Spatial and temporal dynamics of suspended sediment concentrations in coastal waters of South
 China Sea, off Sarawak, Borneo: Ocean colour remote sensing observations and analysis

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#### 12 Abstract

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13 High-quality ocean colour observations are increasingly accessible to support various monitoring and 14 research activities for water quality measurements. In this paper, we present a newly developed 15 regional total suspended solids (TSS) empirical model using MODIS-Aqua's Rrs(530) and Rrs(666) reflectance bands to investigate the spatial and temporal variation of TSS dynamics along the 16 17 southwest coast of Sarawak, Borneo. The performance of this TSS retrieval model was evaluated using 18 error metrics (bias = 1.0, MAE = 1.47, and RMSE = 0.22 in mg/L) with a log10 transformation prior to 19 calculation, as well as a k-fold cross validation technique. The temporally averaged map of TSS 20 distribution, using daily MODIS-Aqua satellite datasets from 2003 until 2019, revealed large TSS 21 plumes detected particularly in the Lupar and Rajang coastal areas on a yearly basis. The average TSS 22 concentration in these coastal waters was in the range of 15 - 20 mg/L, was estimated at these coastal 23 areas. Moreover, the spatial map of TSS coefficient of variation (CV) indicated strong TSS variability 24 (approximately 90%) in the Samunsam-Sematan coastal areas, which could potentially impact nearby 25 coral reef habitats in this region. Our findings-Study on temporal TSS variation provide further 26 evidence that monsoonal patterns drive the TSS release in these tropical water systems, with distinct 27 and widespread TSS plume variations observed between the northeast and southwest monsoon 28 periods. A mHap of relative TSS distribution anomalies revealed strong spatial TSS variations in the 29 Samunsam-Sematan coastal areas, while 2010 recorded a major increase (approximately 100 %) and 30 widespread TSS distribution with respect to the long-term mean. Furthermore, studyour findings on the contribution of river discharge to the TSS distribution showed a weak correlation across time at 31 32 both the Lupar and Rajang river mouth points. The variability of TSS distribution across coastal river 33 points was studied by investigating the variation of TSS pixels at three transect points, stretching from 34 the river mouth into territorial and open water zones, for eight main rivers. ResultsOur findings 35 showed a progressively decreasing pattern of nearly 50 % in relation to the distance from shore, with 36 exceptions in the northeast regions of the study area. Essentially, our findings demonstrate that the 37 TSS levels at the southwest coast of Sarawak are within local water quality standards, promoting 38 various marine and socio-economic activities. This study presents the first observation of TSS 39 distributions at Sarawak coastal systems with the application of remote sensing technologies, to 40 enhance coastal sediment management strategies for the sustainable use of coastal waters and their 41 resources.

42 Keywords: total suspended solids, band-ratio, monsoon, river discharge, Open Data Cube

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### 46 1.0 Introduction

Total Suspended Solids (TSS) play an important role in the aquatic ecosystem as one of the primary 47 48 water quality indicators of coastal and riverine systems (Alcântara et al., 2016; Cao et al., 2018; Chen et al., 2015a; González Vilas et al., 2011; Mao et al., 2012). For example, elevated concentrations of 49 50 TSS in water have an adverse impact on fisheries and biodiversity of the aquatic ecosystem (Bilotta and Brazier, 2008; Chapman et al., 2017; Henley et al., 2000; Wilber and Clarke, 2001). Understanding 51 the impacts of varying water quality in relation to TSS status has been one of the primary concerns 52 53 with respect to a country's growing Blue Economy status and sustainable management of aquatic resources (Lee et al., 2020a; Sandifer et al., 2021; World Bank and United Nations Department of 54 55 Economic and Social Affairs (UNDESA), 2017). With about 40 % of the world's population living within 56 100 km of coastal areas (United Nations, 2017), and with more than 80 % of the population in Malaysia 57 living within 50 km of the coast (Praveena et al., 2012), water quality monitoring and management 58 efforts are important at both regional and global scale.

59 Studying TSS distribution can provide insights into the connections between land and ocean 60 ecosystems (Howarth, 2008; Lemley et al., 2019; Lu et al., 2018). For instance, TSS dynamics allow us 61 to understand the impacts of sediment transport and sediment plumes, particularly in areas 62 experiencing large-scale deforestation, land conversion and damming of rivers (Chen et al., 2007; 63 Espinoza Villar et al., 2013). Sarawak, Malaysian Borneo, experienced significant land use and land cover change activities over the past four decades, with widespread land conversion and deforestation 64 for developments and large-scale plantation activities (Gaveau et al., 2016), as well as building of 65 66 major road infrastructures, such as the Pan-Borneo highway, and hydroelectric dams (Alamgir et al., 2020). As a result, river and coastal systems may potentially drive large TSS loads into downstream 67 68 systems and into the marine and open ocean waters.

Situated at the southern part of the South China Sea, the region of Sarawak, Malaysian
Borneo, has a coastline of about 1035 km where mangrove forests are dominant (Long, 2014). The

coastal regions of Sarawak are rich with marine coastal biodiversity and coral reefs, which can be found at the northeast and southwest part of Sarawak (Praveena et al., 2012). While the coasts of Sarawak provide important socio-economic values to the local communities (Lee et al., 2020b), these coastal areas are potentially facing water quality degradation from TSS riverine outputs in response to land use and land cover change activities.

TSS concentrations are commonly measured through conventional laboratory-based methods to quantify TSS concentrations by field collection of water samples (Ling et al., 2016; Mohammad Razi et al., 2021; Soo et al., 2017; Soum et al., 2021; Tromboni et al., 2021; Zhang et al., 2013). Currently, real-time high-frequency TSS observations using modern optical and bio-sensor systems are also possible (Bhardwaj et al., 2015; Horsburgh et al., 2010). These sensors can be generally found onboard ship and buoy-based observation platforms. Yet, it remains a challenge to quantify TSS concentrations of large spatial coverage and high temporal frequency with these approaches.

83 Ocean colour remote sensing technologies represent an increasingly accessible and powerful 84 tool to provide a synoptic view for short or long-term water quality studies at high temporal and spatial 85 resolutions (Cherukuru et al., 2016; Slonecker et al., 2016; Swain and Sahoo, 2017; Wang et al., 2017; 86 Werdell et al., 2018). Remote sensing can help overcome several constraints of conventional intensive 87 field campaigns such as: (i) costly field campaigns from boat rentals or cruise; (ii) time-consuming and 88 inadequate manpower; and most importantly for this study, (iii) limited spatial and temporal field 89 NASA's Moderate Resolution Imaging Spectroradiometer (MODIS)-Aqua coverage. (https://modis.gsfc.nasa.gov/about/) has a distinctive advantage with its daily revisit time, a spatial 90 91 resolution of 250 – 1000 m, and a large collection of ocean colour data since 2002. Other sensors 92 offering ocean colour measurement capabilities include Landsat-8, which, in comparison with MODIS-93 Aqua, has a 16-day revisit time and high spatial resolution of 30 m. Despite Landsat's powerful ability 94 in capturing higher resolution images, the longer revisit interval may not be suitable for characterizing and studying water bodies with high dynamics of various water constituents. 95

96 Several MODIS-derived models have been developed for TSS retrievals (Chen et al., 2015b; 97 Espinoza Villar et al., 2013; Jiang and Liu, 2011; Kim et al., 2017; Zhang et al., 2010b), including 98 empirical, semi-analytical and machine-learning approaches (Balasubramanian et al., 2020; Jiang et 99 al., 2021). However, the performance of these models proved to be less satisfactory, with recorded 100 low r<sup>2</sup> and high bias and mean absolute error (MAE) values when tested with in situ TSS datasets 101 (Supplementary Materials, Table S1). While these global TSS remote sensing models address the need 102 to improve TSS retrievals and to monitor global TSS trends in various water class types, they tend to underperform in more localised and regional studies (Mao et al., 2012; Ondrusek et al., 2012). The 103 104 coastal waters of Borneo are well-mixed throughout the year and enriched with suspended material and dissolved organic matter (Müller et al., 2016). Various water quality studies of the river systems 105 106 have been actively carried out to assess the dynamics of numerous water quality constituents in 107 response to human activities, with TSS concentrations being one of the primary environmental 108 concerns in this region (Ling et al., 2016; Müller-dum et al., 2019; Tawan et al., 2020). Although studies 109 on the water quality of coastal systems in Borneo have gradually gained much attention (Cherukuru 110 et al., 2021; Limcih et al., 2010; Martin et al., 2018; Soo et al., 2017), there is still much knowledge to 111 gain on the understanding of how coastal waters in the region have been impacted by TSS loadings 112 and transport over large spatial and temporal scales.

113 Here, in this paper, we present a new regional empirical TSS remote sensing model. While 114 various remote sensing models have their own unique computational strengths, this studypaper 115 demonstrates the reliability of band ratio TSS model to be applied within optically complex waters. 116 With the ongoing efforts to address and minimize water quality degradation in coastal systems, as 117 outlined in the United Nation's Sustainability Development Goals no. 14, our study aims to apply the 118 new empirical regional TSS remote sensing model to: (a) investigate the spatial and temporal 119 variability in TSS, (b) identify hotspots of TSS distribution in the coastal waters of Sarawak, Malaysian 120 Borneo, using a long time series of MODIS-Aqua data from year 2003 until 2019, and (c) study the 121 varying monsoonal and river discharge patterns in relation to TSS distribution at river mouths.

## 122 2.0 Methodologies

distribution was mapped using atmospheric-corrected MODIS-Aqua Level 2 product. 124 125 Existing in situ TSS measurements from Martin et al. (2018) and spectral reflectance data from Cherukuru et al. (2021) 126 127 128 Development of regional TSS remote sensing model 129 130 Validation of TSS remote sensing model performance via statistical analysis 131 Application of TSS remote sensing model on Open Data Cube 132 (ODC) platform 133 Retrieval of atmospheric-corrected MODIS-Aqua Level 2 product 134 in ODC platform 135 **Retrieval of Global Precipitation** 136 Measurement (GPM) datasets 137 138 Computation of river discharge using GPM datasets 139 140 Map and plot daily MODIS-Aqua satellite datasets from 2003 till 2019 on: 141 Spatial variation of TSS • 142 Temporal variation of TSS Relationship between TSS and river discharge 143 TSS variability across coastal waters 144

123 Figure below summarizes the processes carried out in this study. Spatial and temporal variation of TSS

Fig. 1: Flowchart summarizing the processes of developing a regional TSS remote sensing model and applying it to analyse
 the spatial and temporal variation of TSS over the study region, using MODIS-Aqua data from year 2003 until 2019. Long term MODIS-Aqua datasets were analysed and mapped on an Open Data Cube (ODC) platform with implementation of robust
 Python libraries and packages.

#### 149

## 150 2.1 Area of study

151 Our study focuses on the southwestern coast of Sarawak (between 1.9° N, 109.65° E and 2.8° N, 111.5° 152 E) in Malaysia, which sits at the northwest part of the Borneo Island. Generally, the island of Borneo (between 3.01° S, 112.18° E and 6.45° N, 117.04° E) contains rich tropical rainforests and biodiversity 153 154 on the lands of Sarawak and Sabah (Malaysia), Brunei, and Indonesia. Typically, Sarawak is a tropical climate region, recording an average ambient temperature of 27.8 °C (variation of 1.8 °C) throughout 155 156 the year. It records high precipitation with an average of 4116.7 mm/yr in Kuching (1.5535° N, 157 110.3593° E), the capital city of Sarawak. Yearly, it experiences both a dry and wet season, which is 158 influenced by: (i) the southwestern monsoon (May to September) and (ii) the northeastern monsoon 159 (November to March). Rivers in Sarawak are connected to the South China Sea and flow through 160 various plantation types, such as oil palm, rubber and sago (Davies et al., 2010). 161 In this study, the southwestern part of Sarawak's coastal regions (Fig. 2), (between 1.9° N, 109.65° E 162 and 2.8° N, 111.5° E) was studied, which comprise several major rivers (e.g. Lupar, Sebuyau, Sematan), 163 as well as the Rajang River, the longest river in Malaysia. Rajang river basin consists in tidally influenced 164 river channel which splits into a northwest (Igan, Lassa and Paloh) and a southwest (Rajang, Belawai) 165 Rajang river delta (Staub et al., 2000). The Rajang river basin drains a dominant area (>50,000km<sup>2</sup>) of 166 sedimentary rocks (Milliman and Farnsworth, 2013; Staub et al., 2000) extending from Belaga to Sibu, 167 with major peatland areas converted into oil palm plantations (Gaveau et al., 2016) as its river flows 168 into the South China Sea (Milliman and Farnsworth, 2013). Major settlements along the Rajang river 169 comprise of Kapit and Kanowit town areas, as well as Sibu city, with a total population size of about 170 388,000 inhabitants (Department of Statistics, 2020). Lupar and Saribas rivers, respectively, comprise

171 <u>a catchment area size of approximately 6500 and 1900 km<sup>2</sup> (Lehner et al., 2006). Situated at the</u>

172 southwest side of the Rajang catchment, Lupar and Saribas rivers surround the Maludam National

173	Park, which is Sarawak's remaining biggest single patch of peat swamp forest (Sarawak Forestry
174	Corporation, 2022). Adjacent to Lupar river mouth is the Sadong river, with an approximate catchment
175	area size of 3500 km <sup>2</sup> (Kuok et al., 2018). Sadong river runs about 150 km and flows through oil palm
176	plantations (Staub and Esterle, 1993). These river systems are associated with increasing land use
177	activities and land cover changes in this region, which essentially transport and connect various
178	biogeochemical water components to the coastal systems of Sarawak.



180 Fig. 2: Map of the study area (© Google Maps), located in the southwestern part of Sarawak, Malaysia (inset). Indicators

<sup>181</sup> show the location of sampling sites used during field expeditions carried out in June and September 2017.

<sup>182 2.2</sup> In situ TSS measurements

<sup>TSS measurements data-were taken from Martin et al. (2018). A total of 35 coastal sites were studied
and are denoted SJ, SS, and SM (see: Table 1 & Fig. 2). These water samples were collected in the
month of June (SJ region) and September (SS and SM regions) in 2017. Water samples were filtered,</sup> 

186 and filters were dried and ashed prior to weighing process. Full details of water sampling and TSS

187 analysis is available in Martin et al. (2018).

188 2.3 Development, calibration and validation of TSS model

In situ remote sensing reflectance spectral data, Rrs(λ), along with 35 measured TSS values, were used
to develop a new remote sensing TSS empirical model for MODIS-Aqua for this case study. Field
measurements of SM, SJ & SS datasets, as shown in Table 1, were used to calibrate the MODIS-Aqua
TSS remote sensing model.

For the in situ remote sensing reflectance,  $Rrs(\lambda)$  readings, a TriOS-RAMSES spectral imaging radiometer was used to measure downwelling irradiance,  $Ed(\lambda)$ , and upwelling radiance,  $Lu(\lambda)$ , with measurement protocols from Mueller et al. (2002). These measurements were recorded under stable sky and sea conditions during the day (10AM to 4PM) with <u>high solar elevation angles</u>. <u>high sun</u> elevation angle condition.

Measurements of reflectance,  $\operatorname{Rrs}(\lambda)$ , were recorded concurrently with the collection of water samples (as described in Section 2.2) and were recorded at wavelength ranging from 280 to 950 nm, which covers the spectrum of ultraviolet, visible and visible/ultraviolet light. These measurements were recorded on a float to capture Lu(0–,  $\lambda$ ) and Ed(0+,  $\lambda$ ), where 0- and 0+ refer to below-surface and above-surface, respectively.

203 Remote sensing reflectance,  $Rrs(\lambda)$ , was computed as follows with reference to Mueller et al. (2002):

$$\operatorname{Rrs}(\lambda, 0+) = \frac{1-p}{n^2} \times \frac{\operatorname{Lu}(0-,\lambda)}{\operatorname{Ed}(0+,\lambda)}$$
(1)

where p = 0.021 refers to the Fresnel reflectance and n = 1.34 is the refractive index of water. Full

205 details of this methodology can be found at Cherukuru et al. (2021).

206 2.3.1 Calibration of empirical model and application to MODIS-Aqua

With the intention to apply <u>a</u>regional TSS remote sensing model to MODIS-Aqua product, a total
 number of 35 <u>different TSS</u> datasets <u>of TSS concentrations</u> were collected in coastal conditions (salinity
 > 15 PSU) and convolved to generate MODIS-Aqua data.

In this study, retrieval of water constituents was established using spectral band ratio combinations which have proven to be a straightforward, yet reliable method for estimating water constituents in optically turbid waters (Ahn and Shanmugam, 2007; Cao et al., 2018; Lavigne et al., 2021; Morel and Gentili, 2009; Neil et al., 2019; Siswanto et al., 2011). Band ratio models help to offset signal noise, such as the effects of atmospheric and irradiance of spectral reflectance to a certain degree (Cherukuru et al., 2016b; Ha et al., 2017; Hu et al., 2012; Liu et al., 2019).

A variety of models using single bands, as well as a combination of MODIS-Aqua's Blue, Green & Red bands (412nm, 440nm, 488nm, 532nm, 555nm & 660nm) were calibrated using field measurements as dependent variable. The calibration process was tested out using various model functions, including linear, power, exponential, and logarithm functions. The best empirical TSS retrieval model was fitted by means of a regression between the in situ TSS data and in situ radiometer values, and can be expressed as follows:

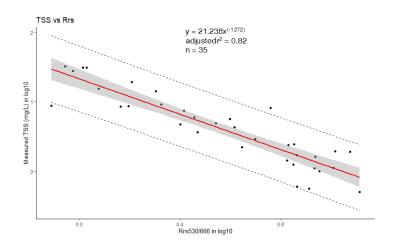
222

 $TSS = 21.238[Rrs(530)/Rrs(666)]^{-1.272}$ (2)

223 This power function model resulted in a coefficient of determination (R<sup>2</sup>) of 0.82 (Fig. 3).

Table 1: Summary statistics of TSS values collected at areas SJ, SS, and SM located within coastal regions in this study, with a total of 35 datasets recorded.

Coastal Area	Minimum	Maximum	Mean	S.D.	C.V.	n
SJ	1.1	19.24	6.89	6.62	96.09	6
SS	0.56	32.1	12.50	11.43	91.45	16
SM	0.5	8.14	2.59	2.70	104.53	13





228 Fig. 3: Empirical relationship of TSS retrieval between in situ Rrs(530)/Rrs(666) bands ratio and measured TSS

- concentration (mg/L), as established via a power law function. Upper and lower dashed lines indicate the 95 % prediction
   interval of the regression.
- 231 2.3.2 Performance assessment and validation of MODIS-Aqua empirical model

An evaluation of the model was performed using a k fold cross validation technique (Refacilzadeh et al., 2020) given the small size of the TSS dataset used in this study (Table 2). A selection of k = 7 was

234 assigned to split the datasets into k groups with an equal number of data points-

An assessment of the performance error of the newly developed TSS model was carried out as per Seegers et al. (2018)'s recommendation for interpreting ocean colour models. These performance metrics used here include the bias, Mean Absolute Error (MAE), Root Mean Squared Error (RMSE), coefficient of variation (CV), as well as the coefficient determination, r<sup>2</sup>, based on the following calculations:

240 Bias = 
$$10^{n} \left[ \frac{\sum_{i=1}^{n} log10 (Mi) - log10(Oi)}{n} \right]$$
 (3)

241 
$$MAE = 10^{n} \left[ \frac{\sum_{i=0}^{n} |\log 10 (Mi) - \log 10(Oi)|}{n} \right]$$
(4)

242 
$$RMSE = \sqrt{\frac{\sum_{i=1}^{n} (log10(Mi) - log10(Oi))^{2}}{n}}$$
(5)

$$CV = \frac{\sigma}{\mu} \times 100\%$$
 (6)

244 where M represents the modelled TSS values, n is the number of samples, and O represents the 245 observed TSS measurements, while  $\sigma$  refers to standard deviation and  $\mu$  represents the mean value. 246 Equations (3), (4) and (5) use a log10-transform of the data as the range of TSS values can span several 247 orders of magnitude. As such, an application of the log-transform prior to error metric calculation 248 allows us to account for uncertainties that are proportional to the concentration values 249 (Balasubramanian et al., 2020; Seegers et al., 2018).

250 251 252 253 Table 2: Calibration and accuracy assessment of the newly derived MODIS-Aqua models in this study for TSS estimations tested using various model functions. Calculation for bias, MAE and RMSE use a log-transform of the data prior to calculation of error metric measurements, as adapted from Seegers et al. (2018) and Balasubramanian et al. (2020). Band ratio Rrs(530)/Rrs(666) is established as function x. Power function model is selected based on low performance metric values.

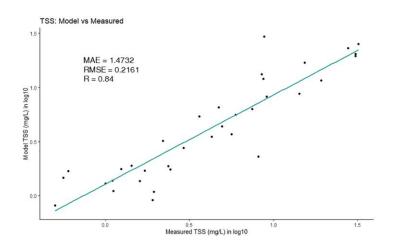
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Model	<u>Function</u>	<u>Bias</u>	MAE	<u>RMSE</u>	<u>CV (%)</u>	<u>R</u>
<u>Power</u>	<u>TSS = 21.238x^-1.272</u>	<u>0.9999</u>	<u>1.4732</u>	<u>0.2161</u>	<u>4.74</u>	<u>0.84</u>
<u>Linear</u>	<u>TSS = -1.8193x + 16.928</u>	<u>1.4463</u>	<u>1.8549</u>	<u>6.7174</u>	<u>20.699</u>	<u>0.6854</u>
<u>Exponential</u>	<u>TSS = 17.784e^-0.296x</u>	<u>1.0791</u>	<u>1.4906</u>	<u>6.3088</u>	<u>3.8920</u>	<u>0.8154</u>
<u>Logarithmic</u>	<u>-8.872ln(x)+19.383</u>	<u>1.1336</u>	<u>1.6177</u>	<u>5.3735</u>	<u>-17.056</u>	<u>0.8128</u>

254

TSS       7       0.85       0.2159       0.1747         able 3: Calibration and accuracy assessment of the newly derived MODIS-Aqua models in this study for TSS- alculation for bias, MAE and RMSE use a log-transform of the data prior to calculation of error metric measing dapted from Seegers et al. (2018) and Balasubramanian et al. (2020).       Parameter       Bands (nm)       Model       Bias       MAE       RMSE       CV (%)		Paramo	<del>xter</del> <del>k-fold (n)</del>	<del>R2</del>	RMSE	MAE	
alculation for bias, MAE and RMSE use a log-transform of the data prior to calculation of error metric meas dapted from Seegers et al. (2018) and Balasubramanian et al. (2020).		<del>TSS</del>	₽	<del>0.85</del>	<del>0.2159</del>	<del>0.1747</del>	_
Parameter Bands (nm) Model Bias MAE RMSE CV (%)							
	Calculation for b	ias, MAE and RMSE	use a log transform of	he data prior			
TSS 21.238[(Rrs530/Rrs 0.9999 1.4732 0.2161 4.74	alculation for b dapted from Sc	ias, MAE and RMSE egers et al. (2018)	Euse a log transform of i and Balasubramanian et Model TSS =	<del>he data prior</del> al. (2020). <b>Bias</b>	to calculatio	<del>n of error m</del>	etric measu

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Fig. 4: Scatterplot of modelled TSS values derived from the proposed model and measured TSS values (mg/L). Check decimals in figure

264 An evaluation of the model was performed using a k-fold cross validation technique (Refaeilzadeh et

al., 2020) given the small size of the TSS dataset used in this study (Table 2). A selection of k = 7 was

assigned to split the datasets into k groups with an equal number of data points.

267 <u>Table 2: Assessment of fitting error for the proposed TSS model, using k-fold cross validation.</u>

Parameter	<u>k-fold (n)</u>	<u>R2</u>	RMSE	MAE
TSS	<u>7</u>	0.85	<u>0.2159</u>	0.1747

269 While these results point to low error levels achieved by the proposed regional TSS retrieval model 270 (Table 3, Fig. 4), caution should be used when applying it to various water types. Water type 271 classification has been thoroughly described by Balasubramanian et al. (2020) where waters are 272 classed into Type I (Blue-Green waters), Type II (Green waters), and Type III (Brown waters). 273 Essentially, the Green-to-Red band ratio is optimised with these datasets corresponding to sediment-274 dominated and yellow-substance loaded water conditions. As highlighted by Morel & Belanger (2006), 275 waters of this type do not have the same spectral characteristics as phytoplankton-rich waters this 276 type of waters is not spectrally consistent to phytoplankton-rich waters (also known as Case 1 waters). 277 In addition to the impact on water clarity, sediment particles (often red-brownish coloured) also tend

to enhance the backscattering and absorption properties, especially at shorter wavelengths (Babin et al., 2003), while the additional presence of coloured dissolved matter (yellow substance) leads to strong absorption properties at short wavelengths. As the TSS retrieval model was developed from samples taken in waters that are bio-optically rich in suspended solids and dissolved organic matter, an application of this TSS model needs to be done cautiously when applying to other water types, particularly those with large concentration of phytoplankton.

#### 284 2.4 Application of TSS retrieval model

285 Daily MODIS-Aqua satellite data from year 2003 to 2019 (total of 6192 individual time slices) were 286 studied with a 2°x 2° spatial resolution (longitude: 109.38, 112.0; latitude: 1.22, 3.35) which covers 287 the southwestern coastal region of Sarawak and southern part of the South China Sea. 288 Atmospherically corrected MODIS-Aqua level 2 reflectance products (Bailey et al., 2010; NASA Official, 289 n.d.) were retrieved for the application of the TSS model proposed in this study. Negative remote 290 sensing reflectance values, possibly due to failure of atmospheric correction, were filtered out before 291 applying the retrieval model, as expressed in Eq. (2), to map the spatial and temporal distribution of TSS estimates. In addition, averaging of spatial and temporal TSS variation maps in this study was 292 293 carried out by filtering TSS values with fewer than 10 valid data points over the whole time series, 294 along with application of sigma clipping operation (refer to: 295 https://docs.astropy.org/en/stable/api/astropy.stats.sigma\_clip.html).

296 2.4.1 Open Data Cube

In this study, the analysis of remote sensing data over large spatial extents and at high temporal resolution was carried out using robust Python libraries and packages run on an Open Data Cube (ODC) platform. Open Data Cube is an open-source advancement in computing technologies and data architectures which addresses the growing volume of freely available Earth Observation (EO) satellite products (Giuliani et al., 2020; Killough, 2019). ODC provides a collection of software which index, manage, and process large EO datasets such as satellite products from the MODIS, Landsat and 303 Sentinel missions (Gomes et al., 2021). These satellite datasets are structured in a multi-dimensional 304 array format, and provide layers of information across latitude and longitude (Open Data Cube, 2021). Leveraging the growing availability of Analysis Ready Data (ARD), and with support from the 305 306 Committee of Earth Observation Satellites (CEOS) (Killough, 2019), the ODC concept has been 307 deployed in many countries across the world. These existing deployments include Digital Earth Africa 308 (https://www.digitalearthafrica.org/), Digital Earth Australia (DEA) (https://www.dea.ga.gov.au/), 309 Vietnam Open Data Cube (http://datacube.vn/), and Brazil Data Cube (https://github.com/brazil-datacube), which provide various time-series datasets of the changing landscape and water content in 310 311 these specific regions (Giuliani et al., 2020; Gomes et al., 2021; Killough, 2019; Lewis et al., 2017). The ecosystem and architecture of ODC is well explained at opendatacube.org. The codes and tools used 312 313 in this application drew upon the information provided in various DEA notebooks (Krause et al. (2021), 314 which can be found at https://github.com/GeoscienceAustralia/dea-notebooks/. 315 2.5 Precipitation data and c-computation of river discharge

316 Monthly precipitation values (mm) over the Lupar and Rajang basins were extracted from the Global 317 Precipitation (GPM) IMERG Measurement Level 3 satellite datasets 318 (https://gpm.nasa.gov/data/imerg) (Supplementary Materials, Fig. S4 7) in order to assess the 319 influence of precipitation in each river basin in relation to TSS concentration at the corresponding river 320 mouth (Supplementary Materails, Fig. S4 - 7).-

Derivation of river discharge (m<sup>3</sup>/s) was computed using total precipitation estimates (mm) over each river basin, and multiplied by a surface discharge runoff factor for the studied region (Sim et al., 2020). The surface runoff was estimated to be 60 % of total precipitation (Staub et al., 2000; Whitmore, 1984). Precipitation data were extracted from the monthly NASA Global Precipitation Measurement (GPM) Level 3 IMERG dataset (https://gpm.nasa.gov/data/imerg). In this study, the Rajang river basin, as well as the combined basins of the Lupar, Sadong, and Saribas rivers (hereafter referred to as the Lupar basin), were studied for their river discharge rates in relation to TSS release.

#### 328 3.0 Results & Discussion

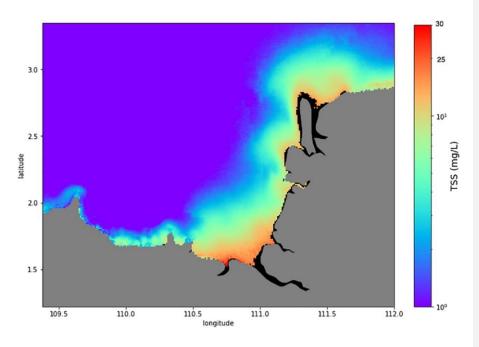
#### 329 3.1 Spatial variation of TSS distribution

Changes in TSS distribution occur across space and time. The regional TSS remote sensing model calibrated in this study was applied to the time series of MODIS-Aqua data to study the variability of spatial TSS distribution and identify potential hotspot areas susceptible to TSS water quality degradation. The map of average TSS for the Sarawak region was generated (Eq. 2) by averaging all the daily MODIS-Aqua TSS images (2003 to 2019) and is presented in Fig. 5. The results show that the waters in the northeast <u>region</u> of the study area, stretching from the Sadong river to the Rajang/Igan river have seen sustained levels of TSS over the 17 years considered in this study.

337 The temporally averaged spatial distribution map (Fig. 5) shows TSS concentrations in the range of 15 338 - 20 mg/L near the river mouth areas, with widespread TSS plumes extending into the South China 339 Sea (Fig. 5). Based on the Malaysia Marine Water Quality Criteria and Standard (Supplementary 340 Materials, Table S2) (Department of Environment, 2019), these coastal waters fall under Class 1 in relation to their TSS (mg/L) status. This classification indicates that these coastal waters support and 341 342 preserve marine life in this local region. Yet, several studies have expressed concerns regarding high 343 TSS loadings in riverine waters owing to the impacts of various land use and land cover changes (LULC) 344 (Ling et al., 2016; Tawan et al., 2020). Among these, the Rajang river has been highlighted to be heavily 345 impacted by various LULC activities such as large-scale deforestation and construction of hydropower 346 dams (Alamgir et al., 2020). In situ water quality studies by Ling et al. (2016) reported on high TSS 347 estimates at one of the upstream tributaries of Rajang river, the Baleh river, with TSS readings up to 348 approximately 100 mg/L. Another study by Tawan et al. (2020) reported a significant TSS release 349 reaching to 940,000 mg per day during wet seasons, with maximum TSS concentrations of 1700 mg/L 350 in the upstream tributaries of the Rajang river, particularly at the Baleh and Pelagus rivers. The 351 majority of the upstream tributary rivers were categorised as Class II (during dry season) and Class III 352 (during wet season) waters according to the National Water Quality Index (Supplementary Materials,

Table S3) (Department of Environment, 2014), due to increased soil erosion from surrounding LULC activities (Tawan et al., 2020). These local in situ findings provide valuable insights on point source TSS estimates in these LULC change regions. Coupled with our spatial map of average TSS captured by remote sensing technologies, our findings seem to suggest that a large portion of TSS loadings from inland and upstream rivers would have settled and deposited in these river channels and were not completely discharged outwards into the coastal areas, which would have caused major water quality

359 degradation in the corresponding coastal systems.



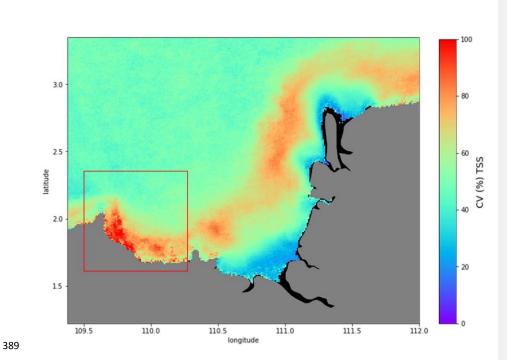
360

361 Fig. 5: Temporally averaged 2°x 2° map of TSS distribution (on a log scale) across the time dimension for each pixel.

Historical patterns of TSS concentration were assessed by comparing annual maps of average TSS distribution (Supplementary Materials; Fig. S1), as well as time series of TSS estimates at the Lupar and Rajang river mouths (Supplementay materials; Fig. S2). From our findings, the annual TSS maps further support the observation where TSS release was evident at Lupar and Rajang/Igan river mouths from 2003 till 2019, which points to Class I of local water quality standards in relation to TSS (mg/L) status. This was found to consistently occur every year. Furthermore, the TSS trend study showed that
both the Lupar and Rajang river mouth points have a gradual increase of TSS concentration over the
17 years (Supplementay materials; Fig. S2). This increasing trend was, however, not statistically
significant (p = 0.43 for Lupar, and p = 0.15 for Rajang).

Moreover, a map of the TSS coefficient of variation (CV) was computed to identify areas with a high degree of relative TSS variation over time (Fig. 6). Here again, the map of CV (%) was produced by aggregation of the daily MODIS-Aqua images (6192 time steps) from 2003 until 2019. Figure 6 shows that the Samunsam-Sematan coastal region (as highlighted by the red box) exhibits an increased level of TSS distribution variability, with a recorded CV of more than 90 %.

376 The Samunsam-Sematan coastal region contains near-pristine mangrove forests which are sheltered 377 from major LULC activities, as compared to other studied sites. Samunsam-Sematan is also well-known 378 locally as a recreational hotspot with coral reefs and various national parks (Sarawak Tourism Board, 379 2021). Data from the Centre for International Forestry Research (CIFOR) Forrest Carbon database 380 (CIFOR, n.d.) revealed that there was more than double the amount of total forest loss (approx. 5,000 381 Ha) recorded in Lundu, a nearby township in the Sematan area in 2011 as compared to the previous 382 years. Deforestation activities, regardless of their scale, can inevitably promote sediment loss and soil 383 leaching into the nearby river systems (Yang et al., 2002). Important information regarding the 384 variability in water quality (as shown in Fig. 6) can provide support to local authorities and relevant 385 agencies in order to identify vulnerable areas that need to be monitored closely, such as the Lundu-386 Sematan region in this case. The CV map thus offers interesting insights into how TSS distribution can 387 vary across large spatial areas, which can ultimately impact local socio-economic activities in this 388 region (Lee et al., 2020b).



390 Fig. 6: Map of CV (%) calculated from the daily time series of MODIS-Aqua satellite images from 2003 until 2019.



#### 401 3.2 Temporal variation of TSS distribution

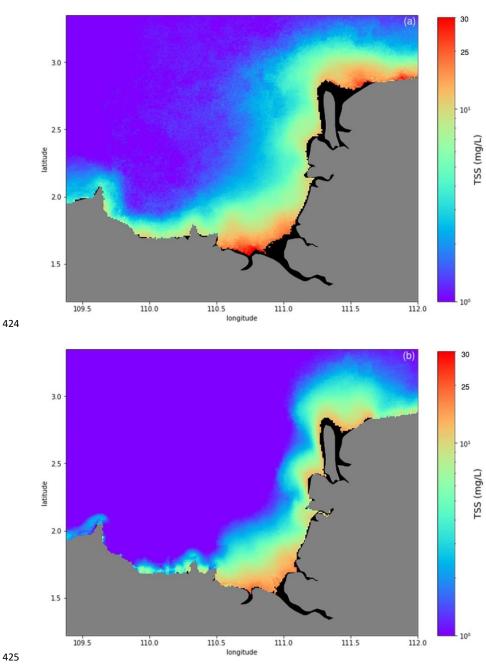
On a temporal scale, the northeast (NE) monsoon period shows a distinct difference in the widespread intensity of TSS distribution as compared to the southwest (SW) monsoon period, along the Sarawak coastline over the 17 years of the considered time series (Fig. 7). Mapping of temporal variations between monsoons using time-series MODIS-Aqua datasets can provide an improved understanding on the intensity of monsoonal patterns in driving the TSS distribution in this region. As shown in Figure 7, TSS release can be seen to extend further into the open ocean South China Sea region during the NE monsoon periods (Fig. 7a) in comparison to the SW monsoon periods (Fig. 7b).

In addition, the differences in TSS release between the NE and SW monsoons ((NE-SW)/NE x 100) were mapped as shown in Fig. 7c. Widespread TSS plumes are detected at Lundu/Sematan region (>80 % relative difference in TSS concentration) on the southwest side of the study area, while substantial TSS plumes are observed in front of the Igan river channel, with more than 50 % relative difference in TSS concentration in comparison to SW monsoon periods. Sadong coastal area is observed to receive considerable TSS loadings (> 30%) during NE monsoon periods.

These coastal areas would thus be more likely to be impacted by the TSS release during the NE monsoon periods. These findings further strengthen the evidence that tropical rivers are majorly impacted by climatic variability such as monsoonal patterns, as highlighted in a study at Baleh river in Sarawak (Chong et al., 2021). This suggests that monsoon rains, which typically last for several months, play an integral role in driving the discharge of TSS in tropical rivers.

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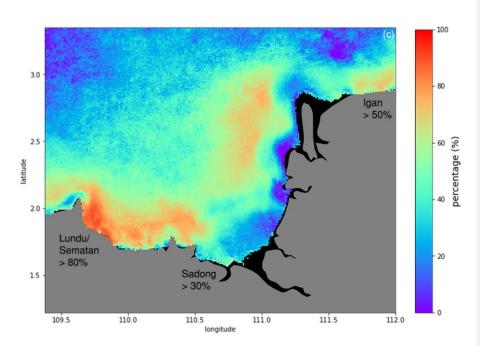




Fig. 7: Long-term average of TSS estimates (mg/L) during the Northeast monsoon (a), and the Southwest monsoon (b). The
map of TSS difference between the Northeast and Southwest monsoon periods, computed in relative percentage (%), is
shown in (c).

Several climatic studies in the Borneo region highlighted 2009 as a year with extreme rainfall events which caused major floods in Sarawak (Dindang et al., 2011; Sa'adi et al., 2017), while drought events were reported in 2014 (Bong and Richard, 2020). Hence, in this study, TSS dynamics for the Lupar and Rajang rivers were studied by assessing the variation of TSS values at selected pixels in relation to monsoonal rainfall patterns in 2009 and 2014 (Supplementary Materials; Fig. S3).

435 Monthly precipitation values (mm) over the Lupar and Rajang basins were extracted from the Global
 436 Precipitation Measurement (GPM) satellite datasets (Supplementary Materials, Fig. S4—7) in order to
 437 assess the influence of precipitation in each river basin in relation to TSS concentration at the
 438 corresponding river mouth.

Generally, the results show fluctuations of TSS concentrations across the NE and SW monsoon periods
in relation to precipitation values (Fig. 8a – d). Based on Fig. 8a, monthly precipitation values recorded

for the Lupar river basin in 2009 showed a clear decreasing trend from the NE monsoon period (wet
season) to the SW monsoon period (dry season), while gradually increasing approaching the year end's
NE monsoonal period. A similar precipitation pattern was observed for the Rajang river basin during
the same year (Fig. 8c).

445 However, these results also show that the TSS distribution (mg/L) at the Lupar river mouth seems to 446 show no distinct trend of decreasing TSS concentration estimates during the SW monsoon period in year 2009 (Fig. 8a) in relation to its precipitation values. Additionally, a sharp rise of TSS release can 447 448 be seen in the month of May (beginning of SW monsoon period), with a near equivalent intensity of 449 TSS release during the NE monsoon period. This observation may potentially be caused by the lag 450 between the time of rainfall events occurring during NE monsoon periods and TSS release entering the coastal river regions. A similar observation was described by Sun et al. (2017a) suggesting that 451 452 riverine outputs could take several days, and even up to one month to reach the coastal river points. 453 Considering the occurrence of extreme rainfall events in 2009, our findings are in agreement with 454 these processes as TSS concentrations generally exhibit a similar intensity throughout the NE and SW 455 monsoonal periods for the Lupar river (Fig. 8a). This result could suggest that the occurrence of extreme rainfall events, as reported for the year 2009, can exert a much larger impact on TSS 456 457 transportation and release in monsoon-driven tropical rivers.

However, our estimates show a generally lower TSS concentration at the Rajang river mouth during
the SW monsoon periods (dry season) in 2009 (Fig. 8c) as compared to the Lupar river mouth (Fig. 8a).
While these observations were recorded during the same year in 2009, when Sarawak experienced
extreme rainfall events, the monsoonal influence between these two different river basins show a
slight difference in terms of TSS distribution. A possible reason for these observations can be explained
by the Rajang river's unique meandering features, which can potentially induce a different
sedimentological behaviour compared to the Lupar river channels, which was discussed by Omorinoye

465 et al. (2021) in a study on geomorphological features of the Sadong river in relation to the

# 466 sedimentation process.

467	Drought events in 2014 can be seen to impact the precipitation values at both the Lupar (Fig. 8b) and
468	Rajang river basins (Fig. 8d). There is no apparent patterns of decreasing precipitation values during
469	the shift of NE to SW monsoonal periods as compared to the year 2009, for either river basin. However,
470	precipitation values were found to increase sharply during the year end NE monsoon period for both
471	river basins. The TSS concentrations at the Lupar coastal river points were found to be the highest
472	during the NE monsoon period earlier in January and February of 2014 (Fig. 8b). This may be due to
473	the temporal lag in the transition of TSS discharge into the coastal systems arising from the prior
474	months (November and December) in the previous year, when higher rainfall events were typically
475	observed in this region (Gomyo and Koichiro, 2009; Tangang et al., 2012). The TSS distribution at both
476	Lupar and Rajang coastal river points showed no distinct trend in relation to the precipitation values
477	throughout a period of ten months until November 2014. These findings suggest that coastal areas in
478	the Borneo region may not be experiencing critical water quality degradation during dry seasons.
479	

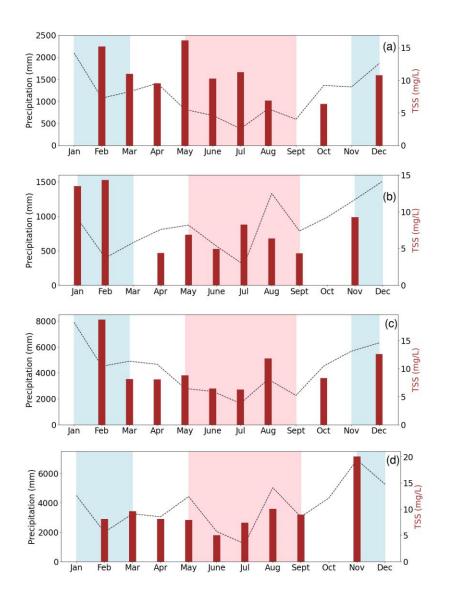


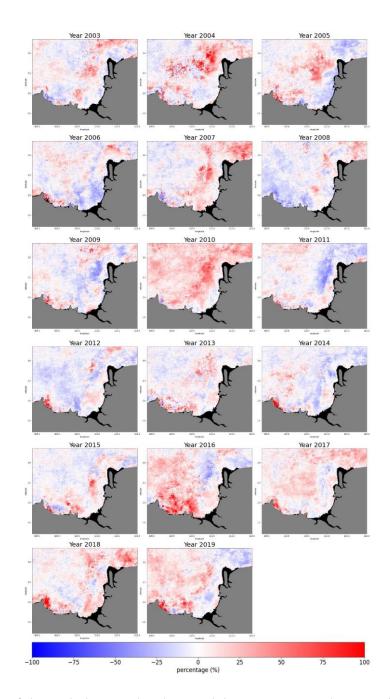
Fig. 8: Temporal analysis of precipitation (mm) from the Lupar and Rajang river basins in relation to TSS concentrations (mg/L)
during the NE and SW monsoon periods at the Lupar ((a): 2009; (c): 2014) and Rajang ((b): 2009; (d): 2014) coastal river point.
The NE monsoon months are highlighted with a blue background; those of the SW monsoon with a pink background, and
intermonsoon periods with a white background.

## 492 <u>3.2.1 Temporal TSS anomalies</u>

493 Considering the temporal variation recorded across monsoons, maps of relative TSS anomalies were 494 calculated for each year as the difference with respect to the long-term TSS mean (Fig. 5), in order to 495 detect changes of TSS distribution occurring annually (Fig. 9). As shown in Figure 9, year 2010 496 experienced a distinct increase of TSS distribution (approximately 100 %), with widespread pattern 497 extending into open ocean waters, in comparison to the long-term TSS mean. This finding provides an interesting insight into the effects of extreme rainfall events as recorded in year 2009, which could 498 499 potentially intensify TSS release into the coastal and open ocean waters. The effects of TSS release can still be seen a year after the extreme rainfall events in this region. This observation could provide 500 501 further evidence that the impacts of the TSS release from the land into rivers and coastal systems may 502 only take effect after a substantial period, as previously observed by Sun et al. (2017a).

503 Figure 9 further reveals an interesting pattern of TSS increase in the Samunsam-Sematan region from 504 year 2004 until 2019, with exceptions during the years 2007 and 2008. As previously highlighted in 505 Section 3.1, the Samunsam-Sematan region has been observed to be a vulnerable coastal area with respect to TSS water quality degradation. From the annual map of TSS anomalies (Fig. 9), we can see 506 507 that the TSS distribution has the tendency to accumulate in the Samunsam-Sematan region, as 508 opposed to being distributed into the open ocean waters. This may be due to the geographical and 509 hydrological characteristics of these coastal regions (Martin et al., 2018), as the TSS release may be sheltered from open ocean waters, and hence induce a higher TSS accumulation in these coastal 510 511 regions.

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518 Fig. 9: Map of relative TSS distribution anomalies with respect to the long-term mean, represented as percentage (%), from

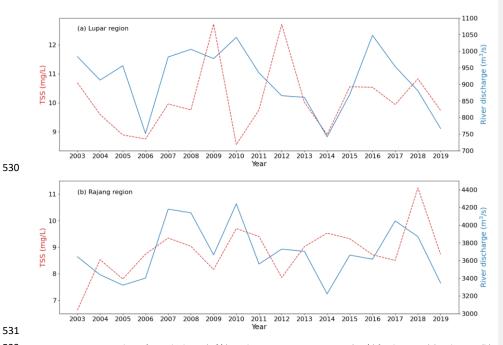
519 year 2003 until 2019.

# 520 3.3 Hydrological factors driving TSS discharge

Apart from the influence of monsoonal patterns, hydrological factors such as the river discharge are
among the dominant drivers in transporting various water constituents in riverine and coastal systems
(Loisel et al., 2014; Petus et al., 2014; Sun, 2017b; Verschelling et al., 2017). In this study, river
discharge from the Lupar and Rajang basins was estimated and investigated.

Yearly river discharge estimates from 2003 until 2019 were investigated to assess its effect on the TSS distribution (Fig. 10) represented by changes in TSS values for pixels located at each Lupar and Rajang coastal river points (Supplementary Materials, Fig. S3). Figure 10a shows that river discharge values in the Lupar basin (750 to 1050 m<sup>3</sup>/s) are approximately twice lower than the Rajang river discharge (Fig. 10)

529 10b), which recorded a range of 3,200 to 4,000 m<sup>3</sup>/s.



532 Fig. 10: Time-series analysis of river discharge (m<sup>3</sup>/s) in relation to TSS concentrations (mg/L) for the Lupar (a) and Rajang (b)

basins from year 2003 to 2019. Note the differing scaling on the ordinate axes in each plot.

534 DFrom our findings, discrepancies between TSS estimates and river discharge were identified in both 535 the Lupar and Rajang coastal regions in these annual time-series, where river discharge was inversely 536 correlated with TSS estimates. These discrepancies are not uncommon, as previously highlighted in a 537 study by Zhan et al. (2019). Especially in 2010 for the Lupar river, Fig. 10a shows a drop in TSS release 538 in relation to the steady increase of river discharge from the river basin. In 2011 and 2012, a negative 539 correlation can be seen between river discharge and TSS estimates, while in subsequent years from 2013 until 2015, there is a clear positive correlation. Although there is no obvious environmental 540 factor that would explain these discrepancies, these findings may imply a complex interaction 541 542 between human interventions, such as damming and deforestation activities, as well as varying hydrological and atmospheric conditions (wind and tidal mixing) in regulating TSS dynamics (Wu et al., 543 544 2012; Zhan et al., 2019).

545 The correlation between TSS release and river discharge at both the Lupar and Rajang coastal areas 546 was further evaluated in this study. Even though river discharge has been shown (in other global 547 studies) to be one of the dominant factors in moderating TSS release (Fabricius et al., 2016; Tilburg et 548 al., 2015; Verschelling et al., 2017; Wu et al., 2012), the TSS distribution at both the Lupar and Rajang river mouths in this study can be seen to be only poorly correlated with river discharge from each river 549 550 basin (Supplementary Materials, Fig. S8a and b). The TSS output from the Lupar basin recorded a 551 correlation coefficient of r = 0.15, while river discharge from the Rajang basin did not substantially 552 influence the TSS release either, with r = 0.27 throughout the seasons. Coupled with tidal mixing processes (Ramaswamy et al., 2004; Zhou et al., 2020) , it is possible that human activities such as 553 554 deforestation, logging, and construction of dams, which are largely occurring within the Rajang basin 555 (Alamgir et al., 2020), are mainly driving TSS release and resuspension in this area. This indicates that although TSS release is regarded to be highly dependent on, and controlled by river discharge patterns, 556 557 this interaction often represents an intricate process linked to local hydrodynamics process and socio-558 economic conditions (Espinoza Villar et al., 2013; Fabricius et al., 2016; Valerio et al., 2018; Zhan et al., 559 2019).

## 560 3.4 Variability of TSS across coastal waters

As previously observed in Fig. 5, varying river plumes of TSS were evidently detected within the coastal regions of the study area. Notably, coastal river plumes represent important factors driving the transport of water constituents and nutrients from coastlines to the open oceanic systems (Petus et al., 2014). To assess this and evaluate the water quality status in coastal zones, the spatial extent of TSS release was investigated along transects covering the territorial (12 nautical miles) and open water areas (24 nautical miles) of the Sarawak region (Fig. 11).

A total of eight coastal points were selected based on the main river mouths located in the southwest region of Sarawak. Transect points are positioned in a line starting at the coastal river points to examine the variations of TSS distribution across different water zones. Daily changes in TSS concentration for each pixel located in front of the river mouths were plotted from 2003 until 2019 (Fig. 12).

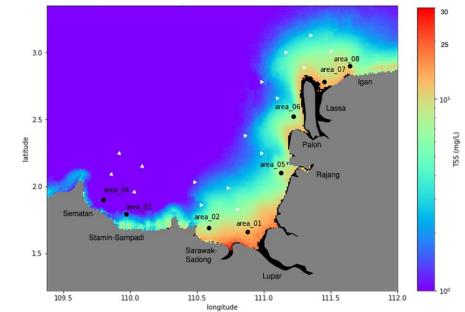


Fig. 11: Map of average TSS estimates (mg/L) with indicators at eight main river mouths and their transect, extending from
coastal waters into territorial and open ocean systems. Indicators of each river mouths are as follows: area\_01 – Lupar river;
area\_02 – Sarawak-Sadong river; area\_03 - Stamin-Sampadi river i; area\_04 – Sematan river; area\_05: Rajang river; area\_06:

576 Paloh river; area\_07: Lassa river; and area\_08: Igan river).

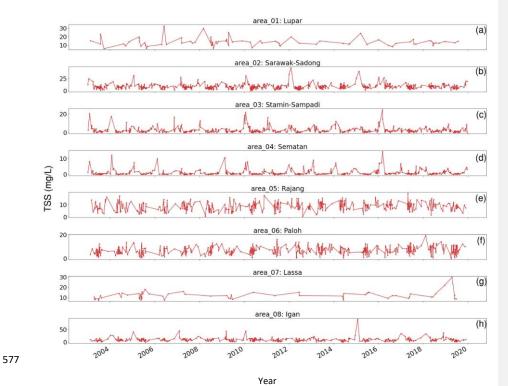


Fig. 12: Graphs of daily TSS estimates (mg/L) recorded at eight river mouth points from 2003 to 2019. Presentation of each
river mouths is as follows: a) area\_01; b) area\_02; c) area\_03; d) area\_04; e) area\_05; f) area\_06; g) area\_07; h) area\_08.
Note the different TSS scales in each plot.

From the high temporal resolution graphs in Fig. 12, no general trend of TSS concentration can be identified over the years at each coastal point. It is worth highlighting that the daily temporal resolution was particularly affected at coastal points located in front of the Lupar (area\_01) and Lassa (area\_07) river mouths due to various pixel data quality issues in these areas. Nonetheless, more than 80 satellite images with minimum cloud coverage at these two locations were processed, while the remaining coastal points had a total of more than 400 satellite images to assess the temporal trend.

Despite the fact that no distinct upward or downward trend was observed, our findings indicate that several river mouths are actively discharging and accumulating substantial TSS amounts over the period of years, while resuspension of bottom sediments induced by wind and tidal cycle is another factor contributing to the variation of TSS values (Park, 2007; Song et al., 2020).

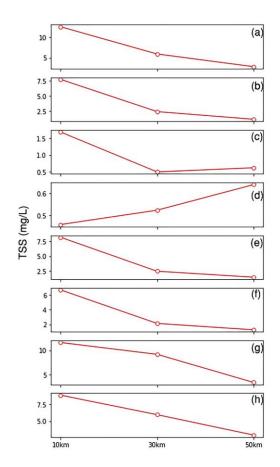
591 The coastal region of the Sarawak-Sadong river (area\_02) shows relatively high TSS distribution 592 patterns with some periods recording an estimate of over 30 mg/L of TSS concentration. This is in agreement with the localised characteristics of the Sarawak river basin which essentially drains 593 594 through the populated Kuching area with high industrial and development activities in the capital city 595 of Sarawak (DID, 2021b). In comparison with other river mouth points, a steady TSS concentration 596 below 20 mg/L was recorded across the Stamin-Sampadi (area\_03), Sematan (area\_04), Rajang (area\_05), and Paloh (area\_06) river mouths. Consistently high TSS values in the daily plots were 597 598 recorded at the Lupar (area\_01) and Pulau Bruit-Lassa (area\_07) river mouths, with estimates of up to 599 30 mg/L on a near-daily basis. Similar high TSS amounts from the Igan (area\_08) river mouth, situated 600 northeast side of the Pulau Bruit-Lassa region, were observed in Cherukuru et al. (2021) and Staub et 601 al. (2000).

Although the daily TSS estimates at each river point are in line with various reported studies (Chen et al., 2011, 2015b; Kim et al., 2017; Mengen et al., 2020; Zhang et al., 2010a), these estimates can be expected to be much higher for sampling points much closer to the river mouths. The selection of coastal river points in this study was made to minimize the gaps with respect to various pixel data quality issues in the MODIS-Aqua datasets, and hence, the use of coastal river points closer to shore would have been impractical.

These findings further suggest that higher TSS loadings within the coastal river areas would have been diluted or deposited while travelling to the open oceanic systems as they are weakly impacted by river discharge in relation to offshore distance (Espinoza Villar et al., 2013). This understanding can be observed in Fig. 13, which shows a progressively decreasing TSS estimates at each transect in relation

	612	to the distance from the shore. Generally, TSS estimates in coastal zones (first transect point) show
	613	considerably higher TSS concentrations. When moving outwards to territorial waters (second transect
	614	point), TSS concentration estimates decrease by nearly 50 % before travelling to open ocean systems
	615	(third transect point), except for the northeast regions (area_07 and area_08) which seem to show
	616	large extension of TSS plumes to the open ocean waters, as also highlighted by Cherukuru et al. (2021).
	617	A reversed trend can be seen in the plot corresponding to the Sematan coastal river systems, although
	618	the absolute increase in TSS estimates across water zones (0.2 mg/L in total) here is only marginal (Fig.
	619	1 <u>3</u> 5d).

620	Nonetheless, it is important to present these synoptic findings to address existing water quality
621	practices and regulations implemented in this local region, in order to improve the monitoring and
622	management of its coastal systems. According to the Malaysia Marine Water Quality Criteria and
623	Standard (Supplementary Materials, Table S2), the coastal areas of Sarawak are dominantly
624	categorised as Class I quality. Our findings in the southwest coastal areas (Sematan and Stamin-
625	Sampadi) showed that the coral reefs there can be well-maintained with negligible impacts from TSS
626	loadings (Fig. 11). While these Sematan and Stamin Sampadi coasts can be seen to be in a healthy
627	water quality state (low average TSS concentration, see e.g. Fig 11), the high coefficient variation
628	reported in these coastal waters, as previously highlighted in Section 3.1 (Fig. 6), clearly stresses the
629	importance of understanding how the quality of coastal systems can vary and be affected by human
630	intervention and changing landscapes over time.



637 Fig. 13: Average TSS estimates (mg/L) computed from 2003 to 2019 for each of the eight rivers (area\_01: (a), area\_02: (b);

638 area\_03: (c); area\_04: (d); area\_05: (e); area\_06: (f); area\_07: (g); and area\_08: (h)) and their relevant transect points with

639 di	istance of 10 km	(coastal waters), 30 km	(territorial waters) and	d 50 km (open ocean waters)	from the shoreline. Note the
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640 varying TSS scales on the ordinate	axes in each plot.
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#### 645 3.5 Discussion of TSS implications for coastal waters

High discharge of TSS into coastal environments can lead to adverse environmental and ecological implications. The presence of TSS affects water transparency and light availability within the surface waters (Dogliotti et al., 2015; Nazirova et al., 2021; Wang et al., 2021). Among others, TSS affects the photosynthesis activities of algae and macrophytes. TSS in water creates a reduction in light penetration, which impacts the primary production of aquatic organisms and hence the support system of marine life (Bilotta and Brazier, 2008; Loisel et al., 2014).

652 Additionally, TSS exerts an influence on zooplankton communities. Reduction in water clarity induces 653 changes in the zooplankton's biomass volume and composition, while TSS may carry a level of toxicity 654 which affects zooplankton through ingestion (Chapman et al., 2017; Donohue and Garcia Molinos, 655 2009). Apart from that, accumulation and deposition of sediments decreases the level of dissolved 656 oxygen (DO) at the bottom of the water column, and subsequently impacts the benthic invertebrate groups (Chapman et al., 2017). Moreover, substantial TSS deposition tends to cause harmful physical 657 658 effects to these benthic groups, such as abrasion, and even clogging by sediment particles (Chapman 659 et al., 2017; Langer, 1980).

As a result of these TSS effects on lower trophic levels, fish communities are critically impacted, with a reduction in diversity and abundance (Kemp et al., 2011). While fish communities learn to adapt to a range of TSS loads (Macklin et al., 2010), increase of TSS concentrations often depletes DO concentrations in the water system and causes stress towards these aquatic communities (Henley et al., 2000). Fish populations tend to decrease, as feeding and growth rates are negatively impacted (Shaw and Richardson, 2001; Sutherland and Meyer, 2007).

Threats to coral reefs have been linked to sediment-induced stress which often leads to a reduction in the coral's growth and metabolic rate, as well as impending mortality (Erftemeijer et al., 2012; Gilmour et al., 2006; Risk and Edinger, 2011). Factors of coral stress are driven by nutrient-rich sediments and microbes which are being carried by TSS, with impacts on the health of coral tissues

(Hodgson, 1990; Risk and Edinger, 2011; Weber et al., 2006). A reduction in light availability impedes the development of corals (Anthony and Hoegh-Guldberg, 2003; Rogers, 1979; Telesnicki and Goldberg, 1995). A combined increase in TSS and nutrient loadings contribute to the decrease of coral species diversity and composition (Fabricius, 2005).

Essentially, presence of TSS in water systems has impacts across various aquatic biota. With the severe implications of decreased fish population, this could lead to a disruption of fisheries activities by local communities, especially considering that more than 80 % of the Sarawak population is living in the coastal areas (DID, 2021a). Coral reefs are important coastal biodiversity assets to the Sarawak region, especially around the Talang-Talang and Satang islands on the southwest coast of Sarawak (Long, 2014). With the use of remote sensing technologies in monitoring Sarawak coastal water quality, the approach presented in this paper provides digital-based solutions to assist relevant authorities and local agencies to better manage the Sarawak coastal waters and their resources. 

### 696 4.0 Conclusion

In this study, a regional empirical TSS retrieval model was developed to analyse TSS dynamics along the southwest coast of Sarawak. The empirical relationship between in situ reflectance values, Rrs( $\lambda$ ), and in situ TSS concentrations was established using a green-to-red band ratio using the MODIS-Aqua Rrs(530) and Rrs(666) reflectance bands. An evaluation of the TSS retrieval model was carried out with error metric assessment, which yielded results of bias = 1.0, MAE = 1.47 and RMSE = 0.22 in mg/L computed in log10-transformed space prior to calculation. A statistical analysis using a k-fold cross validation technique (k = 7) reported  $\frac{1}{2}$  low error metrics (RMSE = 0.2159, MAE = 0.1747).

704 The spatial TSS distribution map shows widespread TSS plumes detected particularly in the Lupar and 705 Rajang coastal areas, with average TSS range of 15 - 20 mg/L estimated at these coastal areas. Based 706 on the spatial map of the TSS coefficient of variation, large TSS variability was identified in the 707 Samunsam-Sematan coastal areas (CV > 90 %). The map of temporal variation of TSS distribution 708 points to a strong monsoonal influence in driving TSS release, with large differences identified 709 between the northeast and southwest monsoon periods in this region. From the annual TSS anomalies maps, the Samunsam-Sematan coastal areas demonstrated strong TSS variation spatially, while 710 711 widespread TSS distribution with nearly 100 % of TSS increase in comparison to long term mean was 712 observed in 2010. Furthermore, our study on river discharge in relation to TSS release demonstrated 713 a weak relationship at both the Lupar and Rajang coastal river points. Study on the TSS variability 714 across coastal river mouths implied that higher TSS loadings in the coastal areas are potentially being 715 deposited or diluted in the process of being transported into the open ocean waters, with varying 716 magnitude at several coastal river points.

'17	Overall, these coastal zones remain within local water quality standards to support various marine and
18	socio economic activities in this regioncoastal areas of Sarawak are dominantly categorised as Class
19	I quality, which remain within local quality standards to support various marine and socio-economic
20	activities in this region. Our findings in the southwest coastal areas (Sematan and Stamin-Sampadi)
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721 showed that the coral reefs there can be well-maintained with negligible impacts from TSS loadings. 722 However, it is important to highlight the various human activities that are widely ongoing in this 723 region, which include deforestation and logging activities (Alamgir et al., 2020; Hon and Shibata, 2013; 724 Vijith et al., 2018). Impacts from these activities in Sarawak can potentially aggravate current soil 725 erosion issues, and ultimately induce more soil leaching and runoff from land to water systems, 726 especially during heavy rainfall events (Ling et al., 2016; Vijith et al., 2018). As a result, human activities 727 may have a greater influence on driving riverine sediments than climatological factors, as reported by Song et al. (2016). As such, this work presents the first observation of TSS distributions at large spatial 728 729 and temporal scales in Sarawak's coastal systems, and of the potential associated impacts on the South China Sea. The findings derived from this work can be used to support local authorities in 730 731 assessing TSS water quality status in the coastal areas of concern, and to enhance coastal management 732 and conservation strategies. The application of remote sensing technologies is of great benefit in the 733 development of sustainable sediment management in the Sarawak coastal region, as demonstrated 734 in this study.

735 Data availability. The dataset related to this study is available as supplement to this paper.

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