

**An analysis of the variability of  $\delta^{13}\text{C}$  in macroalgae from the Gulf of California: indicative of carbon concentration mechanisms and isotope discrimination during carbon assimilation**

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## Abstract

The isotopic composition of carbon in macroalgae ( $\delta^{13}\text{C}$ ) is highly variable, and its prediction is very complex concerning terrestrial plants. ~~The~~ determinants of  $\delta^{13}\text{C}$ -macroalgal ~~variations, we were~~ analyzed ~~in~~ a large stock of specimens that vary in taxa and morphology, ~~collected in~~ shallow marine habitats in the Gulf of California (GC) ~~with~~ ~~featured by~~ distinctive environmental conditions. A large  $\delta^{13}\text{C}$  variability (-34.6‰ to -2.2‰) was observed. ~~Life forms~~ (taxonomy ~~57%~~, morphology, and structural organization ~~34%~~) ~~explains the variability related to the carbon use physiology. Environmental conditions influenced the  $\delta^{13}\text{C}$ -macroalgal values, but did not change the physiology, which is most likely inherently species-specific. Measurements of  $\delta^{13}\text{C}$  were used as indicative of the presence or absence of carbon concentrating mechanisms (CCMs) and as integrative values of the isotope discrimination during carbon assimilation in the lifecycle macroalgae. Based on  $\delta^{13}\text{C}$  signals, macroalgae were classified in four strategies relatives to the capacity of CCM: 1)  $\text{HCO}_3^-$  uptake ( $\delta^{13}\text{C} > -10\text{‰}$ ), 2) using a mix of  $\text{CO}_2$  and  $\text{HCO}_3^-$  uptake ( $-10 < \delta^{13}\text{C} < -30\text{‰}$ ), 3)  $\text{CO}_2$  diffusive entry ( $\delta^{13}\text{C} < -30\text{‰}$ ), and 4) calcifying species ( $\delta^{13}\text{C} > -10\text{‰}$ ). Most species showed a  $\delta^{13}\text{C}$  that indicates a CCM using a mix of  $\text{CO}_2$  and  $\text{HCO}_3^-$  uptake.  $\text{HCO}_3^-$  uptake is also widespread among GC macroalgae, with many *Ochrophyta* species. Few species belonging to *Rhodophyta* relied on  $\text{CO}_2$  diffusive entry exclusively, while calcifying macroalgae species using  $\text{HCO}_3^-$  included only *Amphiroa* and *Jania*. The isotopic signature evidenced the activity of CCM, but it was inconclusive about the preferential uptake of  $\text{HCO}_3^-$  and/or  $\text{CO}_2$  in photosynthesis and the CCM type expressed in macroalgae. In the carbon use strategies study, diverse and species-specific, complementary techniques to the isotopic tools are required.~~

**Keywords:**  $\delta^{13}\text{C}$ -macroalgal, carbon-concentrating mechanisms,  $\text{CO}_2$  diffusive proxy

## 1. Introduction

Macroalgae show a wide diversity of thallus morphologies (e.g., filamentous, articulated, flattened), structural organization (e.g., surface area/volume ratio), and various photosynthetic pigments (e.g., Chlorophyll *a*, *b*, phycocyanin) (Lobban and Harrison, 1994). ~~Based on these features, macroalgae can be classified into only three Phyla, according~~ According to the predominant pigment contents in the thallus, macroalgae are classified in three Phyla. Considering the interaction of morphologies and photosynthetic pigments, are classified ~~or into in~~ dozens of groups ~~considering the interaction of morphologies and photosynthetic pigments~~ (Littler and Littler, 1980; Littler & Arnold, 1982; Balata et al., 2011). For example, the mixture of chlorophyll (*a*, *b*) and carotenoids is dominant in Chlorophyta; chlorophyll (*a*, *c*) and fucoxanthin carotenoid is dominant in Ochrophyta, while Rhodophyta contains chlorophyll (*a*, *d*), carotenoid, and a mixture of phycobilin (e.g., phycocyanin, phycoerythrin, allophycocyanin) (Bold and Wynne, 1978; Masojidek et al., 2004; Gateau et al., 2017). Both traits work as an excellent approximation to explain the fundamentals of metabolism, growth, zonation, and colonization (Littler and Littler, 1980; Littler and Arnold, 1982; Nielsen and Sand-Jensen, 1990; Vásquez-Elizondo and Enríquez, 2017).

~~The thickness of the thallus as a propriety of morphology influences the diffusion boundary layer on surface of the macroalgal, where they carry out the absorption of essential ions and dissolved gases (Hurd, 2000; San Ford and Crawford, 2000).~~ In marine environments, where  $\text{pH} \sim 8.1 \pm 1$ ,  $\text{HCO}_3^-$  accounts for 98% of the total dissolved inorganic carbon (DIC) due to the low diffusion rate of  $\text{CO}_2$  in seawater, ~~that results resulting~~ in a high  $\text{HCO}_3^-:\text{CO}_2$  ratio (150:1) (Sand-Jensen and Gordon, 1984). The limitations to growth imposed by low  $\text{CO}_2$  concentrations in seawater are compensated by carbon concentration mechanisms (CCMs) in most macroalgae that increase the internal inorganic carbon concentration near the site of RuBisCo activity (Giordano et al., 2005). Therefore, the absorption of  $\text{HCO}_3^-$  by most macroalgae is the main-primary source of inorganic carbon for

photosynthesis, but some species depend exclusively on the use of dissolved CO<sub>2</sub> that enters cells by diffusion (Maberly et al., 1992; Beardall and Giordano, 2002; Raven et al., 2002a, b; Giordano et al., 2005). Hence, macroalgal species with productivity limited by lacking CCM's (have low plasticity for carbon inorganic forms uptake) seems to be restricted to subtidal habitats and composed mainly by red macroalgae (but without a morphological patron apparent) (Cornwall et al., 2015, Kübler and Dudgeon, 2015). The rest of the macroalgae with CCM occupies from the intertidal to the deep subtidal.

~~Nevertheless, marine ecosystems have many environmental factors, including~~ The habitat features and environmental conditions in ~~marine ecosystems seawater that~~ modify the main macroalgae photosynthesis drivers, ~~such~~ as light (Anthony et al., 2004; Johansson and Snoeijs, 2014), DIC (Zeebe and Wolf-Gladrow, 2001; Brodeur et al., 2019), and inorganic nutrients (Teichberg et al., 2010; Ochoa-Izaguirre and Soto-Jiménez, 2015). These factors could generate negative consequences for their productivity, principally when they cause resources limitation. Each factor varies from habitat to habitat (e.g., local scale: from intertidal to subtidal and global scale: from temperate to tropical regions), and as in response to these environmental changes, macroalgae can modulate their photosynthetic mechanism (Lapointe and Duke, 1984; Dudgeon et al., 1990; Kübler and Davison 1993, Young ~~et al~~ and Beardall., 2005). The modulation, to increase their photosynthetic activity (up-and-down-regulation processes), implies a physiological acclimation enhancing the transport of DIC (CO<sub>2</sub>, HCO<sub>3</sub><sup>-</sup>) into the cell and its fixation rates (Madsen and Maberly, 2003; Klenell et al., 2004; Zou et al., 2004; Giordano et al., 2005; Enríquez and Rodríguez-Román, 2006; Rautenberger et al., 2015).

The δ<sup>13</sup>C on the thallus of marine macrophytes ~~are indicative of the carbon source used (is a proxy used to identify~~ CO<sub>2</sub> or HCO<sub>3</sub><sup>-</sup>) ~~source~~ in photosynthesis and ~~to allow~~ infer the presence or absence of CCM's (Maberly et al., 1992; Raven et al., 2002a; Giordano et al., 2005). ~~However, the isotopic~~



~~signature may be inconclusive for determine the preference of the carbon source (Roleda and Hurd,~~  
~~2012).~~ Also, the  $\delta^{13}\text{C}$  signal in the algal thallus can be used as an indicator of the physiological state  
of photosynthetic metabolism (Kim et al., 2014; Kübler and Dungeon, 2015). ~~For example,  $\delta^{13}\text{C}$~~   
variability depends, in part, on the life forms as taxonomy, morphology, and structural organization  
(Mercado et al., 2009, Marconi et al., 2011, ~~Roleda and Hurd, 2012;~~ Lovelock et al., 2020~~);~~).  ~~$\delta^{13}\text{C}$~~   
~~but also~~ is also modulated by the interaction to environmental conditions (e.g., light, DIC, and  
nutrients) (Cornelisen et al., 2007; Dudley et al., 2010; Carvalho et al., 2010ab; ~~Roleda and Hurd,~~  
~~2012;~~ Mackey et al., 2015; Rautenberger et al., 2015~~);~~).

In this study, our objective was to investigate the contributions of life forms~~s~~, the changes in the  
habitat features, and environmental conditions to the  $\delta^{13}\text{C}$  macroalgal variability in communities in  
the Gulf of California (GC). To reach our objective, we collected a large stock of macroalgae  
specimens of a diversity of species characterized by ~~a variety of~~various morphological and  
physiological properties. Besides high diversity, in terms of life forms, we selected various shallow  
marine habitats along a latitudinal gradient in the GC or the sample collection, characterized by  
unique and changing environmental factors. The GC features abundant and diverse macroalgae  
populations, ~~which are~~ acclimated and adapted to diverse habitats with environmental conditions,  
determining the light, DIC, and nutrients availability. The  $\delta^{13}\text{C}$  signal from the thallus of macroalgae  
~~was also used to~~were used as indicative of the presence or absence of CCMs and as integrative values  
of the isotope discrimination during carbon assimilation and respiration along lifecycle macroalgae  
~~infer carbon uptake strategies~~ in macroalgae communities in the GC in the function of taxa and  
environmental factors (Maberly et al., 1992; Raven et al., 2002~~a~~; Hepburn et al., 2011; Díaz-Pulido  
et al., 2016). Because the GC is a subtropical zone with high irradiance and specimens were collected  
in the intertidal and shallow subtidal zone, we expect to find a high proportion of species with active

uptake  $\text{HCO}_3^-$  ( $\delta^{13}\text{C} > -10\text{‰}$ ). A third objective was to explore any geographical pattern in the  $\delta^{13}\text{C}$  macroalgal along and between the GC bioregions. Previous studies have indicated changes in the  $\delta^{13}\text{C}$  signal with latitude, mainly related to the light and temperature (Mercado et al., 2009; Marconi et al., 2011; Stepien, 2015; Hofmann and Heesch, 2018; Lovelock et al., 2020). Macroalgae as biomonitors constitute an efficient tool in monitoring programs in large geographical regions (Balata et al., 2011) and for environmental impact assessments (Ochoa-Izaguirre and Soto-Jiménez, 2015).

## **2. Materials and Methods**

### **2.1. Gulf of California description**

The Gulf of California is a subtropical, semi-enclosed sea of the Pacific coast of Mexico, with exceptionally high productivity being the most important fishing regions for Mexico and one of the most biologically diverse worldwide marine areas (Zeitzschel, 1969; Espinosa-Carreón and Valdez-Holguín 2007; Lluch-Cota et al., 2007; Páez-Osuna et al., 2017). GC represents only 0.008% of the area covered by the seas of the planet (265,894 km<sup>2</sup>, 150 km wide, and 1000 km long covering >9 degrees latitude) but has a high physiographic diversity and is biologically mega-diverse with many endemic species, including ~ 766 macrofauna species and/or sub-species where the major number belong to Arthropoda (118 spp) and Mollusca (460) taxa (Brusca et al., 2005; Wilkinson et al., 2009; Espinosa-Carreón and Escobedo-Uriás, 2017) and 116 macroalgae species (Norris, 1975, 1985; Espinoza-Avalos, 1993).

Regionalization criteria of the GC include phytoplankton distribution (Gilbert and Allen, 1943), topography (Rusnak et al., 1964) and depth (Álvarez-Borrego, 1983), oceanographic characteristics (Roden and Emilson, 1979; Álvarez-Borrego, 1983; Marinone and Lavin 2003), biogeography (Santamaría-del-Ángel et al., 1994), and bio-optical characteristics (Bastidas-Salamanca et al.,

2014). The topography is variable along with GC, includes submarine canyons, basins, and variable continental platforms. Besides, GC presents complex hydrodynamic processes, including internal waves, fronts, upwelling, vortices, mixing of tides. The gulf's coastline is divided into three shores: extensive rocky shores, long sandy beaches, numerous scattered estuaries, coastal lagoons, and open muddy bays, tidal flats, and coastal wetlands (Lluch-Cota et al., 2007).

The Gulf of California is different in the north and the south, related to a wide range of physicochemical factors. The surface currents seasonally change direction and flow to the southeast with maximum intensity during the winter and to the northwest in summer (Roden, 1958). The northern part is very shallow (<200 m deep averaged), divided into Upper Gulf, northern Gulf, and Grandes Islas. The surrounding deserts largely influence this region (Norris, 2010) shows marked seasonal changes in coastal surface seawater temperatures (Martínez-Díaz de León et al., 2006; Marinone, 2007). Tidal currents induce a significant cyclonic circulation through June to September and anticyclonic from November to April (Carrillo et al., 2002; Bray, 1988; Velasco-Fuentes and Marinone, 1999; Martínez-Díaz-de-León, 2001). The southern part consists of a series of basins whose depths increase southwards (Fig. 1). The intertidal macroalgae in the southern region are subject to desiccation, mostly during summer. The water column's physicochemical characteristics are highly influenced by the contrasting climatic seasons in the GC, the dry season (nominally from November to May), and the rainy season (from June to October). Annual precipitation (1,080 mm y<sup>-1</sup>) and evaporation (56 mm y<sup>-1</sup>) rates registered during the past 40 years were 881±365 mm y<sup>-1</sup> and 53±7 mm y<sup>-1</sup>, respectively (CNA, 2012).

Previous macroalgae floristic studies of the GC, report around 669 species, including 116 endemic species (Norris, 1975; Espinoza-Avalos, 1993; Pedroche and Senties, 2003). Many endemic species currently have a wide distribution along the Pacific Ocean coast, but with GC origin

(Dreckman, 2002; Aguilar-Rosas et al., 2014). Based on oceanographic characteristics (Roden and Groves, 1959) and in the endemic species distribution (Aguilar Rosas and Aguilar Rosas, 1993; Espinoza-Avalos, 1993), the GC can be classified into three phycofloristic zones: 1) the first zone located from the imaginary line connecting San Francisquito Bay, B.C. to Guaymas, Sonora, with 51 endemic species. 2) the second zone with an imaginary line from La Paz ~~bay~~-Bay (B.C.S.) to Topolobampo (Sinaloa) with 41 endemic species. 3) the third zone is located with an imaginary line from Cabo San Lucas (B.C.S.) to Cabo Corrientes (Jalisco) with ~~ten~~10 endemic species. Besides, 14 endemic species are distributed throughout the GC (Espinoza-Ávalos, 1993). The macroalgal communities are subject to the changing environmental conditions in the diverse habitats in the GC that delimits their zonation, which tolerates a series of anatomical and physiological adaptations to water movement, temperature, sun exposure, and light intensities, low pCO<sub>2</sub>, desiccation (Espinoza-Avalos 1993).

## 2.1 Macroalgae sampling

In this study, the GC coastline (21°-30°N latitude) was divided into six coastal sectors based on the three phycofloristic zones along peninsular and continental GC coastlines (Fig. 1a). In each coastal sector, selected ecosystems and representative habitats were sampled based on macroalgae communities' presence and habitat characterization. Habitats were classified by substrate type (e.g., sandy-rock, rocky shore), hydrodynamic (slow to faster water flows), protection level (exposed or protected sites), and immersion level (intertidal or subtidal) (Fig. 1b).

Based on the local environmental factors, ~~4-5~~ macroalgae specimens (~~4-5~~) of the most representative species were gathered by hand (free diving) during low tide. A total of 809 composite samples were collected from marine habitats along both GC coastlines. The percentages of specimens collected for

the substrate type were sandy-rock 28% and rocky shores 72% based on the habitat features. ~~Related~~  
~~to~~ In the hydrodynamic, 30% of the specimens were collected in habitats with slow to moderate and  
70% with moderate to fast water movement. Regarding the protection level, 57% were exposed  
specimens, and 43% were protected. Finally, 56% were intertidal and 44% subtidal macroalgae  
organisms concerning the emersion level. About half of the protected specimens were collected in  
isolated rock pools, which was noted.

In 4-5 sites of each habitat, we measured *in situ* the salinity, temperature, and pH by using a  
calibrated multiparameter sonde (Y.S.I. 6600V) and the habitat characteristics mentioned above  
noted. Besides, composite water samples were collected for a complementary analysis of nutrients,  
alkalinity (and their chemical components), and  $\delta^{13}\text{C}$ -DIC (data non-included). Briefly, the  
representative habitats were classified by pH levels in  $>9.0$  “alkalinized”, 7.9-8.2 ‘typical’ and  $<7.9$   
“acidified”. Based on the temperature in colder  $<20^\circ\text{C}$ , typical  $20\text{-}25^\circ\text{C}$ , and warmer  $>25^\circ\text{C}$ . 72% of  
the specimens were collected at typical pH values, 22% ~~in~~ alkalinized, and 6% in acidified seawater.  
Regarding the temperature, about 55% of the specimens were collected at typical, 31% at warmer,  
and 14% at colder seawaters. Regarding salinity, most of the ecosystems showed typical values for  
seawater ( $35.4 \pm 0.91$  ups, from 34.5 to 36.1 ups). In this study, the collection surveys were conducted  
during spring (March-April) and dry season (nominally from November to May) from ~~2009-2008~~ to  
2014. Only in few selected ecosystems located at C1, ~~and~~ C2, and C3 sectors, one sampling survey  
was conducted at the end of the rainy season (nominally from June to October in 2014). Thus, these  
ecosystems were possible to include habitat with a salinity range varying from estuarine ( $23.5 \pm 3.0$   
ups) to hypersaline ( $42.7 \pm 7.0$  ups) values. These habitats were mainly isolated rock pools, and only  
a few were sites near tidal channels receiving freshwater discharges. About 95% of the specimens  
were collected at typical seawater salinity (34-36 ups) and only 1.5 and 3.5% in estuarine ( $<30$  ups)

and hypersaline (>37 ups) environments, respectively. Detailed information on the selected shallow marine ecosystems, habitat characterization, and environmental conditions is summarized in the inserted table in Fig. 1.

## **2.2 Macroalgae processing and analysis of the isotopic composition of carbon**

The collected material was washed *in situ* with surface seawater to remove the visible epiphytic organisms, sediments, sand, and debris and then thoroughly rinsed with MilliQ water. The composite samples were double-packed in a plastic bag, labeled with the locality's name and collection date, placed in an ice-cooler box to be kept to 4°C, and immediately transported to the laboratory UAS-Facimar in Mazatlán. In the field, sample aliquots were also preserved in 4% v/v formaldehyde solution for taxonomic identification to the genus or species level (when possible). The following GC macroalgal flora identification manuals were consulted: Dawson, 1944; 1954; 1956; 1961; 1962; 1963; Setchell and Gardner, 1920; 1924; Abbott and Hollenberg, 1976; Ochoa-Izaguirre et al., 2007; Norris, 2010).

In the laboratory, macroalgae samples were immediately frozen at -30°C until analysis. Then, samples were freeze-dried at -38°C and 40 mm Hg for 3 days, upon which they were ground to a fine powder and exposed to HCl vapor for 4 h (acid-fuming) to remove carbonates and dried at 60°C for 6 h (Harris et al. 2001). Aliquots of ~5 mg were encapsulated in tin cups (5x9 mm) and stored in sample trays until analysis. Macroalgae samples were sent to the Stable Isotope Facility (SIF) at the University of California at Davis, CA, USA. Natural <sup>13</sup>C relative abundance relative to <sup>12</sup>C in samples was determined with mass spectrometry, using a Carlo Erba elemental analyzer attached to a Finnigan Delta S mass spectrometer equipped with a Europa Scientific stable isotope analyzer (ANCA-NT 20-20) and a liquid/solid preparation unit (PDZ, Europa, Crews, UK). Isotope ratios of

the samples were calculated using the equation  $\delta (\text{‰}) = (R_{\text{sample}}/R_{\text{standard}} - 1) \times 1000$ , where  $R = {}^{13}\text{C}/{}^{12}\text{C}$ . The  $R_{\text{standard}}$  is relative to the international V-PDB (Vienna PeeDee Belemnite) standard. During the isotopic analysis, the SIF lab used different certified reference materials (e.g., IAEA-600, USGS-40, USGS-41, USGS-42, USGS-43, USGS-61, USGS-64, and USGS-65) for the analytical control quality. The analytical uncertainties reported for the SIF lab were 0.2‰ for  $\delta^{13}\text{C}$  (<https://stableisotopefacility.ucdavis.edu/13cand15n.html>). We also included triplicate aliquots of several specimens of the same species and condition, collected from one patch, or attached to the same substrate, to assess the method error by sampling and processing procedural. The methodological uncertainties were <0.4‰.

### 2.3. Analysis of $\delta^{13}\text{C}$ -macroalgal variability

The variability of  $\delta^{13}\text{C}$  values in macroalgae was analyzed in function of the taxonomy (phylum, genus, and species) and morpho-functional groups (e.g., thallus structure, growth form, branching pattern, and taxonomic affinities; Balata et al. 2011; Ochoa-Izaguirre and Soto-Jiménez, 2015).

The carbon fixation strategies in the macroalgae communities of the GC were ~~identify-identified~~ by  $\delta^{13}\text{C}$  (Hepburn et al., 2011; Díaz-Pulido et al., 2016), in agreement with the Maberly et al. (1992) and Raven et al. (2002a) thresholds. So, macroalgae were classified into four strategies for DIC uptake: 1) CCM-only by active uptake  $\text{HCO}_3^-$  ( $\delta^{13}\text{C} > -10\text{‰}$ ), 2) CCM active uptake  $\text{HCO}_3^-$  and/or diffusive uptake  $\text{CO}_2$  ( $\delta^{13}\text{C} < -11$  to  $-30\text{‰}$ ), 3) Non-CCM,  $\text{CO}_2$  by diffusion only ( $\delta^{13}\text{C} < -30\text{‰}$ ), 4) Calcifying with different carbon-use strategies related to different modes of calcification. The measured  $\delta^{13}\text{C}$ -macroalgal signals are integrative of the discrimination by photosynthesis ( $\Delta^{13}\text{C}_\text{p}$ ) on the carbon source ( $\delta^{13}\text{C}$ -DIC in seawater), respiration ( $\Delta^{13}\text{C}_\text{r}$ ), and probable  $\text{CO}_2$  leak out inside the cell during the CCM process (Sharkey and Berry, 1985; Raven et al., 2005; Carvalho et al., 2009a,b).

Macroalgae were grouped according to their characteristics morpho-functional proposed initially by Littler and Littler (1980) and modified by Balata et al. (2011). Most of the macroalgae species showed a limited distribution along the GC coastlines. Few cosmopolites' species included *Colpomenia tuberculata*, *Sargassum sinicola*, *Padina durvillei*, and *Ulva lactuca*. ~~Not Also, not~~ all morphofunctional groups and taxon were present in every site during each sampling survey, and the sample size in each group varied for taxa, location, and time.

A basic statistical analysis of  $\delta^{13}\text{C}$  values in different macroalgae groups was applied to distribute and calculate the arithmetic mean, standard deviation, minimum and maximum. Because not all macroalgal species were present in sufficient numbers at different collection habitats, several macroalgal groups were not considered for statistical analysis. ~~Regarding the life form, w~~We compared among taxon and morphofunctional groups, collected in the same habitat (within-subjects factor) by multivariate analysis of variance. When differences were noted, a Tukey-Kramer HSD (Honestly Significant Difference) test was performed. Besides, variations of  $\delta^{13}\text{C}$  macroalgal in specimens of the same morpho-functional and taxon collected in different habitats were also investigated with a Kruskal-Wallis test.

~~In this study, t~~The relationships between  $\delta^{13}\text{C}$  with each independent variable ~~related to the~~ inherent to the macroalgae properties (taxon and morphology), biogeographical collection zone (GC coastline and coastal sector), habitat features (substrate, hydrodynamic, protection, and emersion level) and environmental conditions (temperature, pH, and salinity), were examined through simple and multiple linear regression analyses. Excepting temperature, pH, and salinity, most of the independent variables are categorical independent variables. However, these continue variables were also categorized, such as previously was described. ~~Analyses of s~~Simple linear regression analyses were performed to establish the relationships between  $\delta^{13}\text{C}$ -macroalgal with each environmental



parameter analyzed as possible driving factors (e.g., temperature, salinity, [and](#) pH). Multiple linear regression analyses were conducted to evaluate the combined effects of those independent variables (macroalgae properties, biogeographical collection zone, habitat features, and environmental conditions) on the  $\delta^{13}\text{C}$ -macroalgal. In the multivariable regression model, the dependent variable,  $\delta^{13}\text{C}$ -macroalgal, is described as a linear function of the independent variables  $X_i$ , as follows:  $\delta^{13}\text{C}$ -macroalgal =  $a + b_1(X_1) + b_2(X_2) + \dots + b_n(X_n)$  (1). Where  $a$  is regression constant (it is the value of intercept and its value is zero);  $b_1$ ,  $b_2$ , and  $b_n$ , are regression coefficients for each independent variable  $X_i$ . From each one of the fitted regression models, we extracted the estimated regression coefficients for each of the predictor variables (e.g., Bayesian Information Criterion (BIC), Akaike Information Criterion (AIC), root-mean-square error (RMSE), Mallow's  $C_p$  criterion, F Ratio test, p-value for the test ( $\text{Prob} > F$ ), coefficients of determination ( $R^2$ ) and the adjusted  $R^2$  statistics) (SAS Institute Inc., 2018). All regression coefficients were used as indicators of the quality of the regression (Draper and Smith, 1998; Burnham and Anderson, 2002). Kolmogorov-Smirnov normality test was applied for all variables, and all were normally distributed. Most of the  $\delta^{13}\text{C}$  values in each group showed a normal distribution. For all statistical tests, a probability  $P < 0.05$  was used to determine statistical significance. The statistical analysis of the results was done using JMP 14.0 software (SAS Institute Inc.).

### 3. Results

#### 3.1. Taxonomy and morpho-functional groups

Sampled specimens belong to three Phyla, 63 genera, and [170](#) species. The Phyla were identified as Chlorophyta (25%), Ochrophyta (22%), and Rhodophyta (53%). The most representative genus (and their species) were *Ulva* (*U. lactuca*, *U. lobata*, *U. flexuosa*, and *U. intestinalis*), *Codium* (*C.*

*amplivesiculatum* and *C. simulans*), *Chaetomorpha* (*C. antennina*), *Padina* (*P. durvillaei*), *Dictyota* (*D. dichotoma*), *Colpomenia* (*C. tuberculata* and *C. sinuosa*), *Sargassum* (*S. sinicola* and *S. horridum*), *Amphiroa* (*Amphiroa* spp.), *Spyridia* spp, *Polysiphonia* spp., *Gymnogongrus* spp., *Gracilaria* (*G. vermiculophylla*, *G. pacifica* and *G. ~~erispate~~crispata*), *Hypnea* (*H. pannosa* and *H. johnstonii*) *Grateloupia* (*G. filicina* and *G. versicolor*), and *Laurencia* (*L. papillosa* and *L. pacifica*).

~~In our study,~~ The endemic species includes Chlorophyta *Codium amplivesiculatum*, Rhodophyta *Laurencia papillosa*, *Chondracanthus squarrulosa*, *Gracilaria spinigera*, and *~~Graeilaria~~ subsecundata*, and Ochrophyta *Cutleria hancockii*, *Sargassum herphorizum*, *~~Sargassum~~ johnstonii*.

An analysis of the biogeographical diversity among sectors evidenced that P3 (43 genera of 63, 68%) and C3 (63%) at north recorded the highest number of the genus, followed by C1 (38%) and P1 (29%) at the south, and P2 (27%) and C2 (22%). The same pattern was observed in the species diversity, zones P3 (94 of 167 species, 56%) and C3 (52%) at the north, C1 (34%) and P1 (25%) at the south, and C2 and P2 (19-20%) at the center.

The morphofunctional groups identified were 21, ~~of which t.~~ The most common were C-tubular (6 spp., n=69; C-Blade-like (6 spp, n=55); C-Filamentous uniseriate (17 spp, n=49); C-Erect thallus (5 spp, n=33); O-Compressed with branched or divided thallus (19 spp., n=92); O-Thick leathery macrophytes (12 spp., n=104); O-Hollow with spherical or subspherical shape (4 spp, n=87); R-Large-sized corticated (57 spp., n=225); R-Filamentous uniseriate and pluriseriate with erect thallus (9 spp., n=48); and R-Large-sized articulated corallines (6 spp, n=17). The diversity, in terms of presence/absence of the morphofunctional groups, varied among coastline sectors, higher in C3 (16 of 21, 76%) and P3 (71%) at the north, followed by C1 (57%) and P1 (48%) at the south, and C2 and P2 and (42-48%) at the center of both GC coastlines.

### 3.2. $\delta^{13}\text{C}$ -macroalgal variability in function of taxonomy and morpho-functional groups

The variability of  $\delta^{13}\text{C}$  values in macroalgae was analyzed by taxon ~~in the~~ (phylum, genus, species); and morphofunctional groups classified by habitat, coastal sector, and collection season. A complete list of the results of  $\delta^{13}\text{C}$  in 170 macroalgae species is provided in Supporting Information (Table SI-1). Firstly,  $\delta^{13}\text{C}$  values analyzed by phylum showed a unimodal distribution with a peak at  $-14 \pm 1.4\text{‰}$  (Fig 2). ~~where~~ Ochrophyta ( $-21.5$  to  $-2.2\text{‰}$ ,  $-12.5 \pm 3.7\text{‰}$ ), displayed ~~the significantly higher~~ values ~~from than~~ Chlorophyta ( $-25.9$  to  $-5.5\text{‰}$ ,  $-14.5 \pm 3.0\text{‰}$ ) and Rhodophyta ( $-34.6$  to  $-4.5\text{‰}$ ,  $-14.8 \pm 3.9\text{‰}$ ). The  $\delta^{13}\text{C}$ -macroalgal values (average  $\pm$  SD) for the genus of Chlorophyta, Ochrophyta, and Rhodophyta (Fig. 3) varied from  $-33.8 \pm 1.1\text{‰}$  for *Schizymenia* to  $-7.8 \pm 0.7\text{‰}$  for *Amphiroa*. Based on the highest values, specimens of three Phyla ~~with relatively high showed~~  $\delta^{13}\text{C}$  values ( $> -10\text{‰}$ ), evidenced the presence of CCM's by active uptake of  $\text{HCO}_3^-$  (strategy 1) (Fig. 3). For example, *Caulerpa*, *Cladophora*, *Codium*, *Ulva* ~~in for~~ Chlorophyta; *Colpomenia*, *Dictyota*, *Padina*, *Sargassum* ~~for~~ Ochrophyta, and *Hypnea* and *Polysiphonia* ~~for~~ Rhodophyta showed  $\delta^{13}\text{C}$  values  $> -10\text{‰}$ . Likewise, high  $\delta^{13}\text{C}$  values were observed in the calcifying macroalgae ~~species-genus like~~ *Amphiroa* and *Jania*, under -strategy 4 (Fig. 3c).  $\delta^{13}\text{C}$  values lower than  $-30\text{‰}$  that denote uptake of  $\text{CO}_2$  by diffusion (strategy 3), were observed only in Rhodophyta ~~in *Schizymenia* *Schizymenia*, *Halymenia*, and *Gigartina*~~. However, most species showed large  $\delta^{13}\text{C}$  variabilities signals that evidence a mechanism that uses a mix of  $\text{HCO}_3^-$  and  $\text{CO}_2$  for photosynthesis (strategy 2).

~~A-m~~ Multiple comparison analyses revealed significant differences in the  $\delta^{13}\text{C}$ -macroalgal values among genera, ordered as *Schizymenia*  $<$  *Polysiphonia*  $<$  *Ulva*, *Gracilaria* and *Spyridia* ( $-16.1 \pm 0.6\text{‰}$  to  $-15.1 \pm 0.2\text{‰}$ )  $<$  *Gymnogongrus*, *Laurencia*, *Hypnea*, *Cladophora*, *Dictyota*, *Sargasumm*, *Chaetomorpha*, and *Grateloupia* (from  $-15.4 \pm 0.7\text{‰}$  to  $-13.8 \pm 0.8\text{‰}$ )  $<$  *Codium* and *Padina* ( $-12.5 \pm 2.4\text{‰}$  to  $-12.4 \pm 2.5\text{‰}$ )  $<$  *Colpomenia* and *Amphiroa* ( $-9.2 \pm 0.3$  to  $-7.8 \pm 0.7\text{‰}$ ) ( $F=16.81$ ,

338 p<0.001).

339 Aggrupation of  $\delta^{13}\text{C}$  values based on morpho-functional features ~~on-macroalgae-is~~ displayed in Fig.  
340 4. The most representative groups in the phylum Chlorophyta varied from  $-15.8 \pm 0.3\text{‰}$  for C-Tubular  
341 to  $-12.4 \pm 0.5\text{‰}$  for C-~~thallus~~ Thallus erect. The phylum Ochrophyta includes O-Thick leathery with  
342 the lowest mean ( $-14.8 \pm 0.3\text{‰}$ ) and O-Hollow with a spherical or subspherical shape with the highest  
343 values ( $-9.2 \pm 0.3\text{‰}$ ). The lowest and highest  $\delta^{13}\text{C}$  values for Rhodophyta were observed for R-  
344 flattened macrophytes ( $-24.0 \pm 9.6\text{‰}$ ) and R-Larger-sized articulated coralline ( $-7.89 \pm 0.75\text{‰}$ ),  
345 respectively. Significant differences were observed among groups, which were ordered as follows:  
346 R-~~flattened~~ Flattened macrophytes < R-~~blade~~ Blade like < C-Tubular < O-Tick leathery and R-Large  
347 size corticated < C-Blade like and C-Filamentous uniseriate < C-~~Erect thallus~~ Thallus erect and O-  
348 Compressed with branch < O-Hollow with spherical < R-Larger-sized articulated coralline.

349 High intraspecific variability in  $\delta^{13}\text{C}$  signal for the more representative genera of each taxon is  
350 showed in Table 1-3. For *Codium*, *C. brandegeei* ( $11.8 \pm 1.2\text{‰}$ ) and *C. simulans* ( $-11.4 \pm 2.2\text{‰}$ )  
351 showed higher  $\delta^{13}\text{C}$  values than *C. amplivesiculatum* ( $-14.4 \pm 2.7\text{‰}$ ). *Colpomenia* species had higher  
352  $\delta^{13}\text{C}$  values than the other genera, with higher values for *C. tuberculata* ( $-8.7 \pm 3.2\text{‰}$ ) than  
353 *Colpomenia* sp. ( $-10.9 \pm 3.6\text{‰}$ ) and *C. sinuosa* ( $-10.2 \pm 2.9\text{‰}$ ). *Gracilaria* showed comparable  $\delta^{13}\text{C}$   
354 values in the four species (from  $-16.4 \pm 1.6\text{‰}$  for *G. pacifica* to  $-15.5 \pm 2.4\text{‰}$  for *Gracilaria* sp.).  
355 *Hypnea* showed non-significant  $\delta^{13}\text{C}$  differences in three representative species ( $-16.4 \pm 1.7\text{‰}$  for *H.*  
356 *spinella* to  $-14.9 \pm 2.3\text{‰}$  for *Hypnea* sp.). *Laurencia* sp. ( $-12.9 \pm 1.2\text{‰}$ ) was higher than *L. pacifica* ( $-$   
357  $14.9 \pm 2.2\text{‰}$ ), while *Padina* sp. ( $-11.1 \pm 1.5\text{‰}$ ) higher than *P. durvillaei* ( $-13.2 \pm 2.6\text{‰}$ ). *Sargassum* was  
358 one of the most diverse genera studied with six representative species, with  $\delta^{13}\text{C}$  values ordered as  
359 follow: *S. horridum* = *S. sinicola* = *S. johnstonii* ( $-15.5 \pm 2.9$  to  $-15.1 \pm 2.4\text{‰}$ ) < *S. lapazeanum* ( $-$   
360  $14.5 \pm 1.6\text{‰}$ ) = *Sargassum* sp. ( $-14.2 \pm 2.3\text{‰}$ ) < *S. herphorizum* ( $-13.6 \pm 1.6\text{‰}$ ). *Spyridia* sp. ( $-$

17.0±1.2‰) and *S. filamentosa* (-15.8±3.8‰) showed non-significant differences. The six representative species of *Ulva* were divided into two morphological groups, filamentous and laminates. Filamentous species that averaged -16.3±2.0‰ for *U. clathrata*, -16.0±3.6‰ for *U. flexuosa*, -15.7±1.7‰ for *U. acanthophora* and -15.3±2.5‰ for *U. intestinalis* and *Ulva* laminates that included *U. linza* (-15.5±2.4‰) and *U. lactuca* (-14.1±3.1‰). Non-significant differences were observed between morphological groups and among species. A high intra-specific variability, 11-28%, explains average overlapping.

### 3.3. $\delta^{13}\text{C}$ -macroalgal variability in coastal sectors

A variety of macroalgal assemblages were documented along the GC coastlines, with differences in the taxonomic composition according to their fico-floristic region. Multiple comparison analysis of  $\delta^{13}\text{C}$  signals evidenced significant differences between the most common genus and species of macroalgae between and within assemblages grouped by coastal sector, season and collecting year (Supplementary Information Tables SI-2-3). For example, genus *Padina* (e.g., *P. durvillei*) and *Ulva* (e.g., *U. lactuca*), collected in C1 sector during the rainy season, showed lower  $\delta^{13}\text{C}$  values than in other sectors. Differences in the  $\delta^{13}\text{C}$  signal are mainly related to the carbon uptake strategies of the macroalgae (Fig. 5). Even though most species inhabiting the GC coastal sectors displayed domination of strategies based on active CCM's, but the tendencies were different between taxa and coastal regions. The strategy 2 that combined different with mixing DIC sources of DIC were are dominant in all regions and taxa (60-90%). Exceptions were observed in the P1 (68%) and C1 (37%) regions for Ochrophyta, where the specialized strategy 1 of only  $\text{HCO}_3^-$  user ~~were~~ was significant. The strategy 3 based on ~~only the~~ use of  $\text{CO}_2$  was observed in the peninsular coast in P2 and P3 for Rhodophyta with 2-3.3%. Overall, more negative  $\delta^{13}\text{C}$  values ~~in macroalgae specimens' values of~~

~~the same genus~~ were observed at continental (C2) compared to the peninsular coastline (P1-P3) and ~~more negative~~ southward than northward.

### 3.4. $\delta^{13}\text{C}$ -macroalgal variability in function of taxonomy and habitat features and environmental conditions

Variability of  $\delta^{13}\text{C}$  values for the most representative genera was evaluated by multiple comparative analyses in the habitat features' function, including the substrate, hydrodynamic, and emersion level. Large  $\delta^{13}\text{C}$  variability observed between specimens of the same genus collected in the different habits ~~does~~ not show any significant pattern, and non-significant differences were observed. An exception was observed with the emersion level (showed in Fig. 6), where intertidal specimens recorded lesser negative values than subtidal in most macroalgae genus. For example, for *Hydroclathrus* (intertidal  $-5.7 \pm 0.9\text{‰}$ ; subtidal  $-11.4 \pm 5.9\text{‰}$ ), *Amphiroa* (intertidal  $-6.9 \pm 1.5$ ; subtidal  $-9.9 \pm 6.1$ ), *Hypnea* (intertidal  $-13.5 \pm 2.5\text{‰}$ ; subtidal  $-18.6 \pm 1.8\text{‰}$ ), and *Laurencia* (intertidal  $-13.5 \pm 1.3\text{‰}$ ; subtidal  $-17.1 \pm 1.8\text{‰}$ ). Exceptions were observed for *Polysiphonia* (intertidal  $-19.7 \pm 2.2\text{‰}$ , subtidal  $-14.9 \pm 6.7\text{‰}$ ), *Spyridia* (intertidal  $-16.9 \pm 3.3\text{‰}$ , subtidal  $-13.2 \pm 0.7\text{‰}$ ) and *Colpomenia* (intertidal  $-9.4 \pm 3.4\text{‰}$ , subtidal  $-7.7 \pm 1.3\text{‰}$ ).

Non-significant differences were observed for the same genera at different temperatures ranges, except for *Grateloupia* (cold,  $-19.2 \pm 4.7\text{‰}$ , typical  $-14.4 \pm 2.2\text{‰}$ , warm  $-14.5 \pm 2.2\text{‰}$ ) and *Polysiphonia* (cold,  $-21.0 \pm 0.4\text{‰}$ , typical  $-18.1 \pm 5.5\text{‰}$ , warm  $-17.9 \pm 2.3\text{‰}$ ) with more negative values in colder than warmer waters ( $F=6.42$ ,  $p<0.001$ ). Neither significant difference was observed in  $\delta^{13}\text{C}$  values in macroalgae specimens from the different genus in the same temperature range. ~~For example, Colpomenia (cold  $-8.3 \pm 2.4\text{‰}$ , typical  $-9.4 \pm 3.7\text{‰}$ , warm  $-9.2 \pm 2.6\text{‰}$ ), Codium (cold~~

~~11.9±1.9‰, typical 12.5±3.0‰, warm 13.6±0.6‰), and *Padina* (cold 11.3±2.5‰, typical 11.8±1.7‰, warm 13.4±2.7‰)~~ (Fig. 7a).

Significant differences were observed among genus related to the pH level at seawater (Fig. 7b). Under typical pH seawater, *Amphiroa* and *Colpomenia* were 1-2‰ more negatives than in alkaline waters, while *Ulva* and *Spyridia* were 3-5‰ less negative than in acidic waters. *Amphiroa* and *Colpomenia* were not collected in acidic water, and neither *Spyridia* in alkaline waters to compare. Another genus also showed extremes values between alkaline (*Tacanoosca* -7.6±1.0‰) and acidic waters (*Schizymenia*, -32.9±2.0‰). The following order was observed in the genus collect at the three pH ranges: alkaline > typical > acidic. Significant differences were observed for genus *Ahnfeltiopsis*, *Caulerpa*, *Gymnogongrus*, *Padina*, and *Ulva*, with higher values at alkaline than in acidic waters. Values of  $\delta^{13}\text{C}$  for specimens of the same genus collected at typical pH waters are mostly overlapped between those for alkaline and acidic seawaters. Non-significant differences in  $\delta^{13}\text{C}$  values were observed for *Grateloupia*, *Hypnea*, and *Polysiphonia* concerning pH-type waters.

We analyzed the carbon uptake strategies on macroalgal assemblages in the function of environmental factors like temperature, pH, and salinity (Fig. 8). Regarding the  $\delta^{13}\text{C}$  variability for all data set in response to temperature and salinity, a non-significant trend was observed between  $\delta^{13}\text{C}$ -macroalgal in both parameters' function. A poor but significant correlation was observed between  $\delta^{13}\text{C}$  and pH ( $R^2 = 0.04$ ) (Table 4). The proportion of specimens with a strategy of only  $\text{HCO}_3^-$  use was different between environmental factors and taxa (previously described). Ffor example, Ochrophyta showed the highest proportion (35%) in colder temperature, in pH-Alkaline (31%), and at typical salinity regimen (27%), while Chlorophyta enhanced to 30% in acid pH and Rhodophyta recorded 21% at normal seawater. The opposite strategy (only use of dissolved  $\text{CO}_2$ ) ~~that~~ was observed only in Rhodophyta, t. The highest percentage was observed in estuarine salinity

regimen (10%).

### 3.5. Variation latitudinal of $\delta^{13}\text{C}$ -macroalgal

The  $\delta^{13}\text{C}$ -macroalgal variation in the GC biogeography was evaluated by ~~regression~~-linear regression analysis between  $\delta^{13}\text{C}$  values along the nine degrees latitude in both GC coastlines. A non-significant latitudinal trend was observed for datasets, but for the three ~~taxa's~~-Phyla's most representative genera,  $\delta^{13}\text{C}$  values correlated with latitude (Fig. 9). In Chlorophyta, with the higher genera number,  $\delta^{13}\text{C}$  values increased with latitude, with low but significant correlation. Contrarily, in Ochrophyta and Rhodophyta specimens, the  $\delta^{13}\text{C}$  values decreased non-significantly with latitude.

In the most representative morphofunctional groups, sSignificant correlations ( $p < 0.001$ ) were observed for  $\delta^{13}\text{C}$ -macroalgal *versus* latitude ~~in the most representative morphofunctional groups~~ (Fig. 10). Representative morphofunctional groups of Chlorophyta (e.g., C-Tubular, C-Filamentous uniseriate), showed a positive correlation, while those belonging to Ochrophyta (e.g., O-~~T~~thick leathery;) and Rhodophyta (e.g., R-~~large~~-Large sized corticated) showed a negative trend with latitude.

### 3.6. Analyses of $\delta^{13}\text{C}$ macroalgal variability

An analysis of the effects, independent and combined, on the  $\delta^{13}\text{C}$ -macroalgal variability related to life form and environmental factors, ~~was~~ was conducted. Firstly, simple linear regression analyses were performed to evaluate the dependent variable's prediction power ( $\delta^{13}\text{C}$ -macroalgal) in the function of several independent variables controlling the main macroalgae photosynthesis drivers (light, DIC, and inorganic nutrients). Regression coefficients were estimated for each fitted regression model, which is used as indicators of the quality of the regression (Draper and Smith, 1998; Burnham and Anderson, 2002) as was described in Methods; however, our results description focused on the



coefficients of determination ( $R^2$  and adjusted  $R^2$ ). The coefficient  $R^2$  describes the overall relationship between the independent variables  $X_i$  with the dependent variable  $Y$  ( $\delta^{13}\text{C}$ -macroalgal), and it is interpreted as the % of contribution to the  $\delta^{13}\text{C}$  variability. ~~While~~ In comparison, the adjusted  $R^2$  statistics compensate for possible confounding effects between variables.

Results of the analysis of the relationships between  $\delta^{13}\text{C}$  with each independent variable are summarized in Table 4. Regarding the inherent macroalgae properties, Phyla explain only 8% variability, the morphofunctional properties 35%, ~~and taxon by~~ genus 46%, and ~~by~~ species 57%.

~~The biogeographical collection zone,~~ In terms of coastline (continental vs. peninsular) and coastal sectors (C1-C3 and P1-P3), the biogeographical collection zone explained a maximum of 5% variability. ~~Related to the habitat features,~~ Only the emersion level (6%) contributed to the  $\delta^{13}\text{C}$  variability related to the habitat features. The contribution of the seawater's environmental conditions was marginal for pH (4%) and negligible for temperature and salinity. A marginal reduction in the percentage of contribution was observed for Phyla (1%) and morphofunctional properties (1%), but significant for genus (5%) and species (10%).

Multiple regression analyses were also performed to interpret the complex relationships among  $\delta^{13}\text{C}$ -macroalgal, considering the life form (morphofunctional and taxon by genus) and their responses to environmental parameters. Results for the fitted regression models performed for morphofunctional groups (Table 5) and genus (Table 6) evidenced that the effect of the coastal sector and pH ranges on the  $\delta^{13}\text{C}$ -macroalgal increased the contribution by 9-10% each one. The emersion level increased by 5-6%, the contribution respect to individual effect of morphofunctional group and genus, the temperature and pH in 1 and 3%, respectively, while salinity decreased by 1-2%. Adding the effect of the biogeographical collection zone, represented by the coastline sector, to those for morphofunctional group (Table 5) and genus (Table 7), a notable increase of 11-12% was observed.

~~The full model e~~Considering the combined effect of the coastline sector + Habitats features for Morphofunctional group or Genus (Table 7), the full model showed  $R^2$  of 0.60 and 0.71. In contrast, Coastline sector + Environmental conditions + Morphofunctional group or Genus the  $R^2$  increased to 0.62 and 0.72, respectively. The interactive explanations of environmental factors increased the explanation percentage of  $\delta^{13}\text{C}$  variability; however, these contributions were significantly lower than the explained by life forms, such as the morphofunctional properties and taxa by genus and species.

The combined effect of environmental conditions on the  $\delta^{13}\text{C}$  variability was tested for the best-represented genus and morphological groups. Results evidenced that 9 of 21 morphological groups showed significant effects on the  $\delta^{13}\text{C}$  variability (Table 8), five increasing and four decreasing the model constant of  $\delta^{13}\text{C} = -14.2\text{‰}$ . For example, for the O-Hollow with spherical or subspherical shape (+4.9‰) and R-Larger-sized articulated corallines (+6.3‰), the predicted values are  $-7.9 \pm 0.8\text{‰}$  and  $-9.2 \pm 0.4\text{‰}$ . For R-Filamentous uniseriate and pluriseriate with erect thallus (-2.1‰) and C-Tubular (-1.6‰), the predicted values are  $-16.3 \pm 0.5\text{‰}$  and  $-15.8 \pm 0.5\text{‰}$ , respectively. -Regarding taxon, a significant effect was observed only in 13 genera, including *Colpomenia* (+5.4‰), *Amphiroa* (+6.8‰), and *Padina* (+2.2‰) increasing the signal, and *Polysiphonia* (-3.7‰), *Gracilaria* (-0.9‰), and *Spyridia* (-1.4‰) decreasing the signal of the model constant (Table 9). In 33 species was observed a significant effect on the  $\delta^{13}\text{C}$  variability, including *C. tuberculata* +5.9‰, *C. sinuosa* +4.4‰, *H. pannosa* +4.4‰, *H. johnstonii* +4.4‰, and *Amphiroa* spp. (+4.4 to 8.2‰) increasing the model constant  $\delta^{13}\text{C} = -14.6\text{‰}$ , and *Spyridia* sp. (-2.5‰), *G. filicina* (-2.3‰), *P. mollis* (-5.2‰) and *S. pacifica* (-19.2‰) (Table 10).

### 3.7. Preliminary estimations of $\Delta^{13}\text{C}$ -macroalgal

Concurrent analysis of surface seawater for alkalinity, proportions of the chemical species of DIC ( $\text{CO}_2$ ,  $\text{HCO}_3^-$ , and  $\text{CO}_3^{2-}$ ), and  $\delta^{13}\text{C}$ -DIC evidenced that  $\delta^{13}\text{C}$ -DIC in GC seawater averages  $1.4 \pm 0.4\text{‰}$  (-1 to  $4.9\text{‰}$ ) ([Supplementary Information](#) Fig. SI-1). In our preliminary data, the  $\delta^{13}\text{C}$ -DIC seawater slightly (in  $0.5\text{‰}$ ) decreased during the rainy season in those zones influenced by river discharges along the continental coastline, ~~with non~~Non-significant differences were observed among coastal sectors.  $\delta^{13}\text{C}$ -DIC values in GC seawater are comparable to the averages  $1.4$ - $1.6\text{‰}$  reported for the surface seawaters in the Eastern North Pacific in the 1970s-2000s ~~period~~ (Quay et al., 2003; Hinger et al., 2010; Santos et al., 2011).

Based on the subtraction of  $\delta^{13}\text{C}$  macroalgae to  $\delta^{13}\text{C}$ -DIC seawater, the integrative discrimination factor against  $^{13}\text{C}$  averaged  $16.0 \pm 3.1\text{‰}$ ,  $16.8 \pm 4.3\text{‰}$ , and  $14.0 \pm 3.8\text{‰}$  for Phyla Chlorophyta, Rhodophyta, and Ochrophyta, respectively. Five groups were identified in the function of the  $\Delta^{13}\text{C}$  values, one for Chlorophyta ( $\Delta^{13}\text{C} = 16.0 \pm 3.1\text{‰}$ ), two for Rhodophyta ( $16.6 \pm 3.8\text{‰}$  and  $34.6 \pm 1\text{‰}$ ), and two for Ochrophyta ( $9.1 \pm 1.7\text{‰}$  and  $15.7 \pm 2.7\text{‰}$ ) (Fig. S2). Values of  $\Delta^{13}\text{C}$  were comparable to  $\delta^{13}\text{C}$  of the thallus of macroalgae, ~~t. Thus,~~  $\delta^{13}\text{C}$ -macroalgal reflect mainly the discrimination during carbon assimilation. Like  $\delta^{13}\text{C}$ -macroalgal, the  $\Delta^{13}\text{C}$  values were subject to considerable variation.

## 4. Discussions

### 4.1. Explaining the $\delta^{13}\text{C}$ macroalgal variability

~~In t~~This study, ~~results~~ revealed high variability in the  $\delta^{13}\text{C}$  of the large inventory of macroalgae collected along GC coastline between five years period. A linear regression analysis of the effects of life forms s revealed that the  $\delta^{13}\text{C}$ -~~macroalgal~~ variability in the -macroalgal community is mainly explained by taxonomic (genus 46%, species 57%) and morphofunctional groups. This result is

consistent with the report of Lovelock et al. (2020), who found that 66% of  $\delta^{13}\text{C}$  variability was explained by taxonomy. Even so, the variability associated with each genus is not the same and can be classified in three groups: 1) high variability (e.g., *Schizymenia*  $=\pm 19.1\text{‰}$ ), moderate variability (e.g., *Hydroclathrus*  $=\pm 7.3\text{‰}$ ; *Amphiroa*  $=\pm 6.8\text{‰}$ ) and low variability (e.g., *Gracilaria*  $=\pm 0.89$ ; *Spyridia*  $=\pm 1.46\text{‰}$ ). The observed  $\delta^{13}\text{C}$  variability in this study is comparable with those reported in the literature, compiled in Table SI-4.

Most authors studying the isotopic composition of C in macroalgae have reported the high isotopic variability, which has been attributable to the taxon-specific photosynthetic DIC acquisition properties (Raven et al., 2002a, Mercado et al., 2009, Marconi et al., 2011, Stepien, 2015, Díaz-Pulido et al., 2016; Lovelock et al., 2020). ~~In our study, we~~ observed that the intrinsic characteristics of each morpho-functional group of macroalgae (e.g., thallus structure, growth form, branching pattern, and taxonomic affinities) also influence the  $\delta^{13}\text{C}$ -macroalgal signals. The thallus thickness, a morphology propriety, influences the diffusion boundary layer on the surface of the macroalgal, where they carry out the absorption of essential ions and dissolved gases (Hurd, 2000; San-Ford and Crawford, 2000). Thus, morphology can modulate the photosynthesis rates. However, a non-biological or ecological explanation of the  $\delta^{13}\text{C}$  variability, and therefore carbon use physiology, can be given in terms of morphology.

The  $\delta^{13}\text{C}$ -macroalgal depends on the carbon source ( $\delta^{13}\text{C}$ -DIC in seawater), the isotope discrimination during carbon assimilation in the photosynthesis ( $\Delta^{13}\text{C}_p < 29\text{‰}$  in a variable degree), and the plant respiration ( $\Delta^{13}\text{C}_r$  average  $\pm 2.3\text{‰}$ ) (Carvalho et al., 2009a,b, 2010; Carvalho and Eyre, 2011, ~~Rautemberger~~ Rautenberger et al., 2015). Comparatively, the  $\Delta^{13}\text{C}_r$  value is relatively small regarding  $\Delta^{13}\text{C}_p$ . ~~thus~~ Thus,  $\delta^{13}\text{C}$ -macroalgal ~~basically~~ is an integrative value of the isotope discrimination during DIC\_seawater assimilation [ $\Delta^{13}\text{C} = (\delta^{13}\text{C}\text{-DIC seawater} - \delta^{13}\text{C}_{\text{macroalgae}})$ ]

(Carvalho et al., 2009a). Based on the  $\Delta^{13}\text{C}$  values, five groups were identified in our study: one for Chlorophyta ( $\Delta^{13}\text{C}=16.0\pm3.1\text{‰}$ ), two for Rhodophyta ( $16.6\pm3.8\text{‰}$  and  $34.6\pm1\text{‰}$ ), and two for Ochrophyta ( $9.1\pm1.7\text{‰}$  and  $15.7\pm2.7\text{‰}$ ). Values of  $\Delta^{13}\text{C}$  were comparable to  $\delta^{13}\text{C}$  of the thallus of macroalgae. Thus,  $\delta^{13}\text{C}$ -macroalgal reflect mainly the discrimination during carbon assimilation. The  $\delta^{13}\text{C}$ -macroalgal values reflect the discrimination during carbon assimilation attributable to the taxon-specific photosynthetic DIC acquisition properties.  $\Delta^{13}\text{C}$ -macroalgal variability, captured in the  $\delta^{13}\text{C}$ -macroalgal signals, is related to the thickness of the boundary layer around the thallus (Raven et al. 1982), the leakage during carbon uptake (Sharkey and Berry 1985, Maberly et al. 1992), and photosynthetic intensity (Wiencke and Fischer 1990, Kübler and Raven 1995, 1996), and respiration rates (Carvalho et al., 2010; Carvalho and Eyre, 2011, Rautenberger et al., 2015). All intrinsic properties are related to the life form.

Many species that recorded high  $\delta^{13}\text{C}$  values (and low  $\Delta^{13}\text{C}$  values) were fleshy macroalgae that are characterized to be bloom-forming macroalgae belonging to genera *Ulva*, *Gracilaria*, *Cladophora*, *Spyridia*, and *Sargassum* (Páez-Osuna et al., 2013, Valiela et al., 2018). It is not surprising, due to that species with high photosynthetic activity and high relative growth rates (Hiraoka et al., 2020) have high carbon demand that results in lower isotopic discrimination against  $^{13}\text{C}$  (Cornelisen, et al., 2007; Carvalho et al., 2010ab; Kübler and Dungeon, 2015; Rautenberger et al., 2015). Bloom-forming macroalgae (e.g., *Ulva*, *Gracilaria*, *Sargassum*) have been remarks as facultative species with the capacity to switch from C3 to C4 pathway (Valiela et al., 2018). C4 pathway reduces photorespiration, the antagonist process of RuBisCo, enhancing the DIC assimilation in 25-40% and increasing the  $\delta^{13}\text{C}$  values (Ehleringer et al., 1991; Bauwe et al., 2010; Zabaleta et al., 2012). C4 pathway has more energy investment in CCM's than in RuBisCo protein content than C3 pathway (Young et al., 2016). Also, the reports of features of C4 or C4-like pathway

features in algae have increased in the last years (Roberts et al., 2007; Doubnerová and Ryslavá, 2011; Xu et al., 2012, 2013). For example, high activity of key enzymes of C4 metabolisms, such as pyruvate orthophosphate dikinase (PPDK), phosphoenolpyruvate carboxylase (PEPC), and phosphoenolpyruvate carboxykinase (PCK), has been described in many algae species. But the establishment of a true C4 pathway in marine algae is not clear since the massive changes in gene expression patterns seem to be incomplete, and it is suggested that many marine algae have high plasticity to use a combination of CCM to overcome Ci limitations (Roberts et al., 2007; Doubnerová and Ryslavá, 2011; Xu et al., 2012, 2013). A Stepwise model of the path from C3 to C4 photosynthesis is explained by Gowik and Westhoff (2011). More research is required on this topic considering the increasing frequency, intensity, and extension of bloom-forming macroalgae events worldwide (Teichberg et al., 2010; Valiela et al., 2018) and in México (Ochoa-Izaguirre et al. 2007; Ochoa-Izaguirre and Soto-Jiménez 2015; Páez-Osuna et al., 2017).

Changes in the habitat features and environmental conditions, such as light intensity and DIC availability, influencing the growth rate and photosynthetic intensity, have a strong influence on  $\delta^{13}\text{C}$  signal (Carvalho et al., 2007, 2009; Carvalho and Eyre, 2011; Stepien, 2015; Mackey et al., 2015; Rautenberger et al., 2015). The light intensity is the external factor with more influence on the  $\Delta^{13}\text{C}$ -macroalgal due to the regulation of carbon assimilation intensity (Wefer and Killingley 1986, Cooper and DeNiro 1989, Grice et al. 1996; Carvalho et al., 2009a,b). Experimental studies found the light levels as a key factor affecting the  $\delta^{13}\text{C}$  values. f. For example, under saturating light conditions, *Ulva* switched from a carbon uptake of  $\text{HCO}_3^-$  and  $\text{CO}_2$  to increased  $\text{HCO}_3^-$  use (Rautenberger et al., 2015). Furthermore, field studies have shown that species growing in low light habitats as deep subtidal tend to have more negative  $\delta^{13}\text{C}$  values than those in higher light environments (Mercado et al., 2009; Hepburn et al., 2011; Marconi et al., 2011; Stepien 2015; Cornwall et al., 2015, Díaz-

Pulido et al., 2016). In this study, intertidal specimens recorded lesser negative values than subtidal in most macroalgae genus. However, our study did not record the vertical effect in the  $\delta^{13}\text{C}$  signal related to the light limitation ~~was not recorded in our study~~ because only shallow habitats (non-light limited), were ~~considered~~studied.

$\delta^{13}\text{C}$ -DICseawater is reasonably uniform in surface seawater (-4.8 to 3.6‰, median 1.5‰), with  $\delta^{13}\text{C}$  values for  $\text{CO}_2$ ,  $\text{HCO}_3^-$ , and  $\text{CO}_3^{2-}$  nearly -10, -0.5 and 2‰, respectively (Mook et al., 1974; Kroopnick, 1985). Exceptions can be expected where variations in the salinity, alkalinity, and proportions of the chemical species of DIC ( $\text{CO}_2$ ,  $\text{HCO}_3^-$  or  $\text{CO}_3^{2-}$ ) occur (e.g., in coastal environments influenced by river and groundwater discharges) (Mook et al., 1974; Chanton and Lewis 1999; Hinger et al., 2010; Carvalho et al., 2015). Regarding DIC sources for macroalgae in the GC surface seawater, the availability, chemical proportions, and  $\delta^{13}\text{C}$ -DIC, were also relatively constant and uniform. Thus, the influence of the  $\delta^{13}\text{C}$ -DIC variations ~~to~~on the  $\delta^{13}\text{C}$ -macroalgal variability is negligible in the GC.

The effect of other environmental factors, such as salinity and pH, on  $\delta^{13}\text{C}$ -macroalgal signals, were evaluated. Regarding salinity, the influence of freshwater discharge by rivers and groundwater decreases the  $\delta^{13}\text{C}$  signal, which could be explained by the ~~effect of the~~ reduction in the salinity regimen that follows a decrease in  $\delta^{13}\text{C}$ -DIC in water (Hinger et al., 2010; Santos et al., 2011). In our study, a non-significant correlation between  $\delta^{13}\text{C}$ -macroalgal and salinity was observed.

Based on pH, differences in  $\delta^{13}\text{C}$  were found only for a few genera (e.g., *Amphiroa*, *Colpomenia*, *Ulva*, *Spyridia*), with a trend to increase in the  $\delta^{13}\text{C}$  values with pH increase, such as was reported by Maberly et al. (1992) and Raven et al. (2002b). Similar results were reported for Cornwall et al. (2017) in the field study, with the differential response of the  $\delta^{13}\text{C}$  signals to pH among 19 species,

in which only four species were sensitive to pH changes. ~~Based on the complete dataset, a~~ very weak but significant positive linear regression was observed between  $\delta^{13}\text{C}$  and pH. Also, a trend to decrease in the  $\delta^{13}\text{C}$  was recorded in the following order: alkaline > typical > acidic. According to Stepien (2015), the result of meta-analyses between pH ~~values~~ drift experiments and  $\delta^{13}\text{C}$  thresholds was positive only for Rhodophyta and Ochrophyte, but not for Chlorophyta. About 86% of the Stepien metadata met the theoretical CCM assignment based on both parameters, exceptions for species with  $\delta^{13}\text{C} < -30\text{‰}$  that have been capable of raising pH > 9. A strong association between pH compensation point and  $\delta^{13}\text{C}$  was reported by Iñiguez et al. (2009) in three taxa of polar macroalgae. Environmental conditions may influence the  $\delta^{13}\text{C}$ -macroalgal values but not change the carbon use physiology in the macroalgae, which is most likely inherently species-specific.

#### 4.2. Using $\delta^{13}\text{C}$ -macroalgal to indicate the presence of an active CCM

In our study, the  $\delta^{13}\text{C}$  signals from the thallus of macroalgae ~~was~~ were used to evidence the presence of an active CCM ~~infer carbon strategies~~. This tool was first used in macroalgal shallows communities of the Gulf of California. Most macroalgae species displayed  $\delta^{13}\text{C}$  values that exhibit an active CCM's. Then, macroalgae were classified into four strategies for DIC uptake, in agreement with the Maberly et al. (1992) and Raven et al. (2002) thresholds: 1) CCM-only by active uptake  $\text{HCO}_3^-$  ( $\delta^{13}\text{C} > -10\text{‰}$ ), 2) CCM active uptake  $\text{HCO}_3^-$  and/or diffusive uptake  $\text{CO}_2$  ( $\delta^{13}\text{C} < -11$  to  $-30\text{‰}$ ), 3) Non-CCM,  $\text{CO}_2$  by diffusion only ( $\delta^{13}\text{C} < -30\text{‰}$ ), 4) Calcifying with different carbon-use strategies related to different modes of calcification. About of 84% of the total analyzed specimens showed the facultative uptake of  $\text{HCO}_3^-$  and  $\text{CO}_2$ , the most common strategy identified in macroalgal shallow communities (Hepburn et al., 2011; Cornwall et al., 2015; Stepien 2015; Díaz-Pulido et al., 2016). Based on the carbon uptake strategies, the most abundant macroalgae were those able to use both  $\text{HCO}_3^-$  and/or  $\text{CO}_2$  by means of active uptake plus passive diffusion (strategy 2:  ~~$-10 < \delta^{13}\text{C} < 30\text{‰}$~~ ).



633 Macroalgae collected in GC also involved those that are only  $\text{HCO}_3^-$  users (strategy 1:  $\delta^{13}\text{C} > -10\text{‰}$ )  
 634 and those relying on diffusive  $\text{CO}_2$  uptake (strategy 3:  $\delta^{13}\text{C} < -30\text{‰}$ ). Photosynthesis that relies on  
 635  $\text{CO}_2$  uptake (lack of CMM), the most primitive mechanism (Cerling et al., 1993), has fewer energy  
 636 costs than  $\text{HCO}_3^-$  uptake, which requires complex machinery with a high operational cost (Giordano  
 637 et al., 2005; Hopkinson et al., 2011; Hopkinson, 2014; Raven and Beardall, 2016). The energy for  
 638 macroalgae to uptake  $\text{HCO}_3^-$ , cross the plasma membrane, and convert to  $\text{CO}_2$  for photosynthesis, is  
 639 obtained through irradiance (Cornelisen et al., 2007). Based on our sampling effort, focused on  
 640 intertidal and shallow subtidal habitats featured by high-light intensities, we expected high  
 641 proportions of species ~~and specimens~~ with the carbon uptake strategy that uses only  $\text{HCO}_3^-$ . Results  
 642 evidenced that strategy 1 was recorded in specimens belonging to 58 species of 170 total species.  
 643 The higher proportions of CCM species ( $\text{HCO}_3^-$  users), with high-energetic requirements, ~~is~~ are  
 644 explained by those elevated irradiances (Hepburn et al. 2011; Cornwall et al. 2015). Ochrophyta  
 645 showed the highest proportion of species ~~and specimens~~ that depend only on  $\text{HCO}_3^-$  uptake on both  
 646 coastlines in the southern region of GC (P1, C1). ~~These differences can be partially explained by the~~  
 647 Low solubility of  $\text{CO}_2$  ~~is related to~~ due to relatively high temperatures in subtropical waters (Zeebe  
 648 and Wolf-Gladrow, 2007) that impulse the development of CCM (Raven et al., 2002b) and by the  
 649 high affinity to DIC by Ochrophyta, such as has been described before by Diaz-Pulido et al, (2016).  
 650 Only three non-calcifying species (*Schizymenia pacifica*, *Halymenia* sp., *Gigartina* sp.) belonging  
 651 to Rhodophyta were  $\text{CO}_2$  exclusive users ( $\delta^{13}\text{C} = -33.2 \pm 1\text{‰}$ ). Based on measurements of pH drift,  
 652 Murru and Sandgreen (2004), reported ~~to~~ *Schizymenia pacifica* and two species of *Halymenia* (e.g.,  
 653 *H. schizymenioides* and *H. gardner*) as restricted  $\text{CO}_2$  users. Measurements of  $\delta^{13}\text{C}$  in *Halymenia*  
 654 *dilatata* confirmed the  $\text{CO}_2$ -restricted photosynthesis in specimens collected offshore in deep reefs  
 655 of the Great Barrier reef (Díaz-Pulido et al., 2016). Red macroalgae that lack CCM, tend to inhabit

low-light habitats like subtidal or low intertidal and are abundant in cold waters (Kübler et al., 1999; Raven et al., 2002a; Cornwall et al., 2015). According to these authors, approximately 35% of the total red algae tested ~~on a globally scale~~ are strictly CO<sub>2</sub> dependents. The percentage of macroalgae species representative of Arctic and Antarctic ecosystems that lack CCM is 42-60% (Raven et al., 2002b; Iñiguez et al., 2019), 50% for temperate waters of New Zealand (Hepburn et al., 2011), and up to 90% found for a single site of Tasmania, Australia (Cornwall et al., 2015). ~~In our study,~~ sampled 91 red macroalgae species ~~were sampled~~ (of 453 red macroalgae species reported in the GC, Pedroche and Senties, 2003), of which <3% were CO<sub>2</sub> dependents. This low percentage could be related to the fact that deep habitats (>2 m depth low tide) were not explored in our surveys.

~~In our study,~~ Few calcifying macroalgae species using HCO<sub>3</sub><sup>-</sup> and ~~diffusive~~ CO<sub>2</sub> (strategy 4) were also collected, including the genera *Amphiroa* (-7.8±3.7‰) and *Jania* (-9.4±0.7‰), both Rhodophyta with articulated-form. *Padina*, a genus with less capacity to precipitate CaCO<sub>3</sub> (Ilus et al., 2017), displayed relatively high δ<sup>13</sup>C values (-12.5±2.4‰), suggesting the presence of CCM using HCO<sub>3</sub><sup>-</sup> ~~exclusively~~. Some species of *Padina* can use HCO<sub>3</sub><sup>-</sup>, but their efficiency may differ from species to species (Raven et al., 2002a; Enríquez and Rodríguez-Román, 2006). Three genera are ~~very common~~ widespread in the GC. Stepien (2015) reported a global mean of -14.8±1.0‰ for calcifying species compared to -20.1±0.3‰ for non-calcifying species. Calcifying species have a different carbon uptake strategy influenced by the calcifying process that results in elevated δ<sup>13</sup>C signals (Diaz-Pulido et al., 2016). One possibility for high δ<sup>13</sup>C values for calcifying species are related to the excess of H<sup>+</sup> released as residuals products of the calcifying process, ~~a.~~ Also, the acidified boundary layers benefit the HCO<sub>3</sub><sup>-</sup> uptake (McConnaughey and Whelan 1997, Courneau et al., 2012). Another possibility to explain high δ<sup>13</sup>C values can also be related to the highly efficient light properties ~~that are~~ enhanced by the carbonate skeleton, resulting in an optimization of

679 photosynthetic activity (Vásquez-Elizondo et al., 2017). Hofmann and Heesch (2018) reported high  
680  $\delta^{13}\text{C}$  values in eight rhodoliths species (calcifying species) for the organic matter thallus and for  
681 thallus, including  $\text{CaCO}_3$  structure collected in deep habitats (25-40 m) where light availability is  
682 very low/limited. Because the ocean acidification is in progress, negative impacts are expected on  
683 calcifying organisms, more attention as ecological sentinels is warranted in the GC.

684 Measurements of  $\delta^{13}\text{C}$  signals evidence the presence or absence of CCMs in macroalgae and are  
685 indicative of carbon use physiology (Giordano et al., 2005), however, the isotopic signature may be  
686 inconclusive in the determination of the efficient use of one or more DIC species ( $\text{CO}_2$  and/or  $\text{HCO}_3^-$   
687 ) (Roleda and Hurd, 2012). The preferential DIC uptake of macroalgae is assessed by pH drift  
688 experiments (Hepburn et al., 2011; Roleda and Hurd, 2012; Fernandez et al. 2014, 2015; Narvarte et  
689 al., 2020) and it can be determined by simultaneously measuring the  $\text{CO}_2$  uptake and  $\text{O}_2$  production  
690 rates using membrane-inlet mass spectroscopy (MIMS) (Douchi et al., 2019; Burlacot et al., 2020).  
691 Macroalgae that are unable to raise the seawater  $\text{pH}>9.0$  are primarily  $\text{CO}_2$ -users, while those that  
692 can raise the seawater  $\text{pH}>9.0$  (absence of  $\text{CO}_2$ ) are  $\text{HCO}_3^-$ -users (Roleda; Hurd, 2012). Those  
693 differences in the carbon uptake strategies can be easily deduced by pH drift experiments, which  
694 were not done in our study but reported in the literature (Supplementary Information Table SI-4).  
695 Also, the change in  $\delta^{13}\text{C}$  signature within the range specific to a carbon use strategy (e.g., mix  
696  $\text{HCO}_3^-/\text{CO}_2$ -user) can be complemented by simultaneous measurements of  $\text{O}_2$  and  $\text{CO}_2$  produced  
697 and consumed, respectively, during the photosynthetic using MIMS. For example, photosynthetic  
698  $\text{O}_2$  production in a certain macroalgae species with an active CCM preferring (e.g.,  $\text{CO}_2$ ) is about ten  
699 times higher than no active CCM (Burlacot et al., 2020).

700 Based on the  $\delta^{13}\text{C}$  values, it is possible to assume that at least one basal CCM is active, however, it  
701 is impossible to discern what type of CCM is expressed in the organisms (e.g., direct  $\text{HCO}_3^-$  uptake

by the anion-exchange protein AE; Drechsler and Beer 1991; Drechsler et al. 1993) or types of mitochondrial carbonic anhydrase (e.g., internal and external) that enhance the fixation of  $\text{Ci}$  by recycling mitochondrial  $\text{CO}_2$  (Bowes, 1969; Zabaleta et al., 2012; Jensen et al., 2020). Also, the co-existence of different CCMs has been described for the same species (Axelsson et al., 1999, Xu et al., 2012), even that different CCM's can operate simultaneously, generating different  $\text{Ci}$  contributions to RuBisCo internal pool (~~Rautemberger~~-Rautemberger et al., 2015). The variety of CCMs and their combinations could contribute to the high  $\delta^{13}\text{C}$  variability for the same species. In our field study, it is impossible to explain the variations of  $\delta^{13}\text{C}$  or  $\Delta^{13}\text{C}$ -macroalgal relative to CCM or CA activity types. Controlled experiments, ~~as-like~~ those conducted by Carvalho and collaborators (e.g., Carvalho et al. 2009a,b, 2010), are required to obtain this knowledge.

#### 4.3. Variability of $\delta^{13}\text{C}$ macroalgal between the GC bioregions

Changes in the  $\delta^{13}\text{C}$  signal with latitude, mainly related to the light and temperature, have been reported in the literature (Mercado et al., 2009; Marconi et al., 2011; Stepien, 2015; Hofmann and Heesch, 2018; Lovelock et al., 2020). For example, a negative correlation between latitude and  $\delta^{13}\text{C}$ -macroalgal was described by Stepien (2015), concluding that the  $\delta^{13}\text{C}$  signal increased by 0.09‰ for each latitude degree from the Equator. Hofmann and Heesch (2018) recently showed a strong decrease in latitudinal effect ( $R^2 = 0.43$   $\delta^{13}\text{C}_{\text{total}}$  and 0.13, for  $\delta^{13}\text{C}_{\text{organic-tissue}}$ ,  $p=0.001$ ) for rhodolite of the northern hemisphere and macroalgae from coral reefs in Australia. In both cases, the latitude range is higher than we tested ( $30^\circ$  to  $80^\circ$  and from  $10^\circ$  to  $45^\circ$ , respectively). These differences on a big scale tend to be associated with a temperature effect (Stepien, 2015) and their effect on  $\text{CO}_2$  solubility in seawater (Zeebe and Wolf-Gladrow, 2007). However, in our study, any geographical pattern in the  $\delta^{13}\text{C}$  macroalgal was observed. Our linear regression analyzes for latitudes showed a low but significant correlation for the dataset classified by morphofunctional groups and genus,

negative in the cases of Rhodophyta and Ochrophyta groups, and positive for Chlorophyta.

Light is not limited along the GC latitudes, ~~most~~ Most of the shallow habitats occupied by macroalgal communities in the GC were high-light environments. In agreement to literature, the surface seawater temperature across the GC vary in only 1°C annual mean (Escalante et al., 2013, Robles-Tamayo, 2018). However, larger temperature variations of 5-10°C were recorded in the coastal waters across the GC bioregions in both climatic seasons. The combined effect of the coastline sector, habitats feature, or environmental condition for Morphofunctional group or Genus explained 60-62 and 71-72% of the  $\delta^{13}\text{C}$  variability, respectively. Our analysis of variability for the best-represented morphological groups (e.g., R-Filamentous uniseriate and pluriseriate with erect thallus and C-Tubular) and genus (e.g., *Colpomenia*, *Padina*, *Polysiphonia* and *Gracilaria*) revealed that certain life forms are better monitors explaining the variability of  $\delta^{13}\text{C}$ -macroalgal (and  $\Delta^{13}\text{C}$  values) than others. The  $\delta^{13}\text{C}$  variability in morphological groups refers to change within a specific carbon use strategy, but not change in the carbon use physiology that is inherently species-specific. The biological or ecological relevance of the  $\delta^{13}\text{C}$  variability in function of the morphology, in terms of the efficiency in the use of DIC and the isotope discrimination during carbon assimilation and respiration, must be investigated in species of same genus morphologically different or between same morphological structures belonging to a different taxon.

The proportion of specimens with different carbon uptake strategies also showed regional variations. For example, the facultative uptake of  $\text{HCO}_3^-$  and  $\text{CO}_2$  was dominant in the macroalgal shallow communities in the GC (60 to 90% of specimens), with an exception in the P1 region for Ochrophyta where the specialized strategy of only  $\text{HCO}_3^-$  use dominated (68%), and high proportion were observed in C1 with 37%. While the strategy based on only use of  $\text{CO}_2$  was observed in the peninsular coast in P2 and P3 for Rhodophyta with 2-3.3%. Finally, the coastal sector C2 showed

more negative  $\delta^{13}\text{C}$  values in macroalgae specimens of the same genus compared to the peninsular coastline (P1-P3). Small but detectable changes were observed in the Phyla distribution based on environmental conditions. For example, Ochrophyta showed the highest proportion (35%) in colder temperature, in pH-Alkaline (31%), and at typical salinity regimen (27%), while Chlorophyta enhanced to 30% in acid pH and Rhodophyta recorded 21% at normal seawater. The opposite strategy (only use of dissolved  $\text{CO}_2$ ) ~~was that was~~ observed only in Rhodophyta, ~~t.~~ The highest percentage was observed in the estuarine salinity regimen (10%). Again, more research is required to obtain ~~useful~~ valuable information on the physiological and environmental status of macroalgae.

## 5. Conclusions

In conclusion, we observed high  $\delta^{13}\text{C}$ -macroalgal variability in macroalgae communities in the Gulf of California, such as ~~has been~~ reported in other worldwide marine ecosystems. ~~Life~~ The life form is the principal cause of  $\delta^{13}\text{C}$ -macroalgal variability, which explains up to 57% of the variability, ~~respectively~~. Changes in habitat characteristics and environmental conditions also influence the  $\delta^{13}\text{C}$ -macroalgal variability within a specific carbon use strategy. Considering the combined effect of the life form, coastline sector, and environmental conditions, the full model explains up to 72% (genus) of the variability. The effect of the coastal sector, pH ranges, and emersion level were significant, while for salinity and temperature, negligible.

Most macroalgae inhabiting in GC displayed the presence of  $\text{CO}_2$  concentrating mechanisms to uptake  $\text{HCO}_3^-$  for photosynthesis, 84% of the total analyzed specimens were able to use both  $\text{HCO}_3^-$  and/or  $\text{CO}_2$  employing active uptake plus passive diffusion (strategy 2:  $-10 < \delta^{13}\text{C} < -30\text{‰}$ ). Specimens belonging to 58 species of 170 total species showed carbon uptake strategy 1 that use only  $\text{HCO}_3^-$ .

A higher proportion of CCM species ( $\text{HCO}_3^-$  users) was expected because we focused on intertidal and shallow subtidal habitats featured by high-light intensities. Only three non-calcifying species (*Schizymenia pacifica*, *Halymenia* sp., *Gigartina* sp.) belonging to Rhodophyta (3%) were  $\text{CO}_2$  exclusive users (strategy 3:  $\delta^{13}\text{C} < -30\text{‰}$ ). The low percentage of  $\text{CO}_2$  dependents versus 40-90% reported for temperate regions could be related to the shallow habitat sampled in our surveys (<2 m depth low tide). The calcifying macroalgae genera *Amphiroa* and *Jania* using  $\text{HCO}_3^-$  and diffusive  $\text{CO}_2$  influenced by the calcification process (strategy 4) were present in the macroalgal communities along the GC and high  $\delta^{13}\text{C}$  values (~~similar to strategy 1~~). Because the ongoing ocean acidification, these calcifying organisms constitute excellent ecological sentinels in the GC.

Finally, diverse authors have reported significant correlations between  $\delta^{13}\text{C}$  signal and latitude, mainly related to the light and temperature. However, in ~~the our study's~~ latitude range ( $21^\circ$ - $31^\circ\text{N}$ ) ~~in our study~~, the linear regression analyzes showed a low correlation for the  $\delta^{13}\text{C}$ -macroalgal dataset classified by morphofunctional groups and genus, being negative for Rhodophyta and Ochrophyta and positive for Chlorophyta. Non-clear  $\delta^{13}\text{C}$ -macroalgal patterns occur along the GC latitudes. However, detectable changes were observed in the  $\delta^{13}\text{C}$ -macroalgal and the proportion of specimens with different carbon uptake strategies among coastal sectors. For example, the facultative uptake of  $\text{HCO}_3^-$  and  $\text{CO}_2$  was dominant in the macroalgal shallow communities in the GC (60 to 90% of specimens), but in the coastal sector P1 was the specialized strategy of only  $\text{HCO}_3^-$  use the dominant strategy (68%), and significant at C1 (37%).

Our research is the first approximation to understand the  $\delta^{13}\text{C}$ -macroalgal variability in one of the most diverse marine ecosystems in the world, the Gulf of California. We did not pretend to resolve the intricate processes controlling the variations of  $\delta^{13}\text{C}$  or  $\Delta^{13}\text{C}$ -macroalgal during carbon assimilation and respiration and determine the isolated influence of each environmental factor.

793 Despite the large dataset and corresponding statistical analyses, our study faces limitations due to  
794 research design and because no research on  $\delta^{13}\text{C}$ -macroalgal analysis was developed previously in  
795 the GC. The primary deficiency is the lack of pH drift experiments to discriminate  $\delta^{13}\text{C}$  signal  
796 variations to the carbon uptake strategies to determine preferential DIC uptake of macroalgae ( $\text{CO}_2$   
797 or  $\text{HCO}_3^-$ ). The second limitation concerns the lack of controlled experiments to discern what type  
798 of CCM is expressed in macroalgae (e.g., direct  $\text{HCO}_3^-$  uptake by the anion-exchange protein AE,  
799 types of mitochondrial AC, or the co-existence of different CCMs). Also, more research is required  
800 to assess the biological or ecological relevance of the  $\delta^{13}\text{C}$  variability in function of the morphology  
801 (e.g., DIC uptake efficiency and isotope discrimination during carbon assimilation and respiration).  
802 Future studies assessing the ability of macroalgae to use  $\text{CO}_2$  and/or  $\text{HCO}_3^-$  can be assessed by pH  
803 drift experiments and/or MIMS in the cosmopolites' species and within of genus with differences in  
804 the  $\delta^{13}\text{C}$  values between species (e.g., *Ulva* and *Sargassum*). ~~Controlled~~ Finally, controlled  
805 experiments in laboratory and mesocosm type ~~in combination~~ combined with field studies are  
806 required to elucidate what type of CCM is expressed in macroalgae. ~~Even so,~~ the  $\delta^{13}\text{C}$ -macroalgal  
807 was a good indicator to infer the presence or absence of CCM's and identify the macroalgae lineages  
808 that could be in a competitive advantage based on their carbon uptake strategy and identify their  
809 geographical distribution along with GC.

810 Under the current conditions of climate change and their effects as ocean acidification ~~in~~ progresses  
811 and the bloom-forming macroalgae events ~~that~~ increases in México and worldwide, the analysis of  
812  $\delta^{13}\text{C}$ -macroalgal constitutes an excellent tool to help to predict the prevalence and shift of species in  
813 a-macroalgal communities' -focused on carbon metabolism. However, to obtain the maximum  
814 benefit from isotopic tools in the carbon-use strategies study, diverse and species-specific, it is  
815 necessary to use them in combination with other techniques referred to herein.



## 6. Data Availability Statement

Data set are each permanently deposited Soto-Jimenez, [Martin](#) F; Velázquez-Ochoa, Roberto; Ochoa Izaguirre, Maria Julia. Earth and Space Science Open Archive ESSOAr; Washington, Nov 25, 2020. DOI:10.1002/essoar.10504972.1

<https://search.proquest.com/openview/2060de58b217ca47495469b53ae2f347/1?pq-origsite=gscholar&cbl=4882998>

## 7. Author contribution

Velázquez-Ochoa R. participate in the collection, processing, and analysis of the samples as a part of his master's degree thesis. Ochoa-Izaguirre J. also participate in sample collections and identified macroalgae specimens. Soto-Jiménez M.F. coordinated the research, was the [graduate](#) thesis director, and prepared the manuscript with contributions from all co-authors.

## 8. Competing interests

The authors declare that they have no conflict of interest.

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## Figure captions

Fig. 1. Sites collection along the continental (C1-C3) and peninsula (P1-P3) Gulf of California coastlines (A), range of environmental factors supporting or limiting the life processes for the macroalgal communities within a habitat (B), and inserted Table with the features and environmental conditions in the diverse habitats in the GC bioregions that delimits the macroalgal community's zonation.

Fig. 2. Variability of  $\delta^{13}\text{C}$  values for specimens of different macroalgae genera collected along GC coastlines classified by taxon: (A) Chlorophyta, (B) Ochrophyta and (C) Rhodophyta. Shaded background represents the cutoff limits for using  $\text{CO}_2$  Only users and  $\text{HCO}_3^-$  only users, respectively, according to Raven et al., (2002).

Fig. 3. Variability of  $\delta^{13}\text{C}$  values for the genus collected along coastline of the Gulf of California according to their taxon: (A) Chlorophyta, (B) Ochrophyta and (C) Rhodophyta. Genus with  $n=1$  is not shown, and genus  $n=2$  was not considered to the statistical comparison. Different letters indicate significant differences ( $P<0.05$ ):  $a>b>c>d>e$ . Shaded background represent the cutoff limits for using  $\text{CO}_2$  Only users and  $\text{HCO}_3^-$  only users, respectively, according to Raven et al., (2002). For Chlorophyta: Bry= *Bryopsis*, Cau= *Caulerpa*, Cha= *Chaetomorpha*, Cla= *Cladophora*, Cod= *Codium*, Phy= *Phyllocladon*, Str= *Struveopsis*, Ulv= *Ulva*. Phaeophyta: Col= *Colpomenia*, Dic= *Dictyota*, Ect= *Ectocarpus*, End= *Endarachne*, Hyd= *Hydroclathratrus*, Pad= *Padina*, Ros= *Rosenvingeia*, Sar= *Sargassum*, Spa= *Spatoglossum*, Zon= *zonaria*. Rhodophyta: Aca= *Acantophora*, anf= *Anfeliopsis*, Amp= *Amphiroa*, Cen= *Centroceras*, Cer<sup>1</sup>= *Ceramium*, Cer<sup>2</sup>= *Ceratodictyon*, Cho<sup>1</sup>= *Chondracanthus*, Cho<sup>2</sup>= *Chondria*, Das= *Dasya*, Dig= *Digenia*, Euc= *Euchema*, Gel= *Gelidium*, Gig= *Gigartina*, Gra<sup>1</sup>= *Gracilaria*, Gra<sup>2</sup>= *Grateloupia*, Gra<sup>3</sup>=

1278 *Gracilariopsis*, Gym= *Gymnogongrus*, Hal= *Halymenia*, Hyp= *Hypnea*, Jan= *Jania*, Lau=  
 1279 *Laurencia*, Lom= *Lomentaria*, Neo= *Neosiphonia*, Pol= *Polysiphonia*, Pri= *Prionitis*, Rho<sup>1</sup>=  
 1280 *Rhodoglossum*, Rho<sup>2</sup>= *Rhodymenia*, Sch= *Schizymenia*, Spy= *Spyridia*, Tac= *Tacanoosca*. Purple  
 1281 boxplots represent calcifying species group.

1282 Fig. 4. Variability of  $\delta^{13}\text{C}$  values for morphofunctional groups by taxa along coastline of the Gulf  
 1283 of California.

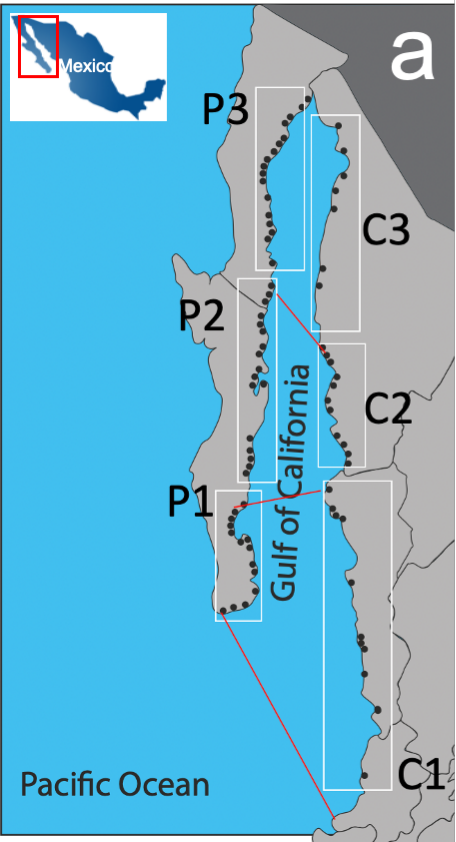
1284 Fig. 5 Proportion of species using different DIC sources according to their carbon uptake  
 1285 strategies:  $\text{HCO}_3^-$  only users ( $\text{CO}_2$  concentrating mechanism active), Users of both sources ( $\text{HCO}_3^-$   
 1286 &  $\text{CO}_2$ ) and  $\text{CO}_2$  only users (non- $\text{CO}_2$  concentrating mechanism active) in function of coast along  
 1287 GC.

1288 Fig. 6. Variability of  $\delta^{13}\text{C}$  values in macroalgae specimens for the most representative genera in  
 1289 function of habitat features (emersion level). Green circles represent genus of Chlorophyta, Brown  
 1290 circles represent genus of Ochrophyta; red circles represent genus Rhodophyta and purple circles  
 1291 represent genus with calcifying capacity.

1292 Fig. 7. Variability of  $\delta^{13}\text{C}$  values in macroalgae specimens for the most representative genus in  
 1293 function of temperature (a) and pH (b) ranges in samples collected along Gulf of California  
 1294 coastline.

1295 Fig. 8. Proportion of species using different DIC sources according to their carbon assimilation  
 1296 strategies:  $\text{HCO}_3^-$  only users ( $\text{CO}_2$  concentrating mechanism active), Users of both sources ( $\text{HCO}_3^-$   
 1297 &  $\text{CO}_2$ ) and  $\text{CO}_2$  only users (non- $\text{CO}_2$  concentrating mechanism active) in function of : (A) pH  
 1298 ranges, (B) temperature ranges and (C) salinity ranges.

1299 Fig. 9. Trends in the  $\delta^{13}\text{C}$ -macroalgal in specimens collected along continental (C1-C3) and  
1300 peninsula (P1-P3) Gulf of California coastline in function of latitudinal gradient.



Habitats features and environmental conditions in sampling sites

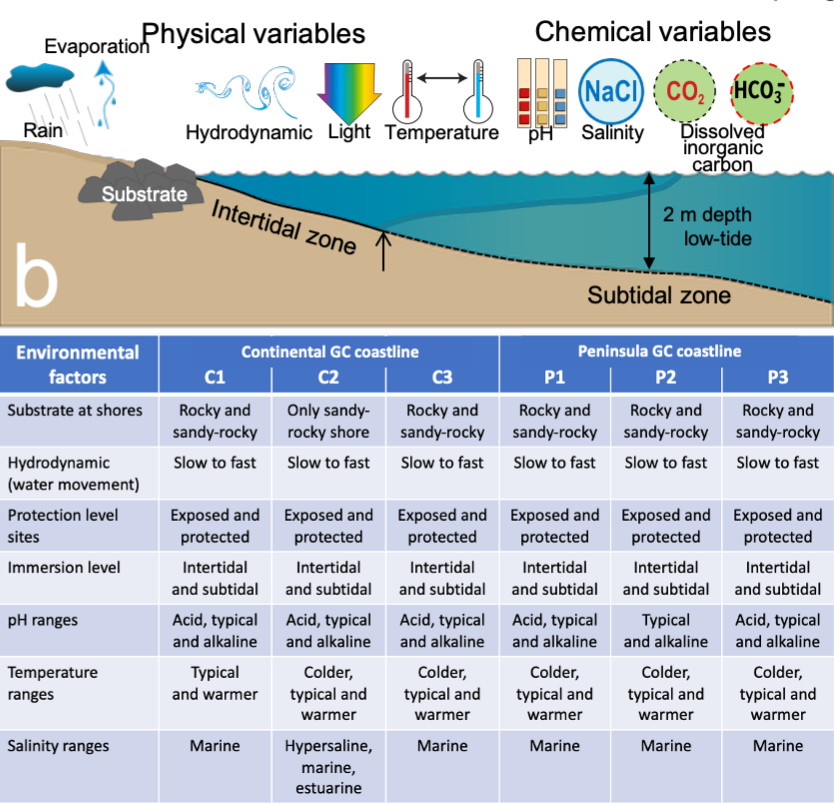


Fig. 1

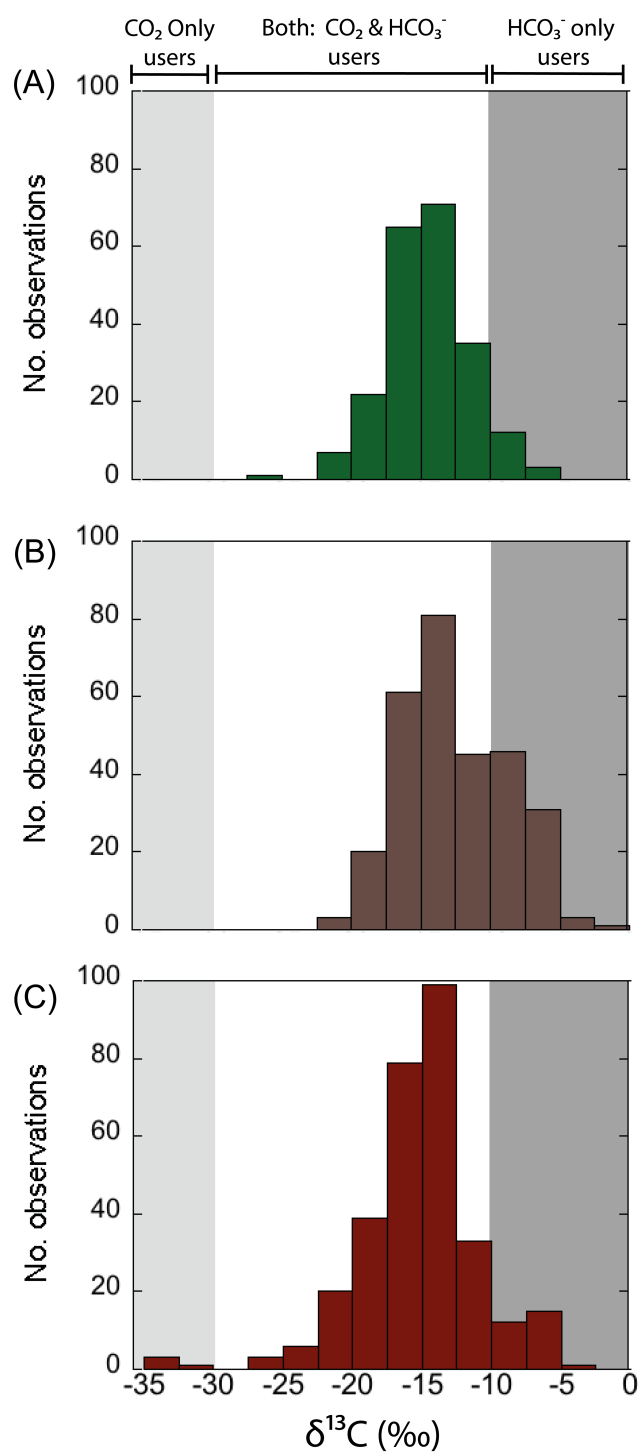
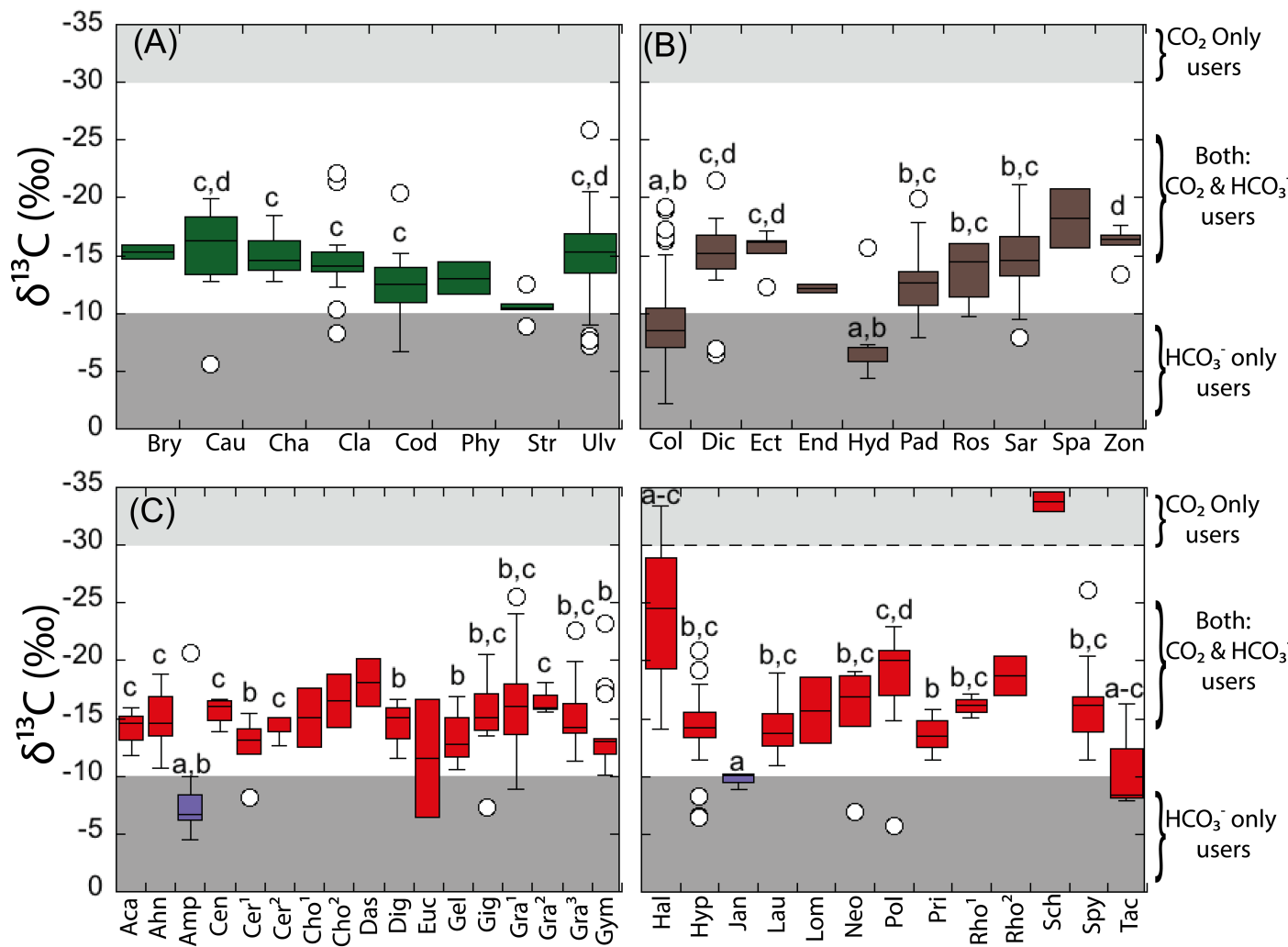
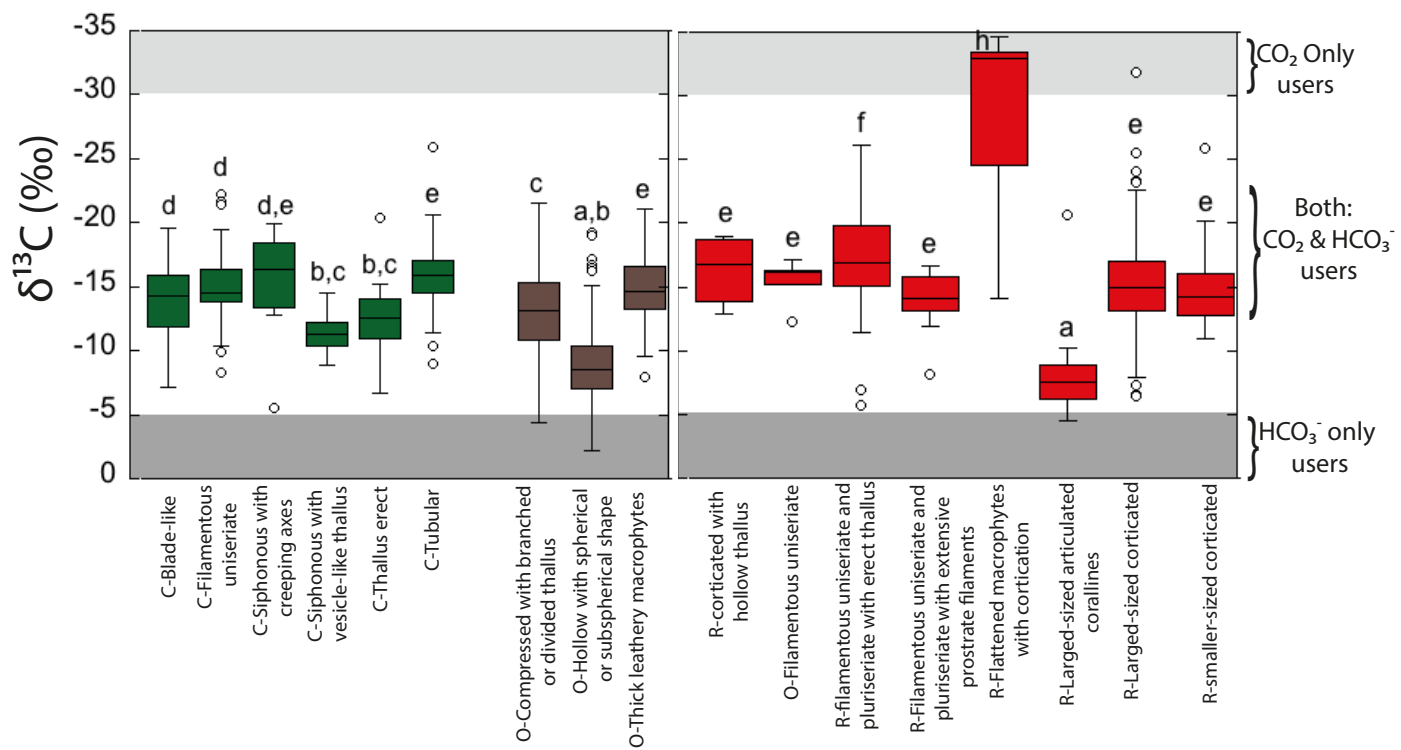


Fig 2

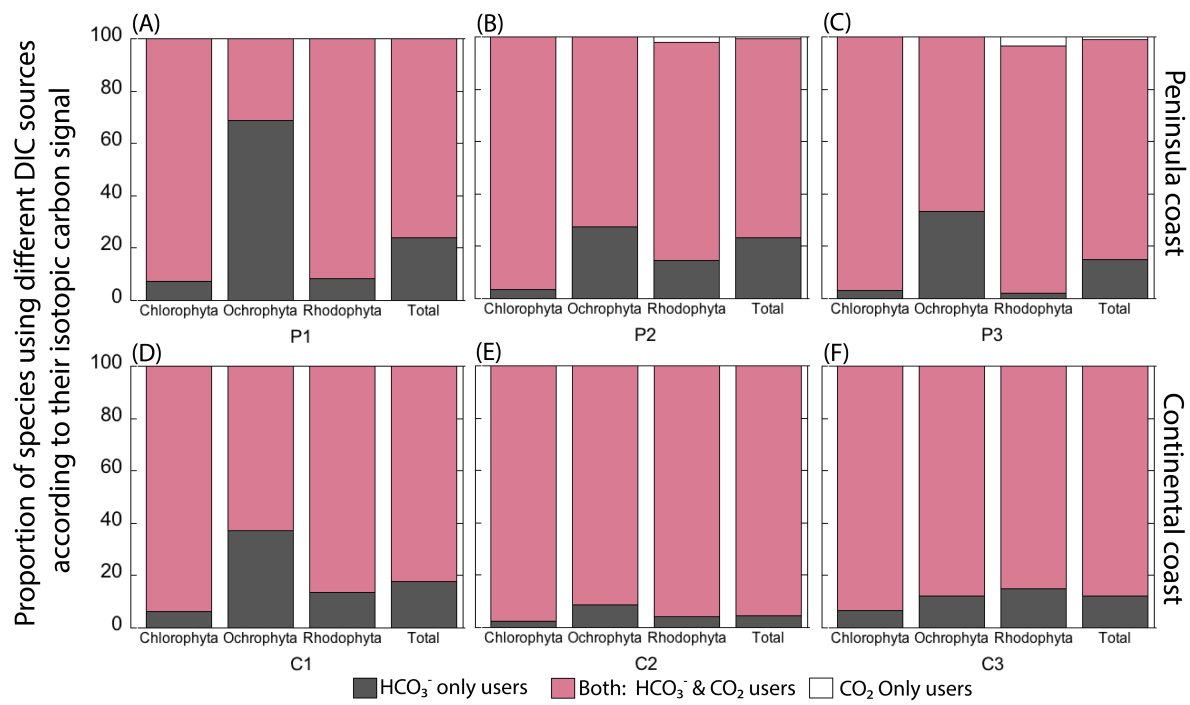


**Fig 3**



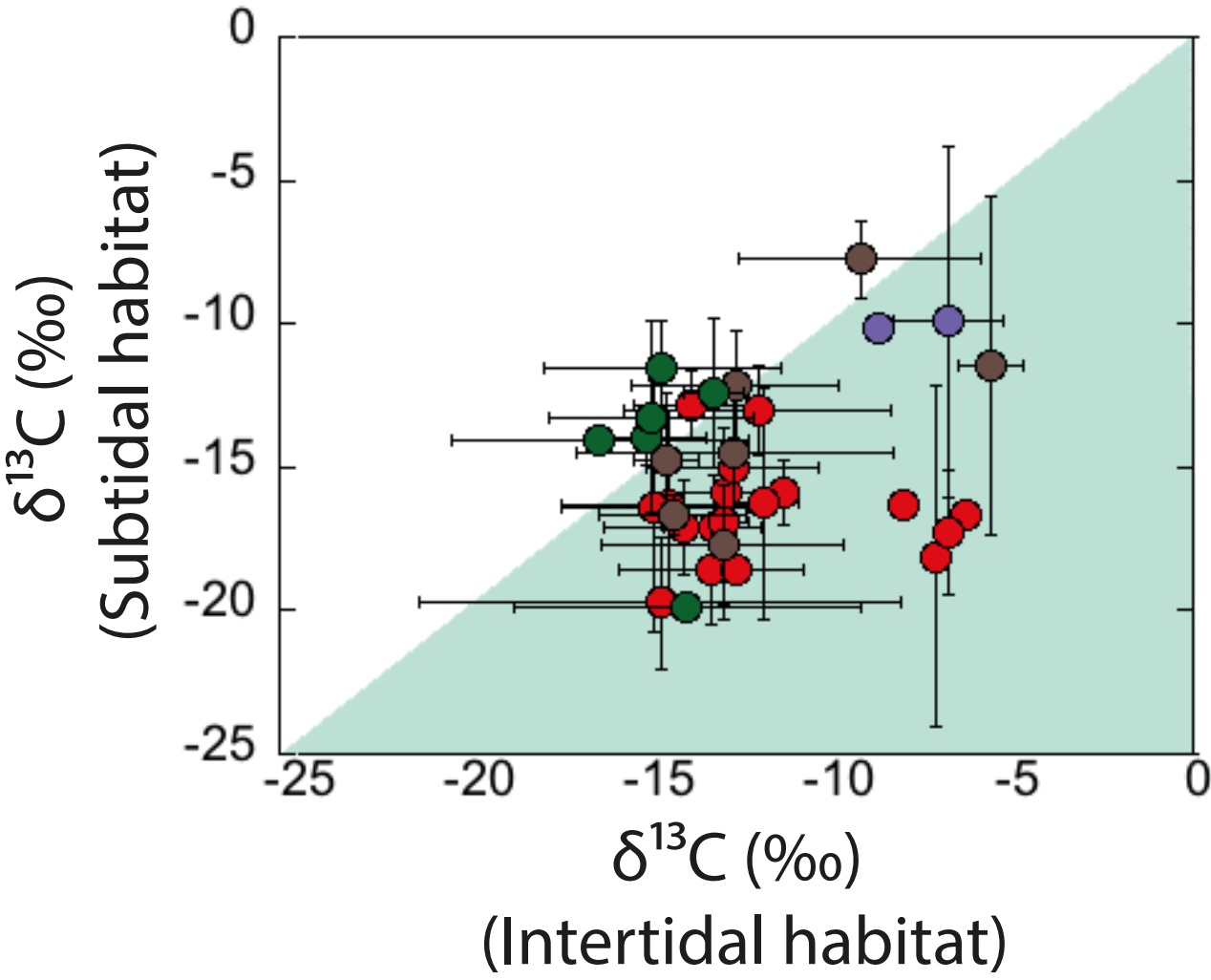
**Fig 4**





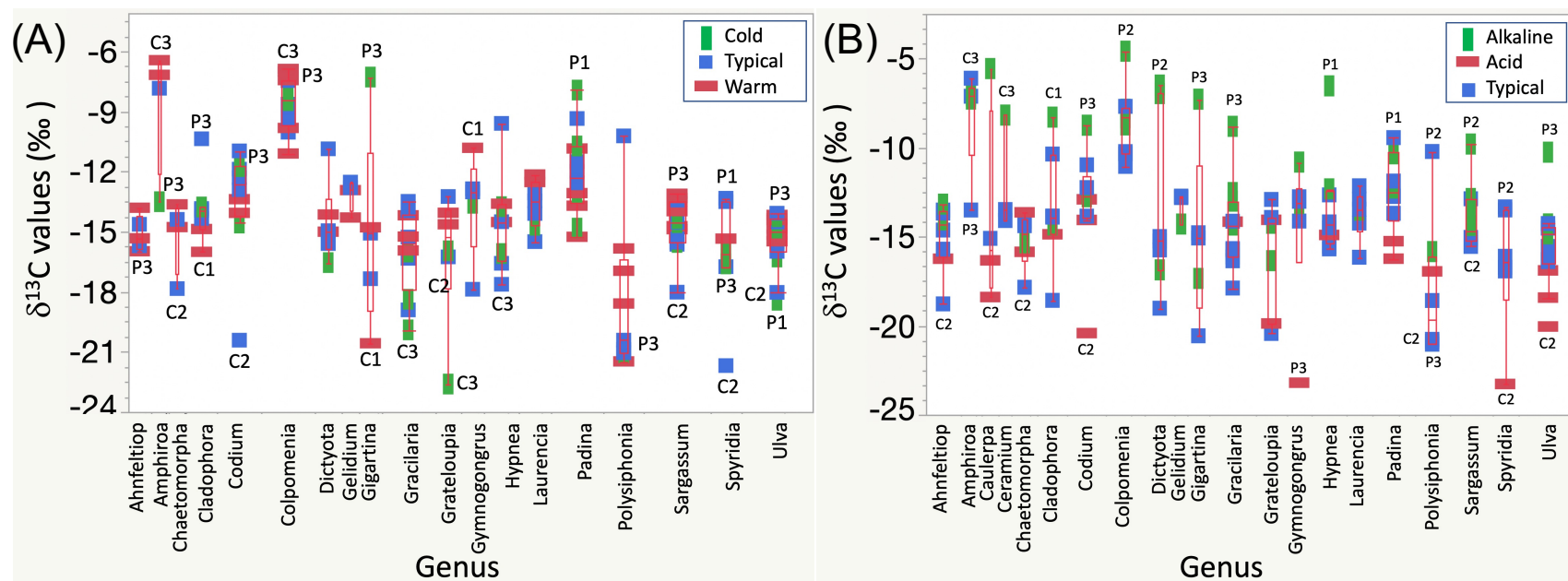
**Fig 5**

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1316 **Fig 6**



**Fig 7**

Proportion of species using different DIC sources according to their isotopic carbon signal

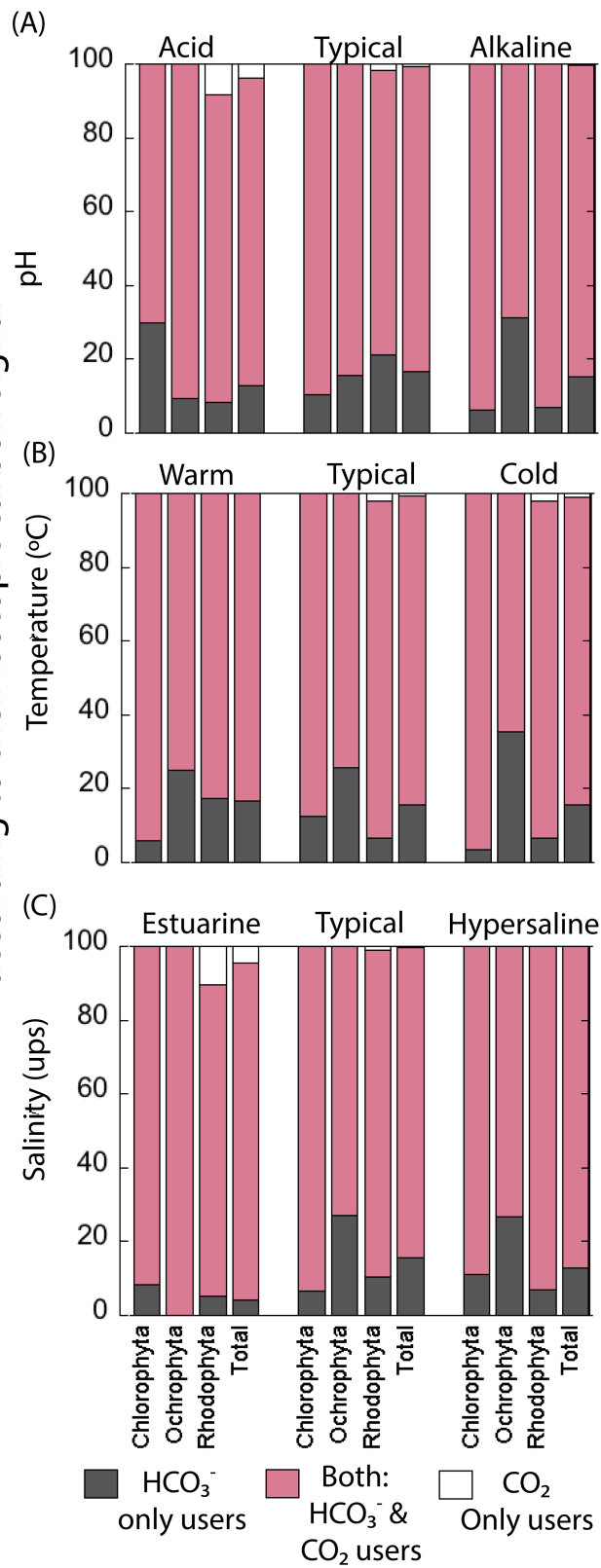
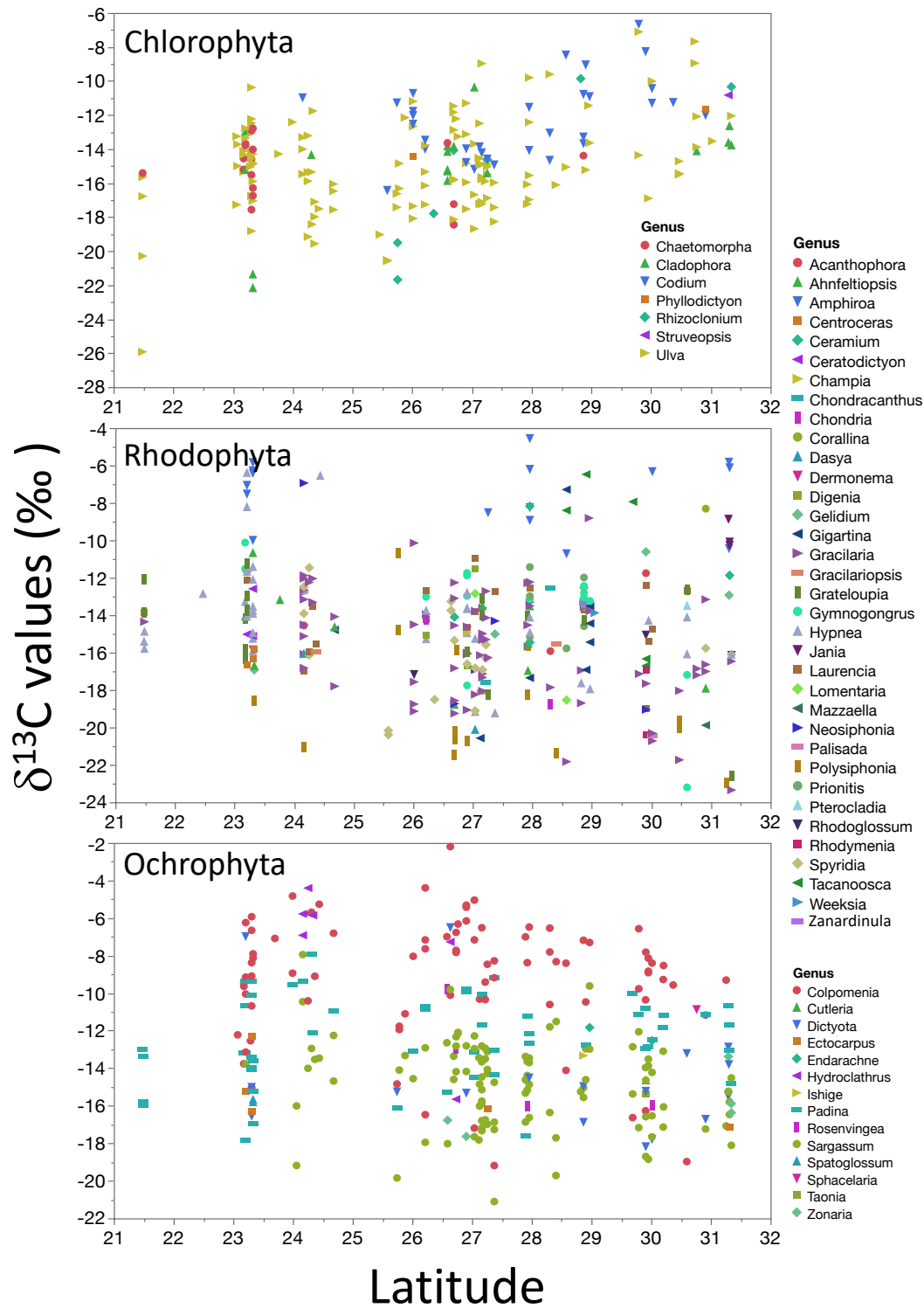
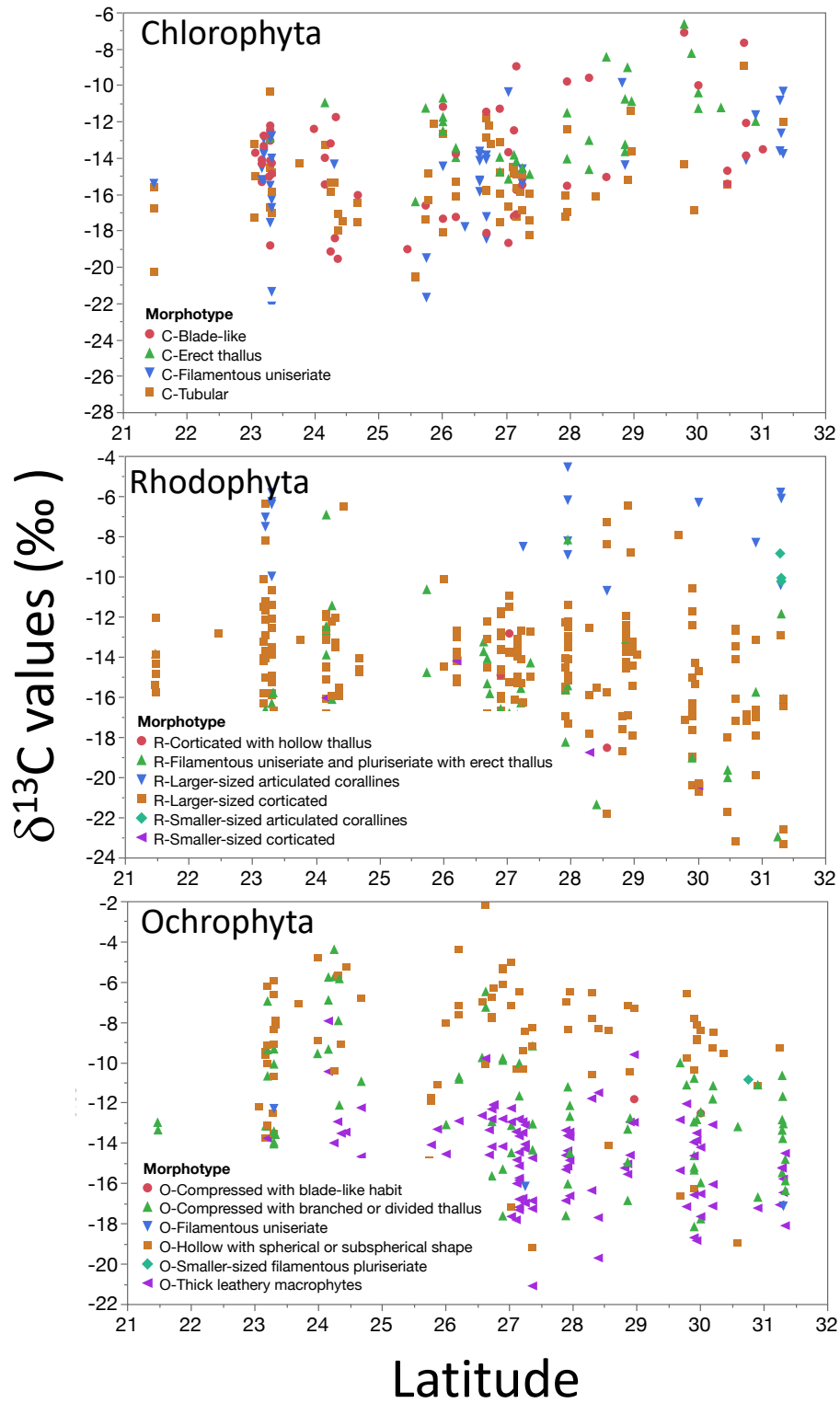


Fig 8



**Fig. 9**



**Fig. 10**

Table 1. Carbon isotopic composition (‰) in species of ~~Phyla~~ Phylum Chlorophyta collected along Gulf of California coastlines.

Species (n composite samples)	$\delta^{13}\text{C} \pm \text{SD}$ (Min to Max, ‰)
<i>Chaetomorpha</i> sp. (3)	-13.7 $\pm$ 0.8 (-14.6 to -12.9)
<i>C. antennina</i> (10)	-14.6 $\pm$ 1.1 (-16.3 to -12.8)
<i>C. linum</i> (5)	-16.8 $\pm$ 1.6 (-18.45 to -14.6)
<i>Codium</i> sp. (5)	-11.6 $\pm$ 3.0 (-14.1 to -6.7)
<i>C. amplivesiculatum</i> (8)	-14.4 $\pm$ 2.7 (-20.4 to -11.3)
<i>C. brandegeei</i> (7)	-11.8 $\pm$ 1.2 (-13.7 to -10.4)
<i>C. fragile</i> (4)	-13.0 $\pm$ 2.7 (-14.8 to -9.0)
<i>C. simulans</i> (9)	-11.4 $\pm$ 2.2 (-14.9 to -8.3)
<i>Ulva</i> sp. (12)	-14.0 $\pm$ 3.9 (-19.2 to -7.1)
<i>U. acanthophora</i> (25)	-15.8 $\pm$ 1.7 (-18.3 to -11.4)
<i>U. clathrata</i> (8)	-16.4 $\pm$ 2.0 (-20.5 to -14.5)
<i>U. compressa</i> (4)	-17.8 $\pm$ 2.4 (-20.6 to -15.4)
<i>U. flexuosa</i> (13)	-16.0 $\pm$ 3.7 (-25.9 to -10.4)
<i>U. intestinalis</i> (16)	-15.3 $\pm$ 2.5 (-20.3 to -8.9)
<i>U. lactuca</i> (31)	-14.1 $\pm$ 3.1 (-19.6 to -7.7)
<i>U. linza</i> (6)	-15.6 $\pm$ 2.4 (-19.4 to -13.2)
<i>U. lobata</i> (5)	-13.2 $\pm$ 1.9 (-15.3 to -11.1)
<i>U. prolifera</i> (3)	-14.2 $\pm$ 1.8 (-15.5 to -12.2)

Table 2. Carbon isotopic composition (‰) in species of ~~Phyla~~Phylum Ochrophyta collected along Gulf of California coastlines.

Species (n composite samples)	$\delta^{13}\text{C} \pm \text{SD}$ (Min to Max, ‰)
<i>Colpomenia</i> sp. (11)	-11.0 $\pm$ 3.7 (-19.0 to -5.4)
<i>C. ramosa</i> (4)	-11.4 $\pm$ 2.6 (-13.8 to -7.8)
<i>C. sinuosa</i> (7)	-10.2 $\pm$ 3.0 (-16.3 to -7.2)
<i>C. tuberculata</i> (64)	-8.7 $\pm$ 3.2 (-19.2 to -2.2)
<i>Padina</i> sp. (15)	-11.1 $\pm$ 1.5 (-13.1 to -7.9)
<i>P. crispata</i> (3)	-11.3 $\pm$ 1.7 (-12.5 to -10.1)
<i>P. durvillaei</i> (36)	-13.2 $\pm$ 2.6 (-20.0 to -9.2)
<i>Sargassum</i> sp. (34)	-14.3 $\pm$ 2.4 (-18.7 to -8.0)
<i>S. herporhizum</i> (7)	-13.7 $\pm$ 1.6 (-16.6 to -11.5)
<i>S. horridum</i> (12)	-15.5 $\pm$ 2.9 (-19.7 to -9.5)
<i>S. johnstonii</i> (10)	-15.4 $\pm$ 2.0 (-17.7 to -11.8)
<i>S. lapazeanum</i> (7)	-14.549 $\pm$ 1.659 (-17.2 to -12.8)
<i>S. sinicola</i> (31)	-15.1 $\pm$ 2.4 (-21.1 to -12.1)

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1327 Table 3. Carbon isotopic composition (‰) in species of ~~Phyla~~ Phylum Rhodophyta collected along  
 1328 Gulf of California coastlines.

Species (n composite samples)	$\delta^{13}\text{C} \pm \text{SD}$ (Min to Max, ‰)
<i>Gracilaria</i> sp. (18)	-15.5 $\pm$ 2.4 (-21.8 to -12.2)
<i>Gracilaria</i> sp.2 (3)	-14.4 $\pm$ 3.7 (-18.7 to -12.3)
<i>G. crispata</i> (7)	-15.1 $\pm$ 3.0 (-19.1 to -10.1)
<i>G. pacifica</i> (6)	-16.5 $\pm$ 1.6 (-18.6 to -13.6)
<i>G. spinigera</i> (3)	-14.9 $\pm$ 3.8 (-17.7 to -12.2)
<i>G. subsecundata</i> (8)	-15.9 $\pm$ 2.8 (-20.3 to -12.8)
<i>G. tepocensis</i> (3)	-15.1 $\pm$ 1.9 (-17.0 to -13.2)
<i>G. textorii</i> (4)	-16.2 $\pm$ 2.6 (-18.1 to -14.3)
<i>G. turgida</i> (5)	-15.3 $\pm$ 3.6 (-20.7 to -12.0)
<i>G. vermiculophylla</i> (16)	-15.9 $\pm$ 3.8 (-23.4 to -8.8)
<i>Hypnea</i> sp. (14)	-14.9 $\pm$ 2.6 (-20.9 to -11.4)
<i>H. johnstonii</i> (5)	-11.2 $\pm$ 3.5 (-13.8 to -6.5)
<i>H. pannosa</i> (5)	-11.8 $\pm$ 3.3 (-15.0 to -6.4)
<i>H. spinella</i> (6)	-16.4 $\pm$ 1.8 (-19.2 to -14.9)
<i>H. valentiae</i> (6)	-15.2 $\pm$ 2.3 (-19.2 to -12.7)
<i>Laurencia</i> sp. (8)	-12.9 $\pm$ 1.2 (-14.7 to -10.5)
<i>L. pacifica</i> (8)	-14.9 $\pm$ 2.2 (-19.0 to -12.7)
<i>L. papillosa</i> (3)	-15.7 $\pm$ 0.3 (-15.9 to -15.6)
<i>Spyrida</i> sp. (5)	-17.1 $\pm$ 1.12 (-19.1 to -16.1)
<i>S. filamentosa</i> (14)	-15.9 $\pm$ 3.8 (-26.2 to -11.5)

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Table 4. Summary of the estimated regression coefficients for each simple linear regression analyses and on the constant of fitted regression models. Estimated regression coefficients includes degrees of freedom for the error (DFE), root-mean-square error (RMSE), coefficients of determination ( $R^2$ ) and the adjusted  $R^2$  statistics, Mallows' Cp criterion (Cp), Akaike Information Criterion (AIC), Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the test (Prob > F). Models information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t| (values \* are significant).

Independent variables	Estimated regression coefficients									Model constant (a)			
	DFE	RMSE	$R^2$	Adjust $R^2$	Cp	AICc	BIC	F ratio	Prob > F	$\delta^{13}\text{C}$ (‰)	SE	t ratio	Prob >  t
Inherent macroalgae properties													
Phyla	806	3.66	0.08	0.07	3	4,401	4,420	33.1	<.0001**	-13.98	0.13	-107.4	<.0001**
Morphofunctional	788	3.10	0.35	0.34	21	4,149	4,251	21.6	<.0001**	-14.21	0.35	-40.80	<.0001**
Genus	746	2.92	0.46	0.41	63	4,104	4,393	10.1	<.0001**	-14.71	0.23	-62.64	<.0001*
Species	641	2.79	0.57	0.46	168	4,195	4,898	5.2	<.0001**	-14.60	0.16	-93.22	<.0001**
Biogeographical collection zone													
GC coastline	807	3.79	0.01	0.01	2	4,456	4,470	7.4	0.0067*	-13.97	0.13	-104.5	<.0001**
Coastal sector	803	3.73	0.05	0.04	6	4,433	4,465	7.9	<.0001*	-14.12	0.16	-90.85	<.0001**
Latitude	807	3.80	0.00	0.00	2	4,462	4,476	1.5	0.23	-12.25	1.41	-8.71	<.0001**
Longitude	807	3.81	0.00	0.00	2	4,463	4,477	0.1	0.80	-15.44	5.83	-2.65	0.0082*
Habitat features													
Substrate	807	3.80	0.00	0.00	2	4,460	4,474	3.2	0.08	-13.82	0.15	-92.06	<.0001*
Hydrodynamic	807	3.80	0.00	0.00	2	4,462	4,476	1.3	0.26	-13.88	0.15	-95.00	<.0001**
Emersion level	807	3.69	0.06	0.06	2	4,412	4,427	52.2	<.0001**	-14.05	0.13	-107.6	<.0001**
Environmental conditions													
Temperature	802	3.70	0.01	0.01	2	4,390	4,404	5.4	0.0207*	-16.11	0.96	-16.78	<.0001*
pH	807	3.73	0.04	0.04	2	4,430	4,444	33.4	<.0001**	-32.45	3.21	-10.13	<.0001**
Salinity	806	3.80	0.00	0.00	2	4,456	4,470	0.9	0.34	-15.77	1.91	-8.27	<.0001**

\*p<0.05, \*\*p<0.0001

Table 5. Summary of the estimated regression coefficients for each multivariate linear regression analyses and on their constant of fitted regression models performed in individuals binned by genus. Estimated regression coefficients include degrees of freedom for the error (DFE), root-mean-square error (RMSE), coefficients of determination ( $R^2$ ) and the adjusted  $R^2$  statistics, Mallows' Cp criterion (Cp), Akaike Information Criterion (AIC), Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the test (Prob > F). Model information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t| (values \* are significant).

Independent variables	DFE	RMSE	Estimated regression coefficients							Prob > F	Model constant (a)				Prob >  t
			R <sup>2</sup>	Adjust R <sup>2</sup>	Cp	AICc	BIC	F ratio	δ <sup>13</sup> C (‰)		SE	t ratio			
Coastal sector	652	2.78	0.57	0.47	157	4,169	4,834	20.0	<.0001*	-17.52	0.64	-27.24	<.0001*		
Substrate	711	2.90	0.49	0.42	98	4,140	4,577	0.4	0.52	-16.35	0.62	-26.20	<.0001*		
Hydrodynamic	714	2.87	0.50	0.43	95	4,120	4,545	0.1	0.78	-16.53	0.64	-25.95	<.0001*		
Emersion level	713	2.77	0.53	0.47	96	4,060	4,489	153.0	<.0001*	-16.65	0.60	-27.85	<.0001*		
Temperature	695	2.81	0.50	0.43	109	4,083	4,564	98.4	<.0001*	-14.60	0.92	-15.91	<.0001*		
Temperature ranges	686	2.87	0.49	0.40	118	4,128	4,645	97.7	<.0001*	-12.91	0.40	-31.97	<.0001*		
pH	701	2.86	0.51	0.43	108	4,134	4,611	156.6	<.0001*	-28.57	2.69	-10.64	<.0001*		
pH ranges	697	2.67	0.57	0.51	112	4,028	4,522	152.2	<.0001*	-16.39	0.58	-28.05	<.0001*		
Salinity	697	2.89	0.50	0.42	111	4,151	4,640	162.2	<.0001*	-17.75	1.63	-10.88	<.0001*		
Salinity ranges	721	2.91	0.47	0.41	86	4,117	4,504	167.8	<.0001*	-17.64	0.74	-23.68	<.0001*		

Table 6. Summary of the estimated regression coefficients for each multivariate linear regression analyses and on their constant of fitted regression models performed in individuals binned by coastline sector and genus. Estimated regression coefficients include degrees of freedom for the error (DFE), root-mean-square error (RMSE), coefficients of determination ( $R^2$ ) and the adjusted  $R^2$  statistics, Mallow's Cp criterion (Cp), Akaike Information Criterion (AIC), Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the test (Prob > F). Model information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t| (values \* are significant).

Independent variables	DFE	RMSE	Estimated regression coefficients						Prob > F	Model constant (a)			
			R <sup>2</sup>	Adjust R <sup>2</sup>	Cp	AICc	BIC	F ratio		$\delta^{13}\text{C}$ (‰)	SE	t ratio	Prob >  t
Substrate	590	2.76	0.62	0.47	219	4,287	5,155	15.8	<.0001*	-17.08	0.66	-25.72	<.0001*
Hydrodynamic	592	2.73	0.62	0.49	217	4,266	5,128	18.6	<.0001*	-17.18	0.67	-25.70	<.0001*
Protection level	590	2.75	0.62	0.48	219	4,285	5,153	20.0	<.0001*	-17.51	0.64	-27.22	<.0001*
Emersion level	603	2.69	0.63	0.50	206	4,217	5,045	18.6	<.0001*	-17.47	0.64	-27.49	<.0001*
Temperature ranges	569	2.74	0.61	0.46	235	4,293	5,202	28.0	<.0001*	-13.73	0.45	-30.32	<.0001*
pH ranges	580	2.50	0.69	0.57	229	4,155	5,051	9.7	0.0019*	-16.88	0.62	-27.15	<.0001*
Salinity ranges	631	2.76	0.58	0.47	176	4,183	4,913	21.2	<.0001*	-18.30	0.79	-23.05	<.0001*

Table 7. Summary of the estimated regression coefficients for each multivariate linear regression analyses and on their constant of fitted regression models performed in individuals binned in coastline sector, habitats features, environmental conditions, and Physiological performed separately by morpho-functional groups and genus. Estimated regression coefficients include degrees of freedom for the error (DFE), root-mean-square error (RMSE), coefficients of determination ( $R^2$ ) and the adjusted  $R^2$  statistics, Mallows' Cp criterion (Cp), Akaike Information Criterion (AIC), Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the test (Prob > F). Model information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t| (values \* are significant).

Full model	Estimated regression coefficients								Prob > F	Model constant (a)			
	DFE	RMSE	R <sup>2</sup>	Adjust R <sup>2</sup>	Cp	AICc	BIC	F ratio		δ <sup>13</sup> C (‰)	SE	t ratio	Prob >  t
Coastline sector + Habitats features + Morphofunctional group													
I-Morpho-functional	593	2.79	0.60	0.46	216	4,301	5,160	20.8	<.0001*	-13.49	0.57	-23.52	<.0001*
Coastline sector + Environmental conditions + Morphofunctional group													
II-Morpho-functional	680	2.90	0.51	0.42	129	4,189	4,750	25.1	<.0001*	-13.42	0.54	-24.74	<.0001*
Coastline sector + Habitat features+ Genus													
I-Genus	482	2.66	0.71	0.51	327	4,565	5,655	15.8	<.0001*	-16.93	0.73	-23.27	<.0001*
Coastline sector + Environmental conditions + Genus													
II-Genus	494	2.49	0.72	0.55	310	4,374	5,438	14.8	0.0001*	-13.55	0.64	-21.17	<.0001*

Table 8. Constant of fitted regression model explaining the  $\delta^{13}\text{C}$  variability by morpho-functional groups. Model information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t|. Only morpho-functional groups with significant effects are enlisted.

Term	Estimated	SE	Razón t	Prob >  t
Model constant	-14.2	0.4	-40.80	<.0001**
R-Smaller-sized articulated corallines	4.5	1.7	2.58	0.0100*
O-Compressed with branched or divided thallus	1.2	0.5	2.66	0.0079*
C-Erect thallus	1.8	0.6	2.84	0.0046*
R-Larger-sized articulated corallines	6.3	0.8	7.95	<.0001*
O-Hollow with spherical or subspherical shape	5.0	0.5	10.51	<.0001*
R-Blade-like with one of few layers of cells	-5.9	3.0	-1.98	0.0476*
C-Tubular	-1.6	0.5	-3.26	0.0012**
R-Filamentous uni&pluriseriate with erect thallus	-2.2	0.6	-3.92	<.0001*
R-Flattened macrophytes with cortication	-8.9	1.3	-7.10	<.0001*

\*p<0.05, \*\*p<0.0001

1386 Table 9. Constant of fitted regression model explaining the  $\delta^{13}\text{C}$  variability by genus. Model  
 1387 information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t|.   
 1388 Only genus with significant effects are enlisted.

Term	Estimated	SE	Razón t	Prob >  t
Model constant	-14.7	0.2	-62.64	<.0001**
<i>Corallina</i>	6.4	2.9	2.22	0.0269*
<i>Tacanoosca</i>	3.5	1.3	2.71	0.0070*
<i>Jania</i>	5.0	1.7	2.97	0.0031*
<i>Struveopsis</i>	4.1	1.3	3.15	0.0017*
<i>Codium</i>	2.3	0.6	4.08	<.0001**
<i>Padina</i>	2.2	0.5	4.8	<.0001**
<i>Hydroclathrus</i>	7.3	1.1	6.59	<.0001**
<i>Amphiroa</i>	6.8	0.8	9.05	<.0001**
<i>Colpomenia</i>	5.4	0.4	14.02	<.0001*
<i>Spyridia</i>	-1.5	0.7	-2.10	0.0361*
<i>Gracilaria</i>	-0.9	0.4	-2.18	0.0294*
<i>Polysiphonia</i>	-3.7	0.8	-4.82	<.0001**
<i>Schizymenia</i>	-19.1	2.1	-9.33	<.0001**

\*p<0.05, \*\*p<0.001

1392 Table 10. Constant of fitted regression model explaining the  $\delta^{13}\text{C}$  variability by species. Model  
 1393 information includes value of the constant a ( $\delta^{13}\text{C}$ , ‰), standard error (SE), t ratio and Prob > |t|.  
 1394 Only genus with significant effects are enlisted.

Term	$\delta^{13}\text{C}$ , ‰ estimated	SE	Razón t	Prob >  t
Model constant	-14.6	0.2	-93.22	<.0001**
<i>Hypnea pannosa</i>	2.8	1.3	2.24	0.0256*
<i>Colpomenia ramosa</i>	3.2	1.4	2.27	0.0237*
<i>Corallina vancouverensis</i>	6.3	2.8	2.27	0.0238*
<i>Caulerpa peltata</i>	3.9	1.6	2.4	0.0165*
<i>Codium</i> sp.	3.0	1.3	2.4	0.0167*
<i>Amphiroa misakiensis</i>	7.1	2.8	2.55	0.0110*
<i>Jania</i> sp.	5.0	2.0	2.56	0.0106*
<i>Codium brandegeei</i>	2.8	1.1	2.63	0.0088**
<i>Hypnea johnstonii</i>	3.4	1.3	2.74	0.0063**
<i>Tacanoosca uncinata</i>	3.4	1.3	2.74	0.0062**
<i>Struveopsis</i> sp.	4.0	1.4	2.86	0.0044**
<i>Padina durvillaei</i>	1.4	0.5	2.87	0.0043**
<i>Amphiroa</i> sp.3	8.2	2.8	2.95	0.0033**
<i>Codium simulans</i>	3.2	0.9	3.41	0.0007**
<i>Amphiroa</i> sp.2	6.6	1.6	4.1	<.0001**
<i>Colpomenia sinuosa</i>	4.4	1.1	4.17	<.0001**



<i>Colpomenia</i> sp.	3.6	0.9	4.27	<.0001**
<i>Padina</i> sp.	3.5	0.7	4.77	<.0001**
<i>Hydroclathrus clathratus</i>	7.2	1.1	6.82	<.0001**
<i>Amphiroa</i> sp.	8.1	0.9	8.67	<.0001**
<i>Colpomenia tuberculata</i>	5.9	0.4	15.45	<.0001**
<i>Spyrida</i> sp.	-2.5	1.3	-1.97	0.0496*
<i>Pyropia thuretii</i>	-5.5	2.8	-1.98	0.0480*
<i>Ulva acanthophora</i>	-1.2	0.6	-2.06	0.0399*
<i>Grateloupia filicina</i>	-2.4	1.1	-2.08	0.0382*
<i>Rhodymenia</i> sp.	-4.1	2.0	-2.08	0.0380*
<i>Ulva compressa</i>	-3.2	1.4	-2.33	0.0203*
<i>Rhizoclonium riparium</i>	-5.1	1.6	-3.15	0.0017**
<i>Polysiphonia</i> sp.	-4.8	1.4	-3.44	0.0006**
<i>Halymenia actinophysa</i>	-9.9	2.8	-3.57	0.0004**
<i>Cladophora microcladioides</i>	-7.2	2.0	-3.64	0.0003**
<i>Polysiphonia mollis</i>	-5.2	1.1	-4.93	<.0001**
<i>Schizymenia pacifica</i>	-19.2	2.0	-9.76	<.0001**

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\*p<0.05, \*\*p<0.001

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