

Climate, land cover and topography: essential ingredients in predicting wetland permanence.

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Abstract. Wetlands in the Prairie Pothole Region (PPR) are forecast to retract in their ranges due to climate change and potholes that typically contain ponded water year-round, which support a larger proportion of biological communities, are most sensitive to climate change. In addition to climate, land use activities and topography also influence ponded water amounts in PPR wetlands. However, topography is not typically included in models forecasting the impacts of climate change on PPR wetlands. Using a combination of variables representing climate, land cover/land use, and topography, we predicted wetland permanence class in the southern Boreal, Parkland and Grassland of the Alberta PPR (N = 40,000 wetlands). We show that while climate and land cover/land use were strong predictors of wetland permanence class, topography was as important, especially in the southern Boreal and Parkland Natural Regions. Our misclassification error rates for the gradient boosting models for each Natural Region were relatively high (43-60) though our learning rates were low (< 0.1) and our maximum tree depths shallow (5-7) to balance bias and overfitting. Clearly, factors in addition to climate, topography, and land cover/land use influence wetland permanence class (e.g., basin size, depth, ground water connectivity, etc.). Despite classification errors, our results indicate that climate was the strongest predictor of wetland permanence class in the Parkland and Grassland Natural Regions, whereas topography was most important in the southern Boreal among the three domains we considered.

1. Introduction

Wetlands provide habitat for diverse communities of flora and fauna (Gibbs, 1993; Loesch et al., 2012; Sundberg et al., 2016) and deliver ecosystem services of disproportionate importance relative to the area they occupy (Mitsch and Gosselink, 2015). The diversity and abundance of flora and fauna in these wetland ecosystems (Daniel et al., 2019; Gleason and Rooney, 2018) is a function of the consistency with which ponded water is available (i.e., pond permanence), which is forecast to decline in amount and duration of presence (i.e., hydroperiod) across the prairie pothole region of North America due to climate change (Euliss et al., 2004; Fay et al., 2016; Steen et al., 2014, 2016). In this region, most wetlands are ponded non-permanently and they support resident biological communities (Daniel et al., 2019; Stewart and Kantrud, 1971) that are sensitive to climate

30 change (Fay et al., 2016; Johnson et al., 2010). Therefore, understanding the relative influence of climate on wetland water levels is critical to improving our understanding of how biological communities in the Prairie Pothole Region (PPR) will respond to climate change.

Alberta lies at the western edge of the Prairie Pothole Region, which encompasses the province's Grassland and Parkland Natural Regions, as well as the southern edge of the Boreal (Schneider, 2013). Given the PPR's semi-arid climate, a
35 decline in wetland hydroperiod is expected because of increases in wetland water deficits (Schneider, 2013; Werner et al., 2013). Simulations for the PPR suggest that the magnitude of change in climatic conditions between 1946 and 2005 were vast enough to drive declines in pond permanence (Werner et al., 2013). Modelling suggests that these wetlands may experience up to a 20% decline in precipitation due to climate change, which could reduce hydroperiods (Fay et al., 2016). Furthermore, forecasts suggest that many of the wetlands in the southern and western PPR may lose their ponded water completely, driven
40 by drier climate conditions in these areas (Johnson et al., 2005, 2010; Reese and Skagen, 2017). Wetlands that contain ponded water year-round will be most sensitive to climate change because they contain water in late summer when they will be subjected to greater evapotranspiration-driven losses (Fay et al. 2016). They are also relatively rare (Ridge et al., 2021). In addition to climate, topography can also affect hydroperiods in PPR wetlands (Johnson et al., 2010; McCauley et al., 2015; Tsai et al., 2012). The potholes, in which these wetlands are located, form a relic of the land's glaciated history and larger
45 catchments contribute more water, resulting in larger water budgets and longer hydroperiods for some pothole wetlands than others (Hayashi et al., 2016; Shaw et al., 2013). Contemporary land-use practices (e.g., filling and ditching) can alter natural topography, affecting flows of surface and groundwater and subsequently wetland hydroperiod. This phenomenon, referred to as consolidation drainage, fully or partially drains upper-watershed wetlands and directs their water to areas lower in the watershed (McCauley et al., 2015). Consolidation drainage is typically done to lower the probability that neighbouring
50 croplands will flood (Schindler and Donahue, 2006; Verhoeven and Setter, 2010), which increases farming efficiency (Wiltermuth and Anteau, 2016).

Changes in land use can influence wetland hydroperiods by more than associated terrain modification. For example, landscapes with a higher proportion of agricultural activities can have longer hydroperiods due to the combination of increased surface run-off and decreased soil infiltration (van der Kamp et al., 2003; Voldseth et al., 2007). Many studies assessing the
55 impacts of climate change on PPR wetlands incorporate land use (Anteau et al., 2016; Vodseth et al., 2009) and there is resounding evidence that wetlands exposed to the same climate regime, but situated among different land-use activities, differ in their sensitivity to climate change (McCauley et al., 2015; Wiltermuth and Anteau, 2016).

While topography is an important predictor of pond permanence (Hayashi et al., 2016; Neff and Rosenberry, 2017; Shaw et al., 2013; Wiltermuth and Anteau, 2016), it is rarely included in studies assessing the impacts of climate change on
60 PPR wetlands and/or biota (Wolfe et al., 2019). Even well-established models (e.g., WETSIM (Poiani and Johnson, 1993), WETLANDSCAPE (Johnson et al., 2010)), applied to the PPR, predict pond permanence in response to climate, but omit topography. However, differences in topography may cause wetlands belonging to the same permanence class to differ in their

sensitivity to climate change. Consequently, our failure to incorporate topography when predicting pond permanence leaves us with an incomplete understanding on how wetland biota are affected by climate change.

65 Incorporating the influence of topography individually and in combination with climate and land cover/land use effects on wetland permanence is a gap we must fill to improve wetland and waterfowl population management across the PPR (Fay et al., 2016). We analyse data collected across multiple field projects and use spatial data, comprising thousands of wetlands across the PPR in Alberta, Canada. Only four permanence classes (of seven) are represented in this study (Table 1), and which permanence class a wetland belongs to is determined by the vegetation zone in the deepest part of the wetland –
70 and this is dictated by its typical hydroperiod/pond permanence over several years (Stewart and Kantrud, 1971). Using these data, we quantify the relative contribution of climate, landcover/land use and topography in predicting different wetland permanence classes of marshes in Alberta’s PPR.

2. Methods

2.1. Study Area

75 The wetlands in our study are in the Albertan extent of the Prairie Pothole Region (PPR) (Figure 1). Wetlands in this region are mainly depressions filled with ponded water, each formed in the last glacial period (Wright, 1972). Spring snow melt is the largest contributor to ponded water amounts, either from direct precipitation into the wetland or as runoff over frozen ground from upland areas (Hayashi et al., 1998). Potholes can differ in the length of time they contain ponded water, which can range from a few weeks after snowmelt to the entire year (Stewart and Kantrud, 1971).

80 The provincial Merged Wetland Inventory (Alberta Merged Wetland Inventory; published by Alberta Environment and Parks) and the Canadian National Wetland Inventory (Canadian Wetlands; published by Environment and Climate Change Canada) do not assign permanence classes or provide measurements (e.g., water volume, depth) that could be used to classify the wetlands in our study region by permanence class. We acquired permanence class data from two smaller wetland inventories (Government of Alberta, 2014) that delineate the location, boundary and permanence class of PPR wetlands based
85 on Stewart and Kantrud’s classification (Stewart and Kantrud, 1971) (Table 1). The two wetland inventories differ in their accuracy (Evans et al., 2017) and include wetlands from the Grassland, Parkland and the southern edge of the Boreal Natural Regions of Alberta. The Grassland comprises mixed-grass prairie, and the Parkland comprises deciduous trees and grasses. Both are semi-arid regions with potential evapotranspiration rates that are greater than annual precipitation (Downing and Pettapiece, 2006). The Parkland, however, experiences more precipitation than the Grassland (Downing and Pettapiece, 2006).
90 While most of the Boreal Natural Region is dominated by coniferous trees and annual precipitation amounts typically exceed evapotranspiration rates (Downing and Pettapiece, 2006), the southern margin of the Boreal in Alberta contains pothole wetlands and more semi-arid to subhumid climate conditions (Brown et al., 2010; Devito et al., 2005). Our study of the Boreal considers only on this southern margin sometimes called the Boreal Transition Zone.

2.2. Wetland Locations and Extents

95 For our analysis, we selected a subset of wetlands from the Merged Alberta Wetland Inventories within each Natural Region (Figure 1). To ensure wetland conditions were indicative of the Natural Regions within which they resided, we excluded those within 500 m of a Natural Region boundary. Then, we randomly selected 12,000 wetlands in the Southern Boreal and Parkland Natural Regions (3,000 per permanence class) and 16,000 in the Grassland (4,000 per permanence class). To ensure spatial independence among sampled wetlands and their relationship to land cover as well as coincide with previous analysis of open
100 water wetlands (Ridge et al. 2017), topography (Branton et al. 2020), and land cover (Evans et al. 2017), we did not select wetlands that were within 1000 m of another selected wetland.

The distribution of wetland sizes was strongly right-skewed across the three Natural Regions of interest (Appendix A). Wetlands were typically small, with Boreal wetlands possessing the largest median size (2.26 ha), followed by Parkland wetlands (1.54 ha) and Grassland wetlands (0.58 ha), though size varies with permanence class (Appendix A). In the Grassland,
105 the largest wetlands tended to be permanently-ponded, whereas the largest wetlands in the Boreal and Parkland tended to be seasonally-ponded (Appendix A). The combination of wetland size and our digital elevation model (DEM) resolution of 25 m suggest that our median wetland sizes would occupy 36, 25, and 9 cells for Boreal, Parkland, and Grassland natural regions, respectively, defining our ability to capture variability among wetland sizes and shape.

2.3. Selecting Variables

110 To select variables representative of climate, land cover/land use and topography that would be useful in testing the relative contribution of these three factors in predicting prairie pothole wetland permanence class, we conducted a literature review using the Web of Science. We limited the search to papers published between 1950 to 2018 with the following key words: 1) Prairie Pothole Region: PPR, Northern Great Plains, Alberta, Saskatchewan, Manitoba and Dakota; 2) weather: climate, temperature and precipitation; 3) disturbance: land use, agriculture, disturbance, oil and gas, grazing and roads and 4) pond
115 permanence: watershed, hydroperiod, permanence class, catchment and wetland. We used “OR” operators between key words under the same class and “AND” operators between each key word class. To characterize topography, we selected variables that are commonly used to describe topographic variations, based on a previous review (Branton and Robinson, 2019). Details and results from this review are reported in Table 2.

2.3.1. Climate

120 We acquired 2013-2014 daily weather data from the AgroClimatic Information Service of Alberta to calculate climate variables. These data include precipitation and temperature measurements from 7,914 weather stations across the province, observed from October 2013 to August 2014. We calculated seasonal precipitation totals and temperature averages from a compilation of proxy variables (Table 2) at each station. Then, using a simple inverse distance weighting (Tarroso et al., 2019), we interpolated climate variables at the center of each wetland in R (R Core Team, 2019).

125 We used annual data on climate variables in this analysis because it was available at a fine spatial resolution and corresponded with the 2014 land cover and topography data we used. Additionally, 2014 was a typical year in terms of climate variables. For example, we found no significant difference in mean annual precipitation between the 1981-2010 climate normal and the annual data from 2013-2014 (paired t-tests Grassland: $t_9 = 1.833$, p -value = 0.652; Parkland: $t_9 = 1.833$, p -value = 0.344) or 2014-2015 (paired t-tests Grassland: $t_9 = 1.833$, p -value = 0.878; Parkland: $t_9 = 1.833$, p -value = 0.315) in either the
130 Grassland or Parkland Natural Regions (cumulative precipitation plots in Appendix B). Importantly, the influence of climate variables on wetland permanence classes will exhibit time lags dependent on site-specific factors, such as soil storage, ground water movement, and vegetation succession within catchments. Consequently, the temporal window of relevant weather would also be site-specific, and we lack a defensible justification on which to base a threshold for including or excluding annual data on climate variables. Coupled with the typical nature of 2014's annual data on climate variables, we elected to use the single
135 year as representative of average conditions in our study area and maximize comparability to our 2014 topography and land use data. We suggest future research could seek to elucidate how legacy effects of climate and land use may influence wetland permanence classes.

2.3.2. Land Cover and Land Use

Prior research in the PPR identified a strong concordance between landcover within 500 m of wetlands and wetland
140 physicochemical conditions (Kraft et al., 2019). Using land cover data from Agriculture and Agri-Food Canada's (AAFC) Annual Crop Inventory for 2014 (Agriculture and Agri-Food Canada, 2014), we calculated the proportion of each land cover class within a 500 m buffer of each wetland (Table 2). In addition to land cover characteristics, we also measured the distance of each wetland centroid to the nearest road using the National Road Network from the Government of Canada (Statistics Canada, 2010). The landscape fragmentation created by road networks has been shown to alter hydrological flow and divert
145 surface runoff (Shaw et al. 2012) such that wetlands in proximity to roads typically have shorter hydroperiods. We estimated these land cover and land use variables in ArcMap 10.4.1 (ESRI, 2012).

2.3.3. Topography

We quantified topographic characteristics of the landscape surrounding each wetland using a 25-m digital elevation model (DEM) for southern and central Alberta (AltaLIS, 2015) (Figure 1). We estimated eight terrain variables (Table 2) using
150 ArcMap 10.4.1 (ESRI, 2012) and SAGA 2.3.2 (Conrad et al., 2015). These variables may be grouped as those with local (e.g., standard deviation of slope) versus global (e.g., terrain surface convexity) application (Branton and Robinson 2019). For local variables, we applied the formula to areas only within 500 m of the wetland boundary. With global variables, we applied a 100 × 100-m moving window and computed the mean value within the 500 m buffers (Table 2).

2.4. Data Analysis

155 We aimed to quantify the relative contribution of annual data on climate variables, land cover/land use and topography for
different wetland permanence classes and determine the ability of these drivers to predict wetland permanence class. Achieving
these two outcomes involved four steps: reducing the number of variables to an orthogonal and parsimonious set for
application; visualizing if wetlands could be partitioned based on their permanence class; parametrizing and calibrating a
predictive model; and then predicting permanence class and assessing model fit. These analyses were performed in R (R Core
160 Team, 2019) and while they quantify a relationship among our independent variables with wetland permanence, they do not
infer causation.

2.4.1. Predicting Wetland Permanence Class

We used an extreme gradient boosting model to predict wetland permanence class for each Natural Region based on a
combination of annual data on climate variables, land cover/land use and topography variables (Appendix C). Extreme gradient
165 boosting is considered a more robust predictive tool than random forest (Sheridan et al., 2016). Like random forest, extreme
gradient boosting creates an ensemble of decision trees that partition data based on a specified grouping (Hastie et al., 2009;
McCune et al., 2002), which in our case is wetland permanence class. In the first decision tree, all observations are equally
weighted (Cutler et al., 2007). The second decision tree attempts to correct for misclassifications derived from the first tree,
assigning a higher weight to observations that were difficult to classify. Each subsequent tree attempts to minimize model error
170 by classifying these error-prone observations (Cutler et al., 2007). The use of the minimum error to build a model ensemble
makes extreme gradient boosting models prone to overfitting (Cutler et al., 2007). To correct for overfitting, extreme gradient
boosting models include a regularized object that penalizes more complex trees (Chen and Guestrin, 2016). We used relatively
low learning rates (<0.1) and restricted tree depths (5-7) to balance overfitting and bias in our model ensembles (Appendix D).

After parametrizing the model, we predicted wetland permanence class in the 1) Southern Boreal, 2) Parkland and 3)
175 Grassland Natural Regions. For each model, we also assessed its performance using test data (70:30 training to test ratio) to
determine the misclassification error rate, comparing results between training and test data. Importantly, misclassification error
rates reflect the proportion of sites classified as a permanence class by the models that differs from the permanence class
assigned it in the wetland inventory and thus assumes that the inventory accurately classifies each wetland. It also does not
differentiate between the misclassification of a temporary wetland as seasonal (perhaps a minor error) and the misclassification
180 of a temporary wetland as permanent (a major error). Consequently, we also broke misclassification rates down by inventory
class for each model. We also evaluated the relative importance of each variable in predicting permanence class by comparing
gain values and assessed under which ranges of each variable a permanence class was more likely to occur with waterfall plots.

3. Results

3.1. Selecting Variables

185 Before predicting wetland permanence class based on land use and landcover, topography and annual data on climate variables, we first determined which metrics were collinear within their metric class. Based on a maximum allowable correlation Pearson correlation of 0.9, we reduced our 30 metrics to 19 that reflected climate (7), land cover/ land use (4) and topography (8) (Table 2). Next, we incorporated these 19 variables into a PCA to explore partitioning of permanence classes in accordance with the annual data on climate, land cover, and topography variables and to facilitate comparison among the three Natural Regions.

190 Wetlands in the Grassland appeared to be better aligned with all three domains than the wetlands in the Southern Boreal and Parkland Natural Regions (Figure 2).

3.2. Model Performance

We built an extreme gradient boosting model for each Natural Region (southern Boreal, Parkland and Grassland) in our study area. Our models had moderate to high error rates for both the training (43-50%) and test datasets (48-61%; Appendix D),

195 which indicates a balance between bias and overfitting. Clearly, annual data on climate, land use/cover and topography alone are not sufficient to perfectly predict wetland permanence class. We conclude that while our models are useful in ranking the relative importance of climate, land cover/land use and topography variables in predicting wetland permanence class, they are not a comprehensive overview of the factors determining permanence class of a given wetland (see Section 4.5). Notably, we focus on the context of each wetland (surrounding topography, land cover/ land use, and climate), rather than wetland-specific

200 properties that would influence permanence class (e.g., basin morphology).

3.3. Relative Importance of Variables in Predicting Wetland Permanence Class Among Natural Regions

In the Parkland and Grassland Natural Regions, annual data on climate explained the greatest amount of variance in wetland permanence class, based on relative gain values (Figure 3A-C). As anticipated, our results suggest that climate conditions vary systematically among the Natural Regions (Figure 4A-D). Among the climate variables included in our analyses, spring

205 temperature (Boreal: $6.87^{\circ}\text{C} \pm 0.425$ SD; Parkland: $6.85^{\circ}\text{C} \pm 0.206$ SD; Grassland: $8.14^{\circ}\text{C} \pm 0.892$ SD) explained the highest magnitude of variance in predicting permanence class in the Grassland (Figure 3A;2C) but was less important in the Southern Boreal and Parkland where values are less extreme (Figure 4A). Winter snowpack (Boreal: 92.15 cm ± 20 SD; Parkland: 67.65 cm ± 14.99 SD; Grassland: 42.14 cm ± 15.06 SD) explained the highest magnitude of variance in predicting permanence class in the southern Boreal and Parkland, and these amounts were distinctly lower in the warmer Grassland (Figure 4C).

210 Land cover/land use was the second most important category of drivers of wetland permanence class, following annual data on climate in the Grassland Natural Region (Figure 3F), but not in the Southern Boreal or Parkland (Figure 3). Yet, unlike climate, land cover/ land use did not vary systematically among the three Natural Regions (Figure 4E-H). Wetlands surrounded

by cropland had shorter hydroperiods in the southern Boreal and Parkland (Figure 5D-E) but wetlands surrounded by natural vegetation had shorter hydroperiods in the Grassland (Figure 5F).

215 Topography was the most important category of drivers of wetland permanence class in the southern Boreal and second most important in the Parkland, and the order of importance for the terrain metrics were nearly the same in both Natural Regions (Figure 3G-I). Though topography metrics were the least important category in the Grassland (Figure 3I), apart from deviation from mean elevation (Figure 4I), variables associated with topography did not systematically vary among Natural Regions (Figure 4J-L).

220 **3.4. Wetland Permanence Class in the Boreal, Parkland and Grassland**

Our findings suggest that wetland permanence class in the Prairie Pothole Region of Alberta correlates with climate, topography and, to a lesser extent, to surrounding land cover/ land use. Generally, across the three Natural Regions, wetlands with shorter hydroperiods (e.g., temporary and seasonal) were typically situated in landscapes with higher spring snowpack amounts (Figure 3A-C) and spring temperatures (e.g., Figure 5A). Longer hydroperiod wetlands were typically situated in
225 landscapes with more summer precipitation and lower spring temperatures (e.g., Figure 5C), occupying relatively low topographic positions with low terrain convexity (e.g., Figure 5G, H), and, in the Grassland, were sometimes surrounded by less natural cover (Figure 5F), though in the southern Boreal they were more common where cropland was less than 25% cover (Figure 5D) and less than 75% in the Parkland (Figure 5E). Interestingly, the relative importance of variables in predicting the occurrence of both shorter and longer-hydroperiod wetlands were shared, and this agreement was strongest between the
230 southern Boreal and Grassland (Appendices F & H).

4. Discussion

Our findings support the assertion of other published studies (e.g., Fay et al., 2016; Johnson et al., 2010a, 2005; Johnson and Poiani, 2016; Reese and Skagen, 2017; Werner et al., 2013; McKenna et al. 2019), which conclude that climate change will affect wetland hydroperiod or permanence class. We anticipate that reduced winter snowpack will dry out temporarily- and
235 seasonally-ponded wetlands, while warmer spring temperatures will reduce the hydroperiod of more permanently ponded wetlands. Yet, annual data on climate is not the only element correlated with wetland permanence class in Alberta's PPR - our analysis used a relatively coarse DEM (25 m), and we nonetheless found that topography was important in predicting permanence class. Consequently, failure to consider topography limits our understanding about the extent to which hydroperiod, and therefore wetland permanence class, may change in response to climate change. We speculate that the use of
240 finer-scale elevation models derived from high resolution LiDAR (e.g., 1 m) or remotely piloted aircraft (e.g., 2-5 cm) will reveal even greater importance of topography in surface runoff and wetland hydroperiod, particularly in the Grassland Natural Region, where wetlands were typically smaller and topographic variation relatively subtle.

4.1. Importance of Climate

The sensitivity of wetland hydroperiods to annual climate data is corroborated in existing literature, which emphasizes that the semi-arid climate drives the region's sensitivity to climate change (Fay et al., 2016; Johnson et al., 2004; Schneider, 2013). In the southern Boreal and Grassland, regions with warmer spring temperatures are likely to experience an earlier onset of spring snowmelt, higher water deficits (Schneider, 2013; Zhang et al., 2011) and lower pond permanence classes for wetlands, whereas cooler peak spring temperatures favour greater pond permanence in these Natural Regions. In the southern Boreal and Parkland, winter snowpack depth was the most important climate variable; and this, we attribute to temporarily- and seasonally-ponded wetlands requiring a minimum threshold of winter snowpack amount to persist, whereas permanently ponded wetlands also benefit from precipitation in other seasons and so can exist at lower winter snowpack amounts (Figure 5B). Because climate forecasts suggest that warmer springs and changes in precipitation timing are likely (Zhang et al., 2011), our finding that climate was the most important domain of variables in predicting permanence class supports previous studies that suggest PPR wetlands are sensitive climate change (Johnson et al., 2010; Paimazumder et al., 2013; Schneider, 2013; Viglizzo et al., 2015; Zhang et al., 2011).

4.2. Importance of Topography

Despite recognition that topography is a useful proxy in wetland mapping (Branton and Robinson, 2019; Los Huertos and Smith, 2013) and that topography must influence surface runoff generating processes that are essential to wetland function (Hayashi et al., 2016; Mushet et al., 2018), the relative importance of topography in hydrological processes is somewhat debated (Devito et al., 2005). Simulations predicting the influence of climate change on the size and isolation of prairie pothole wetlands have focused on climate and land cover/land use (Anteau et al., 2016; Chasmer et al., 2012; Conly et al., 2001; Johnson and Poiani, 2016; McCauley et al., 2015; Steen et al., 2016; Voldseth et al., 2007). Consequently, 1) there is a lack of research quantifying topographic characteristics of wetlands and the landscapes within which they occur, 2) links between topography, vegetation and wetland condition have not been rigorously studied and 3) policy and guidelines on wetland mitigation and compensation prescribe width-to-length ratios and slopes that are characteristic of permanently-ponded wetlands (Environmental Partnerships and Education Branch Alberta, 2007), which are less abundant in all three Natural Regions (Table 1). Despite remaining numerically more abundant, small and more temporarily ponded wetlands are being preferentially lost in Alberta's Prairie Pothole Region (Serran et al., 2017). If we had a better understanding of how topographic structure determines wetland hydrology/function, we could revise policy and regulations governing wetland management to ensure we better match natural landscapes in their frequency and distribution of wetland permanence classes.

4.3. Importance of Land Cover/Land Use

Existing literature identified land cover/land use as the second greatest driver of wetland conditions following climate (Anteau et al. 2016). In the Grassland, the terrain is relatively flat compared to the southern Boreal and Parkland (Alberta Tourism

Parks and Recreation, 2015). Consequently, after the important role of annual data on climate in the more arid Grassland
275 Natural Region (Government et al., 2014), land cover/land use might be a stronger driver of permanence class than topography.
Importantly, the percent cover of natural vegetation is typically low in the Grassland, where most land has been converted to
cropland or pastureland (Alberta Tourism Parks and Recreation, 2015). Combined with the process of consolidation drainage,
which shunts water from scattered low hydroperiod wetlands, concentrating it in larger more permanently ponded wetlands
downstream (McCauley et al., 2015), this leads to Grassland landscapes with more natural cover being more likely to contain
280 temporary and seasonal wetlands. Thus, wetlands surrounded by natural vegetation may have shorter hydroperiods because
cropland resists infiltration and natural vegetation intercepts snow-sourced surface runoff (Anteau, 2012; van der Kamp et al.,
2003; Voldseth et al., 2007), which can account for up 27% of ponded water amounts (van der Kamp et al., 2003).

Because some landscapes in the PPR are flatter than others (Schneider, 2013), and land use activities can modify the
terrain (Anteau 2012; Wiltermuth and Anteau 2016; Anteau et al. 2016), our findings do highlight the importance of
285 considering land use in forecasting the impacts of climate change on PPR wetlands. Boreal and Parkland wetlands have
stronger overlaps in topography metrics and annual data on climate; and, as a result, differences in land use within these regions
may be integral in determining future shifts in the frequency distribution of permanence classes. Forecasts for the province of
Alberta suggest there will be expansions in the agricultural industry within the next decade (Government of Alberta, 2015),
and this suggests that climate impacts on Albertan PPR wetlands will be compounded by land use activities.

290 **4.4. Topographic Position of Wetlands by Permanence Class**

Semi-permanent and permanently-ponded wetlands typically occur in regional or spatial neighbourhood topographic lows (as
opposed to simply local depressions, e.g., perched wetlands), likely because they 1) can hold larger volumes of ponded water
(i.e., larger pond size/volume (Novikmec et al., 2016)) and 2) receive higher volumes of water inputs from the surrounding
landscape (e.g., surface run-off, groundwater Euliss et al., 2004, 2014; LaBaugh et al., 1998; Toth, 1963)). We are unable to
295 partition the natural hydrogeological effects of topographic position on wetland permanence class from the effects of human
alteration of the surrounding landscape, yet the importance of topographic position to wetland permanence class is likely
reinforced by consolidation drainage when wetlands situated higher in the landscape are drained and the water is redirected to
wetlands positioned lower in the landscape (McCauley et al., 2015; Wiltermuth and Anteau, 2016). Because of consolidation
drainage, we may observe increases in hydroperiod of wetlands in topographic lows of the wetlandscape (e.g., sites with low
300 topographic position index values). In the arid but heavily farmed Grassland Natural Region, consolidation drainage can
eliminate temporary and seasonally ponded wetlands from areas with limited remaining natural cover (Serran et al., 2017).
This aligns with our model results, although the probability of observing a permanent or semi-permanent class wetland was
greatest at the lower end of the range of crop cover in our landscapes, the threshold of crop cover above which wetlands were
most probably seasonal or temporary in class was higher in the Grassland, lower in the Parkland, and lowest in the Boreal.
305 Thus, we recommend that future research investigate the role of topographic position on permanence class, in the absence of
human disturbance to control for the influence of consolidation drainage.

4.5. Model Error

Our model misclassification error rates were relatively high (Appendix D), indicating imperfect matching between model-predicted permanence class and inventory-reported permanence class for our study wetlands. One key source of uncertainty in our analysis is that the accuracy of the inventory in assigning wetlands a given permanence class is not validated and in interpreting our model error we must assume that the permanence classes we derived from the inventories are correct, though we know the two inventories differ in their mapping accuracy (Evans et al. 2017). Yet, we hypothesize that our inability to account for soil characteristics (Schneider, 2013) and bathymetry (Huertos and Smith 2013) likely contribute to misclassification by our models (Appendix D). Schneider (2013) stated that within Natural Regions, both elevation (which we did account for), and soil characteristics can vary across the landscape. As such, wetlands situated similarly in the landscape may not have the same soil characteristics, and soil characteristics are understood to influence wetland hydrology by dictating the proportion of incident precipitation that is converted to surface run off (Hayashi et al., 2016). Though Schneider (2013) also mentioned an influence of disturbance history on ecosystems, prior work in our study region reported no temporal lag in wetland environmental conditions and surrounding land cover (Kraft et al., 2019).

The lack of extensive data on basin morphology identifies a gap that would enrich the presented research by enabling direct classification of wetland permanence from raw bathymetric data. Such data would likely reduce the misclassification error rates of our ensemble models, which rely only on annual data on climate, land use, and topography in predicting wetland permanence. Furthermore, these data would provide added value to those conducting research on above- and below-ground hydrologic connectivity and contributing areas (e.g., Chen et al., 2020), as well as those evaluating the impacts of climate change on wetland permanence and subsequently flora and fauna health and resilience (e.g., LaBaugh et al., 2018).. As new technologies for mapping wetland bathymetry become more widely available (e.g., bathymetric LiDAR; Paine et al., 2015; Wang and Philpot, 2007) an opportunity will exist to better understand the link between wetland pattern and process.

Potentially some proportion of model error can be attributed to the use of a single year of climate and land use data as well as our relatively coarse (25 m) digital elevation model. However, it is likely that the contributions of these factors are minimal given that 1) the climate data used (year 2014) is representative of average conditions, coincides with fieldwork, and yielded the strongest among the variables interrogated and therefore improving the quality of its contribution will not change the qualitative outcome of the presented analysis; 2) previous research found that physiochemical conditions in a wetland are quite congruent with surrounding land cover of the same year with only minor differences when catchments were defined with 10 m versus 25 m resolution DEMs (Kraft et al. 2019); and 3) there was no detectible difference in wetland catchment size when they were derived from DEMs of low (10 m) versus high (3 m) resolution (McCauley and Anteau, 2014).

Lastly, wetland permanence classes are ordinal and consequently not all misclassifications are equal. From an ecohydrological perspective, a discrepancy between model-predicted and inventory-reported permanence class can be minor (e.g., temporary vs. seasonal) or major (e.g., temporary vs. permanent), and this is not accounted for in the overall misclassification error rate. When we investigate the class-based misclassification error rates, it is apparent that all models

340 were most successful in classifying wetlands at the extreme ends of the permanence class spectrum, and misclassification error
rates were higher for seasonal and semi-permanent wetlands (Appendix D). Interestingly, the Grassland model tended to
misclassify seasonal wetlands as temporary and semi-permanent wetlands as permanent (i.e., misclassified into adjoining
classes), whereas the Parkland model tended to misclassify seasonal and semi-permanent wetlands more evenly across the
three other permanence classes. Overall, these extreme gradient boosting models are valuable for comparing the relative
345 importance of the climate, topographic, and landscape domains of predictor variables, despite misclassification error rates.

5. Conclusion

Because some landscapes in the PPR are flatter than others (Schneider, 2013), and land use activities can modify topography
(Anteau, 2012; Anteau et al., 2016; Wiltermuth and Anteau, 2016), our findings also highlight the importance of considering
land use in forecasting the impacts of climate change on PPR wetlands. Southern Boreal and Parkland wetlands are most
350 congruent in the relative importance of climate and topography variables; and, as a result, differences in land use within these
regions may be integral in determining future shifts in the frequency distribution of permanence classes. Forecasts for the
province of Alberta suggest expansion in the agricultural industry over the next decade (Government of Alberta, 2015), which
suggests that climate impacts on Alberta's PPR wetlands will be compounded by changes in land use activities.

6. Code/Data availability

355 The data and code for this manuscript are published online with FigShare: <https://doi.org/10.6084/m9.figshare.18945248.v1>

7. Author Contributions

RCR conceptualized the study, acquired funding and resources, supervised and curates the data; DR and JD gathered the data;
JD analysed and visualized the data and wrote the original draft; all authors contributed to investigation and review and editing.

8. Competing Interests

360 The authors declare that there are no competing interests or conflicts of interest.

365 **Table 1.** Descriptions of the four permanence classes included in our study. We describe the typical length of time that these prairie pothole wetlands will contain ponded water, their associated vegetation zones, as described by Stewart and Kantrud (Stewart and Kantrud, 1971), and the number of wetlands belonging to each class in the Alberta Merged Wetland Inventory (Government of Alberta, 2014) that were within the extent of our 25-m Digital Elevation Model.

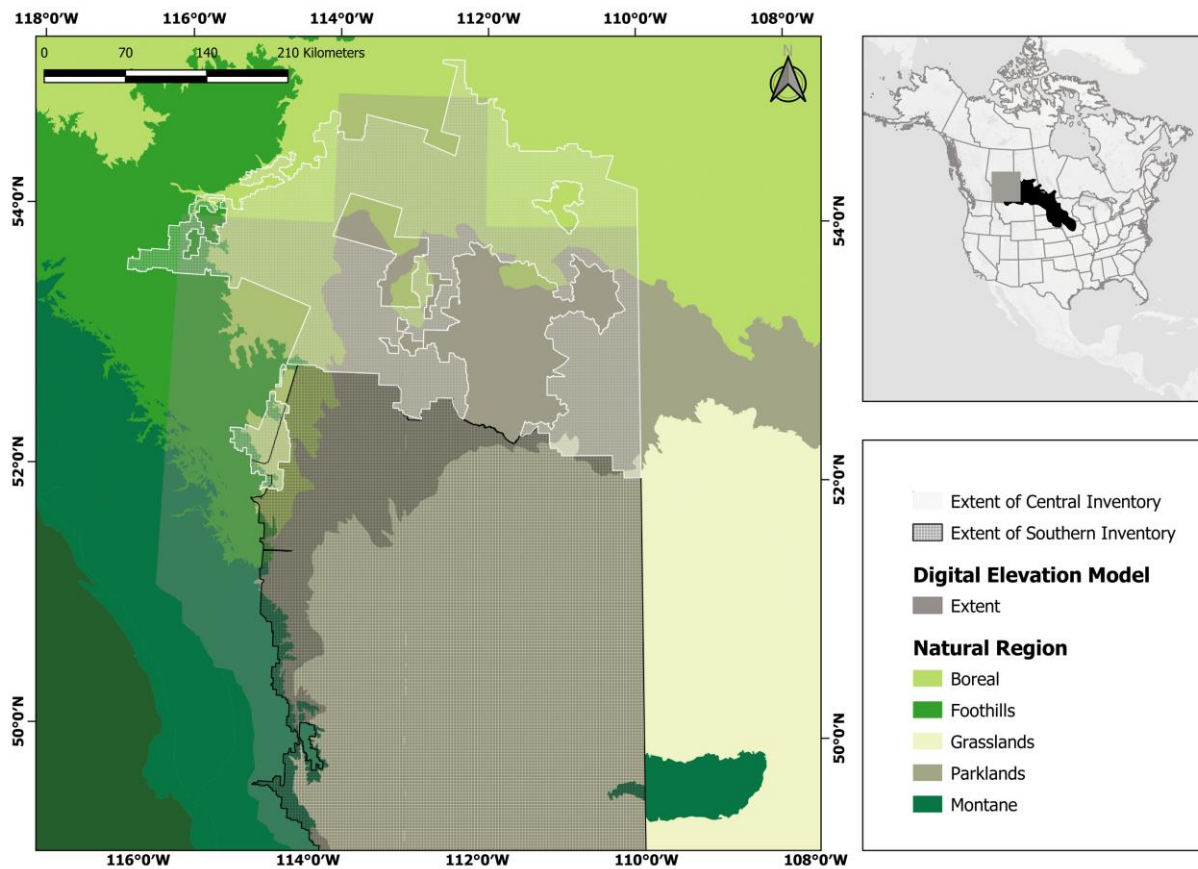
| Permanence class | Typical hydroperiod | Vegetation zones | Natural Region | | |
|------------------|--|--|----------------|----------|-----------|
| | | | Boreal | Parkland | Grassland |
| Temporary | Until mid-spring, typically for four weeks | Wet-meadow (includes wet-meadow emergent), low-prairie, high seepage | 40461 | 51062 | 153872 |
| Seasonal | Late spring to early summer for approximately two months | Shallow-marsh (vegetation zones from shallow to deep: emergent plants, submerged aquatic plants), wet-meadow, low prairie | 30890 | 43836 | 108924 |
| Semi-permanent | Dries fully in drought years only | Deep-marsh (vegetation zones from shallow to deep: emergent vegetation, open-water with bare-soil), shallow-marsh, wet-meadow, low-prairie | 39375 | 47075 | 12240 |
| Permanent | Open water year-round | Open water, deep marsh, shallow-marsh, wet-meadow, low-prairie | 5704 | 10785 | 4952 |

Table 2. List of annual data on climate, land cover and land use, and terrain metrics used to predict wetland permanence class. In this table, we include a description of the significance of each metric for wetland hydroperiod and the proxy metrics we selected. For our analysis, winter months range from November to February, spring April to May and summer June to August. We used Web of Science to conduct this review, limiting the search to papers published between 1950 to 2018, and key words for: 1) the PPR: Prairie Pothole Region, Northern Great Plains, Alberta, Saskatchewan, Manitoba and Dakota; 2) weather: climate, temperature and precipitation; 3) disturbance: land use, agriculture, disturbance, oil and gas, grazing and roads and 4) pond permanence: watershed, hydroperiod, permanence class, catchment and wetland. We used “OR” operators between key words under the same class and “AND” operators between each key word class. For the terrain metrics, we used selected metrics that are commonly used to describe topographic variations, based on a previous review (Branton and Robinson, 2019). Notably, Branton and Robinson (2019) employed controls on collinearity, including PCA.

| Category | Variable | Significance for Wetland Hydroperiod/Formula | Proxy/Class ¹ |
|----------|----------------------------------|--|---|
| Climate | Snowpack/Winter precipitation | Snowpack accounts for 30-60% of ponded water amounts (Hayashi et al., 1998; Tangen and Finocchiaro, 2017) Longer hydroperiods with higher winter precipitation (Collins et al., 2014) | Total Spring Precipitation Total Winter Precipitation Total Precipitation in Winter & Spring Total Spring Snowpack Total Winter Snowpack Total Snowpack in Winter & Spring |
| | Sumer Precipitation | Longer hydroperiods from increased summer precipitation (Clare and Creed, 2014; Eisenlohr, 1972; Euliss et al., 2014; Leibowitz and Vining, 2003) | Total Summer Precipitation |
| | Summer Temperature | Evapotranspiration rates/water losses higher in summer (from June) (Heagle et al., 2007) | Average Maximum Temperature in June Average Maximum Temperature in July Average Maximum Temperature in Summer |
| | Winter/Spring/Summer Temperature | Snowpack may melt too fast with warmer conditions (Crosbie et al., 2013) | Average Maximum Temperature in Spring Average Maximum Temperature in Winter Average Maximum Temperature in Spring & Winter |

¹ This differentiates terrain metrics by global (estimated using a 100 × 100-m moving window and mean value within 500-m buffer recorded) and local (estimated within a 500-m buffer of the wetland).

| Category | Variable | Significance for Wetland Hydroperiod/Formula | Proxy/Class ¹ |
|-----------------------|---|---|--|
| | Precipitation Timing | Fewer wetlands dry up when summer precipitation is earlier in the summer (Meyers, 2018; Vinet and Zhedanov, 2011) | Proportion of Summer Precipitation in June |
| Land Use & Land Cover | Natural Vegetation | Loss of natural cover increases surface runoff (Clare and Creed, 2014) | % Natural Cover |
| | Cropland Cover | Because soil is less porous (more compacted), much of the accumulated water, either from the snowpack or spring/summer precipitation, flows into the wetland – this increases water levels (van der Kamp et al., 2003; Voldseth et al., 2007) | % Cropland Cover |
| | Urban Cover | Longer hydroperiods in urban landscapes, mostly because of higher runoff (when compared to those in croplands) (Fossey and Rousseau, 2016) | % Urban Cover & Bare Ground |
| | Grazing | Grazing lowers snow accumulation (Willms and Chanasyk, 2013), which can increase runoff and hydroperiod (Collins et al., 2014; Niemuth et al., 2010) | % Pastureland |
| | Culverts/Roads | Lowers hydroperiods by blocking surface runoff (Shaw et al., 2012) | Distance to Road |
| | Tilling | Can lower pond area/depth, and by extent hydroperiod, as increases in sedimentation can in fill ponds (Skagen et al., 2016) | % Cropland Cover |
| Terrain Metrics | Mean Elevation (DEM) - Deviation | $ Elevation - Elevation_{mean} ^2$ | Local |
| | Elevation (DEM) - Standard Deviation | $\sqrt{\frac{\sum(Elevation - Elevation_{mean})}{n}}$ | Local |
| | Profile Curvature (PC) - Standard Deviation | $\sqrt{\frac{\sum(Profile\ Curvature - Profile\ Curvature_{mean})}{n}}$ | Local |
| | Slope - Standard Deviation | $\sqrt{\frac{\sum(Slope - Slope_{mean})}{n}}$ | Local |
| | Terrain Surface Convexity | Percentage of upwardly-convex cells within the moving window(Iwahashi and Pike, 2007). | Global |
| | Terrain Surface Texture | Relative frequency of pits and peaks in a 100 × 100-m moving window(Iwahashi and Pike, 2007) | Global |
| | Topographic Position Index | $\frac{Elevation_{mean} - Elevation_{min}}{Elevation_{max} - Elevation_{min}}$ | Local |
| | Slope Variability | $\frac{Slope_{max} - Slope_{min}}{Slope_{max} - Slope_{min}}$ | Local |



380 **Figure 1:** Extents of the Central and Southern Wetland Inventories (Government of Alberta, 2014) used to delineate wetlands
 in our study. We selected wetlands from three Natural Regions – Boreal (12,0000), Parkland (12,0000) and Grassland
 (16,0000); Natural Region boundaries are sourced from the Government of Alberta (Government Alberta, 2016). These
 wetlands are within the southern-Albertan Prairie Pothole Region. There are 356,246 wetlands delineated in the Southern
 Inventory and 253,873 in the Central Inventory. DEM data provided by AltaLIS, (2015).

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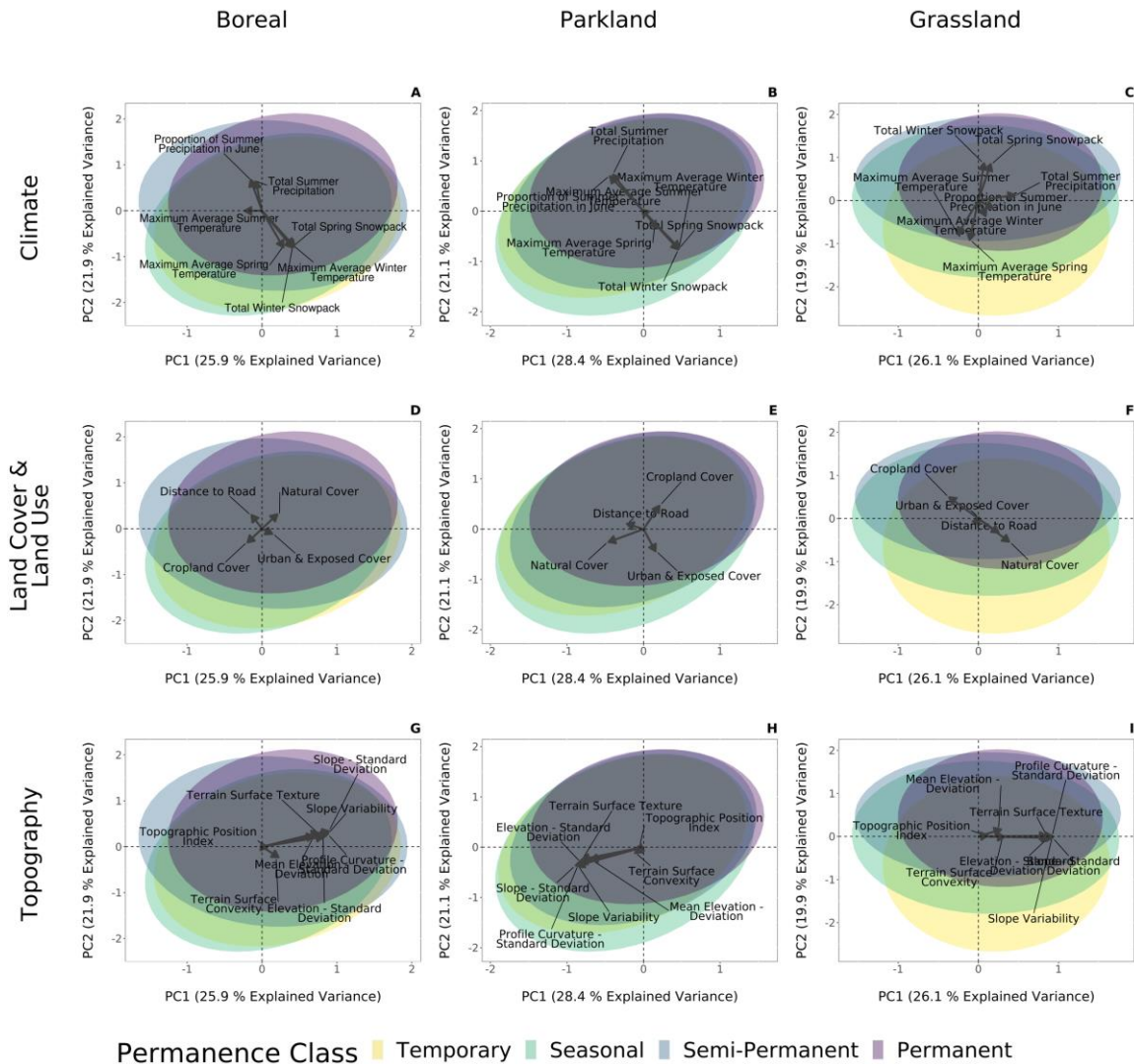
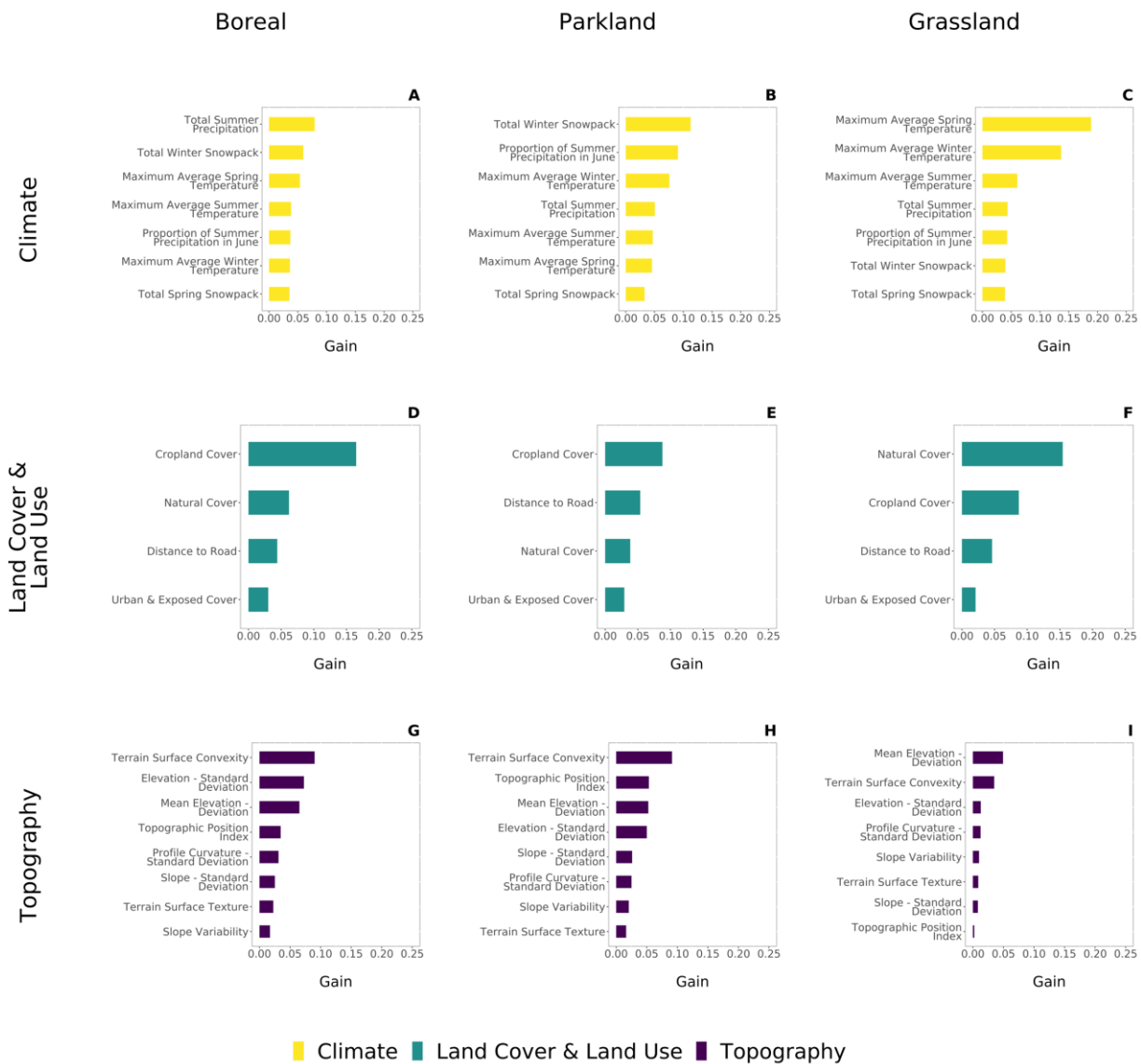


Figure 2. Principal Components Analysis for wetlands delineated in the 1) Boreal (totalling 12,000 wetlands), 2) Parkland (totalling 12,000 wetlands) and 3) Grassland (totalling 16,000 wetlands) Natural Regions. PCAs apply an orthogonal transformation to summarize the data into axes that explain the variance between two correlation matrices. Our data were scaled before implementing the PCA. Vectors on climate (A-C), land use and land cover (D-F) and terrain roughness (G-I) show correlations with both axes. Axis two, for all datasets, represents a hydroperiod gradient and terrain roughness is represented on axis 1.



395 **Figure 3.** Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the 1) Boreal (totalling 12,000 wetlands), 2) Parkland (totalling 12,000 wetlands) and 3) Grassland (totalling 16,000 wetlands) Natural Regions. These variables were proxies for climate (A-C), land cover and land use (D-F) and topography (G-I). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

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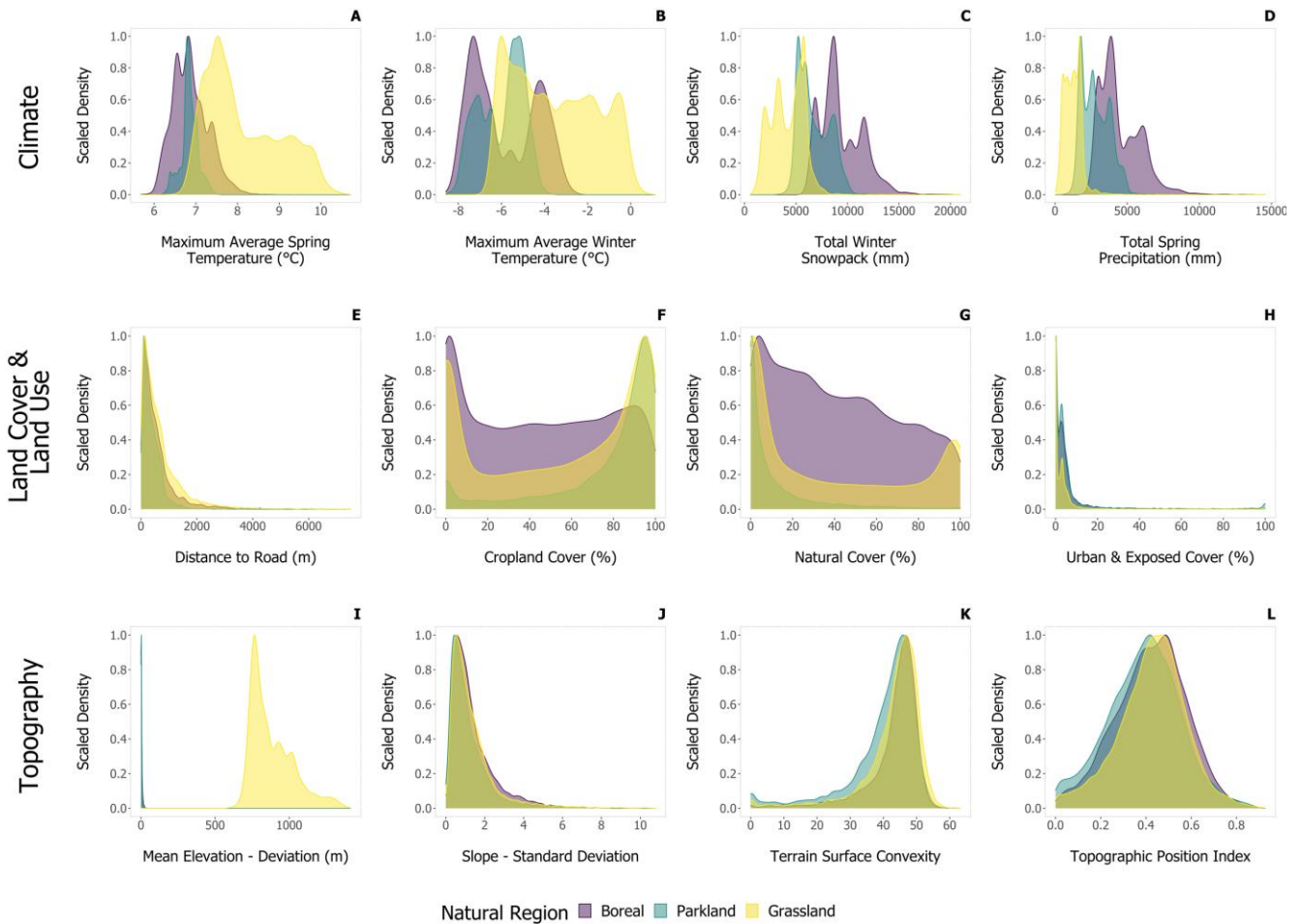


Figure 4. Frequency distribution of the top four climate, land cover and land use and topography variables by Natural Region.

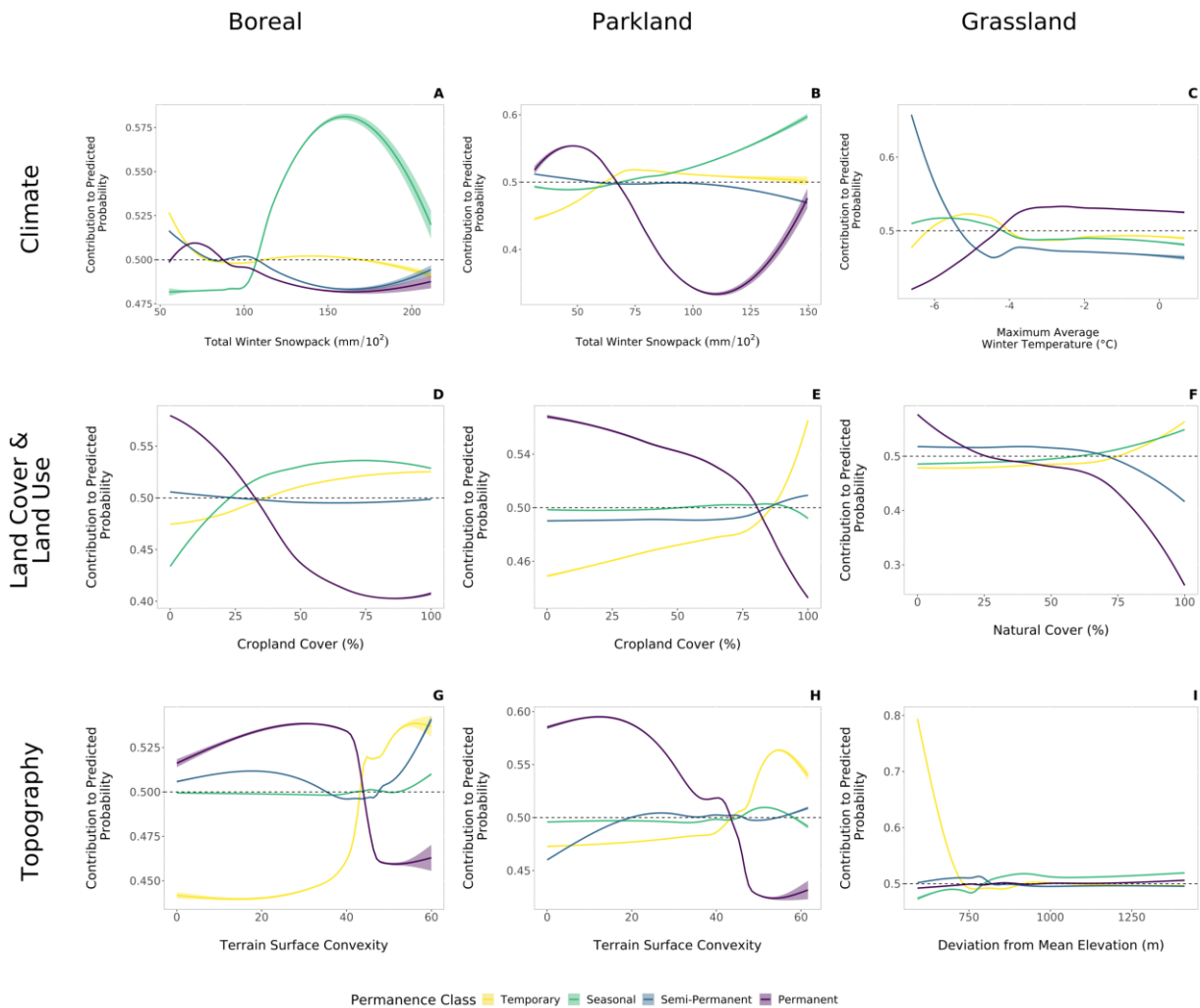
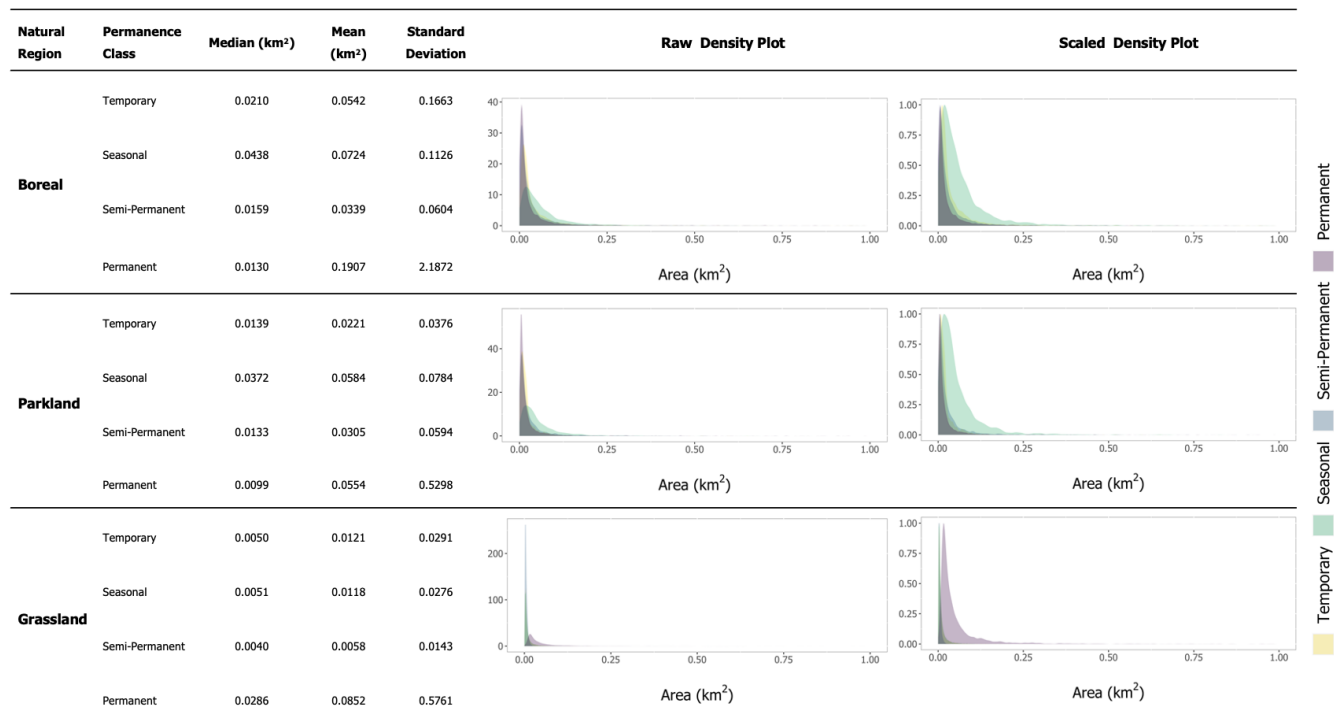


Figure 5. Partial dependence plots for the four wetland classes – temporary, seasonal, semi-permanent and permanent based on top metrics. Predicted probabilities below 0.5 suggest that at this measured value of the metric, observing that permanence class is unlikely. We show 95% confidence intervals and used a generalized additive model-based trend line. Probabilities were derived from extreme gradient boosting models for wetlands delineated in the 1) Boreal (totalling 12,000 wetlands), 2) 410 Parkland (totalling 12,000 wetlands) and 3) Grassland (totalling 16,000 wetlands) Natural Regions.

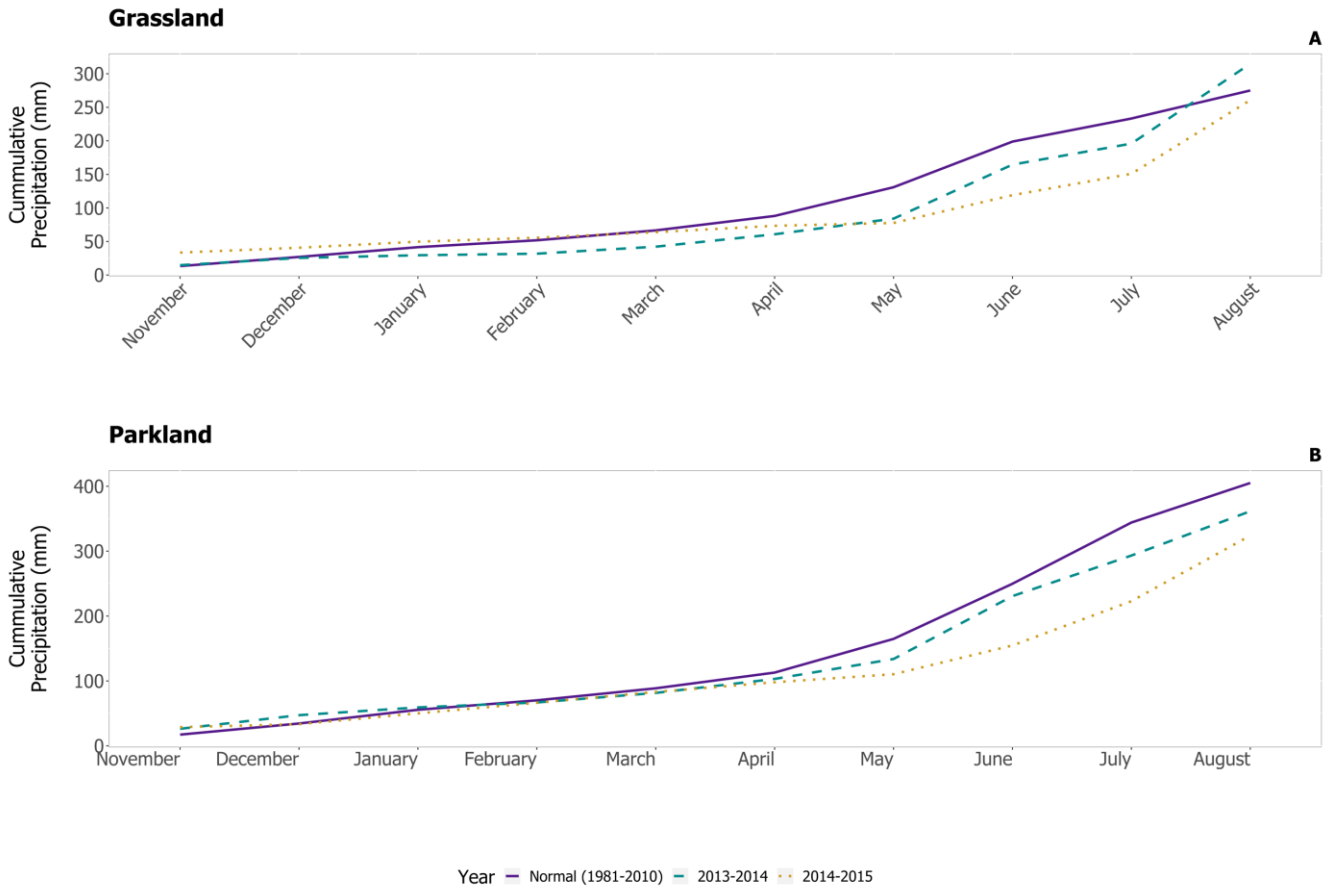
Appendix A.

415 Table 1A. Frequency distribution of wetland sizes in in the Boreal, Grassland and Parkland Natural Regions. Data on wetland sizes were acquired from the Alberta Merged Inventory (Government of Alberta, 2014)



420 **Appendix B.**

Figure 1B. Comparison of cumulative precipitation in the Grassland (panel A) and Parkland Natural Region (panel B) between 2014-2015 to the climate normal. Note data was not available for the southern portion of the Boreal Natural Region of interest in our study.



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Appendix C.

Table 1C. List of parameters tuned for the extreme gradient boosting model, a description of these parameters, their ranges and the ranges evaluated in our cross validation.

| Parameter | Description | Range |
|--|---|---|
| Learning rate | Used to control the contribution of each tree to model. Lower values result in the model being more robust to overfitting. | Typical: 0-1 Model: 0-0.3 Boreal (0.01), Parkland (0.1); Grassland (0.05) |
| Gamma | This controls the complexity of the model. It determines how much loss (difference between prediction and observation) allowable for the formation of a new node. | Typical: 0-20 Model: 0-10 Boreal (8), Parkland (4); Grassland (10) |
| Maximum depth of a tree | This sets the maximum number of nodes that can exist between the tree root and leaves. The larger the value, the more likely a tree is to overfit. | Typical: 1-7 Model: 1-7 Boreal (5), Parkland (7); Grassland (7) |
| Minimum sum of instance weight needed in a child | This sets a minimum weight/purity of data (e.g., number belonging to a given group) for splitting to create a new node in a tree. The higher this number is, the more conservative the algorithm will be. | Typical: 1-7 Model: 1-7 Boreal (5), Parkland (3); Grassland (7) |
| Subsample ratio of the training instance | This sets the number of rows (fractional) that should be included in building a tree. | Typical: 0-1 Model: 0.6-1 Boreal (0.8), Parkland (0.65); Grassland (0.7) |
| Subsample ratio of columns when constructing each tree | This sets the number of predictors (fractional) that should be considered in each tree. | Typical: 0-1 Model: 0.6-1 Boreal (0.8), Parkland (1); Grassland (0.9) |

Appendix D.

Table 1D. Value of parameters used in extreme gradient boosting models for our three datasets, the misclassification error rates and number of trees for our models.

| Parameter | Natural Region | | |
|--|--------------------------------|--------------------------------|-------------------------------|
| | Boreal | Parkland | Grassland |
| Learning rate | 0.01 | 0.1 | 0.05 |
| Gamma | 4 | 6 | 8 |
| Maximum depth of a tree | 6 | 5 | 7 |
| Minimum sum of instance weight needed in a child | 1 | 5 | 7 |
| Subsample ratio of the training instance | 0.8 | 0.90 | 0.70 |
| Subsample ratio of columns when constructing each tree | 1.0 | 1.0 | 0.90 |
| Misclassification error rate | 49.6 (training) 60.6 (test) | 50.1 (training) 59.7 (test) | 42.9 (training) 47.8(test) |
| Number of trees | 37 | 52 | 46 |

Table 2D. Breakdown of misclassification error by permanence class. Data reflect the percent of wetlands classified as a given permanence class by the inventory (row) that the model classified as each permanence class (column). In general, models fared better at classifying temporary and permanent wetlands, and exhibited more misclassification errors in classifying wetlands that the inventory categorized as seasonal or semi-permanent. Semi-permanent wetlands, in particular, tended to be under predicted by the models.

| | | Temporary (n = 3577) | Seasonal (n = 3084) | Semi-Permanent (n = 1770) | Permanent (n = 3569) |
|------------------|------------------------------|-------------------------|------------------------|------------------------------|-------------------------|
| Boreal | Temporary (n = 3000) | 52 | 24 | 10 | 13 |
| | Seasonal (n = 3000) | 26 | 46 | 11 | 17 |
| | Semi-Permanent (n = 3000) | 26 | 18 | 28 | 27 |
| | Permanent (n = 3000) | 14 | 15 | 9 | 62 |
| | | Temporary (n = 3309) | Seasonal (n = 2144) | Semi-Permanent (n = 1976) | Permanent (n = 4571) |
| Parkland | Temporary (n = 3000) | 54 | 14 | 12 | 20 |
| | Seasonal (n = 3000) | 24 | 34 | 14 | 28 |
| | Semi-Permanent (n = 3000) | 23 | 15 | 29 | 34 |
| | Permanent (n = 3000) | 10 | 8 | 11 | 71 |
| | | Temporary (n = 4208) | Seasonal (n = 4025) | Semi-Permanent (n = 2981) | Permanent (n = 4786) |
| Grassland | Temporary (n = 4000) | 60 | 19 | 8 | 14 |
| | Seasonal (n = 4000) | 23 | 53 | 9 | 15 |
| | Semi-Permanent (n = 4000) | 12 | 16 | 45 | 27 |
| | Permanent (n = 4000) | 11 | 13 | 12 | 64 |

Appendix E.

Figure 1E. Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the Boreal (totalling 12,000 wetlands) by permanence class. These variables were proxies for climate (A-D), land cover and land use (E-H) and topography roughness (I-L). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

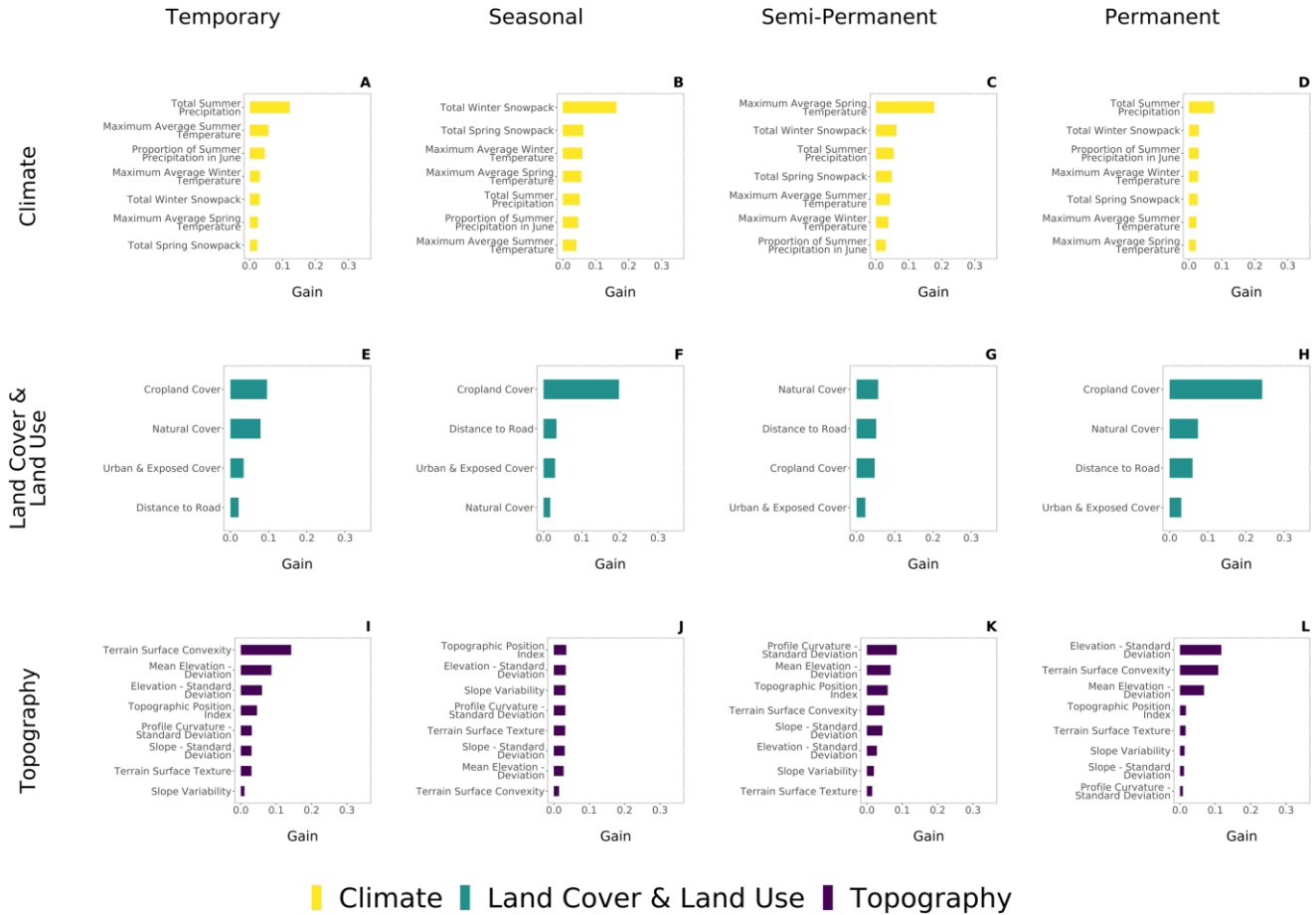
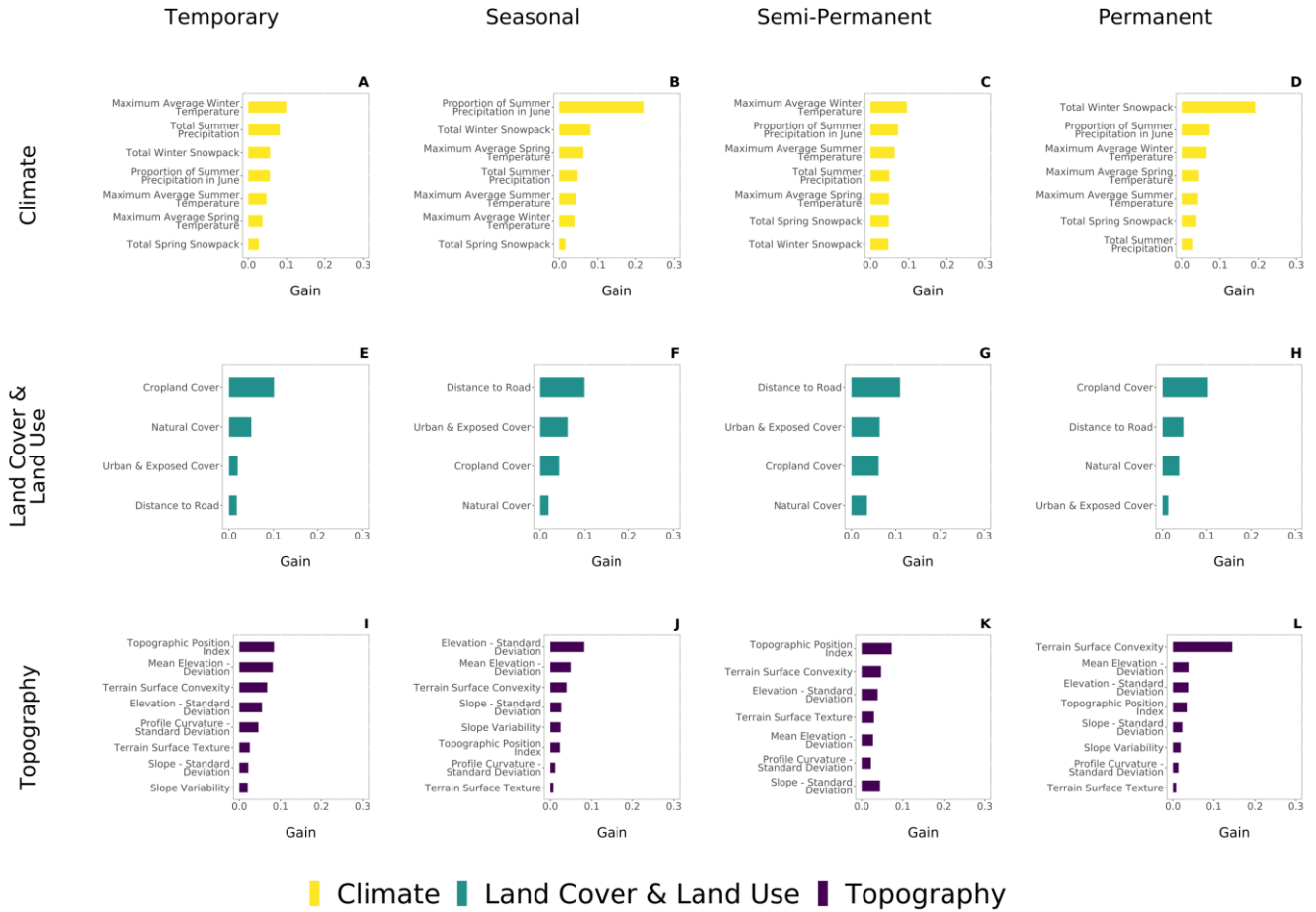


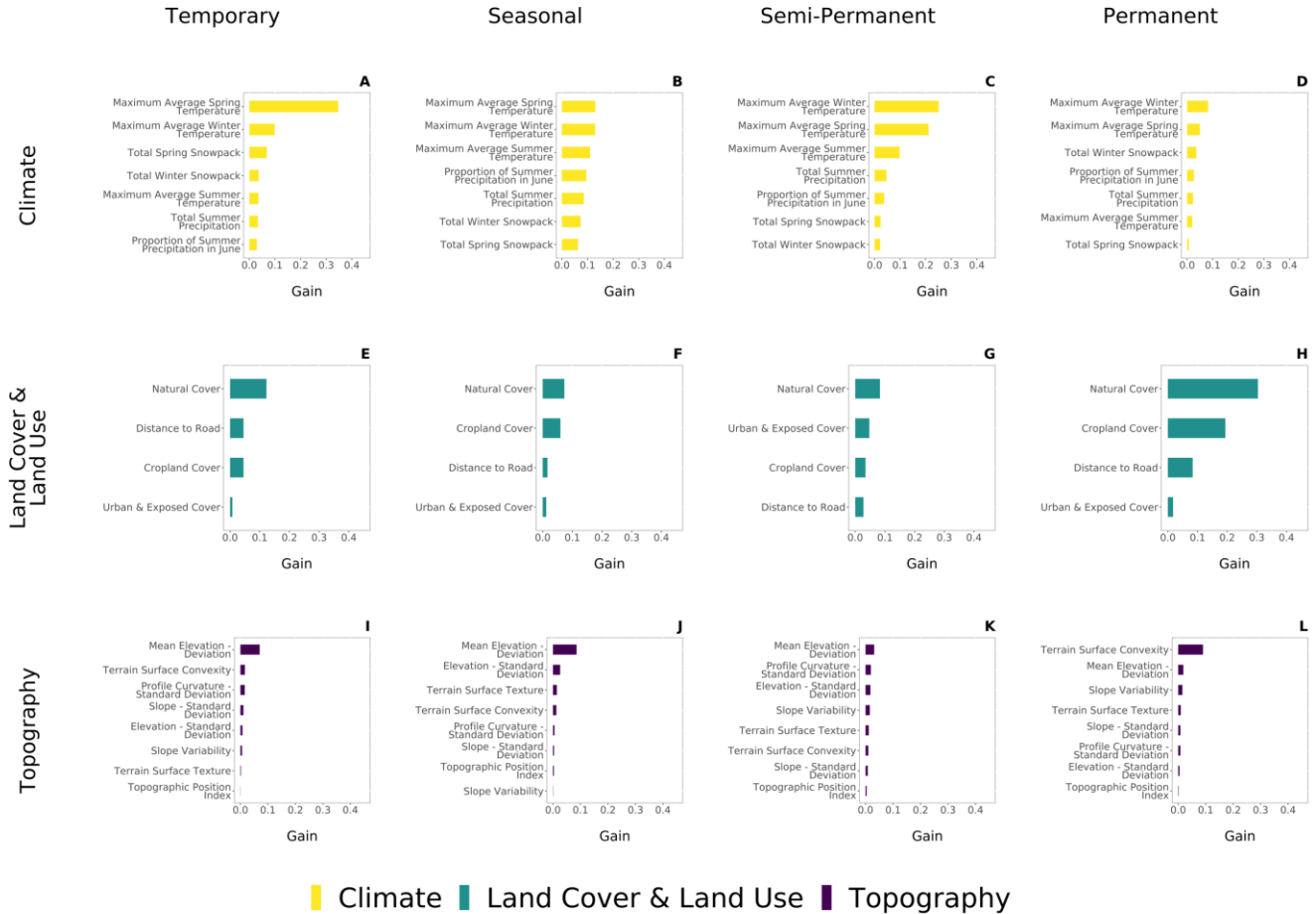
Figure 1F. Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the Parkland (totalling 12,000 wetlands) by permanence class. These variables were proxies for climate (A-D), land cover and land use (E-H) and topography roughness (I-L). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.



Appendix G.

Figure 1G. Variables ranked by their importance in the extreme gradient boosting models for wetlands delineated in the Grassland (totalling 12,000 wetlands) by permanence class. These variables were proxies for climate (A-D), land cover and land use (E-H) and topography roughness (I-L). The gains illustrate the relative contribution of each variables in the model – the higher the value, the greater the importance.

460



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