1	An analysis of the variability of $\delta^{13}$ C in macroalgae from the Gulf of California: indicative of
2	carbon concentration mechanisms and isotope discrimination during carbon assimilation
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## 19 Abstract

The isotopic composition of carbon in macroalgae ( $\delta^{13}$ C) is highly variable, and its prediction 20 21 complex concerning terrestrial plants. The determinants of  $\delta^{13}$ C-macroalgal variations were 22 analyzed in a large stock of specimens that vary in taxa and morphology, collected in shallow 23 marine habitats in the Gulf of California (GC) with distinctive environmental conditions. A large 24  $\delta^{13}$ C variability (-34.6‰ to -2.2‰) was observed. Life forms (taxonomy 57%, morphology and 25 structural organization 34%) explains the variability related to the carbon use physiology. 26 Environmental conditions influenced the  $\delta^{13}$ C-macroalgal values, but did not change the physiology, which is most likely inherently species-specific. Measurements of  $\delta^{13}$ C were used as 27 28 indicative of the presence or absence of carbon concentrating mechanisms (CCMs) and as 29 integrative values of the isotope discrimination during carbon assimilation in the lifecycle macroalgae. Based on  $\delta^{13}$ C signals, macroalgae were classified in four strategies relatives to the 30 capacity of CCM: 1) HCO<sub>3</sub><sup>-</sup> uptake ( $\delta^{13}$ C>-10‰), 2) using a mix of CO<sub>2</sub> and HCO<sub>3</sub><sup>-</sup> uptake (-31  $10 < \delta^{13}C > -30\%$ ), 3) CO<sub>2</sub> diffusive entry ( $\delta^{13}C < -30\%$ ), and 4) calcifying species ( $\delta^{13}C > -10\%$ ). 32 Most species showed a  $\delta^{13}$ C that indicates a CCM using a mix of CO<sub>2</sub> and HCO<sub>3</sub><sup>-</sup> uptake. HCO<sub>3</sub><sup>-</sup> 33 34 uptake is also widespread among GC macroalgae, with many Ochrophyta species. Few species 35 belonging to Rhodophyta relied on CO<sub>2</sub> diffusive entry exclusively, while calcifying macroalgae 36 species using HCO<sub>3</sub><sup>-</sup> included only Amphiroa and Jania. The isotopic signature evidenced the 37 activity of CCM, but it was inconclusive about the preferential uptake of  $HCO_3^-$  and/or  $CO_2$  in 38 photosynthesis and the CCM type expressed in macroalgae. In the carbon use strategies study, 39 diverse and species-specific, complementary techniques to the isotopic tools are required. Keywords:  $\delta^{13}$ C-macroalgal, carbon-concentrating mechanisms, CO<sub>2</sub> diffusive proxy 40

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#### 1. Introduction

43 Macroalgae show a wide diversity of thallus morphologies (e.g., filamentous, articulated, flattened), 44 structural organization (e.g., surface area/volume ratio), and various photosynthetic pigments (e.g., Chlorophyll a, b, phycocyanin) (Lobban and Harrison, 1994). According to the predominant pigment 45 46 contents in the thallus, macroalgae are classified in three Phyla. Considering the interaction of 47 morphologies and photosynthetic pigments, are classified in dozens of groups (Littler and Littler, 48 1980; Littler & Arnold, 1982; Balata et al., 2011). For example, the mixture of chlorophyll (a, b) and 49 carotenoids is dominant in Chlorophyta; chlorophyll (a, c) and fucoxanthin carotenoid is dominant 50 in Ocrophyta, while Rhodophyta contains chlorophyll (a, d), carotenoid, and a mixture of phycobilin 51 (e.g., phycocyanin, phycoerythrin, allophycocyanin) (Bold and Wynne, 1978; Masojidek et al., 52 2004; Gateau et al., 2017). Both traits work as an excellent approximation to explain the 53 fundamentals of metabolism, growth, zonation, and colonization (Littler and Littler, 1980; Littler 54 and Arnold, 1982; Nielsen and Sand-Jensen, 1990; Vásquez-Elizondo and Enríquez, 2017). 55 In marine environments, where  $pH \sim 8.1 \pm 1$ ,  $HCO_3^-$  accounts for 98% of the total dissolved inorganic 56 carbon (DIC) due to the low diffusion rate of CO<sub>2</sub> in seawater, resulting in a high HCO<sub>3</sub><sup>-</sup>:CO<sub>2</sub> ratio 57 (150:1) (Sand-Jensen and Gordon, 1984). The limitations to growth imposed by low  $CO_2$ 58 concentrations in seawater are compensated by carbon concentrating mechanisms (CCMs) in most

59 macroalgae that increase the internal inorganic carbon concentration near the site of RuBisCo 60 activity (Giordano et al., 2005). Therefore, the absorption of  $HCO_3^-$  by most macroalgae is the 61 primary source of inorganic carbon for photosynthesis, but some species depend exclusively on the 62 use of dissolved  $CO_2$  that enters cells by diffusion (Maberly et al., 1992; Beardall and Giordano, 63 2002; Raven et al., 2002a, b; Giordano et al., 2005). Hence, macroalgal species with productivity 64 limited by lacking CCM's (have low plasticity for carbon inorganic forms uptake) seems to be 65 restricted to subtidal habitats and composed mainly by red macroalgae (but without a morphological patron apparent) (Cornwall et al., 2015, Kübler and Dungeon, 2015). The rest of the macroalgae with
CCM occupies from the intertidal to the deep subtidal.

68 The habitat features and environmental conditions in marine ecosystems modify the main 69 macroalgae photosynthesis drivers, such as light (Anthony et al., 2004; Johansson and Snoeijs, 70 2014), DIC (Zeebe and Wolf-Gladrow, 2001; Brodeur et al., 2019), and inorganic nutrients 71 (Teichberg et al., 2010; Ochoa-Izaguirre and Soto-Jiménez, 2015). These factors could generate 72 negative consequences for their productivity, principally when they cause resources limitation. Each 73 factor varies from habitat to habitat (e.g., local scale: from intertidal to subtidal and global scale: 74 from temperate to tropical regions), and as in response to these environmental changes, macroalgae 75 can modulate their photosynthetic mechanism (Lapointe and Duke, 1984; Dudgeon et al., 1990; 76 Kübler and Davison 1993, Young and Beardall., 2005). The modulation, to increase their 77 photosynthetic activity (up-and-down-regulation processes), implies a physiological acclimation 78 enhancing the transport of DIC ( $CO_2$ ,  $HCO_3^{-}$ ) into the cell and its fixation rates (Madsen and 79 Maberly, 2003; Klenell et al., 2004; Zou et al., 2004; Giordano et al., 2005; Enríquez and Rodriguez-80 Román, 2006; Rautenberger et al., 2015).

81 The  $\delta^{13}$ C on the thallus of marine macrophytes are indicative of the carbon source used (CO<sub>2</sub> or 82  $HCO_3$ ) in photosynthesis and allow infer the presence or absence of CCM's (Maberly et al., 1992; 83 Raven et al., 2002a; Giordano et al., 2005). However, the isotopic signature may be inconclusive for 84 determine the preference of the carbon source (Roleda and Hurd, 2012). Also, the  $\delta^{13}$ C signal in the 85 algal thallus can be used as an indicator of the physiological state of photosynthetic metabolism (Kim 86 et al., 2014; Kübler and Dungeon, 2015). For example,  $\delta^{13}$ C variability depends, in part, on the life 87 forms as taxonomy, morphology, and structural organization (Mercado et al., 2009, Marconi et al., 2011, Roleda and Hurd, 2012; Lovelock et al., 2020).  $\delta^{13}$ C is also modulated by the interaction to 88 89 environmental conditions (e.g., light, DIC, and nutrients) (Cornelisen et al., 2007; Dudley et al.,

90 2010; Carvalho et al., 2010ab; Roleda and Hurd, 2012; Mackey et al., 2015; Rautenberger et al., 91 2015). In this study, our objective was to investigate the contributions of life forms, the changes in the habitat features, and environmental conditions to the  $\delta^{13}$ C macroalgal variability in communities 92 93 in the Gulf of California (GC). To reach our objective, we collected a large stock of macroalgae 94 specimens of a diversity of species characterized by various morphological and physiological 95 properties. Besides high diversity, in terms of life forms, we selected various shallow marine habitats 96 along a latitudinal gradient in the GC or the sample collection, characterized by unique and changing 97 environmental factors. The GC features abundant and diverse macroalgae populations, acclimated 98 and adapted to diverse habitats with environmental conditions, determining the light, DIC, and 99 nutrients availability. The  $\delta^{13}$ C signal from the thallus of macroalgae were used as indicative of the 100 presence or absence of CCMs and as integrative values of the isotope discrimination during carbon 101 assimilation and respiration along lifecycle macroalgae in macroalgae communities in the GC in the 102 function of taxa and environmental factors (Maberly et al., 1992; Raven et al., 2002a; Hepburn et 103 al., 2011; Díaz-Pulido et al., 2016). Because the GC is a subtropical zone with high irradiance and 104 specimens were collected in the intertidal and shallow subtidal zone, we expect to find a high 105 proportion of species with active uptake HCO<sub>3</sub><sup>-</sup> ( $\delta^{13}$ C>-10‰). A third objective was to explore any geographical pattern in the  $\delta^{13}$ C macroalgal along and between the GC bioregions. Previous studies 106 have indicated changes in the  $\delta^{13}$ C signal with latitude, mainly related to the light and temperature 107 108 (Mercado et al., 2009; Marconi et al., 2011; Stepien, 2015; Hofmann and Heesch, 2018; Lovelock 109 et al., 2020). Macroalgae as biomonitors constitute an efficient tool in monitoring programs in large 110 geographical regions (Balata et al., 2011) and for environmental impact assessments (Ochoa-111 Izaguirre and Soto-Jiménez, 2015).

### 112 **2. Materials and Methods**

#### 113 **2.1. Gulf of California description**

114 The Gulf of California is a subtropical, semi-enclosed sea of the Pacific coast of Mexico, with 115 exceptionally high productivity being the most important fishing region for Mexico and one of the 116 most biologically diverse worldwide marine areas (Zeitzschel, 1969; Espinosa-Carreón and Valdez-117 Holguín 2007; Lluch-Cota et al., 2007; Páez-Osuna et al., 2017). GC represents only 0.008% of the 118 area covered by the seas of the planet (265,894 km<sup>2</sup>, 150 km wide, and 1000 km long covering >9 119 degrees latitude) but has a high physiographic diversity and is biologically mega-diverse with many 120 endemic species, including  $\sim 766$  macrofauna species and/or sub-species where the major number 121 belong to Arthropoda (118 spp) and Mollusca (460) taxas (Brusca et al., 2005; Wilkinson et al., 122 2009; Espinosa-Carreón and Escobedo-Urías, 2017) and 116 macroalgae species (Norris, 1975, 123 1985; Espinoza-Avalos, 1993).

124 Regionalization criteria of the GC include phytoplankton distribution (Gilbert and Allen, 1943), 125 topography (Rusnak et al., 1964) and depth (Álvarez-Borrego, 1983), oceanographic characteristics 126 (Roden and Emilson, 1979; Álvarez-Borrego, 1983; Marinone and Lavin 2003), biogeography 127 (Santamaría-del-Ángel et al., 1994), and bio-optical characteristics (Bastidas-Salamanca et al., 128 2014). The topography is variable along with GC, includes submarine canyons, basins, and variable 129 continental platforms. Besides, GC presents complex hydrodynamic processes, including internal 130 waves, fronts, upwelling, vortices, mixing of tides. The gulf's coastline is divided into three shores: 131 extensive rocky shores, long sandy beaches, numerous scattered estuaries, coastal lagoons, and open 132 muddy bays, tidal flats, and coastal wetlands (Lluch-Cota et al., 2007).

133 The Gulf of California is different in the north and the south, related to a wide range of 134 physicochemical factors. The surface currents seasonally change direction and flow to the southeast 135 with maximum intensity during the winter and to the northwest in summer (Roden, 1958). The 136 northern part is very shallow (<200 m deep averaged), divided into Upper Gulf, northern Gulf, and 137 Grandes Islas. The surrounding desserts largely influence this region (Norris, 2010) shows marked 138 seasonal changes in coastal surface seawater temperatures (Martínez-Díaz de León et al., 2006; 139 Marinone, 2007). Tidal currents induce a significant cyclonic circulation through June to September 140 and anticyclonic from November to April (Carrillo et al., 2002; Bray, 1988; Velasco-Fuentes and 141 Marinone, 1999; Martínez-Díaz-de-León, 2001). The southern part consists of a series of basins 142 whose depths increase southwards (Fig. 1). The intertidal macroalgae in the southern region are 143 subject to desiccation, mostly during summer. The water column's physicochemical characteristics 144 are highly influenced by the contrasting climatic seasons in the GC, the dry season (nominally from 145 November to May), and the rainy season (from June to October). Annual precipitation (1,080 mm y<sup>-</sup> <sup>1</sup>) and evaporation (56 mm y<sup>-1</sup>) rates registered during the past 40 years were  $881\pm365$  mm y<sup>-1</sup> and 146  $53\pm7$  mm y<sup>-1</sup>, respectively (CNA, 2012). 147

148 Previous macroalgae floristic studies of the GC, report around 669 species, including 116 endemic 149 species (Norris, 1975; Espinoza-Avalos, 1993; Pedroche and Senties, 2003). Many endemic 150 species currently have a wide distribution along the Pacific Ocean coast, but with GC origin 151 (Dreckman, 2002; Aguilar-Rosas et al., 2014). Based on oceanographic characteristics (Roden and 152 Groves, 1959) and in the endemic species distribution (Aguilar Rosas and Aguilar Rosas, 1993; 153 Espinoza-Avalos, 1993), the GC can be classified into three phycofloristic zones: 1) the first zone 154 located from the imaginary line connecting San Francisquito Bay, B.C. to Guaymas, Sonora, with 155 51 endemic species. 2) the second zone with an imaginary line from La Paz Bay (B.C.S.) to 156 Topolobampo (Sinaloa) with 41 endemic species. 3) the third zone is located with an imaginary line 157 from Cabo San Lucas (B.C.S.) to Cabo Corrientes (Jalisco) with ten endemic species. Besides, 14 endemic species are distributed throughout the GC (Espinoza-Ávalos, 1993). The macroalgal communities are subject to the changing environmental conditions in the diverse habitats in the GC that delimits their zonation, which tolerates a series of anatomical and physiological adaptations to water movement, temperature, sun exposure, and light intensities, low pCO2, desiccation (Espinoza-Avalos 1993).

### 163 **2.1 Macroalgae sampling**

In this study, the GC coastline (21°-30°N latitude) was divided into six coastal sectors based on the three phycofloristic zones along peninsular and continental GC coastlines (Fig. 1a). In each coastal sector, selected ecosystems and representative habitats were sampled based on macroalgae communities' presence and habitat characterization. Habitats were classified by substrate type (e.g., sandy-rock, rocky shore), hydrodynamic (slow to faster water flows), protection level (exposed or protected sites), and immersion level (intertidal or subtidal) (Fig. 1b).

170 Based on the local environmental factors, 4-5 macroalgae specimens of the most representative 171 species were gathered by hand (free diving) during low tide. A total of 809 composite samples were 172 collected from marine habitats along both GC coastlines. The percentages of specimens collected for 173 the substrate type were sandy-rock 28% and rocky shores 72% based on the habitat features. In the 174 hydrodynamic, 30% of the specimens were collected in habitats with slow to moderate and 70% with 175 moderate to fast water movement. Regarding the protection level, 57% were exposed specimens, 176 and 43% were protected. Finally, 56% were intertidal and 44% subtidal macroalgae organisms 177 concerning the emersion level. About half of the protected specimens were collected in isolated rock 178 pools, which was noted.



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180 calibrated multiparameter sonde (Y.S.I. 6600V) and the habitat characteristics mentioned above 181 noted. Besides, composite water samples were collected for a complementary analysis of nutrients, alkalinity (and their chemical components), and  $\delta^{13}$ C-DIC (data non-included). Briefly, the 182 183 representative habitats were classified by pH levels in >9.0 "alkalinized", 7.9-8.2 'typical' and <7.9 184 "acidified". Based on the temperature in colder <20°C, typical 20-25°C, and warmer >25°C. 72% of 185 the specimens were collected at typical pH values, 22% alkalinized, and 6% in acidified seawater. 186 Regarding the temperature, about 55% of the specimens were collected at typical, 31% at warmer, 187 and 14% at colder seawaters. Regarding salinity, most of the ecosystems showed typical values for 188 seawater (35.4±0.91 ups, from 34.5 to 36.1 ups). In this study, the collection surveys were conducted 189 during spring (March-April) and dry season (nominally from November to May) from 2008 to 2014. 190 Only in few selected ecosystems located at C1, C2, and C3 sectors, one sampling survey was 191 conducted at the end of the rainy season (nominally from June to October in 2014). Thus, these 192 ecosystems were possible to include habitat with a salinity range varying from estuarine  $(23.5\pm3.0)$ 193 ups) to hypersaline (42.7±7.0 ups) values. These habitats were mainly isolated rock pools, and only 194 a few were sites near tidal channels receiving freshwater discharges. About 95% of the specimens 195 were collected at typical seawater salinity (34-36 ups) and only 1.5 and 3.5% in estuarine (<30 ups) 196 and hypersaline (>37 ups) environments, respectively. Detailed information on the selected shallow 197 marine ecosystems, habitat characterization, and environmental conditions is summarized in the 198 inserted table in Fig. 1.

### 199 2.2 Macroalgae processing and analysis of the isotopic composition of carbon

The collected material was washed *in situ* with surface seawater to remove the visible epiphytic organisms, sediments, sand, and debris and then thoroughly rinsed with MilliQ water. The composite samples were double-packed in a plastic bag, labeled with the locality's name and collection date, placed in an ice-cooler box to be kept to 4°C, and immediately transported to the laboratory UASFacimar in Mazatlán. In the field, sample aliquots were also preserved in 4% v/v formaldehyde
solution for taxonomic identification to the genus or species level (when possible). The following
GC macroalgal flora identification manuals were consulted: Dawson, 1944; 1954; 1956; 1961; 1962;
1963; Setchell and Gardner, 1920; 1924; Abbot and Hollenberg, 1976; Ochoa-Izaguirre et al., 2007;
Norris, 2010).

209 In the laboratory, macroalgae samples were immediately frozen at -30°C until analysis. Then, 210 samples were freeze-dried at -38°C and 40 mm Hg for 3 days, upon which they were ground to a 211 fine powder and exposed to HCl vapor for 4 h (acid-fuming) to remove carbonates and dried at 60°C 212 for 6 h (Harris et al. 2001). Aliquots of ~5 mg were encapsulated in tin cups (5x9 mm) and stored in 213 sample trays until analysis. Macroalgae samples were sent to the Stable Isotope Facility (SIF) at the University of California at Davis, CA, USA. Natural <sup>13</sup>C relative abundance relative to <sup>12</sup>C in 214 215 samples was determined with mass spectrometry, using a Carlo Erba elemental analyzer attached to 216 a Finnigan Delta S mass spectrometer equipped with a Europa Scientific stable isotope analyzer 217 (ANCA-NT 20-20) and a liquid/solid preparation unit (PDZ, Europa, Crewz, UK). Isotope ratios of 218 the samples were calculated using the equation  $\delta$  (‰)=(R<sub>sample</sub>/R<sub>standard</sub>-1)]x1000, where R=<sup>13</sup>C/<sup>12</sup>C. 219 The R<sub>standard</sub> is relative to the international V-PDB (Vienna PeeDee Belemnite) standard. During the 220 isotopic analysis, the SIF lab used different certified reference materials (e.g., IAEA-600, USGS-40, USGS-41, USGS-42, USGS-43, USGS-61, USGS-64, and USGS-65) for the analytical control 221 222 quality. The analytical uncertainties reported for the SIF lab were 0.2‰ for  $\delta^{13}$ C 223 (https://stableisotopefacility.ucdavis.edu/13cand15n.html). We also included triplicate aliquots of 224 several specimens of the same species and condition, collected from one patch, or attached to the 225 same substrate, to assess the method error by sampling and processing procedural. The

226 methodological uncertainties were <0.4‰.

# 227 **2.3.** Analysis of $\delta^{13}$ C-macroalgal variability

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229 genus, and species) and morpho-functional groups (e.g., thallus structure, growth form, branching 230 pattern, and taxonomic affinities; Balata et al. 2011; Ochoa-Izaguirre and Soto-Jiménez, 2015). 231 The carbon fixation strategies in the macroalgae communities of the GC were identified by  $\delta^{13}$ C 232 (Hepburn et al., 2011; Díaz-Pulido et al., 2016), in agreement with the Maberly et al. (1992) and 233 Raven et al. (2002a) thresholds. So, macroalgae were classified into four strategies for DIC uptake: 234 1) CCM-only by active uptake HCO<sub>3</sub><sup>-</sup> ( $\delta^{13}$ C>-10‰), 2) CCM active uptake HCO<sub>3</sub><sup>-</sup> and/or diffusive uptake CO<sub>2</sub> ( $\delta^{13}$ C<-11 to -30‰), 3) Non-CCM, CO<sub>2</sub> by diffusion only ( $\delta^{13}$ C<-30‰), 4) Calcifying 235 with different carbon-use strategies related to different modes of calcification. The measured  $\delta^{13}$ C-236 237 macroalgal signals are integrative of the discrimination by photosynthesis ( $\Delta^{13}C_p$ ) on the carbon 238 source ( $\delta^{13}$ C-DIC in seawater), respiration ( $\Delta^{13}$ C<sub>r</sub>), and probable CO<sub>2</sub> leak out inside the cell 239 during the CCM process (Sharkey and Berry, 1985; Raven et al., 2005; Carvalho et al., 2009a,b). 240 Macroalgae were grouped according to their characteristics morpho-functional proposed initially by

The variability of  $\delta^{13}$ C values in macroalgae was analyzed in function of the taxonomy (phylum,

Littler and Littler (1980) and modified by Balata et al. (2011). Most of the macroalgae species showed a limited distribution along the GC coastlines. Few cosmopolites' species included *Colpomenia tuberculata, Sargassum sinicola, Padina durvillei*, and *Ulva lactuca*. Also, not all morphofunctional groups and taxon were present in every site during each sampling survey, and the sample size in each group varied for taxa, location, and time.

A basic statistical analysis of  $\delta^{13}$ C values in different macroalgae groups was applied to distribute and calculate the arithmetic mean, standard deviation, minimum and maximum. Because not all macroalgal species were present in sufficient numbers at different collection habitats, several macroalgal groups were not considered for statistical analysis. We compared among taxon and morphofunctional groups, collected in the same habitat (within-subjects factor) by multivariate analysis of variance. When differences were noted, a Tukey-Kramer HSD (Honestly Significant Difference) test was performed. Besides, variations of  $\delta^{13}$ C macroalgal in specimens of the same morpho-functional and taxon collected in different habitats were also investigated with a Kruskal-Wallis test.

The relationships between  $\delta^{13}$ C with each independent variable inherent to the macroalgae properties 255 256 (taxon and morphology), biogeographical collection zone (GC coastline and coastal sector), habitat 257 features (substrate, hydrodynamic, protection, and emersion level) and environmental conditions 258 (temperature, pH, and salinity), were examined through simple and multiple linear regression 259 analyses. Excepting temperature, pH, and salinity, most of the independent variables are categorical 260 independent variables. However, these continue variables were also categorized, such as previously 261 was described. Simple linear regression analyses were performed to establish the relationships between  $\delta^{13}$ C-macroalgal with each environmental parameter analyzed as possible driving factors 262 263 (e.g., temperature, salinity, and pH). Multiple linear regression analyses were conducted to evaluate 264 the combined effects of those independent variables (macroalgae properties, biogeographical collection zone, habitat features, and environmental conditions) on the  $\delta^{13}$ C-macroalgal. In the 265 266 multivariable regression model, the dependent variable,  $\delta^{13}$ C-macroalgal, is described as a linear function of the independent variables  $X_i$ , as follows:  $\delta^{13}$ C-macroalgal =  $a + b_1(X_1) + b_2(X_2) + ... +$ 267 268  $b_n(X_n)$  (1). Where a is regression constant (it is the value of intercept and its value is zero);  $b_1$ ,  $b_2$ , 269 and b<sub>n</sub>, are regression coefficients for each independent variable X<sub>i</sub>. From each one of the fitted 270 regression models, we extracted the estimated regression coefficients for each of the predictor 271 variables (e.g., Bayesian Information Criterion (BIC), Akaike Information Criterion (AIC), root-272 mean-square error (RMSE), Mallow's Cp criterion, F Ratio test, p-value for the test (Prob > F), coefficients of determination ( $R^2$ ) and the adjusted  $R^2$  statistics) (SAS Institute Inc., 2018). All 273 274 regression coefficients were used as indicators of the quality of the regression (Draper and Smith, 275 1998; Burnham and Anderson, 2002). Kolmogorov-Smirnov normality test was applied for all 276 variables, and all were normally distributed. Most of the  $\delta^{13}$ C values in each group showed a normal 277 distribution. For all statistical tests, a probability P<0.05 was used to determine statistical significance. The statistical analysis of the results was done using JMP 14.0 software (SAS Institute 278 279 Inc.).

**3. Results** 

## 281 **3.1. Taxonomy and morpho-functional groups**

282 Sampled specimens belong to three Phyla, 63 genera, and 170 species. The Phyla were identified as 283 Chlorophyta (25%), Ochrophyta (22%), and Rhodophyta (53%). The most representative genus (and 284 their species) were Ulva (U. lactuca, U. lobata, U. flexuosa, and U. intestinalis), Codium (C. 285 amplivesiculatum and C. simulans), Chaetomorpha (C. antenina), Padina (P. durvillei), Dictyota 286 (D. dichotoma), Colpomenia (C. tuberculata and C. sinuosa), Sargassum (S. sinicola and S. 287 horridum), Amphiroa (Amphiroa spp.), Spyridia spp, Polysiphonia spp., Gymnogongrus spp., 288 *Gracilaria* (*G. vermiculophylla*, *G. pacifica* and *G. crispata*), *Hypnea* (*H. pannosa* and *H. johnstonii*) 289 Grateloupia (G. filicina and G. versicolor), and Laurencia (L. papillosa and L. pacifica). The 290 endemic species included Chlorophyta Codium amplivesiculatum, Rhodophyta Laurencia papillosa, 291 Chondracanthus squarrulosa, Gracilaria spinigera, and G. subsecundata, and Ochrophyta Cutleria 292 hancockii, Sargassum herphorizum, S. johnstonii.

An analysis of the biogeographical diversity among sectors evidenced that P3 (43 genera of 63, 68%) and C3 (63%) at north recorded the highest number of the genus, followed by C1 (38%) and P1 (29%) at the south, and P2 (27%) and C2 (22%). The same pattern was observed in the species diversity, zones P3 (94 of 167 species, 56%) and C3 (52%) at the north, C1 (34%) and P1 (25%) at the south, and C2 and P2 (19-20%) at the center.

- 298 The morphofunctional groups identified were 21. The most common were C-tubular (6 spp., n=69;
- 299 C-Blade-like (6 spp, n=55); C-Filamentous uniseriate (17 spp, n=49); C-Erect thallus (5 spp, n=33);
- 300 O-Compressed with branched or divided thallus (19 spp., n=92); O-Thick leathery macrophytes (12
- spp., n=104); O-Hollow with spherical or subspherical shape (4 spp, n=87); R-Large-sized corticated
  (57 spp., n=225); R-Filamentous uniseriate and pluriseriate with erect thallus (9 spp., n=48); and RLarge-sized articulated corallines (6 spp, n=17). The diversity, in terms of presence/absence of the
  morphofunctional groups, varied among coastline sectors, higher in C3 (16 of 21, 76%) and P3
  (71%) at the north, followed by C1 (57%) and P1 (48%) at the south, and C2 and P2 and (42-48%)
  at the center of both GC coastlines.

# 307 **3.2.** $\delta^{13}$ C-macroalgal variability in function of taxonomy and morpho-functional groups

The variability of  $\delta^{13}$ C values in macroalgae was analyzed by taxon (phylum, genus, species) and morphofunctional groups classified by habitat, coastal sector, and collection season. A complete list of the results of  $\delta^{13}$ C in 170 macroalgae species is provided in Supporting Information (Table SI-1). Firstly,  $\delta^{13}$ C values analyzed by phylum showed a unimodal distribution with a peak at -14±1.4‰ (Fig 2). Ochrophyta (-21.5 to -2.2‰, -12.5±3.7‰), displayed significantly higher values than Chlorophyta (-25.9 to -5.5‰, -14.5±3.0‰) and Rhodophyta (-34.6 to -4.5‰, -14.8±3.9‰). The  $\delta^{13}$ C-macroalgal values (average±SD) for the genus of Chlorophyta, Ochrophyta, and Rhodophyta 315 (Fig. 3) varied from -33.8±1.1‰ for Schizymenia to -7.8±0.7‰ for Amphiroa. Based on the highest 316 values, specimens of three Phyla showed  $\delta^{13}$ C values>-10‰, evidenced the presence of CCM's by 317 active uptake of HCO<sub>3</sub><sup>-</sup> (strategy 1) (Fig. 3). For example, *Caulerpa*, *Cladophora*, *Codium*, *Ulva* for 318 Chlorophyta Colpomenia, Dictyota, Padina, Sargassum for Ochrophyta, and Hypnea and 319 *Polysiphonia* for Rhodophyta showed  $\delta^{13}$ C values >-10‰. Likewise, high  $\delta^{13}$ C values were observed 320 in the calcifying macroalgae genus Amphiroa and Jania, under strategy 4 (Fig. 3c).  $\delta^{13}$ C values 321 lower than -30‰ that denote uptake of CO<sub>2</sub> by diffusion (strategy 3), were observed only in 322 Rhodophyta Schizymenia, Halvmenia, and Gigartina. However, most species showed large  $\delta^{13}$ C 323 variabilities that evidence a mechanism that uses a mix of HCO<sub>3</sub><sup>-</sup> and CO<sub>2</sub> for photosynthesis 324 (strategy 2).

Multiple comparison analyses revealed significant differences in the  $\delta^{13}$ C-macroalgal values among genera, ordered as *Schizymenia* < *Polysiphonia* < *Ulva*, *Gracilaria* and *Spyridia* (-16.1±0.6‰ to -15.1±0.2‰) < *Gymnogongrus*, *Laurencia*, *Hypnea*, *Cladophora*, *Dictyota*, *Sargasumm*, *Chaetomorpha*, and *Grateloupia* (from -15.4±0.7‰ to -13.8±0.8‰) < *Codium* and *Padina* (-12.5±2.4‰ to -12.4±2.5‰) < *Colpomenia* and *Amphiroa* (-9.2±0.3 to -7.8±0.7‰) (F=16.81, p<0.001).

Aggrupation of  $\delta^{13}$ C values based on morpho-functional features is displayed in Fig. 4. The most representative groups in the phylum Chlorophyta varied from -15.8±0.3‰ for C-Tubular to -12.4±0.5‰ for C-Thallus erect. The phylum Ochrophyta includes O-Thick leathery with the lowest mean (-14.8±0.3‰) and O-Hollow with a spherical or subspherical shape with the highest values (-9.2±0.3‰). The lowest and highest  $\delta^{13}$ C values for Rhodophyta were observed for R-flattened macrophytes (-24.0±9.6‰) and R-Larger-sized articulated coralline (-7.89±0.75‰), respectively. Significant differences were observed among groups, which were ordered as follows: R-Flattened 338 macrophytes < R-Blade like < C-Tubular < O-Tick leathery and R-Large size corticated < C-Blade

339 like and C-Filamentous uniseriate < C- Thallus erect and O-Compressed with branch < O-Hollow

- 340 with spherical < R-Larger-sized articulated coralline.
- High intraspecific variability in  $\delta^{13}$ C signal for the more representative genera of each taxon is 341 342 showed in Table 1-3. For Codium, C. brandegeei (11.8±1.2‰) and C. simulans (-11.4±2.2‰) 343 showed higher  $\delta^{13}$ C values than C. amplivesculatum (-14.4±2.7‰). Colpomenia species had higher 344  $\delta^{13}$ C values than the other genera, with higher values for C. tuberculata (-8.7±3.2‰) than 345 Colpomenia sp. (-10.9±3.6‰) and C. sinuosa (-10.2±2.9‰). Gracilaria showed comparable  $\delta^{13}$ C 346 values in the four species (from  $-16.4\pm1.6\%$  for *G. pacifica* to  $-15.5\pm2.4\%$  for *Gracilaria* sp.). 347 *Hypnea* showed non-significant  $\delta^{13}$ C differences in three representative species (-16.4±1.7‰ for *H*. 348 spinella to -14.9±2.3‰ for Hypnea sp.). Laurencia sp. (-12.9±1.2‰) was higher than L. pacifica (-349 14.9±2.2‰), while Padina sp. (-11.1±1.5‰) higher than P. durvillei (-13.2±2.6‰). Sargassum was 350 one of the most diverse genera studied with six representative species, with  $\delta^{13}$ C values ordered as 351 follow: S. horridum = S. sinicola = S. johnstonii (-15.5 $\pm$ 2.9 to -15.1 $\pm$ 2.4‰) < S. lapazeanum (-352  $14.5\pm1.6\%$ ) = Sargassum sp. (-14.2±2.3‰) < S. herphorizum (-13.6±1.6‰). Spyridia sp. (-353 17.0±1.2‰) and S. filamentosa (-15.8±3.8‰) showed non-significant differences. The six 354 representative species of Ulva were divided into two morphological groups, filamentous and 355 laminates. Filamentous species that averaged -16.3±2.0‰ for U. clathrata, -16.0±3.6‰ for U. 356 *flexuosa*,  $-15.7\pm1.7\%$  for *U. acanthophora* and  $-15.3\pm2.5\%$  for *U. intestinalis* and *Ulva* laminates 357 that included U. linza (-15.5±2.4‰) and U. lactuca (-14.1±3.1‰). Non-significant differences were 358 observed between morphological groups and among species. A high intra-specific variability, 11-359 28%, explains average overlapping.
- **360 3.3.** δ<sup>13</sup>C-macroalgal variability in coastal sectors

361 A variety of macroalgal assemblages were documented along the GC coastlines, with differences in 362 the taxonomic composition according to their fico-floristic region. Multiple comparison analysis of 363  $\delta^{13}$ C signals evidenced significant differences between the most common genus and species of macroalgae between and within assemblages grouped by coastal sector, season and collecting year 364 365 (Supplementary Information Tables SI-2-3). For example, genus Padina (e.g., P. durvillei) and Ulva 366 (e.g., U. lactuca), collected in C1 sector during the rainy season, showed lower  $\delta^{13}$ C values than in 367 other sectors. Differences in the  $\delta^{13}$ C signal are mainly related to the carbon uptake strategies of the 368 macroalgae (Fig. 5). Even though most species inhabiting the GC coastal sectors displayed 369 domination of strategies based on active CCM's, but the tendencies were different between taxa and 370 coastal regions. The strategy 2 with mixing DIC sources are dominant in all regions and taxa (60-371 90%). Exceptions were observed in the P1 (68%) and C1 (37%) regions for Ochrophyta, where the 372 specialized strategy 1 of only  $HCO_3^-$  user was significant. The strategy 3 based on the use of  $CO_2$ 373 was observed in the peninsular coast in P2 and P3 for Rhodophyta with 2-3.3%. Overall, more negative  $\delta^{13}$ C values were observed at continental (C2) compared to the peninsular coastline (P1-374 375 P3) and southward than northward.

# 376 **3.4.** $\delta^{13}$ C-macroalgal variability in function of taxonomy and habitat features and

#### 377 environmental conditions

Variability of  $\delta^{13}$ C values for the most representative genera was evaluated by multiple comparative analyses in the habitat features' function, including the substrate, hydrodynamic, and emersion level. Large  $\delta^{13}$ C variability observed between specimens of the same genus collected in the different habits do not show any significant pattern, and non-significant differences were observed. An exception was observed with the emersion level (showed in Fig. 6), where intertidal specimens recorded lesser negative values than subtidal in most macroalgae genus. For example, for *Hydroclathrus* (intertidal -5.7 $\pm$ 0.9‰; subtidal -11.4 $\pm$ 5.9‰), *Amphiroa* (intertidal -6.9 $\pm$ 1.5; subtidal -9.9 $\pm$ 6.1), *Hypnea* (intertidal -13.5 $\pm$ 2.5‰; subtidal -18.6 $\pm$ 1.8‰), and *Laurencia* (intertidal -13.5 $\pm$ 1.3‰; subtidal -17.1 $\pm$ 1.8‰). Exceptions were observed for *Polysiphonia* (intertidal -19.7 $\pm$ 2.2‰, subtidal -14.9 $\pm$ 6.7‰), *Spyridia* (intertidal -16.9 $\pm$ 3.3‰, subtidal -13.2 $\pm$ 0.7‰) and *Colpomenia* (intertidal -9.4 $\pm$ 3.4‰, subtidal -7.7 $\pm$ 1.3‰).

- Non-significant differences were observed for the same genera at different temperatures ranges, except for *Grateloupia* (cold, -19.2±4.7‰, typical -14.4±2.2‰, warm -14.5±2.2‰) and *Polysiphonia* (cold, -21.0±0.4‰, typical -18.1±5.5‰, warm -17.9±2.3‰) with more negative values in colder than warmer waters (F=6.42, p<0.001). Neither significant difference was observed in  $\delta^{13}$ C
- 393 values in macroalgae specimens from the different genus in the same temperature range (Fig. 7a).

394 Significant differences were observed among genus related to the pH level at seawater (Fig. 7b). 395 Under typical pH seawater, Amphiroa and Colpomenia were 1-2‰ more negatives than in alkaline 396 waters, while Ulva and Spyridia were 3-5% less negative than in acidic waters. Amphiroa and 397 Colpomenia were not collected in acidic water, and neither Spyridia in alkaline waters to compare. 398 Another genus also showed extremes values between alkaline (Tacanoosca -7.6±1.0‰) and acidic 399 waters (Schizymenia -32.9±2.0‰). The following order was observed in the genus collect at the three 400 pH ranges: alkaline > typical > acidic. Significant differences were observed for genus Ahnfeltiopsis, 401 Caulerpa, Gymnogongrus, Padina, and Ulva, with higher values at alkaline than in acidic waters. 402 Values of  $\delta^{13}$ C for specimens of the same genus collected at typical pH waters are mostly overlapped 403 between those for alkaline and acidic seawaters. Non-significant differences in  $\delta^{13}$ C values were 404 observed for Grateloupia, Hypnea, and Polysiphonia concerning pH-type waters.

405 We analyzed the carbon uptake strategies on macroalgal assemblages in the function of

environmental factors like temperature, pH, and salinity (Fig. 8). Regarding the  $\delta^{13}$ C variability for 406 407 all data set in response to temperature and salinity, a non-significant trend was observed between  $\delta^{13}$ C-macroalgal in both parameters' function. A poor but significant correlation was observed 408 between  $\delta^{13}$ C and pH (R<sup>2</sup> = 0.04) (Table 4). The proportion of specimens with a strategy of only 409 410 HCO<sub>3</sub><sup>-</sup> use was different between environmental factors and taxa (previously described). For 411 example, Ochrophyta showed the highest proportion (35%) in colder temperature, in pH-Alkaline 412 (31%), and at typical salinity regimen (27%), while Chlorophyta enhanced to 30% in acid pH and Rhodophyta recorded 21% at normal seawater. The opposite strategy (only use of dissolved CO<sub>2</sub>) 413 414 was observed only in Rhodophyta. The highest percentage was observed in estuarine salinity 415 regimen (10%).

# 416 **3.5. Variation latitudinal of** $\delta^{13}$ **C-macroalgal**

The  $\delta^{13}$ C-macroalgal variation in the GC biogeography was evaluated by linear regression analysis between  $\delta^{13}$ C values along the nine degrees latitude in both GC coastlines. A non-significant latitudinal trend was observed for datasets, but for the three Phyla's most representative genera,  $\delta^{13}$ C values correlated with latitude (Fig. 9). In Chlorophyta, with the higher genera number,  $\delta^{13}$ C values increased with latitude, with low but significant correlation. Contrarily, in Ochrophyta and Rhodophyta specimens, the  $\delta^{13}$ C values decreased non-significantly with latitude.

In the most representative morphofunctional groups, significant correlations (p<0.001) were observed for  $\delta^{13}$ C-macroalgal *versus* latitude (Fig. 10). Representative morphofunctional groups of Chlorophyta (e.g., C-Tubular, C-Filamentous uniseriate), showed a positive correlation, while those belonging to Ochrophyta (e.g., O-Thick leathery;) and Rhodophyta (e.g., R-Large sized corticated) showed a negative trend with latitude.

# 428 **3.6.** Analyses of $\delta^{13}$ C macroalgal variability

429 An analysis of the effects, independent and combined, on the  $\delta^{13}$ C-macroalgal variability related to 430 life form and environmental factors was conducted. Firstly, simple linear regression analyses were 431 performed to evaluate the dependent variable's prediction power ( $\delta^{13}$ C-macroalgal) in the function 432 of several independent variables controlling the main macroalgae photosynthesis drivers (light, DIC, 433 and inorganic nutrients). Regression coefficients were estimated for each fitted regression model, 434 which is used as indicators of the quality of the regression (Draper and Smith, 1998; Burnham and 435 Anderson, 2002) as was described in Methods; however, our results description focused on the coefficients of determination ( $R^2$  and adjusted  $R^2$ ). The coefficient  $R^2$  describes the overall 436 relationship between the independent variables  $X_i$  with the dependent variable Y ( $\delta^{13}$ C-macroalgal), 437 and it is interpreted as the % of contribution to the  $\delta^{13}C$  variability. In comparison, the adjusted  $R^2$ 438 439 statistics compensate for possible confounding effects between variables.

440 Results of the analysis of the relationships between  $\delta^{13}$ C with each independent variable are 441 summarized in Table 4. Regarding the inherent macroalgae properties, Phyla explain only 8% 442 variability, the morphofunctional properties 35%, genus 46%, and species 57%.

In terms of coastline (continental vs. peninsular) and coastal sectors (C1-C3 and P1-P3), the biogeographical collection zone explained a maximum of 5% variability. Only the emersion level (6%) contributed to the  $\delta^{13}$ C variability related to the habitat features. The contribution of the seawater's environmental conditions was marginal for pH (4%) and negligible for temperature and salinity. A marginal reduction in the percentage of contribution was observed for Phyla (1%) and morphofunctional properties (1%), but significant for genus (5%) and species (10%).

449 Multiple regression analyses were also performed to interpret the complex relationships among  $\delta^{13}$ C-

450 macroalgal, considering the life form (morphofunctional and taxon by genus) and their responses to

451 environmental parameters. Results for the fitted regression models performed for morphofunctional 452 groups (Table 5) and genus (Table 6) evidenced that the effect of the coastal sector and pH ranges 453 on the  $\delta^{13}$ C-macroalgal increased the contribution by 9-10% each one. The emersion level increased 454 by 5-6%, the contribution respect to individual effect of morphofunctional group and genus, the 455 temperature and pH in 1 and 3%, respectively, while salinity decreased by 1-2%. Adding the effect of the biogeographical collection zone, represented by the coastline sector, to those for 456 457 morphofunctional group (Table 5) and genus (Table 7), a notable increase of 11-12% was observed. 458 Considering the combined effect of the coastline sector + Habitats features for Morphofunctional group or Genus (Table 7), the full model showed  $R^2$  of 0.60 and 0.71. In contrast, Coastline sector + 459 460 Environmental conditions + Morphofunctional group or Genus the  $R^2$  increased to 0.62 and 0.72, 461 respectively. The interactive explanations of environmental factors increased the explanation 462 percentage of  $\delta^{13}$ C variability; however, these contributions were significantly lower than the 463 explained by life forms, such as the morphofunctional properties and taxa by genus and species.

The combined effect of environmental conditions on the  $\delta^{13}$ C variability was tested for the best-464 465 represented genus and morphological groups. Results evidenced that 9 of 21 morphological groups 466 showed significant effects on the  $\delta^{13}$ C variability (Table 8), five increasing and four decreasing the model constant of  $\delta^{13}$ C=-14.2‰. For example, for the O-Hollow with spherical or subspherical shape 467 468 (+4.9‰) and R-Larger-sized articulated corallines (+6.3‰), the predicted values are -7.9±0.8‰ and 469 -9.2±0.4‰. For R-Filamentous uniseriate and pluriseriate with erect thallus (-2.1‰) and C-Tubular 470 (-1.6‰), the predicted values are -16.3±0.5‰ and -15.8±0.5‰, respectively. Regarding taxon, a 471 significant effect was observed only in 13 genera, including Colpomenia (+5.4‰), Amphiroa 472 (+6.8‰), and Padina (+2.2‰) increasing the signal, and Polysiphonia (-3.7‰), Gracilaria (-0.9‰), and Spyridia (-1.4‰) decreasing the signal of the model constant (Table 9). In 33 species was 473 observed a significant effect on the  $\delta^{13}$ C variability, including C. tuberculate +5.9‰, C. sinuosa 474

475 +4.4‰, *H. pannosa* +4.4‰, *H. johnstonii* +4.4‰, and *Amphiroa spp.* (+4.4 to 8.2‰) increasing the

476 model constant  $\delta^{13}C = -14.6\%$ , and *Spyridia* sp. (-2.5‰), *G. filicina* (-2.3‰), *P. mollis* (-5.2‰) and

477 *S. pacifica* (-19.2‰) (Table 10).

478

# 479 **3.7.** Preliminary estimations of $\Delta^{13}$ C-macroalgal

480 Concurrent analysis of surface seawater for alkalinity, proportions of the chemical species of DIC (CO<sub>2</sub>, HCO<sub>3</sub><sup>-</sup>, and CO<sub>3</sub><sup>2-</sup>), and  $\delta^{13}$ C-DIC evidenced that  $\delta^{13}$ C-DIC in GC seawater averages 481 482 1.4±0.4‰ (-1 to 4.9‰) (Supplementary Information Fig. SI-1). In our preliminary data, the  $\delta^{13}$ C-483 DICseawater slightly (in 0.5%) decreased during the rainy season in those zones influenced by river 484 discharges along the continental coastline. Non-significant differences were observed among coastal sectors.  $\delta^{13}$ C-DIC values in GC seawater are comparable to the averages 1.4-1.6% reported for the 485 486 surface seawaters in the Eastern North Pacific in the 1970s-2000s (Quay et al., 2003; Hinger et al., 487 2010; Santos et al., 2011).

488 Based on the subtraction of  $\delta^{13}$ Cmacroalgae to  $\delta^{13}$ C-DICseawater, the integrative discrimination

489 factor against <sup>13</sup>C averaged 16.0±3.1‰, 16.8±4.3‰, and 14.0±3.8‰ for Phyla Chlorophyta,

490 Rhodophyta, and Ochrophyta, respectively. Five groups were identified in the function of the  $\Delta^{13}$ C

491 values, one for Chlorophyta ( $\Delta^{13}C=16.0\pm3.1\%$ ), two for Rhodophyta ( $16.6\pm3.8\%$  and  $34.6\pm1\%$ ),

492 and two for Ochrophyta (9.1±1.7‰ and 15.7±2.7‰) (Fig. S2). Values of  $\Delta^{13}$ C were comparable to

493  $\delta^{13}$ C of the thallus of macroalgae. Thus,  $\delta^{13}$ C-macroalgal reflect mainly the discrimination during

494 carbon assimilation. Like  $\delta^{13}$ C-macroalgal, the  $\Delta^{13}$ C values were subject to considerable variation.

495

#### 496 **4. Discussions**

# 497 **4.1. Explaining the \delta^{13}C macroalgal variability**

498 This study revealed high variability in the  $\delta^{13}$ C of the large inventory of macroalgae collected along 499 GC coastline between five years period. A linear regression analysis of the effects of life forms 500 revealed that the  $\delta^{13}$ C variability in the -macroalgal community is mainly explained by taxonomic 501 (genus 46%, species 57%) and morphofunctional groups. This result is consistent with the report of 502 Lovelock et al. (2020), who found that 66% of  $\delta^{13}$ C variability was explained by taxonomy. Even 503 so, the variability associated with each genus is not the same and can be classified in three groups: 504 1) high variability (e.g., *Schizymenia* = $\pm 19.1\%$ ), moderate variability (e.g., *Hydroclathrus*= $\pm 7.3\%$ ; 505 Amphiroa = $\pm 6.8\%$ ) and low variability (e.g., Gracilaria = $\pm 0.89$ ; Spyridia = $\pm 1.46\%$ ). The observed 506  $\delta^{13}$ C variability in this study is comparable with those reported in the literature, compiled in Table 507 SI-4.

508 Most authors studying the isotopic composition of C in macroalgae have reported the high isotopic 509 variability, which has been attributable to the taxon-specific photosynthetic DIC acquisition 510 properties (Raven et al., 2002a, Mercado et al., 2009, Marconi et al., 2011, Stepien, 2015, Díaz-511 Pulido et al., 2016; Lovelock et al., 2020). Our study observed that the intrinsic characteristics of 512 each morpho-functional group of macroalgae (e.g., thallus structure, growth form, branching pattern, 513 and taxonomic affinities) also influence the  $\delta^{13}$ C-macroalgal signals. The thallus thickness, a 514 morphology propriety, influences the diffusion boundary layer on the surface of the macroalgal, 515 where they carry out the absorption of essential ions and dissolved gases (Hurd, 2000; San-Ford and 516 Crawford, 2000). Thus, morphology can modulate the photosynthesis rates. However, a non-517 biological or ecological explanation of the  $\delta^{13}$ C variability, and therefore carbon use physiology, can 518 be given in terms of morphology.

519 The  $\delta^{13}$ C-macroalgal depends on the carbon source ( $\delta^{13}$ C-DIC in seawater), the isotope 520 discrimination during carbon assimilation in the photosynthesis ( $\Delta^{13}C_p < 29\%$  in a variable degree), and the plant respiration ( $\Delta^{13}C_r$  average  $\pm 2.3\%$ ) (Carvalho et al., 2009a,b, 2010; Carvalho and Eyre, 521 2011, Rautenberger et al., 2015). Comparatively, the  $\Delta^{13}C_r$  value is relatively small regarding  $\Delta^{13}C_p$ . 522 Thus,  $\delta^{13}$ C-macroalgal is an integrative value of the isotope discrimination during DIC seawater 523 assimilation [ $\Delta^{13}$ C= ( $\delta^{13}$ C-DIC seawater –  $\delta^{13}$ Cmacroalgae)] (Carvalho et al., 2009a). Based on the 524  $\Delta^{13}$ C values, five groups were identified in our study: one for Chlorophyta ( $\Delta^{13}$ C=16.0±3.1‰), two 525 526 for Rhodophyta ( $16.6\pm3.8\%$  and  $34.6\pm1\%$ ), and two for Ochrophyta ( $9.1\pm1.7\%$  and  $15.7\pm2.7\%$ ). Values of  $\Delta^{13}$ C were comparable to  $\delta^{13}$ C of the thallus of macroalgae. Thus,  $\delta^{13}$ C-macroalgal reflect 527 528 mainly the discrimination during carbon assimilation. The  $\delta^{13}$ C-macroalgal values reflect the 529 discrimination during carbon assimilation attributable to the taxon-specific photosynthetic DIC acquisition properties.  $\Delta^{13}$ C-macroalgal variability, captured in the  $\delta^{13}$ C-macroalgal signals, is 530 531 related to the thickness of the boundary layer around the thallus (Raven et al. 1982), the leakage 532 during carbon uptake (Sharkey and Berry 1985, Maberly et al. 1992), photosynthetic intensity 533 (Wiencke and Fischer 1990, Kübler and Raven 1995, 1996), and respiration rates (Carvalho et al., 534 2010; Carvalho and Eyre, 2011, Rautenberger et al., 2015). All intrinsic properties are related to the 535 life form.

536 Many species that recorded high  $\delta^{13}$ C values (and low  $\Delta^{13}$ C values) were fleshy macroalgae that are 537 characterized to be bloom-forming macroalgae belonging to genera *Ulva, Gracilaria, Cladophora,* 538 *Spyridia*, and *Sargassum* (Páez-Osuna et al., 2013, Valiela et al., 2018). It is not surprising that 539 species with high photosynthetic activity and high relative growth rates (Hiraoka et al., 2020) have 540 high carbon demand that results in lower isotopic discrimination against <sup>13</sup>C (Cornelisen, et al., 2007; 541 Carvalho et al., 2010ab; Kübler and Dungeon, 2015; Rautenberger et al., 2015). Bloom-forming 542 macroalgae (e.g., Ulva, Gracilaria, Sargassum) have been remarks as facultative species capable of 543 switching from C3 to C4 pathway (Valiela et al., 2018). C4 pathway reduces photorespiration, the 544 antagonist process of RuBisCo, enhancing the DIC assimilation in 25-40% and increasing the  $\delta^{13}$ C 545 values (Ehleringer et al., 1991; Bauwe et al., 2010; Zabaleta et al., 2012). C4 pathway has more 546 energy investment in CCM's than in RuBisCo protein content than C3 pathway (Young et al., 2016). 547 Also, the reports of C4 or C4-like pathway features in algae have increased in the last years (Roberts 548 et al., 2007; Doubnerová and Ryslavá, 2011; Xu et al., 2012, 2013). For example, high activity of 549 keys enzymes of C4 metabolisms, such as pyruvate orthophosphate dikinase (PPDK), 550 phosphoenolpyruvate carboxylase (PEPC), and phosphoenolpyruvate carboxykinase (PCK), has 551 been described in many algae species. But the establishment of a true C4 pathway in marine algae is 552 not clear since the massive changes in gene expression patterns seem to be incomplete, and it is 553 suggested that many marine algae have high plasticity to use a combination of CCM to overcome Ci 554 limitations (Roberts et al., 2007; Doubnerová and Ryslavá, 2011; Xu et al., 2012, 2013). A Stepwise 555 model of the path from C3 to C4 photosynthesis is explained by Gowik and Westhoff (2011). More 556 research is required on this topic considering the increasing the frequency, intensity, and extension 557 of bloom-forming macroalgae events worldwide (Teichberg et al., 2010; Valiela et al., 2018) and in 558 México (Ochoa-Izaguirre et al. 2007; Ochoa-Izaguirre and Soto-Jiménez 2015; Páez-Osuna et al., 559 2017).

560 Changes in the habitat features and environmental conditions, such as light intensity and DIC 561 availability, influencing the growth rate and photosynthetic intensity, have a strong influence on  $\delta^{13}$ C 562 signal (Carvalho et al., 2007, 2009; Carvalho and Eyre, 2011; Stepien, 2015; Mackey et al., 2015; 563 Rautenberger et al., 2015). The light intensity is the external factor with more influence on the  $\Delta^{13}$ C-564 macroalgal due to the regulation of carbon assimilation intensity (Wefer and Killingley 1986, Cooper 565 and DeNiro 1989, Grice et al. 1996; Carvalho et al., 2009a,b). Experimental studies found the light 566 levels as a key factor affecting the  $\delta^{13}$ C values. For example, under saturating light conditions, *Ulva* 567 switched from a carbon uptake of HCO<sub>3</sub><sup>-</sup> and CO<sub>2</sub> to increased HCO<sub>3</sub><sup>-</sup> use (Rautenberger et al., 2015). 568 Furthermore, field studies have shown that species growing in low light habitats as deep subtidal 569 tend to have more negative  $\delta^{13}$ C values than those in higher light environments (Mercado et al., 570 2009; Hepburn et al., 2011; Marconi et al., 2011; Stepien 2015; Cornwall et al., 2015, Díaz-Pulido 571 et al., 2016). In this study, intertidal specimens recorded lesser negative values than subtidal in most macroalgae genus. However, our study did not record the vertical effect in the  $\delta^{13}$ C signal related to 572 573 the light limitation because only shallow habitats (non-light limited) were studied.

574  $\delta^{13}$ C-DICseawater is reasonably uniform in surface seawater (-4.8 to 3.6%, median 1.5%), with  $\delta^{13}$ C values for CO<sub>2</sub>, HCO<sub>3</sub><sup>-</sup>, and CO<sub>3</sub><sup>2-</sup> nearly -10, -0.5 and 2‰, respectively (Mook et al., 1974; 575 576 Kroopnick, 1985). Exceptions can be expected where variations in the salinity, alkalinity, and proportions of the chemical species of DIC (CO<sub>2</sub>, HCO<sub>3</sub><sup>-</sup> or CO<sub>3</sub><sup>2-</sup>) occur (e.g., in coastal 577 578 environments influenced by river and groundwater discharges) (Mook et al., 1974; Chanton and 579 Lewis 1999; Hinger et al., 2010; Carvalho et al., 2015). Regarding DIC sources for macroalgae in the GC surface seawater, the availability, chemical proportions, and  $\delta^{13}$ C-DIC, were also relatively 580 581 constant and uniform. Thus, the influence of the  $\delta^{13}$ C-DIC variations on the  $\delta^{13}$ C-macroalgal 582 variability is negligible in the GC.

The effect of other environmental factors, such as salinity and pH, on  $\delta^{13}$ C-macroalgal signals, were evaluated. Regarding salinity, the influence of freshwater discharge by rivers and groundwater decreases the  $\delta^{13}$ C signal, which could be explained by the reduction in the salinity regimen that follows a decrease in  $\delta^{13}$ C-DIC in water (Hinger et al., 2010; Santos et al., 2011). In our study, a non-significant correlation between  $\delta^{13}$ C-macroalgal and salinity was observed.

Based on pH, differences in  $\delta^{13}$ C were found only for a few genera (e.g., Amphiroa, Colpomenia, 588 589 *Ulva*, *Spyridia*), with a trend to increase in the  $\delta^{13}$ C values with pH increase, such as was reported 590 by Maberly et al. (1992) and Raven et al. (2002b). Similar results were reported for Cornwall et al. 591 (2017) in the field study, with the differential response of the  $\delta^{13}$ C signals to pH among 19 species, 592 in which only four species were sensitive to pH changes. A very weak but significant positive linear 593 regression was observed between  $\delta^{13}$ C and pH. Also, a trend to decrease in the  $\delta^{13}$ C was recorded in 594 the following order: alkaline > typical > acidic. According to Stepien (2015), the result of metaanalyzes between pH drift experiments and  $\delta^{13}$ C thresholds was positive only for Rhodophyta and 595 596 Ochrophyte but not for Chlorophyta. About 86% of the Stepien metadata met the theoretical CCM 597 assignation based on both parameters, exceptions for species with  $\delta^{13}$ C<-30‰ that have been capable 598 of raising pH>9. A strong association between pH compensation point and  $\delta^{13}$ C was reported by 599 Iñiguez et al. (2009) in three taxa of polar macroalgae. Environmental conditions may influence the 600  $\delta^{13}$ C-macroalgal values but not change the carbon use physiology in the macroalgae, which is most likely inherently species-specific.4.2. Using  $\delta^{13}$ C-macroalgal to indicate the presence of an active 601 602 CCM

603 In our study, the  $\delta^{13}$ C signals from the thallus of macroalgae were used to evidence the presence of 604 an active CCM. This tool was first used in macroalgal shallows communities of the Gulf of California. Most macroalgae species displayed  $\delta^{13}$ C values that exhibit an active CCM's. Then, 605 606 macroalgae were classified into four strategies for DIC uptake, in agreement with the Maberly et al. 607 (1992) and Raven et al. (2002) thresholds: 1) CCM-only by active uptake HCO<sub>3</sub><sup>-</sup> ( $\delta^{13}$ C>-10‰), 2) CCM active uptake HCO<sub>3</sub><sup>-</sup> and/or diffusive uptake CO<sub>2</sub> ( $\delta^{13}$ C<-11 to -30‰), 3) Non-CCM, CO<sub>2</sub> by 608 609 diffusion only ( $\delta^{13}C < 30\%$ ), 4) Calcifying with different carbon-use strategies related to different 610 modes of calcification. About of 84% of the total analyzed specimens showed the facultative uptake

611 of  $HCO_3^-$  and  $CO_2$ , the most common strategy identified in macroalgal shallow communities 612 (Hepburn et al., 2011; Cornwall et al., 2015; Stepien 2015; Díaz-Pulido et al., 2016). Based on the 613 carbon uptake strategies, the most abundant macroalgae were those able to use both  $HCO_3^-$  and/or 614  $CO_2$  by means of active uptake plus passive diffusion (strategy 2).

615 Macroalgae collected in GC also involved those that are only HCO<sub>3</sub><sup>-</sup> users (strategy 1) and those 616 relying on diffusive  $CO_2$  uptake (strategy 3). Photosynthesis that relies on  $CO_2$  uptake (lack of 617 CMM), the most primitive mechanism (Cerling et al., 1993), has fewer energy costs than  $HCO_3^{-1}$ 618 uptake, which requires complex machinery with a high operational cost (Giordano et al., 2005; 619 Hopkinson et al., 2011; Hopkinson, 2014; Raven and Beardall, 2016). The energy for macroalgae to 620 uptake HCO<sub>3</sub><sup>-</sup>, cross the plasma membrane, and covert to CO<sub>2</sub> for photosynthesis, is obtained through 621 irradiance (Cornelisen et al., 2007). Based on our sampling effort, focused on intertidal and shallow 622 subtidal habitats featured by high-light intensities, we expected high proportions of species with the 623 carbon uptake strategy that uses only  $HCO_3^{-}$ . Results evidenced that strategy 1 was recorded in 624 specimens belonging to 58 species of 170 total species. The higher proportions of CCM species 625 (HCO<sub>3</sub><sup>-</sup> users) with high-energetic requirements are explained by those elevated irradiances 626 (Hepburn et al. 2011; Cornwall et al. 2015). Ochrophyta showed the highest proportion of species 627 that depend only on  $HCO_3^-$  uptake on both coastlines in the southern region of GC (P1, C1). Low 628 solubility of CO<sub>2</sub> is related to high temperatures in subtropical waters (Zeebe and Wolf-Gladrow, 629 2007) that impulse the development of CCM (Raven et al., 2002b) and by the high affinity to DIC 630 by Ochrophyta, such as has been described before by Diaz-Pulido et al, (2016).

631 Only three non-calcifying species (Schizymenia pacifica, Halymenia sp., Gigartina sp.) belonging

to Rhodophyta were CO<sub>2</sub> exclusive users ( $\delta^{13}C=-33.2\pm1\%$ ). Based on measurements of pH drift,

633 Murru and Sandgreen (2004) reported Schizymenia pacifica and two species of Halymenia (e.g., H.

634 schizymenioides and H. gardner) as restricted CO<sub>2</sub> users. Measurements of  $\delta^{13}$ C in Halymenia 635 *dilatate* confirmed the CO<sub>2</sub>-restricted photosynthesis in specimens collected offshore in deep reefs 636 of the Great Barrier reef (Díaz-Pulido et al., 2016). Red macroalgae that lack CCM, tend to inhabit low-light habitats like subtidal or low intertidal and are abundant in cold waters (Kübler et al., 1999; 637 638 Raven et al., 2002a; Cornwall et al., 2015). According to these authors, approximately 35% of the 639 total red algae tested globally are strictly CO<sub>2</sub> dependents. The percentage of macroalgae species 640 representative of Arctic and Antarctic ecosystems that lack CCM is 42-60% (Raven et al., 2002b; 641 Iñiguez et al., 2019), 50% for temperate waters of New Zealand (Hepburn et al., 2011), and up to 642 90% found for a single site of Tazmania, Australia (Cornwall et al., 2015). Our study sampled 91 643 red macroalgae species (of 453 red macroalgae species reported in the GC, Pedroche and Sentíes, 644 2003), of which <3% were CO<sub>2</sub> dependents. This low percentage could be related to the fact that 645 deep habitats (>2 m depth low tide) were not explored in our surveys.

646 Few calcifying macroalgae species using  $HCO_3^-$  and  $CO_2$  (strategy 4) were also collected, including 647 the genera Amphiroa (-7.8±3.7‰) and Jania (-9.4±0.7‰), both Rhodophyta with articulated-form. 648 *Padina*, a genus with less capacity to precipitate CaCO<sub>3</sub> (Ilus et al., 2017), displayed relatively high 649  $\delta^{13}$ C values (-12.5±2.4‰), suggesting the presence of CCM using HCO<sub>3</sub><sup>-</sup>. Some species of Padina 650 can use HCO<sub>3</sub>, but their efficiency may differ from species to species (Raven et al., 2002a; Enríquez 651 and Rodríguez-Román, 2006). Three genera are widespread in the GC. Stepien (2015) reported a 652 global mean of -14.8±1.0‰ for calcifying species compared to -20.1±0.3‰ for non-calcifying 653 species. Calcifying species have a different carbon uptake strategy influenced by the calcifying process that results in elevated  $\delta^{13}$ C signals (Diaz-Pulido et al., 2016). One possibility for high  $\delta^{13}$ C 654 655 values for calcifying species are related to the excess of H<sup>+</sup> released as residuals products of the 656 calcifying process. Also, the acidified boundary layers benefit the HCO<sub>3</sub><sup>-</sup> uptake (McConnaughey and Whelan 1997, Courneau et al., 2012). Another possibility to explain high  $\delta^{13}$ C values can also be related to the highly efficient light properties enhanced by the carbonate skeleton, resulting in an optimization of photosynthetic activity (Vásquez-Elizondo et al., 2017). Hofmann and Heesch (2018) reported high  $\delta^{13}$ C values in eight rhodoliths species (calcifying species) for the organic matter thallus and for thallus, including CaCO<sub>3</sub> structure collected in deep habitats (25-40 m) where light availability is limited. Because the ocean acidification in progress, negative impacts are expected on calcifying organisms, more attention as ecological sentinels is warranted in the GC.

664 Measurements of  $\delta^{13}$ C signals evidence the presence or absence of CCMs in macroalgae and are 665 indicative of carbon use physiology (Giordano et al., 2005), however, the isotopic signature may be 666 inconclusive in the determination of the efficient use of one or more DIC species (CO<sub>2</sub> and/or HCO<sub>3</sub><sup>-</sup> 667 ) (Roleda and Hurd, 2012). The preferential DIC uptake of macroalgae is assessed by pH drift 668 experiments (Hepburn et al., 2011; Roleda and Hurd, 2012; Fernandez et al. 2014, 2015; Narvarte et 669 al., 2020) and it can be determined by simultaneously measuring the  $CO_2$  uptake and  $O_2$  production 670 rates using membrane-inlet mass spectroscopy (MIMS) (Douchi et al., 2019; Burlacot et al., 2020). Macroalgae that are unable to raise the seawater pH>9.0 are primarily CO<sub>2</sub>-users, while those that 671 672 can raise the seawater pH>9.0 (absence of CO<sub>2</sub>) are HCO<sub>3</sub>-users (Roleda; Hurd, 2012). Those 673 differences in the carbon uptake strategies can be easily deduced by pH drift experiments, which 674 were not done in our study but reported in the literature (Supplementary Information Table SI-4). 675 Also, the change in  $\delta^{13}$ C signature within the range specific to a carbon use strategy (e.g., mix HCO<sub>3</sub>/CO<sub>2</sub>-user) can be complemented by simultaneous measurements of O<sub>2</sub> and CO<sub>2</sub> produced 676 677 and consumed, respectively, during the photosynthetic using MIMS. For example, photosynthetic 678 O<sub>2</sub> production in a certain macroalgae species with an active CCM preferring (e.g., CO<sub>2</sub>) is about ten 679 times higher than no active CCM (Burlacot et al., 2020).

680 Based on the  $\delta^{13}$ C values, it is possible to assume that at least one basal CCM is active, however, it 681 is impossible to discern what type of CCM is expressed in the organisms (e.g., direct HCO<sub>3</sub><sup>-</sup> uptake 682 by the anion-exchange protein AE; Drechsler and Beer 1991; Drechsler et al. 1993) or types of 683 mitochondrial carbonic anhydrase (e.g., internal and external) that enhance the fixation of Ci by 684 recycling mitochondrial CO<sub>2</sub> (Bowes, 1969; Zabaleta et al., 2012; Jensen et al., 2020). Also, the co-685 existence of different CCMs has been described for the same species (Axelsson et al., 1999, Xu et 686 al., 2012), even that different CCM's can operate simultaneously, generating different Ci 687 contributions to RuBisCo internal pool (Rautenberger et al., 2015). The variety of CCMs and their combinations could contribute to the high  $\delta^{13}$ C variability for the same species. In our field study, it 688 689 is impossible to explain the variations of  $\delta^{13}$ C or  $\Delta^{13}$ C-macroalgal relative to CCM or CA activity 690 types. Controlled experiments, like those conducted by Carvalho and collaborators (e.g., Carvalho 691 et al. 2009a,b, 2010), are required to obtain this knowledge. 4.3. Variability of  $\delta^{13}$ C macroalgal

# 692 between the GC bioregions

693 Changes in the  $\delta^{13}$ C signal with latitude, mainly related to the light and temperature, have been 694 reported in the literature (Mercado et al., 2009; Marconi et al., 2011; Stepien, 2015; Hofmann and 695 Heesch, 2018; Lovelock et al., 2020). For example, a negative correlation between latitude and  $\delta^{13}$ C-696 macroalgal was described by Stepien (2015), concluding that the  $\delta^{13}$ C signal increased by 0.09‰ for 697 each latitude degree from the Equator. Hofmann and Heesch (2018) recently showed a strong decrease in latitudinal effect ( $R^2 = 0.43 \ \delta^{13}C_{total}$  and 0.13, for  $\delta^{13}C_{organic-tissue}$ , p=0.001) for rhodolite 698 699 of the northern hemisphere and macroalgae from coral reefs in Australia. In both cases, the latitude 700 range is higher than we tested (30° to 80° and from 10° to 45°, respectively). These differences on a 701 big scale tend to be associated with a temperature effect (Stepien, 2015) and their effect on CO<sub>2</sub> 702 solubility in seawater (Zeebe and Wolf-Gladrow, 2007). However, in our study, any geographical pattern in the  $\delta^{13}$ C macroalgal was observed. Our linear regression analyzes for latitudes showed a low but significant correlation for the dataset classified by morphofunctional groups and genus, negative in the cases of Rhodophyta and Ochrophyta groups, and positive for Chlorophyta.

706 Light is not limited along the GC latitudes. Most of the shallow habitats occupied by macroalgal 707 communities in the GC were high-light environments. In agreement to literature, the surface seawater 708 temperature across the GC vary in only 1°C annual mean (Escalante et al., 2013, Robles-Tamayo, 709 2018). However, larger temperature variations of 5-10°C were recorded in the coastal waters across 710 the GC bioregions in both climatic seasons. The combined effect of the coastline sector, habitats 711 feature, or environmental condition for Morphofunctional group or Genus explained 60-62 and 71-712 72% of the  $\delta^{13}$ C variability, respectively. Our analysis of variability for the best-represented 713 morphological groups (e.g., R-Filamentous uniseriate and pluriseriate with erect thallus and C-714 Tubular) and genus (e.g., Colpomenia, Padina, Polysiphonia and Gracilaria) revealed that certain 715 life forms are better monitors explaining the variability of  $\delta^{13}$ C-macroalgal (and  $\Delta^{13}$ C values) than 716 others. The  $\delta^{13}$ C variability in morphological groups refers to change within a specific carbon use 717 strategy, but not change in the carbon use physiology that is inherently species-specific. The biological or ecological relevance of the  $\delta^{13}$ C variability in function of the morphology, in terms of 718 719 the efficiency in the use of DIC and the isotope discrimination during carbon assimilation and 720 respiration, must be investigated in species of same genus morphologically different or between 721 same morphological structures belonging to a different taxon.

The proportion of specimens with different carbon uptake strategies also showed regional variations.

- For example, the facultative uptake of  $HCO_3^-$  and  $CO_2$  was dominant in the macroalgal shallow
- communities in the GC (60 to 90% of specimens), with an exception in the P1 region for Ochrophyta
- 725 where the specialized strategy of only  $HCO_3^-$  use dominated (68%), and high proportion were

726 observed in C1 with 37%. While the strategy based on only use of  $CO_2$  was observed in the 727 peninsular coast in P2 and P3 for Rhodophyta with 2-3.3%. Finally, the coastal sector C2 showed 728 more negative  $\delta^{13}$ C values in macroalgae specimens of the same genus compared to the peninsular 729 coastline (P1-P3). Small but detectable changes were observed in the Phyla distribution based on 730 environmental conditions. For example, Ochrophyta showed the highest proportion (35%) in colder 731 temperature, in pH-Alkaline (31%), and at typical salinity regimen (27%), while Chlorophyta 732 enhanced to 30% in acid pH and Rhodophyta recorded 21% at normal seawater. The opposite 733 strategy (only use of dissolved CO<sub>2</sub>) was observed only in Rhodophyta. The highest percentage was 734 observed in the estuarine salinity regimen (10%). Again, more research is required to obtain valuable 735 information on the physiological and environmental status of macroalgae.

736

## 737 **5.** Conclusions

738 In conclusion, we observed high  $\delta^{13}$ C-macroalgal variability in macroalgae communities in the Gulf 739 of California, such as reported in other worldwide marine ecosystems. The life form is the principal 740 cause of  $\delta^{13}$ C-macroalgal variability, which explains up to 57% of the variability. Changes in habitat 741 characteristics and environmental conditions also influence the  $\delta^{13}$ C-macroalgal variability within a 742 specific carbon use strategy. Considering the combined effect of the life form, coastline sector, and 743 environmental conditions, the full model explains up to 72% (genus) of the variability. The effect of 744 the coastal sector, pH ranges, and emersion level were significant, while for salinity and temperature, 745 negligible.

746 Most macroalgae inhabiting in GC displayed the presence of CO<sub>2</sub> concentrating mechanisms to

747 uptake HCO<sub>3</sub><sup>-</sup> for photosynthesis, 84% of the total analyzed specimens were able to use both HCO<sub>3</sub><sup>-</sup>

748 and/or CO<sub>2</sub> employing active uptake plus passive diffusion (strategy 2:  $-10 < \delta^{13}C > -30\%$ ). Specimens 749 belonging to 58 species of 170 total species showed carbon uptake strategy 1 that use only HCO<sub>3</sub><sup>-</sup>. 750 A higher proportion of CCM species (HCO<sub>3</sub><sup>-</sup> users) was expected because we focused on intertidal 751 and shallow subtidal habitats featured by high-light intensities. Only three non-calcifying species 752 (Schizymenia pacifica, Halymenia sp., Gigartina sp.) belonging to Rhodophyta (3%) were CO2 exclusive users (strategy 3:  $\delta^{13}C <-30\%$ ). The low percentage of CO<sub>2</sub> dependents versus 40-90% 753 754 reported for temperate regions could be related to the shallow habitat sampled in our surveys (<2 m 755 depth low tide). The calcifying macroalgae genera Amphiroa and Jania using HCO<sub>3</sub><sup>-</sup> and diffusive 756 CO<sub>2</sub> influenced by the calcification process (strategy 4) were present in the macroalgal communities 757 along the GC and high  $\delta^{13}$ C values. Because the ongoing ocean acidification, these calcifying 758 organisms constitute excellent ecological sentinels in the GC.

759 Finally, diverse authors have reported significant correlations between  $\delta^{13}$ C signal and latitude, 760 mainly related to the light and temperature. However, in our study's latitude range (21°-31°N), the 761 linear regression analyzes showed a low correlation for the  $\delta^{13}$ C-macroalgal dataset classified by 762 morphofunctional groups and genus, being negative for Rhodophyta and Ochrophyta and positive 763 for Chlorophyta. Non-clear  $\delta^{13}$ C-macroalgal patterns occur along the GC latitudes. However, 764 detectable changes were observed in the  $\delta^{13}$ C-macroalgal and the proportion of specimens with 765 different carbon uptake strategies among coastal sectors. For example, the facultative uptake of 766  $HCO_3^-$  and  $CO_2$  was dominant in the macroalgal shallow communities in the GC (60 to 90% of 767 specimens), but in the coastal sector P1 was the specialized strategy of only HCO<sub>3</sub><sup>-</sup> use the dominant 768 strategy (68%), and significant at C1 (37%).

769 Our research is the first approximation to understand the  $\delta^{13}$ C-macroalgal variability in one of the 770 most diverse marine ecosystems in the world, the Gulf of California. We did not pretend to resolve

771 the intricate processes controlling the variations of  $\delta^{13}C$  or  $\Delta^{13}C$ -macroalgal during carbon 772 assimilation and respiration and determine the isolated influence of each environmental factor. 773 Despite the large dataset and corresponding statistical analyses, our study faces limitations due to research design and because no research on  $\delta^{13}$ C-macroalgal analysis was developed previously in 774 775 the GC. The primary deficiency is the lack of pH drift experiments to discriminate  $\delta^{13}$ C signal 776 variations to the carbon uptake strategies to determine preferential DIC uptake of macroalgae ( $CO_2$ 777 or  $HCO_3^{-}$ ). The second limitation concerns the lack of controlled experiments to discern what type 778 of CCM is expressed in macroalgae (e.g., direct  $HCO_3^-$  uptake by the anion-exchange protein AE, 779 types of mitochondrial AC, or the co-existence of different CCMs). Also, more research is required 780 to assess the biological or ecological relevance of the  $\delta^{13}$ C variability in function of the morphology 781 (e.g., DIC uptake efficiency and isotope discrimination during carbon assimilation and respiration). 782 Future studies assessing the ability of macroalgae to use  $CO_2$  and/or  $HCO_3^-$  can be assessed by pH 783 drift experiments and/or MIMS in the cosmopolites' species and within of genus with differences in 784 the  $\delta^{13}$ C values between species (e.g., *Ulva* and *Sargassum*). Finally, controlled experiments in 785 laboratory and mesocosm type combined with field studies are required to elucidate what type of 786 CCM is expressed in macroalgae. Even so, the  $\delta^{13}$ C-macroalgal was a good indicator to infer the 787 presence or absence of CCM's and identify the macroalgae lineages that could be in a competitive 788 advantage based on their carbon uptake strategy and identify their geographical distribution along 789 with GC. Under the current conditions of climate change and their effects as ocean acidification 790 progresses and the bloom-forming macroalgae events increase in México and worldwide, the 791 analysis of  $\delta^{13}$ C-macroalgal constitutes an excellent tool to help to predict the prevalence and shift 792 of species in macroalgal communities' focused on carbon metabolism. However, to obtain the 793 maximum benefit from isotopic tools in the carbon-use strategies study, diverse and species-specific, 794 it is necessary to use them in combination with other techniques referred to herein.

### 795 **6. Data Availability Statement**

- 796 Data set are each permanently deposited Soto-Jimenez, Martin F; Velázquez-Ochoa, Roberto; Ochoa
- 797 Izaguirre, Maria Julia. Earth and Space Science Open Archive ESSOAr; Washington, Nov 25, 2020.
- 798 DOI:10.1002/essoar.10504972.1
- 799 https://search.proquest.com/openview/2060de58b217ca47495469b53ae2f347/1?pq-
- 800 origsite=gscholar&cbl=4882998

#### 801 **7. Author contribution**

- 802 Velázquez-Ochoa R. participate in the collection, processing, and analysis of the samples as a part
- 803 of his master's degree thesis. Ochoa-Izaguirre J. also participate in sample collections and
- 804 identified macroalgae specimens. Soto-Jiménez M.F. coordinated the research, was the graduate
- thesis director and prepared the manuscript with contributions from all co-authors.

#### 806 8. Competing interests

807 The authors declare that they have no conflict of interest.

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## 1229 Figure captions

1230 Fig. 1. Sites collection along the continental (C1-C3) and peninsula (P1-P3) Gulf of California

1231 coastlines (A), range of environmental factors supporting or limiting the life processes for the

- 1232 macroalgal communities within a habitat (B), and inserted Table with the features and
- 1233 environmental conditions in the diverse habitats in the GC bioregions that delimits the macroalgal
- 1234 community's zonation.
- 1235 Fig. 2. Variability of  $\delta^{13}$ C values for specimens of different macroalgae genera collected along GC
- 1236 coastlines classified by taxon: (A) Chlorophyta, (B) Ochrophyta and (C) Rhodophyta. Shaded
- 1237 background represents the cutoff limits for using CO<sub>2</sub> Only users and HCO<sub>3</sub><sup>-</sup> only users,
- 1238 respectively, according to Raven et al., (2002).
- 1239 Fig. 3. Variability of  $\delta^{13}$ C values for the genus collected along coastline of the Gulf of California
- 1240 according to their taxon: (A) Chlorophyta, (B) Ochrophyta and (C) Rhodophyta. Genus with n=1 is
- 1241 not shown, and genus n=2 was not considered to the statistical comparison. Different letters
- 1242 indicate significant differences (P<0.05): a>b>c>d>e. Shaded background represent the cutoff
- 1243 limits for using CO2 Only users and HCO3- only users, respectively, according to Raven et al.,
- 1244 (2002). For Chlorophyta: Bry= *Bryopsis*, Cau=*Caulerpa*, Cha= *Chaetomorpha*, Cla= *Cladophora*,
- 1245 Cod= Codium, Phy= Phyllodictyon, Str= Struveopsis, Ulv=Ulva. Phaeophyta: Col= Colpomenia,
- 1246 Dic= Dictyota, Ect= Ectocarpus, End= Endarachne, Hyd= Hydroclathatrus, Pad= Padina, Ros=
- 1247 Rosenvingea, Sar= Sargassum, Spa= Spatoglossum, Zon= zonaria. Rhodophyta: Aca:
- 1248 *Acantophora*, anf: *Anfeltiopsis*, Amp= *Amphiroa*, Cen= *Centroceras*, Cer<sup>1</sup>= *Ceramium*, Cer<sup>2</sup>=
- 1249 *Ceratodictyon*, Cho<sup>1</sup>= *Chondracanthus*, Cho<sup>2</sup>= *Chondria*, Das= *Dasya*, Dig= *Digenia*, Euc=
- 1250 Euchema, Gel= Gelidium, Gig= Gigartina, Gra<sup>1</sup>= Gracilaria, Gra<sup>2</sup>= Grateloupia, Gra<sup>3</sup>=

- 1251 Gracilariopsis, Gym= Gymnogongrus, Hal= Halymenia, Hyp= Hypnea, Jan= Jania, Lau=
- 1252 Laurencia, Lom= Lomentaria, Neo= Neosiphonia, Pol= Polysiphonia, Pri= Prionitis, Rho<sup>1</sup>=
- 1253 *Rhodoglossum*, Rho<sup>2</sup>= *Rhodymenia*, Sch= *Schizymenia*, Spy= *Spyridia*, Tac= *Tacanoosca*. Purple
- 1254 boxplots represent calcifying species group.
- 1255 Fig. 4. Variability of  $\delta^{13}$ C values for morphofunctional groups by taxa along coastline of the Gulf 1256 of California.
- 1257 Fig. 5 Proportion of species using different DIC sources according to their carbon uptake
- 1258 strategies: HCO<sub>3</sub><sup>-</sup> only users (CO<sub>2</sub> concentrating mechanism active), Users of both sources (HCO<sub>3</sub><sup>-</sup>
- 1259 & CO<sub>2</sub>) and CO<sub>2</sub> only users (non-CO<sub>2</sub> concentrating mechanism active) in function of coast along

1260 GC.

1261 Fig. 6. Variability of  $\delta^{13}$ C values in macroalgae specimens for the most representative genera in

1262 function of habitat features (emersion level). Green circles represent genus of Chlorophyta, Brown

- 1263 circles represent genus of Ochrophyta; red circles represent genus Rhodophyta and purple circles
- 1264 represent genus with calcifying capacity.
- 1265 Fig. 7. Variability of  $\delta^{13}$ C values in macroalgae specimens for the most representative genus in

1266 function of temperature (a) and pH (b) ranges in samples collected along Gulf of California

1267 coastline.

1269 strategies: HCO<sub>3</sub><sup>-</sup> only users (CO<sub>2</sub> concentrating mechanism active), Users of both sources (HCO<sub>3</sub><sup>-</sup>

1270 & CO<sub>2</sub>) and CO<sub>2</sub> only users (non-CO<sub>2</sub> concentrating mechanism active) in function of : (A) pH

1271 ranges, (B) temperature ranges and (C) salinity ranges.

<sup>1268</sup> Fig. 8. Proportion of species using different DIC sources according to their carbon assimilation

- 1272 Fig. 9. Trends in the  $\delta^{13}$ C-macroalgal in specimens collected along continental (C1-C3) and
- 1273 peninsula (P1-P3) Gulf of California coastline in function of latitudinal gradient.



































1297 Fig. 10

Species (n composite samples)	$\delta^{13}C\pm SD$ (Min to Max, ‰)
Chaetomorpha sp. (3)	-13.7±0.8 (-14.6 to -12.9)
C. antennina (10)	-14.6±1.1 (-16.3 to -12.8)
<i>C. linum</i> (5)	-16.8±1.6 (-18.4 to -14.6)
Codium sp. (5)	-11.6±3.0 (-14.1 to-6.7)
C. amplivesiculatum (8)	-14.4±2.7 (-20.4 to -11.3)
C. brandegeei (7)	-11.8±1.2 (-13.7 to -10.4)
C. fragile (4)	-13.0±2.7 (-14.8 to -9.0)
C. simulans (9)	-11.4±2.2 (-14.9 to -8.3)
<i>Ulva</i> sp. (12)	-14.0±3.9 (-19.2 to -7.1)
U. acanthophora (25)	-15.8±1.7 (-18.3 to -11.4)
U. clathrata (8)	-16.4± 2.0 (-20.5 to -14.5)
U. compressa (4)	-17.8±2.4 (-20.6 to -15.4)
U. flexuosa (13)	-16.0±3.7 (-25.9 to -10.4)
U. intestinalis (16)	-15.3±2.5 (-20.3 to -8.9)
U. lactuca (31)	-14.1±3.1 (-19.6 to -7.7)
<i>U. linza</i> (6)	-15.6±2.4 (-19.4 to -13.2)
U. lobata (5)	-13.2±1.9 (-15.3 to -11.1)
U. prolifera (3)	-14.2±1.8 (-15.5 to -12.2)

Table 1. Carbon isotopic composition (‰) in species of Phylum Chlorophyta collected along Gulf of California coastlines.

Species (n composite samples)	$\delta^{13}$ C±SD (Min to Max, ‰)
Colpomenia sp. (11)	-11.0±3.7 (-19.0 to -5.4)
C. ramosa (4)	-11.4±2.6 (-13.8 to-7.8)
C. sinuosa (7)	-10.2±3.0 (-16.3 to -7.2)
C. tuberculata (64)	-8.7±3.2 (-19.2 to -2.2)
Padina sp. (15)	-11.1±1.5 (-13.1 to-7.9)
P. crispata (3)	-11.3±1.7 (-12.5 to -10.1)
P. durvillei (36)	-13.2±2.6 (-20.0 to -9.2)
Sargassum sp. (34)	-14.3±2.4 (-18.7 to -8.0)
S. herporhizum (7)	-13.7±1.6 (-16.6 to -11.5)
S. horridum (12)	-15.5±2.9 (-19.7 to-9.5)
S. johnstonii (10)	-15.4±2.0 (-17.7 to -11.8)
S. lapazeanum (7)	-14.5±1.6 (-17.2 to -12.8)
S. sinicola (31)	-15.1±2.4 (-21.1 to -12.1)

Table 2. Carbon isotopic composition (‰) in species of Phylum Ochrophyta collected along Gulf of California coastlines.

1300 Table 3. Carbon isotopic composition (‰) in species of Phylum Rhodophyta collected along Gulf

1301 of California coastlines.

Species (n composite samples)	$\delta^{13}C\pm SD$ (Min to Max, ‰)
Gracilaria sp. (18)	-15.5±2.4 (-21.8 to -12.2)
Gracilaria sp.2 (3)	-14.4±3.7 (-18.7 to -12.3)
G. crispata (7)	-15.1±3.0 (-19.1 to -10.1)
G. pacifica (6)	-16.5±1.6 (-18.6 to -13.6)
G. spinigera (3)	-14.9±3.8 (-17.7 to -12.2)
G. subsecundata (8)	-15.9±2.8 (-20.3 to -12.8)
G. tepocensis (3)	-15.1±1.9 (-17.0 to -13.2)
G. textorii (4)	-16.2±2.6 (-18.1 to -14.3)
G. turgida (5)	-15.3±3.6 (-20.7 to -12.0
G. vermiculophylla (16)	-15.9±3.8 (-23.4 to -8.8)
Hypnea sp. (14)	-14.9±2.6 (-20.9 to -11.4)
H. johnstonii (5)	-11.2±3.5 (-13.8 to -6.5)
H. pannosa (5)	-11.8±3.3 (-15.0 to -6.4)
H. spinella (6)	-16.4±1.8 (-19.2 to -14.9)
H. valentiae (6)	-15.2±2.3 (-19.2 to -12.7)
Laurencia sp. (8)	-12.9±1.2 (-14.7 to -10.5)
L. pacifica (8)	-14.9±2.2 (-19.0 to -12.7)
L. papillosa (3)	-15.7±0.3 (-15.9 to -15.6)
<i>Spyrida</i> sp. (5)	-17.1±1.12 (-19.1 to -16.1)
S. filamentosa (14)	-15.9±3.8 (-26.2 to -11.5)

1305 Table 4. Summary of the estimated regression coefficients for each simple linear regression

1306 analyses and on the constant of fitted regression models. Estimated regression coefficients includes

- 1307 degrees of freedom for the error (DFE), root-mean-square error (RMSE), coefficients of
- 1308 determination (R<sup>2</sup>) and the adjusted R<sup>2</sup> statistics, Mallow's Cp criterion (Cp), Akaike Information
- 1309 Criterion (AIC), Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the
- 1310 test (Prob > F). Models information includes value of the constant a ( $\delta^{13}C$ , ‰), standard error (SE),
- 1311 t ratio and Prob > |t| (values \* are significant).

Independent			J	Estimated Adjust	regressio	Model constant (a) $\delta^{13}C$							
variables	DFE	RMSE	$\mathbb{R}^2$	R <sup>2</sup>	Ср	AICc	BIC	ratio	Prob > F	(‰)	SE	t ratio	Prob >  t
					Inherent	macroalga	ae proper	ties					
Phyla	806	3.66	0.08	0.07	3	4,401	4,420	33.1	<.0001**	-13.98	0.13	-107.4	<.0001**
Morphofunctional	788	3.10	0.35	0.34	21	4,149	4,251	21.6	<.0001**	-14.21	0.35	-40.80	<.0001**
Genus	746	2.92	0.46	0.41	63	4,104	4,393	10.1	<.0001**	-14.71	0.23	-62.64	<.0001*
Species	641	2.79	0.57	0.46	168	4,195	4,898	5.2	<.0001**	-14.60	0.16	-93.22	<.0001**
					Biogeogra	aphical co	llection 2	zone					
GC coastline	807	3.79	0.01	0.01	2	4,456	4,470	7.4	0.0067*	-13.97	0.13	-104.5	<.0001**
Coastal sector	803	3.73	0.05	0.04	6	4,433	4,465	7.9	<.0001*	-14.12	0.16	-90.85	<.0001**
Latitude	807	3.80	0.00	0.00	2	4,462	4,476	1.5	0.23	-12.25	1.41	-8.71	<.0001**
Longitude	807	3.81	0.00	0.00	2	4,463	4,477	0.1	0.80	-15.44	5.83	-2.65	0.0082*
					Н	abitat feat	tures						
Substrate	807	3.80	0.00	0.00	2	4,460	4,474	3.2	0.08	-13.82	0.15	-92.06	<.0001*
Hydrodynamic	807	3.80	0.00	0.00	2	4,462	4,476	1.3	0.26	-13.88	0.15	-95.00	<.0001**
Emersion level	807	3.69	0.06	0.06	2 Enviro	4,412 nmental c	4,427 onditions	52.2 s	<.0001**	-14.05	0.13	-107.6	<.0001**
Temperature	802	3.70	0.01	0.01	2	4,390	4,404	5.4	0.0207*	-16.11	0.96	-16.78	<.0001*
pH	807	3.73	0.04	0.04	2	4,430	4,444	33.4	<.0001**	-32.45	3.21	-10.13	<.0001**
Salinity	806	3.80	0.00	0.00	2	4,456	4,470	0.9	0.34	-15.77	1.91	-8.27	<.0001**
*p<0.05, **p<0.0001													

1313	Table 5. Summary of the estimated regression coefficients for each multivariate linear
1314	regression analyses and on their constant of fitted regression models performed in
1315	individuals binned by genus. Estimated regression coefficients include degrees of freedom
1316	for the error (DFE), root-mean-square error (RMSE), coefficients of determination (R <sup>2</sup> ) and
1317	the adjusted R <sup>2</sup> statistics, Mallow's Cp criterion (Cp), Akaike Information Criterion (AIC),
1318	Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the test (Prob
1319	> F). Model information includes value of the constant a ( $\delta^{13}$ C, ‰), standard error (SE), t



Estimated regression coefficients									Model constant (a)						
				Adjust					Prob >	δ13C			Prob >		
Independent variables	DFE	RMSE	R <sup>2</sup>	R <sup>2</sup>	Ср	AICc	BIC	F ratio	F	(‰)	SE	t ratio	t		
Coastal sector	652	2.78	0.57	0.47	157	4,169	4,834	20.0	<.0001*	-17.52	0.64	-27.24	<.0001*		
Substrate	711	2.90	0.49	0.42	98	4,140	4,577	0.4	0.52	-16.35	0.62	-26.20	<.0001*		
Hydrodynamic	714	2.87	0.50	0.43	95	4,120	4,545	0.1	0.78	-16.53	0.64	-25.95	<.0001*		
Emersion level	713	2.77	0.53	0.47	96	4,060	4,489	153.0	<.0001*	-16.65	0.60	-27.85	<.0001*		
Temperature	695	2.81	0.50	0.43	109	4,083	4,564	98.4	<.0001*	-14.60	0.92	-15.91	<.0001*		
Temperature ranges	686	2.87	0.49	0.40	118	4,128	4,645	97.7	<.0001*	-12.91	0.40	-31.97	<.0001*		
pH	701	2.86	0.51	0.43	108	4,134	4,611	156.6	<.0001*	-28.57	2.69	-10.64	<.0001*		
pH ranges	697	2.67	0.57	0.51	112	4,028	4,522	152.2	<.0001*	-16.39	0.58	-28.05	<.0001*		
Salinity	697	2.89	0.50	0.42	111	4,151	4,640	162.2	<.0001*	-17.75	1.63	-10.88	<.0001*		
Salinity ranges	721	2.91	0.47	0.41	86	4,117	4,504	167.8	<.0001*	-17.64	0.74	-23.68	<.0001*		

1327 Table 6. Summary of the estimated regression coefficients for each multivariate linear regression

- 1328 analyses and on their constant of fitted regression models performed in individuals binned by
- 1329 coastline sector and genus. Estimated regression coefficients include degrees of freedom for the
- 1330 error (DFE), root-mean-square error (RMSE), coefficients of determination (R<sup>2</sup>) and the adjusted
- 1331 R<sup>2</sup> statistics, Mallow's Cp criterion (Cp), Akaike Information Criterion (AIC), Bayesian
- 1332 Information Criterion (BIC) minimum, F Ratio test, and p-value for the test (Prob > F). Model
- 1333 information includes value of the constant a ( $\delta^{13}$ C,  $\infty$ ), standard error (SE), t ratio and Prob > |t|
- 1334 (values \* are significant).

			Est	timated	regre	ssion coe		Model constant (a)					
				Adjust					Prob >	δ <sup>13</sup> C			Prob >
Independent variables	DFE	RMSE	R <sup>2</sup>	R <sup>2</sup>	Ср	AICc	BIC	F ratio	F	(‰)	SE	t ratio	t
Substrate	590	2.76	0.62	0.47	219	4,287	5,155	15.8	<.0001*	-17.08	0.66	-25.72	<.0001*
Hydrodynamic	592	2.73	0.62	0.49	217	4,266	5,128	18.6	<.0001*	-17.18	0.67	-25.70	<.0001*
Protection level	590	2.75	0.62	0.48	219	4,285	5,153	20.0	<.0001*	-17.51	0.64	-27.22	<.0001*
Emersion level	603	2.69	0.63	0.50	206	4,217	5,045	18.6	<.0001*	-17.47	0.64	-27.49	<.0001*
Temperature ranges	569	2.74	0.61	0.46	235	4,293	5,202	28.0	<.0001*	-13.73	0.45	-30.32	<.0001*
pH ranges	580	2.50	0.69	0.57	229	4,155	5,051	9.7	0.0019*	-16.88	0.62	-27.15	<.0001*
Salinity ranges	631	2.76	0.58	0.47	176	4,183	4,913	21.2	<.0001*	-18.30	0.79	-23.05	<.0001*

1337 Table 7. Summary of the estimated regression coefficients for each multivariate linear regression

1338 analyses and on their constant of fitted regression models performed in individuals binned in

1339 coastline sector, habitats features, environmental conditions, and Physiological performed

1340 separately by morpho-functional groups and genus. Estimated regression coefficients include

- 1341 degrees of freedom for the error (DFE), root-mean-square error (RMSE), coefficients of
- 1342 determination (R<sup>2</sup>) and the adjusted R<sup>2</sup> statistics, Mallow's Cp criterion (Cp), Akaike Information
- 1343 Criterion (AIC), Bayesian Information Criterion (BIC) minimum, F Ratio test, and p-value for the
- 1344 test (Prob > F). Model information includes value of the constant a ( $\delta^{13}$ C, ‰), standard error (SE),
- 1345 t ratio and Prob > |t| (values \* are significant).

			Est	Model constant (a)									
			1	Adjust	0				Prob >	δ <sup>13</sup> C			Prob >
Full model	DFE	RMSE	$\mathbb{R}^2$		Ср	AICc	BIC	F ratio	F	(‰)	SE	t ratio	t
Coastline sector + Habita	ats feature	es + Morph	hofunctio	onal gro	oup								
I-Morpho-functional	593	2.79	0.60	0.46	216	4,301	5,160	20.8	<.0001*	-13.49	0.57	-23.52	<.0001*
Coastline sector + Envire	onmental	conditions	s + Morp	hofunc	tional	group							
II-Morpho-functional	680	2.90	0.51	0.42	129	4,189	4,750	25.1	<.0001*	-13.42	0.54	-24.74	<.0001*
Coastline sector + Habita	at features	s+ Genus											
I-Genus	482	2.66	0.71	0.51	327	4,565	5,655	15.8	<.0001*	-16.93	0.73	-23.27	<.0001*
Coastline sector + Envire	onmental	conditions	s + Genu	s									
II-Genus	494	2.49	0.72	0.55	310	4,374	5,438	14.8	0.0001*	-13.55	0.64	-21.17	<.0001*
1347 Table 8. Constant of fitted regression model explaining the  $\delta^{13}$ C variability by morpho-functional

1348 groups. Model information includes value of the constant a ( $\delta^{13}$ C,  $\infty$ ), standard error (SE), t ratio

1349 and Prob > |t|. Only morpho-functional groups with significant effects are enlisted.

Term	Estimated	SE	Razón t	Prob >  t
Model constant	-14.2	0.4	-40.80	<.0001**
R-Smaller-sized articulated corallines	4.5	1.7	2.58	0.0100*
O-Compressed with branched or divided thallus	1.2	0.5	2.66	0.0079*
C-Erect thallus	1.8	0.6	2.84	0.0046*
R-Larger-sized articulated corallines	6.3	0.8	7.95	<.0001*
O-Hollow with spherical or subspherical shape	5.0	0.5	10.51	<.0001*
R-Blade-like with one of few layers of cells	-5.9	3.0	-1.98	0.0476*
C-Tubular	-1.6	0.5	-3.26	0.0012**
R-Filamentous uni&pluriseriate with erect thallus	-2.2	0.6	-3.92	<.0001*
R-Flattened macrophytes with cortication	-8.9	1.3	-7.10	<.0001*
p<0.03, p<0.0001				

1359 Table 9. Constant of fitted regression model explaining the  $\delta^{13}$ C variability by genus. Model

1360 information includes value of the constant a ( $\delta^{13}C$ , ‰), standard error (SE), t ratio and Prob > |t|.

Term	Estimated	SE	Razón t	Prob >  t
Model constant	-14.7	0.2	-62.64	<.0001**
Corallina	6.4	2.9	2.22	0.0269*
Tacanoosca	3.5	1.3	2.71	0.0070*
Jania	5.0	1.7	2.97	0.0031*
Struveopsis	4.1	1.3	3.15	0.0017*
Codium	2.3	0.6	4.08	<.0001**
Padina	2.2	0.5	4.8	<.0001**
Hydroclathrus	7.3	1.1	6.59	<.0001**
Amphiroa	6.8	0.8	9.05	<.0001**
Colpomenia	5.4	0.4	14.02	<.0001*
Spyridia	-1.5	0.7	-2.10	0.0361*
Gracilaria	-0.9	0.4	-2.18	0.0294*
Polysiphonia	-3.7	0.8	-4.82	<.0001**
Schizymenia	-19.1	2.1	-9.33	<.0001**

1361 Only genus with significant effects are enlisted.

1365 Table 10. Constant of fitted regression model explaining the  $\delta^{13}$ C variability by species. Model

1366 information includes value of the constant a ( $\delta^{13}$ C, ‰), standard error (SE), t ratio and Prob > |t|.

1367 Only genus with significant effects are enlisted.

Term	$\delta^{13}$ C, ‰ estimated	SE	Razón t	Prob >  t
Model constant	-14.6	0.2	-93.22	<.0001**
Hypnea pannosa	2.8	1.3	2.24	0.0256*
Colpomenia ramosa	3.2	1.4	2.27	0.0237*
Corallina vancouverensis	6.3	2.8	2.27	0.0238*
Caulerpa peltata	3.9	1.6	2.4	0.0165*
Codium sp.	3.0	1.3	2.4	0.0167*
Amphiroa misakiensis	7.1	2.8	2.55	0.0110*
Jania sp.	5.0	2.0	2.56	0.0106*
Codium brandegeei	2.8	1.1	2.63	0.0088**
Hypnea johnstonii	3.4	1.3	2.74	0.0063**
Tacanoosca uncinata	3.4	1.3	2.74	0.0062**
Struveopsis sp.	4.0	1.4	2.86	0.0044**
Padina durvillei	1.4	0.5	2.87	0.0043**
Amphiroa sp.3	8.2	2.8	2.95	0.0033**
Codium simulans	3.2	0.9	3.41	0.0007**
Amphiroa sp.2	6.6	1.6	4.1	<.0001**
Colpomenia sinuosa	4.4	1.1	4.17	<.0001**

<i>Colpomenia</i> sp.	3.6	0.9	4.27	<.0001**
Padina sp.	3.5	0.7	4.77	<.0001**
Hydroclathrus clathratus	7.2	1.1	6.82	<.0001**
Amphiroa sp.	8.1	0.9	8.67	<.0001**
Colpomenia tuberculata	5.9	0.4	15.45	<.0001**
<i>Spyrida</i> sp.	-2.5	1.3	-1.97	0.0496*
Pyropia thuretii	-5.5	2.8	-1.98	0.0480*
Ulva acanthophora	-1.2	0.6	-2.06	0.0399*
Grateloupia filicina	-2.4	1.1	-2.08	0.0382*
Rhodymenia sp.	-4.1	2.0	-2.08	0.0380*
Ulva compressa	-3.2	1.4	-2.33	0.0203*
Rhizoclonium riparium	-5.1	1.6	-3.15	0.0017**
Polysiphonia sp.	-4.8	1.4	-3.44	0.0006**
Halymenia actinophysa	-9.9	2.8	-3.57	0.0004**
Cladophora microcladioides	-7.2	2.0	-3.64	0.0003**
Polysiphonia mollis	-5.2	1.1	-4.93	<.0001**
<u>Schizymenia pacifica</u> *p<0.05, **p<0.001	-19.2	2.0	-9.76	<.0001**



1373 Fig. S1. Histogram representing the distribution of  $\delta^{13}$ C-DIC values in surface seawater in the Gulf

<sup>1374</sup> of California.



1378Fig. S2. Histograms representing the distribution of  $\Delta^{13}$ C-macroalgal in macroalgae collected in the1379Gulf of California for Phyla a) Chlorophyta, b) Rhodophyta, and c) Ochrophyta.