Climate and topography: two essential ingredients in predicting wetland permanence.

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Abstract. Wetlands in the Prairie Pothole Region (PPR) are forecast to retract in their ranges due to climate change and potholes that typically contain ponded water year-round, which support a larger proportion of biological communities, are

10   most sensitive to climate change. In addition to climate, land use activities and topography also influence ponded water amounts in PPR wetlands. However, topography is not typically included in models forecasting the impacts of climate change on PPR wetlands. Using a combination of variables representing climate, land cover and land use, and topography, we predicted wetland permanence class in the southern Boreal, Parkland and Grassland of the Alberta PPR. We show that while climate is the strongest predictor of wetland permanence class in each Natural Region, topography was nearly as

15   important in the Parkland and Southern Boreal.

1. Introduction

Wetlands provide habitat for diverse communities of flora and fauna (Gibbs, 1993; Loesch et al., 2012; Sundberg et al., 2016) and deliver ecosystem services of disproportionate importance relative to the area they occupy (Mitsch and Gosselink, 2015). The diversity and abundance of flora and fauna in these wetland ecosystems (Daniel et al., 2019; Gleason and Rooney, 2018),

20   is a function of the consistency with which ponded water is available (i.e., pond permanence), which is forecast to decline in amount and duration of presence (i.e., hydroperiod) across the prairie pothole region of North America due to climate change (Euliss et al., 2004; Fay et al., 2016; Steen et al., 2014, 2016). In this region, the majority of wetlands are ponded non-permanently and they support resident biological communities (Daniel et al., 2019; Stewart and Kantrud, 1971) that are sensitive to climate change (Fay et al., 2016; Johnson et al., 2010). Therefore, understanding the relative influence of climate

25   on wetland water levels is critical to improving our understanding of how biological communities in the Prairie Pothole Region (PPR) will respond to climate change.

Alberta lies at the western edge of the Prairie Pothole Region, which encompasses the province's Grassland and Parkland Natural Regions, as well as the southern edge of the Boreal (Schneider, 2013). Given the PPR's semi-arid climate, a decline in wetland hydroperiod is expected because of increases in wetland water deficits (Schneider, 2013; Werner et al.,

30 2013). Simulations for the PPR suggest that the magnitude of change in climatic conditions between 1946 and 2005 were vast enough to drive declines in pond permanence (Werner et al., 2013). Modelling suggests that these wetlands may experience up to a 20% decline in precipitation due to climate change, which could reduce hydroperiods (Fay et al., 2016). Furthermore, forecasts suggest that many of the wetlands in the southern and western PPR may lose their ponded water completely, driven by drier climate conditions in these areas (Johnson et al., 2005, 2010; Reese and Skagen, 2017). Wetlands that contain ponded

35 water year-round will be most sensitive to climate change because they contain water in late summer, when they will be subjected to greater evapotranspiration-driven losses (Fay et al. 2016). They are also relatively rare (Ridge et al., 2021). In addition to climate, topography can also affect hydroperiods in PPR wetlands (Johnson et al., 2010; McCauley et al., 2015; Tsai et al., 2012). The potholes, in which these wetlands are located, form a relic of the land's glaciated history and larger catchments contribute more water resulting in larger water budgets and longer hydroperiods for some pothole wetlands than

40 others (Hayashi et al., 2016; Shaw et al., 2013). Contemporary land-use practices (e.g., filling and ditching) also alter topography, affecting flows of surface and groundwater and subsequently wetland hydroperiod. This phenomenon, referred to as consolidation drainage, fully or partially drains upper-watershed wetlands and directs their water to areas lower in the watershed (McCauley et al., 2015). Consolidation drainage is typically done to lower the probability that neighbouring croplands will flood (Schindler and Donahue, 2006; Verhoeven and Setter, 2010), which increases farming efficiency

45 (Wiltermuth and Anteau, 2016).

Changes in land use can influence wetland hydroperiods by more than associated terrain modification. For example, landscapes with a higher proportion of agricultural activities can have longer hydroperiods due to the combination of increased surface run-off and decreased soil infiltration (van der Kamp et al., 2003; Voldseth et al., 2007). Many studies assessing the impacts of climate change on PPR wetlands incorporate land use (Anteau et al., 2016; Vodseth et al., 2009) and there is

50 resounding evidence that wetlands exposed to the same climate regime, but situated among different land-use activities, differ in their sensitivity to climate change (McCauley et al., 2015; Wiltermuth and Anteau, 2016).

While topography is an important predictor of pond permanence (Hayashi et al., 2016; Neff and Rosenberry, 2017; Shaw et al., 2013; Wiltermuth and Anteau, 2016), it is rarely included in studies assessing the impacts of climate change on PPR wetlands and/or biota (Wolfe et al., 2019) Even well-established models (e.g., WETSIM (Poiani and Johnson, 1993),

55 WETLANDSCAPE (Johnson et al., 2010)), applied to the PPR, predict pond permanence in response to climate, but omit topography. For example, differences in topography may cause wetlands belonging to the same permanence class to differ in their sensitivity to climate change. Consequently, our failure to incorporate topography when predicting pond permanence leaves us with an incomplete understanding on how wetland biota are affected by climate change.

Incorporating the influence of topography individually and in combination with climate and land cover/land use

60 effects on wetland permanence is a gap we must fill to improve wetland and waterfowl population management across the PPR (Fay et al., 2016). To overcome this gap, we analyse data collected across multiple field projects and use spatial data, comprising thousands of wetlands across the PPR in Alberta, Canada. Only four permanence classes (of seven) are represented in this study (Table 1), and which permanence class a wetland belongs to is determined by the vegetation zone in the deepest

part of the wetland – and this is dictated by its typical hydroperiod/pond permanence over several years (Stewart and Kantrud, 1971). Using these data, we quantify the relative contribution of climate, land cover/land use and topography for different wetland permanence classes. We also determine the ability of these drivers to predict wetland permanence class.

2. Methods

2.1. Study Area

The wetlands in our study are in the Albertan extent of the Prairie Pothole Region (PPR) (Figure 1). Wetlands in this region are called potholes because they are depressions filled with ponded water, each formed in the last glacial period (Wright, 1972). Spring snow melt is the largest contributor to ponded water amounts, either from direct precipitation into the wetland or as runoff over frozen ground from upland areas (Hayashi et al., 1998). Potholes can differ in the length of time they contain ponded water, which can range from a few weeks after snowmelt to the entire year (Stewart and Kantrud, 1971).

Publicly available wetland data (Canadian Wetlands; published by Environment and Climate Change Canada) for our study area lack definition of permanence class or measurements (e.g., water volume, depth) that could be used to classify the wetlands by permanence type. Despite this challenge, we acquired data from two wetland inventories (Government of Alberta, 2014) that delineated the location, boundary and permanence class of PPR wetlands based on Stewart and Kantrud's classification (Stewart and Kantrud, 1971) (Table 1). The two wetland inventories differ in their accuracy (Evans et al., 2017) and include wetlands from the Grassland, Parkland and the southern edge of the Boreal Natural Regions of Alberta. The Grassland comprises mixed-grass prairie, and the Parkland comprises deciduous trees and grasses. Both are semi-arid regions with potential evapotranspiration rates that are greater than annual precipitation (Downing and Pettapiece, 2006). The Parkland, however, experiences more precipitation than the Grassland (Downing and Pettapiece, 2006). While the larger Boreal Natural Region is dominated by coniferous trees and annual precipitation amounts typically exceed evapotranspiration rates (Downing and Pettapiece, 2006), the southern margin of the Boreal in Alberta contains pothole wetlands and more semi-arid to subhumid climate conditions (Brown et al., 2010; Devito et al., 2005).

2.2. Wetland Locations and Extents

For our analysis, we selected a subset of wetlands from the Merged Alberta Wetland Inventories within each Natural Region (Figure 1). To ensure wetland conditions were indicative of the natural regions within which they resided, we excluded those within 500 m of a Natural Region boundary. Then, we randomly selected 12,000 wetlands in the Southern Boreal and Parkland Natural Regions (3,000 per permanence class) and 16,000 in the Grassland (4,000 per permanence class). To ensure spatial independence among sampled wetlands and their relationship to land cover as well as coincide with previous analysis of open water wetlands (Ridge et al. 2017), topography (Branton et al. 2020), and land cover (Evans et al. 2017), we did not select wetlands that were within 1000 m of another selected wetland.

The distribution of wetland sizes was strongly right-skewed across the three Natural Regions of interest. Wetlands were typically small, with Boreal wetlands possessing the largest median size (2.26 ha), followed by Parkland wetlands (1.54 ha) and Grassland wetlands (0.58 ha, Appendix A). The combination of wetland size and our digital elevation model (DEM) resolution of 25 m suggest that our median wetland sizes would occupy 904, 616, and 232 cells for Boreal, Parkland, and Grassland natural regions, respectively, and demonstrate our ability to capture variability among wetland sizes and shape.

2.3. Selecting Variables

To select variables representative of climate, land cover/land use and topography that would be useful in testing the relative contribution of these three factors in predicting prairie pothole wetland permanence class, we conducted a literature review using the Web of Science. We limited the search to papers published between 1950 to 2018 with the following key words: 1) Prairie Pothole Region: PPR, Northern Great Plains, Alberta, Saskatchewan, Manitoba and Dakota; 2) weather: climate, temperature and precipitation; 3) disturbance: land use, agriculture, disturbance, oil and gas, grazing and roads and 4) pond permanence: watershed, hydroperiod, permanence class, catchment and wetland. We used "OR" operators between key words under the same class and "AND" operators between each key word class. To characterize topography, we selected variables that are commonly used to describe topographic variations, based on a previous review (Branton and Robinson, 2019). Details and results from this review are reported in Table 2.

2.3.1. Climate

We acquired 2013-2014 daily weather data from the AgroClimatic Information Service of Alberta to calculate climate variables. These data include precipitation and temperature measurements from 7,914 weather stations across the province, observed from October 2013 to August 2014. We calculated seasonal precipitation totals and temperature averages from a compilation of proxy variables (Table 2) at each station. Then, using a simple inverse distance weighting (Tarroso et al., 2019), we interpolated climate variables at the center of each wetland in R (R Core Team, 2019).

We used annual data on climate variables in this analysis because it was available at a fine spatial resolution and corresponded with the 2014 land cover and topography data we used. Additionally, 2014 was also a typical year in terms of climate variables. For example, we found no significant difference in mean annual precipitation between the 1981-2010 climate normal and the annual data from 2013-2014 (paired t-tests Grassland: $t_9 = 1.833$, p-value $= 0.652$; Parkland: $t_9 = 1.833$, p-value $= 0.344$) or 2014-2015 (paired t-tests Grassland: $t_9 = 1.833$, p-value $= 0.878$; Parkland: $t_9 = 1.833$, p-value $= 0.315$) in either the Grassland or Parkland Natural Regions (cumulative precipitation plots in Appendix B). Importantly, the influence of climate variables on wetland permanence classes will exhibit time lags dependent on site-specific factors, such as soil storage, ground water movement, and vegetation succession within catchments. Consequently, the temporal window of relevant weather would also be site-specific, and we lack a defensible justification on which to base a threshold for including or excluding annual data on climate variables. Coupled with the typical nature of 2014's annual data on climate variables, we elected to use the single year as representative of average conditions in our study area and maximize comparability to our 2014

topography and land use data. We suggest future research could seek to elucidate how legacy effects of climate and land use may influence wetland permanence classes.

2.3.2. Land Cover and Land Use

Prior research in the PPR identified a strong concordance between landcover within 500 m of wetlands and wetland

130   psychochemical conditions (Kraft et al., 2019). Using land cover data from Agriculture and Agri-Food Canada's (AAFC) Annual Crop Inventory for 2014 (Agriculture and Agri-Food Canada, 2014), we calculated the proportion of each land cover class within a 500 m buffer of each wetland (Table 2). In addition to land cover characteristics, we also measured the distance of each wetland centroid to the nearest road using the National Road Network from the Government of Canada (Statistics Canada, 2010). The landscape fragmentation created by road networks has been shown to alter hydrological flow and divert

135   surface runoff (Shaw et al. 2012) such that wetlands in proximity to roads typically have shorter hydroperiods. We estimated these land cover and land use variables in ArcMap 10.4.1 (ESRI, 2012).

2.3.3. Topography

We quantified topographic characteristics of the landscape surrounding each wetland using a 25-m digital elevation model (DEM) for southern and central Alberta (AltaLIS, 2015) (Figure 1). We estimated eight terrain variables (Table 2) using

140   ArcMap 10.4.1 (ESRI, 2012) and SAGA 2.3.2 (Conrad et al., 2015). These variables may be grouped as those with local (e.g., standard deviation of slope) versus global (e.g., terrain surface convexity) application (Branton and Robinson 2019). For local variables, we applied the formula to areas only within 500 m of the wetland boundary. With global variables, we applied a 100 × 100-m moving window and computed the mean value within the 500 m buffers (Table 2).

2.4. Data Analysis

145   We aimed to quantify the relative contribution of annual data on climate variables, land cover/land use and topography for different wetland permanence classes and determine the ability of these drivers to predict wetland permanence class. Achieving these two outcomes involved four steps: reducing the number of variables to an orthogonal and parsimonious set for application; visualizing if wetlands could be partitioned based on their permanence class; parametrizing and calibrating a predictive model; and then predicting permanence class and assessing model fit. These analyses were performed in R (R Core

150   Team, 2019) and while they quantify a relationship among our independent variables with wetland permanence, they do not infer causation.

2.4.1. Predicting Wetland Permanence Class

We used an extreme gradient boosting model to predict wetland permanence class. Extreme gradient boosting is considered a more robust predictive tool than random forest (Sheridan et al., 2016). Like random forest, extreme gradient boosting creates

155   an ensemble of decision trees that partition data based on a specified grouping (Hastie et al., 2009; McCune et al., 2002), which

in our case is wetland permanence class. In the first decision tree, all observations are equally weighted (Cutler et al., 2007). The second decision tree attempts to correct for misclassifications derived from the first tree, assigning a higher weight to observations that were difficult to classify. Each subsequent tree attempts to minimize model error by classifying these error-prone observations (Cutler et al., 2007). The use of the minimum error to build a model ensemble makes extreme gradient boosting models prone to overfitting (Cutler et al., 2007). To correct for overfitting, extreme gradient boosting models include a regularized object that penalizes more complex trees (Chen and Guestrin, 2016).

After parametrizing the model for each Natural region, we predicted wetland permanence class in the 1) Southern Boreal, 2) Parkland and 3) Grassland Natural Regions using a combination of annual data on climate variables, land cover/land use and topography variables (Appendix C). For each model, we also assessed its performance using test data (70:30 training to test ratio) to determine the misclassification error rate, comparing results between training and test data. We also evaluated the relative importance of each variable in predicting permeance class by comparing gain values and assessed under which ranges of each variable a permanence class was more likely to occur with waterfall plots.

3. Results

3.1. Selecting Variables

Before predicting wetland permeance class based on land use and landcover, topography and annual data on climate variables, we first determined which metrics were collinear within their metric class. Based on a maximum allowable correlation Pearson correlation of 0.9, we reduced our 30 metrics to 19. Next, we incorporated these 19 variables into a PCA to explore partitioning of permanence classes in accordance with the annual data on climate, land cover, and topography variables and to facilitate comparison among the three Natural Regions. Wetlands in the Grassland appeared to be better aligned with all three domains than the wetlands in the Boral and Parkland (Figure 2).

3.2 Model Performance

We built an extreme gradient boosting model for each Natural Region (southern Boreal, Parkland and Grassland) in our study area. We selected 19 variables that reflected climate (7), land cover/use (4) and topography (8) (Table 2).
Our models had moderate to high error rates (Appendix D), which suggests that annual data on climate, land use/cover and topography alone are not sufficient variables in predicting permanence class.

3.2. Relative Importance of Variables in Predicting Wetland Permanence Class Among Natural Regions

In each Natural Region, annual data on climate explained the greatest amount of variance in wetland permanence class, based on relative gain values (Figure 3A-C). As anticipated, our results suggest that climate conditions vary systematically among the Natural Regions (Figure 4A-D). Among the climate variables included in our analyses, spring temperature explained the

185 highest magnitude of variance in predicting permanence class in the Southern Boreal and Grassland (Figure 3A;2C) but was less important in the Parkland where values are less extreme (Figure 4A).

Land cover/land use was the second most important category of drivers of wetland permanence class, following annual data on climate in the Grassland Natural Region (Figure 3F), but not in the Southern Boreal or Parkland (Figure 3). Yet, unlike climate, land cover/ land use did not vary systematically among the three Natural Regions (Figure 4E-H). Wetlands surrounded

190 by natural vegetation may have shorter hydroperiods (Figure 4F).

In the southern Boreal and Parkland, topography was the second most important category of drivers of wetland permeance class, and the order of importance for the terrain metrics were nearly the same (Figure 3G-I). Though terrain metrics were the least important category in the Grassland (Figure 3I), apart from deviation from mean elevation (Figure 4I), variables associated with topography did not systematically vary among Natural regions (Figure 4J-L).

195 **3.3. Wetland Permeance Class in the Boreal, Parkland and Grassland**

Our findings suggest that wetland permanence class in the Prairie Pothole Region of Alberta correlates with climate, terrain and, to a lesser extent, to surrounding land cover/ land use. Generally, across the three Natural Regions, wetlands with shorter hydroperiods (e.g., temporary and seasonal) were typically situated in landscapes with higher spring snowpack amounts (Figure 3A-C) and spring temperatures (e.g., Figure 5A). Longer hydroperiod wetlands were typically situated in landscapes

200 with more summer precipitation and lower spring temperatures (e.g., Figure 5C), occupying relatively low topographic positions with low terrain convexity (e.g., Figure 5G, H), and were surrounded by less natural cover (e.g., Figure 5D,F). Interestingly, the relative importance of variables in predicting the occurrence of both shorter and longer-hydroperiod wetlands were shared, and this agreement was strongest between the Southern Boreal and Grassland (Appendices F & H).

4. Discussion

205 Our findings support the assertion of other published studies (e.,g., Fay et al., 2016; Johnson et al., 2010a, 2005; Johnson and Poiani, 2016; Reese and Skagen, 2017; Werner et al., 2013: McKenna et al. 2019), which conclude that climate change will affect wetland hydroperiod or permanence class. We anticipate that reduced winter snowpack will dry out temporarily and seasonally ponded wetlands, while warmer spring temperatures will reduce the hydroperiod of more permanently ponded wetlands. Yet, annual data on climate is not the only element correlated with wetland permanence class in Alberta's PPR - our

210 analysis used a relatively coarse DEM (25 m), and we nonetheless found that topography was important in predicting permanence class. Consequently, failure to consider topography limits our understanding about the extent to which hydroperiod, and therefore wetland permanence class, may change in response to climate change. We speculate that the use of finer-scale elevation models derived from high resolution LiDAR (e.g., 1 m) or remotely piloted aircraft (e.g., 2-5 cm) will reveal the importance of topography in surface runoff and wetland hydroperiod.

4.1. Importance of Climate

The sensitivity of wetland hydroperiods to annual climate data is corroborated in existing literature, which emphasizes that the semi-arid climate drives the region's sensitivity to climate change (Fay et al., 2016; Johnson et al., 2004; Schneider, 2013). In the Southern Boreal and Grassland, regions with warmer spring temperatures are likely to experience an earlier onset of spring snowmelt, higher water deficits (Schneider, 2013; Zhang et al., 2011) and lower pond permanence classes for wetlands, whereas cooler peak spring temperatures favour greater pond permanence in these Natural Regions. In the Parkland, winter snowpack depth was the most important climate variable; and this, we attribute to temporarily and seasonally ponded wetlands requiring a minimum threshold of winter snowpack amount to persist, whereas permanently ponded wetlands also benefit from precipitation in other seasons and so can exist at lower witner snowpack amounts (Figure 5B). Because climate forecasts suggest that warmer springs and changes in precipitation timing are likely (Zhang et al., 2011), our finding that climate was the most important domain of variables in predicting permanence class supports previous studies that suggest PPR wetlands are sensitive climate change (Johnson et al., 2010; Paimazumder et al., 2013; Schneider, 2013; Viglizzo et al., 2015; Zhang et al., 2011).

4.2. Importance of Topography

Despite recognition that topography is a useful proxy in wetland mapping (Branton and Robinson, 2019; Los Huertos and Smith, 2013) and that topography must influence surface runoff generating processes that are essential to wetland function (Hayashi et al., 2016; Mushet et al., 2018), the relative importance of topography in hydrological processes is somewhat debated (Devito et al., 2005). Simulations predicting the influence of climate change on the size and isolation of prairie pothole wetlands have focused on climate and land cover/land use (Anteau et al., 2016; Chasmer et al., 2012; Conly et al., 2001; Johnson and Poiani, 2016; McCauley et al., 2015; Steen et al., 2016; Voldseth et al., 2007). Consequently, 1) there is a lack of research quantifying topographic characteristics of wetlands and the landscapes within which they occur, 2) links between topography, vegetation and wetland condition have not been rigorously studied and 3) policy and guidelines on wetland mitigation and compensation prescribe width-to-length ratios and slopes that are characteristic of permanently-ponded wetlands, which are rarer in the Grassland (Environmental Partnerships and Education Branch Alberta, 2007). Thus, the classified frequency distribution of wetland permanence classes across Alberta's Prairie Pothole Region has skewed toward permanently ponded wetlands (Serran et al., 2017). If we had a better understanding of how topographic structure determines wetland hydrology/function, we could revise policy and regulations governing wetland management to ensure we better match natural landscapes in their frequency and distribution of wetland permanence classes.

4.3. Importance of Land Cover/Land Use

Existing literature identified land cover/land use as the second greatest driver of wetland conditions following climate (Anteau et al. 2016). In the Grassland, the terrain is relatively flat compared to the southern Boreal and Parkland (Alberta Tourism

Parks and Recreation, 2015). Consequently, after the important role of annual data on climate in the more arid Grassland Natural Region (Government et al., 2014), land cover/land use would be a stronger driver of permanence class than topography. Importantly, the percent cover of natural vegetation is typically low in the Grassland, where most land has been converted to cropland or pastureland (Alberta Tourism Parks and Recreation, 2015). Combined with the process of consolidation drainage, which shunts water from scattered low hydroperiod wetlands, concentrating it in larger more permanently ponded wetlands downstream (McCauley et al., 2015), this leads to Grassland landscapes with more natural cover being more likely to contain temporary and seasonal wetlands. Thus, wetlands surrounded by natural vegetation may have shorter hydroperiods because cropland resists infiltration and natural vegetation intercepts snow-sourced surface runoff (Anteau, 2012; van der Kamp et al., 2003; Voldseth et al., 2007), which can account for up 27% of ponded water amounts (van der Kamp et al., 2003).

Because some landscapes in the PPR are flatter than others (Schneider, 2013), and land use activities can modify the terrain (Anteau 2012; Wiltermuth and Anteau 2016; Anteau et al. 2016), our findings do highlight the importance of considering land use in forecasting the impacts of climate change on PPR wetlands. Boreal and Parkland wetlands have stronger overlaps in terrain metrics and annual data on climate; and, as a result, differences in land use within these regions may be integral in determining future shifts in the frequency distribution of permanence classes. Forecasts for the province of Alberta suggest there will be expansions in the agricultural industry within the next decade (Government of Alberta, 2015), and this suggests that climate impacts on Albertan PPR wetlands will be compounded by land use activities.

4.4. Topographic Position of Wetlands by Permanence Class

Semi-permanent and permanently-ponded wetlands typically occur in regional or spatial neighourhood topographic lows (as opposed to simply local depressions, e.g., perched wetlands), likely because they 1) can hold larger volumes of ponded water (i.e., larger pond size/volume (Novikmec et al., 2016)) and 2) receive higher volumes of water inputs from the surrounding landscape (e.g., surface run-off, groundwater Euliss et al., 2004, 2014; LaBaugh et al., 1998; Toth, 1963)). We are unable to partition the natural hydrogeological effects of topographic position on wetland permanence class from the effects of human alteration of the surrounding landscape, yet the importance of topographic position to wetland permanence class is likely reinforced by consolidation drainage:. when wetlands situated higher in the landscape are drained and the water is redirected to wetlands positioned lower in the landscape (McCauley et al., 2015; Wiltermuth and Anteau, 2016). Because of consolidation drainage, we may observe increases in hydroperiod of wetlands in topographic lows of the wetlandscape (e.g., sites with low topographic position index values. In the arid but heavily farmed Grassland Natural Region, consolidation drainage can eliminate temporary and seasonally ponded wetlands from areas with limited remaining natural cover (Serran et al., 2017). This aligns with our model results, as though the probability of observing a permanent or semi-permanent class wetland was greatest at the lower end of the range of crop cover in our landscapes, the threshold of crop cover above which wetlands were most probably seasonal or temporary in class was higher in the Grassland, lower in the Parkland, and lowest in the Boreal. Thus, we recommend that future work investigate the role of topographic position on permanence class, in the absence of human disturbance to control for the influence of consolidation drainage.

4.5. Model Error

280 We hypothesize that our inability to account for soil characteristics (Schneider, 2013) and bathymetry (Huertos and Smith 2013) could explain these high error rates (Appendix D). Schneider (2013) stated that within Natural Regions, both elevation (which we did account for), and soil characteristics can vary across the landscape. As such, wetlands situated similarly in the landscape may not have the same soil characteristics, and soil characteristics are understood to influence wetland hydrology by dictating the proportion of incident precipitation that is converted to surface run of (Hayashi et al., 2016). Though Schneider

285 (2013) also mentioned an influence of disturbance history on ecosystems, prior work in our study region reported no temporal lag in wetland environmental conditions and surrounding land cover (Kraft et al., 2019).

The lack of extensive bathymetric data identifies a gap that would enrich the presented research by enabling direct classification of wetland permanence from raw bathymetric data. Furthermore, these data would provide added value to those conducting research on above and below ground hydrologic connectivity and contributing areas (e.g., citations), evaluation of

290 the impacts of climate change on wetland permanence and subsequently flora and fauna health and resilience (e.g., citations), as well as reduce the error in our analysis of the contributions of climate, land use, and topography on wetland permanence. As new technologies for mapping wetland bathymetry become more widely available (e.g., bathymetric LiDAR Paine et al., 2015; Wang and Philpot, 2007) an opportunity will exist to better understand the link between wetland pattern and process.

It is also likely that some proportion of model error can be attributed to the use a single year of climate and land use

295 data as well as our relatively course (25 m) digital elevation model. However, it is likely that the contributions of these factors are minimal given that 1) the climate data used (year 2014) is representative of average conditions, coincides with fieldwork, and yielded the strongest among the variables interrogated and therefore improving the quality of its contribution will not change the qualitative outcome of the presented analysis; and 2) previous research found that physiochemical conditions in a wetland are quite congruent with surrounding land cover of the same year with only minor differences when catchments were

300 defined with 10 m versus 25 m resolution DEMs (Kraft et al. 2019).

5. Conclusion

Because some landscapes in the PPR are flatter than others (Schneider, 2013), and land use activities can modify topography (Anteau, 2012; Anteau et al., 2016; Wiltermuth and Anteau, 2016), our findings also highlight the importance of considering land use in forecasting the impacts of climate change on PPR wetlands. Southern Boreal and Parkland wetlands are most

305 congruent in the relative importance of climate and topography variables; and, as a result, differences in land use within these regions may be integral in determining future shifts in the frequency distribution of permanence classes. Forecasts for the province of Alberta suggest expansion in the agricultural industry over the next decade (Government of Alberta, 2015), which suggests that climate impacts on Alberta's PPR wetlands will be compounded by changes in land use activities.

6. Code/Data availability

310 On acceptance of the manuscript, the code and data will be uploaded to FigShare for archiving and a DOI provided.

7. Author Contributions

RCR conceptualized the study, acquired funding and resources, supervised and curates the data; DR and JD gathered the data; JD analysed and visualized the data and wrote the original draft; all authors contributed to investigation and review and editing.

8. Competing Interests

315 The authors declare that there are no competing interests or conflicts of interest.

Table 1. Descriptions of the four permanence classes included in our study. We describe the typical length of time that these prairie pothole wetlands will contain ponded water, their associated vegetation zones, as described by Stewart and Kantrud (Stewart and Kantrud, 1971), and the number of wetlands belonging to each class in the Alberta Merged Wetland Inventory 320 (Government of Alberta, 2014) that were within the extent of our 25-m Digital Eleveation Model.

Permanence class	Typical hydroperiod	Vegetation zones	Natural Region		
			Boreal	Parkland	Grassland
Temporary	Until mid-spring, typically for four weeks	Wet-meadow (includes wet-meadow emergent), low-prairie, high seepage	40461	51062	153872
Seasonal	Late spring to early summer for approximately two months	Shallow-marsh (vegetation zones from shallow to deep: emergent plants, submerged aquatic plants), wet-meadow, low prairie	30890	43836	108924
Semi-permanent	Dries fully in drought years only	Deep-marsh (vegetation zones from shallow to deep: emergent vegetation, open-water with bare-soil), shallow-marsh, wet-meadow, low-prairie	39375	47075	12240
Permanent	Open water year-round	Open water, deep marsh, shallow-marsh, wet-meadow, low-prairie	5704	10785	4952

Table 2. List of annual data on climate, land cover and land use, and terrain metrics used to predict wetland permanence class. In this table, we include a description of the significance of each metric for wetland hydroperiod and the proxy metrics we selected. For our analysis, winter months range from November to February, spring April to May and summer June to August. We used Web of Science to conduct this review, limiting the search to papers published between 1950 to 2018, and key words for: 1) the PPR: Prairie Pothole Region, Northern Great Plains, Alberta, Saskatchewan, Manitoba and Dakota; 2) weather: climate, temperature and precipitation; 3) disturbance: land use, agriculture, disturbance, oil and gas, grazing and roads and 4) pond permanence: watershed, hydroperiod, permanence class, catchment and wetland. We used "OR" operators between key words under the same class and "AND" operators between each key word class. For the terrain metrics, we used selected metrics that are commonly used to describe topographic variations, based on a previous review (Branton and Robinson, 2019). Notably, Branton and Robinson (2019) employed controls on collinearity, including PCA.

Category	Variable	Significance for Wetland Hydroperiod/Formula	Proxy/Class[1]
Climate	Snowpack/Winter precipitation	Snowpack accounts for 30-60% of ponded water amounts(Hayashi et al., 1998; Tangen and Finocchiaro, 2017) Longer hydroperiods with higher winter precipitation (Collins et al., 2014)	Total Spring Precipitation Total Winter Precipitation Total Precipitation in Winter & Spring Total Spring Snowpack Total Winter Snowpack Total Snowpack in Winter & Spring
	Sumer Precipitation	Longer hydroperiods from increased summer precipitation(Clare and Creed, 2014; Eisenlohr, 1972; Euliss et al., 2014; Leibowitz and Vining, 2003)	Total Summer Precipitation
	Summer Temperature	Evapotranspiration rates/water losses higher in summer (from June)(Heagle et al., 2007)	Average Maximum Temperature in June Average Maximum Temperature in July Average Maximum Temperature in Summer
	Winter/Spring/Summer Temperature	Snowpack may melt too fast with warmer conditions(Crosbie et al., 2013),13]	Average Maximum Temperature in Spring Average Maximum Temperature in Winter Average Maximum Temperature in Spring & Winter

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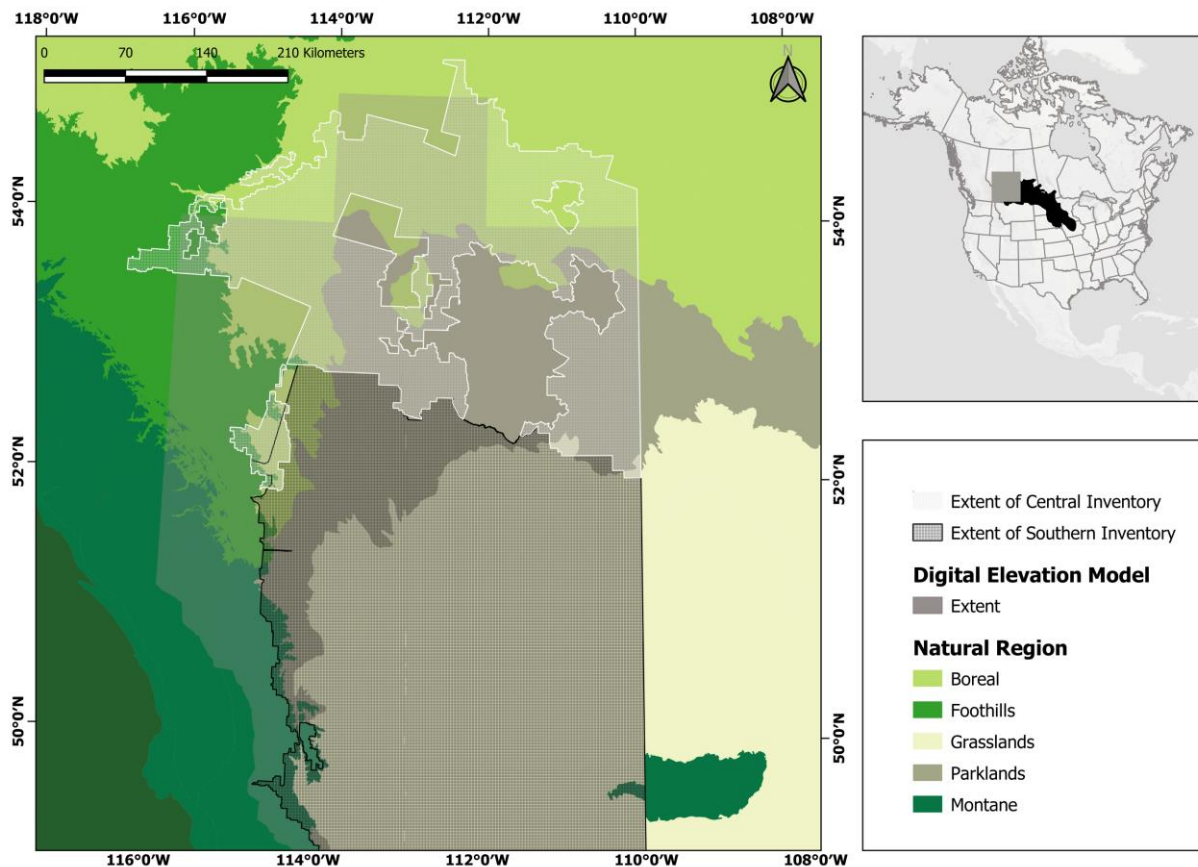
[1] This differentiates terrain metrics by global (estimated using a 100×100-m moving window and mean value within 500-m buffer recorded) and local (estimated within a 500-m buffer of the wetland.

Category	Variable	Significance for Wetland Hydroperiod/Formula	Proxy/Class[1]
	Precipitation Timing	Fewer wetlands dry up when summer precipitation is earlier in the summer(Meyers, 2018; Vinet and Zhedanov, 2011)	Proportion of Summer Precipitation in June
Land Use & Land Cover	Natural Vegetation	Loss of natural cover increases surface runoff (Clare and Creed, 2014)	% Natural Cover
	Cropland Cover	Because soil is less porous (more compacted), much of the accumulated water, either from the snowpack or spring/summer precipitation, flows into the wetland – this increases water levels (van der Kamp et al., 2003; Voldseth et al., 2007)	% Cropland Cover
	Urban Cover	Longer hydroperiods in urban landscapes, mostly because of higher runoff (when compared to those in croplands)(Fossey and Rousseau, 2016)	% Urban Cover & Bare Ground
	Grazing	Grazing lowers snow accumulation (Willms and Chanasyk, 2013), which can increase runoff and hydroperiod(Collins et al., 2014; Niemuth et al., 2010)	% Pastureland
	Culverts/Roads	Lowers hydroperiods by blocking surface runoff(Shaw et al., 2012)	Distance to Road
	Tilling	Can lower pond area/depth, and by extent hydroperiod, as increases in sedimentation can in fill ponds(Skagen et al., 2016)	% Cropland Cover
Terrain Metrics	Mean Elevation (DEM) - Deviation	$$\|Elevation - Elevation_{mean}\|^2$$	Local
	Elevation (DEM) - Standard Deviation	$$\sqrt{\frac{\sum(Elevation - Elevation_{mean})}{n}}$$	Local
	Profile Curvature (PC) - Standard Deviation	$$\sqrt{\frac{\sum(Profile\ Curvature - Profile\ Curvature_{mean})}{n}}$$	Local
	Slope - Standard Deviation	$$\sqrt{\frac{\sum(Slope - Slope_{mean})}{n}}$$	Local
	Terrain Surface Convexity	Percentage of upwardly-convex cells within the moving window(Iwahashi and Pike, 2007).	Global
	Terrain Surface Texture	Relative frequency of pits and peaks in a 100×100-m moving window(Iwahashi and Pike, 2007)	Global
	Topographic Position Index	$$\frac{Elevation_{mean} - Elevation_{min}}{Elevation_{max} - Elevation_{min}}$$	Local
	Slope Variability	$$Slope_{max} - Slope_{min}$$	Local

335 **Figure 1:** Extents of the Central and Southern Wetland Inventories (Government of Alberta, 2014) used to delineate wetlands in our study. We selected wetlands from three Natural Regions – Boreal (12,0000), Parkland (12,0000) and Grassland (16,0000); Natural Region boundaries are sourced from the Government of Alberta (Government Alberta, 2016). These wetlands are within the southern-Albertan Prairie Pothole Region. There are 356,246 wetlands delineated in the Southern Inventory and 253,873 in the Central Inventory. DEM data provided by AltaLIS, (2015).

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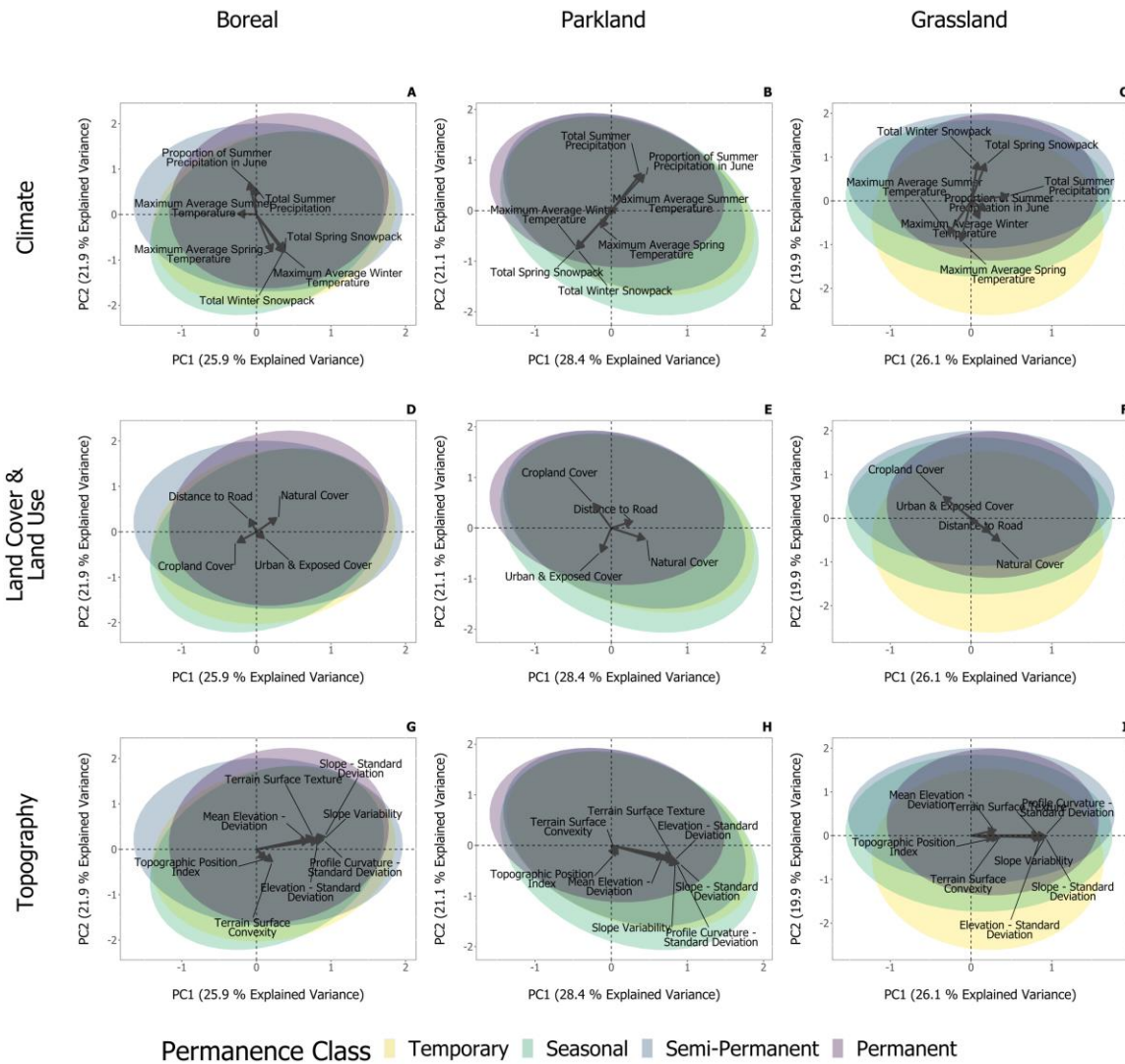
Figure 2. Principal Components Analysis for wetlands delineated in the 1) Boreal (totalling 12,000 wetlands), 2) Parkland (totalling 12,000 wetlands) and 3) Grassland (totalling 16,000 wetlands) Natural Regions. PCAs apply an orthogonal transformation to summarize the data into axes that explain the variance between two correlation matrices. Our data were scaled before implementing the PCA. Vectors on climate (A-C), land use and land cover (D-F) and terrain roughness (G-I) show correlations with both axes. Axis two, for all datasets, represents a hydroperiod gradient and terrain roughness is represented on axis 1.
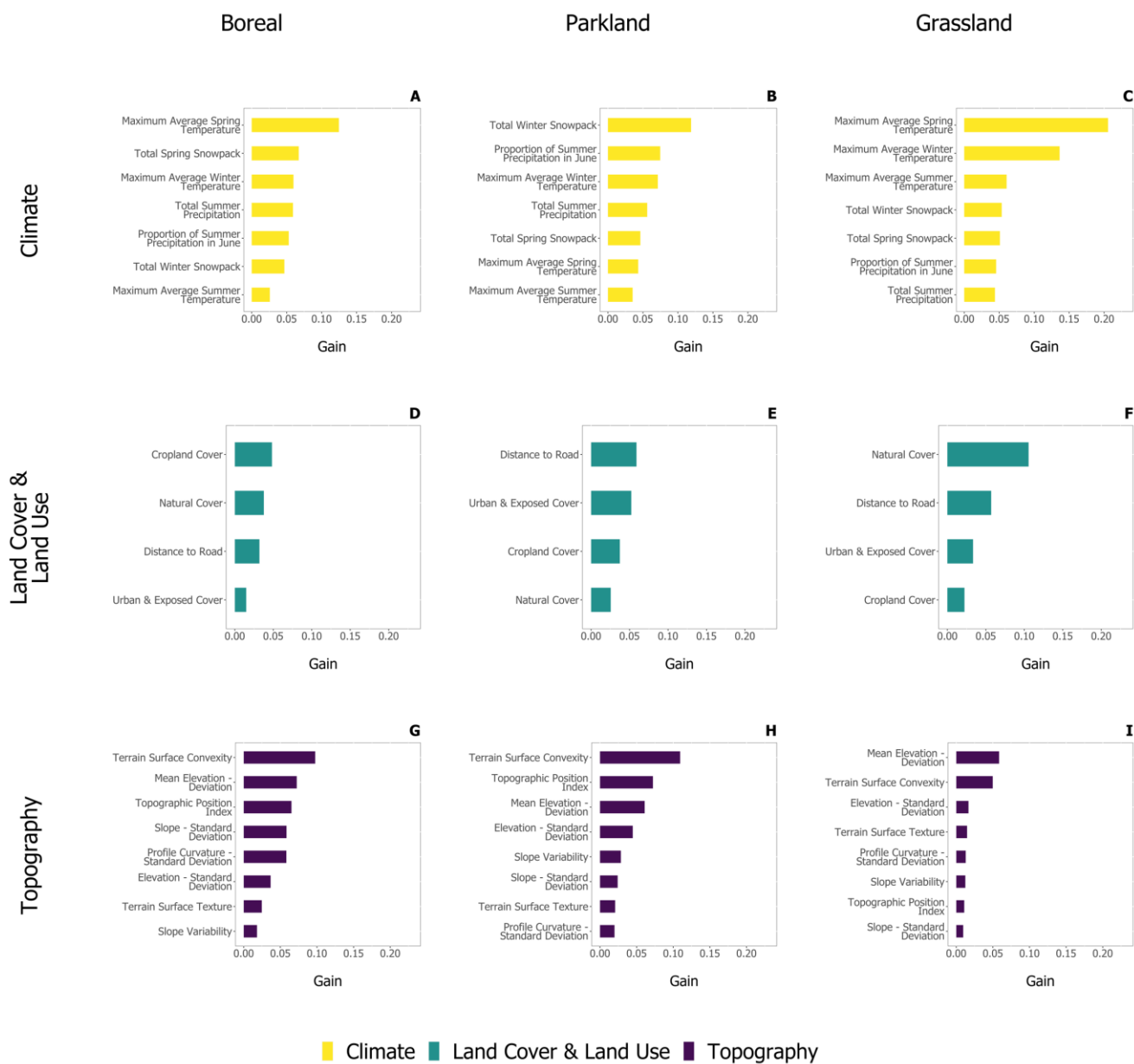
Figure 3. Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the 1) Boreal (totalling 12,000 wetlands), 2) Parkland (totalling 12,000 wetlands) and 3) Grassland (totalling 16,000 wetlands) Natural Regions. These variables were proxies for climate (A-C), land cover and land use (D-F) and topography (G-I). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

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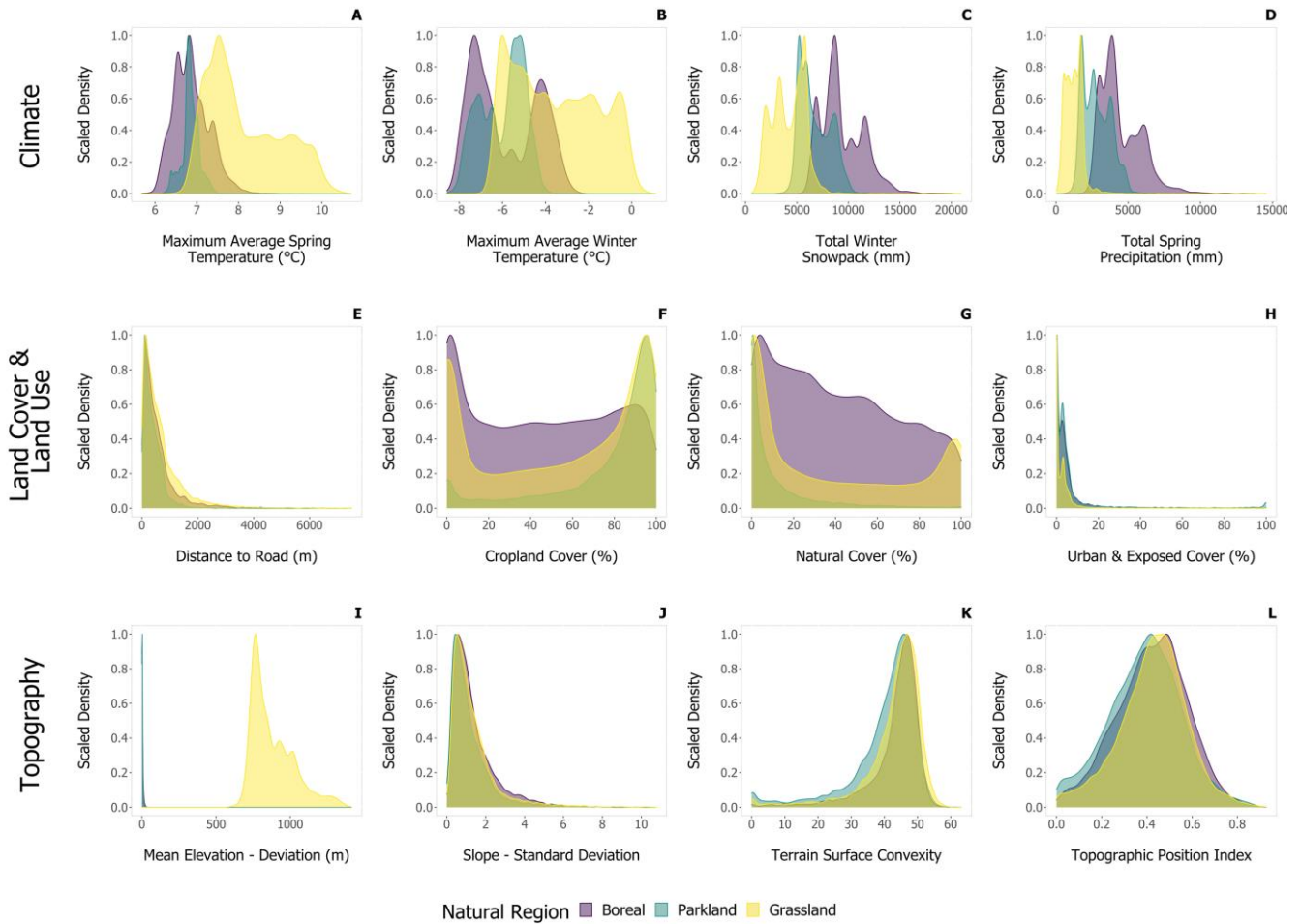
Figure 4. Frequency distribution of the top four climate, land cover and land use and topography variables by Natural Region.
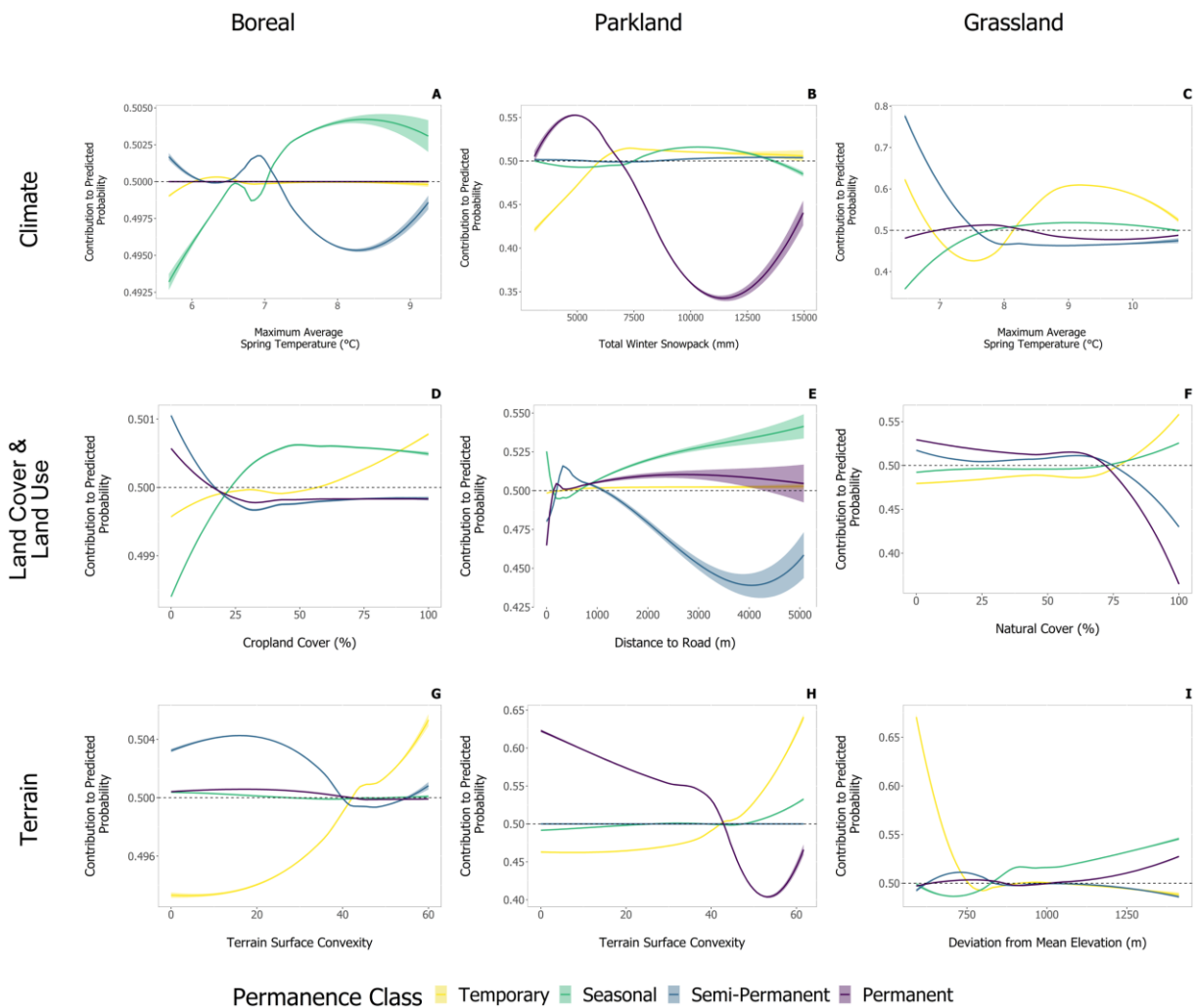
Figure 5. Partial dependence plots for the four wetland classes – temporary, seasonal, semi-permanent and permanent based on top metrics. Predicted probabilities below 0.5 suggest that at this measured value of the metric, observing that permeance class is unlikely. We show 95 % confidence intervals and used a generalized additive model-based trend line. Probabilities were derived from extreme gradient boosting models for wetlands delineated in the 1) Boreal (totalling 12,000 wetlands), 2) Parkland (totalling 12,000 wetlands) and 3) Grassland (totalling 16,000 wetlands) Natural Regions.

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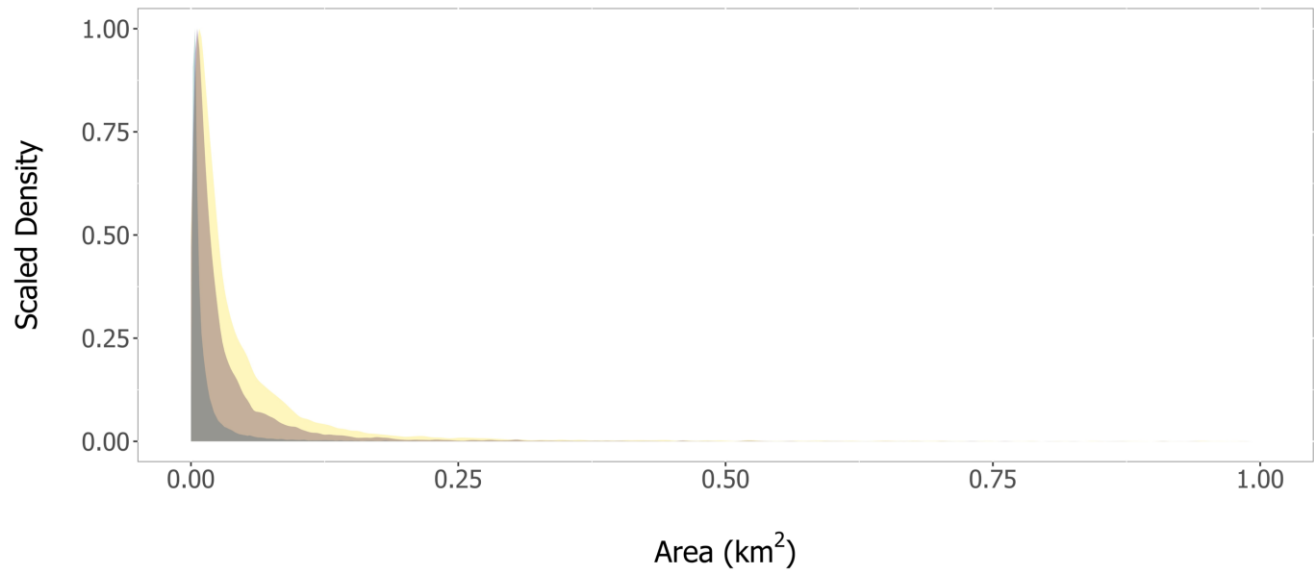
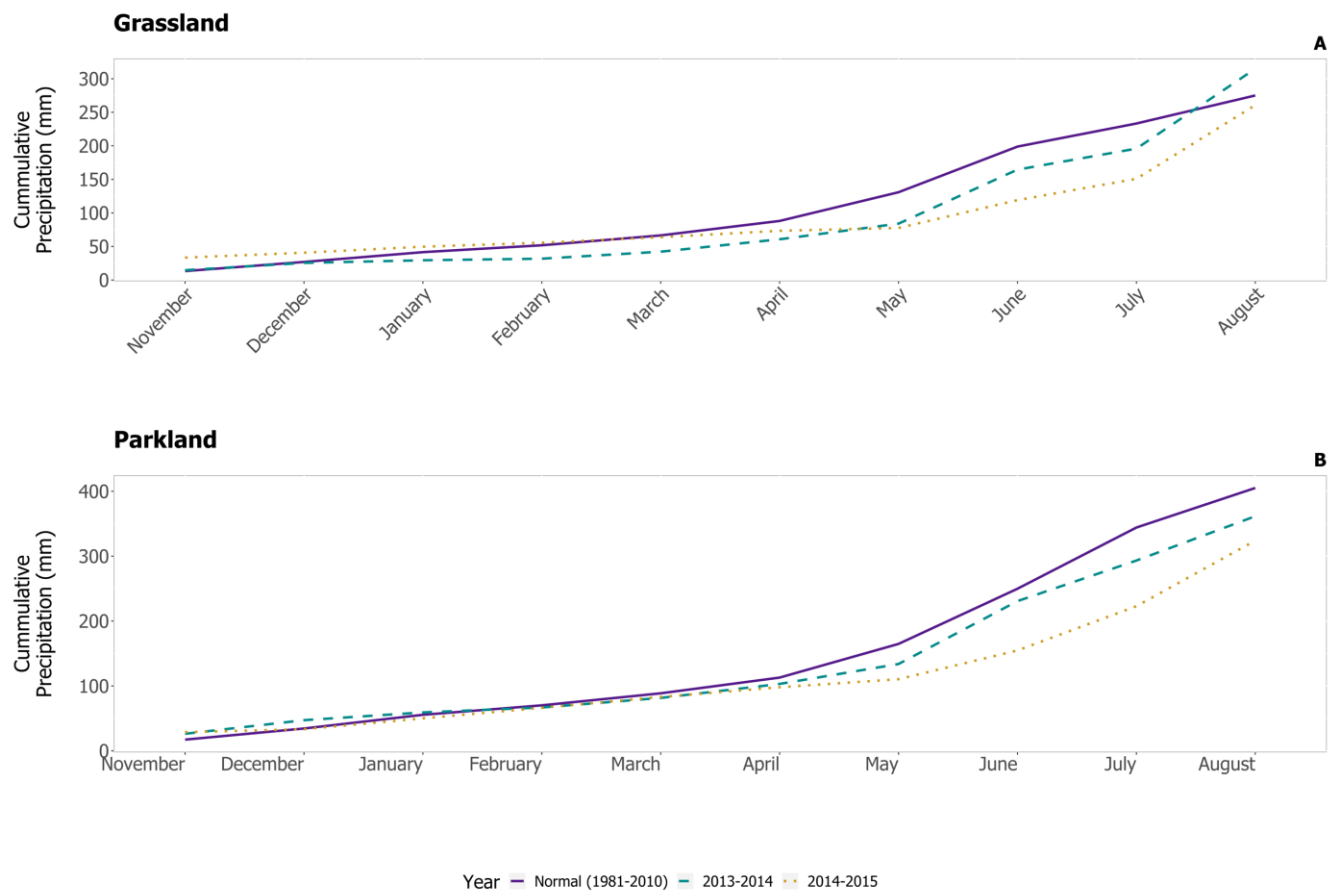
Appendix A. Frequency distribution of wetland sizes in in the Boreal, Grassland and Parkland Natural Regions. Data on wetland sizes were acquired from the Alberta Merged Inventory (Government of Alberta, 2014).

Natural Region	Median	Mean	Standard Deviation
Boreal	0.0226	0.08008	0.95645
Grassland	0.0058	0.02209	0.23276
Parkland	0.0154	0.04341	0.30779

370

Appendix B. Comparison of cumulative precipitation in the Grassland and Parkland Natural Region between 2014-2015 to the climate normal.

Appendix C. List of parameters tuned for the extreme gradient boosting model, a descript of these parameters, their ranges
375 and the ranges evaluated in our cross validation.

Parameter	Description	Range
Learning rate	Used to control the contribution of each tree to model. Lower values result in the model being more robust to overfitting.	Typical: 0-1 Model: 0-0.3 Boreal (0.01), Parkland (0.1); Grassland (0.05)
Gamma	This controls the complexity of the model. It determines how much loss (difference between prediction and observation) allowable for the formation of a new node.	Typical: 0-20 Model: 0-10 Boreal (8), Parkland (4); Grassland (10)
Maximum depth of a tree	This sets the maximum number of nodes that can exists between the tree root and leaves. The larger the value, the more likely a tree is to overfit.	Typical: 1-7 Model: 1-7 Boreal (5), Parkland (7); Grassland (7)
Minimum sum of instance weight needed in a child	This sets a minimum weight/purity of data (e.g., number belonging to a given group) for spiting to create a new node in a tree. The higher this number is, the more conservative the algorithm will be.	Typical: 1-7 Model: 1-7 Boreal (5), Parkland (3); Grassland (7)
Subsample ratio of the training instance	This sets the number of rows (fractional) that should be included in building a tree.	Typical: 0-1 Model: 0.6-1 Boreal (0.8), Parkland (0.65); Grassland (0.7)
Subsample ratio of columns when constructing each tree	This sets the number of predictors (fractional) that should be considered in each tree.	Typical: 0-1 Model: 0.6-1 Boreal (0.8), Parkland (1); Grassland (0.9)

Appendix D. Value of parameters used in extreme gradient boosting models for our three datasets, the misclassification error rates and number of trees for our models.

Parameter	Natural Region		
	Boreal	**Parkland**	**Grassland**
Learning rate	0.01	0.1	0.05
Gamma	4	6	8
Maximum depth of a tree	5	5	7
Minimum sum of instance weight needed in a child	5	5	7
Subsample ratio of the training instance	0.8	0.90	0.70
Subsample ratio of columns when constructing each tree	0.8	1.0	0.90
Misclassification error rate	49.6 (training)	52.6 (training)	45.3 (training)
	56.3 (test)	59.7 (test)	50.1 (test)
Number of trees	37	52	46

380 **Appendix E.** Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the Boreal (totalling 12,000 wetlands) by permanence class. These variables were proxies for climate (A-D), land cover and land use (E-H) and topography roughness (I-L). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

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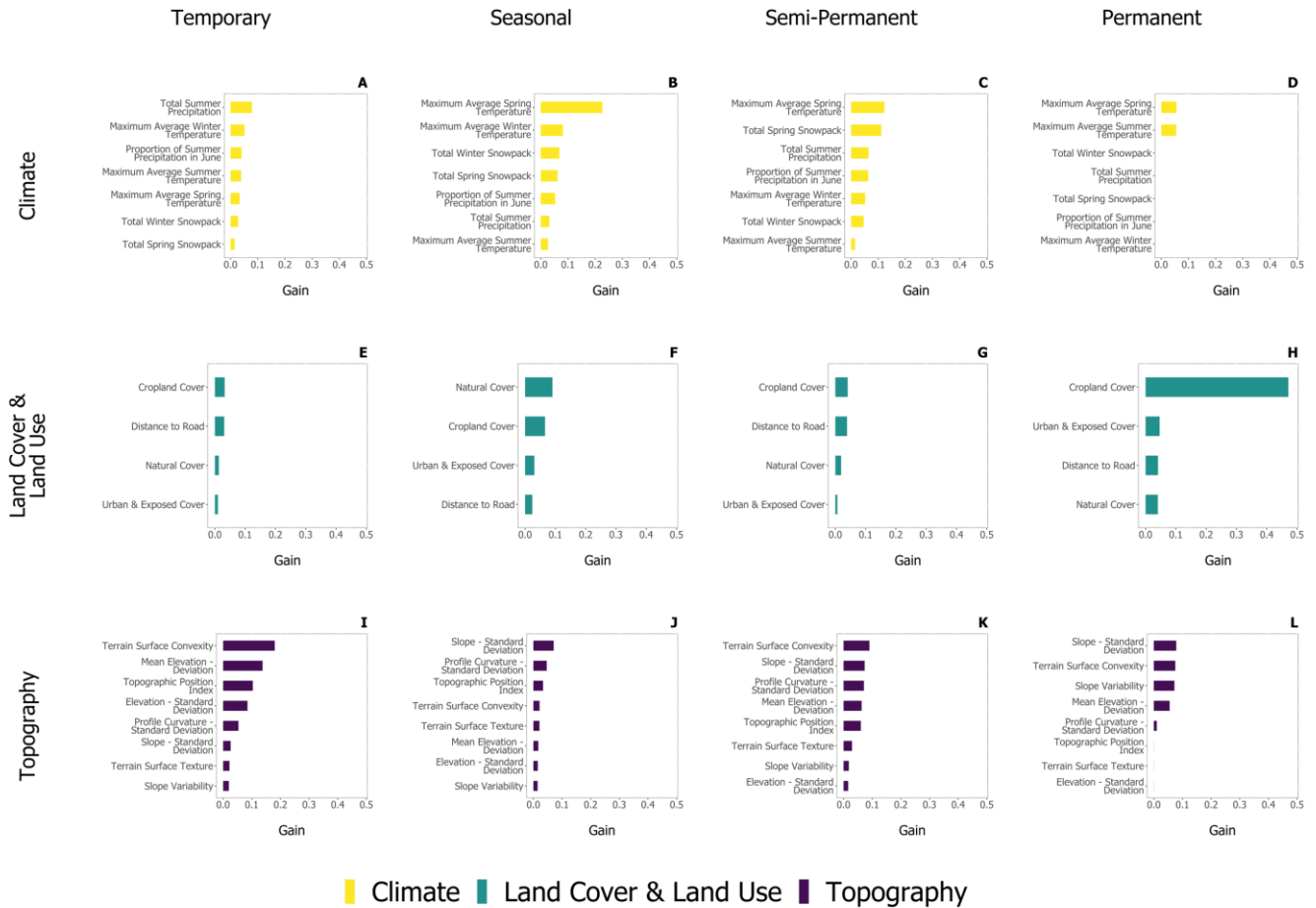
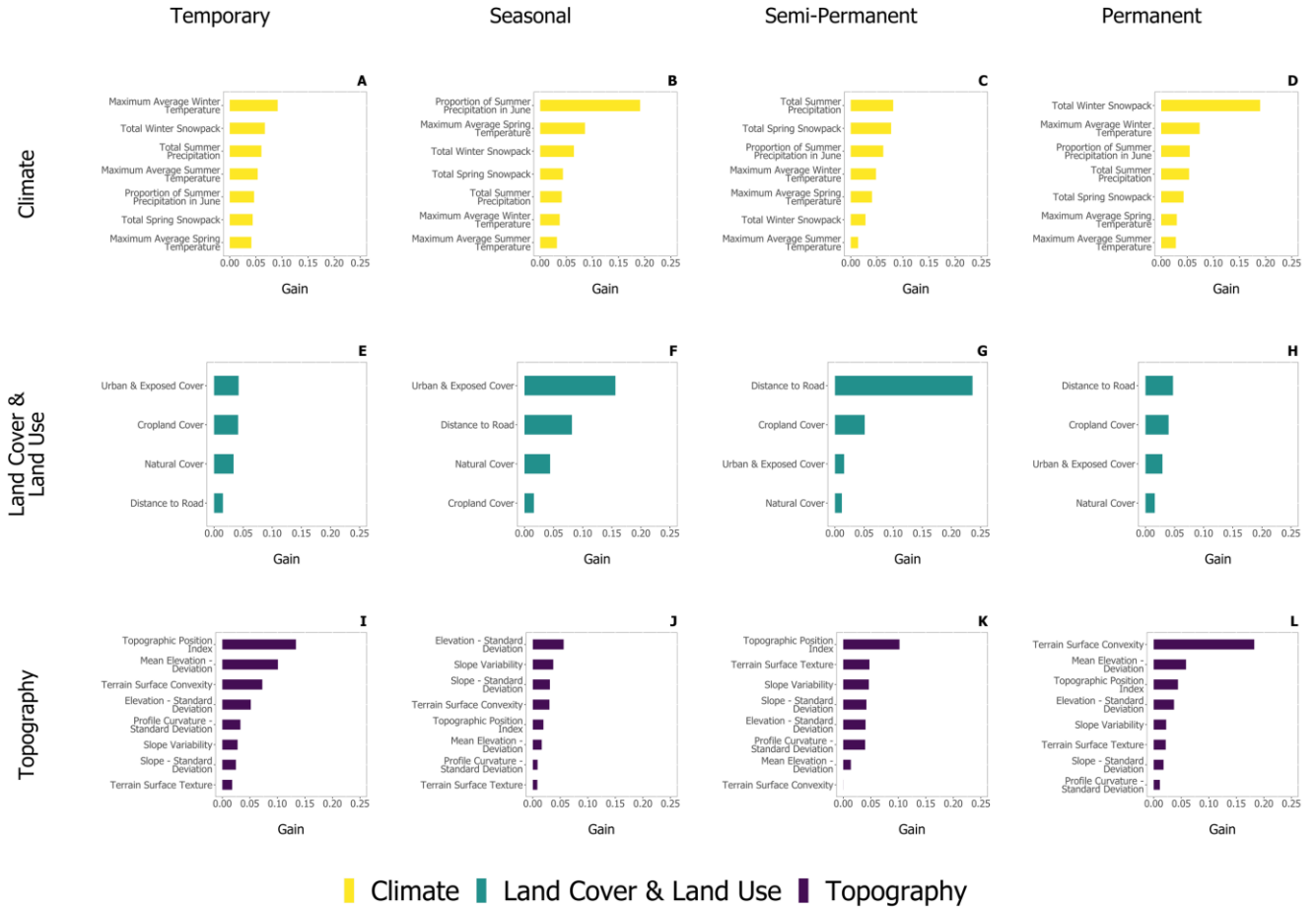

Appendix F. Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the Parkland (totalling 12,000 wetlands) by permanence class. These variables were proxies for climate (A-D), land cover and land use (E-H) and topography roughness (I-L). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

390

Appendix G. Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the Grassland (totalling 12,000 wetlands) by permanence class. These variables were proxies for climate (A-D), land cover and land use (E-H) and topography roughness (I-L). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

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