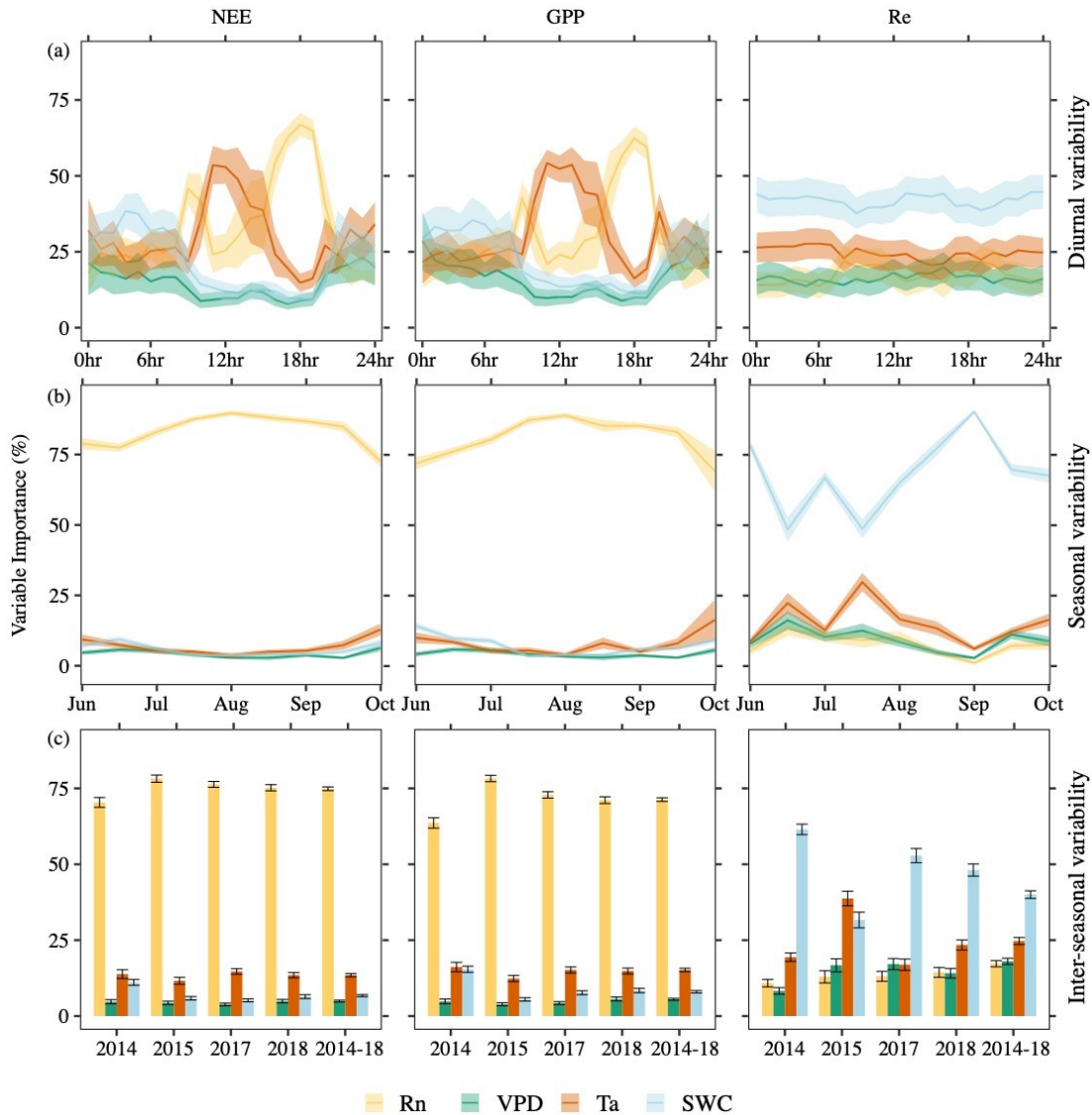
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258 **Figure 4. Seasonal variability of daily 5-day moving average daily NEE, R_e , and GPP in the swamp alpine**
 259 **meadow.**

260 3.3 The importance of environmental forcing controlling C fluxes

261 Our data processed by a machine learning technique suggest that the relative importance of the primary
 262 environmental drivers (R_n , T_a , VPD, and SWC) regulating C fluxes varies on different time scales (i.e.,
 263 diurnal and seasonal scales) in this swamp meadow of the QTP (Fig. 5(a), (b), (c)). The diurnal variability

264 of NEE and GPP was mostly driven by T_a (Fig. 5(a)), especially in the central hours of the day between
265 11:00 and 15:00, while R_e showed a fairly lower temperature dependence compared to NEE and GPP
266 (Fig. 5(a)). SWC was relatively more important than air temperature controlling the diurnal variability
267 of respiration (Fig. 5(a)). The seasonal variability shaping the terrestrial C fluxes are regulated not only
268 by meteorological variables but also by plant phenology. To separate the role of meteorological variables
269 from phenology, we carried out a random forest analysis every fortnight and assumed that plant
270 phenology changed little during this time span (Fig. 5(b); seasonal variability). The analyses based on
271 random forest revealed a distinct seasonal pattern for different C fluxes from June to September, pointing
272 to a marked contribution of net radiation over NEE and GPP (Fig. 5(b)). R_e , however, was found mostly
273 regulated by SWC. The contribution patterns of each environmental driver to the variations of C fluxes
274 on inter-seasonal scale (Fig. 5(c)) are similar to the ones found at the seasonal scale (Fig. 5(b)). The inter-
275 seasonal variability of NEE and GPP were explained clearly by R_n , while SWC revealed a stronger
276 relative importance over R_e . Overall, SWC dynamics seem to be the most important variable explaining
277 the variability observed in the R_e data, suggesting that soil moisture plays an essential role on diurnal
278 and seasonal basis in this cold swamp meadow ecosystem. Note also that T_a played a secondary role in
279 regulating R_e at all assessed time scales (Fig. 5(a), (b), (c)).



280

281 **Figure 5. Contribution to diurnal and seasonal variation of NEE, GPP, and Re from different environmental**
 282 **drivers (Rn (yellow), Ta (orange), SWC (blue), and VPD (green)). Solid lines with shades (diurnal and**
 283 **seasonal variability) and bars with error bars (inter-seasonal variability) both illustrate the average \pm**
 284 **standard deviation of the importance across 1000 decision trees. Inter-seasonal variability refers to the**
 285 **variability of the integrated growing season of 2014, 2015, 2017, and 2018.**

286 **4. Discussion**

287 Since NEE is the difference between Re and GPP, environmental variables affecting Re and GPP could
 288 affect NEE indirectly (Song et al., 2011). Radiation affects the magnitude of plant photosynthesis and
 289 controls temperature, one of the key factors related to C fluxes. Abundant radiation benefits
 290 photosynthesis and respiration and thus directly affects the C sink strength of an alpine wetland

291 ecosystem in the Qinghai Lake basin (Cao et al., 2017). In this study, radiation has been identified as the
292 most important factor that regulating diurnal and seasonal variations of GPP and NEE, but soil moisture
293 contributed the most to the variation of Re. On annual scale, Niu et al. (2017) show that 99% of the
294 interannual variation of NEE in an alpine swamp meadow can be well explained by temperature
295 conditions, precipitation and radiation. The results of this study demonstrate that ecosystem C
296 sequestration is regulated not only by radiation and temperature but also by soil moisture in the alpine
297 swamp meadow site studied herein. Given there were no significant differences in net radiation among
298 the four years we studied, the effects of SWC and temperature on C fluxes on diurnal and seasonal scales
299 are therefore discussed in detail below.

300 **4.1 Low soil moisture is associated with enhanced ecosystem respiration**

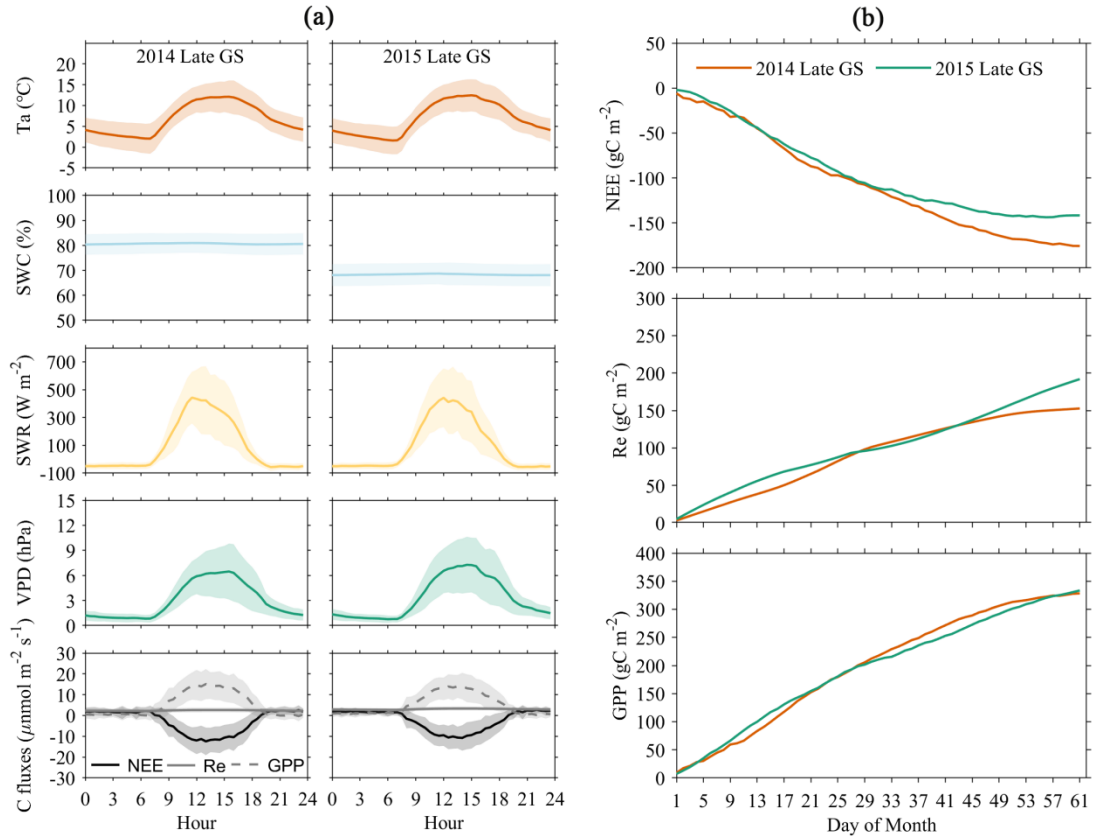
301 A previous study in alpine meadow ecosystems found that water stress may be the key limiting factor
302 leading to a decline in the photosynthetic rate at noon with a low SWC of 6–21% (Zhang et al., 2018).
303 In this alpine swamp meadow, the soil layer maintained a relatively high SWC due to the frequent
304 precipitation during the growing season. SWC was always greater than 70% during the entire study
305 period (Fig. 2(b)). Therefore, microbial activity, and thus heterotrophic respiration were likely
306 suppressed by the anaerobic environment due to saturated soil water condition (Chimner and Cooper,
307 2003; Sun et al., 2021). At this site, SWC was found to be a more important factor than temperature
308 controlling the variability of Re at different time aggregations (Fig. 5(a), (b), (c)). This result suggests
309 that even under water-saturated conditions, the C dynamics of this alpine swamp meadow are still highly
310 sensitive to changes in soil moisture and could therefore be significantly influenced by future changes in
311 water supply (Li et al., 2015). In fact, previous studies have stressed that soil moisture will likely interact
312 with temperature to affect Re (Han et al., 2013) and therefore modify the overall C sink strength.

313 To better understand the underlying mechanisms around how SWC interacts with the C fluxes in the
314 studied alpine swamp meadow ecosystem, we selected two late growing season periods, which have
315 significant differences in SWC but no significant difference in Ta (Fig. 6(a); Table S2).

316 The most significant difference in C fluxes between the late growing season of 2014 and 2015 was
317 observed in Re ($p < 0.05$). Additionally, on both diurnal and seasonal scales, a 15.4% decrease of SWC in
318 the late growing season of 2015 resulted in a 25.7% increase of Re and a 19.4% decrease in net C uptake

319 compared to 2014 (Figs. 6(a), (b); Table S2). According to Jansson and Hofmockel (2020), the
320 intensification of anaerobic conditions due to water saturated soil can be responsible for weaker
321 respiratory losses. This finding suggests that drier conditions in 2015 likely prevented this alpine swamp
322 meadow from water-logged states, thereby strengthening soil respiration due to improved soil aeration
323 (Wang et al., 2014).

324 Although the SWC in September 2015 was much greater than the 6–21% range reported by Zhang et al.
325 (2018), our data suggest that a 22.2% reduction in SWC in September 2015 resulted in a 51.6% decline
326 in net C uptake compared to September 2014 (Fig. S1; Table S3). There is evidence from the literature
327 that the rates of net C uptake in alpine wetlands during the growing season can be lower under drier
328 conditions (Hao et al., 2011), indicating that this alpine swamp meadow ecosystem may be adapted to
329 high levels of SWC (Li et al., 2015). Higher SWC may limit the diffusion of oxygen from the atmosphere
330 to the soil, inhibiting the activity of microorganisms and reducing the decomposition rate of soil organic
331 matter (Chimner and Cooper, 2003). Our comparisons suggest that drying can weaken the overall C sink
332 strength in this alpine swamp meadow ecosystem. Wetlands are predicted to experience lower water
333 tables due to permafrost degradation in the Tibetan Plateau and, therefore, permafrost thaw-induced
334 wetland drying could enhance the response of C emissions to climate warming (Yu et al., 2020).



335

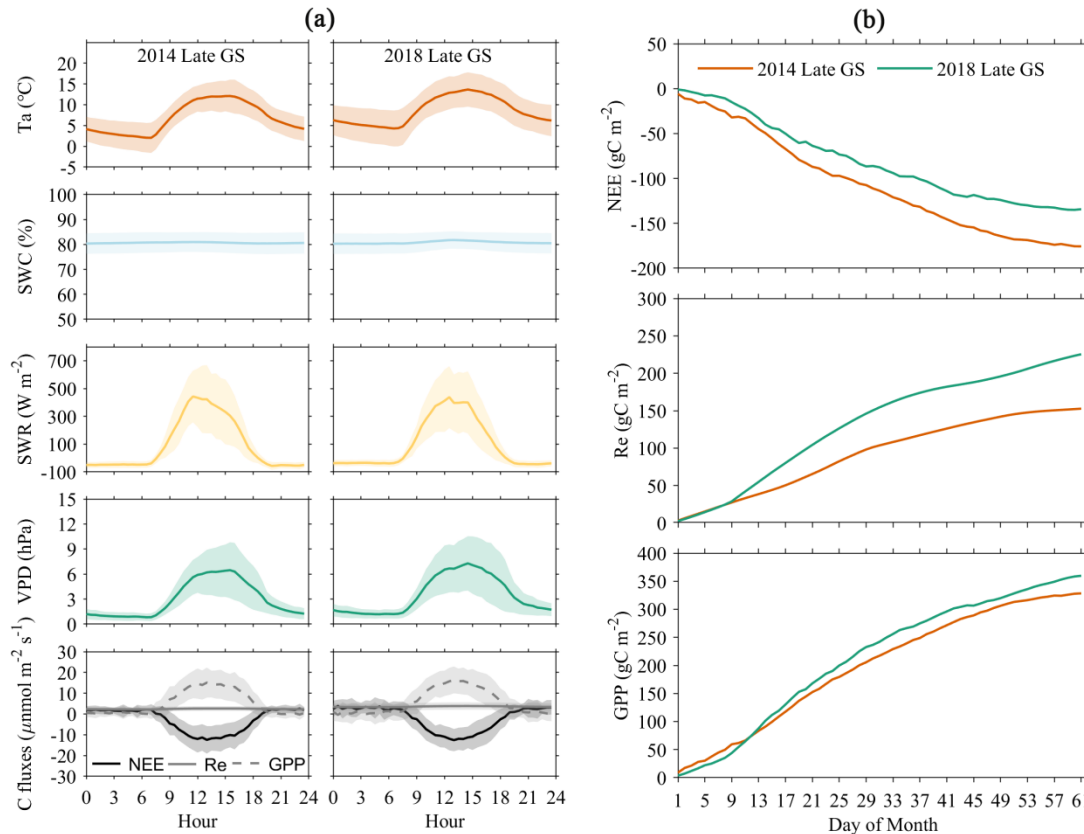
336 **Figure 6. (a) Comparisons of the diurnal variations of environmental drivers (Ta, SWC, Rn, and VPD) and**
 337 **C fluxes (NEE, Re, and GPP) between the late growing season of 2014 and 2015. The shading represents the**
 338 **mean ± standard deviation of the presented variables. (b) Comparisons of the daily accumulated C fluxes**
 339 **(NEE, Re, and GPP) between the late growing season of 2014 and 2015. Note: late GS represents late (Aug.–**
 340 **Sep.) growing season.**

341 4.2 Temperature increase leads to higher C losses

342 The important role played by temperature controlling C exchange has been extensively found in alpine
 343 marshland across the QTP (Qi et al., 2021). For example, Zhao et al. (2010) and Zhao et al. (2005) show
 344 that Re follows the exponential variation of soil temperature. Zhu et al. (2020) also suggested that soil
 345 temperature plays the most important role in the change of monthly Re in the alpine wetland at Luanhaizi,
 346 northeastern Qinghai-Tibetan Plateau. We therefore explored another comparison between the late
 347 growing season of 2014 and 2018 (Fig. 7(a), (b); Table S2) when phenology, radiation, and SWC were
 348 almost identical at both periods but temperature differed. Compared to 2014, a 25% increase in Ta in the
 349 late growing season of 2018 led to joint larger GPP and Re fluxes (Fig. 7a, b; Table S2). Although both
 350 GPP and Re increased, the intensification in Re was greater than GPP, indicating that warmer

351 temperatures had a stronger impact on Re, resulting in a decrease of the net C uptake (Fig. 7(b); Table
352 S2). To evaluate if this finding is also consistent at an annual scale, we further analyzed annual
353 aggregated data. An annual comparison was made between 2014, 2017, and 2018 when SWC were found
354 insignificantly different while temperatures in 2017 and 2018 were 44.4% higher than in 2014 (Table
355 S4). Again,, this 44.4% increase in Ta in 2017 and 2018 both led to stronger GPP and Re (Table S4).
356 Although both GPP and Re increased, the intensity in Re was greater than GPP, indicating that warmer
357 temperatures have a stronger impact on Re at this site, resulting in an approximately 50% decrease of the
358 net C uptake (Table S4).

359 This comparison suggests that future warming could weaken the overall C sink strength in this alpine
360 swamp meadow ecosystem. Similar climate sensitivities have also been found in recent studies. For
361 example, a study performed by Niu et al. (2017) in an alpine swamp meadow on the central Tibetan
362 Plateau suggests Re was more sensitive to increased temperature than GPP. This suggests that global
363 warming may exacerbate future C releases in the alpine wetlands of the QTP (Gao et al., 2019; Niu
364 et al., 2017; Zhu et al., 2020). Liu et al. (2018) concluded that warming has a significant inhibitory effect
365 on GPP and minor effect on Re, resulting in a weaker carbon sequestration capacity of their studied
366 alpine wetland ecosystem. However, other researchers have also reached different conclusions. For
367 instance, Qi et al. (2021) found that GPP is consistently more sensitive than Re to changes of temperature
368 at daily, seasonal, and annual scales, suggesting that cold condition can act as strong constraint on C
369 uptake in alpine marshlands. Wei et al. (2021) also found that the uptake of C by plants will exceed the
370 amount of C release under warmer and wetter climate conditions based on manipulative experiments and
371 model simulations for the Tibetan Plateau. Their study is based on a longer-term trend while our study
372 only covers 4-years of year-round observations thus site-specific differences in time and space scales
373 may explain this variability. These inconsistent ecosystem responses suggest that there are still large
374 uncertainties regulating the responses of C fluxes to temperature variation and further work is still crucial.



375

376 **Figure 7. (a) Comparisons of the diurnal variations of environmental drivers (T_a , SWC, Rn, and VPD) and**
 377 **C fluxes (NEE, Re, and GPP) between the late growing season of 2014 and 2018. The shading represents the**
 378 **mean \pm standard deviation of the presented variables. (b) Comparisons of the daily accumulated C fluxes**
 379 **(NEE, Re, and GPP) between the late growing season of 2014 and 2018. Note: late GS represents late (Aug.–**
 380 **Sep.) growing season.**

381 4.3 Combined effects of temperature and soil moisture on C exchange dynamics

382 The QTP experienced a higher rate of temperature increase than that of the Northern Hemisphere average
 383 (Zhang et al., 2013). The effects triggered by climate-induced warming over NEE in this area have been
 384 argued to either increase or decrease the net C balance NEE, or even have no effect whatsoever
 385 (Ganjurjav et al., 2018; Li et al., 2020; Wu et al., 2011; Zhu et al., 2017). These inconsistent responses
 386 could be due to water limitations offsetting the C balance or even reversing the effect of elevated
 387 temperatures, which change the decomposition and photosynthetic processes (Wu et al., 2011; Yu et al.,
 388 2013; Zhao et al., 2019a). Alpine swamp meadows of the QTP have recently attracted much attention
 389 because they hold 5.9% (~ 1.98 Pg C) of the total grassland soil organic C (~ 33.52 Pg C). Such
 390 ecosystems have the highest organic C density (~ 50 kg C m^{-2}) and play an important role in the global

391 C cycle (Niu et al., 2017). To test whether the observed SWC effects in this study were representative of
392 other sites on the QTP and put it into a broader perspective, we examined the temperature and
393 precipitation (as a proxy for SWC) impacts on NEE (Liu et al., 2016).

394 The NEE observations from this study were within the NEE ranges of previous studies in similar
395 ecosystems located across the QTP ($-255.5 - 173.2 \text{ g C m}^{-2} \text{ y}^{-1}$) (Table 2). According to Wei et al. (2021),
396 there are six observational studies about C flux around our study site, three of them are focused on alpine
397 swamp meadows. Among them, one study had one-year dataset (Zhang et al., 2008), and the other two
398 characterized the same location (Zhao et al., 2005, 2010). The three studies were reported as a net C
399 source, while our 4-year dataset revealed that this alpine swamp meadow functioned as a net C sink of -
400 $168.0 \pm 62.5 \text{ g C m}^{-2} \text{ y}^{-1}$ at a 3571 m asl.. The different directions of C exchange suggest that there are
401 still uncertainties in our understanding of C exchange in alpine swamp meadows, and further efforts are
402 still needed to improve our projection of C balance change of this ecosystem under changing climate.

403 In addition, the NEE estimates of this alpine swamp meadow show a stronger C sink strength than those
404 from alpine meadows ($-161.3 - 85.4 \text{ g C m}^{-2} \text{ y}^{-1}$) (Chai et al., 2017; Wang et al., 2017; Wu et al., 2020),
405 alpine steppes ($-30 - 21.8 \text{ g C m}^{-2} \text{ y}^{-1}$) (Wang et al., 2018; Wang et al., 2020a; Wang et al., 2020b; Wu
406 et al., 2010), and alpine shrublands ($-14 - -67 \text{ g C m}^{-2} \text{ y}^{-1}$) (Zhao et al., 2005, 2006). This is likely a result
407 of the inhibiting effects of the nearly-saturated soil condition over soil respiration (Sun et al., 2021). In
408 permanently or seasonally inundated swamp meadows, high SWC may have triggered lower C loss rates
409 and further benefited C preservation. At our site, the higher C sink strength was likely attributed to higher
410 precipitation (and therefore higher SWC) and lower temperature, which created colder and more humid
411 conditions than other sites (Table 2). It has been demonstrated that cold and humid conditions favour
412 stronger C sinks in alpine meadow ecosystems (Fu et al., 2009).

413 The interannual comparison of the sites presented in Table 2 shows that under low annual precipitation
414 conditions ($\sim 300 \text{ mm}$), the joint effects of warming and reduced precipitation weakened the net C uptake
415 at Damxung (Niu et al., 2017), and even turned the C sink of Xiaobo Lake wetland into a C source when
416 comparing 2015 with 2012 and 2013 (Cao et al., 2017; Wu et al., 2018). Under relatively high annual
417 precipitation ($\sim 500 \text{ mm}$), the joint effects of warming and increased precipitation enhanced C release in
418 Haibei^a site when comparing 2006 to 2004 and 2005 (Zhao et al., 2010). This indicates that net C uptake
419 under warming conditions can be weakened even under high annual precipitation rates.

420

421 **Table 2 Comparison of annual NEE (g C m⁻² y⁻¹) at different sites in the QTP.**

Site	Altitude (m)	Ecosystem	Year	Ta (°C)	P (mm)	Annual NEE	Reference
Haibei ^a (37°35'N, 101°20'E)	3200	AWM	2004	2.3	493.5	101.1	Zhao et al., 2010
			2005	2.2	475.2	44.0	
			2006	3.6	562.4	173.2	
Xiaopo Lake (36°42'N, 100°46'E)	3228	AWM	2015	2.8	304.3	54.6	Wu et al., 2018
			2012	1.2	357.0	-225.6	Cao et al., 2017
			2013	1.2	357.0	-255.5	
Damxung (30°28'N, 91°4'E)	4285	ASM	2009	3.4	208.9	-148.5	Niu et al., 2017
			2011	2.6	393.3	-190.8	
Haibei ^b (37°35'N, 100°00'E)	3571	ASM	2014	-0.9	662.8	-240.3	This study
			2015	-0.6	521.4	-200.1	
			2017	-0.5	661.2	-118.1	
			2018	-0.3	624.3	-113.4	

422 Note: P denotes precipitation, AWM denotes Alpine Wetland Meadow and ASM represents Alpine Swamp Meadow,

423 Haibei^a and Haibei^b denotes different sites.

424 **5. Conclusions**

425 The alpine swamp meadow from the QTP presented in this study has been found to act as a consistent
 426 and strong sink of CO₂ (-168.0 ± 62.5 g C m⁻² y⁻¹). The results from a novel machine learning technique
 427 revealed that air temperature is the most important variable driving NEE and GPP on a diurnal scale,
 428 while net radiation has a stronger importance controlling the seasonal variability of the same fluxes. Soil
 429 moisture, however, has the largest influence over Re variability on diurnal and seasonal scales,
 430 suggesting that SWC is a key control on Re. In addition, air temperature played a less important role in
 431 regulating Re. This study reveals that both drying and warming can suppress net C uptake in water-
 432 saturated alpine swamp meadow ecosystems by enhancing Re. The response of net C uptake to climate
 433 warming further indicates that the forecasted warming in the QTP will not always increase the net C sink
 434 strength. Our results not only highlight the contributions of soil moisture in regulating C sequestration

435 under high water conditions but also support future process-based modelling initiatives focusing on
436 alpine swamp meadow ecosystem C dynamics.

437 **Data availability.** Post-processed data and scripts used in this paper are available from the authors upon
438 request (xyli@bnu.edu.cn).

439 **Author contribution**

440 JQ, XL, LL, TRC, XW, and ELB designed the research. JQ, LL, and ELB processed the data and
441 performed the analyses. YJ and HY helped collect the data. JQ and ELB created the visualization of the
442 outputs. JQ, XL, and ELB prepared the manuscript with contributions (writing, review, and edition) from
443 LL, TRC, and ZJ.

444 **Competing interests.** The authors declare that they have no conflict of interest.

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