

1 Lagrangian-Eulerian time and length scales of mesoscale ocean chlorophyll
2 from Bio-Argo floats and satellites

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15 **Format of this document:**

16 Black = Referee Comments (posted online); Blue = Author Comments (posted online);

17 Green = Description of how manuscript has changed, following the Author Comments

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19 **Author Comments in response to Referee #1**

20 The manuscript analyses decorrelation in time and space from both a Lagrangian and Eulerian
21 perspective with the ultimate aim to estimate how well Argo float act as Lagrangian platforms.
22 The motivation for the paper is sound and it addresses some very important questions. I'm
23 excited to use the results from a published version of the MS in future studies and believe it to
24 have a wide potential utility. There are, however, a couple of major questions/concerns I need
25 resolved before recommending publication.

26 Thank you for your close read and evaluation of our manuscript.

27 Before we address each of your comments individually, we would like to preface with an
28 overview. The matters you bring up in your comments 1-3 are related by a single overarching
29 point that we perhaps did not make clear enough in our original manuscript. Our intention in this
30 study was to analyze time- and length-scales of mesoscale ocean chlorophyll variability (and
31 velocity). This choice of scale dictated our choices of data products and filtering. Given this, our
32 responses to your comments 1-3 are related.

33 1. The first equation suggests, to my understanding, that the Chl field is fixed in in space. This is
34 a bold assumption that needs to be carefully motivated. I would have expected the advection
35 decorrelation term to be applied to the Eulerian observer since Chl is advected with the velocity
36 field. One could possibly argue that biomass might originate from stationary processes at for
37 example seam mounds, but this is rather the exception than the rule. As a consequence, I expect
38 that a Lagrangian sampling platform in general, with some specific exceptions, experiences
39 longer temporal decorrelation time scales compared to the Eulerian observer. I'm willing to
40 admit that I might have misunderstood Eq1 and the reasoning around it, but I don't think I'm the
41 only one if so. This need to either be explained better or changed.

42 Equation (1) is simply a material derivative “budget” for chlorophyll and does not make any
43 assumptions about the properties of the chlorophyll field. In our reading, nothing about the
44 presentation of Equation (1) prescribes a behavior for the chlorophyll field. When the equation is
45 scaled (equation (5)), Eulerian scales are indeed used for the advection term. The discussion that
46 introduces Equation (1) (lines 84-107) is based on the velocity field, for which the theoretical
47 and observational studies cited show that, for velocity, it is true that $T_L < T_E$. However, prior to
48 conducting this study, it was not known whether there was any systematic relationship between
49 $T_{L,Chl}$ and $T_{E,Chl}$. By scaling Equation (1) with scales derived from the chlorophyll fields and
50 Lagrangian or Eulerian chlorophyll time series, we are able to consider how the movement of an
51 observer relative to “movement” of the chlorophyll fields (u' versus $L_{E,Chl}/T_{E,Chl}$) influences the
52 Lagrangian decorrelation time.

53 Like you, we were surprised to find $T_{L,Chl} \leq T_{E,Chl}$. We suspect there are several possible origins
54 for this behavior. Firstly, this may be the manifestation of an observer moving across existing
55 gradients in the chlorophyll field, as would happen when a mesoscale eddy stirs a horizontal
56 gradient. The empirical curves in Figure 6 (from equation 11) would support this, as described in
57 Section 4.5. Another possibility is that chlorophyll may actually be conserved for longer along a
58 trajectory than our results would indicate: if patches are organized in small scale filaments that
59 are not fully resolved in an 0.25° product, the inability for a drifter-projected time series to
60 resolve such near-constant chlorophyll levels along a filament will result in an early temporal
61 decorrelation. The result that the ratio $T_{L,Chl}/T_{E,Chl}$ is approximately 1 relative to the smoothed
62 subtrahend (where sub-pixel variability probably dominates) while the ratio is less than 1 relative
63 to the climatology subtrahend (where larger and/or slower processes dominate) supports this
64 interpretation. We plan to add the preceding discussion to the manuscript.

65 **Changes made:** We added a section 4.6 which discusses how the relationship between $T_{l, Chl}$ and
66 $T_{e, Chl}$ may depend on the resolution of satellite data (basically, that these are results of mesoscale
67 variance). We did not change the presentation of Equation (1) after carefully reviewing section 2
68 and concluding that as currently written there is no presupposition of a relationship between $T_{l,}$
69 chl and $T_{e, chl}$ nor an assumption that the chlorophyll field is fixed.

70 2. The use of Chl fields with a 0.25° spatial resolution and the removal of sub- and mesoscale
71 variability weakens the study significantly. It is abundantly clear that submesoscale processes are
72 of first-order importance in controlling the variability of Chl, as mentioned in the MS and cited
73 publications by Amala Mahadevan or Marina Levy. A general analysis of decorrelation time-
74 and length scales can get away with using coarser grids by defining the domain of interest
75 carefully but this study doesn't have that luxury. One specific aim, as I understand, is to evaluate
76 the utility of float which requires the use of the highest resolution possible. I would have
77 preferred that a 1km product had been used (OC-CCI at 1km is for example available from
78 Plymouth Marine Laboratory), but I understand if a 4km product is used out of necessity. Aren't
79 the results quite dependent of the rather arbitrarily chosen 0.25° pixel size? How much would the
80 results differ if 0.125° , 0.5 , or 1° pixels were used instead?

81 We acknowledge that submesoscale processes are of first-order importance in driving surface
82 chlorophyll variability. Further, we do believe that our results are dependent on the ocean color
83 pixel size. That being said, the choice of 0.25° was not arbitrary, and we believe our results are
84 still novel, meaningful, and useful. As indicated by the title (though perhaps the manuscript
85 needs to more clearly convey this), our interest is in studying the mesoscale chlorophyll field.

86 Our motivation for this interest is both a practical and intellectual matter.

87
88 As a practical matter, there is a tradeoff between resolving more variance and dealing with
89 increased gaps when moving to a finer resolution ocean color product. For the purpose of this
90 study, we chose to prioritize data coverage, leading us to select a blended, 0.25° product, and a
91 focus on the mesoscales. As an additional practical matter, given the relatively sluggish motion
92 of floats (as quantified in this study), they may not capture the full spectrum of submesoscale
93 processes. Given that we wanted to incorporate in-situ data in this study, we felt it best to focus
94 on mesoscale variations. The choice of product is also consistent with the grid size of the
95 Eulerian velocity field (0.25° altimetric geostrophic currents), and we aimed for consistency
96 since we compare the two variables.

97
98 Intellectually speaking, combined Lagrangian-Eulerian scales of chlorophyll are unknown at any
99 scale, and we believe that contributions at the mesoscale are useful. New results are still being
100 gleaned about geostatistics of the mesoscale chlorophyll field and their biophysical origin (e.g.,
101 Eveleth et al., 2021). We believe the results here stand on their own and our mesoscale study
102 may lay the groundwork for follow-up studies targeting the submesoscale, either utilizing a more
103 spatially or temporally expansive drifter and ocean color dataset or a high-resolution model.

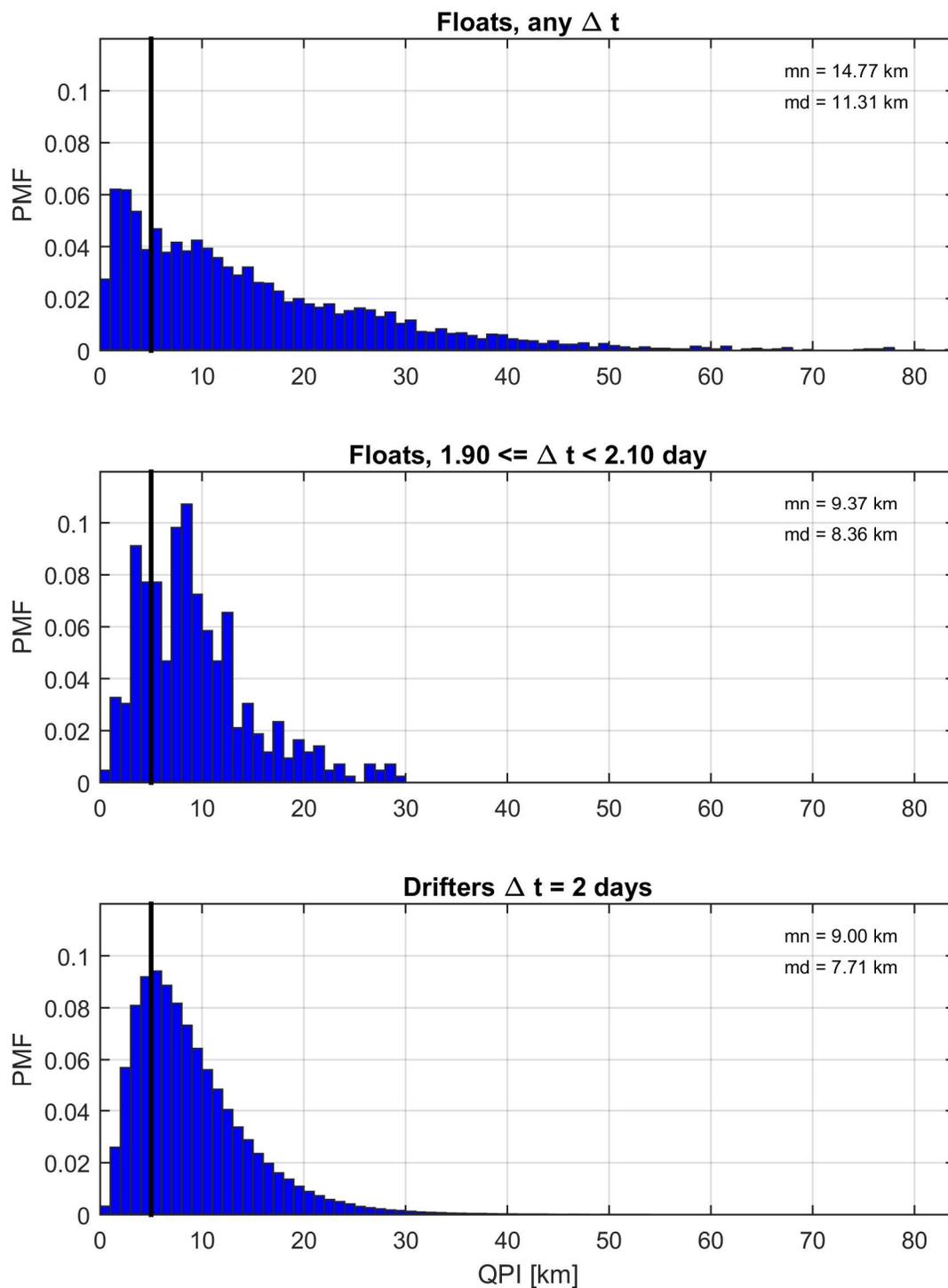
104
105 Finally, we do suspect that our results are dependent on the choice of ocean color product
106 resolution. We suspect that the major consequence is that $T_{L,Chl} \leq T_{E,Chl}$ for the reasons outlined
107 in our response to your comment (1).

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109 We plan to include a more detailed discussion of why a 0.25° product was utilized and
110 specifically delineate what the limitations of this choice are and how it likely influences our
111 results (incorporating the last paragraph of our response to your comment (1)). In a revised
112 introduction we plan to clearly motivate an analysis of mesoscale variability as done here, and in
113 a revised conclusion we plan to recommend subsequent studies of submesoscale variability as
114 done here.

115 Changes made: We continued to work with the 0.25° GlobColour fields. We updated the text in
116 the following manners. We updated the last paragraph of the introduction (Section 1) to clarify
117 that this is a study of mesoscale variance and to motivate that choice of scale. We updated
118 Section 3.2.2 to clarify our choice of product. We wrote a new subsection of Results and
119 Discussions (4.6) where we discuss how our results are influenced by the choice of data products
120 and filtering (following the main points of our first Author Comment) and point out that an
121 analysis of submesoscale-resolving data may lead to different conclusions. We updated the
122 conclusions (Section 5) to reinforce that our results are indicative of mesoscale variance and to
123 suggest that future studies of submesoscale statistics are warranted.

124 3. The use of geostrophic velocities to estimate QPI is problematic. There are many processes
125 that attribute to Lagrangian decorrelation missing from these fields- I'm not even sure if Ekman
126 drift is included? Many of these forces are also likely to affect the upper ocean to larger extent,
127 creating an even further biasing when being omitted. One easy test is to calculate QPI for the
128 drifters the same way as the floats to see how representative the geostrophic velocity fields are.
129 Another option is to conduct the exercise in a high resolution ocean model using virtual drifters
130 and floats.

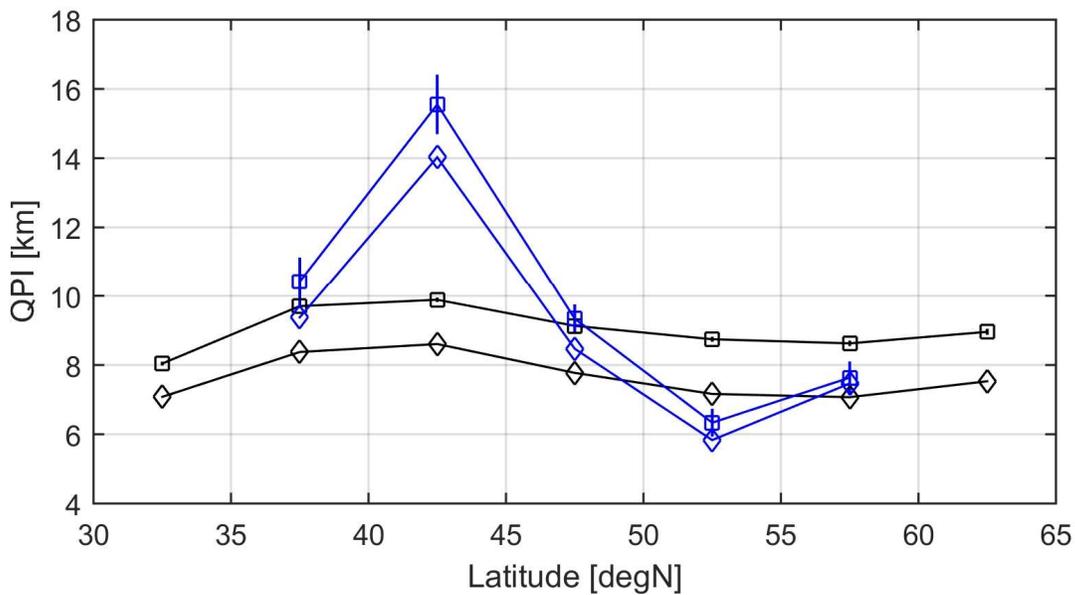
131 The QPI is calculated using trajectories computed from the global altimetry product, which
132 includes a geostrophic term based on sea level anomalies and nothing else. The study of Della
133 Penna et al. (2015) that developed the QPI compared distributions of QPI for SVP (real) drifters
134 using trajectories calculated from different altimetry products (see their Supplementary
135 Information Figure 4), including a global altimetry product (geostrophy only), a regional
136 altimetry product (geostrophy only), and a regional Ekman corrected product (geostrophy +
137 Ekman). Their conclusion was that “[using] different products does not alter significantly the
138 shape and the extent of the [distribution of QPI], yet differences in the distributions can be
139 observed in the tails”. Though their study was performed in the Southern Ocean and though they
140 compare trajectories in a slightly different manner than we do, we took this to mean a
141 geostrophic term would likely dominate the trajectories, especially in the vicinity of the Gulf
142 Stream and North Atlantic Current where a geostrophic balance is generally reasonable. Our
143 choice was further motivated by our desire to study mesoscale variations in the velocity fields,
144 which the altimetric geostrophic fields are known to capture reasonably well.



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 146 **Figure R1:** Probability mass functions (PMF) of QPI for all floats (top panel), floats with $\Delta t \approx 2$
 147 days (middle panel), and drifters with $\Delta t = 2$ days (bottom). Vertical lines represent 5 km.

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We computed the QPI for all of our drifter returns that fall in the study domain of [30N, 65N, 300E, 340E] and [2003-01-01, 2016-12-31] and used our daily subsampling, so that $\Delta t = 2$ days since trajectories are compared at $[t_{i-1}, t_i, t_{i+1}]$. The distributions of drifter QPI are shown in Figure R1 (third panel). We found two things surprising. Firstly, the median over all drifters is larger than expected at about 8 km. Secondly, the distribution of float measured QPI when restricting to profiles with $\Delta t \approx 2$ days is very similar (Figure R1, second panel), with only a slightly larger median. Inspecting QPI as a function of latitude (averaged in 5° bins) for the two platforms reveals that, while the distributions are similar when including all samples, the latitudinal variations are different. Each platform has a maximum near the Gulf Stream (40-45°N) and a minimum at 50-55°N or 55-60°N, but the QPI is more variable for the floats, with a maximum larger than that for drifters and a minimum that is smaller (even though, presumably, floats are less Lagrangian with respect to the surface flow) (Figure R2). As we know that floats tend to lag the surface flow, the energetic and sheared currents of the Gulf Stream may exaggerate this difference, causing the very large QPI there. On the other hand, deeper mixed layers and more sluggish currents at higher latitudes may cause a relatively smaller QPI for the floats at higher latitudes. Another possibility is that Ekman transports become important farther north away from the Gulf Stream, and the deeper floats are sheltered from this flow, instead primarily feeling the geostrophic flow and yielding a relatively smaller QPI compared to drifters, who feel the total current. However, given the relatively stable distribution of drifter QPI with latitude, we suspect lack of including an Ekman term amounts to a small difference, in line with the findings of Della Penna et al. (2015). Instead, deviations for the drifters are probably primarily due to sub-map-grid scale processes or altimetric geostrophic currents generally underestimating surface flow due to the finite differences being computed from a product that has been mapped from the actual altimetry swaths (Ascani et al., 2013; Sudre and Morrow, 2008).



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Figure R2: QPI as a function of latitude for drifters (black) and floats with $\Delta t \approx 2$ days (blue) in 5° latitude bins. Squares with vertical line are means ± 1 standard error. Diamonds are medians.

178 All that said, we still think that the QPI as we calculate it (deviations from a surface geostrophic
 179 trajectory) is reasonable for our study mainly because, as indicated (though we will clarify), we
 180 have set out to conduct a study of mesoscale Eulerian-Lagrangian time and length scales. For
 181 example, the velocity scale analysis (Figure 5) is based on Eulerian scales from satellite
 182 altimetric geostrophic currents and drifter velocity time series filtered to remove super-inertial
 183 variability. If we take that geostrophic altimetric fields are a reasonable approximation for the
 184 mesoscale flow, then it is reasonable to us to emphasize float segments whose trajectories are
 185 similar to trajectories subject to that flow.

186 **Changes made:** The above results from our Author Comment are summarized in a paragraph
 187 added to Appendix A (the appendix that explains the QPI). The QPI is unchanged.

188 4. I'm not happy with how the Chl data for the floats is handled. The mean Chl concentration in
 189 the mixed layer is not what is observed by satellite. This is of particular importance in regions
 190 with deep Chl maxima where most Chl is close to the base of the mixed layer and not visible
 191 from space. This issue can easily be amplified in this study if there is MLD variability over short
 192 timescales or if the isolines are sloping. Each case could lead to spurious variability in Chl
 193 observed by the float, compared to the drifters. The correct approach would be to use attenuation
 194 or PAR from the float (or Kd_{490} from satellite in if not available on the float) and average the
 195 Chl data down to the first optical depth. An even better approach would be to match satellite Chl
 196 to the floats the same way as to the drifters. I don't see any benefits in using In-situ observations
 197 for one platform and satellite-derived data for the other when comparing the two.

198 We have computed an alternate depth-reduced chlorophyll series from the floats that is meant to
 199 better approximate what the satellites see. About 90% of the satellite-measured chlorophyll
 200 signal in the open ocean comes from a depth of $1/Kd_{490}$, and it is exponentially weighted in the
 201 vertical (Gordon and McCluney, 1975). Our new approach is to conduct a weighted average over
 202 one attenuation depth. First, we estimate Kd_{490} from the floats following Morel et al. (2007)
 203 (their equation 8),

$$204 \quad Kd_{490} = 0.0166 + 0.0773[\text{Chl}]^{0.6715}, \quad (\text{R1.1})$$

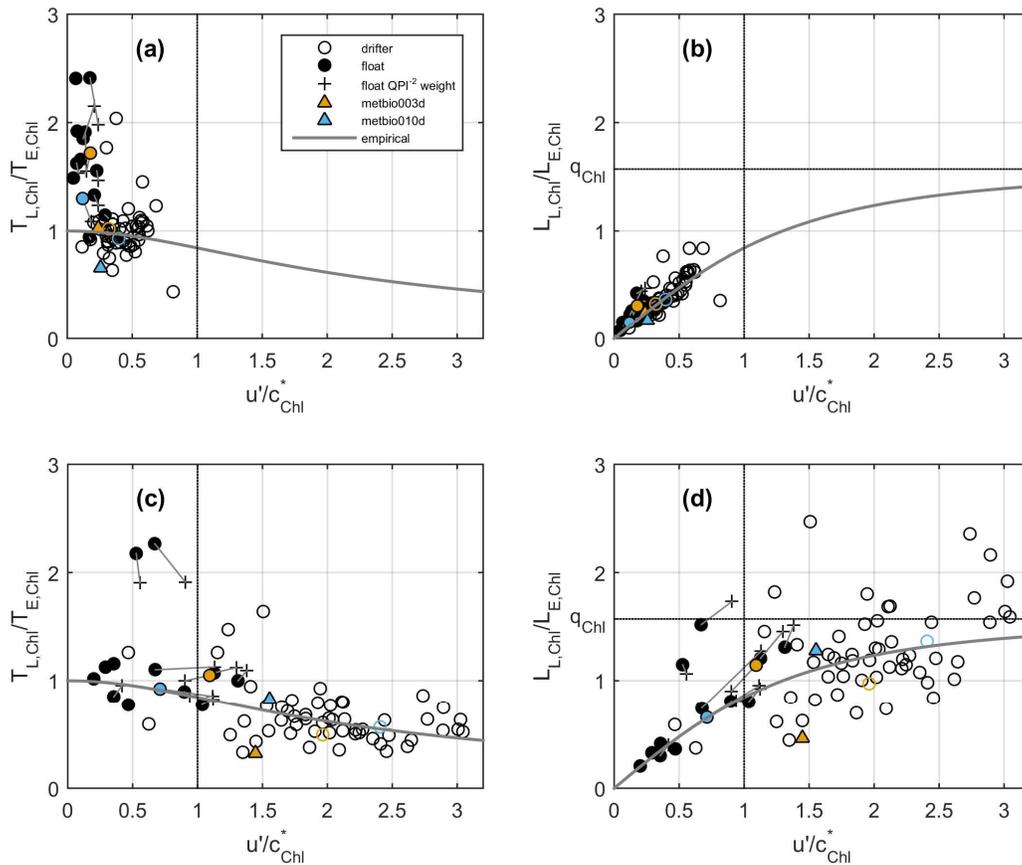
205 where we take $[\text{Chl}]$ as the mixed-layer average chlorophyll. Then, we take a weighted vertical
 206 average at each time step as

$$207 \quad \text{Chl}_{\text{float}}(t) = \frac{\sum_{z=1/Kd_{490}}^{z=\text{surface}} \text{Chl}(z,t) \exp(-2Kd_{490}(t)z)}{\sum_{z=1/Kd_{490}}^{z=\text{surface}} \exp(-2Kd_{490}(t)z)}. \quad (\text{R1.2})$$

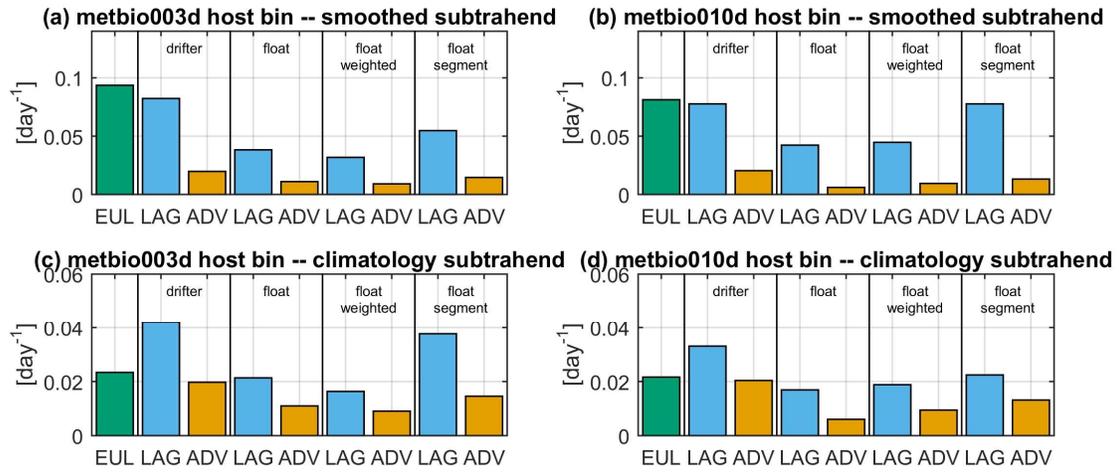
208 We use a weighted sum instead of integral because some profiles have missing data near the
 209 surface. The series is then log transformed and filtered as before. In general, the two time series
 210 (R1.2 and the MLD-average used in the original manuscript) are very similar with some
 211 discrepancies at daily to subdaily fluctuations; however, these are generally removed with the
 212 subdaily filter (compare Figure B1 to attached revised Figure R7). We then reran all scales and
 213 provide here a complete set of figures (equivalents to Figures 6-9, B1, D1). The results are not
 214 appreciably different. We would be fine with using the new method of depth reduction.

215 As for using float-measured data instead of projecting satellite data onto float trajectories, we see
 216 great value in using in-situ observations. Firstly, it is not possible to do this analysis with satellite

217 data projected onto the floats, given the limited number of floats available. As you can see in
 218 Figure B1 (squares), there are many gaps when projecting satellite ocean color onto a float
 219 trajectory, even when using a nearest neighbor approach as is done in that figure, which is more
 220 generous than the preferred bilinear interpolation used for drifters in the paper. This issue is not
 221 prohibitive when working with the drifters because there is a tremendous number of them, so
 222 even though individual segments are sparse and offer few pairs at a given lag, a composite ACF
 223 can be constructed from many sparse segments. Secondly, we feel it is instructive to demonstrate
 224 what can be learned from a near-continuous, in-situ time series, as this represents as complete a
 225 dataset as is possible and is what most float users will work with.

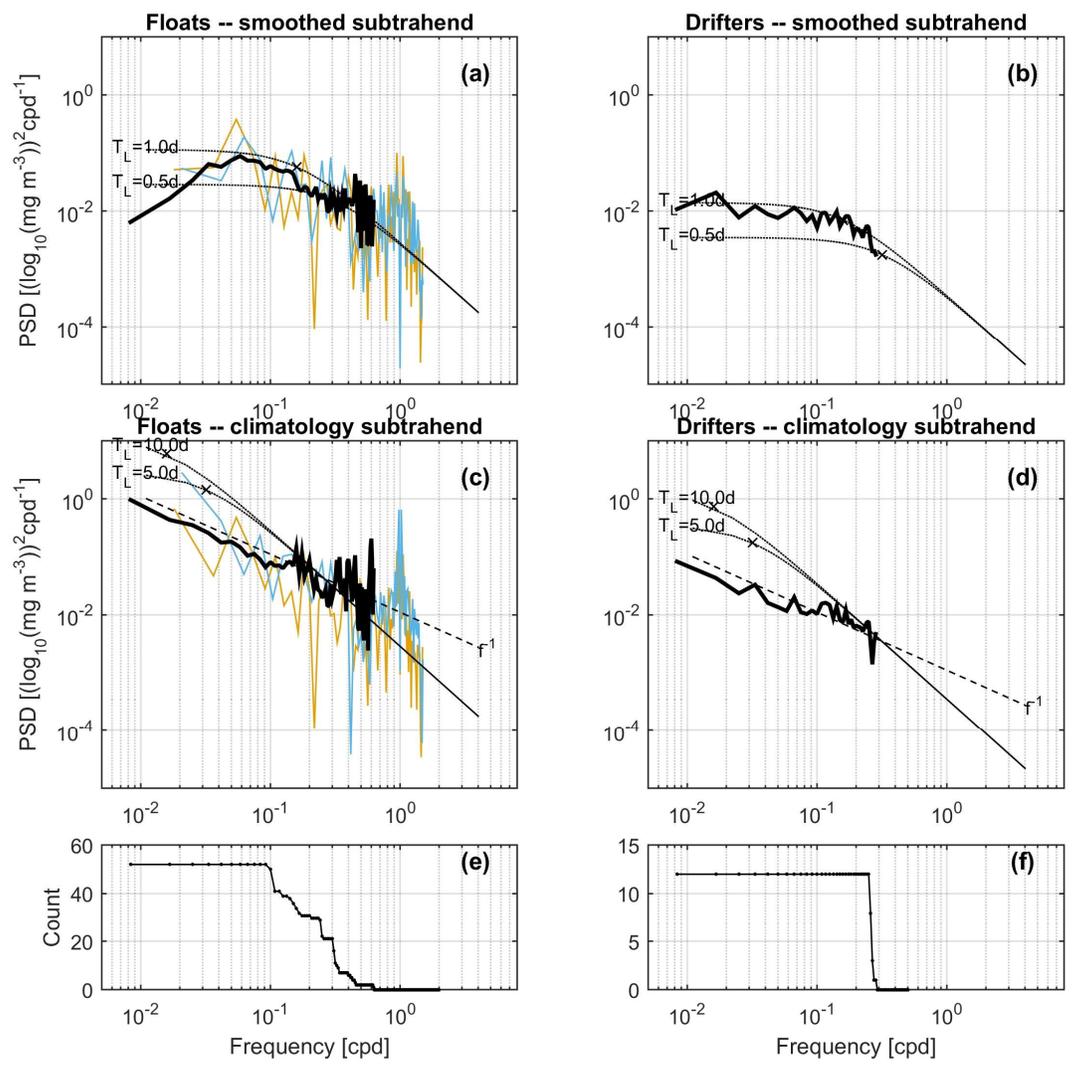


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 227 **Figure R3:** Revised Figure 6 from manuscript using new definition of float chlorophyll series.
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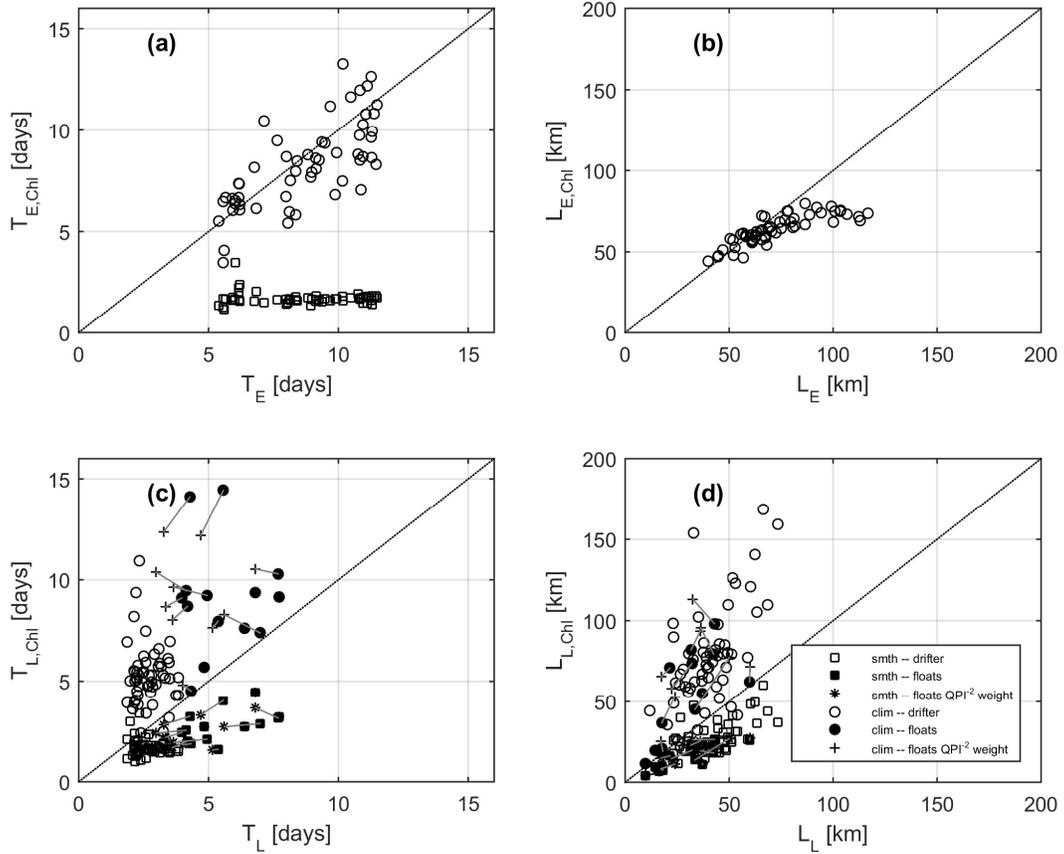
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Figure R4: Revised Figure 7 from manuscript using new definition of float chlorophyll series.



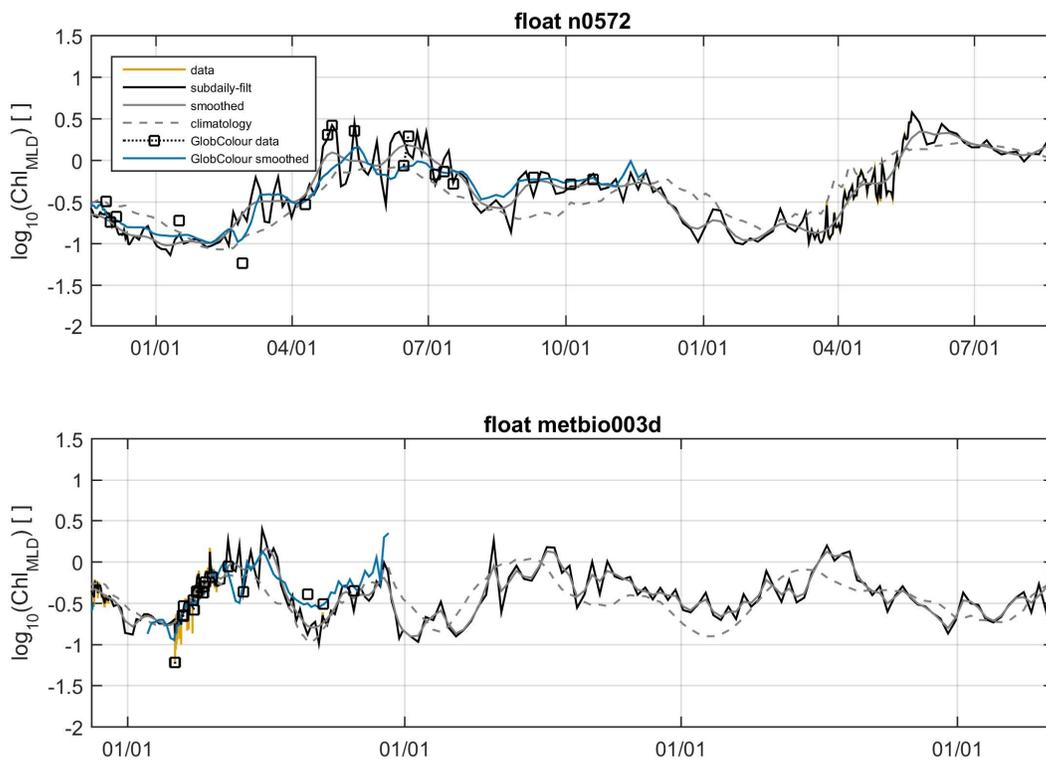
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Figure R5: Revised Figure 8 from manuscript using new definition of float chlorophyll series.



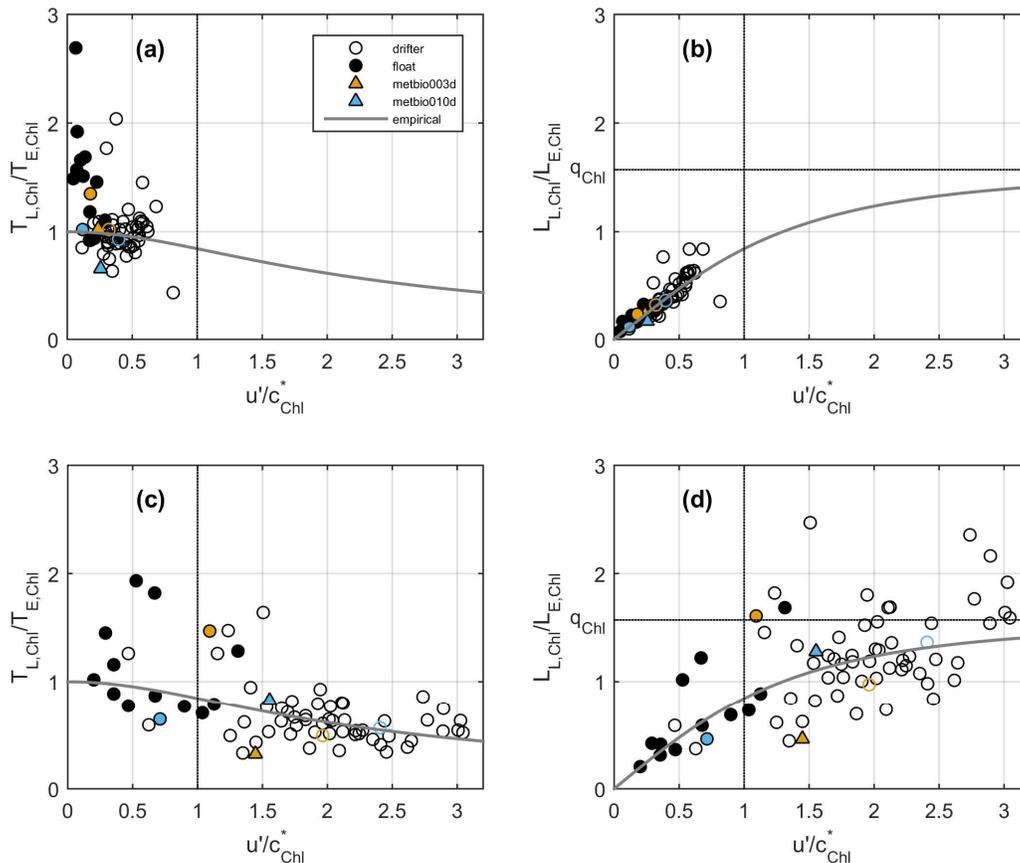
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Figure R6: Revised Figure 9 from manuscript using new definition of float chlorophyll series.



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Figure R7: Revised Figure B1 from manuscript using new definition of float chlorophyll series.



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 242 **Figure R8:** Revised Figure D1 from manuscript using new definition of float chlorophyll series.
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244 Changes made: We continue to work with in-situ float data out of necessity. The depth-reduced
 245 chlorophyll series for floats is now calculated using equations R1.1-R1.2 given in our Author
 246 Comment (instead of the MLD-average) and all calculations and figures are revised. Section
 247 3.2.1 is updated with a development and presentation of the equation (following our Author
 248 Comment). New section 4.6 mentions that the results and conclusions are not sensitive to this
 249 choice.

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 251 5. Finally, while the formalism in the MS is thorough and impressive, I think it might scare many
 252 potentially readers away. Cleaning up the text by explaining the reasoning in a way that can be
 253 understood by a wide audience and move a portion of the analytical description to an appendix
 254 would probably increase the readership statistics and potential of citations.

255 This is a good suggestion for such a technical paper. Already, in our initial submission we've
 256 made great effort to move non-essential information to appendices (there are already five). In our
 257 reading, though there is a deep theoretical exposition in Section 2.1 and a lot of methodological
 258 detail in Sections 3.3-3.5, we believe that the information retained in the main text is essential for
 259 evaluation of the paper. However, we agree that the exposition can be improved. Referee #2
 260 made some good suggestions on how to enhance the clarity of (and reduce length of) sections 3.3

261 and 3.4. We believe that by addressing those issues and by more clearly motivating our analyses,
262 readability will improve.

263 Changes made: Section 3.4 is shortened and much simpler. Section 3.3 was updated for clarity.
264 No new appendices were added.

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266 Author Comments in response to Referee #2

267 This manuscript presents extensive work evaluating Eulerian and Lagrangian time and length
268 scales of velocity and chlorophyll, as well as discussion about how they correlate. The proper
269 interpretation of drifting phytoplankton observed in a Eulerian fashion is a longstanding
270 paradigm in ocean ecology. However, estimates of Lagrangian phytoplankton statistics and
271 comparisons with Eulerian counterparts are rare. This study represents a significant contribution
272 towards best understanding how to interpret phytoplankton/chlorophyll measured in both
273 Eulerian and Lagrangian platforms. The authors are very thorough in their analysis and
274 description of the results. Nonetheless, I have a few comments to be addressed prior to
275 recommending publication.

276 Thank you for your close read and evaluation of our manuscript.

277 Major comment: There are several data limitations that guide methodological decisions in an
278 analysis of this type (e.g., the broad spatial averaging, chlorophyll averaging in the MLD). While
279 some of the issues arising from these are mentioned briefly throughout the text, I would prefer to
280 see a dedicated discussion section with the limitations and caveats.

281 The averaging of scales (or compositing of ACFs) over $[5^\circ \times 5^\circ]$ space bins is meant to enhance
282 the quality of the estimates by averaging over a region that is relatively spatially homogenous.
283 Other authors doing a similar analysis of velocity in this region used $[10^\circ \times 10^\circ]$ space bins and
284 found this adequate to describe spatial variability in Lagrangian scales (Lumpkin et al., 2002).
285 We chose to use the same $[5^\circ \times 5^\circ]$ space bins as Glover et al. (2018), who calculated variograms
286 of satellite ocean color in each bin much like we compute ACFs and found these bins good to
287 resolve spatial variability. We will better motivate this in the text.

288 As for the depth reduction of chlorophyll, we had indeed used a simple average over the mixed
289 layer since other authors had done this and demonstrated good agreement with satellite ocean
290 color when describing seasonal variability in the region (Yang et al., 2020). At the suggestion of
291 Referee #1, we computed an alternate depth-reduced chlorophyll series from the floats that is
292 meant to better approximate what the satellites see. Please refer to our Author Comment to
293 Referee #1 for full details, but briefly, we utilize the fact that 90% of the satellite-measured
294 chlorophyll signal in the open ocean comes from a depth of $1/Kd_{490}$, and it is exponentially-
295 weighted (Gordon and McCluney, 1975). We estimate Kd_{490} from the floats following Morel et
296 al. (2007) (their equation 8),

$$297 Kd_{490} = 0.0166 + 0.0773[\text{Chl}]^{0.6715}$$

298 where we take $[\text{Chl}]$ as the mixed-layer average chlorophyll, and then take a weighted vertical

299 average at each time step as

$$300 \text{Chl}_{\text{float}}(t) = \frac{\sum_{z=1/Kd_{490}}^{z=\text{surface}} \text{Chl}(z,t) \exp(-2Kd_{490}(t)z)}{\sum_{z=1/Kd_{490}}^{z=\text{surface}} \exp(-2Kd_{490}(t)z)} .$$

301 The series is then log transformed and filtered as before. After rerunning all scales (please refer
302 to set of figures in Author Comment to Referee #1), the results are not appreciably different. We
303 will include a description of this comparison in the text.

304 As for the choice of ACF parameters in Table 1, please refer to our response to your comment
305 below.

306 We plan to consolidate all of the above matters (spatial averaging, depth averaging, ACF
307 parameters) into a subsection of the Discussions, as you suggest.

308 Changes made: At your suggestion, we wrote a new Section 4.6 (Methodological decisions) that
309 includes all the issues mentioned in the Author Comment and how they might influence our
310 results: depth-reduction of float data, ACF parameters and spatial averaging, choice of ocean
311 colour product and filtering.

312 Specific comments:

313 I find that, while technically correct, talking about Lagrangian-Eulerian “statistics” in the title
314 and throughout the text can be misleading. Why not refer to the specific statistics that are
315 included in the analysis? i.e., Lagrangian-Eulerian time and length scales.

316 We felt that use of the word “statistics” made for a more compact title, with the meaning
317 becoming clear after reading the abstract. But we do not object to changing the title to:
318 “Lagrangian-Eulerian time and length scales of mesoscale ocean chlorophyll from Bio-Argo
319 floats and satellites”.

320 Changes made: We suggest that the title be changed (substituting “time and length scales” for
321 “statistics”) if the Editor allows. On a few occasions, we continue to use “statistics” for brevity,
322 noting that the first sentences of the Introduction (Section 1) and Conclusions (Section 5) define
323 our use of “statistics” as “time and length scales”.

324 The notation of upper case L for both Lagrangian and length-scale can be a bit confusing. I
325 suggest using upper and lower case or a different notation to improve readability.

326 We agree about the confusion. We thought about using upper and lower case letters but this can
327 become problematic since a lower case “l” can look like a number 1 or capital “I” or something
328 else. We propose to maintain “T” and “L” for scales and replace subscripts “L” and “E” with
329 either “l” and “e” or “LAG” and “EUL” for Lagrangian and Eulerian, respectively, depending on
330 which of the two looks best.

331 Changes made: All subscripts “E” for Eulerian and “L” for Lagrangian are changed to “e” and
332 “l”, respectively, in all text, equations, and figures.

333 Equation 5. Terminology becomes confusing here too when calling the nominators Lagrangian,
334 Eulerian and Spatial (chlorophyll) scales. Is there a different name that could be more
335 appropriate and less confusing? This is essentially a change in chlorophyll, correct?

336 These are effectively standard deviations of chlorophyll computed in different frames: from
337 Lagrangian time series (subscript “LAG” or “l”) from Eulerian time series (subscript “EUL” or
338 “e”), or from spatial maps (subscript “spatial”). Though less than satisfactory, we cannot think of
339 a better notation to express this point. However, we could add to the text the literal definition of
340 each term as supplied here in our Author’s Comment document.

341 **Changes made:** We could not find a better terminology here. To help, we now indicate in the text
342 that angle brackets indicate standard deviations, and we note that these terms are defined in
343 Table 1.

344 Table 1 is also confusing. Why are ACF bins different? Why are time windows for Eulerian and
345 Lagrangian different? Does that have any effect on the comparison? (I think it would if you were
346 calculating other statistics). Where does the 27.8km ACF bin for Eulerian length scale come
347 from? I probably missed it.

348 As indicated in your Major Comment above, we plan to better motivate these choices and
349 consolidate them into a subsection of Discussions that will cover all methodological choices.
350 Briefly, we address your specific questions here.

351 Regarding temporal ACF bin sizes: Ideally, one would use a bin size that matches the sampling
352 interval of the time series because this is the smallest lag that can be resolved. For this reason,
353 the ACFs based on satellite altimetry, satellite ocean color, or drifters use a bin size of 1 day. The
354 floats have a variable profiling interval. While they sometimes profile with a frequency of about
355 1 per day, they generally profile less frequently and we found a bin size of 5 days to be a
356 reasonable choice (with smaller bin sizes, many segments would offer no pairs). As we state, the
357 two metbio* float segments are given special attention because they profile more frequently, and
358 for that reason we were able to use a finer bin size of 1 day. As a general statement, choosing a
359 larger bin size for the ACF causes structure (curvature) of the ACF to be poorly resolved at short
360 lag and biases time scales large. This point is brought up in section 4.2.

361 Regarding temporal segment lengths for ACF analysis: The Lagrangian segments should be kept
362 as short as possible because as a platform moves it may encounter different environmental
363 (physical or otherwise) conditions, and we found 120 days was a reasonable length of time. For
364 Eulerian segments, this is not an issue, and, since seasonal variability is removed, length of the
365 segment is generally unimportant. Given that, we used 365-366 day segments for chlorophyll out
366 of convenience since the data were stored as yearly files.

367 Regarding spatial ACF bin size: Related to our point about temporal ACF bin sizes, it would be
368 best to use a bin size equal to the data spacing. We chose 27.8 km as that approximately
369 corresponds to the 0.25° resolution of the data in the latitudinal direction. Obviously, pairs
370 spaced zonally may have a separation less than that distance and would fall into the first bin.

371 **Changes made:** The above discussion from our Author Comment is included in our new Section
372 4.6 (Methodological decisions).

373 Figure 1d. orange profiles: QPI<5km; blue: all others (i.e., not total)

374 You are correct. The total height of the bars represents the total number of profiles, but the blue
375 region represents only the portion with QPI > 5 km. We will fix this.

376 Changes made: The caption of Figure 1 has been corrected accordingly.

377 Line 180. Please specify the convention for flagging. It is my understanding that BGCArgo
378 flagging may have changed through the years and between institutions. (I've used Sprofs where
379 3 means bad).

380 We will provide a brief description here and include reference to the relevant Argo user manual
381 for more information.

382 Changes made: The flag levels have been defined and a citation to Argo Data Management Team
383 (2019) added.

384 How does the GlobColour product compare to other products? Why is this one selected over
385 others? (OCCCI, for instance). I suggest including a brief sentence.

386 We did not compare how different satellite products affect the scales that are calculated. We
387 chose GlobColour because it is blended from all available satellites and is therefore probably a
388 most complete product in terms of space-time coverage without interpolation. Further, the study
389 of Zhang et al. (2019) demonstrated that GlobColour data projected onto surface drifter tracks
390 resolve realistic Lagrangian behavior in terms of (sub)mesoscale dynamics, so we conclude that
391 their space-time information is biophysically accurate. We plan to include this information in the
392 manuscript, in the Methods section where we introduce the data. As for why we chose to use a
393 25 km product, that requires a more nuanced discussion and we refer you to please see our
394 Author Comment to Referee #1. We propose to include that discussion in a revised Discussions
395 section.

396 Changes made: The motivations for using a 0.25° ocean colour product, the rationale for why
397 GlobColour specifically was chosen, and how the choice of this product may influence our
398 results are all discussed in the new Section 4.6 as it seemed to fit better there.

399 Section 3.3 could be simplified. Two sets of chlorophyll anomalies are estimated: 1. Anomalies
400 with respect to a 31-day smoothing filter, and 2. Anomalies with respect to the climatology. I
401 would suggest stating something like that to start, and then continue with the details.

402 This is a reasonable suggestion, and we can modify the opening sentence of section 3.3
403 accordingly.

404 Changes made: The opening paragraph of Section 3.3 has been updated following your
405 suggestion.

406 The climatology is based on the same 31-day filter + a boxcar function? This is not exactly what
407 comes to my mind when “climatology” or “repeating annual cycle” is mentioned.

408 We apologize for the confusion here. This is an admittedly technical point so we left the details
409 in the Appendix B, but perhaps we need to clarify the main text. Essentially, the “smoothed”
410 subtrahend is from a 3-D convolution with a filter kernel that is a 2-D Gaussian in space and a
411 31-day Hamming window in time. The “climatology” subtrahend comes from first stacking the
412 arrays by day of year in a 4th dimension so that the convolution is with a 4-D kernel that is a 2-D
413 Gaussian in space, a 31-day Hamming window in day-of-year (like a Julian day, not absolute

414 calendar date), and a boxcar (equal weights) across years. That way, as we say in the Appendix
415 B, “[for] example, January 1 of every year is regarded as having the same time coordinate”. The
416 end result is a set of maps for each day-of-year, hence making it a repeating annual cycle. We
417 can move the illustrative sentence (reproduced here) to the main text for clarity.

418 **Changes made:** The third paragraph of Section 3.3 has been updated to clarify how the
419 “climatology” subtrahend is constructed. We follow the outline given in our Author Comment
420 and use text from Appendix B.

421 I don’t like the use of the satellite-based “subtrahend” to estimate chlorophyll anomalies from the
422 MLD-averaged chlorophyll from the float. How does the MLD average compare to the satellite?
423 I think some type of bias correction may be needed. You mention that the subtrahend is
424 regressed against float data. Do you mean you corrected a bias? That should be included in a
425 supplement.

426 It is not possible to construct a “climatology” subtrahend from the floats because there is not
427 enough interannual coverage of floats over the spatial footprint of the horizontal component of
428 the filter at any given time step. For this reason, we need to turn to climatological fields
429 constructed from the satellite data. To illustrate how the subtrahends look (and how float and
430 satellite data compare), we included Figure B1. We think that figure illustrates that the satellite-
431 constructed “climatology” subtrahend is reasonable to compare with the float data. As you point
432 out, the regression effectively serves as a bias correction so that the mean and range of the
433 subtrahend (once it is projected onto the floats) is comparable to the mean and range of the float-
434 measured chlorophyll. The details of the procedure are described in the existing Appendix B. If it
435 is helpful, regression coefficients can be included.

436 **Changes made:** We continue to use the satellite data to construct the “climatology” subtrahend
437 for the floats. We updated Appendix B to refer to the regression as serving like a “bias
438 correction” and include the equation.

439 Line 240. Why aren’t Eulerian and Lagrangian segments equal?

440 This is a matter of convenience. The Lagrangian segments should be kept as short as possible
441 because as a platform moves it may encounter different environmental (physical or otherwise)
442 conditions, and we found 120 days was a reasonable length of time. However, this is not
443 important for Eulerian segments, especially since low frequency (such as seasonal) variability
444 has been removed. Since the Eulerian data are stored in annual files, it was easiest to work with
445 year-long segments.

446 **Changes made:** This answer is given in the new Section 4.6 (Methodological decisions).

447 Section 3.4 could be simplified as well. If I understand correctly, you tested two approaches to
448 estimate spatially averaged scales. In lines 272-275 you mention you use one or the other. When
449 and why you use each one should be clearer.

450 This is basically correct. When possible, we apply both methods (e.g., compare Figures 6 and
451 D1), but only equation 8 is an option for any scales derived from ocean color due to the large
452 number of gaps. In simplest terms: “All scales are derived by averaging in space (from
453 integrating Eq. (7) and averaging), except any scales involving satellite ocean color, where large
454 numbers of gaps require computing scales from space-composited ACFs (from integrating Eq.

455 (8)).” We can open the discussion on lines 272-281 with the preceding simple sentence and then
456 eliminate much of the redundant (and less clear) text that follows.

457 Changes made: Section 3.4 has been rewritten following our Author Comment above.

458 Line 292. Picks?

459 Sorry: “picks” should read “scales”.

460 Changes made: The typo has been corrected.

461 Lines 319-320: “If we take ...” this sentence is confusing.

462 We apologize for the confusion here. Our intention is to draw some contrast between float
463 profiles where the QPI is “small” and “large”. While the distribution in Figure 2 is continuous
464 and there is no real threshold, we noted that there is a mode of profiles between zero and 5 km,
465 so we chose this threshold for display purposes. As we mention in the text, 5 km is a good
466 compromise between having a large amount of profiles and having a QPI that is small, so it
467 serves as a reasonable threshold between a “small” and “large” QPI for the purposes of display in
468 Figures 2-3. Other than for display purposes in those figures, though, QPI is only used for
469 weighting averaged scales and there is no use of a threshold in Figures 4-9 or their
470 interpretations. We can update the text with the information supplied in this Author’s Comment
471 document to clarify where the 5 km threshold comes from and when and why it is used.

472 Changes made: The opening paragraph of Section 4.1 has been rewritten following the outline
473 given in our Author Comment above to make it clear that the threshold of 5 km is arbitrary and
474 for display purposes only.

475 I probably missed this. Are the results in figures 5 to 9 based on all profiles or only QPI<5km?

476 We apologize for the confusion here. Figures 5 to 9 display results based on all float profiles. We
477 can update the captions to convey this. The filled circles treat all float segments equally in the
478 averages whereas the crosses weight by segment-median QPI^{-2} so that segments with smaller
479 QPI count more.

480 Changes made: The caption in Figure 5 has been updated. In our reading, the edits to Section 4.1
481 now make it clear that the threshold of 5 km is for display purposes only, and that the threshold
482 has no bearing on Figures 5-9 or their discussion.

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