

1 Spatial and temporal dynamics of suspended sediment concentrations in coastal waters of South  
2 China Sea, off Sarawak, Borneo: Ocean colour remote sensing observations and analysis

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## 12 Abstract

13 High-quality ocean colour observations are increasingly accessible to support various monitoring and  
14 research activities for water quality measurements. In this paper, we present a newly developed  
15 regional total suspended solids (TSS) empirical model using MODIS-Aqua's Rrs(530) and Rrs(666)  
16 reflectance bands to investigate the spatial and temporal variation of TSS dynamics along the  
17 southwest coast of Sarawak, Borneo. The performance of this TSS retrieval model was evaluated using  
18 error metrics (bias = 1.0, MAE = 1.47, and RMSE = 0.22 in mg/L) with a log<sub>10</sub> transformation prior to  
19 calculation, as well as a k-fold cross validation technique. The temporally averaged map of TSS  
20 distribution, using daily MODIS-Aqua satellite datasets from 2003 until 2019, revealed large TSS  
21 plumes detected particularly in the Lupar and Rajang coastal areas on a yearly basis. The average TSS  
22 concentration in these coastal waters was in the range of 15 – 20 mg/L. Moreover, the spatial map of  
23 TSS coefficient of variation (CV) indicated strong TSS variability (approximately 90 %) in the  
24 Samunsam-Sematan coastal areas, which could potentially impact nearby coral reef habitats in this  
25 region. Study on temporal TSS variation provide further evidence that monsoonal patterns drive the  
26 TSS release in these tropical water systems, with distinct and widespread TSS plume variations  
27 observed between the northeast and southwest monsoon periods. A map of relative TSS distribution  
28 anomalies revealed strong spatial TSS variations in the Samunsam-Sematan coastal areas, while 2010  
29 recorded a major increase (approximately 100 %) and widespread TSS distribution with respect to the  
30 long-term mean. Furthermore, study on the contribution of river discharge to the TSS distribution  
31 showed a weak correlation across time at both the Lupar and Rajang river mouth points. The variability  
32 of TSS distribution across coastal river points was studied by investigating the variation of TSS pixels  
33 at three transect points, stretching from the river mouth into territorial and open water zones, for  
34 eight main rivers. Results showed a progressively decreasing pattern of nearly 50 % in relation to the  
35 distance from shore, with exceptions in the northeast regions of the study area. Essentially, our  
36 findings demonstrate that the TSS levels at the southwest coast of Sarawak are within local water  
37 quality standards, promoting various marine and socio-economic activities. This study presents the  
38 first observation of TSS distributions at Sarawak coastal systems with the application of remote  
39 sensing technologies, to enhance coastal sediment management strategies for the sustainable use of  
40 coastal waters and their resources.

41 **Keywords:** total suspended solids, band-ratio, monsoon, river discharge, Open Data Cube

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## 45 1.0 Introduction

46 Total Suspended Solids (TSS) play an important role in the aquatic ecosystem as one of the primary  
47 water quality indicators of coastal and riverine systems (Alcântara et al., 2016; Cao et al., 2018; Chen  
48 et al., 2015a; González Vilas et al., 2011; Mao et al., 2012). For example, elevated concentrations of  
49 TSS in water have an adverse impact on fisheries and biodiversity of the aquatic ecosystem (Bilotta  
50 and Brazier, 2008; Chapman et al., 2017; Henley et al., 2000; Wilber and Clarke, 2001). Understanding  
51 the impacts of varying water quality in relation to TSS status has been one of the primary concerns  
52 with respect to a country's growing Blue Economy status and sustainable management of aquatic  
53 resources (Lee et al., 2020a; Sandifer et al., 2021; World Bank and United Nations Department of  
54 Economic and Social Affairs (UNDESA), 2017). With about 40 % of the world's population living within  
55 100 km of coastal areas (United Nations, 2017), and with more than 80 % of the population in Malaysia  
56 living within 50 km of the coast (Praveena et al., 2012), water quality monitoring and management  
57 efforts are important at both regional and global scale.

58 Studying TSS distribution can provide insights into the connections between land and ocean  
59 ecosystems (Howarth, 2008; Lemley et al., 2019; Lu et al., 2018). For instance, TSS dynamics allow us  
60 to understand the impacts of sediment transport and sediment plumes, particularly in areas  
61 experiencing large-scale deforestation, land conversion and damming of rivers (Chen et al., 2007;  
62 Espinoza Villar et al., 2013). Sarawak, Malaysian Borneo, experienced significant land use and land  
63 cover change activities over the past four decades, with widespread land conversion and deforestation  
64 for developments and large-scale plantation activities (Gaveau et al., 2016), as well as building of  
65 major road infrastructures, such as the Pan-Borneo highway, and hydroelectric dams (Alamgir et al.,  
66 2020). As a result, river and coastal systems may potentially drive large TSS loads into downstream  
67 systems and into the marine and open ocean waters.

68 Situated at the southern part of the South China Sea, the region of Sarawak, Malaysian  
69 Borneo, has a coastline of about 1035 km where mangrove forests are dominant (Long, 2014). The

70 coastal regions of Sarawak are rich with marine coastal biodiversity and coral reefs, which can be  
71 found at the northeast and southwest part of Sarawak (Praveena et al., 2012). While the coasts of  
72 Sarawak provide important socio-economic values to the local communities (Lee et al., 2020b), these  
73 coastal areas are potentially facing water quality degradation from TSS riverine outputs in response  
74 to land use and land cover change activities.

75 TSS concentrations are commonly measured through conventional laboratory-based methods  
76 to quantify TSS concentrations by field collection of water samples (Ling et al., 2016; Mohammad Razi  
77 et al., 2021; Soo et al., 2017; Soum et al., 2021; Tromboni et al., 2021; Zhang et al., 2013). Currently,  
78 real-time high-frequency TSS observations using modern optical and bio-sensor systems are also  
79 possible (Bhardwaj et al., 2015; Horsburgh et al., 2010). These sensors can be generally found onboard  
80 ship and buoy-based observation platforms. Yet, it remains a challenge to quantify TSS concentrations  
81 of large spatial coverage and high temporal frequency with these approaches.

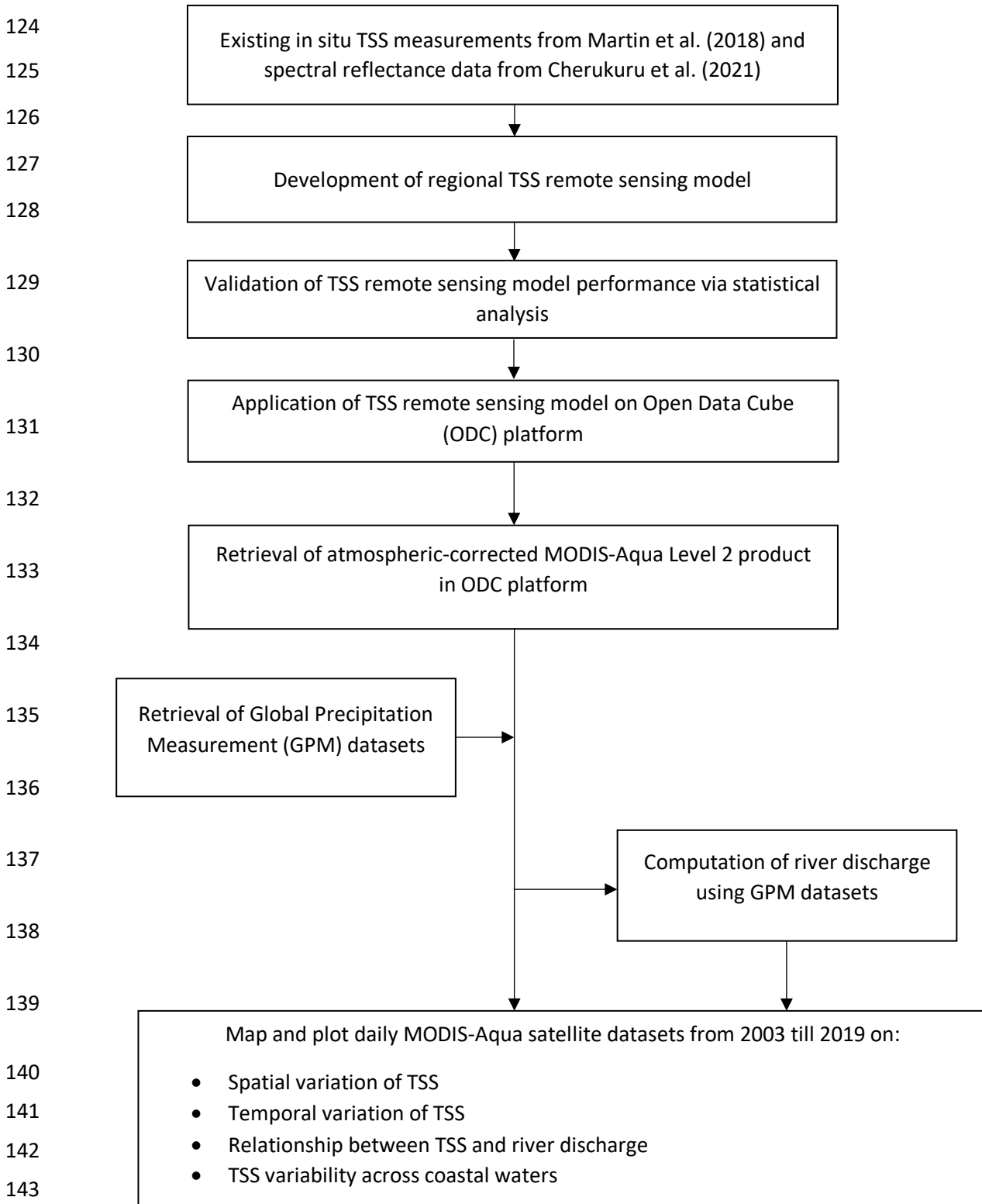
82 Ocean colour remote sensing technologies represent an increasingly accessible and powerful  
83 tool to provide a synoptic view for short or long-term water quality studies at high temporal and spatial  
84 resolutions (Cherukuru et al., 2016a; Slonecker et al., 2016; Swain and Sahoo, 2017; Wang et al., 2017;  
85 Werdell et al., 2018). Remote sensing can help overcome several constraints of conventional intensive  
86 field campaigns such as: (i) costly field campaigns from boat rentals or cruise; (ii) time-consuming and  
87 inadequate manpower; and most importantly for this study, (iii) limited spatial and temporal field  
88 coverage. NASA's Moderate Resolution Imaging Spectroradiometer (MODIS)-Aqua  
89 (<https://modis.gsfc.nasa.gov/about/>) has a distinctive advantage with its daily revisit time, a spatial  
90 resolution of 250 – 1000 m, and a large collection of ocean colour data since 2002. Other sensors  
91 offering ocean colour measurement capabilities include Landsat-8, which, in comparison with MODIS-  
92 Aqua, has a 16-day revisit time and high spatial resolution of 30 m. Despite Landsat's powerful ability  
93 in capturing higher resolution images, the longer revisit interval may not be suitable for characterizing  
94 and studying water bodies with high dynamics of various water constituents.

95           Several MODIS-derived models have been developed for TSS retrievals (Chen et al., 2015b;  
96 Espinoza Villar et al., 2013; Jiang and Liu, 2011; Kim et al., 2017; Zhang et al., 2010b), including  
97 empirical, semi-analytical and machine-learning approaches (Balasubramanian et al., 2020; Jiang et  
98 al., 2021). However, the performance of these models proved to be less satisfactory, with recorded  
99 low  $r^2$  and high bias and mean absolute error (MAE) values when tested with in situ TSS datasets  
100 (Supplementary Materials, Table S1). While these global TSS remote sensing models address the need  
101 to improve TSS retrievals and to monitor global TSS trends in various water class types, they tend to  
102 underperform in more localised and regional studies (Mao et al., 2012; Ondrusek et al., 2012). The  
103 coastal waters of Borneo are well-mixed throughout the year and enriched with suspended material  
104 and dissolved organic matter (Müller et al., 2016). Various water quality studies of the river systems  
105 have been actively carried out to assess the dynamics of numerous water quality constituents in  
106 response to human activities, with TSS concentrations being one of the primary environmental  
107 concerns in this region (Ling et al., 2016; Müller-dum et al., 2019; Tawan et al., 2020). Although studies  
108 on the water quality of coastal systems in Borneo have gradually gained much attention (Cherukuru  
109 et al., 2021; Limchih et al., 2010; Martin et al., 2018; Soo et al., 2017), there is still much knowledge to  
110 gain on the understanding of how coastal waters in the region have been impacted by TSS loadings  
111 and transport over large spatial and temporal scales.

112           Here, in this paper, we present a new regional empirical TSS remote sensing model. While  
113 various remote sensing models have their own unique computational strengths, this study  
114 demonstrates the reliability of band ratio TSS model to be applied within optically complex waters.  
115 With the ongoing efforts to address and minimize water quality degradation in coastal systems, as  
116 outlined in the United Nation's Sustainability Development Goals no. 14, our study aims to apply the  
117 new empirical regional TSS remote sensing model to: (a) investigate the spatial and temporal  
118 variability in TSS, (b) identify hotspots of TSS distribution in the coastal waters of Sarawak, Malaysian  
119 Borneo, using a long time series of MODIS-Aqua data from year 2003 until 2019, and (c) study the  
120 varying monsoonal and river discharge patterns in relation to TSS distribution at river mouths.

121 2.0 Methodologies

122 Figure below summarizes the processes carried out in this study. Spatial and temporal variation of TSS  
123 distribution was mapped using atmospheric-corrected MODIS-Aqua Level 2 product.



144 Fig. 1: Flowchart summarizing the processes of developing a regional TSS remote sensing model and applying it to analyse  
145 the spatial and temporal variation of TSS over the study region, using MODIS-Aqua data from year 2003 until 2019. Long-  
146 term MODIS-Aqua datasets were analysed and mapped on an Open Data Cube (ODC) platform with implementation of robust  
147 Python libraries and packages.

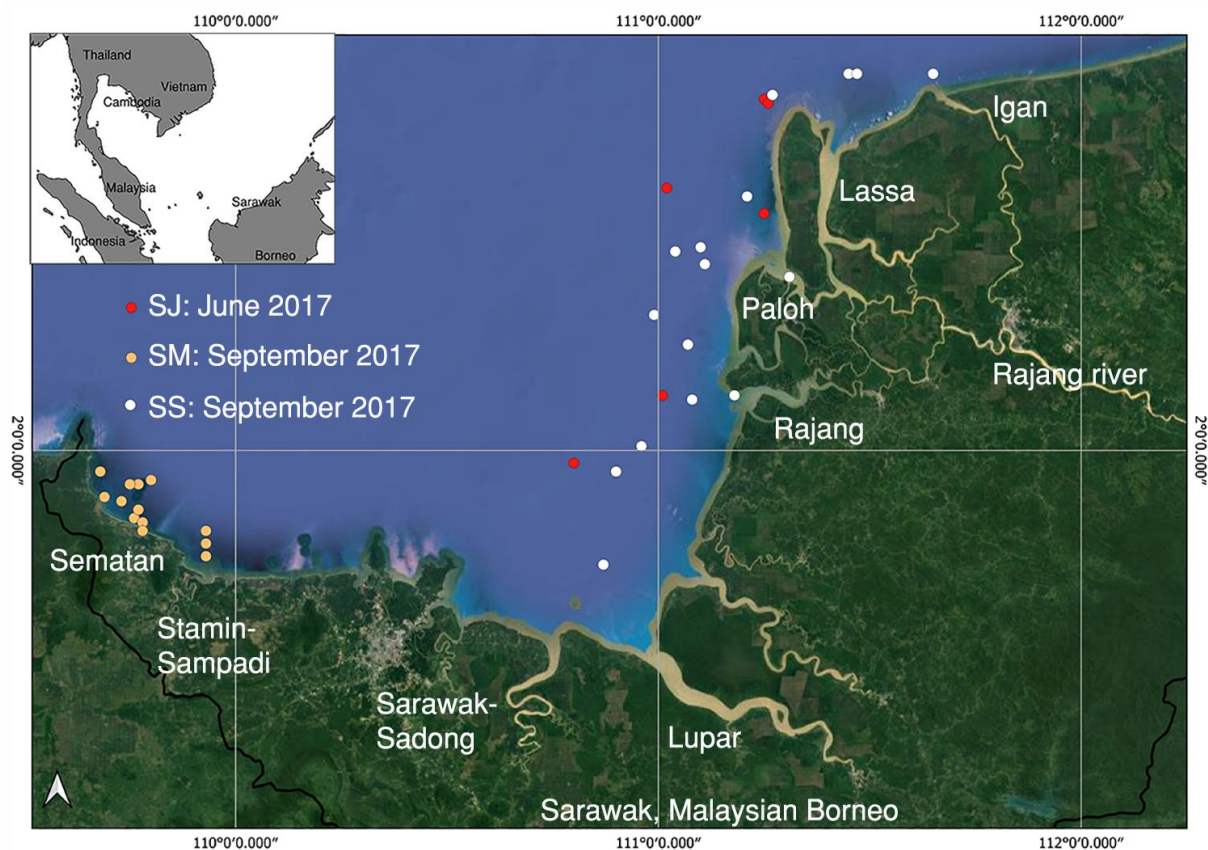
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## 149 2.1 Area of study

150 Our study focuses on the southwestern coast of Sarawak (between 1.9° N, 109.65° E and 2.8° N, 111.5°  
151 E) in Malaysia, which sits at the northwest part of the Borneo Island. Generally, the island of Borneo  
152 (between 3.01° S, 112.18° E and 6.45° N, 117.04° E) contains rich tropical rainforests and biodiversity  
153 on the lands of Sarawak and Sabah (Malaysia), Brunei, and Indonesia. Typically, Sarawak is a tropical  
154 climate region, recording an average ambient temperature of 27.8 °C (variation of 1.8 °C) throughout  
155 the year. It records high precipitation with an average of 4116.7 mm/yr in Kuching (1.5535° N,  
156 110.3593° E), the capital city of Sarawak. Yearly, it experiences both a dry and wet season, which is  
157 influenced by: (i) the southwestern monsoon (May to September) and (ii) the northeastern monsoon  
158 (November to March). Rivers in Sarawak are connected to the South China Sea and flow through  
159 various plantation types, such as oil palm, rubber and sago (Davies et al., 2010).

160 In this study, the southwestern part of Sarawak's coastal regions (Fig. 2), (between 1.9° N, 109.65° E  
161 and 2.8° N, 111.5° E) was studied, which comprise several major rivers (e.g. Lupar, Sebuyau, Sematan),  
162 as well as the Rajang River, the longest river in Malaysia. Rajang river basin consists in tidally influenced  
163 river channel which splits into a northwest (Igan, Lassa and Paloh) and a southwest (Rajang, Belawai)  
164 Rajang river delta (Staub et al., 2000). The Rajang river basin drains a dominant area (>50,000km<sup>2</sup>) of  
165 sedimentary rocks (Milliman and Farnsworth, 2013; Staub et al., 2000) extending from Belaga to Sibul,  
166 with major peatland areas converted into oil palm plantations (Gaveau et al., 2016) as its river flows  
167 into the South China Sea (Milliman and Farnsworth, 2013). Major settlements along the Rajang river  
168 comprise of Kapit and Kanowit town areas, as well as Sibul city, with a total population size of about  
169 388,000 inhabitants (Department of Statistics, 2020). Lupar and Saribas rivers, respectively, comprise  
170 a catchment area size of approximately 6500 and 1900 km<sup>2</sup> (Lehner et al., 2006). Situated at the  
171 southwest side of the Rajang catchment, Lupar and Saribas rivers surround the Maludam National

172 Park, which is Sarawak's remaining biggest single patch of peat swamp forest (Sarawak Forestry  
173 Corporation, 2022). Adjacent to Lupar river mouth is the Sadong river, with an approximate catchment  
174 area size of 3500 km<sup>2</sup> (Kuok et al., 2018). Sadong river runs about 150 km and flows through oil palm  
175 plantations (Staub and Esterle, 1993). These river systems are associated with increasing land use  
176 activities and land cover changes in this region, which essentially transport and connect various  
177 biogeochemical water components to the coastal systems of Sarawak.



178  
179 Fig. 2: Map of the study area (© Google Maps), located in the southwestern part of Sarawak, Malaysia (inset). Indicators  
180 show the location of sampling sites used during field expeditions carried out in June and September 2017.

## 181 2.2 In situ TSS measurements

182 TSS measurements were taken from Martin et al. (2018). A total of 35 coastal sites were studied and  
183 are denoted SJ, SS, and SM (see: Table 1 & Fig. 2). These water samples were collected in the month  
184 of June (SJ region) and September (SS and SM regions) in 2017. Water samples were filtered, and

185 filters were dried and ashed prior to weighing process. Full details of water sampling and TSS analysis  
186 is available in Martin et al. (2018).

### 187 2.3 Development, calibration and validation of TSS model

188 In situ remote sensing reflectance spectral data,  $R_{rs}(\lambda)$ , along with 35 measured TSS values, were used  
189 to develop a new remote sensing TSS empirical model for MODIS-Aqua for this case study. Field  
190 measurements of SM, SJ & SS datasets, as shown in Table 1, were used to calibrate the MODIS-Aqua  
191 TSS remote sensing model.

192 For the in situ remote sensing reflectance,  $R_{rs}(\lambda)$  readings, a TriOS-RAMSES spectral imaging  
193 radiometer was used to measure downwelling irradiance,  $E_d(\lambda)$ , and upwelling radiance,  $L_u(\lambda)$ , with  
194 measurement protocols from Mueller et al. (2002). These measurements were recorded under stable  
195 sky and sea conditions during the day (10AM to 4PM) with high solar elevation angles.

196 Measurements of reflectance,  $R_{rs}(\lambda)$ , were recorded concurrently with the collection of water samples  
197 (as described in Section 2.2) and were recorded at wavelength ranging from 280 to 950 nm, which  
198 covers the spectrum of ultraviolet, visible and visible/ultraviolet light. These measurements were  
199 recorded on a float to capture  $L_u(0-, \lambda)$  and  $E_d(0+, \lambda)$ , where 0- and 0+ refer to below-surface and  
200 above-surface, respectively.

201 Remote sensing reflectance,  $R_{rs}(\lambda)$ , was computed as follows with reference to Mueller et al. (2002):

$$R_{rs}(\lambda, 0+) = \frac{1 - p}{n^2} \times \frac{L_u(0-, \lambda)}{E_d(0+, \lambda)} \quad (1)$$

202 where  $p = 0.021$  refers to the Fresnel reflectance and  $n = 1.34$  is the refractive index of water. Full  
203 details of this methodology can be found at Cherukuru et al. (2021).

#### 204 2.3.1 Calibration of empirical model and application to MODIS-Aqua



205 With the intention to apply a regional TSS remote sensing model to MODIS-Aqua product, a total of  
 206 35 different datasets of TSS concentrations were collected in coastal conditions (salinity > 15 PSU)  
 207 and convolved to generate MODIS-Aqua data.

208 In this study, retrieval of water constituents was established using spectral band ratio combinations  
 209 which have proven to be a straightforward, yet reliable method for estimating water constituents in  
 210 optically turbid waters (Ahn and Shanmugam, 2007; Cao et al., 2018; Lavigne et al., 2021; Morel and  
 211 Gentili, 2009; Neil et al., 2019; Siswanto et al., 2011). Band ratio models help to offset signal noise,  
 212 such as the effects of atmospheric and irradiance of spectral reflectance to a certain degree  
 213 (Cherukuru et al., 2016b; Ha et al., 2017; Hu et al., 2012; Liu et al., 2019).

214 A variety of models using single bands, as well as a combination of MODIS-Aqua’s Blue, Green & Red  
 215 bands (412nm, 440nm, 488nm, 532nm, 555nm & 660nm) were calibrated using field measurements  
 216 as dependent variable. The calibration process was tested out using various model functions, including  
 217 linear, power, exponential, and logarithm functions. The best empirical TSS retrieval model was fitted  
 218 by means of a regression between the in situ TSS data and in situ radiometer values, and can be  
 219 expressed as follows:

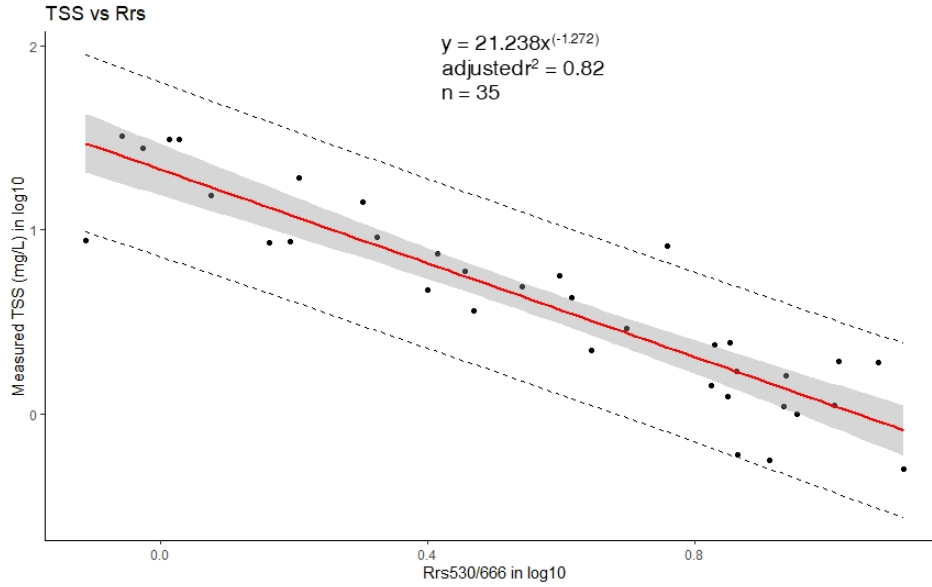
$$TSS = 21.238[Rrs(530)/Rrs(666)]^{-1.272} \quad (2)$$

221 This power function model resulted in a coefficient of determination ( $R^2$ ) of 0.82 (Fig. 3).

222 Table 1: Summary statistics of TSS values collected at areas SJ, SS, and SM located within coastal regions in this study, with a  
 223 total of 35 datasets recorded.

<b>Coastal Area</b>	<b>Minimum</b>	<b>Maximum</b>	<b>Mean</b>	<b>S.D.</b>	<b>C.V.</b>	<b>n</b>
<b>SJ</b>	1.1	19.24	6.89	6.62	96.09	6
<b>SS</b>	0.56	32.1	12.50	11.43	91.45	16
<b>SM</b>	0.5	8.14	2.59	2.70	104.53	13

224



225

226 Fig. 3: Empirical relationship of TSS retrieval between in situ Rrs(530)/Rrs(666) bands ratio and measured TSS  
 227 concentration (mg/L), as established via a power law function. Upper and lower dashed lines indicate the 95 % prediction  
 228 interval of the regression.

229 2.3.2 Performance assessment and validation of MODIS-Aqua empirical model

230 An assessment of the performance error of the newly developed TSS model was carried out as per  
 231 Seegers et al. (2018)'s recommendation for interpreting ocean colour models. These performance  
 232 metrics used here include the bias, Mean Absolute Error (MAE), Root Mean Squared Error (RMSE),  
 233 coefficient of variation (CV), as well as the coefficient determination,  $r^2$ , based on the following  
 234 calculations:

235 
$$\text{Bias} = 10^{\wedge} \left[ \frac{\sum_{i=1}^n \log_{10}(M_i) - \log_{10}(O_i)}{n} \right] \quad (3)$$

236 
$$\text{MAE} = 10^{\wedge} \left[ \frac{\sum_{i=0}^n | \log_{10}(M_i) - \log_{10}(O_i) |}{n} \right] \quad (4)$$

237 
$$\text{RMSE} = \sqrt{\frac{\sum_{i=1}^n (\log_{10}(M_i) - \log_{10}(O_i))^2}{n}} \quad (5)$$

238 
$$\text{CV} = \frac{\sigma}{\mu} \times 100\% \quad (6)$$

239 where M represents the modelled TSS values, n is the number of samples, and O represents the  
 240 observed TSS measurements, while  $\sigma$  refers to standard deviation and  $\mu$  represents the mean value.

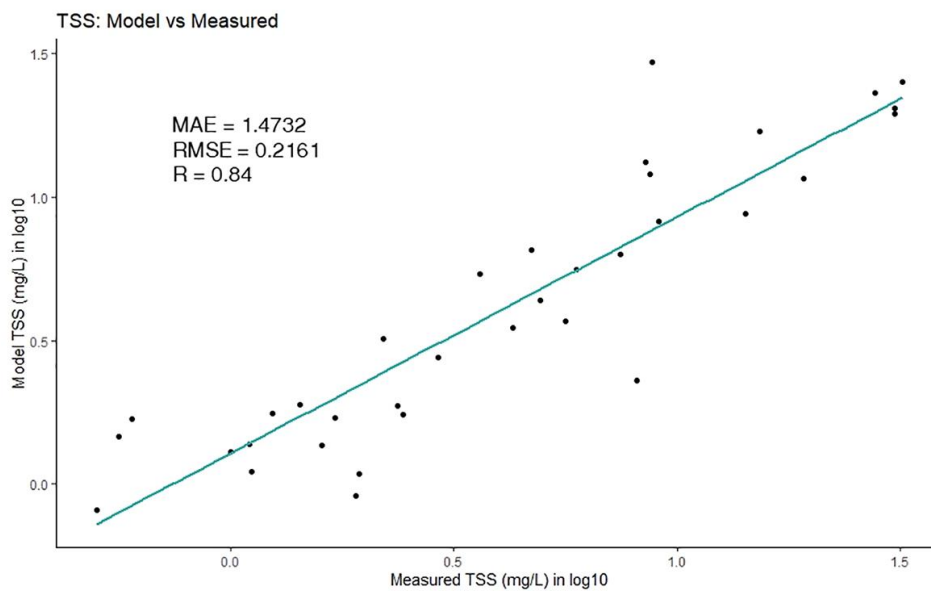
241 Equations (3), (4) and (5) use a log<sub>10</sub>-transform of the data as the range of TSS values can span several  
 242 orders of magnitude. As such, an application of the log-transform prior to error metric calculation

243 allows us to account for uncertainties that are proportional to the concentration values  
 244 (Balasubramanian et al., 2020; Seegers et al., 2018).

245 Table 2: Calibration and accuracy assessment of the newly derived MODIS-Aqua models in this study for TSS estimations  
 246 tested using various model functions. Calculation for bias, MAE and RMSE use a log-transform of the data prior to calculation  
 247 of error metric measurements, as adapted from Seegers et al. (2018) and Balasubramanian et al. (2020). Band ratio  
 248 Rrs(530)/Rrs(666) is established as function x. Power function model is selected based on low performance metric values.

Model	Function	Bias	MAE	RMSE	CV (%)	R
Power	$TSS = 21.238x^{-1.272}$	0.9999	1.4732	0.2161	4.74	0.84
Linear	$TSS = -1.8193x + 16.928$	1.4463	1.8549	6.7174	20.699	0.6854
Exponential	$TSS = 17.784e^{-0.296x}$	1.0791	1.4906	6.3088	3.8920	0.8154
Logarithmic	$-8.872\ln(x)+19.383$	1.1336	1.6177	5.3735	-17.056	0.8128

249  
 250  
 251



252  
 253 Fig. 4: Scatterplot of modelled TSS values derived from the proposed model and measured TSS values (mg/L).

254 An evaluation of the model was performed using a k-fold cross validation technique (Refaeilzadeh et  
 255 al., 2020) given the small size of the TSS dataset used in this study (Table 2). A selection of k = 7 was  
 256 assigned to split the datasets into k groups with an equal number of data points.

257 Table 2: Assessment of fitting error for the proposed TSS model, using k-fold cross validation.

Parameter	k-fold (n)	R2	RMSE	MAE
TSS	7	0.85	0.2159	0.1747

258  
 259 While these results point to low error levels achieved by the proposed regional TSS retrieval model  
 260 (Table 3, Fig. 4), caution should be used when applying it to various water types. Water type  
 261 classification has been thoroughly described by Balasubramanian et al. (2020) where waters are  
 262 classed into Type I (Blue-Green waters), Type II (Green waters), and Type III (Brown waters).  
 263 Essentially, the Green-to-Red band ratio is optimised with these datasets corresponding to sediment-  
 264 dominated and yellow-substance loaded water conditions. As highlighted by Morel & Belanger (2006),  
 265 waters of this type do not have the same spectral characteristics as phytoplankton-rich waters (also  
 266 known as Case 1 waters). In addition to the impact on water clarity, sediment particles (often red-  
 267 brownish coloured) also tend to enhance the backscattering and absorption properties, especially at  
 268 shorter wavelengths (Babin et al., 2003), while the additional presence of coloured dissolved matter  
 269 (yellow substance) leads to strong absorption properties at short wavelengths. As the TSS retrieval  
 270 model was developed from samples taken in waters that are bio-optically rich in suspended solids and  
 271 dissolved organic matter, an application of this TSS model needs to be done cautiously when applying  
 272 to other water types, particularly those with large concentration of phytoplankton.

#### 273 2.4 Application of TSS retrieval model

274 Daily MODIS-Aqua satellite data from year 2003 to 2019 (total of 6192 individual time slices) were  
 275 studied with a 2°x 2° spatial resolution (longitude: 109.38, 112.0; latitude: 1.22, 3.35) which covers  
 276 the southwestern coastal region of Sarawak and southern part of the South China Sea.  
 277 Atmospherically corrected MODIS-Aqua level 2 reflectance products (Bailey et al., 2010; NASA Official,

278 n.d.) were retrieved for the application of the TSS model proposed in this study. Negative remote  
279 sensing reflectance values, possibly due to failure of atmospheric correction, were filtered out before  
280 applying the retrieval model, as expressed in Eq. (2), to map the spatial and temporal distribution of  
281 TSS estimates. In addition, averaging of spatial and temporal TSS variation maps in this study was  
282 carried out by filtering TSS values with fewer than 10 valid data points over the whole time series,  
283 along with application of sigma clipping operation (refer to:  
284 [https://docs.astropy.org/en/stable/api/astropy.stats.sigma\\_clip.html](https://docs.astropy.org/en/stable/api/astropy.stats.sigma_clip.html)).

#### 285 2.4.1 Open Data Cube

286 In this study, the analysis of remote sensing data over large spatial extents and at high temporal  
287 resolution was carried out using robust Python libraries and packages run on an Open Data Cube (ODC)  
288 platform. Open Data Cube is an open-source advancement in computing technologies and data  
289 architectures which addresses the growing volume of freely available Earth Observation (EO) satellite  
290 products (Giuliani et al., 2020; Killough, 2019). ODC provides a collection of software which index,  
291 manage, and process large EO datasets such as satellite products from the MODIS, Landsat and  
292 Sentinel missions (Gomes et al., 2021). These satellite datasets are structured in a multi-dimensional  
293 array format, and provide layers of information across latitude and longitude (Open Data Cube, 2021).  
294 Leveraging the growing availability of Analysis Ready Data (ARD), and with support from the  
295 Committee of Earth Observation Satellites (CEOS) (Killough, 2019), the ODC concept has been  
296 deployed in many countries across the world. These existing deployments include Digital Earth Africa  
297 (<https://www.digitalearthafrika.org/>), Digital Earth Australia (DEA) (<https://www.dea.ga.gov.au/>),  
298 Vietnam Open Data Cube (<http://datacube.vn/>), and Brazil Data Cube ([https://github.com/brazil-data-](https://github.com/brazil-data-cube)  
299 [cube](https://github.com/brazil-data-cube)), which provide various time-series datasets of the changing landscape and water content in  
300 these specific regions (Giuliani et al., 2020; Gomes et al., 2021; Killough, 2019; Lewis et al., 2017). The  
301 ecosystem and architecture of ODC is well explained at [opendatacube.org](https://opendatacube.org). The codes and tools used

302 in this application drew upon the information provided in various DEA notebooks (Krause et al. (2021),  
303 which can be found at <https://github.com/GeoscienceAustralia/dea-notebooks/>.

## 304 2.5 Precipitation data and computation of river discharge

305 Monthly precipitation values (mm) over the Lupar and Rajang basins were extracted from the Global  
306 Precipitation Measurement (GPM) Level 3 IMERG satellite datasets  
307 (<https://gpm.nasa.gov/data/imerg>) in order to assess the influence of precipitation in each river basin  
308 in relation to TSS concentration at the corresponding river mouth (Supplementary Materails, Fig. S4 –  
309 7).

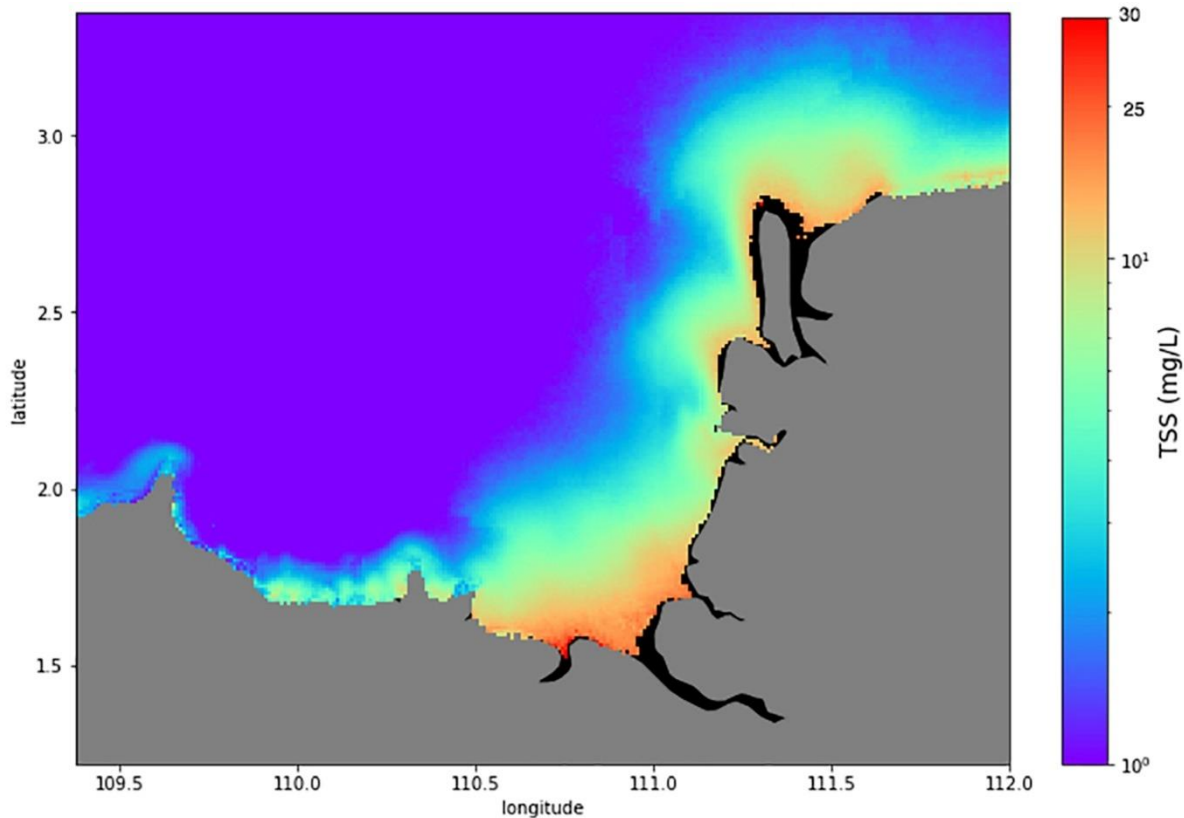
310 Derivation of river discharge ( $\text{m}^3/\text{s}$ ) was computed using total precipitation estimates (mm) over each  
311 river basin, and multiplied by a surface discharge runoff factor for the studied region (Sim et al., 2020).  
312 The surface runoff was estimated to be 60 % of total precipitation (Staub et al., 2000; Whitmore,  
313 1984). In this study, the Rajang river basin, as well as the combined basins of the Lupar, Sadong, and  
314 Saribas rivers (hereafter referred to as the Lupar basin), were studied for their river discharge rates in  
315 relation to TSS release.

## 316 3.0 Results & Discussion

### 317 3.1 Spatial variation of TSS distribution

318 Changes in TSS distribution occur across space and time. The regional TSS remote sensing model  
319 calibrated in this study was applied to the time series of MODIS-Aqua data to study the variability of  
320 spatial TSS distribution and identify potential hotspot areas susceptible to TSS water quality  
321 degradation. The map of average TSS for the Sarawak region was generated (Eq. 2) by averaging all  
322 the daily MODIS-Aqua TSS images (2003 to 2019) and is presented in Fig. 5. The results show that the  
323 waters in the northeast region of the study area, stretching from the Sadong river to the Rajang/Igan  
324 river have seen sustained levels of TSS over the 17 years considered in this study.

325 The temporally averaged spatial distribution map (Fig. 5) shows TSS concentrations in the range of 15  
326 – 20 mg/L near the river mouth areas, with widespread TSS plumes extending into the South China  
327 Sea (Fig. 5). Based on the Malaysia Marine Water Quality Criteria and Standard (Supplementary  
328 Materials, Table S2) (Department of Environment, 2019), these coastal waters fall under Class 1 in  
329 relation to their TSS (mg/L) status. This classification indicates that these coastal waters support and  
330 preserve marine life in this local region. Yet, several studies have expressed concerns regarding high  
331 TSS loadings in riverine waters owing to the impacts of various land use and land cover changes (LULC)  
332 (Ling et al., 2016; Tawan et al., 2020). Among these, the Rajang river has been highlighted to be heavily  
333 impacted by various LULC activities such as large-scale deforestation and construction of hydropower  
334 dams (Alamgir et al., 2020). In situ water quality studies by Ling et al. (2016) reported on high TSS  
335 estimates at one of the upstream tributaries of Rajang river, the Baleh river, with TSS readings up to  
336 approximately 100 mg/L. Another study by Tawan et al. (2020) reported a significant TSS release  
337 reaching to 940,000 mg per day during wet seasons, with maximum TSS concentrations of 1700 mg/L  
338 in the upstream tributaries of the Rajang river, particularly at the Baleh and Pelagus rivers. The  
339 majority of the upstream tributary rivers were categorised as Class II (during dry season) and Class III  
340 (during wet season) waters according to the National Water Quality Index (Supplementary Materials,  
341 Table S3) (Department of Environment, 2014), due to increased soil erosion from surrounding LULC  
342 activities (Tawan et al., 2020). These local in situ findings provide valuable insights on point source TSS  
343 estimates in these LULC change regions. Coupled with our spatial map of average TSS captured by  
344 remote sensing technologies, our findings seem to suggest that a large portion of TSS loadings from  
345 inland and upstream rivers would have settled and deposited in these river channels and were not  
346 completely discharged outwards into the coastal areas, which would have caused major water quality  
347 degradation in the corresponding coastal systems.



348

349 Fig. 5: Temporally averaged 2°x 2° map of TSS distribution (on a log scale) across the time dimension for each pixel.

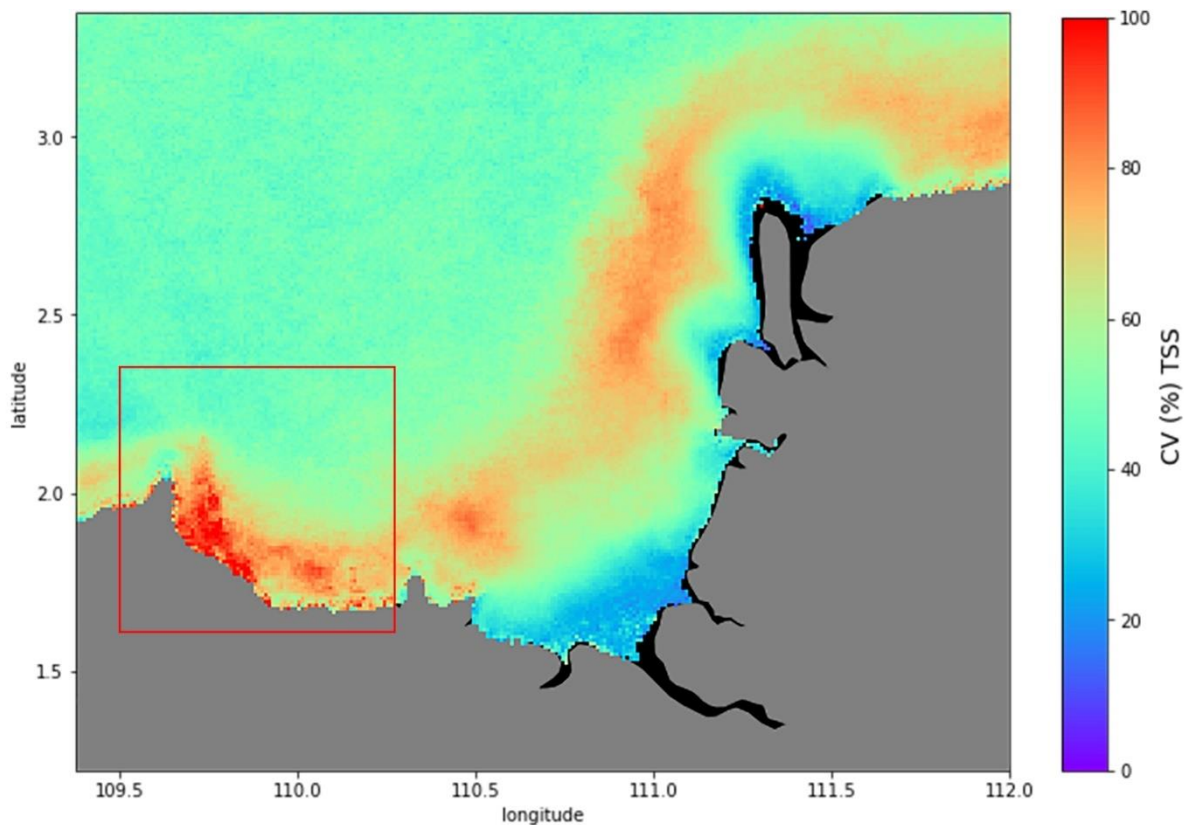
350 Historical patterns of TSS concentration were assessed by comparing annual maps of average TSS  
 351 distribution (Supplementary Materials; Fig. S1), as well as time series of TSS estimates at the Lupar  
 352 and Rajang river mouths (Supplementary materials; Fig. S2). From our findings, the annual TSS maps  
 353 further support the observation where TSS release was evident at Lupar and Rajang/Igan river mouths  
 354 from 2003 till 2019, which points to Class I of local water quality standards in relation to TSS (mg/L)  
 355 status. This was found to consistently occur every year. Furthermore, the TSS trend study showed that  
 356 both the Lupar and Rajang river mouth points have a gradual increase of TSS concentration over the  
 357 17 years (Supplementary materials; Fig. S2). This increasing trend was, however, not statistically  
 358 significant ( $p = 0.43$  for Lupar, and  $p = 0.15$  for Rajang).

359 Moreover, a map of the TSS coefficient of variation (CV) was computed to identify areas with a high  
 360 degree of relative TSS variation over time (Fig. 6). Here again, the map of CV (%) was produced by  
 361 aggregation of the daily MODIS-Aqua images (6192 time steps) from 2003 until 2019. Figure 6 shows



362 that the Samunsam-Sematan coastal region (as highlighted by the red box) exhibits an increased level  
363 of TSS distribution variability, with a recorded CV of more than 90 %.

364 The Samunsam-Sematan coastal region contains near-pristine mangrove forests which are sheltered  
365 from major LULC activities, as compared to other studied sites. Samunsam-Sematan is also well-known  
366 locally as a recreational hotspot with coral reefs and various national parks (Sarawak Tourism Board,  
367 2021). Data from the Centre for International Forestry Research (CIFOR) Forrest Carbon database  
368 (CIFOR, n.d.) revealed that there was more than double the amount of total forest loss (approx. 5,000  
369 Ha) recorded in Lundu, a nearby township in the Sematan area in 2011 as compared to the previous  
370 years. Deforestation activities, regardless of their scale, can inevitably promote sediment loss and soil  
371 leaching into the nearby river systems (Yang et al., 2002). Important information regarding the  
372 variability in water quality (as shown in Fig. 6) can provide support to local authorities and relevant  
373 agencies in order to identify vulnerable areas that need to be monitored closely, such as the Lundu-  
374 Sematan region in this case. The CV map thus offers interesting insights into how TSS distribution can  
375 vary across large spatial areas, which can ultimately impact local socio-economic activities in this  
376 region (Lee et al., 2020b).



377

378 Fig. 6: Map of CV (%) calculated from the daily time series of MODIS-Aqua satellite images from 2003 until 2019.

379 3.2 Temporal variation of TSS distribution

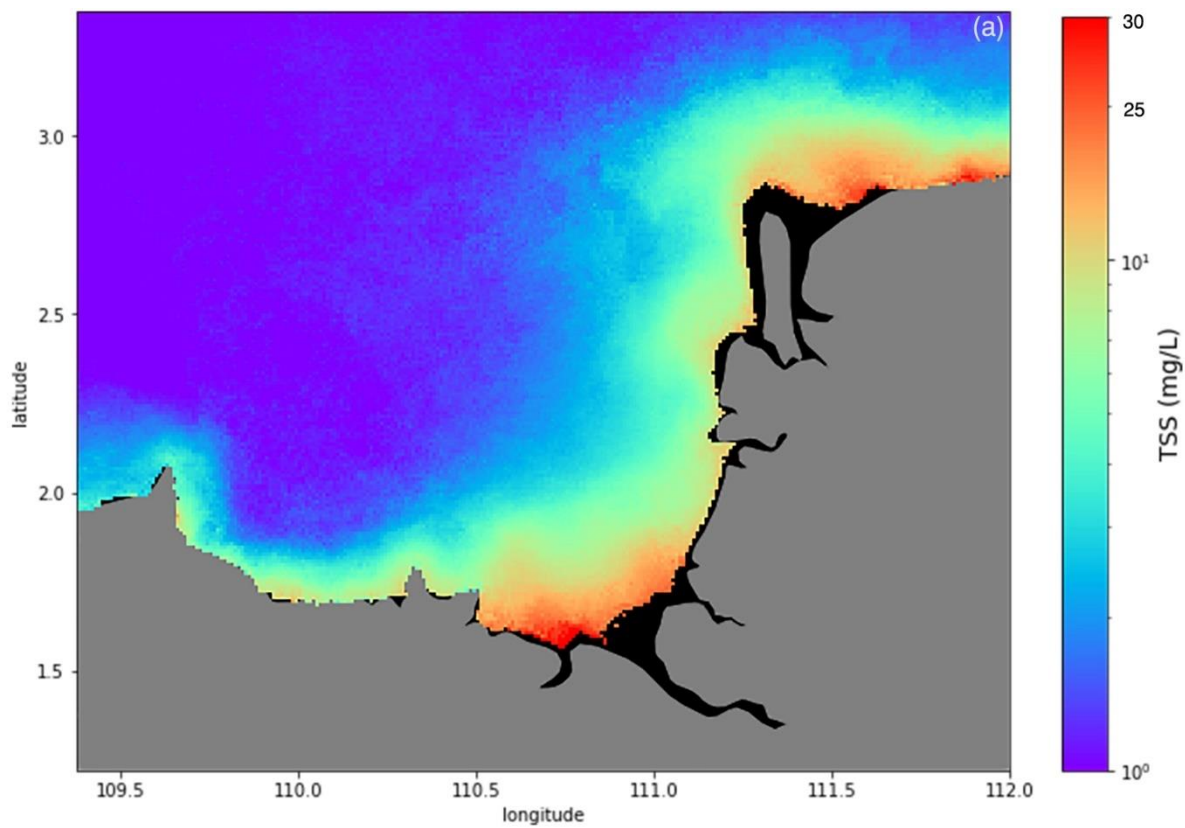
380 On a temporal scale, the northeast (NE) monsoon period shows a distinct difference in the widespread  
 381 intensity of TSS distribution as compared to the southwest (SW) monsoon period, along the Sarawak  
 382 coastline over the 17 years of the considered time series (Fig. 7). Mapping of temporal variations  
 383 between monsoons using time-series MODIS-Aqua datasets can provide an improved understanding  
 384 on the intensity of monsoonal patterns in driving the TSS distribution in this region. As shown in Figure  
 385 7, TSS release can be seen to extend further into the open ocean South China Sea region during the  
 386 NE monsoon periods (Fig. 7a) in comparison to the SW monsoon periods (Fig. 7b).

387 In addition, the differences in TSS release between the NE and SW monsoons  $((NE-SW)/NE \times 100)$  were  
 388 mapped as shown in Fig. 7c. Widespread TSS plumes are detected at Lundu/Sematan region ( $>80\%$   
 389 relative difference in TSS concentration) on the southwest side of the study area, while substantial  
 390 TSS plumes are observed in front of the Igan river channel, with more than  $50\%$  relative difference in

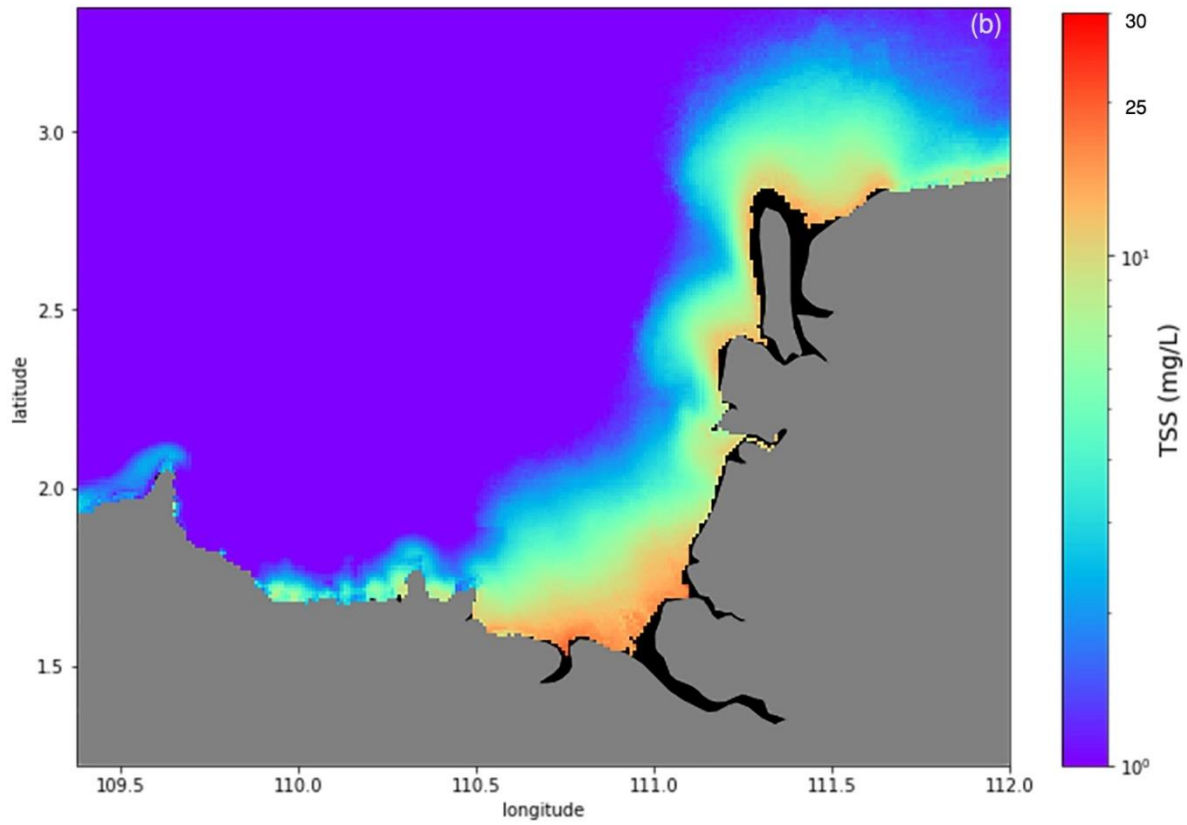
391 TSS concentration in comparison to SW monsoon periods. Sadong coastal area is observed to receive  
392 considerable TSS loadings (> 30%) during NE monsoon periods.

393 These coastal areas would thus be more likely to be impacted by the TSS release during the NE  
394 monsoon periods. These findings further strengthen the evidence that tropical rivers are majorly  
395 impacted by climatic variability such as monsoonal patterns, as highlighted in a study at Baleh river in  
396 Sarawak (Chong et al., 2021). This suggests that monsoon rains, which typically last for several months,  
397 play an integral role in driving the discharge of TSS in tropical rivers.

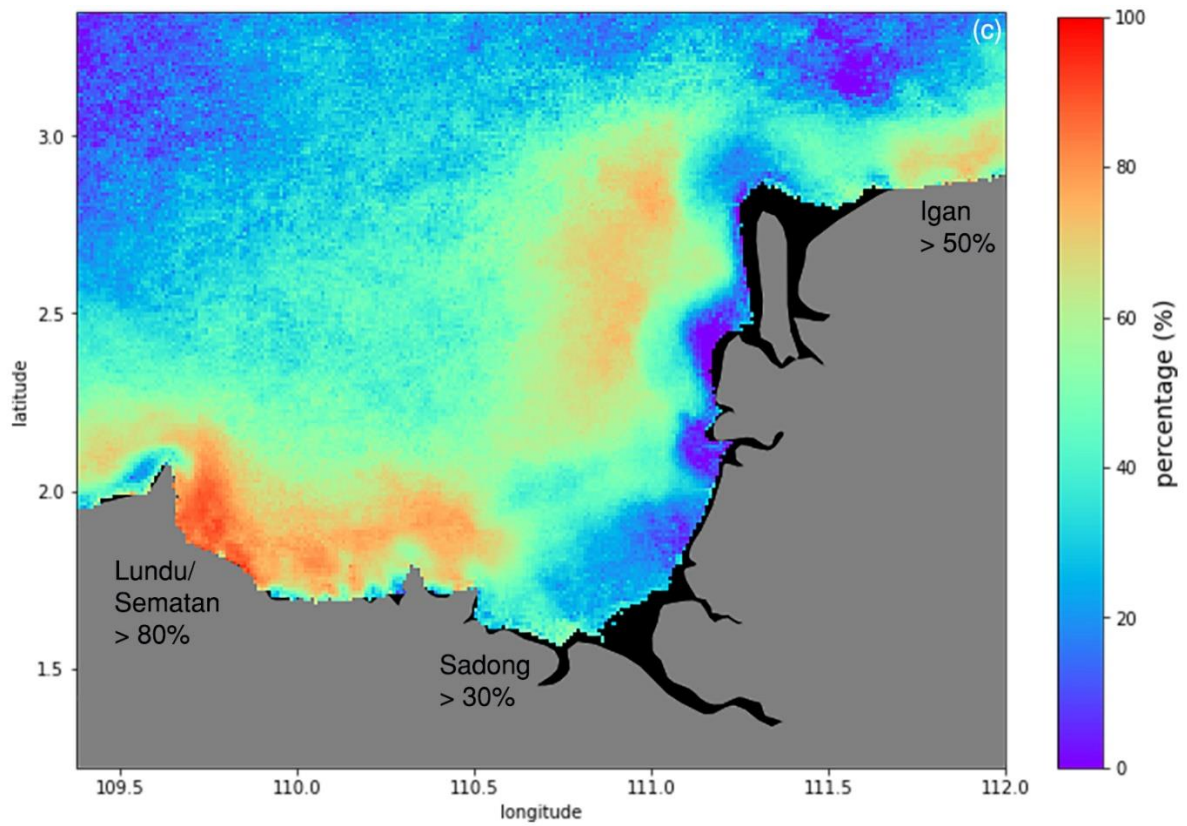
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Fig. 7: Long-term average of TSS estimates (mg/L) during the Northeast monsoon (a), and the Southwest monsoon (b). The

403 map of TSS difference between the Northeast and Southwest monsoon periods, computed in relative percentage (%), is  
404 shown in (c).

405 Several climatic studies in the Borneo region highlighted 2009 as a year with extreme rainfall events  
406 which caused major floods in Sarawak (Dindang et al., 2011; Sa'adi et al., 2017), while drought events  
407 were reported in 2014 (Bong and Richard, 2020). Hence, in this study, TSS dynamics for the Lupar and  
408 Rajang rivers were studied by assessing the variation of TSS values at selected pixels in relation to  
409 monsoonal rainfall patterns in 2009 and 2014 (Supplementary Materials; Fig. S3).

410 Generally, the results show fluctuations of TSS concentrations across the NE and SW monsoon periods  
411 in relation to precipitation values (Fig. 8a – d). Based on Fig. 8a, monthly precipitation values recorded  
412 for the Lupar river basin in 2009 showed a clear decreasing trend from the NE monsoon period (wet  
413 season) to the SW monsoon period (dry season), while gradually increasing approaching the year end's  
414 NE monsoonal period. A similar precipitation pattern was observed for the Rajang river basin during  
415 the same year (Fig. 8c).

416 However, these results also show that the TSS distribution (mg/L) at the Lupar river mouth seems to  
417 show no distinct trend of decreasing TSS concentration estimates during the SW monsoon period in  
418 year 2009 (Fig. 8a) in relation to its precipitation values. Additionally, a sharp rise of TSS release can  
419 be seen in the month of May (beginning of SW monsoon period), with a near equivalent intensity of  
420 TSS release during the NE monsoon period. This observation may potentially be caused by the lag  
421 between the time of rainfall events occurring during NE monsoon periods and TSS release entering  
422 the coastal river regions. A similar observation was described by Sun et al. (2017a) suggesting that  
423 riverine outputs could take several days, and even up to one month to reach the coastal river points.  
424 Considering the occurrence of extreme rainfall events in 2009, our findings are in agreement with  
425 these processes as TSS concentrations generally exhibit a similar intensity throughout the NE and SW  
426 monsoonal periods for the Lupar river (Fig. 8a). This result could suggest that the occurrence of

427 extreme rainfall events, as reported for the year 2009, can exert a much larger impact on TSS  
428 transportation and release in monsoon-driven tropical rivers.

429 Drought events in 2014 can be seen to impact the precipitation values at both the Lupar (Fig. 8b) and  
430 Rajang river basins (Fig. 8d). There is no apparent patterns of decreasing precipitation values during  
431 the shift of NE to SW monsoonal periods as compared to the year 2009, for either river basin. However,  
432 precipitation values were found to increase sharply during the year end NE monsoon period for both  
433 river basins. The TSS concentrations at the Lupar coastal river points were found to be the highest  
434 during the NE monsoon period earlier in January and February of 2014 (Fig. 8b). This may be due to  
435 the temporal lag in the transition of TSS discharge into the coastal systems arising from the prior  
436 months (November and December) in the previous year, when higher rainfall events were typically  
437 observed in this region (Gomyo and Koichiro, 2009; Tangang et al., 2012). The TSS distribution at both  
438 Lupar and Rajang coastal river points showed no distinct trend in relation to the precipitation values  
439 throughout a period of ten months until November 2014. These findings suggest that coastal areas in  
440 the Borneo region may not be experiencing critical water quality degradation during dry seasons.

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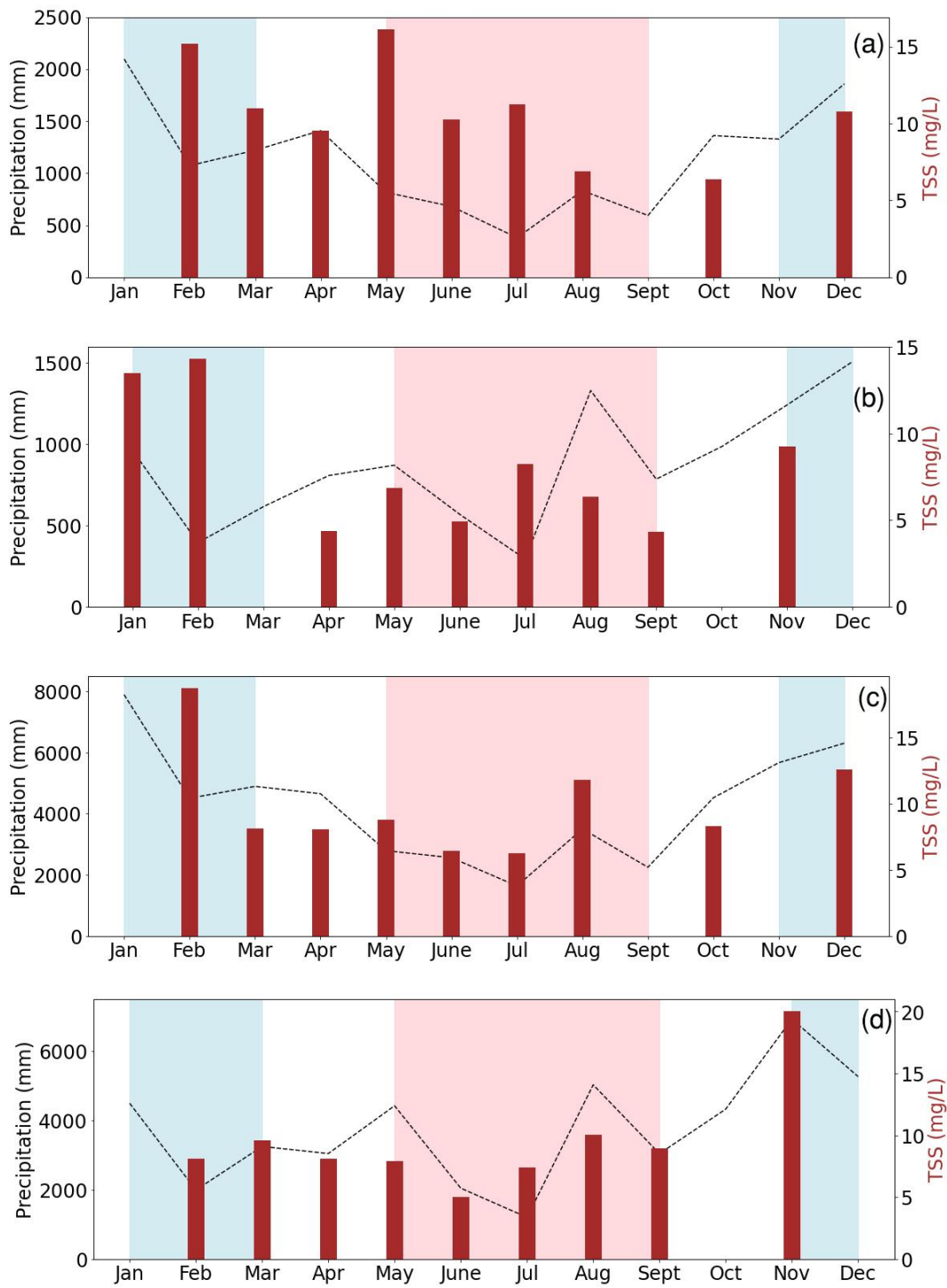
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450 Fig. 8: Temporal analysis of precipitation (mm) from the Lupar and Rajang river basins in relation to TSS concentrations (mg/L)

451 during the NE and SW monsoon periods at the Lupar ((a): 2009; (c): 2014) and Rajang ((b): 2009; (d): 2014) coastal river point.

452 The NE monsoon months are highlighted with a blue background; those of the SW monsoon with a pink background, and

453 intermonsoon periods with a white background.

### 454 3.2.1 Temporal TSS anomalies

455 Considering the temporal variation recorded across monsoons, maps of relative TSS anomalies were  
456 calculated for each year as the difference with respect to the long-term TSS mean (Fig. 5), in order to  
457 detect changes of TSS distribution occurring annually (Fig. 9). As shown in Figure 9, year 2010  
458 experienced a distinct increase of TSS distribution (approximately 100 %), with widespread pattern  
459 extending into open ocean waters, in comparison to the long-term TSS mean. This finding provides an  
460 interesting insight into the effects of extreme rainfall events as recorded in year 2009, which could  
461 potentially intensify TSS release into the coastal and open ocean waters. The effects of TSS release  
462 can still be seen a year after the extreme rainfall events in this region. This observation could provide  
463 further evidence that the impacts of the TSS release from the land into rivers and coastal systems may  
464 only take effect after a substantial period, as previously observed by Sun et al. (2017a).

465 Figure 9 further reveals an interesting pattern of TSS increase in the Samunsam-Sematan region from  
466 year 2004 until 2019, with exceptions during the years 2007 and 2008. As previously highlighted in  
467 Section 3.1, the Samunsam-Sematan region has been observed to be a vulnerable coastal area with  
468 respect to TSS water quality degradation. From the annual map of TSS anomalies (Fig. 9), we can see  
469 that the TSS distribution has the tendency to accumulate in the Samunsam-Sematan region, as  
470 opposed to being distributed into the open ocean waters. This may be due to the geographical and  
471 hydrological characteristics of these coastal regions (Martin et al., 2018), as the TSS release may be  
472 sheltered from open ocean waters, and hence induce a higher TSS accumulation in these coastal  
473 regions.

474

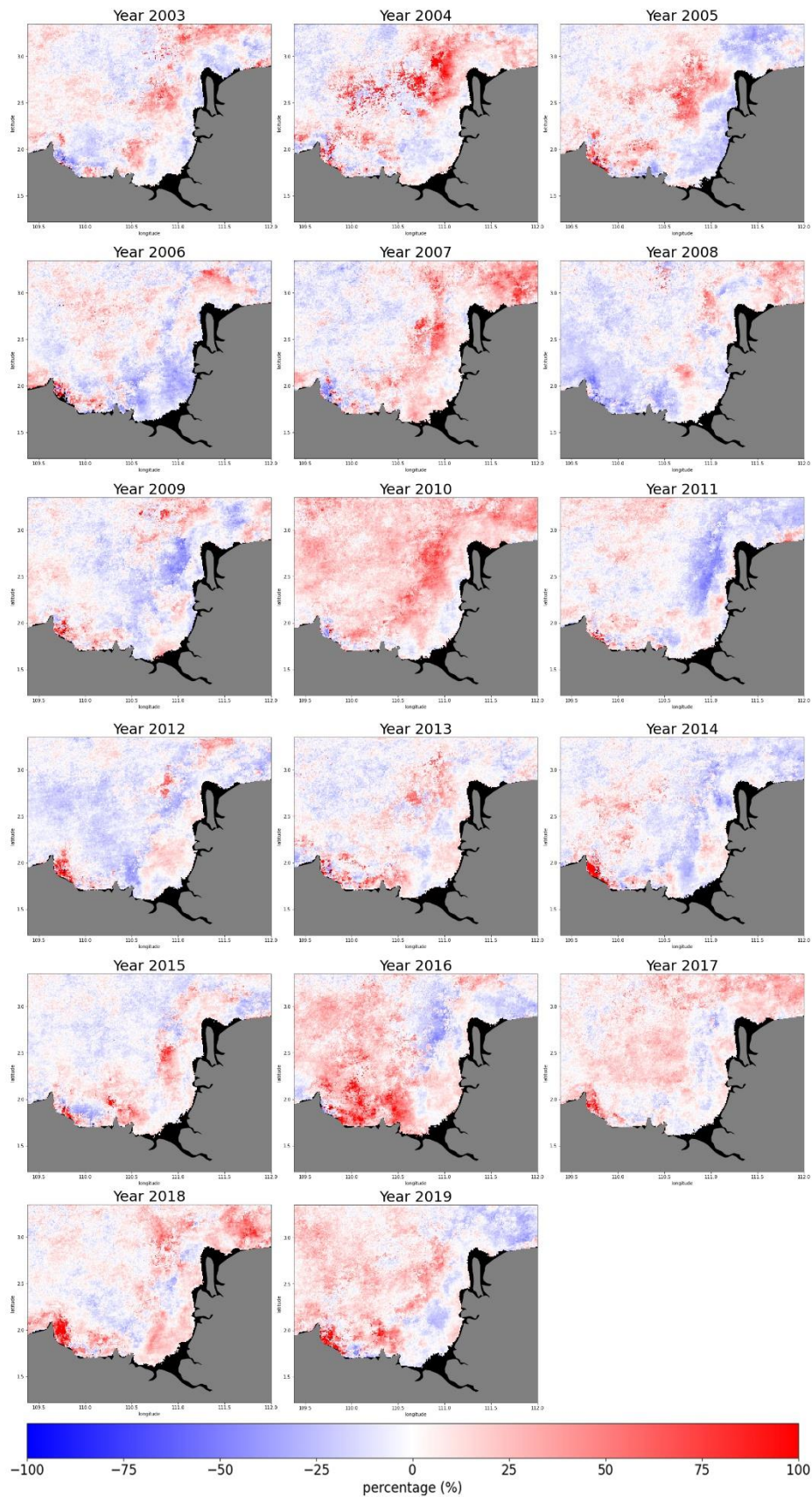
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480 Fig. 9: Map of relative TSS distribution anomalies with respect to the long-term mean, represented as percentage (%), from  
 481 year 2003 until 2019.

482 3.3 Hydrological factors driving TSS discharge

483 Apart from the influence of monsoonal patterns, hydrological factors such as the river discharge are  
484 among the dominant drivers in transporting various water constituents in riverine and coastal systems  
485 (Loisel et al., 2014; Petus et al., 2014; Sun, 2017b; Verschelling et al., 2017). In this study, river  
486 discharge from the Lupar and Rajang basins was estimated and investigated.

487 Yearly river discharge estimates from 2003 until 2019 were investigated to assess its effect on the TSS  
488 distribution (Fig. 10) represented by changes in TSS values for pixels located at each Lupar and Rajang  
489 coastal river points (Supplementary Materials, Fig. S3). Figure 10a shows that river discharge values in  
490 the Lupar basin (750 to 1050 m<sup>3</sup>/s) are approximately twice lower than the Rajang river discharge (Fig.  
491 10b), which recorded a range of 3,200 to 4,000 m<sup>3</sup>/s.



492

493

494 Fig. 10: Time-series analysis of river discharge (m<sup>3</sup>/s) in relation to TSS concentrations (mg/L) for the Lupar (a) and Rajang (b)

495 basins from year 2003 to 2019. Note the differing scaling on the ordinate axes in each plot.

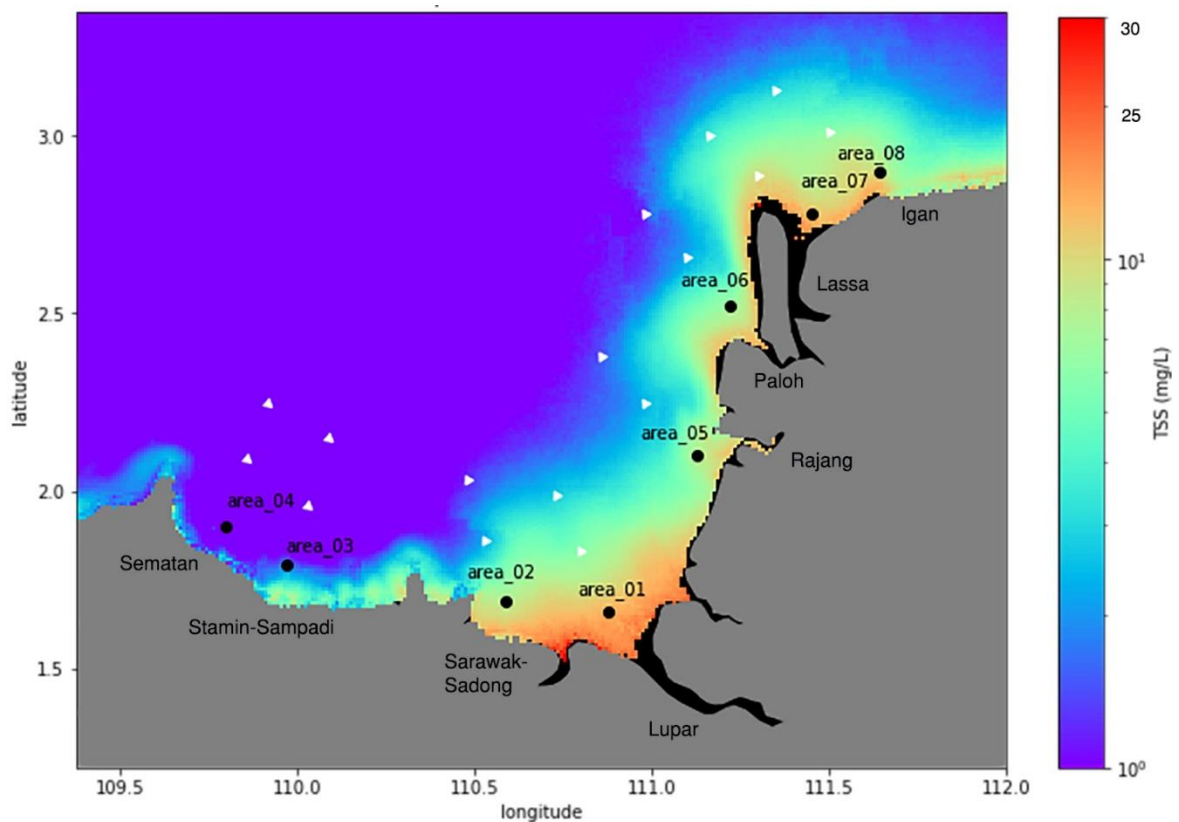
496 Discrepancies between TSS estimates and river discharge were identified in both the Lupar and Rajang  
497 coastal regions in these annual time-series, where river discharge was inversely correlated with TSS  
498 estimates. These discrepancies are not uncommon, as previously highlighted in a study by Zhan et al.  
499 (2019). Especially in 2010 for the Lupar river, Fig. 10a shows a drop in TSS release in relation to the  
500 steady increase of river discharge from the river basin. In 2011 and 2012, a negative correlation can  
501 be seen between river discharge and TSS estimates, while in subsequent years from 2013 until 2015,  
502 there is a clear positive correlation. Although there is no obvious environmental factor that would  
503 explain these discrepancies, these findings may imply a complex interaction between human  
504 interventions, such as damming and deforestation activities, as well as varying hydrological and  
505 atmospheric conditions (wind and tidal mixing) in regulating TSS dynamics (Wu et al., 2012; Zhan et  
506 al., 2019).

507 The correlation between TSS release and river discharge at both the Lupar and Rajang coastal areas  
508 was further evaluated in this study. Even though river discharge has been shown (in other global  
509 studies) to be one of the dominant factors in moderating TSS release (Fabricius et al., 2016; Tilburg et  
510 al., 2015; Verschelling et al., 2017; Wu et al., 2012), the TSS distribution at both the Lupar and Rajang  
511 river mouths in this study can be seen to be only poorly correlated with river discharge from each river  
512 basin (Supplementary Materials, Fig. S8a and b). The TSS output from the Lupar basin recorded a  
513 correlation coefficient of  $r = 0.15$ , while river discharge from the Rajang basin did not substantially  
514 influence the TSS release either, with  $r = 0.27$  throughout the seasons. Coupled with tidal mixing  
515 processes (Ramaswamy et al., 2004; Zhou et al., 2020) , it is possible that human activities such as  
516 deforestation, logging, and construction of dams, which are largely occurring within the Rajang basin  
517 (Alamgir et al., 2020), are mainly driving TSS release and resuspension in this area. This indicates that  
518 although TSS release is regarded to be highly dependent on, and controlled by river discharge patterns,  
519 this interaction often represents an intricate process linked to local hydrodynamics process and socio-  
520 economic conditions (Espinoza Villar et al., 2013; Fabricius et al., 2016; Valerio et al., 2018; Zhan et al.,  
521 2019).

522 3.4 Variability of TSS across coastal waters

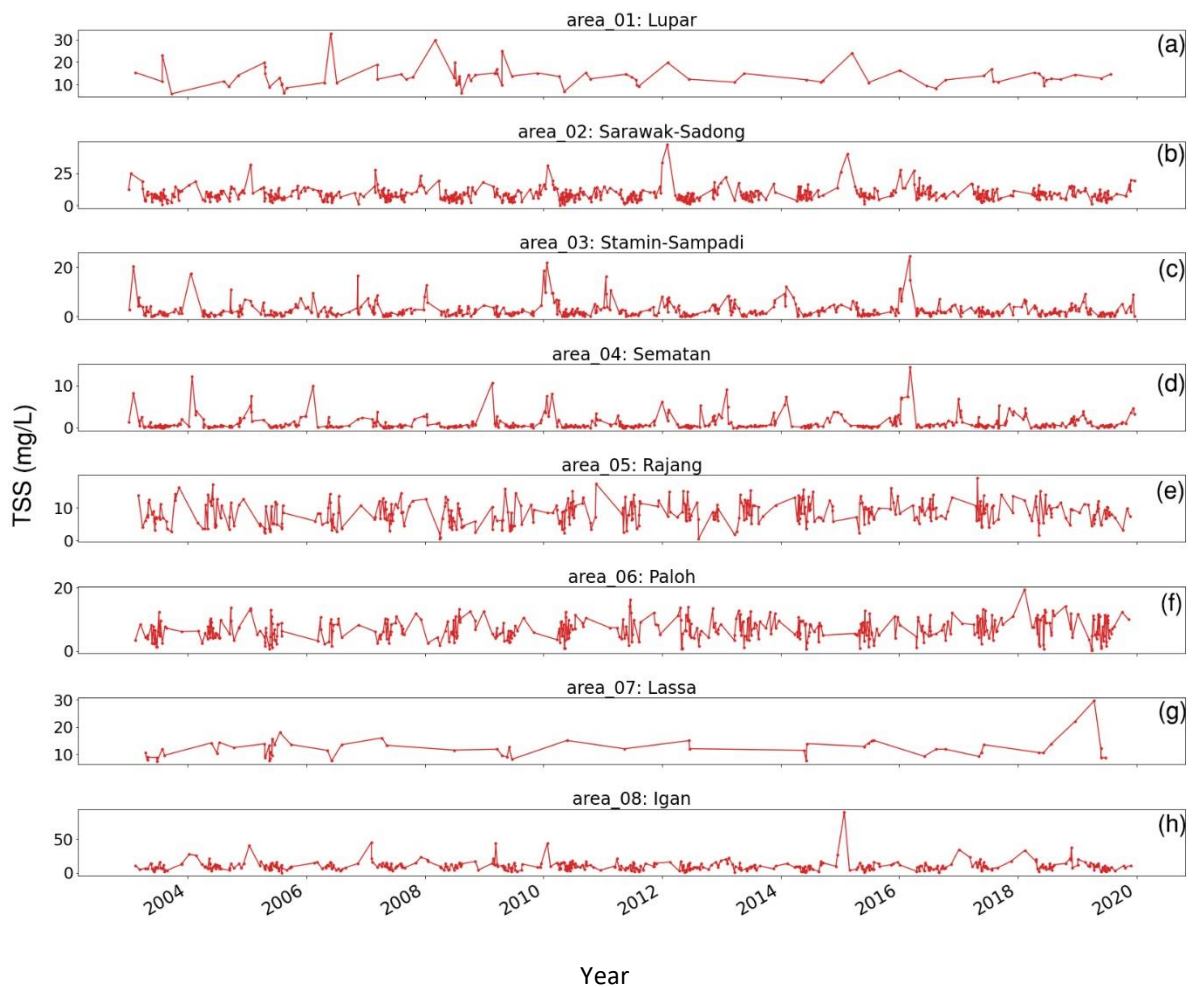
523 As previously observed in Fig. 5, varying river plumes of TSS were evidently detected within the coastal  
524 regions of the study area. Notably, coastal river plumes represent important factors driving the  
525 transport of water constituents and nutrients from coastlines to the open oceanic systems (Petus et  
526 al., 2014). To assess this and evaluate the water quality status in coastal zones, the spatial extent of  
527 TSS release was investigated along transects covering the territorial (12 nautical miles) and open water  
528 areas (24 nautical miles) of the Sarawak region (Fig. 11).

529 A total of eight coastal points were selected based on the main river mouths located in the southwest  
530 region of Sarawak. Transect points are positioned in a line starting at the coastal river points to  
531 examine the variations of TSS distribution across different water zones. Daily changes in TSS  
532 concentration for each pixel located in front of the river mouths were plotted from 2003 until 2019  
533 (Fig. 12).



534

535 Fig. 11: Map of average TSS estimates (mg/L) with indicators at eight main river mouths and their transect, extending from  
 536 coastal waters into territorial and open ocean systems. Indicators of each river mouths are as follows: area\_01 – Lupar river;  
 537 area\_02 – Sarawak-Sadong river; area\_03 - Stamin-Sampadi river i; area\_04 – Sematan river; area\_05: Rajang river; area\_06:  
 538 Paloh river; area\_07: Lassa river; and area\_08: Igan river).



539  
 540 Fig. 12: Graphs of daily TSS estimates (mg/L) recorded at eight river mouth points from 2003 to 2019. Presentation of each  
 541 river mouths is as follows: a) area\_01; b) area\_02; c) area\_03; d) area\_04; e) area\_05; f) area\_06; g) area\_07; h) area\_08.  
 542 Note the different TSS scales in each plot.

543 From the high temporal resolution graphs in Fig. 12, no general trend of TSS concentration can be  
 544 identified over the years at each coastal point. It is worth highlighting that the daily temporal  
 545 resolution was particularly affected at coastal points located in front of the Lupar (area\_01) and Lassa  
 546 (area\_07) river mouths due to various pixel data quality issues in these areas. Nonetheless, more than  
 547 80 satellite images with minimum cloud coverage at these two locations were processed, while the  
 548 remaining coastal points had a total of more than 400 satellite images to assess the temporal trend.

549 Despite the fact that no distinct upward or downward trend was observed, our findings indicate that  
550 several river mouths are actively discharging and accumulating substantial TSS amounts over the  
551 period of years, while resuspension of bottom sediments induced by wind and tidal cycle is another  
552 factor contributing to the variation of TSS values (Park, 2007; Song et al., 2020).

553 The coastal region of the Sarawak-Sadong river (area\_02) shows relatively high TSS distribution  
554 patterns with some periods recording an estimate of over 30 mg/L of TSS concentration. This is in  
555 agreement with the localised characteristics of the Sarawak river basin which essentially drains  
556 through the populated Kuching area with high industrial and development activities in the capital city  
557 of Sarawak (DID, 2021b). In comparison with other river mouth points, a steady TSS concentration  
558 below 20 mg/L was recorded across the Stamin-Sampadi (area\_03), Sematan (area\_04), Rajang  
559 (area\_05), and Paloh (area\_06) river mouths. Consistently high TSS values in the daily plots were  
560 recorded at the Lupar (area\_01) and Pulau Brait-Lassa (area\_07) river mouths, with estimates of up to  
561 30 mg/L on a near-daily basis. Similar high TSS amounts from the Igan (area\_08) river mouth, situated  
562 northeast side of the Pulau Brait-Lassa region, were observed in Cherukuru et al. (2021) and Staub et  
563 al. (2000).

564 Although the daily TSS estimates at each river point are in line with various reported studies (Chen et  
565 al., 2011, 2015b; Kim et al., 2017; Mungen et al., 2020; Zhang et al., 2010a), these estimates can be  
566 expected to be much higher for sampling points much closer to the river mouths. The selection of  
567 coastal river points in this study was made to minimize the gaps with respect to various pixel data  
568 quality issues in the MODIS-Aqua datasets, and hence, the use of coastal river points closer to shore  
569 would have been impractical.

570 These findings further suggest that higher TSS loadings within the coastal river areas would have been  
571 diluted or deposited while travelling to the open oceanic systems as they are weakly impacted by river  
572 discharge in relation to offshore distance (Espinoza Villar et al., 2013). This understanding can be  
573 observed in Fig. 13, which shows a progressively decreasing TSS estimates at each transect in relation

574 to the distance from the shore. Generally, TSS estimates in coastal zones (first transect point) show  
575 considerably higher TSS concentrations. When moving outwards to territorial waters (second transect  
576 point), TSS concentration estimates decrease by nearly 50 % before travelling to open ocean systems  
577 (third transect point), except for the northeast regions (area\_07 and area\_08) which seem to show  
578 large extension of TSS plumes to the open ocean waters, as also highlighted by Cherukuru et al. (2021).  
579 A reversed trend can be seen in the plot corresponding to the Sematan coastal river systems, although  
580 the absolute increase in TSS estimates across water zones (0.2 mg/L in total) here is only marginal (Fig.  
581 13d).

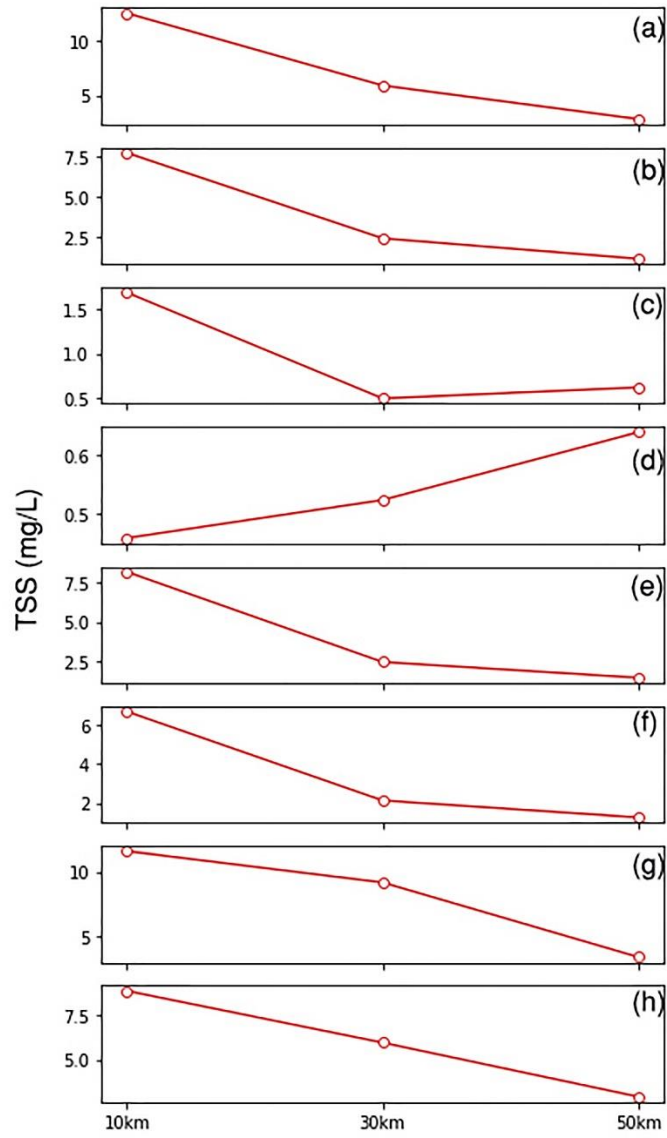
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588 Fig. 13: Average TSS estimates (mg/L) computed from 2003 to 2019 for each of the eight rivers (area\_01: (a), area\_02: (b);  
 589 area\_03: (c); area\_04: (d); area\_05: (e); area\_06: (f); area\_07: (g); and area\_08: (h)) and their relevant transect points with  
 590 distance of 10 km (coastal waters), 30 km (territorial waters) and 50 km (open ocean waters) from the shoreline. Note the  
 591 varying TSS scales on the ordinate axes in each plot.

592

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595



### 596 3.5 Discussion of TSS implications for coastal waters

597 High discharge of TSS into coastal environments can lead to adverse environmental and ecological  
598 implications. The presence of TSS affects water transparency and light availability within the surface  
599 waters (Dogliotti et al., 2015; Nazirova et al., 2021; Wang et al., 2021). Among others, TSS affects the  
600 photosynthesis activities of algae and macrophytes. TSS in water creates a reduction in light  
601 penetration, which impacts the primary production of aquatic organisms and hence the support  
602 system of marine life (Bilotta and Brazier, 2008; Loisel et al., 2014).

603 Additionally, TSS exerts an influence on zooplankton communities. Reduction in water clarity induces  
604 changes in the zooplankton's biomass volume and composition, while TSS may carry a level of toxicity  
605 which affects zooplankton through ingestion (Chapman et al., 2017; Donohue and Garcia Molinos,  
606 2009). Apart from that, accumulation and deposition of sediments decreases the level of dissolved  
607 oxygen (DO) at the bottom of the water column, and subsequently impacts the benthic invertebrate  
608 groups (Chapman et al., 2017). Moreover, substantial TSS deposition tends to cause harmful physical  
609 effects to these benthic groups, such as abrasion, and even clogging by sediment particles (Chapman  
610 et al., 2017; Langer, 1980).

611 As a result of these TSS effects on lower trophic levels, fish communities are critically impacted, with  
612 a reduction in diversity and abundance (Kemp et al., 2011). While fish communities learn to adapt to  
613 a range of TSS loads (Macklin et al., 2010), increase of TSS concentrations often depletes DO  
614 concentrations in the water system and causes stress towards these aquatic communities (Henley et  
615 al., 2000). Fish populations tend to decrease, as feeding and growth rates are negatively impacted  
616 (Shaw and Richardson, 2001; Sutherland and Meyer, 2007).

617 Threats to coral reefs have been linked to sediment-induced stress which often leads to a reduction  
618 in the coral's growth and metabolic rate, as well as impending mortality (Erftemeijer et al., 2012;  
619 Gilmour et al., 2006; Risk and Edinger, 2011). Factors of coral stress are driven by nutrient-rich  
620 sediments and microbes which are being carried by TSS, with impacts on the health of coral tissues

621 (Hodgson, 1990; Risk and Edinger, 2011; Weber et al., 2006). A reduction in light availability impedes  
622 the development of corals (Anthony and Hoegh-Guldberg, 2003; Rogers, 1979; Telesnicki and  
623 Goldberg, 1995). A combined increase in TSS and nutrient loadings contribute to the decrease of coral  
624 species diversity and composition (Fabricius, 2005).

625 Essentially, presence of TSS in water systems has impacts across various aquatic biota. With the severe  
626 implications of decreased fish population, this could lead to a disruption of fisheries activities by local  
627 communities, especially considering that more than 80 % of the Sarawak population is living in the  
628 coastal areas (DID, 2021a). Coral reefs are important coastal biodiversity assets to the Sarawak region,  
629 especially around the Talang-Talang and Satang islands on the southwest coast of Sarawak (Long,  
630 2014). With the use of remote sensing technologies in monitoring Sarawak coastal water quality, the  
631 approach presented in this paper provides digital-based solutions to assist relevant authorities and  
632 local agencies to better manage the Sarawak coastal waters and their resources.

#### 633 4.0 Conclusion

634 In this study, a regional empirical TSS retrieval model was developed to analyse TSS dynamics along  
635 the southwest coast of Sarawak. The empirical relationship between in situ reflectance values,  $R_{rs}(\lambda)$ ,  
636 and in situ TSS concentrations was established using a green-to-red band ratio using the MODIS-Aqua  
637  $R_{rs}(530)$  and  $R_{rs}(666)$  reflectance bands. An evaluation of the TSS retrieval model was carried out with  
638 error metric assessment, which yielded results of bias = 1.0, MAE = 1.47 and RMSE = 0.22 in mg/L  
639 computed in log<sub>10</sub>-transformed space prior to calculation. A statistical analysis using a k-fold cross  
640 validation technique (k = 7) reported low error metrics (RMSE = 0.2159, MAE = 0.1747).

641 The spatial TSS distribution map shows widespread TSS plumes detected particularly in the Lupar and  
642 Rajang coastal areas, with average TSS range of 15 – 20 mg/L estimated at these coastal areas. Based  
643 on the spatial map of the TSS coefficient of variation, large TSS variability was identified in the  
644 Samunsam-Sematan coastal areas (CV > 90 %). The map of temporal variation of TSS distribution  
645 points to a strong monsoonal influence in driving TSS release, with large differences identified

646 between the northeast and southwest monsoon periods in this region. From the annual TSS anomalies  
647 maps, the Samunsam-Sematan coastal areas demonstrated strong TSS variation spatially, while  
648 widespread TSS distribution with nearly 100 % of TSS increase in comparison to long term mean was  
649 observed in 2010. Furthermore, our study on river discharge in relation to TSS release demonstrated  
650 a weak relationship at both the Lupar and Rajang coastal river points. Study on the TSS variability  
651 across coastal river mouths implied that higher TSS loadings in the coastal areas are potentially being  
652 deposited or diluted in the process of being transported into the open ocean waters, with varying  
653 magnitude at several coastal river points.

654 Overall, these coastal areas of Sarawak are dominantly categorised as Class I quality, which remain  
655 within local quality standards to support various marine and socio-economic activities in this region.  
656 Our findings in the southwest coastal areas (Sematan and Stamin-Sampadi) showed that the coral  
657 reefs there can be well-maintained with negligible impacts from TSS loadings. However, it is important  
658 to highlight the various human activities that are widely ongoing in this region, which include  
659 deforestation and logging activities (Alamgir et al., 2020; Hon and Shibata, 2013; Vijith et al., 2018).  
660 Impacts from these activities in Sarawak can potentially aggravate current soil erosion issues, and  
661 ultimately induce more soil leaching and runoff from land to water systems, especially during heavy  
662 rainfall events (Ling et al., 2016; Vijith et al., 2018). As a result, human activities may have a greater  
663 influence on driving riverine sediments than climatological factors, as reported by Song et al. (2016).  
664 As such, this work presents the first observation of TSS distributions at large spatial and temporal  
665 scales in Sarawak's coastal systems, and of the potential associated impacts on the South China Sea.  
666 The findings derived from this work can be used to support local authorities in assessing TSS water  
667 quality status in the coastal areas of concern, and to enhance coastal management and conservation  
668 strategies. The application of remote sensing technologies is of great benefit in the development of  
669 sustainable sediment management in the Sarawak coastal region, as demonstrated in this study.

670 Data availability. The dataset related to this study is available as supplement to this paper.

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