Mediterranean seagrasses as carbon sinks: Methodological and regional differences

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Abstract. The increasing rates of CO_2 due to anthropogenic activities are causing important potential climate threats for the Mediterranean Sea: ocean acidification and warming. In this region, two seagrass species, *Posidonia oceanica* and *Cymodocea nodosa* can play a crucial role in climate change mitigation. Seagrasses can act as carbon sinks; buffer lowering pH values during the day and store carbon in the sediment underneath their meadows. In this study, we analysed published and unpublished data (collected for this study) on seagrass community metabolism to compare two methodologies: benthic chambers and multiparametric sensors and evaluate trends through time for these two species. Furthermore, we analysed seasonal trends of both seagrass species´ metabolic rates and their variation between the Eastern and Western Mediterranean basins. Most evaluated meadows, 80.9%, were autotrophic. Calculated metabolic rates differ between methodologies, with multiparametric sensors estimating rates almost an order of magnitude higher, 143.22 ± 28.21 (SE) mmol O_2 m^{-2} day^{-1} for Net Community Production (NCP) compared to an average of 18.75 ± 3.80 (SE) mmol O_2 m^{-2} day^{-1} for measurements with benthic chambers. However, sensors are not able to differentiate between habitats and only useful to assess seagrass metabolism at a broader community level, whereas benthic chambers are capable to evaluate rates at species level and confirm that *P. oceanica* is more productive compared to *C. nodosa*. We found similar metabolic rates in the Eastern and Western Mediterranean regions for *P. oceanica* with the benthic chamber technique and higher NCP in the West based on sensor measurements.

1 Introduction

A fifth of the global carbon sequestration in marine sediments (Duarte et al., 2004; Kennedy et al., 2010) can be attributed to seagrass meadows, despite the fact that they cover only a 0.1% of the ocean surface. This "blue carbon", which is defined as organic carbon buried in sediments underneath marine vegetation (Duarte et al., 2004; Kennedy et al., 2010; Mcleod et al., 2011; Greiner et al., 2013) is the result of the combination of intense metabolic activity of the vegetation, high trapping capacity of allochthonous matter and an effective carbon preservation in sediments underneath meadows (Cebrian, 1999). Due to the enhanced deposition rates caused by the physical presence of the canopies in the water-column seagrass meadows capture suspended organic matter, which accumulates as organic matter in the sediment (Romero et al., 1994; Pergent et al., 1997; Mateo et al., 2006; Hendriks et al., 2008; Kennedy et al., 2010). However, also the *in situ* plant growth, for which productivity can be a proxy, contributes to organic matter accumulation in the sediment (Greiner et al., 2013). Even though carbon dioxide (CO_2) uptake through metabolic activity is on much shorter timescales compared to carbon storage, seagrass productivity and growth contributes to the sequestration and storage of a considerable amount of carbon in underlying sediments (Mcleod et al., 2011; Fourqurean et al., 2012). There are species specific differences in carbon burial rates and stocks; for *Cymodocea nodosa,* the annual carbon burial from productivity at a specific site has been estimated as 4.4 g C m^{-2} y^{-1} whereas *P. oceanica* accumulated 66.4 g C m^{-2} y^{-1} at the same location (Cebrián et al., 1997). In general, seagrass communities tend to be autotrophic, although primary production varies depending on the evaluated temporal scales: daily, seasonally (Bay, 1984; Alcoverro et al., 1995; Gobert et al., 2006; Hendriks et al., 2014) and annually (Champenois et al.,

45 2012, 2019). Furthermore, primary production and related carbon storage also depend on many other variables such as structural complexity (Trevathan-Tackett et al., 2015), size of the meadow (Ricart et al., 2017) local nutrient dynamics (Armitage et al., 2016), hydrodynamics (Samper-Villarreal et al., 2016) and water depth (Serrano et al., 2014). A consistent methodology to estimate seagrass productivity is crucial to estimate its contribution to the global carbon sink capacity and to approximate the economic and ecological consequences of the decline of this ecosystem worldwide (Orth et al., 2006;

50 Waycott et al., 2009). Despite their importance, little is known about the effect of increasing rates of atmospheric CO_2 on these ecosystems and predictions for the open ocean may not reflect future conditions in coastal zones (Hendriks et al., 2010; Hofmann et al., 2011; Kelly et al., 2013; Lacoue-Labarthe et al., 2016), the main habitat for seagrasses. Increasing dissolved CO_2 in the water column generates pH changes and contributes to the complex biogeochemical dynamics driving coastal ecosystems (Aufdenkampe et al., 2011), which are often also affected by human activities. Dynamics of the carbonate

55 system in coastal regions are frequently influenced by benthic ecosystems that have the capacity to buffer physical and chemical conditions of the environment (Gutiérrez et al., 2011). By capturing CO_2 and releasing oxygen (O_2), seagrass ecosystems drive fluctuations in pH and dissolved oxygen concentrations in the adjacent water column that follow daily and seasonal patterns modulated by metabolic activity (Duarte et al., 2013; Hendriks et al., 2014). Primary production is hence an important component in the assessment of pH variation in coastal ecosystems. Through their photosynthetic activity, pH

60 modification of the adjacent water mass by seagrasses attenuates ocean acidification, possibly providing a temporary refuge for calcifying organisms while oxygenating the water column through O_2 production (Hendriks et al., 2014). Seagrass metabolism can be conditioned by abiotic parameters such as temperature and can be assessed through the changes in concentration of dissolved oxygen in water. Water temperature has an effect on the air-water exchange by modifying oxygen solubility and affects ecosystem metabolism (Brown et al., 2004; Vaquer-Sunyer et al., 2013). In addition, anthropogenic

65 pressures such as eutrophication (Paerl, 2006) or climate change are affecting oxygen cycling (Keeling et al., 2002; Conley et al., 2009; Keeling et al., 2010) and dissolved oxygen is one of the environmental parameters that has changed more drastically in a short period of time (Diaz et al., 1995; Diaz, 2001), with potential catastrophic consequences for marine life (Vaquer-Sunyer et al., 2008). In coastal ecosystems, increased nutrient inputs contribute to higher organic production and oxygen demand with a consequently greater likelihood of hypoxia (Karim et al., 2003; Zhang et al., 2010). Measuring

70 dissolved oxygen concentrations in seagrass meadows allows to infer metabolic parameters such as Gross Primary Production (GPP) for the meadow community. Metabolic community status is defined by the Net Community Production (NCP), the difference between GPP and Community Respiration (CR). When a community exports organic carbon to adjacent systems, it is considered as net autotrophic (NCP > 0), on the other hand if the community gets external organic carbon inputs to sustain its own metabolism, it is considered as net heterotrophic (NCP < 0) (Champenois et al., 2012).

75 The dominant seagrass species in the Mediterranean Sea are *Cymodocea nodosa* and *Posidonia oceanica*, inhabiting a region that has been defined as a "hotspot for climate change" (Giorgi, 2006) with warming rates of two- to four-fold higher than in other regions (Vargas-Yáñez et al., 2008; Burrows et al., 2011). An increase in the seasonal average temperature of 2.2ºC in winter and 3.4ºC in summer is projected for the end of this century for the Western Basin of the Mediterranean Sea in a scenario with moderated greenhouse gasses emissions (Jordà et al., 2012). Furthermore, extreme thermal events are expected

80 to be more intense and frequent in the Mediterranean region (IPCC, 2021) while the basin is not subject to the same rates of warming, with indications that the Eastern Mediterranean is warming faster than the Western Mediterranean (Amitai et al., 2020; Nykjaer, 2009). Climate warming can impact *P. oceanica* meadows negatively, as higher temperatures stress the species physiologically (Marbà et al., 2010), with shoot mortality increasing during heat waves exceeding 28°C at the end of summer (Diaz-Almela et al., 2007). The total surface area occupied by *P.oceanica* meadows is estimated to range between 1

85 to 2% of the total surface area of the Mediterranean Sea (Béthoux et al., 1986; Pasqualini et al., 1998), although this number is uncertain (Bonacorsi et al., 2013). The distribution has been estimated as 510.710 ha in the Western and 713.992 ha in the Eastern Mediterranean basin (Telesca et al., 2015), which might be conservative due to lack of data, with much more data

available in the Western basin compared with the Eastern part where absence of data is common. Contrary to other European seagrasses with decreasing loss rates (de los Santos et al., 2019), the extent of *P. oceanica* meadows is decreasing considerably with between 13 as-50% areal extent lost since the 1960s (Boudouresque et al., 2009; Marba et al., 2014; Telesca et al., 2015). *C. nodosa* meadows have a higher thermal tolerance and are supposed to cope better with increasing temperatures (Egea et al., 2018), even though high temperatures during heat waves over the coming decades might increase the annual mean temperature with 4ºC and will probably exceed the limit beyond which *C. nodosa* losses can be expected in the Mediterranean Sea (Olsen et al., 2012; Chefaoui et al., 2018). Also, all Mediterranean water bodies are affected by anthropogenic CO_2 emissions, however more so in the Western basin, demonstrated by decreasing values of pH ranging with -0.14 to -0.005 pH units drops since the beginning of the industrial era to 2001, higher than elsewhere in the open ocean (Touratier et al., 2011). Therefore, ocean acidification is a climate change indicator that has been characterized as one of the most important for the Mediterranean Sea, together with temperature and UV radiation (Micheli et al., 2013). In addition to climate change, the Mediterranean Sea is likely to be more impacted by human disturbances (e.g., overfishing, increasing pollution from wastewater outfalls, riverine, farmland, runoffs, fish farming and the introduction of alien species (Lejeusne et al., 2010) than other seas (Giorgi et al., 2008; Richon et al., 2019). With these multiple pressures increasing it is crucial to evaluate the functions and services provided by key coastal ecosystems such as seagrasses.

Seagrass metabolism has classically been measured using closed benthic chambers (Duarte et al., 2010a). However, the spatial heterogeneity of these ecosystems (Gazeau et al., 2005) and their high temporal variability cannot be easily estimated with this approach (Karl et al., 2003). For this purpose, using sensors can be more suitable as dissolved oxygen concentrations can be evaluated for longer periods of time. Additionally, the aquatic eddy covariance technique can give even more precise values for productivity, although spatially limited, nevertheless so far there has only been one study including *P. oceanica* in the Mediterranean Sea (Koopmans et al., 2020). GPP values obtained with the use of benthic chambers could provide an underestimate as a result of photorespiration, while the use of multiparametric probes measuring oxygen in the canopy probably provides more realistic GPP values (Champenois et al., 2012). The use of multiparametric probes to measure O_2 also provides the opportunity to measure metabolic rates without damaging roots or rhizomes. In river and lake ecosystems, the measurement of metabolism by oxygen probes and loggers is a commonly used method (Cole et al., 2000; Coloso et al., 2008), while it is not as wide spread in coastal waters (Odum et al., 1958; Odum et al., 1962; Ziegler et al., 1998; Vaquer-Sunyer et al., 2012) due to the higher lateral transport rates of water in these systems.

The aim of this study is to evaluate the potential carbon capture of the dominant Mediterranean Sea seagrass (*Posidonia oceanica* and *Cymodocea nodosa*) communities through their metabolic activity, comparing two methodologies (benthic chambers and multiparametric probes) and evaluate the spatial and temporal differences between Mediterranean regions. We do so by conducting field measurements amended by published data compiled from the literature.

2 Methods

2.1 Data compilation

Data for the metabolic parameters was extracted from the literature, through a literature search on SCOPUS and the Web of Science using the keywords "Posidonia", OR "Cymodocea", OR "Seagrass", AND "Productivity", OR "Metabolism" and manually screened for data on metabolism in the Mediterranean basin. This database was extended with submitted data and data from dedicated sampling campaigns in 2016 in Mallorca (Western Mediterranean) and 2017 in the Eastern basin (Crete and Cyprus, see Table 1, Fig. 1). We also compiled data from multiparametric sensors collected during different periods ranging from 2011 to 2019 (for details see Table 1). While data using the benthic chambers methodology had a higher

number of literature studies, with a total of 12 publications with data for *P. oceanica* and/or *C. nodosa* meadows (for details see Table 2), and a wider temporal cover with studies carried out from 1982 to 2019. Importantly, this study adds new data on Mediterranean seagrasses metabolism in the Eastern Mediterranean Basin (Crete, Cyprus; Table 1), where little data has been published before. Data available as oxygen concentration over time was processed and analysed to obtain the metabolic parameters, when this was not available, we used the reported metabolic rates.

2.2 Site description

Data from multiparametric probes in either *Posidonia oceanica* and/or *Cymodocea nodosa* meadows came from one site in France (Corsica, literature), eight sites in Spain (Mallorca, of which one collected for this study), three in Crete (Greece) and one in Cyprus (Republic of Cyprus) the last four all collected for this study, (see Table 1). All were shallow sites, ranging from 2.9 metres depth (Punta Negra, Mallorca) to 15.7 metres depth (Cap Enderrocat, Mallorca). Extracted data from Corsica came from Cape Revellata, a protected area. The site in Limassol (Cyprus) is adjacent to the second largest city in Cyprus and it is considered an impacted area affected by high anthropogenic pressures related to intense tourism and the construction of extensive coastal infrastructures. Marathi and Kalami are located close to each other next (< 10 km) to the Port of Souda in Western Crete, both sampling sites have similar environmental conditions and are impacted by notably sewage discharge, agriculture and industrial/chemical pollution (Simboura et al. (2016). Maridati, in Eastern Crete, is a pristine bay but affected by intermittent discharges of an ephemeral stream (Wesselmann et al., 2021). In Mallorca, multiparametric probes were

 in sites with a range of environmental impacts, encompassing sites with protection like Cap Enderrocat, which forms part of an SPA (*Special Protection Area*) under the *Birds* Directive and is a SIC (Site of Community Importance, Natura 2000), as well as Son Verí and Cala Blava, which are also protected and host 11.5% of the total Posidonia meadows within the Marine Protected Area of Cap Enderrocat - Cap Blanc. Even though the locations are protected, the bay of Palma is influenced by various outlets of waste water treatment plants. On the opposite site in the bay of Palma Punta Negra is considered as a Natural Area of Special Interest (SPA and a natural space protected by law by the Government of the Balearic Islands). Magaluf (Mallorca, Bay of Palma) is in front of a touristic beach but the location of the sensors was sheltered behind an island (Isla Sa Porrassa). Sant Elm (Mallorca) is located in a relatively pristine area but near a sewage plant effluent. Pollença (Mallorca) is in an enclosed bay affected by considerable organic input from the s´Albufereta wetlands, an effluent of the sewage plant, nearby harbour and urban area. The least impacted site would be Sta. Maria, a bay located on the coast of Cabrera island and the most pristine sampling area. Cabrera island is part of a Maritime and Terrestrial National Park located at the Cabrera Archipelago, and recognized internationally as special protection zones with importance for the Mediterranean and SAC (Special Areas of Conservation). The sampling sites where sensors were deployed in the Mediterranean therefore include sites with different degrees of human impact and protected areas with very low anthropogenic impact.
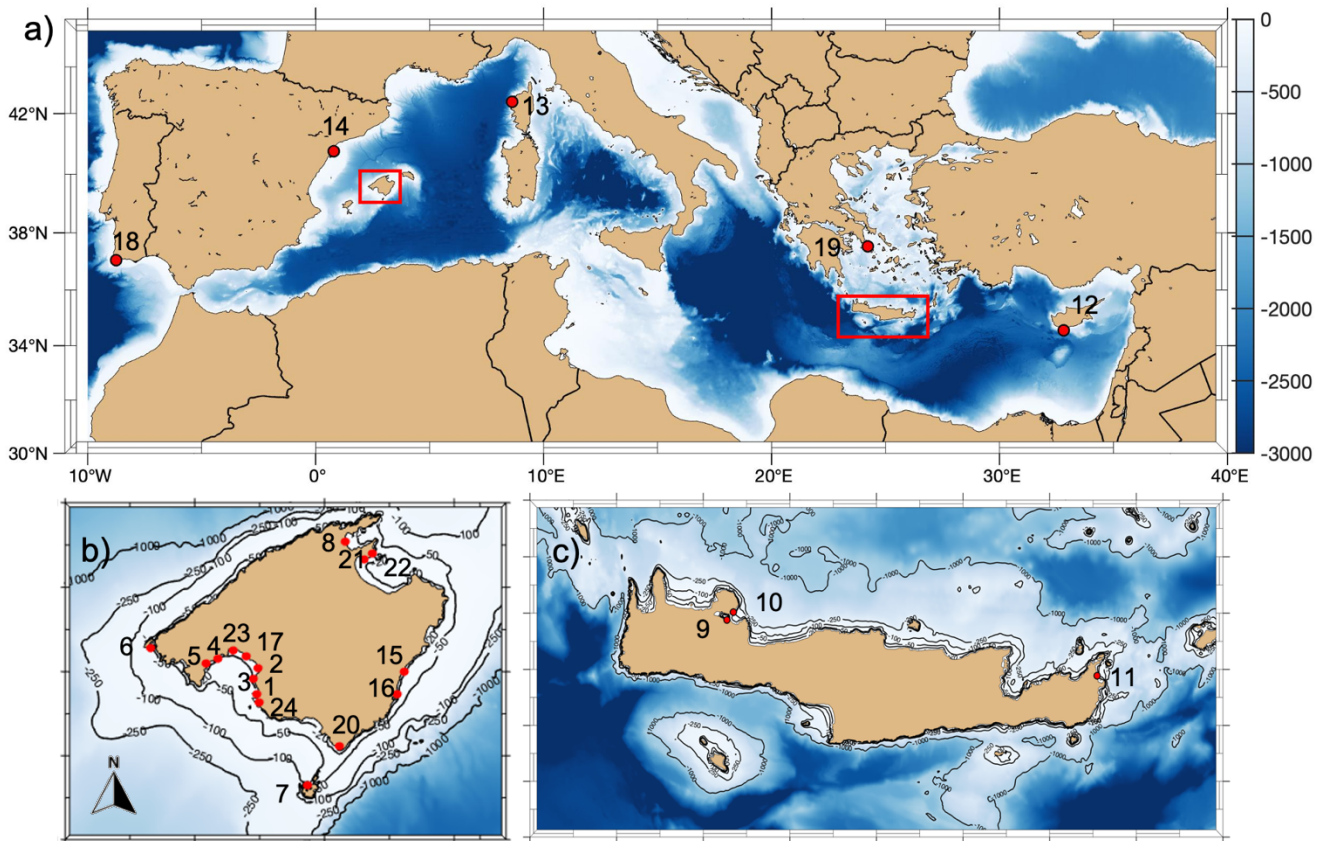
Figure 1. Location of sampling sites with numbered ID as described in Table 1 and Table 2 included in this compilation (map generated with Matlab (2012), M_Map: A mapping package for Matlab).

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Table 1. Characteristics of sampling stations where multiparametric probes were deployed. Temperature and salinity represent average values during the deployment.

Region	Station	ID	Date	Species	Depth (m)	Temperature (°C)	Salinity
Mallorca	Cap Enderrocat[1]	1	25/8/2016-26/8/2016	*Posidonia oceanica*	14.6	26.6	40.1
			18/8/2016-29/8/2016	*Cymodocea nodosa*	15.7	26.6	38.8
Mallorca	Son Verí[2]	2	5/06/2012-11/06/2012	*Posidonia oceanica*	7.3	23.4	40.8
			5/06/2012-8/06/2012		5.4	23.4	40.8
Mallorca	Cala Blava[2]	3	6/06/2012-12/06/2012	*Posidonia oceanica*	5.9	23.8	38.9
			5/6/2012-11/6/2012		4.4	23.8	39
Mallorca	Pta. Negra[3]	4	2/07/2019-3/07/2019	*Cymodocea nodosa*	2.9	15.4	36.8
			11/04/2019-12/04/2019	*Posidonia oceanica*	3.3	15.3	37.1
Mallorca	Magalluf[2]	5	20/9/2011-23/9/2011	*Posidonia oceanica*	6.3	26.3	40.5
Mallorca	St. Elm[2]	6	13/09/2011-16/09/2011	*Posidonia oceanica*	9.4	26.8	40.3
Mallorca	Cabrera[2]	7	6/9/2011-9/9/2011	*Posidonia oceanica*	7.2	26.6	40.2
Mallorca	Pollença[4]	8	16/10/2018-17/10/2018	*Cymodocea nodosa*	6.4	23	38.6
			16/7/2018-17/7/2018	*Posidonia oceanica*	6.1	24	39
			15/1/2019-16/1/2019	*Posidonia oceanica*	7.1	13.2	36.9
			15/1/2019-16/1/2019	*Cymodocea nodosa*	7.7	13.2	37

Region	Station	ID		Species	Depth (m)	Temperature (°C)	Salinity
			18/4/2018-19/4/2018	*Cymodocea nodosa*	6.8	16.1	37.7
			18/4/2018-19/4/2018	*Posidonia oceanica*	6.5	16.1	38.4
			25/06/2015-30/06/2015	*Cymodocea nodosa*	8	25.7	40.6
			25/06/2015-1/7/2015	*Posidonia oceanica*	4.5	25.8	40.9
Crete	Marathi[1]	9	18/7/2017-20/7/2017	*Posidonia oceanica*	4.7	26.3	40.5
	Kalami[1]	10	18/7/2017-20/7/2017	*Cymodocea nodosa*	5.4	27	40
	Maridati[1]	11	21/7/2017-23/7/2017	*Cymodocea nodosa*	6.2	25.2	40.5
			21/7/2017-23/7/2017	*Posidonia oceanica*	8.9	25.1	40.7
Cyprus	Limassol[1]	12	4/09/2017-7/09/2017	*Cymodocea nodosa*	3.2	27.3	40.2
Corsica	Revelatta[5]	13	2006-2016	*Posidonia oceanica*	NA	18.7	NA
Corsica	Revelatta[6]	13	2006-2009	*Posidonia oceanica*	NA	18.6	NA

Source: 1. Unpublished data., 2. Hendriks et al., 2014, 3. Marx et al., 2021, 4. Hendriks et al., submitted., 5. Champenois et al., 2019, 6. Champenois et al., 2012.

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Table 2. Characteristics of sampling stations with data for benthic chamber deployments.

Region	Station	ID	Season	Year	Species	Depth (m)	Temperature (°C)	Salinity
Corsica	Bay of Calvi[1]	13	Annual	1982-1984	*Posidonia oceanica*			
Spain	Ebro Delta[2]	14	Summer	2000	*Cymodocea nodosa*			
Mallorca	Magalluf[3]	5	Annual	2001	*Posidonia oceanica*			
Mallorca	Magalluf[4]	5	Summer/Spring	2001	*Posidonia oceanica*		27.5	
Mallorca	Cabrera[5]	7	Summer	2000	*Cymodocea nodosa*		18	38
Mallorca	Sa Paret[5]	15	Summer	2000	*Posidonia oceanica*		18	38
Mallorca	Porto Colom[5]	16	Summer	2000	*Posidonia oceanica*		18	38
Mallorca	Sta Maria[5]	7	Summer	2000	*Posidonia oceanica*		18	38
Mallorca	Magalluf[3]	5	Annual	2002	*Posidonia oceanica*			
Mallorca	Magalluf[3]	5	Annual	2002	*Posidonia oceanica*			
Mallorca	Cap Encerrocat[6]	1	Annual	2002	*Posidonia oceanica*			
Mallorca	Bay of Palma[6]	17	Annual	2002	*Posidonia oceanica*			
Portugal	Ria Formosa[7]	18	Summer	2002	*Cymodocea nodosa*			
Greece	Sounion[8]	19	Annual	2006	*Posidonia oceanica*			
Mallorca	Es Cargol[9]	20	Annual	2006	*Posidonia oceanica*			
Greece	Sounion[8]	19	Winter/Spring	2007	*Posidonia oceanica*			
Corsica	Revellata[10]	13	Summer/Winter	2007-2009	*Posidonia oceanica*			
Corsica	Revellata[11]	13	Fall	2012	*Posidonia oceanica*			
Mallorca	Alcanada[12]	21	Fall	2012	*Posidonia oceanica*		18	36.6
Mallorca	Alcanada[12]	21	Winter	2012	*Posidonia oceanica*		13	35.94
Mallorca	Albufera[12]	22	Summer	2012	*Posidonia oceanica*		25.5	38.27
Mallorca	Calanova[12]	23	Summer	2012	*Posidonia oceanica*		23.5	38.14
Mallorca	Alcanada[12]	21	Summer	2012	*Posidonia oceanica*		25.25	38.23
Mallorca	Arenal[12]	24	Summer	2013	*Posidonia oceanica*		24.25	38.08
Mallorca	Alcanada[12]	21	Spring	2013	*Posidonia oceanica*		18.8	37.7
Mallorca	Arenal[12]	24	Summer	2013	*Posidonia oceanica*		27.6	37.7
Mallorca	Calanova[12]	23	Summer	2013	*Posidonia oceanica*		28.3	37.6
Mallorca	Albufera[12]	22	Summer	2013	*Posidonia oceanica*		23.8	38
Mallorca	Alcanada[12]	21	Summer	2013	*Posidonia oceanica*		23.5	38

Mallorca	Pt. Negra[13]	4	Summer	2019	*Posidonia oceanica*	2.79	26.06	37.79
Mallorca	Pt. Negra[13]	4	Spring	2019	*Cymodocea nodosa*	5.66	15.85	37.01
Mallorca	Pt. Negra[13]	4	Spring	2019	*Cymodocea nodosa*	3.40	15.33	37.42
Mallorca	Pt. Negra[13]	4	Spring	2019	*Cymodocea nodosa*	3.40	15.85	37.02
Mallorca	Pt. Negra[13]	4	Summer	2019	*Cymodocea nodosa*	2.84	26.19	37.75
Mallorca	Pt. Negra[13]	4	Summer	2019	*Cymodocea nodosa*	2.79	26.06	37.79
Mallorca	Pt. Negra[13]	4	Summer	2019	*Cymodocea nodosa*	2.79	26.06	37.79
Mallorca	Pt. Negra[13]	4	Summer	2019	*Cymodocea nodosa*	2.84	26.19	37.75
Mallorca	Pt. Negra[13]	4	Spring	2019	*Posidonia oceanica*	3.40	15.33	37.42
Mallorca	Pt. Negra[13]	4	Spring	2019	*Cymodocea nodosa*	3.40	15.33	37.42
Mallorca	Pt. Negra[13]	4	Fall	2019	*Posidonia oceanica*			
Mallorca	Pt. Negra[13]	4	Summer	2019	*Posidonia oceanica*	2.84	26.19	37.75
Mallorca	Pt. Negra[13]	4	Summer	2019	*Posidonia oceanica*	2.84	26.19	37.75
Mallorca	Pt. Negra[13]	4	Spring	2019	*Posidonia oceanica*	3.40	15.85	37.02
Mallorca	Pt. Negra[13]	4	Spring	2019	*Posidonia oceanica*	3.40	15.85	37.02
Mallorca	Pt. Negra[13]	4	Summer	2019	*Posidonia oceanica*	2.79	26.06	37.79
Mallorca	Pt. Negra[13]	4	Spring	2019	*Posidonia oceanica*	3.40	15.33	37.42
Mallorca	Pt. Negra[13]	4	Fall	2019	*Cymodocea nodosa*			

Source: 1. Frankignoulle et al., 1987, 2. Barrón et al., 2004, 3. Barrón et al., 2009, 4. Barrón et al., 2006, 5. Holmer et al., 2004, 6. Gazeau et al., 2005, 7. Santos et al., 2004, 8. Apostolaki et al., 2010, 9. Gacia et al., 2012, 10. Champenois et al., 2012, 11. Olivé et al., 2016, 12. Agawin et al., 2017, 13. Marx et al., 2021.

2.3 Data analysis

Reported values from the literature were used for benthic chambers, in these articles Net Community Production (NCP) was generally estimated from changes in dissolved oxygen using the Winkler titration spectrophotometric method (Labasque et al., 2004). Benthic chambers enclose a section of the seagrass meadow, and flexible fitted plastic bags, not permeable for gases, assure the possibility of movement of the shoots inside, see details in the method section of each paper for the exact construction used. NCP, GPP and CR data were extracted from literature as well as accompanying biotic parameters. The benthic chamber methodology has been more generally used to assess metabolism of seagrass meadows and the database of this study contains a total of 100 NCP estimations. We compare the data obtained between both methodologies, however for the paper reporting sensor data, we have calculated metabolism from raw oxygen profiles obtained with the multiparametric sensors where possible (including for published and new data described in 2.1), and only used directly reported productivity values for the data obtained from Champenois et al., 2012 and 2019.

2.3.1 Metabolic rate calculations from multiparametric probes

Where available, we used time series of dissolved oxygen (DO, in mg/L), pH_{NBS}, salinity and temperature (°C) measured in *P. oceanica* and/or *C. nodosa* meadows with multiparametric sensors (OTT Hydrolab DSX5 and HL4). pH accuracy for both sensors is ± 0.2, while the accuracy for dissolved oxygen is ± 0.2 mg/L. The duration of the data collection was different depending on the site, from 1 full day to 4 consecutive days (see Table 1), while 24-h periods were used for calculations. Sensors were deployed 0.2 m above the seafloor in seagrass meadows. Data was recorded every 15 minutes except in Cap Enderrocat where readings were taken every 10 minutes (Table 1). Biological metadata detailing habitat traits, was obtained following the methodology described by (Hendriks et al., 2014)

Sensors were calibrated before each deployment with a two-point pH calibration, with 7.00 and 10.00 NIST traceable pH buffers (Hendriks et al., 2014). Oxygen sensors (Hach LDOTM) were calibrated using the water saturated air method

calibration. For validation of salinity, specific conductance calibrations were performed with 50.000uS/cm buffers. For depth measurements, pressure readings were corrected for specific conductance.

Meteorological data for the deployment period was obtained from the Agencia Estatal de Meteorología (AEMET) for the stations in Mallorca, from the Cyprus Department of Meteorology for Cyprus sampling sites and from the Hellenic National Meteorological Service for the locations in Crete (see Table S1 in the Supplement).

With input parameters dissolved oxygen (DO), temperature (°C) and salinity we calculated the metabolic rates of the seagrass habitats using a modification of the model of Coloso et al., (2008) implemented in MATLAB (version 7.5. the Mathworks Inc.) explained in detail in Vaquer-Sunyer et al., (2012). For each station, we manually introduced the Mixed Layer Depth (MLD), the latitude, and year, day and time as day fraction. For wind speed (m/s) we used the $k660$ calculations based on Kihm et al., (2010). Net community production (NCP) was calculated as Gross Primary Production (GPP) - Community Respiration (CR), taking into account diffusive exchange with the atmosphere (D) following Eq. (1):

$$DO = NCP + D + A. \tag{1}$$

The diffusion with the atmosphere is regulated by the difference in DO concentration linked to atmospheric equilibrium (DO sat) and the air-sea gas velocity transfer for oxygen (k) at a given temperature according to Eq. (2).

$$D = k \, (DOsat - DO). \tag{2}$$

where D can be positive (DO addition to the system) or negative (DO removal from the system). Wind speed was estimated at each station for 15 minutes intervals (10 minutes for the Cap Enderrocat station) to predict $k660$ (air-sea gas transfer velocity for O_2 at 20º C and salinity 35) based on Kihm et al., (2010) and Cole et al., (1998). Schmidt number equations for seawater according to Wanninkhof (1992) were used for the k calculation from $k660$. As the cubic model equals the model proposed by Wanninkhof et al., (1999) for short-term winds this parameterization by Kihm et al., (2010) is used.

The model assumes that the only metabolic activity during the night is respiration (Community Respiration; CR) as in the absence of sunlight there is no photosynthetic production. CR can be extracted from the change in O_2 concentration during the night (Net Community Production at night = CR), from 1 h past sunset to 1 h before sunrise. During the day light period Net Community Production (NCP) is considered to be the result of the balance between Gross Primary Production (GPP) and CR. NCP was calculated using the rate of change of DO within the interval of 24 hours, while CR is calculated from night values and converted to hourly rates. As is general, we assumed that CR rates during the light period equal those at night and use the equivalent hourly values of CR for the light period to estimate GPP, adding the oxygen consumption (CR) for this period to NCP to get GPP. GPP and CR could be underestimated since it is likely that CR during daytime exceeds CR at night (Grande et al., 1989; Pace et al., 2005; Pringault et al., 2007) but this underestimation would not affect NCP values (Cole et al., 2000). Individual estimates of CR, NCP and GPP within the measured intervals obtained from the multiparametric probes were calculated over a 24h period for each day and station. As we did not dispose of vertical profiles of Conductivity, Temperature and Depth (CTD) for each station to calculate the Mixed Layer Depth, we used the model of Condie et al., (2001) to calculate the MLD following Eq. (3):

$$S = \frac{\rho C p U_2}{g \alpha H Q}.$$

(3)

where S represents the non-dimensional parameter of the ratio of the input of kinetic energy by the wind to the input of potential energy by solar radiation. ρ is the density, calculated from the salinity, temperature and pressure collected by the *in situ* multiparametric sensor following the formula of Fofonoff et al., (1983). Cp represents the specific heat, considered here to be 3850 J kg $^{-1}$ °C^{-1} as the relative value for seawater. U_2 refers to the diurnally averaged wind speed specified here to be measured 2 m above the water body (Simpson et al., 1974; Holloway, 1980). In our case, the wind data was measured at 10 m above the upstream edge of the water body and was converted, according to the wind profile power law. Eq. (4):

$$\frac{\vartheta}{\vartheta r} = \left(\frac{z}{zr}\right)^{\alpha}. \tag{4}$$

with ϑ as the wind speed (in meters per second) at a determined height z (metres) and ϑr the speed that is known at a reference height (zr). The exponential α is a coefficient derived empirically which varies upon the stability of the atmosphere. In our case, neutral stability is assumed and within those conditions α is approximately $0.143.g\ (Eq.\ (3))$, corresponding to the gravitational acceleration (9.8 m. s^{-1}). $\alpha\ (Eq.\ (3))$ represents the thermal expansion coefficient which was calculated as a function of the absolute salinity, *in situ* temperature and pressure. This function is included in the Gibbs-SeaWater (GSW) Oceanographic Toolbox (McDougall et al., 2011), and evaluates the thermal expansion coefficient α^t in respect with the *in situ* temperature (t), from equation (2.18.1) of the TEOS-10 Manual (IOC, 2015) following Eq. (5):

$$\alpha^t = \alpha^t(S_{A.}\ t.p) = -\frac{1}{v}\frac{\partial\rho}{v\partial T}\Big|S_A.p = \frac{1}{v}\frac{\partial v}{v\partial T}\Big|s_A.p = \frac{g_{Tp}}{g_P}. \tag{5}$$

This function uses the full TEOS-10 Gibbs function $g\ (SA.\ t.\ p)$ of IOC (2015) as the sum of the IAPWS (2008) and IAPWS (2009) Gibbs functions.

Finally, H. Eq. (3) equals the average water depth (in m) and Q Eq. (3) the diurnally averaged shortwave radiative heat flux (Wm^{-2}). Then, the surface mixed layer (z_s) was approximated following Eq.(6) by Condie et al., (2001):

$$z_s = \left((2.9 - 0.20\ ln\ S) \pm 0.04\right). \tag{6}$$

To be as accurate as possible, we tried different intervals for the wind and radiation data. Taking in account that stratification starts to develop between 09:00 and 10:00 hours and reaches its maximum around 16:00 hours (Condie et al., 2001), we decided to take the interval from 9:00 to 16:00 hours into account for the wind data. This interval resulted to be the more accurate (highest R^2 obtained for that interval, $R^2=0.9$) for the wind data, based on the linear regressions between zH/S obtained with the different wind and radiation intervals.

2.3.2 Statistical analysis

We used mixed linear models with package lme4 in the R environment (R core team, 2021) to evaluate differences between methods, regions and species. To reflect the variability between study approaches and sampling procedures and therefore variability in the precision of outcome of each study, we used a linear model where publication was included as random effect unless specified differently. We also analysed abiotic (wind, pH, depth) parameters related to sensor data as there was more additional data associated to these measurements. As the data was not normally distributed according to the Shapiro-Wilk test, we log transformed data for GPP, and CR before analysis. NCP could not be log transformed due to negative values.

3 Results

We compiled 133 CR, 141 GPP and 168 NCP estimates, most of them (86.4 %) restricted to seagrass meadows located in the Western Mediterranean (Table S2). The studied meadows were situated at water depths between 0.5 and 22 m (Table 1, 2) where water salinity ranged from 35.94 to 40.98 (Table 1, 2). The majority (50.0 %) of the metabolic rates were assessed in summer, 19.9% of data was sampled in spring, 10.6% in fall and 7.2% in winter, with seawater temperatures during the measurements varying between 13 °C and 28.5 °C (average 23.2°C $\pm$ 4.4 SD) between locations (Table 1, 2). Estimates of seagrass metabolic rates in Mediterranean seagrass meadows span from 1982 to 2019 (Table 1, 2), but most measurements were conducted after the year 2007 for benthic chambers and even later, after 2015 for sensors. Benthic chamber deployments were concentrated in the Western basin (72 metabolic measurements) with only 6 measurements in the Eastern basin. While 55 metabolic measurements were made in the Western basin with multiparametric sensors, compared to 14 in the Eastern basin.

Table 1. Summary of metabolic rates (mmol O_2 m^{-2} d^{-1}), for the different methodologies, per region, season, and when possible per species.

			NCP	SE	CR	SE	GPP	SE	P/R	SE
benthic chambers			**18.75**	*3.80*	**41.18**	*4.55*	**55.31**	*6.39*	**1.57**	*0.12*
Cymodocea nodosa			-9.18	*5.29*	30.64	*7.48*	22.01	*5.21*	1.49	*0.34*
	WEST		-9.18	*5.29*	30.64	*7.48*	22.01	*5.21*	1.49	*0.34*
		Fall	-1.40	*0.00*	3.50	*0.00*	2.00	*0.00*	0.57	*0.00*
		Spring	-0.44	*1.01*	6.86	*3.42*	6.43	*2.62*	1.05	*0.19*
		Summer	-12.23	*7.05*	39.37	*9.06*	27.90	*6.35*	1.68	*0.46*
Posidonia oceanica			25.47	*4.21*	45.45	*5.55*	68.77	*7.94*	1.60	*0.10*
	EAST		39.36	*5.61*	52.60	*11.73*	93.20	*17.67*	1.86	*0.14*
		Fall	27.04		47.92		74.96		1.56	
		Spring	46.03		42.94		88.97		2.07	
		Summer	48.64	*15.22*	82.38	*24.27*	134.74	*43.20*	1.62	*0.05*
		Winter	32.92	*3.85*	29.99	*4.44*	62.91	*0.59*	2.14	*0.30*
	WEST		24.33	*4.51*	44.40	*6.15*	65.19	*8.66*	1.56	*0.11*
		Anual	20.55		95.55		115.75		1.21	
		Fall	16.47	*8.78*	57.96	*18.26*	79.64	*24.01*	1.44	*0.32*
		Spring	27.82	*7.68*	31.92	*9.31*	53.80	*18.17*	1.68	*0.16*
		Summer	19.88	*4.62*	45.40	*9.75*	68.51	*13.58*	1.61	*0.20*
		Winter	48.79	*29.72*	31.35	*9.29*	40.52	*9.49*	1.39	*0.17*
sensors			**143.22**	*28.21*	**229.91**	*25.57*	**329.21**	*29.91*	**2.11**	*0.26*
	EAST		-236.12	*19.83*	429.01	*69.27*	192.90	*61.66*	0.41	*0.08*
		Summer	-236.12	*19.83*	429.01	*69.27*	192.90	*61.66*	0.41	*0.08*
	WEST		182.12	*23.20*	207.78	*24.98*	341.60	*31.61*	2.31	*0.27*
		Fall	133.62	*47.09*	151.13	*30.72*	185.86	*36.18*	1.59	
		Spring	-62.19	*47.08*	444.88	*76.32*	382.68	*41.28*	0.89	*0.46*
		Summer	228.98	*25.05*	187.77	*26.82*	408.11	*52.01*	2.68	*0.08*
		Winter	182.52	*30.52*	125.40	*23.30*	325.75	*26.41*	2.65	*0.35*

The amount of CR estimates assessed with multiparametric probes (66) and benthic chambers (67) was similar whereas GPP was more often estimated using multiparametric sensors (74) compared to benthic chambers (67). For net community productivity (NCP) benthic chambers (99) were preferred over sensors (69 measurements). However, we found negative respiration rates (oxygen production) at night for many sensor deployments, an indication for the influence of lateral advection and passing of different water masses. Therefore, we trimmed the dataset to contain only measurements where this influence was not detected (see number of measurements in each Figure). Since benthic chambers and multiparametric sensors yielded such different CR with 41.2 $\pm$ 4.55 (SE) mmol O_2 m^{-2} d^{-1} for benthic chambers and 229.9 $\pm$ 25.57 mmol O_2

305    $m^{-2}\,d^{-1}$ for sensors ($t_{df=84.86}=9.57$, $p<0.0001$), with almost an order of magnitude difference for NCP as well (18.8 ± 3.80 and 143.2 ± 28.21 mmol O_2 $m^{-2}\,d^{-1}$ for benthic chambers and sensors respectively; $t_{df=25.61}=2.78$, $p<0.001$) and GPP (55.3 ± 6.39 and 329.2 ± 29.91 mmol O_2 $m^{-2}\,d^{-1}$ for chambers and sensors; $t_{df=101.05}=11.14$, $p<0.0001$) in a mixed model, with as only factor methodology and as random effect study (Fig. 2) we decided to analyse the metabolic rates estimated using benthic chambers and multiparametric sensors separately.

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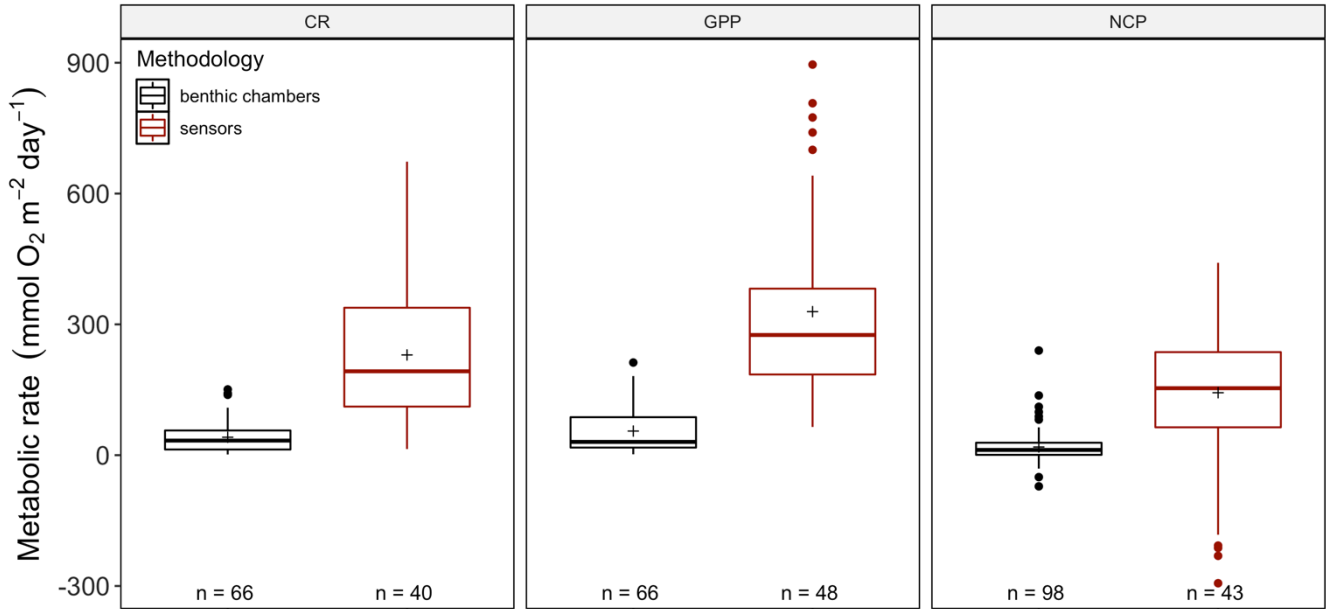


Figure 2. Comparison between benthic chamber- (black) and sensor data (red) for CR, GPP and NCP (mmol O_2 m^{-2} day^{-1}) for the full
315   dataset. Upper and lower hinges correspond to the upper and lower quartiles, the line inside the boxes correspond to the median, the mean is marked by a crosshair and the error bars are based on minimum and maximum standard deviation for each parameter. We found differences for CR, GPP and NCP ($p<0.0.001$).

320   **3.1 Multiparametric sensors**

Sensor data were collected in the water column, with lateral movement between habitats of water masses, and there were no significant differences, in GPP ($t_{df=31.75}=-0.16$, $p=0.87$), CR ($t_{df=32.46}=0.91$ $p=0.37$) and NCP ($t_{df=32.30}=0.21$, $p=0.84$), between the two species (*P. oceanica* and *C. nodosa*), tested in a mixed model with "Site" as random factor, including depth, region and seasons. Therefore, we didn´t divide the sensor data for the two species. GPP ($t_{df=2.93}=0.84$, $p=0.47$), and CR ($t_{df=3.78}=-$
325   1.37, $p=0.25$) were similar between the Eastern and Western Mediterranean basins (Fig. 3, Table 1), but NCP was higher the Western basin ($t_{df=3.63}=3.25$, $p=0.04$). No significant influence of depth was identified for any of the metabolic parameters, nor season for CR or GPP, however NCP was lower in spring ($t_{df=17.72}=-3.60$, $p<0.01$). In spring, we also found the highest GPP rates (Mean $\pm$ SE) with 382.68 ± 41.28 mmol O_2 m^{-2} day^{-1}, however, the corresponding CR rates for spring were 444.88 ± 76.32 mmol O_2 m^{-2} day^{-1} respectively, leading to a negative NCP of -62.19 ± 47.08 mmol O_2 m^{-2} day^{-1}. For the rest of the
330   seasons productivity was higher than respiration, reflected in positive averaged NCP rates and confirming that seagrass meadows normally tend to be autotrophic ecosystems with a mean P/R ratio above 1 (2.1 ± 0.26), confirming the tendency of net autotrophy. In the Eastern Mediterranean basin, only data recorded with sensors in summer was available to compare to measurements in the Western basin. Average GPP in the Western basin in summer was 408.11 ± 52.01 (SE) mmol O_2 m^{-2}

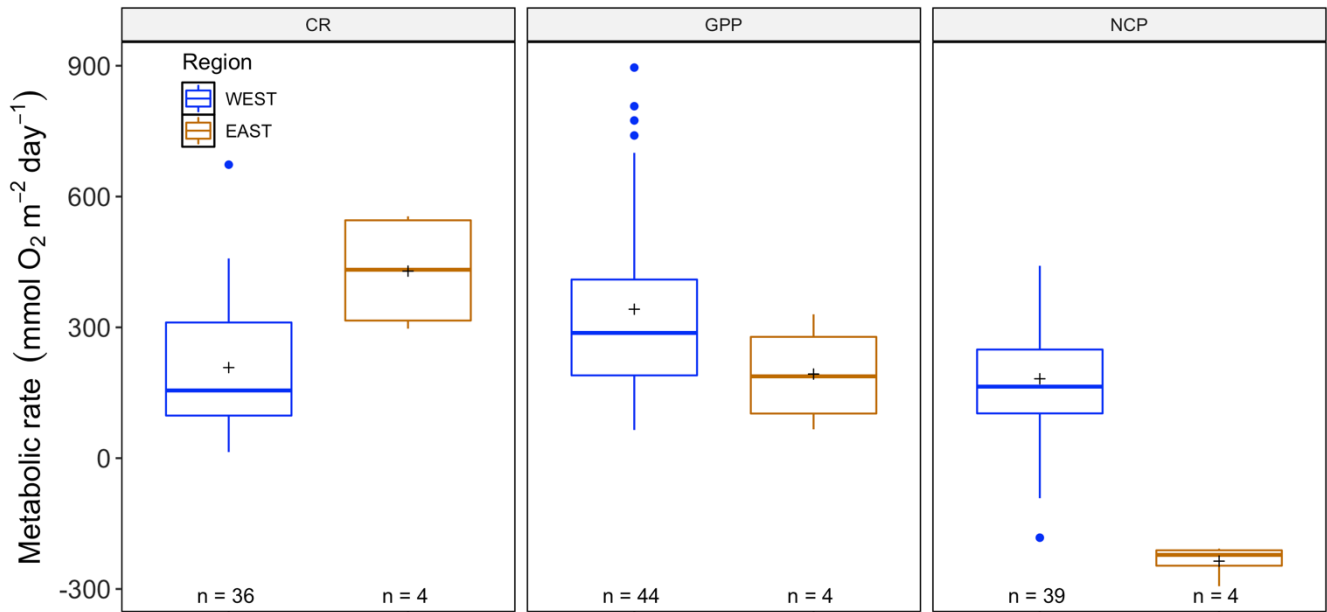
day^{-1}, and while the average yearly GPP did not differ between basins, the productivity in summer is significantly different
335 from the Eastern basin, with a lower average of 192.90 ± 61.66 mmol O$_2$ m^{-2} day^{-1}. Average NCP in the Eastern basin
(summer) was negative with -236.12 ± 19.83 mmol O$_2$ m^{-2} day^{-1}, based on only 4 measurements, and in the Western basin
positive with 182.12 ± 23.20 mmol O$_2$ m^{-2} day^{-1} averaged over the year and 228.98 ± 25.05 mmol O$_2$ m^{-2} day^{-1} in summer.
The negative NCP rates in the Eastern basin are the result of the high CR measured there, with 429.01 ± 69.27 mmol O$_2$ m^{-2}
day^{-1}, more than twice the CR measured in summer for the Western basin of 187.77 ± 26.82 mmol O$_2$ m^{-2} day^{-1}. This led to
340 low P/R ratios, with on average 0.41 ± 0.08 in the Eastern basin in summer. In the Western basin in summer, the seagrass
communities tended to be net autotrophic, reflected in an average P/R ratio of 2.31 ± 0.27. In general, for both basins, the
threshold GPP where the ecosystem switched between net heterotrophy and net autotrophy (NCP=0) was 254.03 mmol O$_2$ m^{-2} day^{-1}.

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Figure 3. Averaged GPP, NCP and CR (mmol O$_2$ m^{-2} day^{-1}) values for Eastern (blue) and Western (orange) Mediterranean basins calculated from measurements of multiparametric sensors. Upper and lower hinges correspond to the upper and lower quartiles, the lines inside the boxes correspond to the median, the mean is marked by a crosshair and the error bars are based on minimum and maximum standard deviation for each parameter. We found differences with p<0.05 for NCP between the regions.

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The temperature recorded during the highest GPP measurement in *P. oceanica* in the Western basin was 26.6ºC, which is
close, even though a bit higher, to the optimal value reported for *P. oceanica* of 25.8 ºC (Savva et al., 2018). For the Eastern
Mediterranean basin, the highest GPP obtained was 329.94 mmol O$_2$ m^{-2} day^{-1} at Maridati (Crete) during July and the *in situ*
temperature registered at that moment was 25.0 ºC, which was not the highest temperature registered in the Eastern basin
355 (28.5ºC) and lower than the mean temperature in the Eastern basin during the summer sampling campaign (25.9 ± 0.8 ºC).
The lowest GPP values found in the Western and Eastern regions were similar, we found the lowest value for GPP of 64.72
mmol O$_2$ m^{-2} day^{-1} for the Western basin in the the bay of Pollença (Mallorca) during fall whereas the lowest GPP value in
the Eastern basin was 66.19 mmol O$_2$ m^{-2} day^{-1} in Maridati (Crete) in summer; temperatures during both measurements were
different with 2,5 degree Celsius of difference between them; 25.4 ºC in Maridati station (Crete) and 22.9ºC in Pollença
360 (Mallorca).

We tested with individual regression models for the effect of temperature, which did not significantly affect GPP ($t_{df=40.5}$=-1.22, p=0.23), however NCP increased ($t_{df=63}$=2.46, p=0.02) with increasing temperatures while, surprisingly, CR decreased ($t_{df=34.5}$=-2.24, p=0.03) (See Supplement, Fig. S1). Windspeed did not drive metabolic rates with $t_{df=10.7}$=-0.69, p=0.51; $t_{df=14.5}$=-0.86, p=0.41 and $t_{df=7.1}$=1.07, p=0.32 respectively for GPP, NCP and CR. Over the 12 year of data available for

365 sensor measurements (2007 – 2019), we found an increase of CR ($t_{df=12.8}$=2.82, $p<0.05$), but no significant changes in GPP ($t_{df=31.18}$=1.23, p=0.23) nor NCP ($t_{df=28.9}$=-0.57, p=0.57; Figure A6).

3.2 Benthic chambers

We found significant differences for CR ($t_{df=15.3}$=4.05, $p<0.01$, and GPP ($t_{df=15.3}$=5.2, $p<0.001$) but not NCP ($t_{df=22.7}$=1.19,

370 p=0.25; Fig. 4) between *P. oceanica* and *C. nodosa* productivity, in a model including depth and season, and study as random factor. At a seasonal scale, there were no significant differences for NCP, or CR for *C. nodosa* with NCP (χ^2=0.22, p=0.90), CR (χ^2=2.63, p=0.27) but there was for GPP, with higher values during summer (χ^2=6.29, p=0.04). GPP was lower than CR during all seasons for which measurements were available, reflected in the averaged NCP, with a negative rate (-9.2 ±5.29 SE mmol O_2 m^{-2} day^{-1}), revealing that the *C. nodosa* community tends to be net heterotrophic. The threshold GPP to

375 switch from a heterotrophic to an autotrophic system (NCP=0) was 23.12 O_2 m^{-2} day^{-1} for *C. nodosa*.
There were no significant differences between NCP (χ^2=3.87, p=0.42) and CR (χ^2=8.57, p=0.07) across seasons (with year as random factor) for *P. oceanica,* but there were for GPP (χ^2=17.43, $p<0.01$, Figure S2). As we did not have *C. nodosa* data for the Eastern Mediterranean basin we only examined *P. oceanica* to distil patterns between Eastern and Western Mediterranean regions. There were no significant differences for NCP ($t_{df=7.48}$=-0.32, p=0.76), GPP $t_{df=8.04}$=-0.65 p=0.54) or

380 CR $t_{df=8.04}$=-0.71, p=0.50) in *Posidonia* incubations between Eastern and Western regions (Fig. 5), due to the high variability between sites, which was incorporated in the model as a random factor. For the Western basin, averaged NCP was 24.33 ± 4.51 SE mmol O_2 m^{-2} day^{-1}, with the average GPP (65.19 ± 8.66 mmol O_2 m^{-2} day^{-1}) higher than the CR rate (44.4 ± 6.15 mmol O_2 m^{-2} day^{-1}), reflecting the tendency of *P. oceanica* communities to be net autotrophic. This was also the case for the Eastern basin, with average NCP even higher at 39.36 ± 5.61 SE mmol O_2 m^{-2} day^{-1}, average GPP (93.20 ± 17.67 mmol O_2

385 m^{-2} day^{-1}), higher than the CR rate (52.60 ± 11.73 mmol O_2 m^{-2} day^{-1}). The threshold GPP to switch from a heterotrophic to an autotrophic system (NCP=0) was 36.64 O_2 m^{-2} day^{-1} for *P. oceanica*, with both basins combined. Contrary to the sensor data, temperature was not correlated with any metabolic rate, with NCP ($t_{df=17.2}$=0.02, p=0.84), GPP ($t_{df=16.14}$=1.65, p=0.12), and CR ($t_{df=16.58}$1.29, p=0.21; Figure S1). For chamber incubations we found a tendency over time, albeit not significant, (See Supplement Fig. S3) of decreasing GPP ($t_{df=5.34}$=-2.39, p=0.06) and CR ($t_{df=3.02}$=-3.117, p=0.05) but not NCP

390 ($t_{df=7.96}$=0.40, p=0.70).
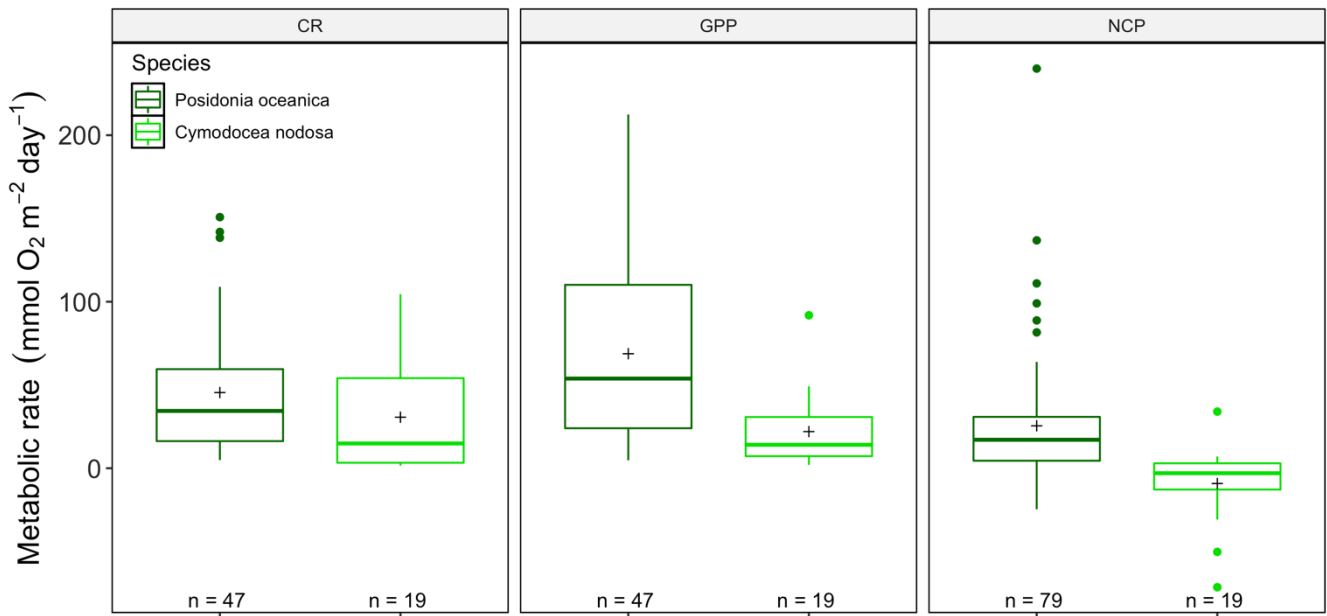
Figure 4. Metabolic rates GPP, CR and NCP (mmol O_2 m^{-2} day^{-1}) for *Cymodocea nodosa (*light green) and *Posidonia oceanica* (dark green) for the benthic chambers dataset in the Western Mediterranean basin. Upper and lower hinges correspond to the upper and lower quartiles, the line inside the boxes correspond to the median, the mean is marked by a crosshair and the error bars are based on minimum and maximum standard deviation for each parameter (Rates were different with *p*<0.001 for GPP and CR p<0.01).
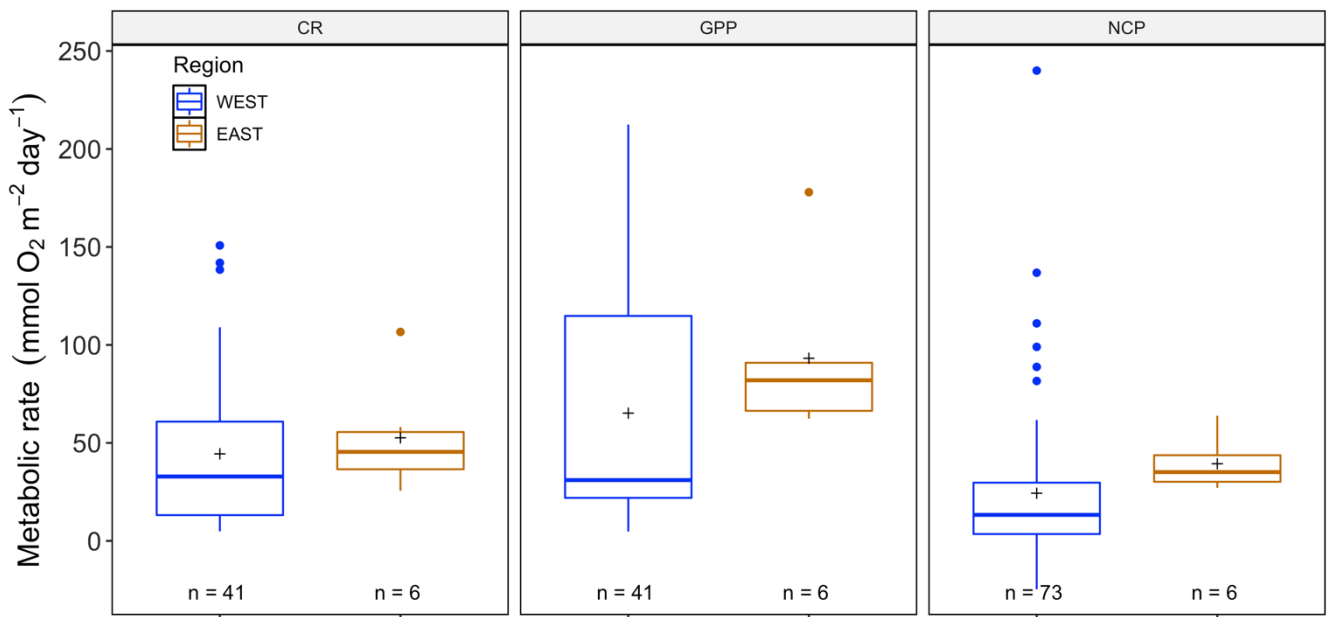
Figure 5. Metabolic rates GPP, NCP and CR (mmol O_2 m^{-2} day^{-1}) for the Eastern (blue) and Western (orange) Mediterranean as obtained from the benthic chamber dataset for *Posidonia oceanica*. Upper and lower hinges correspond to the upper and lower quartiles. The line inside the boxes correspond to the median, the mean is marked by a crosshair and the error bars are based on minimum and maximum standard deviation for each parameter.

4 Discussion

In this study we found significant differences in metabolic rates between *Cymodocea nodosa* and *Posidonia oceanica* in benthic chambers, while no species-specific pattern could be distilled with sensor data. Only significant differences in NCP estimated with sensor data were found between the Eastern and Western Mediterranean basins, which might be due to the

persistence of a much lower number of observations in the Eastern basin, with subsequent loss of statistical power to detect differences. The bias toward a higher number of observations in summer compared to other seasons could have prevented detection of seasonal patterns, with only lower NCP in spring detected with sensors and a higher GPP in summer with benthic chambers and similar metabolic rates for the rest of the seasons. Over half of NCP measurements was done in summer with 55.1% for benthic chamber data and 69.8% for sensor data. Over three quarter (80.4%) of the data for *P. oceanica* and *C. nodosa* communities show that these seagrass meadows are net autotrophic in almost all seasons and locations and are capable to act as carbon sinks and modify pH on diurnal an annual time scale. When this productivity is buried as biomass, carbon could be fixed over long time scales, highlighting the role of these seagrass meadows in climate change mitigation. The threshold GPP where the ecosystem switched between net heterotrophy and net autotrophy (NCP=0) was 254.03 mmol O_2 m^{-2} day^{-1} for measurements from sensors, which is higher than the estimate of 186 mmol O_2 m^{-2} day^{-1} for different seagrass species and regions pooled by Duarte et al. (2010). For estimates from benthic chambers, this threshold was much lower, 36.64 O_2 m^{-2} day^{-1} for *P. oceanica* and 23.12 O_2 m^{-2} day^{-1} for *C. nodosa*.

4.1 Multiparametric Sensors

Due to the effect of lateral advection and mixing of water masses, masking a species-specific signal, it was not possible to demonstrate differences in the metabolic rates between *Cymodocea nodosa* and *Posidonia oceanica* meadows since the sensors are measuring a composed signal in the water column. Due to logistic constraints, sensor deployment in *C. nodosa* and *P. oceanica* meadows when both seagrass species were present in a same site, was separated by a distance of less than 10 m, which adds to the lack of differentiation of values found between the two studied species. Measurements with multiparametric probes should therefore be interpreted as measurement at an ecosystem level as the influence of oxygen dynamics of macrophytes near the measuring site cannot be separated. The influence of phytoplankton and other primary producers may affect sensors as well as benthic chamber measurements. During a spring bloom a relationship between Chl *a* in the water column and GPP has been shown for the bay of Palma (Gazeau et al., 2005). However, not always a correlation with planktonic Chl *a* can be demonstrated; for instance, during a study of annual patterns in the Bay of Revellata (Italy), the highest GPP values recorded in a *P. oceanica* meadow where found when planktonic Chl *a* was particularly low and the highest values of Chl *a* did not reflect an increase in GPP and NCP values (Champenois et al. (2012). Sensor measurements in meadows of both seagrass species allowed us to estimate the metabolic activity of the whole ecosystem and compare between regions, showing similar rates for CR and GPP between the Eastern and Western Mediterranean basins, while NCP was higher in the Western basin, with negative values for the Eastern basin (Table 1). This difference is caused by the relatively high CR rates in the Eastern basin of on average 429.0 ± 69.27 (SE) mmol O_2 m^{-2} day^{-1}, two times higher than the averaged CR rate in the Western basin 207.78 ± 24.98 (SE) mmol O_2 m^{-2} day^{-1}. No significant differences were found ~~for~~ GPP, which has relatively low values in the Eastern basin of 192.90 ± 61.66 (SE) mmol O_2 m^{-2} day^{-1} compared to a yearly average of 341.60 ± 31.61 (SE) mmol O_2 m^{-2} day^{-1} for the Western Mediterranean due to the low sample size (n=4), limited measurements over different seasons and high variability of measurements in the East. When we use GPP of the Western basin in summer only, there is a significant difference in GPP between the basins. More measurements are necessary in the Eastern basin, during all seasons in order to obtain a more robust comparison.

The highest GPP was recorded in Cap Enderrocat (Mallorca, Spain) during summer in a *Cymodcea nodosa* meadow with a production of 895.78 mmol O_2 m^{-2} day^{-1}. Even if these values are high, this is lower than the 1338.0 mmol O_2 m^{-2} day^{-1} measured at the bay of Revellata (Corsica) by Champenois et al. (2012). These authors suggest that extreme GPP values in *P. oceanica* meadows may be rare events that are hardly captured by the classic benthic chambers methodology and the presence of high values measured with multiparametric sensors in this database might reflect that these events could be less uncommon than previously thought. We find the highest GPP values during the summer months, and the highest CR rates in spring (Fig. A5), with a positive relationship with increasing temperatures up to an optimum temperature (Fig. S1),

confirming that increasing temperatures enhance metabolic rates (Brown et al., 2004) until a threshold is reached. Abiotic and biotic factors that drive seagrass community metabolism differ between regions. In the Eastern basin, GPP was affected by temperature ($p<0.05$, Fig. S1). No effect was found of water depth on metabolic rates, while depth determines light availability, which in turn determines seagrass distribution, biomass and productivity (Dennison, 1987). This lack of effect is probably due to the limited depth range of the sites in the database (max. 22m, min. 0.6m depth), with measurements mainly in shallow sites, with an average depth of 8.7 ± 0.61 (SE) m. There was a relation between temperature and NCP and CR for the sensor data, with increasing Net Production with rising temperatures, but surprisingly lower Community Respiration rates.

4.2 Benthic chambers

Due to the effective enclosure with benthic chambers, it was possible to assess the productivity of *Posidonia oceanica* and *Cymodocea nodosa* communities separately, confirming higher CR and GPP in *P. oceanica* communities (Fig. 4). Estimates are for the whole community and include epiphyte and bacterial communities associated to each seagrass species. The higher Gross Productivity and Community Respiration, calculated per surface area (m^2) is logical as *P. oceanica* in general has a higher biomass per m^2 and therefore is more productive relative to *C. nodosa.* Caution should be taken in interpreting these results, as no data for *C. nodosa* from the Eastern basin was available and the database contained a higher number of *P. oceanica* estimates (n=42 for CR, GPP; n=79 for NCP), compared to *C. nodosa* (n=19)

There was a tendency for GPP and CR to decrease over the progressing years (2001-2019) for incubations of *P. oceanica* (Figure A6), however, our analysis with mixed models and each study as random factor did not indicate significant differences. With a simpler model (linear model) the decrease in CR would have been significant ($t_{df=45}=-6.08$, $p<0.001$) as well as for GPP $t_{df=45}=-6.19$, $p<0.001$) but not for the longer time series (1982-2019) of NCP. The decrease in CR is in contrast with the increase in CR found through time with sensor data (2007-2019), these differences may be due to the fact that sensors also detect trends from other photosynthetic organisms, or limitations for seagrass communities in benthic chambers as water renewal is limited (Champenois et al., 2012, 2019). No differences could be demonstrated in NCP and CR for both species through the seasons, however GPP was higher in summer. There were remarkable individual differences between the two species with an average positive NCP for *P. oceanica*, with only positive values found in benthic incubations and clear net autotrophic communities during the sampling periods in *P. oceanica* meadows. These meadows appear more productive than *C. nodosa* meadows, in agreement with previous studies (Duarte et al., 2010b; Champenois et al., 2012, 2019), which had a higher incidence of heterotrophic communities.

4.3 Sensors vs Benthic chambers

Significant different GPP, NCP and CR were obtained for the different methods, with values almost an order of magnitude larger when estimated from sensor data (Table 1) compared to estimations from benthic chambers as previously reported by Champenois et al. (2012). This difference may be due to a possible underestimation of the metabolic rates assessed by the benthic chambers methodology or an overestimation associated to water renovation at the placement of sensors. There are some limitations linked to the methodology using benthic chambers as, even if most incubations use flexible material for the bags, which allows for movement and some mixing, there is no real interchange with the water column and nutrient limitation could occur. However the effect of oxygen or nutrient limitation should be limited when incubations are short (24h) (Barrón et al., 2009). The reduction in water motion could lead to the increase of the width of the diffusive boundary layer (DBL) between a seagrass leaf and the water column and slower exchange of nutrients and CO_2 with the water column,

since water velocity determines DBL boundary thickness (Enríquez et al., 2006; Hendriks et al., 2017). Another possible explanation for the underestimation in benthic chambers of metabolic rates could be the fact that the insertion of the base of the benthic chambers into the sediment may cut the roots and rhizomes but this should be considered a rare event as most of the biomass of the belowground tissues of *P. oceanica* and *C. nodosa* is located at deeper strata. For the seagrass physiology, rhizomes play an important role as they translocate resources between shoots (Marbà et al., 2002), affecting therefore seagrass metabolism if they are severed. Another reason that may explain a possible underestimation in metabolic rates is the fact that pH may increase, together with O_2 during the day; those two factors, together with a high irradiance conducts the Ribulose-1.5. biphosphate-carboxylase-oxygenase enzyme to change from carboxylase to oxygenase (Heber et al., 1996). Under this reaction there is a higher consumption of O_2 and a CO_2 exudation which may conduce to a lower GPP estimation from the change in O_2 (Champenois et al., 2012). Nonetheless, benthic chambers can provide measurements for meadows of a single species when enclosures are properly selected. Multiparametric sensors measure a composed signal affected by all surrounding photosynthetic (and heterotrophic) organisms and measured metabolic rates reflect an ecosystem estimation. Measurements with the eddy covariance methodology, providing estimates on shorter time scales and therefore limiting the effect of lateral transport of water masses, have led to estimations of NCP for *P. oceanica* ranging from 85 to 119 mmol O_2 m^{-2} d^{-1} in nearshore meadows (Koopmans et al., 2020). These estimates are closer to our estimates obtained from sensors, which on average were between 1.7 and 1.2 times higher, while the values obtained from benthic chambers are much lower, with Eddy Covariance estimates 4.5 – 6.3 times higher than the NCP obtained from benthic chambers. The used method should therefore be selected depending on the study objectives taking the focus of the study into account, whether it is more important to attribute productivity a certain species or surface or an ecosystem estimate is required.

Independent of the method used, two main limitations remain: the lower amount of available data within the Eastern Mediterranean and the higher sampling frequency during summer compared to other seasons. More than half (55% for benthic chambers and 70% for sensor data) of the data has been collected in summer, due to the logistic restraints of underwater observations. The lack of data for the Eastern basin is specifically urgent to solve, as climate change does not act homogeneously on the two regions and warming is faster for the Eastern basin (Amitai et al. (2020); Nykjaer (2009). Failing to collect actual data on metabolic rates could be detrimental for our knowledge on the rates of change in the Mediterranean in the future. We also highlight the lack of data for *C. nodosa* compared to *P. oceanica*. This is probably due to the fact that *C. nodosa* has not had a marked historical presence in many locations, especially in the Wester Mediterranean. Their current expanse and apparent less sensitivity to higher temperatures compared to *P. oceanica* convert *C. nodosa* meadows in potential important players governing coastal metabolic signals deserving a closer look into productivity of these meadows and controlling factors.

4.4 Seagrass metabolism and carbon burial

More than three quarters, 80.9% of the NCP values were positive, reflecting the strong capacity of seagrass meadows to act as carbon sinks, which is also exemplified by the high number of P/R ratios above one for both methodologies (Table 1). *P. oceanica* communities tended to be net autotrophic. The exception for this net autotrophic averaged NCP values was found in the Eastern basin, in Maridati (Crete, Greece) where the averaged NCP was -236.1 mmol O_2 m^{-2} day^{-1} during July 2017. This value is in agreement with the low decadal carbon burial rates and stocks measured by Wesselmann et al., (2021) at the same site. This station is located nearby a temporary stream and receive its discharges which implies an extra nutrient input into this area. In 1980's there was a massive sediment input (Wesselmann et al., 2021) that may have caused mortality and physiological stress to the seagrass meadow and this may be still reflected in the net heterotrophic metabolism observed 2-3 decades later. We also saw heterotrophic meadows in the bay of Pollença (Mallorca, Spain) during spring 2018 (on average -137.1 mmol O_2 m^{-2} day^{-1}). This sampling site is located nearby a nutrient source with high input of organic matter from the

s´Albufereta coastal lagoon. Additionally, the organic matter input from the nearby harbour and the sewage systems from urban areas and hotels may have affected the metabolism of the seagrasses in this site. Sampling for additional parameters like nutrient values in the water column and sedimentation of organic matter and nutrients at the sites of the metabolism measurements could add important information enabling the explanation of deviant metabolic rates. Seagrass ecosystems are mainly net autotrophic and hence act as carbon sinks but might be threatened and disappear due to high organic inputs and other anthropic threats (e.g., mechanic destruction of the seagrass meadows).

5 Conclusions

Seagrass metabolic rates (Net Community Production, Gross Primary Production and Community Respiration) are significantly different depending on the methodology used. The rates obtained with benthic chambers are lower than those obtained with multiparametric sensors. With the benthic chamber methodology seagrass metabolism at a species level can be compared, with demonstrated differences between *Posidonia oceanica* and *Cymodocea nodosa* for GPP and CR. With *P. oceanica* the more productive species compared to *C. nodosa*, but also with higher respiration rates. Multiparametric sensors can assess metabolism at an ecosystem level, and showed NCP was higher in the Western Mediterranean basin compared to the Eastern basin. Benthic chamber measurements could not demonstrate a link between temperature and metabolic rates, while sensor data showed an increase in NCP and an unexpected decrease in CR with higher temperatures. When we plot our data according to the year, we see an increase in CR calculated from sensor data while benthic chamber rates have been decreasing for CR and GPP over the years. However, care should be taken in interpreting these trends as this in inconclusive due to the variability in seasonal measurements and measurement sites.

There is a publication bias with a higher number of observations in the Western region and a more elevated number of observations for summer compared to other seasons. The possibility to deploy multiparametric sensors for longer periods than benthic chambers allows for longer observational periods and detection of high GPP values that could be underestimated due to possible nutrient limitation or missed due to the shorter incubation time of the benthic chamber technique. Sampling during different time periods confirmed the switch between negative and positive NCP values, with changes between autotrophy and heterotrophy in a same location during different periods of the year, therefore reinforcing the importance of monitoring during the whole year and not only summer. The high percentage of autotrophic meadows highlights their key role for climate change mitigation, by acting as carbon sinks through growth as well as through accumulation of allochthonous carbon through particle retention. Therefore, it is important to augment the knowledge on seagrass metabolism in regions and seasons where there is little data available to prevent the deterioration of seagrass meadows in the context of climate change where they play an essential role.

Code availability

565 Metabolic rates of the seagrass meadows were calculated using a modification of the model of Cole et al. (2000), implemented in MATLAB (version 7.5, the Mathworks Inc.) explained in detail in Vaquer-Sunyer et al. (2012)

Data availability

The full dataset will be publicly available at digital.csic.es/handle/xxxxx/xxxxxx.

570 **Executable research compendium (ERC)**

Sample availability

NA

575 **Supplement link**

Author contribution

Conceptual idea IEH and NM. Data collection MW, SF, RVS, IEH, NM. Analyses IEH, AEM, SF. Literature compilation IEH and AEM. All authors have contributed to the writing of the article.

580 **Competing interests**

The authors declare that they have no conflict of interest.

Disclaimer

Special issue statement

Sample availability

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