- 1 Deposit_-feeding of <u>Nonionellina labradorica (</u>-foraminifera) from an
- 2 Arctic methane seep site and possible association with a
- 3 methanotroph revealed by transmission electron microscopy
- 4 Christiane Schmidt^{1,2,3}, Emmanuelle Geslin¹, Joan M Bernhard⁴, Charlotte LeKieffre^{1,5}, Mette
- 5 Marianne Svenning^{2,6}, Helene Roberge^{1,7}, Magali Schweizer¹, Giuliana Panieri²
- ¹LPG, Laboratoire de Planétologie et de GéodynamiqueGéosciences, Univ<u>of</u>. Angers, <u>Nantes</u> Université <u>de Nantes</u>,
 <u>Le Mans Univ</u> CNRS, LPG, SFR QUASAV, Angers, 49000, France
- ⁸ ²CAGE, Centre for Arctic Gas Hydrate, Environment and Climate, UiT, The Arctic University of Norway, Tromsø,
 ⁹ 9010, Norway
- 10 ³ZMT, Leibniz Centre for Tropical Marine Research, Bremen, 28359, Germany
- 11 ⁴Woods Hole Oceanographic Institution, Geology & Geophysics Department, Woods Hole, 02543, MA, USA
- ⁵Cell and Plant Physiology Laboratory, CNRS, CEA, INRAE, IRIG, Université Grenoble Alpes, Grenoble, 38054
 France

- 14 ⁶Department of Arctic and Marine Biology, UiT, The Arctic University of Norway, Tromsø, 9037, Norway
- 15 ⁷Université de Nantes, CNRS, Institut des Matériaux Jean Rouxel, IMN, Nantes, 44000 France
- 16
- 17 Correspondence to Christiane Schmidt christiane.schmidt@leibniz-zmt.de
- 18

20 Abstract. Several foraminifera are deposit feeders that consume organic detritus (-dead particulate 21 organic material along with entrained bacteria). However, the role of such foraminifera in the 22 benthic food-web remains understudied. FAs foraminifera may associate withfeeding on methanotrophic bacteria, which are ¹³C-depleted, feeding on them has been suggested tomay cause 23 negative cytoplasmic and/or calcitic δ^{13} C values δ^{13} C values in the foraminiferal cytoplasm and/or 24 25 calcite. To test whether the foraminiferal diet includes methanotrophs, we performed a short-term 26 (20-h1-d) feeding experiment with Nonionellina labradorica from an active Arctic methaneemission site (Storfjordrenna, Barents Sea) using the marine methanotroph Methyloprofundus 27 28 sedimenti, and analyzed N. labradorica cytology via Transmission Electron microscopy (TEM). 29 We hypothesized that M. sedimenti would be visible, post experiment in degradation vacuolesas 30 evidenced by their ultrastructure, in degradation vacuoles after this feeding experiment, as 31 evidenced by their ultrastructure. Sediment grains (mostly clay) occurred inside one or several degradation vacuoles in all foraminifers. In 24% of the specimens from the feeding experiment 32 33 degradation vacuoles also contained bacteria, although none could be confirmed to be the offered M. sedimenti. Observations of the area adjacent to the apertural areae after 20-h incubation 34 35 revealed three putative methanotrophs, close to clay particles, based on bacterial. These methanotrophs were identified based on internal-ultrastructural characteristics, such as a type I 36 37 stacked intracytoplasmic membranes (ICM), storage granules (SG) and gram negative cell walls 38 (GNCW): Furthermore, more, N. labradorica specimens were examined for specific adaptations to this active Arctic methane-emission site; we noted the absence of bacterial endobionts in all 39 specimens examined N. labradorica but confirmed the presence of kleptoplasts, which were often 40 partially degraded. Based on these observations In sum, we suggest that M. sedimenti can be 41 consumed by N. labradorica via untargeted grazing in seeps and that N. labradorica can be 42 43 generally classified as a deposit feeder at this Arctic site. These results suggest that if 44 methanothrophs are available to the foraminifera in their habitat, their non-selective uptake could make a substantial contribution to altering δ^{+3} C_{test} values. This in turn may impact metazoans 45 grazing on benthic foraminifera by altering their 8⁻¹³C signature. 46

47

48 benthic foraminifera – feeding experiment – grazing - marine methanotrophs – Arctic methane
49 seeps- transmission electron microscopy – ultrastructure – kleptoplasts- protist – molecular
50 identification

51 1. Introduction

52 In methane seep sites, the upward migration of methane affects the pore-water chemistry of nearsurface sediments, where benthic foraminifera inhabiting the sediment interface have been shown 53 to live (e.g. Dessandier et al., 2019). Extremely light isotopic signals of δ^{13} C have been measured 54 in seep-associated foraminiferal calcite tests (Wefer et al., 1994; Rathburn et al., 2003; Hill et al., 55 2004b; Panieri et al., 2014). Studies specifically looking at living (rose bengal stained) 56 57 for a miniferal tests support the hypothesis that the carbon isotopic composition is strongly 58 influenced by the porewater DIC (McCorkle et al., 1990a). Interspecific δ^{13} C differences between 59 species with similar depth indicate sometimes taxon-specific "vital" effects (McCorkle et al., 60 1990a). One Those "vital" effects describe the biology of the different species, which could reflect different feeding patterns. It has been suggested that Nonionella auris is an indicator of methane 61 release and possibly explanation of low $\delta^{13}C$ signals in foraminifera could be due to the 62 ingestioningests of ¹³C-depleted methanotrophse oxidizing bacteria —(Wefer et al., 1994). 63 Recently, specimens of the foraminifer Melonis barlegeanus (Williamson, 1858) collected from 64 an active methane seep site were was found to be closely associated with a putative methanotrophs 65 at their apertural regionreasoning the need to examine feeding habits of foraminifera living on or 66 around methane seeps (Bernhard and Panieri, 2018), providing impetus to examine feeding habits 67 68 of foraminifera living in or around methane seeps. The observation by Bernhard and Panieri (2018) brought to light the need to examine feeding 69 70 habits of foraminifera living on or around methane seeps. The species M. barleeanus could feed 71 on aerobic methane-oxidizing bacteria (methanothrophs), which are abundant in the water column 72 around methane seeps (Tavormina et al., 2010). Methanotrophs produce the biomarker diplopterol, which has an extremely light δ^{13} C signature (-60 ‰) and makes methanotrophs isotopically very 73 74 light themselves (Hinrichs et al., 2003). If-Our hypothesis is that if foraminifera accidentally or intentionally ingest methanotrophs, δ^{13} C values of foraminiferal cytoplasm should be altered by 75 such phagocytosistheir diet. However, experimental evidence was inconclusive whether isotope 76 labelling of food can influence foraminiferal calcite, as no new calcite was produced during 77 experiments using the foraminifera_Haynesina germanica and Ammonia beccari (Mojtahid et al., 78 79 2011). Experiments using a a novel high-pressure culturing system incubator on Cibicides wuellerstorfi illustratedrevealed the difficulty to measure the sensitive relationship between 80 81 methane exposure and the foraminifera Cibicides wuellerstorfi. However, it was shown in onethis

82	epineti gizelatunduesuevatieth Urfmiliklik Contica vaielatele Coleevatien pattiepostatuken per Cinelleposen daatoon haisen ba	
83	(Wollenburg et al., 2015). It is also not yet conclusive if diet can influence foraminiferal calcite,	
84	as new calcite did not form during experiments (Mojtahid et al., 2011).	
85	Several studies found that the lightest isotopic $\delta^{43}C$ values were measured in tests coated by	
86	methane-derived authigenic carbonate (MDAC) overgrowth (Torres et al., 2010; Panieri et al.,	
87	2014; Consolaro et al., 2015; Panieri et al., 2017; Schneider et al., 2017). MDACs represent a	
88	diagenetic alteration of the foraminiferal test that alters the δ^{13} C of the foraminiferal isotope record	
89	It can form high Mg calcite coatings contributing to the bulk of foraminiferal carbonate up to 58	
90	wt% MgCO (Schneider et al., 2017). MDACs are formed at the SMTZ, the sulfate-methane-	
91	transition zone (SMTZ), near the sediment-water interface where the upward flow of methane	
92	encounters the downward diffusion of sulfate from overlying seawater (Bian et al., 2001;	
93	Schneider et al., 2017).	
94	For aminifera play an important role in the carbon cycle on the deep seafloor (Nomaki et al., 2005)	
95	where feeding behavior and food preference vary with species (Nomaki et al., 2006). Selected	
96	species of deep-sea benthic foraminifera have been shown to feed selectively on ¹³ C-labeled algae	
97	from sedimentary organic matter, but unselectively on ¹³ C-labeled bacteria of the strain Vibrio	
98	(Nomaki et al., 2006). A study from the seafloor around Adriatic seeps suggested that $\delta^{13}C$ of	
99	foraminiferal cytoplasm could be influenced by feeding on the sulfur-oxidizing bacterium	
100	Beggiatoa, whose abundance was also positively correlated with foraminiferal densities Panieri,	
101	(Panieri, 2006). Generally, some foraminifera can ingest dissolved organic matter (DOM); some	
102	are herbivorous, carnivorous, suspension feeders and most commonly deposit feeders reviewerd	
103	in-(reviewed in Lipps, 1983). [Goldstein, 1994 #1903] Deposit feeders are omnivorous, gathering	
104	fine-grained sediment (e.g., clay) and associated bacteria, organic detritus (dead particulate organic	
105	material) and, if present, diatom cells using their pseudopodia. Hence, bacteria are involuntarily	
106	part of the "food-mix" (Levinton, 1989). Based on the ultrastructure of the diet found in vacuoles	
107	serveral species of foraminifera from different habitats have already been classified to be deposit	
108	feeders (Goldstein and Corliss, 1994).	
109	The fact that bacteria are sometimes part of the "food mix" made usHere we investigate if	
110	Nonionellina labradorica associated would feed in a short-term feeding experiment on the marine	
111	with-methanotrophs, e.g. Metyloprofundus sedimenti, in a short-term feeding experiment and	
112	compare its ultrastructure on experimental specimens and field specimens Nonionellina	

Formatted

114	with N. digitata in Svalbard fjord sediments (Hald and Korsun, 1997; Shetye et al., 2011; Fossile
115	et al., 2020). (Carrier et al., 2020) Next In addition to its wide distribution, it is an especially interesting experimental species.
116	for feeding studies because it hosts kleptoplasts, i.e. sequestered chloroplasts, of diatom origin
117	insideitscytoplasm(Cedingen, 1991; Jauffiaisetal, 2018) (Cedragen, 1991; Jauffiaisetal, 2019b <u>) SEMirrageset Noniorelling labradorica's</u> aperture
118	shows a specific ornamentation, possibly a morphological adaptation to this "predatory" mode of
119	life for obtaining the kleptoplasts (Bernhard and Bowser, 1999). Denitrification has been
120	speculated for <i>N. labradorica</i> (reviewed in Charrieau et al., 2019), <u>bBecause t</u> The <u>foraminiferal</u> genus
121	$Nonionella is potentially capable to can denitify, which was demonstrated on two the species { $
122	CINE AND CITE DATE AND A USE OF THE OWNER OF T
123	degradation vacuoles of this species from an active methane-emitting site in the Arctic
123 124	degradation vacuoles of this species from an active methane-emitting site in the Arctic (Storfjordrenna, Barents Sea) before and after a feeding experiment
124	(Storfjordrenna, Barents Sea) before and after a feeding experiment
124 125	(Storfjordrenna, Barents Sea) before and after a feeding experiment 2. Materials and methods
124 125 126	 (Storfjordrenna, Barents Sea) before and after a feeding experiment 2. Materials and methods 2.1. Site description and sampling living foraminifera
124 125 126 127	 (Storfjordrenna, Barents Sea) before and after a feeding experiment 2. Materials and methods 2.1. Site description and sampling living foraminifera The sampling site was located app. 50 km south of Svalbard at 382m water depth at the mouth of

the dia Abindra being a stranger of the second state of the second s

113

- 131 et al., 2017; Hong et al., 2018). GHP3 is a ~500-m diameter, 10-m tall mound that actively vents 132 methane (Fig. 1). Marine sediment samples were collected during CAGE cruise 18-05 supported 133 by the research vessel Kronprins Haakon ion in October 2018 and sampled from the seafloor by the Remotely Operated Vehicle (ROV) Ægir. A blade corer (surface dimensions 27 x 19 cm, Fig. 134 135 1c) was used to sample living foraminifera; it was placed directly in the vicinity of bacterial mats. The blade corer containing the sediment sample was opened immediately once onboard. A small 136
- 137 aquarium hose was used to sample the upper most surface layer (0-1 cm). The wet sediment was
- collected in petri dishes and wet sieved to a size range of 250-500 µm, which served as source of 138
- 139 living (cytoplasm containing) foraminifera. The species N. labradorica, which was the visibilly
- 140 abundant, was subsequently used for a feeding experiments described in detail below. A previous
- study on GHP1 in Storfjordrenna also showed also N. labradorica is also occurreding in other 141
- 142 sediment cores (MC_902 and MC_919) in the top 2 cm (Carrier et al., 2020).

Field Code Changed

Formatted: Font: Italic **Field Code Changed** Field Code Changed

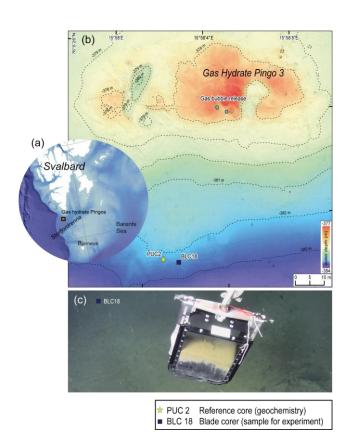


Figure 1. Description of the sampling site Gas Haydrate Pingo 3 (GHP3), a gas-hydrate bearing moundt, which actively vents methane, located in Storfjordrenna Barents Sea. (a) Map illustrating Svalbard Aarchipelago and the distance towards the sampling site_is app. 50 km offshore. (b) Map of sampling site GHP3, active gas bubble release is marked on the top of the underwater mount, yellow star indicates location of push corer PUC2 (taken for geochemical analyseis), black squared box indicates the location of the blade corer BLC18 (from which the sediment was derived source for the experiment). (c) Underwater image of retrieval of BLC18 taken by <u>ROV</u> camera of ROV (remotely operated vehicle) illustrating the coloration of sediment withand the sea-floor visible in background.

144 2.2. Geochemistry of the study site

145 For geochemical analysis of the study site a push corer (PUC2; henceforth) was used (referred to as geochemistry core) was taken to obtain measurements of $\frac{1}{2} \delta^{13}C_{DIC}$ and sulfate, because blade 146 corer (BLC18) did not allow those measurements. PUC2 was taken in close vicinity to BLC18, 147 148 ~5m apart (see Figure S1Fig-1). Pore-water samples were taken from PUC2 using rhizons that 149 were inserted through pre-drilled holes in the core tube at intervals of 1 cm (Table S1). Acid washed 20-ml syringes were attached to the rhizons for pore water collection. Depending on the 150 151 amount of pore water collected, the samples were split for $\delta^{13}C_{DIC}$ and sulfate measurements. To 152 the samples, 10 μ L of saturated HgCl₂ (aqueous) was added to stop microbial activity, and stored 153 in cold conditions (5°C). δ^{13} C_{DIC} was determined using a<u>A</u> ThermoScientific Gasbench II coupled 154 to a ThermoScientific MAT 253 IRMS at the Stable Isotope Laboratory (SIL) at CAGE, UIT was used to determine δ^{13} C_{DIC} of the pore-water. Anhydrous phosphoric acid was added to small glass 155 156 vials (volume 4.5 mL), that were closed and flushed with helium 5.0 gas before the pore--water 157 sub-sample was measured. A pore-water sub-sample (volume 0.5 mL) was then added through the 158 septa with a syringe <u>needle</u>, followed by equilibration for 24 h at 24°C to liberate the CO₂ gas. 159 Three solid calcite standards with a range of +2 to -49 % were used for normalization to $\delta^{13}C$ -160 VPDB. Correction of measured δ^{13} C by -0.1 ‰, was done to account for fractionation between (gas) and (aqueousar) in sample vials. Instrument precision for δ^{13} C on a MAT253 IRMS was $\frac{1}{3}$ 161 162 +/- 0.1 ‰ (SD). Sulfate was measured with a Metrohm ion chromatography instrument equipped with column Metrosep A sup 47 and eluted with 1.8 mmol/L Na₂CO₃ + 1.7 mmol/L NaHCO₃ at 163 164 the University of Bergen.

165 2.3. Culturing of the marine methanotroph M. sedimenti

Methyloprofundus sedimenti PKF-14 had been previously isolated from a water-column sample 166 collected at Prins Karls Forland, Svalbard in the laboratory at UiT in Tromsø. Methyloprofundus 167 sedimenti were cultured in 10-ml batches of a 35:65 mix of 1/10 Nitrate Mineral Salt medium 168 (NMS) and sterile filtered sea water using 125-mL Wheaton® serum bottles with butyl septa and 169 170 aluminum crimp caps (Teknolab[®]). Methane was injected to give a headspace of 20% methane in 171 air, and the bottles were incubated without shaking at 15°C in darkness. Purity of the cultures and 172 cell integrity was verified by microscopy and by absence of growth on agar plates with a general 173 medium for heterotrophic bacteria (tryptone, yeast extract, glucose and agar).

Transmission Electron Microscopy was performed on culture aliquots to allow morphological 174 comparison to previously published work (Tavormina et al., 2015). Methyloprofundus sedimenti 175 176 strain PKF 14 cells have a gram negative cell wall, coccoid to slightly elongated shape and characteristic stacked intracytoplasmic membrane (ISM) and storage granules (SG) (Fig 2c). 177 178 Additionally, 16S rRNA gene sequencing was performed (data not shown) to confirm it to be 179 similar to the published Methyloprofundus sedimenti (Tavormina et al., 2015). 180 On the ship, Nonionellina labradorica (Fig. 2a,b) specimens showing-a-dark greenish brown 181 cytoplasm were picked using sable artist brushes under a stereomicroscope immediately after wet

182 sieving the sediment using natural seawater delivered from the ship pump. Living specimens had

183 a partly inorganic covering surrounding the test, which was gently removed using fine artist

184 brushes. Those so-called cysts are nothing unusal with many foraminiferan taxa (Heinz et al.,

185 2005). Another Nonionellidae, Nonionella iridea, was similarly embedded with a cyst / covering
 186 in sediment

Our specimens were subsequently rinsed twice in filtered artificial seawater to remove any sediment before placing them into the experimental petri dishes. Care was taken that those were minimally exposed to light during preparation of the experiment, as kleptoplasts are known to be

190 highly light sensitive in this foraminifer (Jauffrais et al., 2019b).

191 The experiment with *M. sedimenti* was conducted for athe total duration of 20-h to resemble previous experiments on N. labradorica usingon transmission electron microscopy and 192 193 nanometre-scale secondary ion mass spectrometry isotopic imaging (TEM-NanoSIMS) (Jauffrais 194 et al., 2019b)u (Jauffrais et al., 2019), and included two more time points at 4 and 8 -h, where 195 incubations were terminated. A short pre-experimental phase (2-4 h) was included before the initial 196 start of the feeding experiment, to allow specimens to acclimate. During the pre-experimental 197 phase specimens were not fed and resided in the petri dishes to adjust to the experimental 198 conditions. The 20 h feeding experiment with M. sedimenti started after a short starvation phase 199 where organisms resided in petri dishes with ASW for 2-4 h and were not fed or manipulated 200 during this time. The feeding experiment consisted of several small petri dishes (3.5 cm \emptyset , 3 mL) 201 each containing five foraminifera-N. labradorica in ASW at ambient salinity 35 (Red Sea Salt). 202 Petri dishes were sealed with Parafilm[®] and covered with aluminum foil and placed inside the 203 incubator in complete darkness. Temperature inside the chamber was maintained at 2-3°C, which 204 is within the range of the site's bottom-water temperature $(-1.8 - 4.6^{\circ}C)$ (Hong et al., 2017). The feeding of *M. sedimenti* was performed once at the beginning of the experiment by adding 100 μL
of culture to 3 mL of artificial seawater to produce a final concentration of ~1E10⁶ bacteria / mL
in-<u>each petri dishthe experiment</u>. Previously conducted feeding studies were used as guides:
Muller and Lee (1969) used 1E10⁴ bacteria/mL seawater and Mojtahid et al. (2011) used 4E10⁸
bacteria/mL seawater.

210 Five foraminifera, which served as initial/field specimens (Table 1),_-were fixed without M.

211 sedimenti incubation. The respective petri dishes, were incubated for 4, 8 and 20 h to determine if

212 incubation duration influenced response of the foraminifera to the methanotroph. One petri dish

213 containing five foraminifera, which were un-fed and fixed at 20 h, served as a negative "control".

214 After the end of the respective incubation times, each foraminifer was picked with a sterilized fine

215 artist brush, which was cleaned in 70% ethanol between each specimen, and placed individually

216 into a fixative solution (4% glutaraldehyde and 2% paraformaldehyde dissolved in ASW).

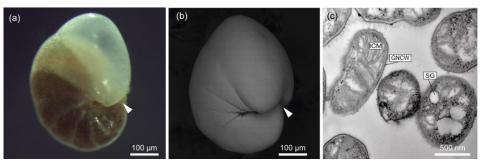


Figure 2 Exemplary illustration of *Nonionellina labradorica*, utilized in this study. (a) Reflected light microscopy image from a specimen directly after sampling, white arrowhead indicates aperture location. (b) Scanning electron image from a specimen before molecular analysis was performed, white arrowhead indicates aperture location. (c) Transmission electron microscopy image of a culture of *Metyloprofundus sedimenti*, the marine methanotroph used in the feeding experiment. The Characteristic features for methanotroph identification include is the typical type I ICM=intracytoplasmic membranes<u>s (ICM). Furthermore, other internal structures visible are</u>, SG=storage granules (SG), and <u>a GNCW=gram-negative cell wall (GNCW)</u>.

217 2.5. Transmission Electron microscopy (TEM) preparation

218 Samples of *N. labradorica* preserved in fixative solution were transported to the University of

219 Angers, where they were prepared for ultrastructural analysis using established protocols

220 (Lekieffre et al., 2018). Four embedded foraminiferal cells per treatment Embedded foraminiferal

221 cells-were sectioned using an ultramicrotome (Leica® Ultracut S) equipped with a diamond knife

(Diatome[®], ultra 45°). Grids were stained using UranyLess[®] EM Stain (EMS, USA). Ultra-thin
 sections (70 nm) were observed with a JEOL JEM-1400 TEM at the SCIAM facility, University
 of Angers.

- 225 To document the ultrastructure of Methyloprofundus sedimenti, a sub-sample of the culture used 226 for experiments was imaged with TEM (Fig. 2c). To do so, an exponentially growing culture was 227 collected, centrifuged, pre-fixed with 2.5 % (w/v) glutaraldehyde in growth medium overnight, 228 washed in PBS (Phosphate Buffered Saline), then post fixed with 1% (w/v) aqueous osmium 229 tetroxide for 1.-5 hours at room temperature. After dehydration in an ethanol series, the samples 230 were embedded in an Epon equivalent (Serva) epoxy resin. Ultra-thin sections were cut on a Leica 231 EM UC6 ultramicrotome, and stained with 3 % (w/v) aqueous uranyl acetate followed by staining 232 with lead citrate (Reynolds, 1963) at 20 °C for 4-5 min. The samples were examined with a JEOL 233 JEM-1010 transmission electron microscope at an accelerating voltage of 80 kV with a Morada 234 camera system at the Advanced Microscopy Core Facility (AMCF), Faculty of Health Science,
- 235 UiT The Arctic University of Norway.

236 2.6. Foraminifera ultrastructural observation and image processing

Four specimens per experimental time point (initials, 4, 8 and -20 h) plus one un-fed (control) 237 238 specimen were examined with the TEM. From each specimen, a minimum of 50 TEM images was 239 taken, including images detailing the degradation vacuoles (5-27 images of degradation vacuoles 240 per specimen). The ultrastructure was examined at different parts of the images sections focusing 241 (a) in the cell interior to document vitality, (b) on degradation vacuoles to determine their content, 242 and (c) at the exterior to survey for microbes entrained in remnant "reticulopodial trunk" material, 243 which can be extended outside foraminiferal tests during feeding and locomotion (Anderson and 244 Leef9)Ahays saklin televised of EMACAEAAHD Christel And Christel and a statistic televised and the statistic of the statistic of

245 2.7. Molecular genetics and morphology

DNA metabarcoding and morphological documentation were performed on 13 specimens of *N. labradorica.* Briefly, live specimens were dried on micropaleontological slides and transported in
a small container, cooled with ice-pads to the University of Angers. All specimens were imaged
for morphological analysis using a Scanning Electron Microscope (SEM; EVOLS10, ZEISS, Fig.
S1) followed by individually extracting total DNA in DOC buffer (Pawlowski, 2000). To amplify
foraminiferal DNA, a hot start PCR (2 min. at 95°C) was performed in a volume of 25µl with 40

cycles of 30 s at 95°C, 30 s at 50°C and 2 min at 72°C, followed by 10 min at 72°C for final 252 253 extension. Primers s14F3 and sB were used for the first PCR and 30 cycles at an annealing 254 temperature of 52°C (other parameters unchanged) for the nested PCR with primers s14F1 and J2 255 (Pawlowski, 2000; Darling et al., 2016). Positive amplifications were sequenced directly with the 256 Sanger method at Eurofins Genomics (Cologne, Germany). For taxonomic identification, DNA sequences were compared first with BLAST (Basic Local Alignment Search Tool) (Altschul et al., 257 258 1997) and then within an alignment comprising other Nonionids implemented in SeaView (Gouy 259 et al., 2010) and corrected manually.

260 **3. Results**

261 3.1. Sample description and geochemistry of the study site

262 The visual observation of the sediments within the blade corer BLC18 immediately after sampling 263 (Fig. 1c) indicated that the sediment appeareds light grey – yellowish in the upper part until app. 264 13 cm and dark brown from app. 13 cm to the bottom. TAt approximately 13 cm the sulfate 265 measured in the pore water of the geochemistry core (PUC2) declined from ~2750 ppm at the 266 sediment-water interface to ~706 ppm at approximately 13 cm (see Fig. S1, Table S1). A decline in sulfate concentration indicates that the anaerobic oxidation of methane (AOM) occurred at app. 267 13 cm depth. The SMTZ (Sulfate Methane Transition Zone) characterized by <u>a a reduced δ^{13} C</u> 268 269 DIC value of -32‰ at app. 13 cm sediment depth can be considered shallow on the global average 270 (Egger et al., 2018).

271 3.2. Ultrastructure of methanotroph culture used in the feeding experiment

272 Transmission Electron Microscopy was performed on culture aliquots to allow morphological

- 273 comparison to previously published work (Tavormina et al., 2015). Methyloprofundus sedimenti
- strain PKF-14 cells are coccoid to slightly elongated shape and is characterized by typical type I
- 275 <u>stdatacythnimentmet(M)fi23huMshprinlsahufidnatiadagiuhpilistalIntdahunhm(M)staggants</u>Contsgingam
- 276 negative cell wall (GNCW), which are not uniquely charactersitic of methanotrophs (GNCW) (Fig. 2c).
- 277 Additionally, 16S rRNA gene sequencing was performed (data not shown) to confirm it to be
- 278 similar to the published Methyloprofindus sedimenti (Tavormina et al., 2015). These features were used to identify M. sedimenti,

Formatted

279 3.3. Foraminiferal ultrastructure from an Arctic seep environment

280 3.3.1 General ultrastructure

281 All 17 specimens examined for ultrastructure were considered living at the time of observation 282 (Fig. 3), as the mitochondria had characteristic double membranes and occasionally visible cristae 283 (Nomaki et al., 2016). Cytoplasm exhibited several vacuoles and kleptoplasts concentrated in the 284 youngest chambers (Fig. 3a) and, in some specimens, the a nucleus with nucleoli was visible (Fig. 285 3b). Kleptoplasts were numerous throughout the cytoplasm and occurred in the form of a single 286 chloroplast (Fig. 3a-b), or as double chloroplasts (Fig. S2a-d). Not all kleptoplasts were intacting 287 some showed peripheral degradation of the membranes indicated by an increasing number of white 288 areas between pyrenoid, lamella and thylakoids (Fig. S2a-d). The mitochondria occurred often in 289 small clusters of two to five throughout the cytoplasm and were oval, round or kidney-shaped in 290 cross section (Fig. 3e-f). Peroxisomes in N. labradorica occurred mostly as pairs (Fig. 3c) or small 291 clusters of 3-4 spherical organelles (Fig. S3a1a b). The mitochondria occurred often in small 292 elusters of two to five throughout the cytoplasm and were oval, round or kidney shaped in cross 293 section (Fig. 3c f).-Sometimes, but not always, peroxisomes were associated with endoplasmic 294 reticulum (Fig. S34be) but could also occur alone. Golgi apparatus (Fig 3d) had intact membranes,

295 often occurring near mitochondria.

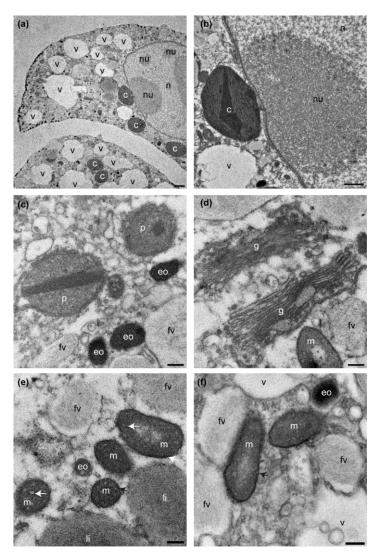


Figure 3 Transmission electron micrographs showing cellular ultrastructure of *N. <u>Labradorica</u>.* (a) Cytoplasm showing parts of two chambers, with nucleus with nucleoli, vacuoles and several kleptoplasts, (b) nuclear envelope, nucleoli, and kleptoplasts, (c) peroxisomes and electron opaque bodies, (d) Golgi, (e-f) mitochondria. V=vacuole, c=kleptoplast, nu=nucleoli, n=nucleus p=peroxisome, eo=electron opaque body, m=mitochondrion, fv=fibrillar vesicle, li=lipid droplet. Scales: (a) 2 µm, (b) 1 µm, (c-f) 200 nm

297 3.3.2 Ultrastructure of aperture-associated bacteria

298 In total, three putative methanotrophs were identified in the vicinity of two foraminifer specimens 299 (sample E39, Fig. 4; E37, Fig. 5). Theose microbes were identified adjacent next to reticulopodial 300 remains in the cross section (Fig. 4b). As an aid for identification of M. sedimenti we used the 301 characteristics shown in the literature (Tavormina (Tavormina et al., 2015) et al. 2015) and a our 302 own TEM observation obtained from M. sedimenti culture (Fig. 2c). As noted, Methyloprofundus 303 sedimenti is characterized by a typical type I intracellular stacked-ytoplasmic stacked membrane 304 (ISSM). Other characteristics, which athatare not specific for methanotrophs, were included storage 305 granules (SG) and a typical gram-negative cell wall (GNCW) (Fig. 2c). On specimen E39 from 306 the 20 h treatment, we found the methanotroph exhibiting the clearest internal structure, having 307 both typical type I stacked intracytoplasmic stacked membranes (ISCM) and +SG) and a second putative methanotroph showing SG+GNCW_(Fig. 4c). Specimen E36, from the 20 h treatment, 308

309 hosted another putative methanotroph showing three large SG (Fig. 5). Storage granules occur

310 throught this putative methanotroph (Fig. 5c).

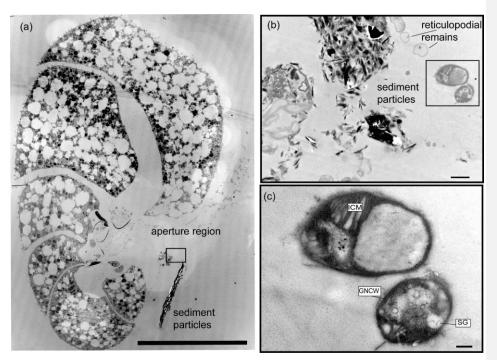


Figure 4 Transmission electron micrographs of *N. labradorica* from 20 h treatment (sample E39) (a) Stitched cross section of TEM images showing location of methanotroph at the aperture region (black <u>rectangle rectangle is the location of image shown in panel b</u>) (b) Location of two putative methanotrophs next to sediment particles and putative reticulopodial remains <u>(black rectangle is location of image shown in panel c)</u> -(c) Close up of two putative methanotrophs revealing detailed feature for identification, such as <u>typical type I stacked intracytoplasmic stacked</u> membranes (ICSM), <u>and other characteristics, such as</u> storage granules (SG), and gram-negative cell wall (GNCW), scale bars: a: 100 µm, b: 1 µm, c: 200 nm.

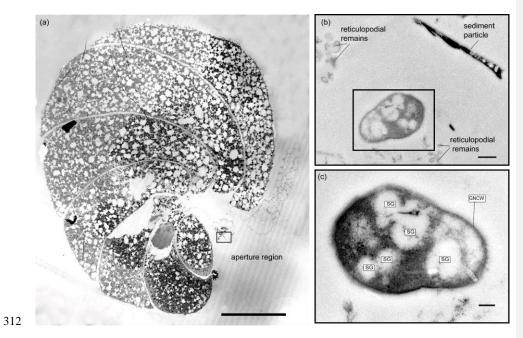


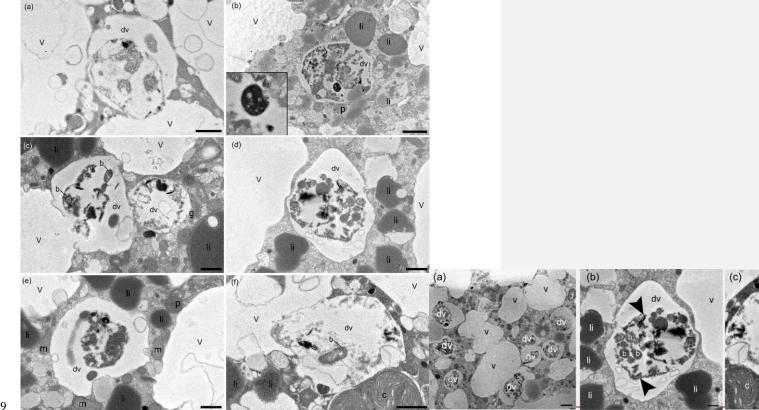
Figure 5 Transmission electron micrographs of *N. labradorica* from 20 h treatment (sample E37) (a) Stitched cross section of TEM images showing location of putative methanotroph (black rectangle) at the aperture region. (b) Location of the putative methanotroph next to sediment particles and sections of the putative reticulopodial remains (c) Close up of putative methanotroph showing several SG throughout its cell, scale bars: a: 100 μ m, b: 0.5 μ m, c: 200 nm.

313

314 3.3.3 Contents of degradation vacuoles

315	Digestive vacuoles and food vacuoles are often summarized as degradation vacuoles in the
316	literature (Lekieffre et al., 2018) and this makes sense for our study as well. A degradation vacuole
317	is a vacuole where enzymatic activities degrade contents, often making them unidentifiable (Bé et
318	al., 1982; Hemleben et al., 2012). Sediment particles were present in many degradation vacuoles.
319	The sediment grains wereare easy to recognize in the TEM image as angular grains spiking out of
320	inside the vacuoles, next to organic debris, which can have many different shapes. Each specimen
321	had at least one degradation vacuole and mostly several, which were degradation vacuole filled
322	with sediment particles present (Table 1). If a sediment particle was visible, the vacuole was
323	defined as a degradation vacuole (dv), and if it was not and empty then it was defined as a standard
324	vacuole (v) (Fig. 6). The observed entrained sSediment particles were platelets, are likely the
325	remains of clay-grains from the seafloor, and hence show that the vacuole must contain eell-foreign
i.	

- 326 objects, around which degradation processes have started. Next to sediment particles, Four4 out of
- 327 17 specimens examined (23%) had one or morea few bacteria of various sizes inside their degradation
- 328 vacuoles <u>next to sediment particles</u> (Fig 6 <u>c, fb-e</u>).



329

Figure 6 TEM micrographs of *N. labradorica* showing degradation vacuoles containing miscellaneous items, including bacteria (b), inorganics (clav platelets) and unidentifiable remains after 14 incubation, which are shown enlarged in the left side of the image in a zoom window (a,b; specimens E27, E28, respectively); after 8h incubation (c,d; specimen E14), after 20h incubation (e,f; specimens E36, E37, respectively). v=vacuole, dv=degradation vacuole, c=kleptoplast, p=peroxisome, m=mitochondrion, li=lipid, g= Golgi. Scales: (a, c-f) 1 μm, (b) 2 μm.TEM micrographs of N. labradoriea. (a) Overview of degradation vacuoles (dv) in relation to empty vacuoles (v) in the youngest chambers of specimen E5 (field). (b) Bacteria in degradation vacuoles (white b) next to clay particles (black arrow) in specimen E14 (8 h incubation), (c) Elongated bacterium inside degradation vacuole adjacent to clay particles of specimen E37 (20 h incubation), scale bars: a: 2 μm, b,e: 0.5 μm.

330 3.4. Foraminiferal genetics

331 Six of 13 specimens analyzed for genetics were positively amplified and sequenced (Fig. S_{43}^{43}). 332 The sequences are deposited in GenBank under the accession numbers MN514777 to MN514782. 333 When comparing them via BLAST, they were between 98.6% and 99.6% identical to published 334 sequences belonging to foraminifera identified as the morphospecies N. labradorica, from the 335 Skagerrak, Svalbard and the White Sea (Holzmann and Pawlowski, 2017; Jauffrais et al., 2019b). 336 Sequences were also included in an alignment comprising other nonionids implemented in 337 Seaview (not shown) and corrected manually to check the BLAST search. This step confirmed the 338 BLAST identification.

339 4. Discussion

340 4.1. Sampling site and geochemistry

341 The sampling site of blade corer BLC18 was in close proximity (~50 m) to an active methane-vent 342 releasing methane bubbles at the gas hydrate pPingo (GHP3) (Serov et al., 2017). At such sites 343 with high methane fluxes, the SMTZ (sulfate methane transition zone) is shallow, as sulfate infrom 344 the sediment is readily consumed in the first tens of <u>centimeters</u> (Barnes and Goldberg, 1976; 345 Iversen and Jørgensen, 1993) by sulfate-reducing bacteria (SRB) (reviewed in Carrier et al., 2020). 346 Geochemical analysis of PUC2, revealed an SMTZ at app. 13 cm, which. The depth of 13 cm is 347 rather shallow (Egger et al., 2018), as it can also be several meters deep in other sites (reviewed in 348 Panieri et al., 2017). Similar gGeochemical characteristics can be considered similar at the 349 sampling location of living specimens (BLC18) given the close proximity of the two locationsand the core taken for geochemistry (PUC2). The geochemical data at PUC2 allows us conclude that 350 351 the site, where living foraminifera were collected, can be classified as an active methane emission 352 site.

353

354 4.2. <u>Possible aAssociationssociation</u> with putative methanotrophs

355 The possible association of *N. labradorica* with the three putative-methanotrophs could bewas

356 identified documented via presence of two putative methanothrophs, based on microbial

357 <u>ultrastructure for</u>on two foraminifera specimens based on comparing internal bacterial

358 characteristics to published literature (Tavormina et al., 2015)... Transmission electron microscopy

Formatted: Font: Italic

359	i sporefile local latur dite a til dh'an i lidyplan Ilabumation fi <u>podi</u> zsaiowip tiven hantqi <mark>lskiyoji i i gidto i a tefan</mark> I
360	
361	field-remains, to preserve its microbiom Another benthic for a minifer, Melonis barleeanus, has been noted to have clumps of
362	putative methanotrophs at the apertural opening of field-collected specimens Ourare similar to field collected (Bernhard and
363	<u>Print NSvelvenierikeltentetentetentetentetentetentetentete</u>
364	described in this study for this species for the first time, shows that methanotrophs may be are ingested via
365	untargeted grazing-in seeps, as N. labradorica appears to be a non-selective feeder.
366	4.3. Feeding on other bacteria and contents of Ddegradation vacuoles show large number of
367	sediment particles and few bacteria
368	Our results of the feeding experiment and experimental specimens-show that only-23% of the
369	examined N. labradorica specimens contained bacteria inside their degradation vacuoles. That is
370	not a large quantity-proportion compared to presence of sediment particles, which occurred in
371	100% of the examined degradation vacuoles for a minifers. From this result, however, we We infer
372	that N. labradorica at this site is a deposit feeder, feeding on organic detritus and associated
373	bacteria. The bacteria observed in the degradation vacuoles resembled those from other deep-sea
374	foraminifera (Globobulimina pacifica and Uvigerina peregrina) and the shallow-dwelling genus
375	Ammonia (Goldstein and Corliss, 1994). Salt-marsh foraminifera also feed on bacteria and
376	detritus, as observed in TEM studies (Frail-Gauthier et al., 2019). Scavenging on bacteria has also
377	been observed by other foraminifera from intertidal environments such as Ammonia tepida or
378	Haynesina germanica (Pascal et al., 2008) and is a logical consequence from detritus feeding.
379	Certain foraminifera have been shown to selectively ingest algae/bacteria according to strain (Lee
380	et al., 1966; Lee and Muller, 1973). From laboratory cultures we know that several foraminifera
381	cultures require bacteria to reproduce, as antibiotics inhibited reproduction (Muller and Lee, 1969).
382	Future studies will need to employ additionally molecular tools to additionally determine the food
383	contents inside the cytoplasm (e.g. (e.g. Salonen et al., 2019). For example, aA recent study-by
384	used metabarcoding to assess the contribution of bacterial - <u>eukaryotic</u> OTUs associated with
385	intertidal foraminifera,-and revealinged that Ammonia sp. T6 preys on metazoans, can predate_on
386	metazoan taxa, whereas Elphidium sp. S5 and Haynesina sp. S16 wereare more likely to ingest
387	diatom <u>s</u> a (Chronopoulou et al., 2019).
1	

4.4. General ultrastructure of *N. labradorica* collected in a seep environment

389 Our observations also included the intact nature of all major organelle typess of thise species, as this 390 was essential to conclude vitality after the experiment (Nomaki et al., 2016). Mitochondria and 391 kleptoplasts were generally homogeneously distributed throughout the cytoplasm confirming 392 previous observations of six N. labradoricia_-from the Gullmar Fjord (Lekieffre et al., 2018; 393 Jauffrais et al., 2019b). If mitochondria are concentrated predominately under pore plugs, it can 394 be an indicator that the electron acceptor oxygen is scarce in their environment, as the pores are 395 the direct connection from the cell to the environment. This has been observed in several other 396 studies where mitochondria were accumulated under pores in N. stella (Leutenegger and Hansen,

397 1979) and *Bolivina pacifica* (Bernhard et al., 2010).

398 For the specimenssamples from our particular site, we also observed kleptoplasts abundantly and

399 evenly distributed throughout the cytoplasm, confirming previous TEM studies on the species from 400 fjord sediments (Cedhagen, 1991; Jauffrais et al., 2018). Occasionally, Even though our study did 401 not focus on kleptoplasts, we could observe that kleptoplasts -were occasionally degraded, which 402 could have happened; a) during sampling, b) due to exposure to microscope lights or c) due to the 403 age and condition of kleptoplasts inside the host. Kleptoplasts in N. labradorica have been studied 404 in detail describing their diatom origin_(Cedhagen, 1991), sensitivity to light and missing 405 photosynthetic functionality (e.g. (Jauffrais et al., 2019b). (Jauffrais et al., 2019a)It has been 406 suggested that kleptoplasts could function as a seasonal energy reservoir, for example, (e.g. in 407 winter) (Jauffrais et al., 2016).

Formatted: Font: Italic,

408 5. Conclusions

409 Based on the content of degradation vacuoles-observed, we conclude that N. labradorica from 410 GHP3 our study site, an active methane emmitting site in the Barents Sea, -is a deposit---feeder.-411 Lit ingests large amounts of sediment particles together with bacteria-as part of consuming detritus 412 detrivorous diet living on the sea floor. On two specimens of the feeding experiment, putative 413 methanotrophs were observed near the N. labradorica aperture, suggesting ingestion of M. sedimenti-At the aperture region of two different foraminifera specimens, next to reticulopodial 414 415 remains and sediment particles, we observed three putative marine methanotrophs after 20 h 416 incubation. One of the putative methanotrophs had characteristic ISM, which resemble the methanotroph M. sedimenti in culture. We conclude that it is possible that N. labradorica may 417 418 ingests M. sedimenti via "untargeted grazing"-in this seep sites. Further studies are needed on

419 feeding strategies of several-other paleo-oceanographically relevant foraminifera to detangle the

420 relationship between δ¹³C measured in of foraminiferal calcite, their cytoplasm and dietary compositioneon tribution to their diet.

421 6. Data availability

- 422 Data in form of TEM images will be deposited at PANGAEA (-under doi:
- 423 Molecular data will be is deposited before publication at Genbank.

424 7. Sample availability

425 Samples are available upon request and TEM thinsections archived at the University of Angers.

426 8. Acknowledgments

We thank the captains, crew members and scientists onboard R/V Kronprins Haakon and ROV 427 Ægir Team for their assistance; Anne-Grethe Hestnes for growing the methanotroph culture. 428 429 Florence Manero, Romain Mallet and Rodolphe Perrot at the SCIAM microscopy facility 430 University of Angers are to thank for their expertise with the TEM and SEM. We thank Sunil 431 Vadakkepuliyambatta for helping to prepare the map presented in Figure 1; Sophie Quinchard 432 (LPG-BIAF) for supporting the molecular analysis. Funding was received through the Research 433 Council of Norway, CAGE (Center for Excellence in Arctic Gas Hydrate Environment and 434 Climate, project number 223259) and NORCRUST (project number 255150) to GP, EG, and CS. 435 CS position was funded through the MOPGA (Make Our Planet Great Again) fellowship by 436 CAMPUS France, the NORCRUST project and the University of Angers. JMB was partially 437 supported by US NSF 1634469, WHOI's Investment in Science Program, and by the Région Pays de la Loire through the FRESCO Project. 438

439 Author Contributions

Designed the project and experiment: GP, EG, CS; Collected samples: CS, EG; Performed
experiment: CS; Sample preparation: CS, HR; TEM observations and interpretations: CS, JMB,
EG, CL; Conducted molecular genetics: MSc; Wrote the paper: CS, GP, JMB; Provided critical
review and edits to the manuscript: EG, CL, MSv, MSc, HR; Contributed
reagents/materials/analysis tools: MSv, MSc, CL.

Table I. Summary of TEM observations of Nonionellina labradorica comparing field specimens

and experimental specimens. Field specimens (initials) were not fed, nor was a non-fed control

preserved after a 20 h incubation. The only putative methanotrophs were observed and imaged in

specimens from the 20 h incubation. Bacteria of unknown origin were described as rod shaped

cells in the degradation vacuoles.

Duration of experiment (h)/field samples	Food provided (yes (x)/no)	Sample ID	Cytoplasm: Degradation vacuole Contents		Aperture region: (putative) Methanotrophs
•			bacteria	Clay/in- organics	
Field	No	E1	no	X	no
samples	No	E3	no	х	no
(Initials)	No	E5	no	х	no
	No	E6	no	х	no
4	Х	E25	no	х	no
	х	E27	х	х	no
	х	E28	no	х	no
	х	E29	no	х	no
8	х	E14	х	х	no
	х	E15	no	х	no
	х	E16	no	Х	no
	х	E17	no	Х	no
20	х	E36	х	Х	1 x
	х	E37	х	Х	no
	х	E38	no	Х	no
	х	E39	no	Х	2 x
Control (20)	no	E44	no	Х	no

455 **References-:**

- 457 Altschul, S. F., Madden, T. L., Schäffer, A. A., Zhang, J., Zhang, Z., Miller, W., and Lipman, D.
- J.: Gapped BLAST and PSI-BLAST: a new generation of protein database search programs, 458
- Nucleic Acids Res., 25, 3389-3402, https://doi.org/10.1093/nar/25.17.3389, 1997. 459
- 460 Barnes, R. O. and Goldberg, E. D.: Methane production and consumption in anoxic marine
- sediments, Geology, 4, 297-300, https://doi.org/10.1130/0091-461
- 7613(1976)4<297:MPACIA>2.0.CO;2, 1976. 462
- Bé, A. W. H., Spero, H. J., and Anderson, O. R.: Effects of symbiont elimination and reinfection 463
- on the life processes of the planktonic foraminifer Globigerinoides sacculifer, Marine Biology, 464 70, 73-86, https://doi.org/10.1007/BF00397298, 1982. 465
- Bernhard, J. M. and Bowser, S. S.: Benthic foraminifera of dysoxic sediments: chloroplast 466
- sequestration and functional morphology, Earth-Sci. Rev., 46, 149-165, 467
- https://doi.org/10.1016/s0012-8252(99)00017-3, 1999. 468
- 469 Bernhard, J. M. and Panieri, G.: Keystone Arctic paleoceanographic proxy association with
- putative methanotrophic bacteria, Sci Rep-Uk, 8, 10610, https://doi.org/10.1038/s41598-018-470 28871-3, 2018. 471
- 472 Bernhard, J. M., Goldstein, S. T., and Bowser, S. S.: An ectobiont-bearing foraminiferan,
- 473 Bolivina pacifica, that inhabits microxic pore waters: cell-biological and paleoceanographic 474
- insights, Environmental Microbiology, 12, 2107-2119, 10.1111/j.1462-2920.2009.02073.x, 475 2010.
- 476 Carrier, V., Svenning, M. M., Gründger, F., Niemann, H., Dessandier, P.-A., Panieri, G., and
- 477 Kalenitchenko, D.: The Impact of Methane on Microbial Communities at Marine Arctic Gas 478 Hydrate Bearing Sediment, Frontiers in Microbiology, 11, 10.3389/fmicb.2020.01932, 2020.
- 479 Cedhagen, T.: Retention of chloroplasts and bathymetric distribution in the sublittoral
- 480 foraminiferan Nonionellina labradorica, Ophelia, 33, 17-30,
- https://doi.org/10.1080/00785326.1991.10429739, 1991. 481
- 482 Charrieau, L. M., Ljung, K., Schenk, F., Daewel, U., Kritzberg, E., and Filipsson, H. L.: Rapid
- 483 environmental responses to climate-induced hydrographic changes in the Baltic Sea entrance, 484
- Biogeosciences, 16, 3835-3852, 10.5194/bg-16-3835-2019, 2019.
- 485 Choquel, C., Geslin, E., Metzger, E., Filipsson, H. L., Risgaard-Petersen, N., Launeau, P.,
- 486 Giraud, M., Jauffrais, T., Jesus, B., and Mouret, A.: Denitrification by benthic foraminifera and
- 487 their contribution to N-loss from a fjord environment, Biogeosciences, 18, 327-341, 10.5194/bg-
- 488 18-327-2021, 2021.

- 489 Chronopoulou, P.-M., Salonen, I., Bird, C., Reichart, G.-J., and Koho, K. A.: Metabarcoding
- 490 insights into the trophic behavior and identity of intertidal benthic foraminifera, Frontiers in
- 491 microbiology, 10, 1169, <u>https://doi.org/10.3389/fmicb.2019.01169</u>, 2019.
- 492 Consolaro, C., Rasmussen, T., Panieri, G., Mienert, J., Bünz, S., and Sztybor, K.: Carbon isotope
- 493 (δ 13C) excursions suggest times of major methane release during the last 14 kyr in Fram Strait,
- 494 the deep-water gateway to the Arctic, Clim. Past, 11, 669-685, <u>https://doi.org/10.5194/cp-11-</u>
- 495 <u>669-2015</u>, 2015.
- 496 Darling, K. F., Schweizer, M., Knudsen, K. L., Evans, K. M., Bird, C., Roberts, A., Filipsson, H.
- 497 L., Kim, J.-H., Gudmundsson, G., Wade, C. M., Sayer, M. D. J., and Austin, W. E. N.: The
- 498 genetic diversity, phylogeography and morphology of Elphidiidae (Foraminifera) in the
- 499 Northeast Atlantic, Mar. Micropaleontol., 129, 1-23,
- 500 <u>https://doi.org/10.1016/j.marmicro.2016.09.001</u>, 2016.
- 501 Dessandier, P.-A., Borrelli, C., Kalenitchenko, D., and Panieri, G.: Benthic Foraminifera in
- 502 Arctic Methane Hydrate Bearing Sediments, Frontiers in Marine Science, 6,
- 503 https://doi.org/10.3389/fmars.2019.00765, 2019.
- 504 Egger, M., Riedinger, N., Mogollón, J. M., and Jørgensen, B. B.: Global diffusive fluxes of
- 505 methane in marine sediments, Nature Geoscience, 11, 421-425, 10.1038/s41561-018-0122-8, 506 2018.
- 507 Fossile, E., Nardelli, M. P., Jouini, A., Lansard, B., Pusceddu, A., Moccia, D., Michel, E., Péron,
- 508 O., Howa, H., and Mojtahid, M.: Benthic foraminifera as tracers of brine production in
- 509 Storfjorden "sea ice factory", Biogeosciences, 17, <u>https://doi.org/10.5194/bg-17-1933-2020</u>, 510 2020.
- 511 Frail-Gauthier, J. L., Mudie, P. J., Simpson, A. G. B., and Scott, D. B.: Mesocosm and
- Microcosm Experiments On the Feeding of Temperate Salt Marsh Foraminifera, J. Foraminifer.
 Res., 49, 259-274, https://doi.org/10.2113/gsjfr.49.3.259, 2019.
- 514 Goldstein, S. T. and Corliss, B. H.: Deposit feeding in selected deep-sea and shallow-water
- benthic foraminifera, Deep Sea Research Part I: Oceanographic Research Papers, 41, 229-241,
 https://doi.org/10.1016/0967-0637(94)90001-9, 1994.
- 517 Gouy, M., Guindon, S., and Gascuel, O.: SeaView version 4: a multiplatform graphical user
- interface for sequence alignment and phylogenetic tree building, Mol. Biol. Evol., 27, 221-224,
 https://doi.org/10.1093/molbev/msp259, 2010.
- 520 Hald, M. and Korsun, S.: Distribution of modern benthic foraminifera from fjords of Svalbard,
- 521 European Arctic, The Journal of Foraminiferal Research, 27, 101-122,
- 522 https://doi.org/10.2113/gsjfr.27.2.101, 1997.

- 523 Heinz, P., Geslin, E., and Hemleben, C.: Laboratory observations of benthic foraminiferal cysts, Mar. Biol. Res., 1, 149-159, 2005. 524
- 525 Hemleben, C., Spindler, M., and Anderson, O. R.: Modern planktonic foraminifera, Springer 526 Science & Business Media2012.
- 527 Herguera, J. C., Paull, C. K., Perez, E., Ussler Iii, W., and Peltzer, E.: Limits to the sensitivity of
- 528 living benthic foraminifera to pore water carbon isotope anomalies in methane vent
- 529 environments, Paleoceanography, 29, 273-289, https://doi.org/10.1002/2013PA002457, 2014.
- 530 Hill, R., Schreiber, U., Gademann, R., Larkum, A. W. D., Kuhl, M., and Ralph, P. J.: Spatial
- heterogeneity of photosynthesis and the effect of temperature-induced bleaching conditions in 531
- three species of corals, Marine Biology, 144, 633-640, https://doi.org/10.1007/s00227-003-1226-532 <u>1</u>, 2004a. 533
- 534 Hill, T. M., Kennett, J. P., and Valentine, D. L.: Isotopic evidence for the incorporation of
- 535 methane-derived carbon into foraminifera from modern methane seeps, Hydrate Ridge,
- Northeast Pacific, Geochimica et Cosmochimica Acta, 68, 4619-4627, 536
- https://doi.org/10.1016/j.gca.2004.07.012, 2004b. 537
- Hinrichs, K.-U., Hmelo, L. R., and Sylva, S. P.: Molecular fossil record of elevated methane 538
- levels in late Pleistocene coastal waters, Science, 299, 1214-1217, 539
- 540 https://doi.org/10.1126/science.1079601, 2003.
- 541 Holzmann, M. and Pawlowski, J.: An updated classification of rotaliid foraminifera based on
- ribosomal DNA phylogeny, Mar. Micropaleontol., 132, 18-34, 542
- https://doi.org/10.1016/j.marmicro.2017.04.002, 2017. 543
- Hong, W.-L., Torres, M. E., Carroll, J., Crémière, A., Panieri, G., Yao, H., and Serov, P.: 544
- Seepage from an arctic shallow marine gas hydrate reservoir is insensitive to momentary ocean 545 warming, Nat. Commun., 8, 15745, https://doi.org/10.1038/ncomms15745, 2017. 546
- 547 Hong, W. L., Torres, M. E., Portnov, A., Waage, M., Haley, B., and Lepland, A.: Variations in
- 548 gas and water pulses at an Arctic seep: fluid sources and methane transport, Geophys. Res. Lett., 45, 4153-4162, https://doi.org/10.1029/2018GL077309, 2018. 549
- 550 Iversen, N. and Jørgensen, B. B.: Diffusion coefficients of sulfate and methane in marine
- sediments: Influence of porosity, Geochimica et Cosmochimica Acta, 57, 571-578, 551
- https://doi.org/10.1016/0016-7037(93)90368-7, 1993. 552
- Jauffrais, T., LeKieffre, C., Schweizer, M., Jesus, B., Metzger, E., and Geslin, E.: Response of a 553
- kleptoplastidic foraminifer to heterotrophic starvation: photosynthesis and lipid droplet 554
- biogenesis, FEMS Microbiol. Ecol., 95, 10.1093/femsec/fiz046, 2019a. 555

- 556 Jauffrais, T., LeKieffre, C., Schweizer, M., Geslin, E., Metzger, E., Bernhard, J. M., Jesus, B.,
- Filipsson, H. L., Maire, O., and Meibom, A.: Kleptoplastidic benthic foraminifera from aphotic 557
- habitats: insights into assimilation of inorganic C, N and S studied with sub-cellular resolution, 558 Environmental microbiology, 21, 125-141, https://doi.org/10.1111/1462-2920.14433, 2019b.
- 559
- Lee, J. J. and Muller, W. A.: Trophic dynamics and niches of salt marsh foraminifera, Am. Zool., 560 13, 215-223, 1973. 561
- 562 Lee, J. J., McEnery, M., Pierce, S., Freudenthal, H., and Muller, W.: Tracer experiments in feeding littoral foraminifera, The Journal of Protozoology, 13, 659-670, 1966. 563
- 564 LeKieffre, C., Bernhard, J. M., Mabilleau, G., Filipsson, H. L., Meibom, A., and Geslin, E.: An
- 565 overview of cellular ultrastructure in benthic foraminifera: New observations of rotalid species in 566 the context of existing literature, Mar. Micropaleontol., 138, 12-32,
- 567 https://doi.org/10.1016/j.marmicro.2017.10.005, 2018.
- Leutenegger, S. and Hansen, H. J.: Ultrastructural and radiotracer studies of pore function in 568 foraminifera, Marine Biology, 54, 11-16, 10.1007/BF00387046, 1979. 569
- 570 Lipps, J. H.: Biotic Interactions in Benthic Foraminifera, in: Biotic Interactions in Recent and
- Fossil Benthic Communities, edited by: Tevesz, M. J. S., and McCall, P. L., Springer US, 571
- Boston, MA, 331-376, 10.1007/978-1-4757-0740-3_8, 1983. 572
- Mackensen, A.: On the use of benthic foraminiferal $\delta 13C$ in palaeoceanography: constraints 573
- from primary proxy relationships, Geological Society, London, Special Publications, 303, 121-574 133, https://doi.org/10.1144/SP303.9, 2008. 575
- 576 Mojtahid, M., Zubkov, M. V., Hartmann, M., and Gooday, A. J.: Grazing of intertidal benthic
- foraminifera on bacteria: Assessment using pulse-chase radiotracing, J. Exp. Mar. Biol. Ecol., 577 399, 25-34, https://doi.org/10.1016/j.jembe.2011.01.011, 2011. 578
- Muller, W. A. and Lee, J. J.: Apparent Indispensability of Bacteria in Foraminiferan Nutrition, 579
- The Journal of Protozoology, 16, 471-478, https://doi.org/10.1111/j.1550-7408.1969.tb02303.x, 580 581 1969.
- 582 Nomaki, H., Heinz, P., Nakatsuka, T., Shimanaga, M., and Kitazato, H.: Species-specific
- ingestion of organic carbon by deep-sea benthic foraminifera and meiobenthos: In situ tracer 583
- experiments, Limnol. Oceanogr., 50, 134-146, https://doi.org/10.4319/lo.2005.50.1.0134, 2005. 584
- 585 Nomaki, H., Heinz, P., Nakatsuka, T., Shimanaga, M., Ohkouchi, N., Ogawa, N. O., Kogure, K.,
- Ikemoto, E., and Kitazato, H.: Different ingestion patterns of C-13-labeled bacteria and algae by 586 deep-sea benthic foraminifera, Marine Ecology-Progress Series, 310, 95-108, 587
- https://doi.org/10.3354/meps310095, 2006. 588

- 589 Nomaki, H., Bernhard, J. M., Ishida, A., Tsuchiya, M., Uematsu, K., Tame, A., Kitahashi, T.,
- 590 Takahata, N., Sano, Y., and Toyofuku, T.: Intracellular Isotope Localization in Ammonia sp.
- 591 (Foraminifera) of Oxygen-Depleted Environments: Results of Nitrate and Sulfate Labeling
- 592 Experiments, Frontiers in Microbiology, 7, https://doi.org/10.3389/fmicb.2016.00163, 2016.
- 593 Panieri, G.: Foraminiferal response to an active methane seep environment: A case study from
- the Adriatic Sea, Mar. Micropaleontol., 61, 116-130,
- 595 https://doi.org/10.1016/j.marmicro.2006.05.008, 2006.
- 596 Panieri, G., James, R. H., Camerlenghi, A., Westbrook, G. K., Consolaro, C., Cacho, I., Cesari,
- 597 V., and Cervera, C. S.: Record of methane emissions from the West Svalbard continental margin
- 598 during the last 23.500yrs revealed by δ 13C of benthic foraminifera, Global and Planetary
- 599 Change, 122, 151-160, <u>https://doi.org/10.1016/j.gloplacha.2014.08.014</u>, 2014.
- 600 Panieri, G., Lepland, A., Whitehouse, M. J., Wirth, R., Raanes, M. P., James, R. H., Graves, C.
- 601 A., Crémière, A., and Schneider, A.: Diagenetic Mg-calcite overgrowths on foraminiferal tests in 602 the vicinity of methane seeps, Earth and Planetary Science Letters, 458, 203-212,
- 603 https://doi.org/10.1016/j.epsl.2016.10.024, 2017.
- Pascal, P.-Y., Dupuy, C., Richard, P., and Niquil, N.: Bacterivory in the common foraminifer
- Ammonia tepida: Isotope tracer experiment and the controlling factors, J. Exp. Mar. Biol. Ecol., 359, 55-61, <u>https://doi.org/10.1016/i.jembe.2008.02.018</u>, 2008.
- Pawlowski, J.: Introduction to the molecular systematics of foraminifera, Micropaleontology, 46,1-12, 2000.
- 609 Rathburn, A. E., Pérez, M. E., Martin, J. B., Day, S. A., Mahn, C., Gieskes, J., Ziebis, W.,
- 610 Williams, D., and Bahls, A.: Relationships between the distribution and stable isotopic
- 611 composition of living benthic foraminifera and cold methane seep biogeochemistry in Monterey
- 612 Bay, California, Geochemistry, Geophysics, Geosystems, 4, 2003.
- 613 Risgaard-Petersen, N., Langezaal, A. M., Ingvardsen, S., Schmid, M. C., Jetten, M. S. M., Op
- 614 den Camp, H. J. M., Derksen, J. W. M., Piña-Ochoa, E., Eriksson, S. P., Peter Nielsen, L., Peter
- 615 Revsbech, N., Cedhagen, T., and van der Zwaan, G. J.: Evidence for complete denitrification in a
- 616 benthic foraminifer, Nature, 443, 93, <u>https://doi.org/10.1038/nature05070</u>, 2006.
- 617 Salonen, I. S., Chronopoulou, P.-M., Bird, C., Reichart, G.-J., and Koho, K. A.: Enrichment of
- 618 intracellular sulphur cycle-associated bacteria in intertidal benthic foraminifera revealed by 16S
- 619 and aprA gene analysis, Sci Rep-Uk, 9, 1-12, <u>https://doi.org/10.1038/s41598-019-48166-5</u>, 2019.
- 620 Schneider, A., Crémière, A., Panieri, G., Lepland, A., and Knies, J.: Diagenetic alteration of
- benthic foraminifera from a methane seep site on Vestnesa Ridge (NW Svalbard), Deep Sea
 Research Part I: Oceanographic Research Papers, 123, 22-34,
- 623 https://doi.org/10.1016/j.dsr.2017.03.001, 2017.

- 624 Serov, P., Vadakkepuliyambatta, S., Mienert, J., Patton, H., Portnov, A., Silyakova, A., Panieri,
- 625 G., Carroll, M. L., Carroll, J., Andreassen, K., and Hubbard, A.: Postglacial response of Arctic
- Ocean gas hydrates to climatic amelioration, Proceedings of the National Academy of Sciences,
 114, 6215-6220, 10.1073/pnas.1619288114, 2017.
- 628 Shetye, S., Mohan, R., Shukla, S. K., Maruthadu, S., and Ravindra, R.: Variability of
- Nonionellina labradorica Dawson in Surface Sediments from Kongsfjorden, West Spitsbergen,
 Acta Geologica Sinica English Edition, 85, 549-558, <u>https://doi.org/10.1111/j.1755-</u>
- 631 <u>6724.2011.00450.x</u>, 2011.
- $031 \quad 0724.2011.00430.x, 2011.$
- 632 Tavormina, P. L., Hatzenpichler, R., McGlynn, S., Chadwick, G., Dawson, K. S., Connon, S. A.,
- and Orphan, V. J.: Methyloprofundus sedimenti gen. nov., sp. nov., an obligate methanotroph from ocean sediment belonging to the 'deep sea-1'clade of marine methanotrophs, Int. J. Syst.
- 635 Evol. Microbiol., 65, 251-259, <u>https://doi.org/10.1099/ijs.0.062927-0</u>, 2015.
- 636 Torres, M. E., Martin, R. A., Klinkhammer, G. P., and Nesbitt, E. A.: Post depositional alteration

637 of foraminiferal shells in cold seep settings: New insights from flow-through time-resolved

- analyses of biogenic and inorganic seep carbonates, Earth and Planetary Science Letters, 299,
- 639 10-22, <u>https://doi.org/10.1016/j.eps1.2010.07.048</u>, 2010.
- Wefer, G., Heinze, P. M., and Berger, W. H.: Clues to ancient methane release, Nature, 369, 282,
 https://doi.org/10.1038/369282a0, 1994.
- 642 Wollenburg, J. E., Raitzsch, M., and Tiedemann, R.: Novel high-pressure culture experiments on
- 643 deep-sea benthic foraminifera—Evidence for methane seepage-related δ13C of Cibicides
 644 wuellerstorfi, Mar. Micropaleontol., 117, 47-64, 2015.
- 645