

Dear Editor,

below we provide a point-by-point response to all individual comments of the reviewers. We thank the reviewers for their constructive comments that helped improving the overall quality of our revised manuscript.

We indicate the reviewer's comments by black text, our answers are given in blue text, and new text that now appears in the revised manuscript is given in green text. In addition, we include at the end versions of the revised manuscript and the revised supplementary information with all changes indicated.

We hope that our revised manuscript now meets with your approval.

Sincerely yours,

Thomas Koop

Answer to reviewer 1

Eickhoff et al. present valuable results with a potential significant contribution to a better understanding of the microphysical processes in the ocean and the (polar) atmosphere.

However, in my opinion, some chapters could benefit from some shortening/rewriting, while the major findings should be rather put in a nutshell. Furthermore, I think, this manuscript is still lacking a look on the bigger picture and a critical discussion of its (atmospheric) significance and implications. See below for more specific comments.

We thank the reviewer for highlighting the importance of our contribution for the polar ocean and atmosphere and for the very detailed comments that have improved the revised version of the manuscript.

1. Introduction

L34-39 and L61-71.: Could you be more specific, which of the findings hold true for the Arctic, the Antarctic or both polar regions? This appears very important to me to not mix them up, since certain features of the Arctic and Antarctic are quite different.

The comments on the *Fragilariopsis cylindrus* hold for both the Arctic and the Antarctic and we have updated that sentence accordingly.

The diatom *Fragilariopsis cylindrus* (see Fig. 1) is widespread in polar environments and is one of the predominant species within the Arctic and Antarctic microbial assemblages (Kang and Fryxell, 1992; Poulin et al., 2011; van Leeuwe et al., 2018).

Regarding atmospheric INP, we have added the following new paragraph at the end of the introduction to discuss the differences in more detail:

There are some differences regarding the relevance of INPs in the Arctic and Antarctic polar regions. While in both polar latitudes the absolute concentrations of INPs are low, the influence of anthropogenic aerosols and INPs is much larger in the Arctic due to long-range transport during the Arctic winter (Šantl-Temkiv et al., 2019; Šantl-Temkiv et al., 2020; Ekman and Schmale, 2022). During the Arctic summer, aerosol lifetimes are shorter due to increased wet removal preventing long range transport and thus increasing the importance of locally emitted INPs. In the Antarctic, the influence of anthropogenic aerosols and INPs is generally much smaller (Stohl and Sodemann, 2010; Ekman and Schmale, 2022). During winter, blowing snow from the sea ice is the main aerosol source in the southern polar region, while DMS and other organic compounds from algae bloom are the main source during summer.

L34-35 Please add a reference.

The references were added as requested.

The diatom *Fragilariopsis cylindrus* (see Fig. 1) is widespread in polar environments and is one of the predominant species within the Arctic and Antarctic microbial assemblages (Kang and Fryxell, 1992; Poulin et al., 2011; van Leeuwe et al., 2018).

L40-42 Please recheck, if Wilson et al. (2015) or Aslam et al. (2018) are suitable references for supporting the statement of “the production of so-called ice-binding proteins (IBPs)”

The two references referred to only the second part of the sentence “and of other EPS also found in other diatom species”. We have now added a reference for the first part of the sentence “the production of so-called ice-binding proteins (IBPs)” and separated the two parts with a comma:

One prominent example is the production of so-called ice-binding proteins (IBPs) (Bayer-Giraldi et al., 2011), and of other EPS that are also found in other diatom species (Wilson et al., 2015; Aslam et al., 2018).

L52-54 Which reference states that EPS do have good ice-binding properties? Is there any knowledge under which conditions (chemical composition) EPS do have these properties?

We have added the appropriate references and added the info that these EPS are mainly polysaccharides and proteins:

The very good ice-binding properties of *fcIBP* and EPS (mainly polysaccharides and proteins) under sea ice brine conditions have been reported in previous studies (Krembs et al., 2002; Bayer-Giraldi et al., 2011; Krembs et al., 2011).

L46-60 It might be important to clarify the relationship between the terms “ice-binding protein”, “antifreeze protein” and “ice-nucleating protein” here. Do I get it right that less efficient ice-nucleating substances (in the sense that they are active at lower temperatures) can be at the same time antifreeze-substances at higher temperatures?

We have added some text to clarify the relationships of the different terms as requested:

L59-60 “Here we explore whether a similar ice-nucleating effect does occur also for IBPs from *F. cylindrus*” -> What is the outcome? This might be an important result that deserves to be mentioned in chapter “5. Summary and Conclusions”

The results of the ice nucleation experiments are now mentioned in the Summary:

For the ice-binding (antifreeze) protein *fcIBP11*, we did not observe any evidence for promoting ice nucleation at low temperatures.

2. Material and methods

L76 Please add the months, when ANT XVI/3 took place.

The information has been added:

The investigated *F. cylindrus* cells belong to the strain TM99 isolated in 1999 from the sea ice of the Weddel Sea, Antarctica, by Thomas Mock (*Polarstern* ANT XVI/3 expedition, which took place in the early spring from March to May 1999).

L76-81 When did the laboratory steps happen after the isolation in 1999? Back in 1999 or just recently, just before the ice nucleation experiments started? I just wonder how many cells (of those 108 cells) were still alive after a storage of approximately 20 years.

The cells were kept in culture since the expedition, the laboratory steps took place just recently, immediately before the experiments started. This information was added to the manuscript.

Since then, stock cultures were kept in *f/2* medium (Guillard and Ryther, 1962) set up with Antarctic water and cultivated at 0°C and under continuous illumination of approximately 25 $\mu\text{E m}^{-2} \text{s}^{-1}$. Before the experiment, cell numbers of the *F. cylindrus* cultures were monitored using a Coulter Counter, and cells were harvested during the exponential growth phase.

L84-94 This section could be shortened, since it contains much redundant information. It might be enough to refer to Tabel S1 in regard of the exact composition of the artificial sea salt.

Section has been shortened as suggested by removing the detailed information on the salts and referring to table S1.

The composition of the salts and their concentrations are given in Supplemental Information Table S1.

L 94, 100, 104, 110, 118, ... If you only used one type of filter throughout this study, it might be enough to mention the filter type just once in the beginning.

We now mention the filter type only once with a more generalized statement:

The artificial seawater was filtered through a syringe filter (0.22 μm , Polyethersulfone, SimplePure) in order to exclude any effect of suspended dust particles on ice nucleation. This filter has been used for all filtrations in this study unless otherwise mentioned.

L103 I guess replacing "*F. cylindrus* cells" with "*F. cylindrus* samples" might make it more accurate.

Changed as suggested.

L106 Any reference that states that fcIBP11 is a soluble macromolecule detached from the cell? Or could it be connected to the cell surface of the diatoms as well?

According to Bayer-Giraldi et al., 2011, all analysed fcIBP11 isoforms are neither intracellular nor transmembrane, but secreted into the extracellular space. We have added the reference to the sentence.

L109-110 Is it possible to give an approximate estimate of the extend of cell loss?

We have added an approximate estimate as requested:

From the comparison of the frozen fraction curves obtained with the sample with those of unfiltered samples (see below) our best estimate of the concentration is about 2×10^6 cells per mL (estimated uncertainty range $1 \times 10^6 - 1 \times 10^7$ cells per mL).

L136 “...belongs to the DUF3494 IBP family,....” .. was already mentioned in the introduction (L 43). Why is it relevant to mention it here again?

We have removed the sentence in question as requested.

L145-184 Is it possible to shorten these sections onto the relevant information since all these methods have been published before? Of course, the main principle and deviations from the original protocol should be mentioned. If you want to keep all details, maybe it is possible to shift part of these descriptions to the SI? Instead I would appreciate a short explanation, why you chose two different setups (2.3.1, 2.3.2) for the ice nucleation experiments.

We have shortened these paragraphs by shifting parts of the text to the SI. We have also added a short explanation, why we have used two different methods for our ice nucleation experiments:

The DSC experiment has been used as a simple and direct method to check whether *F. cylindrus* diatoms are potential ice nucleators or not. The method does not allow for the observation of single droplets, and we can only study cell fragments but not intact cells because the latter are disrupted during the emulsion preparation process. Therefore, we have used the WISDOM microfluidic device, which is described below, as the main experimental method in this study.

L185. You should check the appropriate use of “INP” throughout the manuscript. Maybe “ice nucleating molecules” or “ice nucleating entities” could be more correct here? I recommend to check the terminology proposed by (Vali et al., 2015).

We have used “INP”, because we think “particle” is the most accurate term for the diatom cells and their fragments. This usage is also in line with many papers in the literature (including those of authors of the Vali et al. (2015) paper), in which diatoms were termed INPs, e.g.: (Creamean et al., 2021; McCluskey et al., 2018).

L186-188 Is this true? I think the methodology in the literature is just different to yours, where bigger droplets with higher volume per droplet were used. E.g. Budke and Koop (2015) used in the BINARY 1 μ L droplets. Why did you choose for these small volumes of 380 pL (l 190) then?

We think there is a misunderstanding here. What we mean is that \sim microliter droplet experiments usually employ a larger number of INPs per droplet (not INP concentration per volume). Moreover, these large droplet experiments are often subject to heterogeneous ice nucleation by impurities and the supporting surfaces at low temperatures, thus making it nearly impossible to study INPs with very low activity that nucleate only at temperatures just above homogeneous nucleation temperature. (The median freezing temperature of ‘pure’ water in the BINARY device is about -30°C , while the freezing temperature of the diatom experiments presented in Fig. 2 is between -32°C and -37°C .) For this purpose, small volume experiments such as the \sim nanoliter microfluidic device used here are much more suitable, as they allow to reach homogeneous ice nucleation because the smaller droplets usually contain less impurities per droplet (The median freezing temperature of artificial seawater in the WISDOM device is about -40°C , see Fig.2). We have rewritten the text to make this clearer.

Ice nucleation studies using larger-volume droplet arrays usually employ relatively high concentrations of INPs per droplet, e.g. mineral dust particles or bacterial cells (Budke and Koop, 2015; Hiranuma et al., 2015; Wex et al., 2015; DeMott et al., 2018; Hiranuma et al., 2019; Kunert et al., 2019; Ickes et al., 2020), to ensure that freezing is induced at a temperature that is higher than that triggered by the

supporting surface or minute amounts of impurities contained in the water. In the present study, the total amount of INPs was small due to the limited availability of *F. cylindrus* cells, suggesting the use of small droplet methods which require less total INP material. We investigated droplets with a diameter of 90 μm , corresponding to a volume of about 380 μL . Another, probably more important advantage of using these small droplet volumes is that we can measure ice nucleation down to the homogenous freezing temperature of water (Riechers et al., 2013; Reicher et al., 2018; Tarn et al., 2021), enabling also the investigation of rather poor ice nucleators.

L189-192: I find this statement quite surprising, since you started your experiments with a solution of a highly concentrated culture (see 2.1). Then you performed several dilutions. And now you are stating that the number of INP was not enough for your experimental setup. Why did you perform dilutions then?

This question probably arises from the same misunderstanding discussed in the previous comment. We did these dilution experiments to obtain more accurate cumulative numbers of ice active sites per mass n_{m_total} and over a broader temperature range such that we can compare the data better to other literature data (see Figure 7).

L199-308 Is it possible to shift this section (or parts of it) into the SI?

As suggested, we have shifted the entire section (former lines L199-308) into the SI.

3. Results and Discussion

General:

Is it possible to bring this part more on a point? It feels like reading and rereading the same or similar aspects.

We have shortened this part by shifting the detailed description of DSC experiments including the former Fig.4 into the SI. We have also rewritten some sentences or parts and now term this section "Results" (see your comment on section names further below).

L 326 How can freezing at -44°C be still relevant, when water droplets usually freeze at -38°C homogeneously? Is it because of the small volume of the droplets in your setup?

The droplets in the DSC experiments (and also in the microfluidic experiments) are all large enough not to be affected by the Gibbs-Thomson effect. The homogeneous freezing temperature of water droplets in the DSC experiments is indeed -38°C , as defined by the onset of the exothermic signal. The corresponding onset temperatures for the two samples, sea-water and sea-water containing *F. cylindrus*, are shifted to lower temperatures (-39°C to -40°C) due to the colligative freezing point depression effect of the salts contained in the seawater. We added this information to this text (now in the SI):

Because of the colligative freezing point depression of the seawater, the freezing temperatures of the reference and the sample are shifted to lower temperatures, compared to pure water.

Figure 5: Just as another example, where sentences can be shortened. Instead of: “b): Freezing temperatures of the same samples shown in panel (a), but after filtration with a pore size of 0.22 μm .” you could write: “b): Freezing temperatures of the filtered (0.22 μm) samples.”

Changed as suggested.

Line 369 “all diatoms” (?) or all cells?

We agree and changed the term to ‘all *F. cylindrus* cells’ to make this point more clear. We also made corresponding changes at several other places throughout the text where appropriate.

L378 Within the whole manuscript, proteins are proposed as likely ice nucleating molecules. What about polysaccharides? Any tests performed into this direction?

Yes, it is known from the literature that polysaccharides can also act as ice-nucleating molecules (e.g. Dreischmeier et al. (2017)), so we added the term polysaccharides to the sentence in question:

We suggest that their freezing is induced by cell fragments or by INPs released by the *F. cylindrus* diatoms, e.g. soluble species from the EPS such as proteins or polysaccharides.

However, as the experiments on soluble components described further below in the paper did not reveal any ice nucleating effect, we did not analyse our samples for polysaccharide content.

L379 “macromolecules” or maybe “polysaccharides”, to be more specific. Why is the comparison of diatoms with birch pollen relevant here?

We used the comparison to birch pollen here just to mention that by filtering the *F. cylindrus* samples, it is entirely possible that ice nucleation activity can remain, when either smaller fragments or soluble ice-nucleating molecules passes through the filter, thus resulting in n_N values larger than 1. This is exactly what happens when pollen suspended in water are filtered and the remaining washing water contains ice nucleating molecules. Also in the case of pollen, the n_N values reach values far above 1 (about 10^4 per pollen grain to be precise).

4. Discussion and Implications

General:

What is the difference between the chapters “3. Results and Discussion” and “4. Discussion and Implications”? It appears to me that the text (or at least parts) of chapter 4 in the current version of the manuscript might still represent a subchapter of “3. Results and Discussion”.

As suggested we have renamed chapters 3-5 and also shifted some of their contents. Chapter 3 is now simply termed “Results”. Chapter 4 remains “Discussion and Implications” because we compare our results to those of other diatoms and combine these datasets into a single parameterization. Moreover, we added two new paragraphs to discuss implications of our results for the atmosphere and compare them to (the very sparse) Southern Ocean diatom and INP measurements (see your next comments).

This study was mainly motivated with the atmospheric relevance of INPs (e.g. lines 64-72, lines 534-536). However, a critical discussion of these (new) findings for atmospheric implications are still missing and could fit here. Some of the following aspects should be discussed in this section:

Which atmospheric residence time would you expect for diatoms, their fragments and exudates? Whole marine diatoms are rather big and might precipitate within few seconds or minutes. Can it be expected that complete cells/fragments will make it into the atmospheric layers relevant for mixed-phase clouds or even cirrus clouds? (as implied in lines 66-70)

We have added some of this information to the text.

Which are the atmospheric concentrations of diatoms/*Fragilariopsis cylindrus* or fragments in the ambient air?

As far as we now there are no such measurements.

In Figure 10, you nicely compare the ice nucleating activity of *Fragilariopsis cylindrus* with several diatoms from other studies. However, a rating of the importance of diatoms as INP in the polar regions/Southern Ocean in comparison to other types of INP (such as marine bacteria, mineral dust, ...) in regard of abundance and/or ice nucleating efficiency is missing.

The Antarctic is known for a low number of efficient atmospheric INPs in comparison to the Arctic (e.g. (McCluskey et al., 2018; Wex et al., 2019; Hartmann et al., 2021; Zeppenfeld et al., 2021)). Considering this fact, how would you evaluate the importance of your findings for a better understanding of the Antarctic environment/atmosphere?

We have added two new paragraphs relating our data on *F. cylindrus* and sea ice diatoms to marine INP data and also try to compare the potential importance of our experimental data to Southern Ocean INP abundances. As there are only very sparse data on diatoms and INP in the Southern Ocean and the Antarctic marine environment these can only be regarded as order of magnitude estimates. Clearly, more field experiments in these regions are highly desirable. The new text reads:

In the following, we put the ice nucleation data of *F. cylindrus* and the other sea ice diatoms into context by comparing to field studies. Wilson et al. (2015) provided experimental evidence for a marine biogenic source of ice nucleating particles and suggested that exudates and fragments of diatoms as a source of the ice nucleating material located in the sea surface microlayer. Their low-temperature freezing data reveals a cumulative number of ice nucleating active sites per total organic carbon mass n_{m_TOC} of $\sim 1.3 \times 10^{10} \text{ g}^{-1}$ at -27°C (calculated from the equation given in the caption of their Fig. 2), which is the low-temperature end of their data, and the most relevant to the present study. To compare this value to the n_{m_total} values given in Fig. 7, we estimated that the organic carbon content of their samples varies between 39.32% (representing the organic carbon content of *F. cylindrus* cells, see above) or 100% (representing a purely organic carbon composition), resulting in a range of n_{m_total} of $\sim 5.0 \times 10^9 - 1.3 \times 10^{10} \text{ g}^{-1}$ for their Arctic sea surface microlayer samples. These are compared to n_{m_total} values of $8.2 \times 10^7 \text{ g}^{-1}$ (2σ prediction bands: $2.8 \times 10^7 - 2.4 \times 10^8 \text{ g}^{-1}$) for *F. cylindrus* and of $5.8 \times 10^8 \text{ g}^{-1}$ (2σ prediction bands: $7.0 \times 10^7 - 4.8 \times 10^9 \text{ g}^{-1}$) for sea ice diatoms, respectively, at -27°C , indicating that *F. cylindrus* and other sea ice diatoms may contribute to the marine INP in the Southern Oceans and Antarctic seawater, assuming the Wilson et al. parameterization applies also to these areas.

In another comparison, we use measurements of insoluble aerosol particles made at Amsterdam Island in the Southern Indian Ocean (Gaudichet et al., 1989). These measurements show that marine biogenic particles make up between 8 and 28% of the number of detected particles and that these were predominantly assigned to *Radiolaria* and diatom fragments (identified as amorphous silicates), with about 27 % or $2.7 \times 10^4 \text{ m}^{-3}$ particles observed in the southern winter (July) and fewer in fall (May, 8 %, $2.4 \times 10^4 \text{ m}^{-3}$) and spring (September, 7 %, $1.8 \times 10^3 \text{ m}^{-3}$). If we assume that all *Radiolaria* and diatom fragments can be attributed to *F. cylindrus* diatoms, we can calculate the mass concentration of *F. cylindrus* diatom cells per cubic meter of air from the mass per individual cell ($m_{\text{total}} = 4.5 \times 10^{-11} \text{ g}$, see above), yielding values of $1.2 \times 10^{-6} \text{ g m}^{-3}$ air (July), $1.1 \times 10^{-6} \text{ g m}^{-3}$ air (May), and $8.1 \times 10^{-8} \text{ g m}^{-3}$ air (September). Using the parametrization of the cumulative number of ice nucleating active sites per mass *F. cylindrus* in Eq. (S10), we calculate a $n_{m_{\text{total}}}$ value of $8.2 \times 10^7 \text{ g}^{-1}$ (2σ prediction bands: $2.8 \times 10^7 - 2.4 \times 10^8 \text{ g}^{-1}$) at -27°C , see above, from which we can derive the $\sim 88 \text{ INP m}^{-3}$ air (2σ : 3-250) at -27°C in fall (May). This value can be compared to *in situ* total INP measurements in the Southern Ocean south of Australia in fall (March-April) yielding values between 34 and 207 INP m^{-3} air at -27°C (McCluskey et al., 2018). Although the above calculations are order of magnitude estimates, the comparison shows that it is not unreasonable that sea ice diatoms such as *F. cylindrus* and their fragments may constitute a significant fraction of the INP in the Southern Ocean and Antarctic waters.

Specific:

L494 You performed own elementary analysis for obtaining the carbon content in your samples? Could you please add the method to chapter “2. Materials and Methods”? Few lines might be sufficient.

As requested, we have added a few lines on the elemental analysis in the new section 2.4:

The total carbon content of the *F. cylindrus* samples has been determined using elemental analysis. For this purpose, an amount of 0.7 mg *F. cylindrus* diatoms was combusted at a high temperature ($T > 1000^\circ \text{C}$) in a Tin-crucible and the composition was analysed using a commercially available elemental analyser (EuroVector, Euro EA).

Figure 10: Is it anyhow possible to still extend this figure with the experimental freezing results on the diatom *Thalassiosira pseudonana* by (Wilson et al., 2015) or (Knopf et al., 2011)?

Knopf et al. (2011) were the first to show that marine diatoms can act as INPs. However, as they presented only freezing temperature data, but no n_m or n_N values, we could not include these data in our comparison in Figure 7 (former Fig.10)

The same is true for the *Thalassiosira pseudonana* data in Wilson et al., 2015. Nevertheless, we now make a numerical comparison in the text to the sea surface microlayer freezing data of Wilson et al., 2015 at -27°C , which is the low temperature end for the cumulative number of INPs of their study (their Fig.2b) and the high temperature end of our *F. cylindrus* data. Note, that the direct comparison is difficult, as they provide their data as cumulative number of INPs per gram of **total organic carbon** (TOC). As the source of the ice nucleator and the species distribution in the sea surface microlayer is not known to us, we had to estimate the mass fraction of the organic carbon. The related text now reads:

Wilson et al. (2015) provided experimental evidence for a marine biogenic source of ice nucleating particles and suggested that exudates and fragments of diatoms as a source of the ice nucleating material located in the sea surface microlayer. Their low-temperature freezing data reveals a cumulative number of ice nucleating active sites per total organic carbon mass n_{m_TOC} of $\sim 1.3 \times 10^{10} \text{ g}^{-1}$ at -27°C (calculated from the equation given in the caption of their Fig. 2), which is the low-temperature end of their data, and the most relevant to the present study. To compare this value to the n_{m_total} values given in Fig. 7, we estimated that the organic carbon content of their samples varies between 39.32% (representing the organic carbon content of *F. cylindrus* cells, see above) or 100% (representing a purely organic carbon composition), resulting in a range of n_{m_total} of $\sim 5.0 \times 10^9$ – $1.3 \times 10^{10} \text{ g}^{-1}$ for their Arctic sea surface microlayer samples. These are compared to n_{m_total} values of $8.2 \times 10^7 \text{ g}^{-1}$ (2σ prediction bands: 2.8×10^7 – $2.4 \times 10^8 \text{ g}^{-1}$) for *F. cylindrus* and of $5.8 \times 10^8 \text{ g}^{-1}$ (2σ prediction bands: 7.0×10^7 – $4.8 \times 10^9 \text{ g}^{-1}$) for sea ice diatoms, respectively, at -27°C , indicating that *F. cylindrus* and other sea ice diatoms may contribute to the marine INP in the Southern Oceans and Antarctic seawater, assuming the Wilson et al. parameterization applies also to these areas.

Figure 10: Now you show values which are normalized on mass (total mass?) At which part did you include the carbon content (L 494-495) then?

Originally, we only know the number of diatom cells per mL of water, which is why we calculated the cumulative number n_N of ice nucleating sites per number of *F. cylindrus* diatoms. In the literature, many studies (including field studies) provide the cumulative number n_N of ice nucleating sites per mass. Hence, to compare our freezing data to other studies in the literature the mass concentration of the *F. cylindrus* diatoms was required and, thus, the mass per individual *F. cylindrus* diatom cell m_{total} . However, in the literature we only found a value for the mass of carbon per *F. cylindrus* diatom. Hence, we needed the mass fraction of carbon in the *F. cylindrus* diatoms (which we could determine experimentally, see new section 2.4). To make this procedure more clear, we added the calculation of m_{total} in the manuscript:

For the *F. cylindrus* samples investigated here, we used the total carbon mass per *F. cylindrus* cell from the literature (Kang and Fryxell, 1992) and performed elemental analysis to obtain the carbon content of our samples, resulting in a value of 39.32 % to calculate the average total mass per individual *F. cylindrus* diatom cell of $m_{total} = 4.5 \times 10^{-11} \text{ g}$. Using these values and our experimental data in Eq. (1), we have calculated the ice nucleating active sites n_{m_total} of the *F. cylindrus* diatoms, see the blue circles in Fig. 7.

L505-507: “All temperatures were corrected for the freezing point depressions of different buffers...” How did you do it? Did you follow the approach by Koop and Zobrist (2009)?

For the data from Xi et al. (2021), we only received freezing temperature data that were already corrected for the freezing point depression by these authors. Hence, we have used their data for our comparison plot.

For the data from Ickes et al., we got their raw data with both the uncorrected freezing temperatures and those corrected for the freezing point depression. Hence, we have used their corrected ones for our comparison plot.

For our own data, we have obtained the freezing point depression by directly measuring the shift in ice nucleation temperature between pure distilled water and the artificial seawater that we used. We have corrected all measured nucleation temperatures of *F. cylindrus* diatoms in seawater by this temperature shift. We have added the following text:

To allow a direct comparison of ice nucleation of the different diatoms, which were studied in different types of aqueous solutions, all the ice nucleation temperatures shown in Fig. 7 have been corrected

(either by the original authors or by us) for the colligative solute effect and represent diatom ice nucleation in pure water. We have corrected the freezing temperatures of the *F. cylindrus* samples by the measured difference between the T_{50} of pure double-distilled water and pure artificial seawater without any diatoms.

L508-510 Is it necessary to mention this in the main text? It could be sufficient to add this information as a footnote in Figure 10.

Changed as suggested.

L528 Is there any reason, why you convert °C to K at this late part of the manuscript? You have not done it before, so why here?

We removed the additional information on the Kelvin temperature range to avoid confusion.

5. Conclusions

General:

The current version of the text rather represents a summary of the manuscript. However, conclusions are still sparse in this section. I'd recommend adding some real conclusions and a renaming of this chapter "Summary" or "Summary and Conclusions".

We have renamed this Chapter to "Summary and Conclusions" as requested and also made some text changes to that part.

Minor comments:

L67 "can be transported"

We inserted missing word "be".

L538: Remove "s" from "seas-ice"

The extra letter has been removed.

L538 Check for a consistent writing of "sea-ice diatoms" versus "sea ice diatoms" throughout the manuscript

Checked and unified to „sea ice diatoms“.

References:

- Aslam, S. N., Strauss, J., Thomas, D. N., Mock, T., and Underwood, G. J. C.: Identifying metabolic pathways for production of extracellular polymeric substances by the diatom *Fragilariopsis cylindrus* inhabiting sea ice, *The ISME journal*, 12, 1237–1251, doi:10.1038/s41396-017-0039-z, 2018.
- Bayer-Giraldi, M., Weikusat, I., Besir, H., and Dieckmann, G.: Characterization of an antifreeze protein from the polar diatom *Fragilariopsis cylindrus* and its relevance in sea ice, *Cryobiology*, 63, 210–219, doi:10.1016/j.cryobiol.2011.08.006, 2011.
- Budke, C. and Koop, T.: BINARY: an optical freezing array for assessing temperature and time dependence of heterogeneous ice nucleation, *Atmos. Meas. Tech.*, 8, 689–703, doi:10.5194/amt-8-689-2015, 2015.
- Creamean, J. M., Cenicerros, J. E., Newman, L., Pace, A. D., Hill, T. C. J., DeMott, P. J., and Rhodes, M. E.: Evaluating the potential for Haloarchaea to serve as ice nucleating particles, *Biogeosciences*, 18, 3751–3762, doi:10.5194/bg-18-3751-2021, 2021.
- DeMott, P. J., Möhler, O., Cziczo, D. J., Hiranuma, N., Petters, M. D., Petters, S. S., Belosi, F., Bingemer, H. G., Brooks, S. D., Budke, C., Burkert-Kohn, M., Collier, K. N., Danielczok, A., Eppers, O., Felgitsch, L., Garimella, S., Grothe, H., Herenz, P., Hill, T. C. J., Höhler, K., Kanji, Z. A., Kiselev, A., Koop, T., Kristensen, T. B., Krüger, K., Kulkarni, G., Levin, E. J. T., Murray, B. J., Nicosia, A., O'Sullivan, D., Peckhaus, A., Polen, M. J., Price, H. C., Reicher, N., Rothenberg, D. A., Rudich, Y., Santachiara, G., Schiebel, T., Schrod, J., Seifried, T. M., Stratmann, F., Sullivan, R. C., Suski, K. J., Szakáll, M., Taylor, H. P., Ullrich, R., Vergara-Temprado, J., Wagner, R., Whale, T. F., Weber, D., Welti, A., Wilson, T. W., Wolf, M. J., and Zenker, J.: The Fifth International Workshop on Ice Nucleation phase 2 (FIN-02): laboratory intercomparison of ice nucleation measurements, *Atmos. Meas. Tech.*, 11, 6231–6257, doi:10.5194/amt-11-6231-2018, 2018.
- Dreischmeier, K., Budke, C., Wiehemeier, L., Kottke, T., and Koop, T.: Boreal pollen contain ice-nucleating as well as ice-binding 'antifreeze' polysaccharides, *Scientific reports*, 7, 41890, doi:10.1038/srep41890, 2017.
- Ekman, A. M. and Schmale, J.: Aerosol processes in high-latitude environments and the effects on climate, in: *Aerosols and climate*, Carslaw, K. S. (Ed.), Elsevier, Amsterdam, Kidlington, Cambridge, MA, 651–706, 2022.
- Gaudichet, A., Lefèvre, R., Gaudry, A., Ardouin, B., Lambert, G., and Miller, J.: Mineralogical composition of aerosols at Amsterdam Island, *Tellus B*, 41B, 344–352, doi:10.1111/j.1600-0889.1989.tb00313.x, 1989.
- Guillard, R. R. L. and Ryther, J. H.: Studies of marine planktonic diatoms: I. *Cyclotella nana* hustedt, and *Detonula convolvacea* (cleve) gran, *Can. J. Microbiol.*, 8, 229–239, doi:10.1139/m62-029, 1962.
- Hiranuma, N., Adachi, K., Bell, D. M., Belosi, F., Beydoun, H., Bhaduri, B., Bingemer, H., Budke, C., Clemen, H.-C., Conen, F., Cory, K. M., Curtius, J., DeMott, P. J., Eppers, O., Grawe, S., Hartmann, S., Hoffmann, N., Höhler, K., Jantsch, E., Kiselev, A., Koop, T., Kulkarni, G., Mayer, A., Murakami, M., Murray, B. J., Nicosia, A., Petters, M. D., Piazza, M., Polen, M., Reicher, N., Rudich, Y., Saito, A., Santachiara, G., Schiebel, T., Schill, G. P., Schneider, J., Segev, L., Stopelli, E., Sullivan, R. C., Suski, K., Szakáll, M., Tajiri, T., Taylor, H., Tobo, Y., Ullrich, R., Weber, D., Wex, H., Whale, T. F., Whiteside, C. L., Yamashita, K., Zelenyuk, A., and Möhler, O.: A comprehensive characterization of ice nucleation by three different types of cellulose particles immersed in water, *Atmos. Chem. Phys.*, 19, 4823–4849, doi:10.5194/acp-19-4823-2019, 2019.
- Hiranuma, N., Augustin-Bauditz, S., Bingemer, H., Budke, C., Curtius, J., Danielczok, A., Diehl, K., Dreischmeier, K., Ebert, M., Frank, F., Hoffmann, N., Kandler, K., Kiselev, A., Koop, T., Leisner, T., Möhler, O., Nillius, B., Peckhaus, A., Rose, D., Weinbruch, S., Wex, H., Boose, Y., DeMott, P. J., Hader, J. D., Hill, T. C. J., Kanji, Z. A., Kulkarni, G., Levin, E. J. T., McCluskey, C. S., Murakami, M.,

- Murray, B. J., Niedermeier, D., Petters, M. D., O'Sullivan, D., Saito, A., Schill, G. P., Tajiri, T., Tolbert, M. A., Welti, A., Whale, T. F., Wright, T. P., and Yamashita, K.: A comprehensive laboratory study on the immersion freezing behavior of illite NX particles: a comparison of 17 ice nucleation measurement techniques, *Atmos. Chem. Phys.*, 15, 2489–2518, doi:10.5194/acp-15-2489-2015, 2015.
- Ickes, L., Porter, G. C. E., Wagner, R., Adams, M. P., Bierbauer, S., Bertram, A. K., Bilde, M., Christiansen, S., Ekman, A. M. L., Gorokhova, E., Höhler, K., Kiselev, A. A., Leck, C., Möhler, O., Murray, B. J., Schiebel, T., Ullrich, R., and Salter, M. E.: The ice-nucleating activity of Arctic sea surface micro-layer samples and marine algal cultures, *Atmos. Chem. Phys.*, 20, 11089–11117, doi:10.5194/acp-20-11089-2020, 2020.
- Kang, S.-H. and Fryxell, G.: *Fragilariopsis cylindrus* (Grunow) Krieger: The most abundant diatom in water column assemblages of Antarctic marginal ice-edge zones, *Polar Biol*, 12, doi:10.1007/BF00236984, 1992.
- Knopf, D. A., Alpert, P. A., Wang, B., and Aller, J. Y.: Stimulation of ice nucleation by marine diatoms, *Nature Geosci*, 4, 88–90, doi:10.1038/ngeo1037, 2011.
- Krembs, C., Eicken, H., and Deming, J. W.: Exopolymer alteration of physical properties of sea ice and implications for ice habitability and biogeochemistry in a warmer Arctic, *Proceedings of the National Academy of Sciences of the United States of America*, 108, 3653–3658, doi:10.1073/pnas.1100701108, 2011.
- Krembs, C., Eicken, H., Junge, K., and Deming, J.: High concentrations of exopolymeric substances in Arctic winter sea ice: implications for the polar ocean carbon cycle and cryoprotection of diatoms, *Deep Sea Research Part I: Oceanographic Research Papers*, 49, 2163–2181, doi:10.1016/S0967-0637(02)00122-X, 2002.
- Kunert, A. T., Pöhlker, M. L., Tang, K., Krevert, C. S., Wieder, C., Speth, K. R., Hanson, L. E., Morris, C. E., Schmale III, D. G., Pöschl, U., and Fröhlich-Nowoisky, J.: Macromolecular fungal ice nuclei in *Fusarium*: effects of physical and chemical processing, *Biogeosciences*, 16, 4647–4659, doi:10.5194/bg-16-4647-2019, 2019.
- McCluskey, C. S., Hill, T. C. J., Humphries, R. S., Rauker, A. M., Moreau, S., Strutton, P. G., Chambers, S. D., Williams, A. G., McRobert, I., Ward, J., Keywood, M. D., Harnwell, J., Ponsonby, W., Loh, Z. M., Krummel, P. B., Protat, A., Kreidenweis, S. M., and DeMott, P. J.: Observations of Ice Nucleating Particles Over Southern Ocean Waters, *Geophys. Res. Lett.*, 45, 11,989–11,997, doi:10.1029/2018GL079981, 2018.
- Poulin, M., Daugbjerg, N., Gradinger, R., Ilyash, L., Ratkova, T., and Quillfeldt, C. von: The pan-Arctic biodiversity of marine pelagic and sea-ice unicellular eukaryotes: a first-attempt assessment, *Mar Biodiv*, 41, 13–28, doi:10.1007/s12526-010-0058-8, 2011.
- Reicher, N., Segev, L., and Rudich, Y.: The Welzmann Supercooled Droplets Observation on a Microarray (WISDOM) and application for ambient dust, *Atmos. Meas. Tech.*, 11, 233–248, doi:10.5194/amt-11-233-2018, 2018.
- Riechers, B., Wittbracht, F., Hütten, A., and Koop, T.: The homogeneous ice nucleation rate of water droplets produced in a microfluidic device and the role of temperature uncertainty, *Physical chemistry chemical physics PCCP*, 15, 5873–5887, doi:10.1039/c3cp42437e, 2013.
- Šantl-Temkiv, T., Lange, R., Beddows, D., Rauter, U., Pilgaard, S., Dall'Osto, M., Gunde-Cimerman, N., Massling, A., and Wex, H.: Biogenic Sources of Ice Nucleating Particles at the High Arctic Site Vil-lum Research Station, *Environmental science & technology*, 53, 10580–10590, doi:10.1021/acs.est.9b00991, 2019.
- Šantl-Temkiv, T., Sikoparija, B., Maki, T., Carotenuto, F., Amato, P., Yao, M., Morris, C. E., Schnell, R., Jaenicke, R., Pöhlker, C., DeMott, P. J., Hill, T. C. J., and Huffman, J. A.: Bioaerosol field measurements: Challenges and perspectives in outdoor studies, *Aerosol Science and Technology*, 54, 520–546, doi:10.1080/02786826.2019.1676395, 2020.

- Stohl, A. and Sodemann, H.: Characteristics of atmospheric transport into the Antarctic troposphere, *J. Geophys. Res.*, 115, doi:10.1029/2009JD012536, 2010.
- Tarn, M. D., Sikora, S. N. F., Porter, G. C. E., Shim, J.-U., and Murray, B. J.: Homogeneous Freezing of Water Using Microfluidics, *Micromachines*, 12, doi:10.3390/mi12020223, 2021.
- Vali, G., DeMott, P. J., Möhler, O., and Whale, T. F.: Technical Note: A proposal for ice nucleation terminology, *Atmos. Chem. Phys.*, 15, 10263–10270, doi:10.5194/acp-15-10263-2015, 2015.
- van Leeuwe, M. A., Tedesco, L., Arrigo, K. R., Assmy, P., Campbell, K., Meiners, K. M., Rintala, J.-M., Selz, V., Thomas, D. N., and Stefels, J.: Microalgal community structure and primary production in Arctic and Antarctic sea ice: A synthesis, *Elementa: Science of the Anthropocene*, 6, doi:10.1525/elementa.267, 2018.
- Wex, H., Augustin-Bauditz, S., Boose, Y., Budke, C., Curtius, J., Diehl, K., Dreyer, A., Frank, F., Hartmann, S., Hiranuma, N., Jantsch, E., Kanji, Z. A., Kiselev, A., Koop, T., Möhler, O., Niedermeier, D., Nillius, B., Rösch, M., Rose, D., Schmidt, C., Steinke, I., and Stratmann, F.: Intercomparing different devices for the investigation of ice nucleating particles using Snomax[®] as test substance, *Atmos. Chem. Phys.*, 15, 1463–1485, doi:10.5194/acp-15-1463-2015, 2015.
- Wilson, T. W., Ladino, L. A., Alpert, P. A., Breckels, M. N., Brooks, I. M., Browse, J., Burrows, S. M., Carslaw, K. S., Huffman, J. A., Judd, C., Kilhau, W. P., Mason, R. H., McFiggans, G., Miller, L. A., Nájera, J. J., Polishchuk, E., Rae, S., Schiller, C. L., Si, M., Temprado, J. V., Whale, T. F., Wong, J. P. S., Wurl, O., Yakobi-Hancock, J. D., Abbatt, J. P. D., Aller, J. Y., Bertram, A. K., Knopf, D. A., and Murray, B. J.: A marine biogenic source of atmospheric ice-nucleating particles, *Nature*, 525, 234–238, doi:10.1038/nature14986, 2015.
- Xi, Y., Mercier, A., Kuang, C., Yun, J., Christy, A., Melo, L., Maldonado, M. T., Raymond, J. A., and Bertram, A. K.: Concentrations and properties of ice nucleating substances in exudates from Antarctic sea-ice diatoms, *Environmental science. Processes & impacts*, 23, 323–334, doi:10.1039/d0em00398k, 2021.

Answer to reviewer 2

This is a sound set of experiments showing an increase of 7.2 °C in the ice nucleation temperatures for seawater containing *F. cylindrus* diatoms when compared to pure seawater. The laboratory study seem carried out well and the literature review is state of the art. There are two important aspect to be considered before this paper can be accepted.

We thank the reviewer for this positive evaluation of our study and for the detailed comments that has improved the revised version of the manuscript.

The paper does not read well and it seems a bit too long. I suggest to merge the results discussion implication or to short by half all the text in the last two section. Once have a feeling there are many sentences saying the same and not really giving a clear simple message. Clean up and make a simple clear concise message.

We have significantly shortened the manuscript by deleting several parts of the Materials and Methods (subsections 2.3.1, 2.3.2, and nearly entirely 2.3.3 including Figs. 2 and 3) and also the entire subsection 3.1.1 of the Experimental Results (including Fig. 4) and moving them into the Supplementary Information. Furthermore, we have edited the sections 3. Results, 4. Discussions and Implications, and 5. Summary and Conclusions. Following suggestions by reviewer 1, we renamed these sections and also added some text on the atmospheric relevance of our experimental results on the ice nucleation of *F. cylindrus* diatom cells and fragments, as well as sea ice diatoms in general to section 4.

It seems to be that in the abstract and conclusion, and also in the introduction (well written) one of the main result is the results "that *F. cylindrus* diatom cells as well as cell fragments suspended in seawater can induce heterogeneous ice nucleation, while icebinding proteins produced by *F. cylindrus* such as *fclBP11* have negligible ice nucleation activity.". This is important and also compared with the literature, but what is the reason? Any literature support any speculation and possible reasons? This is in stark contrast with other literature supporting the idea of proteins being important in INP, but little is discuss in the text of this paper. I suggest to expand this extensively given it seems a major result. It is also important to give possible biogeochemical reasons of cell fragments being more important than proteins.

We have now stated this result also in the Summary and Conclusion section. The reason why some ice-binding 'antifreeze' proteins act as at least moderate ice nucleators (such as the *Tenebrio molitor* *tmAFP*) and others show only minute or now ice nucleation activity (such as the *fclBP11* studied here) is not entirely clear, but a recent modelling study suggest that this may have to do with the ice planes that they usually bind to (see statement at the end of section 3). However, typical ice-nucleating proteins are usually much larger than these ice-binding 'antifreeze' proteins and their large ice-active site may therefore be much better suited in supporting a newly forming ice embryo. We have enhanced the section explaining the differences and similarities between ice-binding proteins, antifreeze proteins and ice-nucleating proteins that may also help understanding these phenomena:

The very good ice-binding properties of *fclBP* and EPS (mainly polysaccharides and proteins) under sea ice brine conditions have been reported in previous studies (Krembs et al., 2002; Bayer-Giraldi et al., 2011; Krembs et al., 2011). Ice-binding proteins (IBPs) bind to ice crystal surfaces and by doing so can control the crystal growth rate, inhibit ice recrystallization or help to adhere their host to ice (Davies, 2014; Bar Dolev et al., 2016; Guo et al., 2017). Originally, IBPs were known as antifreeze

(glyco)proteins, which protect fish and insects by thermal hysteresis, i.e. by depressing the temperature where active crystal growth occurs to below the equilibrium melting point temperature (Bar Dolev et al., 2016). However, not all IBPs have such thermal hysteresis antifreeze properties. For example, a recently discovered IBP from the Antarctic bacterium *Marinomonas primoryensis* binds its bacterial host to diatoms and the Antarctic sea ice layer (Guo et al., 2017). Furthermore, even ice-nucleating proteins are sometimes considered to be a subgroup of IBPs, because their active sites appear to be structurally similar, just much larger, than those of regular IBPs with antifreeze properties (Davies, 2014; Bar Dolev et al., 2016; Eickhoff et al., 2019; Hudait et al., 2019). These considerations may imply that the much smaller ice-binding sites of ‘antifreeze’ IBPs could also stabilize the formation of small ice embryos and thereby promote the nucleation of ice from liquid water, however, only at very low temperatures (Davies, 2014; Bar Dolev et al., 2016; Eickhoff et al., 2019; Hudait et al., 2019). Indeed, it has been shown both experimentally as well as in molecular dynamics simulations that the ice-binding antifreeze proteins of the mealworm beetle *Tenebrio molitor* (*tmAFP*) can also trigger the nucleation of new ice crystals just a few degree Celsius above the homogenous freezing temperature of water or an aqueous solution (Eickhoff et al., 2019; Hudait et al., 2019). Here, we explore whether a similar ice-nucleating effect also occurs for IBPs from *F. cylindrus*.

References:

- Bar Dolev, M., Braslavsky, I., and Davies, P. L.: Ice-Binding Proteins and Their Function, *Annual review of biochemistry*, 85, 515–542, doi:10.1146/annurev-biochem-060815-014546, 2016.
- Bayer-Giraldi, M., Weikusat, I., Besir, H., and Dieckmann, G.: Characterization of an antifreeze protein from the polar diatom *Fragilariopsis cylindrus* and its relevance in sea ice, *Cryobiology*, 63, 210–219, doi:10.1016/j.cryobiol.2011.08.006, 2011.
- Davies, P. L.: Ice-binding proteins: a remarkable diversity of structures for stopping and starting ice growth, *Trends in Biochemical Sciences*, 39, 548–555, doi:10.1016/j.tibs.2014.09.005, 2014.
- Eickhoff, L., Dreischmeier, K., Zipori, A., Sirotinskaya, V., Adar, C., Reicher, N., Braslavsky, I., Rudich, Y., and Koop, T.: Contrasting Behavior of Antifreeze Proteins: Ice Growth Inhibitors and Ice Nucleation Promoters, *The journal of physical chemistry letters*, 10, 966–972, doi:10.1021/acs.jpcclett.8b03719, 2019.
- Guo, S., Stevens, C. A., Vance, T. D. R., Olijve, L. L. C., Graham, L. A., Campbell, R. L., Yazdi, S. R., Escobedo, C., Bar-Dolev, M., Yashunsky, V., Braslavsky, I., Langelaan, D. N., Smith, S. P., Allingham, J. S., Voets, I. K., and Davies, P. L.: Structure of a 1.5-MDa adhesin that binds its Antarctic bacterium to diatoms and ice, *Science advances*, 3, e1701440, doi:10.1126/sciadv.1701440, 2017.
- Hudait, A., Qiu, Y., Odendahl, N., and Molinero, V.: Hydrogen-Bonding and Hydrophobic Groups Contribute Equally to the Binding of Hyperactive Antifreeze and Ice-Nucleating Proteins to Ice, *Journal of the American Chemical Society*, 141, 7887–7898, doi:10.1021/jacs.9b02248, 2019.
- Krembs, C., Eicken, H., and Deming, J. W.: Exopolymer alteration of physical properties of sea ice and implications for ice habitability and biogeochemistry in a warmer Arctic, *Proceedings of the National Academy of Sciences of the United States of America*, 108, 3653–3658, doi:10.1073/pnas.1100701108, 2011.
- Krembs, C., Eicken, H., Junge, K., and Deming, J.: High concentrations of exopolymeric substances in Arctic winter sea ice: implications for the polar ocean carbon cycle and cryoprotection of diatoms, *Deep Sea Research Part I: Oceanographic Research Papers*, 49, 2163–2181, doi:10.1016/S0967-0637(02)00122-X, 2002.

Revised manuscript with changes indicated

Ice nucleating properties of the sea ice diatom *Fragilariopsis cylindrus* and its exudates

Lukas Eickhoff¹, Maddalena Bayer-Giraldi², Naama Reicher³, Yinon Rudich³, Thomas Koop¹

¹Faculty of Chemistry, Bielefeld University, Universitätsstraße 25, 33615 Bielefeld, Germany

5 ²Alfred-Wegener-Institut Helmholtz-Zentrum für Polar- und Meeresforschung (AWI), Bremerhaven, Germany

³Department of Earth and Planetary Sciences, Weizmann Institute of Science, 76100 Rehovot, Israel

Correspondence to: Thomas Koop (thomas.koop@uni-bielefeld.de)

10 **Abstract.** In this study, we investigated the ice nucleation activity of the Antarctic sea ice diatom *Fragilariopsis cylindrus*. Diatoms are the main primary producers of organic carbon in the Southern Ocean and the Antarctic sea ice diatom *F. cylindrus* is one of the predominant species. This psychrophilic diatom is abundant in open waters and within sea ice, ~~and it.~~ It has developed several mechanisms to cope with the extreme conditions of its environment, for example, the production of ice-binding proteins (IBP) and extracellular polymeric substances, known to alter the structure of ice. Here, we investigated the ice nucleation activity of *F. cylindrus* using a microfluidic device containing individual sub-nanoliter~~e~~ (~90 µm) droplet samples. The experimental method and a newly implemented Poisson statistics-based data evaluation procedure applicable to samples with low ice nucleating particle concentrations were validated by comparative ice nucleation experiments with well-investigated bacterial samples from *Pseudomonas syringae* (Snomax). The experiments reveal an increase of 7.2 °C in the ice nucleation temperatures for seawater containing *F. cylindrus* diatoms when compared to pure seawater. Moreover, also *F.*

15 *cylindrus* fragments show ice-nucleation activity, while experiments with *F. cylindrus* ice-binding protein (*fcIBP*) show no significant ice nucleation activity. A comparison with experimental results from other diatoms suggests a universal behaviour of polar sea ice diatoms, and we provide a diatom mass-based parameterization of their ice-nucleation activity for use in models.

1 Introduction

25 Sea ice is a two-phase medium, composed predominantly of crystalline ice with embedded liquid channels and pockets (inclusions) where active life can take place. As seawater freezes, dissolved sea salt ions are segregated from the growing ice lattice and accumulate in liquid brine inclusions, which have a lower freezing point due to their high salinity. Its porous structure makes sea ice a habitat for various organisms and enables life within the liquid brine network. Higher irradiance levels in sea ice when compared to the seawater column represent an advantage for photosynthetically active microorganisms

30 populating the pore space (Eicken, 1992). During sea ice formation, most microorganisms from the water column remain

entrapped within the ice or are scavenged by floating ice crystals (Ackley and Sullivan, 1994). Species composition changes with the aging of ice and the stabilization of the brine channel system (Krembs and Engel, 2001), resulting in a dominance of diatom species producing “sticky” extracellular polymeric substances (EPS) with ice-adhering functions (Raymond et al., 1994).

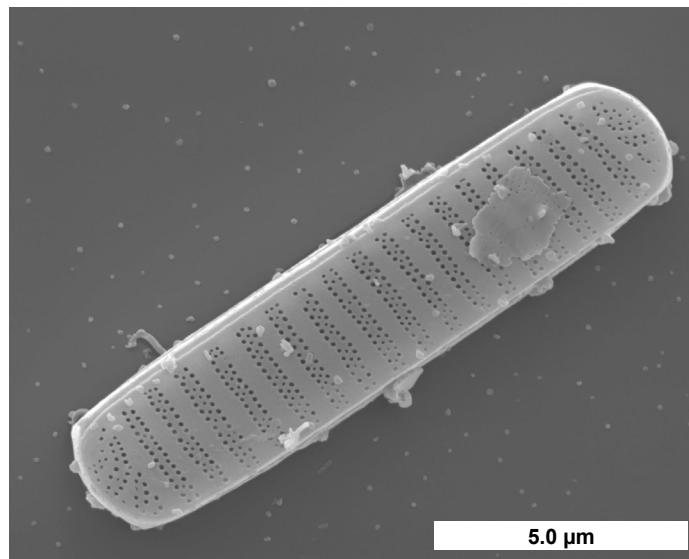
35

The diatom *Fragilariopsis cylindrus* (see Fig. 1) is widespread in polar environments and is one of the predominant species within the Arctic and Antarctic microbial assemblages (Kang and Fryxell, 1992; Poulin et al., 2011; van Leeuwe et al., 2018). The species thrives within sea ice, where it can be found ~~distributed~~ along the sea ice column (Bartsch, 1989; Garrison and Buck, 1989; Günther and Dieckmann, 2001; Poulin et al., 2011). It is, therefore, considered ~~as~~ an indicator of sea ice extent ~~d~~ in paleo-environmental studies for reconstructions of past variations (Gersonde and Zielinski, 2000). *F. cylindrus* is also abundant in the water column, for example in the proximity of the sea ice-edge zone (Kang and Fryxell, 1992; Lizotte, 2001) and in ice-covered waters (Garrison and Buck, 1989). *F. cylindrus* has developed a range of mechanisms s for coping with the extreme conditions occurring within sea ice (Mock et al., 2017). One prominent example is the production of so-called ice-binding proteins (IBPs) (Bayer-Giraldi et al., 2011), and of other EPS that are also found in other diatom species (Wilson et al., 2015; Aslam et al., 2018). *F. cylindrus* produces several IBP isoforms (*fc*IBPs), all of which belong to the broadly extended DUF3494 IBP family (Vance et al., 2019). It was shown that *fc*IBP isoform 11 affects the microstructure, i.e., the shape and size, of ice crystals (Bayer-Giraldi et al., 2011; Bayer-Giraldi et al., 2018). Moreover, EPS offer a protective environment to *F. cylindrus* in order to cope with the conditions of the sea ice habitat (Aslam et al., 2012a; Aslam et al., 2012b; Aslam et al., 2018). It has been suggested that *fc*IBPs accumulate in EPS and, in contact with the icy walls of brine inclusions, alter the pore space resulting in an increased habitability (Bayer-Giraldi et al., 2011).

40

45

50



55 **Figure 1:** *F. cylindrus* cell visualized by scanning electron microscopy. (Image courtesy of Henrik Lange and Friedel Hinz, Alfred Wegener Institute, Germany).

The very good ice-binding properties of *fc*IBP and EPS ~~may imply that the corresponding ice binding sites (mainly polysaccharides and proteins) under sea ice brine conditions have been reported in previous studies~~ (Krembs et al., 2002; Bayer-Giraldi et al., 2011; Krembs et al., 2011). Ice-binding proteins (IBPs) bind to ice crystal surfaces and by doing so can control the crystal growth rate, inhibit ice recrystallization or help to adhere their host to ice (Davies, 2014; Bar Dolev et al., 2016; Guo et al., 2017). Originally, IBPs were known as antifreeze (glyco)proteins, which protect fish and insects by thermal hysteresis, i.e. by depressing the temperature where active crystal growth occurs to below the equilibrium melting point temperature (Bar Dolev et al., 2016). However, not all IBPs have such thermal hysteresis antifreeze properties. For example, a recently discovered IBP from the Antarctic bacterium *Marinomonas primoryensis* binds its bacterial host to diatoms and the Antarctic sea ice layer (Guo et al., 2017). Furthermore, even ice-nucleating proteins are sometimes considered to be a subgroup of IBPs, because their active sites appear to be structurally similar, just much larger, than those of regular IBPs with antifreeze properties (Davies, 2014; Bar Dolev et al., 2016; Eickhoff et al., 2019; Hudait et al., 2019). These considerations may imply that the much smaller ice-binding sites of ‘antifreeze’ IBPs could also stabilize the formation of small ice embryos and thereby promote the nucleation of ice from liquid water. ~~The rationale behind this proposal is the fact that the active sites for ice binding and those for the promotion of ice nucleation appear to be quite similar and to match those of crystalline ice. And indeed, however, only at very low temperatures~~ (Davies, 2014; Bar Dolev et al., 2016; Eickhoff et al., 2019; Hudait et al., 2019). Indeed, it has been shown both experimentally as well as in molecular dynamics simulations that the ice-binding antifreeze proteins of the mealworm beetle *Tenebrio molitor* (*tmAFP*), ~~which normally prevent the growth of existing ice crystals at temperatures just below 0 °C,~~ can also trigger the nucleation of new ice crystals at lower temperatures just a few degree Celsius above the homogenous freezing temperature of water or an aqueous solution (Eickhoff et al., 2019; Hudait et al., 2019). Here, we explore whether a similar ice-nucleating effect ~~does occur~~ also occurs for IBPs from *F. cylindrus*.

Many ~~particles of~~ biological ~~origin~~ particles such as bacteria, viruses, or diatoms have been detected in the sea surface microlayer as well as in the thawing permafrost (Leck and Bigg, 2005; Wilson et al., 2015; Irish et al., 2017; Creamean et al., 2020; Ickes et al., 2020; Roy et al., 2021). Some of these biological particles ~~are able to~~ can increase the ice nucleation temperature of small water droplets and act as ice-nucleating particles INPs (DeMott et al., 2016; Ickes et al., 2020; Welti et al., 2020; Creamean et al., 2021; Hartmann et al., 2021; Roy et al., 2021). These biological al particles can be transported to the atmospheric boundary layer by sea spray aerosol droplets (Irish et al., 2019; Steinke et al., 2022). In the polar atmosphere, they can be transported over long distances (Šantl-Temkiv et al., 2019; Šantl-Temkiv et al., 2020). Sea spray aerosol contributes to ice nucleation under mixed-phase cloud conditions as well as at cirrus temperatures in the upper troposphere (DeMott et al., 2016; Hartmann et al., 2021; Wagner et al., 2021). Further experiments on diatoms and their EPS show that they ~~are able to~~ can promote ice nucleation in small droplets of water or seawater (Knopf et al., 2011; Wilson et al., 2015; Ickes et al., 2020; Xi et al., 2021). Thus, diatoms like *F. cylindrus* may ~~have effects on~~ affect ice nucleation in cloud droplets.

There are some differences regarding the relevance of INPs in the Arctic and Antarctic polar regions. While in both polar latitudes the absolute concentrations of INPs are low, the influence of anthropogenic aerosols and INPs is much larger in the Arctic due to long-range transport during the Arctic winter (Šantl-Temkiv et al., 2019; Šantl-Temkiv et al., 2020; Ekman and Schmale, 2022). During the Arctic summer, aerosol lifetimes are shorter due to increased wet removal preventing long range transport and thus increasing the importance of locally emitted INPs. In the Antarctic, the influence of anthropogenic aerosols and INPs is generally much smaller (Stohl and Sodemann, 2010; Ekman and Schmale, 2022). During winter, blowing snow from the sea ice is the main aerosol source in the southern polar region, while DMS and other organic compounds from algae bloom are the main source during summer.

2 Material and methods

2.1 Sampling and cultivation of the *F. cylindrus* diatoms

The investigated *F. cylindrus* cells belong to the strain TM99 isolated in 1999 from the sea ice of the Weddel Sea, Antarctica, by Thomas Mock (*Polarstern* ANT XVI/3 expedition). ~~Stock, which took place in the early spring from March to May 1999.~~ Since then, stock cultures were kept in *f/2* medium (Guillard and Ryther, 1962) set up with Antarctic water and cultivated at 0°C and under continuous illumination of approximately 25 $\mu\text{E m}^{-2} \text{s}^{-1}$. ~~Cell~~ Before the experiment, cell numbers of the *F. cylindrus* cultures were monitored using a Coulter Counter, and cells were harvested during the exponential growth phase. Cell cultures were distributed in 50 mL Falcon tubes each containing about 1×10^8 cells, and they were centrifuged at 0°C at 3220 g for 30 minutes. The clear spent *f/2* medium was carefully separated from the cell pellet by pipetting, and both were shock-frozen in liquid nitrogen and stored at -80°C.

2.2 Sample preparation

2.2.1 Preparation of artificial seawater

110 For the ice nucleation experiments, we used artificial seawater that mimics the natural conditions in the habitat of Antarctic *F. cylindrus* diatoms. The salinity in the Antarctic region is about 34.5, which corresponds to 34.5 g salts per 1000 g seawater (Roy-Barman and Jeandel, 2016), and we prepared artificial seawater of this salinity for dispersing the diatoms and as a reference for the ice nucleation experiments. For preparing the seawater, the six most important ions were considered, i.e., the cations Sodium, Potassium, Magnesium and Calcium and the anions Chloride and Sulphate, which together make up for about

115 99.4 % of the dissolved ions in seawater (Roy-Barman and Jeandel, 2016). ~~The required composition was obtained by dissolving 11.8446 g (202.68 mmol) NaCl, 0.3758 g (5.04 mmol) KCl, 5.3280 g (26.21 mmol) MgCl₂·6H₂O, 4.4902 g (13.94 mmol) Na₂SO₄·10H₂O and 0.7460 g (5.07 mmol) CaCl₂·2H₂O in 477.23 g (26.49 mol) double distilled water. More detailed information on~~ The composition of the salts and their concentrations are given in Supplemental Information Table S1. The artificial seawater was filtered through a syringe-filter (0.22 μm, Polyethersulfone, SimplePure) in order to exclude any effect

120 of suspended dust particles on ice nucleation. This filter has been used for all filtrations in this study unless otherwise mentioned. The samples were stored at a temperature of -18 °C before use.

2.2.2 Preparation of *F. cylindrus* samples

The initial *F. cylindrus* samples contained about 10⁸ diatoms per tube, see Sect. 2.1. These samples were placed in a micro reaction tube and were filled up with the filtered artificial seawater to a volume of 2 mL. The resulting stock suspension of

125 5×10⁷ cells per mL was used in all experiments. By further dilution with filtered artificial seawater, we generated several more dilute suspensions with concentrations of 1×10⁷, 2×10⁶, 1×10⁶ and 5×10⁵ cells per mL. For ice nucleation experiments on the fragments and exudates of the *F. cylindrus* cells, we have ~~also~~ filtered these five samples ~~using a syringe filter (0.22 μm, Polyethersulfone, SimplePure).~~

130 In order to identify the ice-nucleating entities of the *F. cylindrus* ~~cells~~ samples, we separated the different components by means of filtration and centrifugation. We filtered a ~~2×10⁶~~ 1×10⁷ cells per mL *F. cylindrus* suspension ~~using a syringe filter (0.22 μm, Polyethersulfone, SimplePure).~~, such that the *F. cylindrus* cells should remain in the filter while smaller fragments of destroyed cells and any soluble species such as soluble ice-binding protein *fcIBP11* should be able to pass the filter, see Fig. S1 in the Supplemental Information for details. Thereafter, we recovered the filter cake containing the whole *F. cylindrus* cells and

135 larger cell-fragments by shaking the filter in a vial with artificial seawater. Although we used the same volume of artificial seawater as for the preparation of the original cell suspension, we surmise that the concentration of the resuspended diatoms is lower than the initial concentration. From the comparison of the frozen fraction curves obtained with the sample with those of unfiltered samples (see below) our best estimate of the concentration is about 2×10⁶ cells per mL (estimated uncertainty range 1×10⁶ – 1×10⁷ cells per mL). Finally, the cell suspension was filtered again ~~(0.22 μm, Polyethersulfone, SimplePure)~~

140 for comparison with the pure artificial seawater sample. To verify the method, all steps were also done with a vial of pure artificial seawater without suspended *F. cylindrus* cells.

We also performed ice nucleation experiments on fresh *f/2* medium (Guillard and Ryther, 1962) as well as on the spent *f/2* 145 medium, in which the *F. cylindrus* diatoms were actually grown. The sample preparation procedure is described in detail in Supplemental Information Fig. S2. The spent *f/2* medium should not contain many cells, because they were separated by centrifugation. Nevertheless, we filtered the medium ~~with a syringe filter (0.22 μ m, Polyethersulfone, SimplePure),~~ such that only small fragments and soluble proteins (e.g., *fcIBP11*) should have remained in the filtrate (Bayer-Giraldi et al., 2011). In the next step, this sample was centrifuged using a 100 kDa centrifugal filter (Polyethersulfone, satorius Vivaspin 500, 15000g) 150 such that the remaining solution should not contain any diatom fragments but only smaller soluble molecules such as the soluble *fcIBP11* protein. For comparison, we also applied the identical centrifugation step with freshly prepared *f/2* medium that had never been in contact to any diatoms.

2.2.53 Preparation of *P. syringae* samples

In additional experiments, we verified our Poisson evaluation procedure (see Sect. 2.3.3). For this purpose, we used well- 155 studied bacterial cells of *P. syringae*, commercially available as Snomax[®], from the same batch as investigated in previous studies (Budke and Koop, 2015; Wex et al., 2015). The molecular mass of the individual ice-nucleating proteins in the bacteria is about 150 kDa (Wolber et al., 1986; Govindarajan and Lindow, 1988). A suspension of *P. syringae* with a concentration of 4 mg per mL was prepared from dry Snomax with double-distilled water. By diluting this stock suspension with further double-distilled water, we also prepared additional more dilute suspensions with concentrations of 1×10^{-2} , 2×10^{-3} and 1×10^{-3} mg per 160 mL. Using an average value of the cell number density of 1.4×10^9 cells per mg (Wex et al., 2015), these mass concentrations correspond to cell concentrations of 1.4×10^7 , 2.8×10^6 and 1.4×10^6 cells per mL.

2.2.64 Preparation of *fcIBP11*

Previous studies suggest that *fcIBP11* plays a major role in the response of *F. cylindrus* to freezing conditions (Bayer-Giraldi et al., 2010), by binding to ice and affecting ice crystal growth (Bayer-Giraldi et al., 2011; Bayer-Giraldi et al., 2018). ~~The *fcIBP11* protein belongs to the DUF3494 IBP family, which presently constitutes the most broadly spread IBP family often found in sea ice microorganisms.~~ 165 For our experiments, we used the recombinant *fcIBP* isoform 11 (EMBL Heidelberg), GenBank accession no. DR026070. The protein was expressed as previously described (Bayer-Giraldi et al., 2011) and resuspended in Tris-HCl buffer (pH 7.0). For determining the ice nucleation activity of *fcIBP11*, we prepared a stock solution with a *fcIBP11* concentration of 0.1 mmol L^{-1} . We diluted this sample by a factor of ten to a concentration of 0.01 mmol L^{-1} 170 using Tris-HCl buffer (pH 7.0) and performed ice nucleation experiments on both sample solutions with the modified WISDOM microfluidic experiment (Reicher et al., 2018; Eickhoff et al., 2019), see below.

2.3 Experimental methods for ice nucleation experiments

2.3.1 Differential scanning calorimetry

175 A classic method for the investigation of homogeneous and heterogeneous ice nucleation is differential scanning calorimetry (DSC) of emulsified droplets (Rasmussen and MacKenzie, 1972; Koop, 2004). Here, we used a DSC apparatus (TA-Instruments, DSC-Q100), which was described in detail previously including its calibration procedure (Riechers et al., 2013). As bulk samples notoriously suffer from unwanted impurities, we performed measurements of inverse water-in-oil emulsion samples containing micrometre-sized droplets. As many thousands of droplets are investigated simultaneously, such samples
180 allow the detection of very reproducible exothermic heterogeneous ice nucleation signals down to the homogeneous ice nucleation temperature of about -38°C (Pinti et al., 2012; Riechers et al., 2013; Dreischmeier et al., 2017). Further information on the emulsion preparation procedure is given in the Supplemental Information.

~~The principle preparation procedure for the water in oil emulsion (w/o) samples was almost identical to the method described earlier. 1 mL of 7 wt% emulsifier Span[®]65 (Merck) dissolved in 93 wt% of a mixture of 50 vol% methyleclopentane (Aeros Organics, 99 %) and 50 vol% methylecyclohexane (Aeros Organics, 95 %) was used as the organic phase. The aqueous phase consisted of 1 mL of an *F. cylindrus* suspension with a concentration of 1×10^7 cells per mL, see Sect. 2.2.2 above, or alternatively of 1 mL of pure artificial seawater for comparison. The mixtures of the organic and aqueous phase were subsequently emulsified by stirring with a high speed disperser (IKA Ultra-Turrax T25 basic) for 10 min at 20'000 rpm. For a DSC measurement, about 10 mg of such an emulsion was filled into an aluminium pan that was sealed hermetically and then
185 transferred into the calorimeter. The samples were cooled at a rate of -5°C per min down to -60°C , and subsequently reheated, first at 5°C per min and then, in the temperature range between -20°C and $+5^{\circ}\text{C}$, at 1°C per min.~~

~~The DSC experiment has been used as a simple and direct method to check whether *F. cylindrus* diatoms are potential ice nucleators or not. The method does not allow for the observation of single droplets, and we can only study cell fragments but not intact cells because the latter are disrupted during the emulsion preparation process. Therefore, we have used the WISDOM
190 microfluidic device, which is described below, as the main experimental method in this study.~~

2.3.2 WISDOM microfluidic device

Most of the ice nucleation experiments presented in this study were carried out using droplet microfluidics. In particular, we used a microfluidic device based upon the WISDOM (Weizmann Supercooled Droplets Observation on a Microarray) experiment (Reicher et al., 2018; Reicher et al., 2019), with some minor modifications for a setup operated at Bielefeld
200 University, including adapted temperature and heating rate calibrations, see a previous in-detail description (Eickhoff et al., 2019). These modifications and the general procedure for the sample preparation are given in the Supplemental Information.

For the droplet generation, we used two syringe pumps (neMESYS NEM-B101-02 E), one filled with the aqueous sample and another with an organic phase consisting of 2 wt% Span[®]80 (Merck) dissolved in 98 wt% of a mineral oil (Sigma-Aldrich, mineral oil M3516). The microfluidic chip was connected to the pumps with PTFE tubes. The droplets generated within the chip had diameters of $90\ \mu\text{m} \pm 5\ \mu\text{m}$.

For the freezing experiments, we placed the microfluidic chip after the droplet production on a temperature-controlled cold-stage (Linkam, BCS 196) attached to an optical microscope (Olympus, BX51 TRF). The temperature of the droplets in the chip was calibrated with respect to the cooling (or heating) rate as well as to the absolute temperature, and is described in detail in a previous study. The freezing of the droplets was observed using the transmission mode of the microscope and we recorded the images with a digital camera (Q-Imaging, MicroPublisher 5.0 RTV) for later analysis by a LabView routine that detects a freezing event from the change in grey values of a particular droplet upon freezing. Typical changes in the droplets' grey values during freezing experiments are depicted in Supplemental Information Fig. S3. In each individual experiment, between about 45 to 70 droplets were observed simultaneously, depending upon the percentage of droplet filled microcells within the droplet array of the chip.

For all *F. cylindrus* measurements, the chip was first cooled to a temperature of $-20\ ^\circ\text{C}$ at a rate of $5\ ^\circ\text{C}$ per min, because no freezing events were detected in this temperature range. After equilibration at this temperature for 2 min, the samples were then cooled at a slower rate of $1\ ^\circ\text{C}$ per min to $-45\ ^\circ\text{C}$, at which all droplets were frozen. Thereafter, the chip was heated relatively quickly at a rate of $5\ ^\circ\text{C}$ per min, until $-10\ ^\circ\text{C}$, and after two minutes of equilibration, it was then heated to $5\ ^\circ\text{C}$ at $1\ ^\circ\text{C}$ per min. The detailed temperature profiles for each type of experiment are listed in Supplemental Information Table S2.

2.3.3 Evaluation procedure for samples with small INP concentrations

Ice nucleation studies using larger-volume droplet arrays usually employ relatively high concentrations of INPs per droplet, e.g. mineral dust particles or bacterial cells, with (Budke and Koop, 2015; Hiranuma et al., 2015; Wex et al., 2015; DeMott et al., 2018; Hiranuma et al., 2019; Kunert et al., 2019; Ickes et al., 2020), to ensure that freezing is induced at a large number of INPs per droplet.

temperature that is higher than that triggered by the supporting surface or minute amounts of impurities contained in the water.

In the present study, the INP concentrations were much lower total amount of INPs was small due to the limited availability of *F. cylindrus* cells, suggesting the use of small droplet methods which require less total INP material. We investigated droplets with a diameter of $90\ \mu\text{m}$, corresponding to a volume of about 380 pL. Another, probably more important advantage of using these small droplet volumes is that we can measure ice nucleation down to the homogenous freezing temperature of water (Riechers et al., 2013; Reicher et al., 2018; Tarn et al., 2021), enabling also the investigation of rather poor ice nucleators. As the concentrations c of *F. cylindrus* cells varied between 5×10^5 and 5×10^7 cells mL^{-1} , the corresponding average INP concentrations were ranged between about only 0.19 up to and 19 diatom cells per droplet. It becomes immediately clear that when the average INP concentration λ is smaller than 1, i.e. on average less than one cell per droplet, there must be droplets devoid of any cells, because the number of cells in an individual droplet can only be an integer (assuming only whole cells –

235 without fragments – being present). In such a case, heterogeneous ice nucleation cannot be triggered in every droplet, but only
in those containing at least one cell. Hence, homogeneous ice nucleation is to be expected to occur in the ‘empty’ droplets.
Moreover, even if the average INP concentration λ is exactly one per droplet, there will be a few droplets that contain two or
more INPs and, thus, other droplets that do not contain any INPs. The distribution of INPs among microfluidic droplets at
small average INP concentration can be described using Poisson statistics (Huebner et al., 2007; Köster et al., 2008; Edd et al.,
240 2009; Collins et al., 2015). ~~The following Poisson distribution can be used to describe the probability $P_x(k)$ that an individual
droplet contains exactly k INPs when the average concentration is λ INPs per droplet. The detailed documentation of this
procedure is given in the Supplemental Information.~~

$$P_x(k) = \frac{\lambda^k}{k!} \exp(-\lambda). \quad (1)$$

Note that the derivation of the Poisson distribution contains a simplification that require a larger number of droplets and hence
245 Eq. (1) becomes more accurate as the number of investigated droplets increases. For the microfluidic experiments performed
in this work with more than a hundred droplets investigated for each sample the simplification applies.

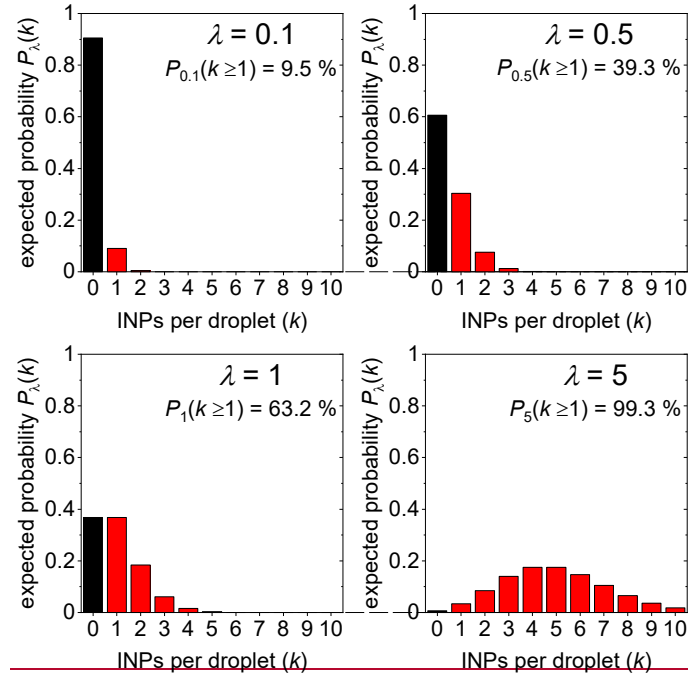
The average number of INPs per droplet, λ , is easily calculated from the concentration c of INPs in the stock solution and the
volume V of an individual microfluidic droplet:

$$\lambda = c \cdot V_{\text{droplet}}. \quad (2)$$

250 Furthermore, the droplet volume V_{droplet} can be expressed by the droplet’s radius r or alternatively by its diameter d :

$$V_{\text{droplet}} = \frac{4}{3} \pi \cdot r^3 = \frac{\pi}{6} \cdot d^3. \quad (3)$$

Figure 2 shows the calculated Poisson distributions of the number of cells per droplet for four different values of λ in a
concentration range relevant to this study. For lower values of λ , the histograms exhibit the tilted shape typical of Poisson
distributions, while for larger values of λ , the Poisson distribution approaches the more symmetrical shape of a normal
255 distribution.



260

Figure 2: Calculated probability $P_x(k)$ of the number k of INPs per droplet for different values λ of the average cell concentration per droplet. The black coloured bars indicate the probability for the occurrence of droplets without any INPs, while the red coloured bars indicate the combined probability $P_x(k \geq 1)$ for droplets containing at least one INP. The corresponding values for $P_x(k \geq 1)$ are annotated in each panel for different values of λ .

265

For the ice nucleation experiments considered here, only those droplets containing at least one INP and those without any INPs are relevant, as this determines whether they are subject to heterogeneous or homogeneous nucleation, respectively. Whether a droplet contains one, two or more INPs is of less importance, as long as every INP is identical and, thus, induces heterogeneous ice nucleation at the same temperature. The probability that a droplet does not contain any INPs can be calculated easily by inserting $k = 0$ into Eq. (1):

$$P_x(0) = \frac{\lambda^0}{0!} \exp(-\lambda) = \frac{1}{1} \exp(-\lambda) = \exp(-\lambda). \quad (4)$$

270

$P_x(0)$ is shown as the black coloured bar in each panel of Fig. 2. The probability that a droplet contains at least one INP, $P_x(k \geq 1)$, is given by the combined probability of all red coloured bars in each panel of Fig. 2, and it can be calculated using the fact that the sum of all probabilities $P_x(k)$ for k from 0 to ∞ must become 1 (see Eq. (5)):

$$P_x(k) = \sum_{k=0}^{\infty} \frac{\lambda^k}{k!} \exp(-\lambda) = 1. \quad (5)$$

Hence, $P_x(k \geq 1)$ can be calculated from the following difference

$$P_x(k \geq 1) = \sum_{k=1}^{\infty} \frac{\lambda^k}{k!} \exp(-\lambda) = \sum_{k=0}^{\infty} \frac{\lambda^k}{k!} \exp(-\lambda) - \sum_{k=0}^0 \frac{\lambda^k}{k!} \exp(-\lambda) = 1 - P_x(0) = 1 - \exp(-\lambda). \quad (6)$$

275 Since λ can be expressed by the product of the droplets' diameter and the known concentration of INPs in the stock solution, ϵ , (see Eq. (2) and Eq. (3)) this yields:

$$P_x(k \geq 1) = 1 - \exp\left(-\frac{\pi}{6} \cdot c \cdot d^3\right). \quad (7)$$

280 **The equations above have been derived for applications where the average concentration c of INPs in solution is known. However, in ice nucleation experiments of natural samples, the concentration c of INPs per volume is often unknown a priori and other values such as the organic carbon content has to be used for comparison. In such cases, Eq. (7) can be used to obtain a rough estimate of λ and, thus, c from ice nucleation experiments when a plateau in the experimental frozen fraction curve is observed. The frozen fraction is defined as the number of frozen droplets relative to the number of all droplets, at a given temperature. Here, we term the value of the frozen fraction at the plateau as f_{ice}^t . If a sufficiently large number of droplets is investigated, then f_{ice}^t corresponds to the fraction of droplets that froze heterogeneously and thus may be equated with that fraction of droplets containing at least one INP.**
285 2.4 Elemental analysis

The total carbon content of the *F. cylindrus* samples has been determined using elemental analysis. For this purpose, an amount of 0.7 mg *F. cylindrus* diatoms was combusted at a high temperature ($T > 1000$ °C) in a Tin-crucible and the composition was analysed using a commercially available elemental analyser (EuroVector, Euro EA).

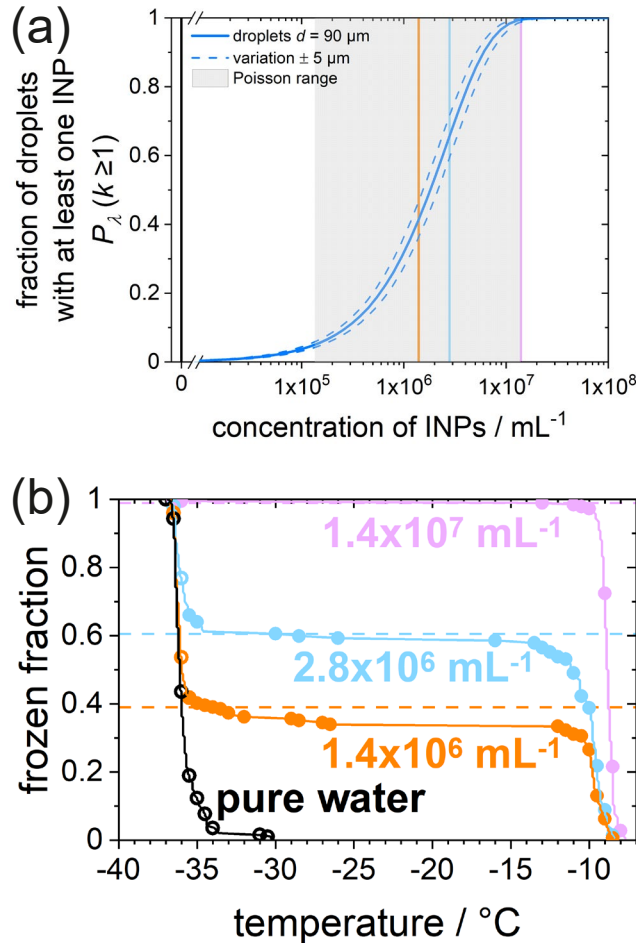
290 ~~There are~~ two underlying assumptions for experimentally obtaining f_{ice}^t . First, every droplet containing at least one INP freezes heterogeneously, which appears entirely reasonable. Secondly, every droplet containing one or more INPs freezes at a higher temperature than those droplets without any INP, i.e. the difference between the heterogeneous and homogeneous ice nucleation temperature is large enough to be easily distinguished in the experiment. With these two assumptions a plateau in the frozen fraction curve can be interpreted as follows: the fraction of droplets below the plateau froze heterogeneously and contain at least one INP, and the fraction of droplets above the plateau froze homogeneously (when their freezing temperature is consistent with homogenous freezing) and, thus, do not contain INPs. In practise, this evaluation procedure does not work if none of the droplets froze heterogeneously or if all droplets froze heterogeneously at the same temperature without any obvious plateau, i.e. it is only applicable for intermediate average INP concentrations in what we term the “Poisson relevant range”.

300 We define this “Poisson relevant range” as the range of average INP concentrations, in which both droplets without any INP as well as droplets containing one or more INPs occur and, thus, both can be observed readily in the corresponding freezing experiments. For the experiments presented here, we establish the Poisson relevant range as the area between $P_x(k \geq 1)$ values of 5.0 % and 99.5 %. The lower limit was set at 5.0% in order to avoid any influence of the freezing of a minor fraction of droplets induced by impurities, the upper limit corresponds to about one out of 200 droplets not containing any INP and thus freezing homogeneously. For higher concentrations, when every droplet contains at least one INP, the above Poisson evaluation is not needed and the classic method can be used, and so this upper limit sets an endpoint for the Poisson based

305

evaluation. The classic method indeed assumes that every observed droplet contains at least one INP and it has been described in detail previously.

310 To demonstrate the concentration range suitable for the Poisson method, i.e. the Poisson relevant range, the latter is indicated in Fig. 3a as the grey shaded area. The solid blue curve shows the values of $P_x(k \geq 1)$ calculated using Eq. (7) as a function of the average INP concentration c of the studied sample and a droplet diameter of $90 \mu\text{m}$. The two dashed lines show the changes for a deviation of $\pm 5 \mu\text{m}$ in droplet diameter.



315 **Figure 3:** (a) Fraction of droplets containing at least one INP, $P_x(k \geq 1)$ as a function of INP concentration in the investigated *P. syringae* sample. The solid blue curve represents the values of $P_x(k \geq 1)$ for the droplets in the WISDOM experiment with a diameter of $90 \mu\text{m}$, calculated using Eq. (7), the dashed curves indicate the uncertainty for a variation of $\pm 5 \mu\text{m}$ in droplet diameter. Eq. The grey shaded area shows the Poisson relevant range, with the lower and upper limits at the concentrations corresponding to $P_x(k \geq 1) = 0.050$ and $P_x(k \geq 1) = 0.995$, respectively. The coloured vertical bars mark the experimentally investigated concentrations of *P. syringae*: $1.4 \times 10^6 \text{ mL}^{-1}$ (orange), $2.8 \times 10^6 \text{ mL}^{-1}$ (blue), and $1.4 \times 10^7 \text{ mL}^{-1}$ (purple) and pure water (black). A comparable plot for the *F. cylindrus* diatoms can be found in Fig. S4. (b) Fraction of frozen droplets as a function of temperature for different concentrations of *P. syringae* bacteria in double distilled water (coloured) and pure double distilled water (black) for reference. The horizontal lines mark the values for $P_x(k \geq 1)_{\text{measured}}$, see text.

320

325 Data points of frozen fractions are binned in temperature intervals of 0.5 °C (intervals without freezing events are not shown). Filled circles represent droplets containing *P. syringae* cells (based on calculations for $P_x(k \geq 1)$ with Eq. (7)), in which freezing was induced heterogeneously. Open circles represent droplets that should not contain *P. syringae* according to the calculations and, thus froze homogeneously.

To verify the procedure, we investigated aqueous suspensions of the well studied ice nucleating bacterium *Pseudomonas syringae* in the form of the commercial product Snomax. The ice nucleation temperatures of each about 165 ± 15 droplets, from 330 three single measurements with 45 to 70 droplets each, containing either pure double distilled water or three different concentrations of *P. syringae* were investigated, see Supplemental Information Table S3. These concentrations are also marked in Fig. 3a as vertical lines. A similar plot for the *F. cylindrus* diatoms can be found in Fig. S4. The resulting experimental frozen fraction curves of *P. syringae* are shown in Fig. 3b. Double distilled water (black open symbols) shows a steep increase in frozen fraction below about -34.0 °C, in agreement with homogeneous ice nucleation rates of droplets of such diameter. 335 Following this observation, all droplets of the *P. syringae* samples that froze at around or below this temperature are assumed to have nucleated homogeneously, i.e. they are considered to contain no INPs in the analysis below.

For all *P. syringae* samples, the first freezing events occur at much higher temperatures of about -8 to -9 °C, and the frozen fraction curve in each case initially increases strongly before reaching a plateau, and subsequently the remaining liquid droplets freeze only at very low temperatures. In each sample, the plateau occurs at a different value of the frozen fraction, e.g. f_{ice}^t is 340 higher the larger the *P. syringae* concentrations (pink > blue > orange). We determined the corresponding f_{ice}^t values, as defined above, from the experimentally obtained frozen fraction curve as the value of the frozen fraction at -34.0 °C, i.e. at the threshold between heterogeneous and homogenous ice nucleation as defined above. The resulting f_{ice}^t values for the three concentrations were 0.99, 0.61, and 0.39, respectively, indicated as the dashed horizontal lines in Fig. 3b. These f_{ice}^t values correspond to $P_x(k \geq 1)_{measured}$ and can be used to infer the average INP concentration from Eq. (7). Because in the current 345 experiments the INP concentrations are known (i.e., 1.4×10^7 , 2.8×10^6 , and 1.4×10^6 mL⁻¹), these experimentally derived f_{ice}^t values can be compared to the expected f_{ice}^t values, corresponding to $P_x(k \geq 1)_{calculated}$ values calculated from Eq. (7), yielding values of 1.00 ± 0.01 , 0.66 ± 0.06 , and 0.41 ± 0.05 , respectively. These theoretical values are in good agreement (within experimental uncertainty) with the measured values and thus confirm our approach and the inferred INP concentrations of 1.2×10^7 , 2.5×10^6 , and 1.3×10^6 mL⁻¹ (see Supplementary Table S3) deviate by about 14%, 11% and 7% from the prepared 350 concentrations, which is very good given that INP concentrations can vary by orders of magnitude. For further validation that the Poisson distribution is necessary for a proper evaluation in the above mentioned concentration range, the cumulative number of active ice nucleating sites n_N per number of *P. syringae* bacteria was evaluated and discussed in the text describing Supplementary Fig. S5.

3 Results and Discussion

355 3.1 Ice nucleation of *F. cylindrus*

3.1.1 Differential Scanning Calorimetry

As an initial experiment initially, the ice nucleation activity of *F. cylindrus* diatom cells was studied by differential scanning calorimetry. For these measurements, an inverse emulsion (DSC). We found that fragments or exudates of pure artificial seawater *F. cylindrus* diatoms are potential ice nucleators as a reference was compared with an emulsion of artificial seawater the sample containing 1×10^7 *F. cylindrus* cells per mL, see Fig. 4.

360

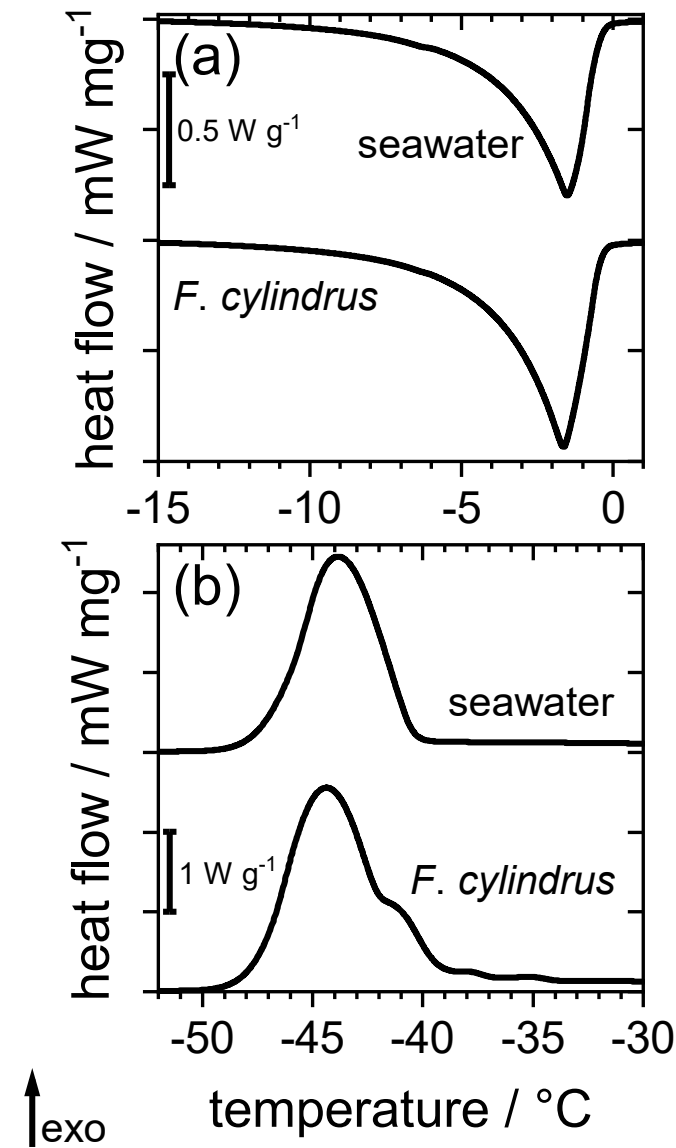


Figure 4: Comparison of DSC thermograms of water in oil emulsions containing pure artificial seawater and artificial seawater with *F. cylindrus* cells. (a) The endothermic melting signals are almost identical for pure seawater and seawater containing diatoms. (b) Exothermic diatoms induce freezing signals for pure seawater and seawater containing diatoms. While the seawater emulsion shows only one freezing signal, the emulsion containing *F. cylindrus* shows the same signal but with a shoulder and smaller signals at higher temperature, indicative of diatom-induced heterogeneous ice nucleation.

First, the endothermic ice melting signals of the reference and the sample in Fig. 4a show almost the same signal, indicating that any colligative effect of the diatoms is negligible when compared to the amount of the dissolved ions in the artificial seawater. This similarity in the ice melting signals also implies no change in water activity of the artificial seawater upon the addition of the diatoms and, thus, no colligative effect on the homogeneous ice nucleation (freezing) signals is to be expected. The exothermic freezing signals for both emulsions are shown in Fig. 4b. For the seawater reference, one distinct nearly symmetrical freezing signal is revealed with a maximum at about $-44\text{ }^{\circ}\text{C}$ and an onset at about $-40\text{ }^{\circ}\text{C}$. In contrast, the *F. cylindrus* sample shows the same maximum, but in addition a second exothermic signal in the form of a shoulder at about $-42\text{ }^{\circ}\text{C}$, with an onset at a somewhat higher temperature of $-39\text{ }^{\circ}\text{C}$ when compared to the reference, and with small signals as high as $-34\text{ }^{\circ}\text{C}$.

The larger broad signal in both emulsion samples corresponds to the homogeneous ice nucleation temperature of artificial seawater. This signal is also observed in the *F. cylindrus* sample because many of the emulsion droplets in that sample do not contain diatoms. The exothermic shoulder of the signal, which is not present in the reference, is most likely due to the freezing of droplets containing a diatom cell or fragment, and the shift of the onset to higher temperature is a first indication for the heterogeneous ice nucleation activity of the diatoms.

Because of the fact that the diatoms are of similar size as the emulsion droplets and the potential of mechanical disruption of diatom cells during the fast stirring of the disperser during emulsion preparation, these emulsion. The results are described in detail in the Supplemental Information and in Fig. S8. These experiments appear to us as not suitable for a quantitative analysis of the ice nucleation activity of *F. cylindrus*. Thus, we employed non-invasive methods in the experiments described below.

3.1.2 Droplet Microfluidics

initiated a more detailed study using the WISDOM-microfluidic device. First, we investigated the ice-nucleating properties of samples containing *F. cylindrus* diatoms, diatom cells, as well as fragments and exudates, at different concentrations suspended in artificial seawater. ~~For this purpose, we made use of~~ We used the droplet microfluidic devices described in Sect. 2.3.2 above. The results of these experiments are presented in Fig. 5a2a, which shows, as a function of temperature, the frozen fraction of droplets f_{ice} , commonly defined as the cumulative number of droplets frozen when cooled to a certain temperature relative to the total number of droplets (Murray et al., 2012). Thus, f_{ice} is practically independent of the total number of droplets investigated in a particular experiment. In our case, the number of droplets varied between 45 and 70 droplets per single measurement, and typically three single measurements per sample were performed. Figure 5a2a shows that the freezing temperatures of all *F. cylindrus* samples (coloured symbols) are higher than that of the artificial seawater reference sample (grey symbols), hence supporting the observations from the DSC experiments above that the *F. cylindrus* diatoms promote ice

nucleation. To compare the different samples, we use the T_{50} temperature, which is defined as that temperature at which half of the observed droplets are frozen, i.e. $f_{ice} = 0.5$. For the artificial seawater, we measured a T_{50} of -40.1 °C, and T_{50} of the *F. cylindrus* suspensions ~~is~~ shifted to a higher temperature by between about 2.8 °C to 7.2 °C with increasing diatom concentration. Detailed information on the increase in T_{50} of the different concentrations is given in ~~the~~ Supplemental Information Table S4. This significant concentration dependence of the T_{50} shift reveals that not all diatoms nucleate ice at exactly the same temperature and implies a distribution of the ice nucleation efficiency as has been observed previously also for other ice nucleators (Herbert et al., 2014; Budke and Koop, 2015).

cell concentration / mL⁻¹ # frozen droplets

●	5×10^7	●	1×10^6	○	○	○	○	○					
●	1×10^7	●	5×10^5	○	5	○	15	○	25	○	35	○	45
●	2×10^6	○	artificial seawater										

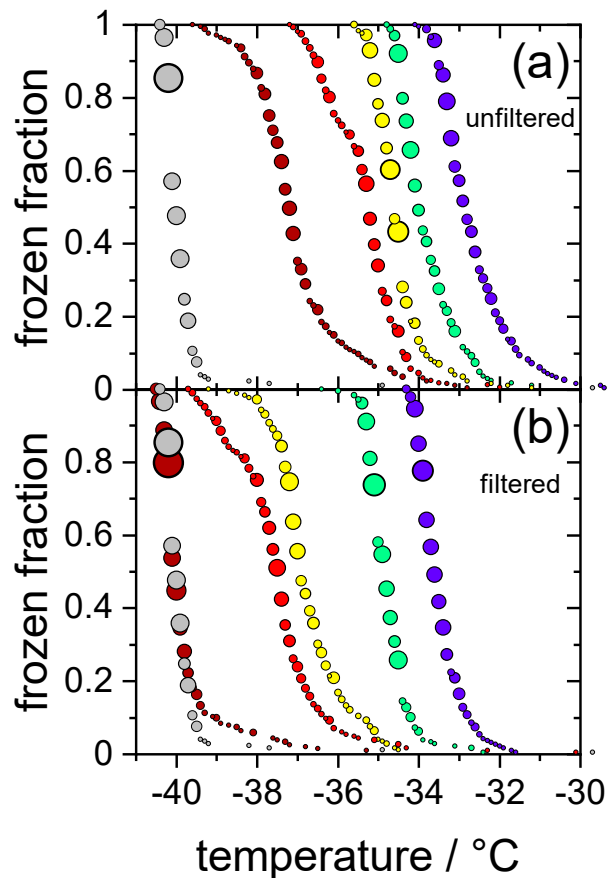
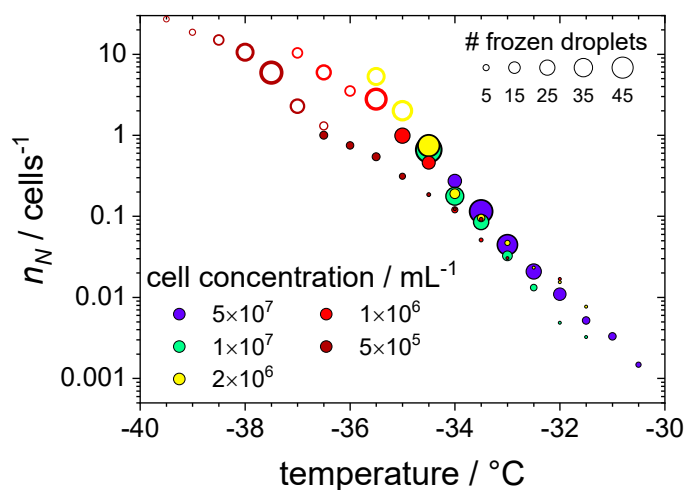


Figure 52: Cumulative fraction of frozen droplets as a function of temperature for different *F. cylindrus* cell concentrations (coloured circles) and pure artificial seawater (grey circles) as a reference. The size of the circles indicates the number of droplets frozen within the same temperature interval (0.1 °C). EachEvery dataset combines three individual measurements containing each between 45 and 70 droplets. **(a):** Frozen fraction curves for the five *F. cylindrus* samples, containing mostly whole diatoms and, probably, some fragments. **(b):** Freezing temperatures of the ~~same samples shown in panel (a), but after filtration with a pore size of 0.22 μm. filtered (0.22 μm) samples.~~ These

410 samples, thus, contain no whole cells but fragments as well as proteins and other soluble components. Note that the concentrations refer to the diatom concentrations before filtration. The seawater reference (grey circles) is the same in both panels.

This ~~fact can be~~ visualized better by plotting the cumulative number n_N of ice nucleating sites per number of *F. cylindrus* ~~diatoms~~ diatom cells, defined in Eq. (S2S9), as a function of freezing temperature, see Fig. 63. This n_N value is independent of the concentration of investigated INP and of the size of the investigated droplets, but can be measured for a wide range of temperatures using different concentrations, and allows for the comparison with results from other experimental techniques (see discussion below). Figure 63 reveals that at $-30.0\text{ }^{\circ}\text{C}$, $\sim 0.1\%$ of the *F. cylindrus* ~~diatoms~~ diatom cells promote ice nucleation, which increases to $\sim 1\%$ at $-32.0\text{ }^{\circ}\text{C}$ and $\sim 10\%$ at $-33.5\text{ }^{\circ}\text{C}$. Between about $-35.0\text{ }^{\circ}\text{C}$ and $-36.5\text{ }^{\circ}\text{C}$ all ~~diatoms~~ *F. cylindrus* cells trigger the nucleation of ice, i.e. $n_N = 1$. By definition, n_N values larger than one should not be possible, because it would imply that one diatom can induce the freezing of more than one droplet, which is unreasonable. The highest n_N values occur at the lowest diatom concentrations and, therefore, we must consider ~~the Poisson range defined above~~ statistics, i.e. whether or not each droplet ~~does~~ indeed contains a diatom cell. Following the treatise mentioned in Sect. 2.3.3 and outlined in detail in the Supplemental Information, and using Eq. (7S7), we indicate in Fig. 63 all the droplets that contain at least one diatom as filled circles, while all droplets that do not contain any *F. cylindrus* diatom cells are displayed as open circles. This analysis reveals a relatively sharp transition between filled and unfilled circles at n_N values of about one ice nucleating active site per diatom cell. All droplets frozen at $n_N \geq 1$ (and lower temperatures) do not contain intact *F. cylindrus* ~~diatoms~~ cells. We suggest that their freezing is induced by cell fragments or by INPs released by the *F. cylindrus* diatoms, e.g. soluble species from the EPS such as proteins or polysaccharides. A similar behaviour has been observed previously for birch pollen that release about 10^4 ice nucleators per pollen particle, which turned out to be ice-nucleating macromolecules (Pummer et al., 2012; Augustin et al., 2013; Pummer et al., 2015; Dreischmeier et al., 2017).



430 **Figure 6:** ~~Cumulative~~ The cumulative number of ice nucleating sites n_N per number of *F. cylindrus* diatom cells as a function of temperature, obtained from the data shown in Fig. 5a2a with the help of Eq. (S2S9). The original data were binned into intervals of $0.5\text{ }^{\circ}\text{C}$. The size of the ~~circle~~ symbols indicates the absolute number of droplets frozen in a particular bin, and the cell concentrations per mL are

435 indicated by colour. The filled circles represent the droplets that contain whole *F. cylindrus* cells, while Poisson statistics suggest that the open circles should not contain any ~~whole~~intact diatoms but probably some cell fragments, see text.

To verify the above interpretation, we performed experiments in which the samples from the measurements shown in Fig. ~~5a2a~~ and Fig. ~~63~~ were filtered with a pore size of 0.22 μm . This procedure ~~should remove~~removes intact whole diatoms, ~~whose size is which are~~ about 4.5 to 74 μm for the apical axis and 2.4 to 4 μm for the transapical axis (Lundholm and Hasle, 2008; Cefarelli et al., 2010). In Fig. ~~5b2b~~, the cumulative fraction of frozen droplets of these filtered samples is shown. The symbol colours represent the same suspensions as shown in Fig. ~~5a2a~~, but ~~this time~~ filtered, and the artificial seawater reference data is identical to that in panel (a). All frozen fraction curves are shifted to lower temperatures when compared to the unfiltered samples, suggesting a significant but not entire removal of INPs. Only the filtrate of the suspension with the ~~small~~lowest concentrations reveals a T_{50} that is the same as the seawater reference (-40.1 $^{\circ}\text{C}$), suggesting that this sample does not contain any significant concentration of INPs after filtration. All other filtrated suspensions show T_{50} values that are higher by between 2.6 $^{\circ}\text{C}$ and 445 6.4 $^{\circ}\text{C}$ relative to the seawater. For further information on the T_{50} shifts, see Supplementary Table S4. Together these results imply that ~~indeed~~ either fragments of *F. cylindrus* or molecules released by the diatoms can nucleate ice, but with a significantly reduced efficiency than intact diatoms. Moreover, these results can also explain the observations in Fig. ~~63~~ of ice nucleation of droplets at $n_N \gtrsim 1$ that ~~should do~~ not contain any ~~diatoms~~full diatom cells. Below, we present further experiments to investigate the nature of the ice-nucleating particles.

450 **3.32 Ice nucleation of resuspended *F. cylindrus* cells**

In the following experiments, we ~~tried to separate diatoms~~separated diatom cells from their fragments or ~~any~~ released INPs. For this purpose, the sample suspension of *F. cylindrus* with a concentration of ~~1x10⁷~~ cells per mL, which was shown already in Fig. ~~5 above 2~~, was analysed further, and the results are presented in Fig. ~~74~~. The green data points are those of the unfiltered sample and is identical to that shown in Fig. ~~5a2a~~, and the magenta data points is are identical to the filtered solution already 455 presented in Fig. ~~5b2b~~ (there as green data points). This ~~sample~~ suspension should contain only INPs smaller than 0.22 μm . Next, most (but not all) of the diatom cells and fragments contained in the filter cake of that filtration procedure were resuspended in artificial seawater. Thus, the concentration of the resuspended cells is ~~probably significantly smaller than about~~ 2×10^6 cells per mL (estimated uncertainty range $1 \times 10^6 - 1 \times 10^7$ cells per mL). The frozen fraction of that sample is shown as the orange data points in Fig. ~~74~~ and shows the same ~~onset~~ ice nucleation onset temperature of about -32.5 $^{\circ}\text{C}$ as the original 460 unfiltered suspension (green); ~~h~~ However, the curve is ~~much~~ broader, suggesting that it ~~indeed contain much~~contains less of the most active ice nucleators. ~~In order to~~ To verify that all fragments smaller than 0.22 μm had been leached out during the first filtration step, this resuspended filter cake sample was filtered again with a 0.22 μm filter. The results of this procedure on the freezing behaviour is are shown as the blue circles in Fig. ~~7. These 4. The~~ frozen fraction data are is practically identical to that of the artificial seawater, ~~strongly~~ suggesting that ~~indeed~~ filtration of the pure whole cells had ~~s~~ been successful and 465 hardly any fragments smaller than 0.22 μm are left in the filtrate. This analysis also implies that the ice nucleation of the unfiltered suspension is due to whole cells as well as cell fragments; but not due to ice-nucleating molecules released from the

diatoms. The T_{50} shift upon filtration of about 1.5 °C is similar in magnitude to the effect of reducing the concentration of the unfiltered diatoms from 5×10^7 cells per mL to 1×10^7 cells per mL, i.e. by a factor of 5. This similarity may indicate that fragments make up about 10-20% of the INPS in the unfiltered samples, which ~~is in agreement~~ agrees with the fact that some ice nucleation is observed for values of $n_N \gtrsim 1$, see Fig. 63.

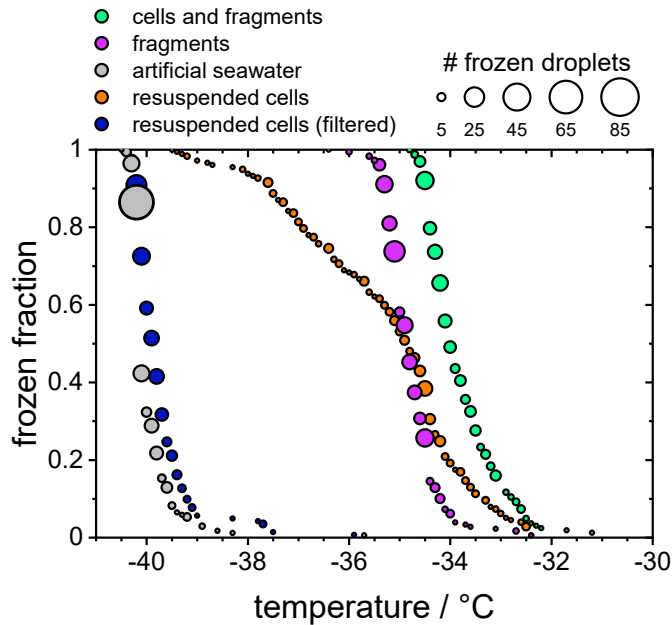


Figure 74: The frozen fraction of a sample with 1×10^7 *F. cylindrus* diatoms per mL after different treatments. The symbol size indicates the total number of droplets frozen at that temperature. The green coloured data are the untreated sample and are the same as those in Fig. 5a2a. The magenta data are the filtered sample that should just contain fragments of the diatoms. It is the same data as the green data in Fig. 5b2b. The grey data points show the freezing of the artificial seawater for reference (also replotted from Fig. 5.2). The orange data show the freezing of the diatoms that were resuspended from the filter into artificial seawater. Its concentration is likely smaller than 1×10^7 cells per mL, because not all cells could be resuspended. The blue data points represent the freezing of the droplets consisting of the resuspended cell suspension after renewed filtration: it should not contain any diatoms or fragments.

3.43 Ice nucleation of spent medium and of purified *fcIBP11*

We also investigated the spent *f/2* medium (Guillard and Ryther, 1962), i.e., the medium in which the *F. cylindrus* diatoms were cultivated before they were separated by centrifugation to investigate their ice nucleating effects. Separation of the diatoms from the spent *f/2* medium by centrifugation is not perfect and, hence, smaller fragments, as well as soluble macromolecules such as proteins, may remain in the spent medium. These may be potential ice nucleators, as it has been shown previously that even smaller ice-binding antifreeze proteins can act as ice nucleators at lower temperatures (Eickhoff et al., 2019).

In Fig. 85 we compare the frozen fraction curve for the spent *f/2* medium (light green circles) with that of a freshly prepared *f/2* medium, which never had been in contact ~~to~~with any *F. cylindrus* diatoms (olive circles). Clearly, the spent medium, even

after centrifuging off the diatoms, shows significant ice nucleation with a T_{50} of about $-35.7\text{ }^{\circ}\text{C}$, while the T_{50} of the fresh medium is much lower at $-40.0\text{ }^{\circ}\text{C}$. In additional experiments, the spent medium has been filtered in two further steps, first by using a $0.22\text{ }\mu\text{m}$ syringe filter (light blue circles) and then by using a 100 kDa centrifugation filter (pink circles). For comparison the fresh medium has been also filtered with a 100 kDa centrifugation filter (purple circles). Obviously, filtration of the spent medium with a $0.22\text{ }\mu\text{m}$ filter shows hardly any effect on ice nucleation as its T_{50} is shifted to $-36.0\text{ }^{\circ}\text{C}$, which is the same as the unfiltered sample within the temperature uncertainty of our setup of $\pm 0.3\text{ }^{\circ}\text{C}$.

495

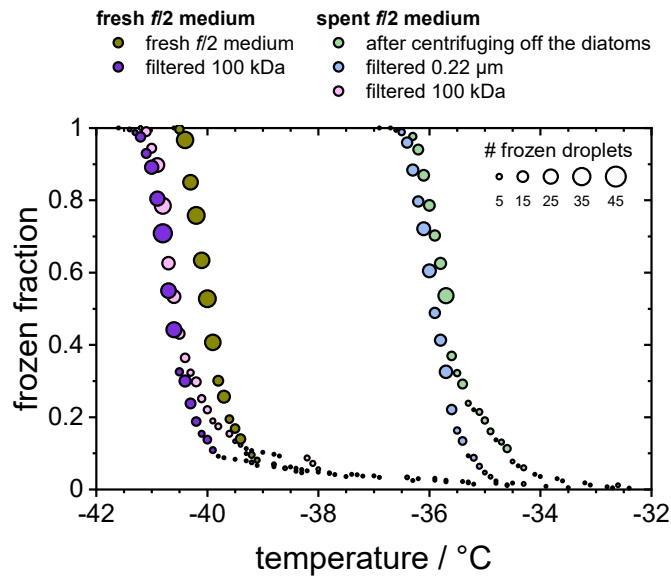


Figure 85: Frozen fraction of differently treated *f/2* media as a function of temperature. The olive and purple circles belong to a fresh *f/2* medium that is untreated (olive) or had been filtered using a 100 kDa filter (purple). The green, blue and pink circles belong to the untreated, $0.22\text{ }\mu\text{m}$ filtered and 100 kDa filtered spent medium, in which the *F. cylindrus* diatoms had grown before they were centrifuged and separated from the medium.

In contrast, filtration with a 100 kDa filter ~~resulted in a strongly~~substantially reduced ~~the~~ ice nucleation with a T_{50} value of $-40.6\text{ }^{\circ}\text{C}$, which is the same as that of the filtrated fresh medium of $-40.7\text{ }^{\circ}\text{C}$, suggesting that the 100 kDa filter removed all remaining ice nucleators present in the spent medium. This observation suggests that any macromolecules smaller than 100 kDa that were present in the spent medium are not ice nucleation active, because otherwise they ~~had~~would have passed the filter and led to an increased T_{50} when compared to the fresh medium. The ice-binding proteins present in and/or released from *F. cylindrus* are similar in size to the well characterized *fcIBP11*, which is about 26 kDa (Bayer-Giraldi et al., 2011). Thus, ice-binding proteins released by the *F. cylindrus* into the spent medium should have passed the filter and could have induced ice nucleation, if they had significant ice nucleation activity. However, the results shown in Fig. 85 do not reveal any ice nucleation activity ~~and, thus, can~~. This may be interpreted as follows. Either, ~~any~~ proteins remaining in the filtrate do not promote ice nucleation or, ~~alternatively,~~ *F. cylindrus* does not release any proteins into the spent medium. ~~In order to~~To shed

510

further light on the ice-nucleating ability of ice-binding proteins from *F. cylindrus*, we studied purified *fc*IBP11 samples in additional experiments. We studied the ice nucleation activity of two *fc*IBP11 solutions of different concentration as well as concentrations and that of the pure Tris-HCl buffer for comparison. The results are presented in Fig. 96. The two *fc*IBP11 samples with concentrations of 0.1 mM (dark blue circles) and 0.01 mM (light blue circles) concentrations reveal T_{50} values of -39.8 °C and -39.4 °C, which are equal to the $T_{50} = -39.7$ °C of the buffer reference (black circles) within experimental temperature uncertainty (± 0.3 °C). Thus, no significant shift in the freezing temperature is observed, and even when considering the increased ice nucleation temperature of the *fc*IBP11 at frozen fractions below about 25% it appears that *fc*IBP11 is not an efficient ice nucleator with relevance for atmospheric or biospheric processes, owing to its unnaturally high concentration in the droplet samples investigated here. These observations are in a good agreement with recent theoretical studies, which suggest that moderate 'antifreeze' IBPs show no nucleation of ice perpendicular to the basal and prismatic ice planes (Cui et al., 2022). And indeed, these basal and prismatic planes are exactly those planes, at which the moderate *fc*IBP11 binds to ice (Kondo et al., 2018).

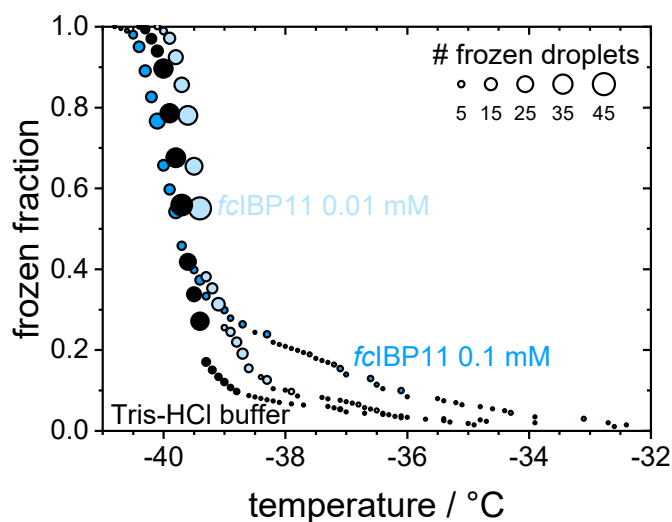


Figure 96: Cumulative frozen fractions as a function of temperature of droplets containing *fc*IBP11 solutions with concentrations of 0.1 mmol per L (dark blue) and 0.01 mmol per L (light blue). The black circles show the freezing of the Tris-HCl buffer for reference. The circle area indicates the number of droplets frozen at a particular temperature.

Overall, the results show that *F. cylindrus* diatom cells as well as cell fragments suspended in seawater can induce heterogeneous ice nucleation, while ice-binding proteins produced by *F. cylindrus* such as *fc*IBP11 have negligible ice nucleation activity.

4 Discussion and Implications

Here, we put the ~~above~~ results ~~obtained above~~ in the context of previous ice nucleation studies on diatoms. Triggered by the pioneering initial laboratory studies of marine diatom-induced ice nucleation (Alpert et al., 2011; Knopf et al., 2011) modelling studies have shown that in some regions of the atmosphere, marine diatoms may ~~indeed~~ contribute to the atmospheric INP (Burrows et al., 2013; Ickes et al., 2020). ~~In order to~~. To use laboratory ice nucleation data in such models, the data ~~have to~~ must be evaluated and parameterized appropriately. For example, a direct comparison of T_{50} or f_{ice} originating from different laboratory studies on different types of INPs it not meaningful, as different sample volumes, INP concentrations, buffer concentrations, etc. may have been used. Therefore, it is preferable to compare the cumulative number of ice nucleating active sites per mass, surface area or the number of the INPs. Here, we make a comparison based on total INP mass, using the following definition of the cumulative number of ice nucleating active sites per mass n_{m_total} (Murray et al., 2012; Hiranuma et al., 2015; Hiranuma et al., 2019; Xi et al., 2021).

$$n_{m_total} = \frac{-\ln(1-f_{ice})}{c_{m_total} \cdot V} \quad (81)$$

Here, V is the volume of an individual droplet in the experiment and c_{m_total} is the total mass of biological material per droplet. For the *F. cylindrus* samples investigated here, we used the total carbon mass per *F. cylindrus* cell from the literature (Kang and Fryxell, 1992) and ~~used elementary performed elemental~~ analysis to obtain the carbon content of our samples, resulting in a value of 39.32 % ~~to calculate the average total mass per individual *F. cylindrus* diatom cell of $m_{total} = 4.5 \times 10^{-11}$ g.~~ Using these values and our experimental data in Eq. (81), we have calculated the ice nucleating active sites n_{m_total} of the *F. cylindrus* diatoms, see the blue circles in Fig. 407. (We have fitted this data set and provide a corresponding parameterization, see Supplementary Fig. S6S9 and Eq. (S3S10).) Also shown in Fig. 407 are n_{m_total} data of other the sea ice diatoms *Melosira arctica* (blue squares) and *Nitzschia stellata* (blue triangles) and of the temperate diatom *Skeletonema marinoi* (open ~~red~~orange circles) from previous studies (Ickes et al., 2020; Xi et al., 2021).

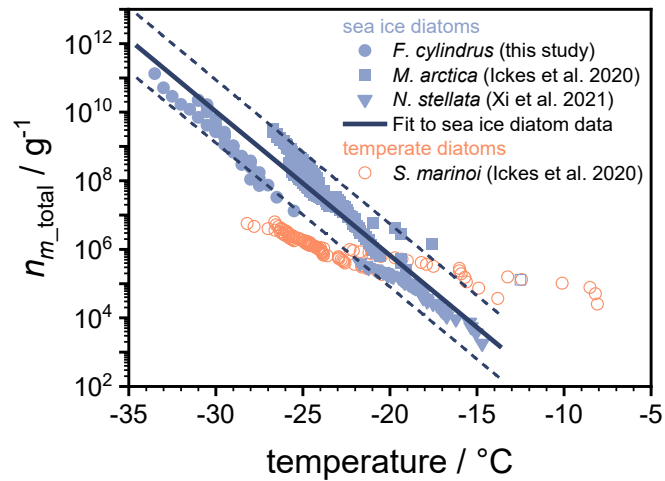


Figure 107: Experimental data of n_{m_total} , i.e. the number of ice active sites per total mass of *F. cylindrus* diatom cells (blue circles) and other sea ice diatoms (blue squares and triangles) from previous studies, as well as n_{m_total} data for one temperate diatom species (open red/orange circles) (Ickes et al., 2020; Xi et al., 2021). The solid line represents a fit of the n_{m_total} values for the three sea ice diatom species (see Eq. (92)), while the dashed lines indicate the 2σ upper and lower prediction bands of this fit. All temperatures were corrected for the freezing point depressions of different buffers and solutes, so that they represent the ice nucleation induced by the diatoms in pure water.

The n_{m_total} values for *N. stellata* were provided by the authors (Xi et al., 2021). For *M. arctica* and the *S. marinoi*, we calculated n_{m_total} from the total number of cells given in the original work and provided by the authors (Ickes et al., 2020), and assume cell volumes of $653 \mu\text{m}^3$ and $125 \mu\text{m}^3$ and a cell density of 1 mg mL^{-1} (Olenina et al., 2006; Xi et al., 2021). ~~In order to~~

To allow a direct comparison of ice nucleation of the different diatoms, which were studied in different types of aqueous solutions, all the ice nucleation temperatures shown in Fig. 107 have been corrected (either by the original authors or by us) for the colligative solute effect and represent diatom ice nucleation in pure water. We have corrected the freezing temperatures of the *F. cylindrus* samples by the measured difference between the T_{50} of pure double-distilled water and pure artificial seawater without any diatoms.

The comparison in Fig. 107 reveals that the curves of the three sea ice diatoms complement one another as n_{m_total} values of different magnitudes have been obtained over different temperature ranges. Interestingly, while there are some offsets exist between the different data sets, their slopes are quite similar. In contrast, the slope of the n_{m_total} data of the temperate diatom is significantly smaller. The observed similarities of the sea ice diatom data sets suggest a more generalized description of their behaviour in models. For this purpose, we fitted these data sets to provide a parametrization of n_{m_total} as a function of temperature. The three different data sets consist of different numbers of data points, which ~~was/were~~ taken into account in order to give each data set the same statistical weight. We further note that one strongly deviating data point from the *M. arctica* data set (indicated as an open square in Fig. 107) was excluded from the fitting procedure. The resulting parameterization is given as:

$$\log_{10}(n_{m_total} \text{ g}^{-1}) - \log_{10}(n_{m_total} \text{ g}^{-1}) = -0.420053 \text{ } ^\circ\text{C}^{-1} \cdot T - 2.57818 \quad (92)$$

where T is temperature to be entered in units of $^\circ\text{C}$. For numerical code verification, Eq. (92) should result in a value for n_{m_total} of $6.7 \times 10^5 \text{ g}^{-1}$ at a temperature of $-20.0 \text{ } ^\circ\text{C}$. This parametrization is valid over the temperature range between $-13.7 \text{ } ^\circ\text{C}$ to $-34.5 \text{ } ^\circ\text{C}$ (i.e., 259.45 to 238.65 K). The parameterization is shown as the thick solid line in Fig. 407, and the upper and lower 2σ prediction bands are given as dashed lines. In summary, Fig. 407 shows that the parameterization line and its prediction bands are an appropriate representation of the ice nucleation activity of three types of sea ice diatoms suitable for use in atmospheric or biogeosciences model applications.

5 Conclusions

F.
 In the following, we put the ice nucleation data of *F. cylindrus* and the other sea ice diatoms into context by comparing to field studies. Wilson et al. (2015) provided experimental evidence for a marine biogenic source of ice nucleating particles and suggested that exudates and fragments of diatoms as a source of the ice nucleating material located in the sea surface microlayer. Their low-temperature freezing data reveals a cumulative number of ice nucleating active sites per total organic carbon mass n_{m_TOC} of $\sim 1.3 \times 10^{10} \text{ g}^{-1}$ at $-27 \text{ } ^\circ\text{C}$ (calculated from the equation given in the caption of their Fig. 2), which is the low-temperature end of their data, and the most relevant to the present study. To compare this value to the n_{m_total} values given in Fig. 7, we estimated that the organic carbon content of their samples varies between 39.32% (representing the organic carbon content of *F. cylindrus* cells, see above) or 100% (representing a purely organic carbon composition), resulting in a range of n_{m_total} of $\sim 5.0 \times 10^9$ – $1.3 \times 10^{10} \text{ g}^{-1}$ for their Arctic sea surface microlayer samples. These are compared to n_{m_total} values of $8.2 \times 10^7 \text{ g}^{-1}$ (2σ prediction bands: 2.8×10^7 – $2.4 \times 10^8 \text{ g}^{-1}$) for *F. cylindrus* and of $5.8 \times 10^8 \text{ g}^{-1}$ (2σ prediction bands: 7.0×10^7 – $4.8 \times 10^9 \text{ g}^{-1}$) for sea ice diatoms, respectively, at $-27 \text{ } ^\circ\text{C}$, indicating that *F. cylindrus* and other sea ice diatoms may contribute to the marine INP in the Southern Oceans and Antarctic seawater, assuming the Wilson et al. parameterization applies also to these areas.

In another comparison, we use measurements of insoluble aerosol particles made at Amsterdam Island in the Southern Indian Ocean (Gaudichet et al., 1989). These measurements show that marine biogenic particles make up between 8 and 28% of the number of detected particles and that these were predominantly assigned to *Radiolaria* and diatom fragments (identified as amorphous silicates), with about 27 % or $2.7 \times 10^4 \text{ m}^{-3}$ particles observed in the southern winter (July) and fewer in fall (May, 8 %, $2.4 \times 10^4 \text{ m}^{-3}$) and spring (September, 7 %, $1.8 \times 10^3 \text{ m}^{-3}$). If we assume that all *Radiolaria* and diatom fragments can be

610 attributed to *F. cylindrus* diatoms, we can calculate the mass concentration of *F. cylindrus* diatom cells per cubic meter of air
from the mass per individual cell ($m_{\text{total}} = 4.5 \times 10^{-11}$ g, see above), yielding values of 1.2×10^{-6} g m⁻³ air (July), 1.1×10^{-6} g m⁻³
air (May), and 8.1×10^{-8} g m⁻³ air (September). Using the parametrization of the cumulative number of ice nucleating active
sites per mass *F. cylindrus* in Eq. (S10), we calculate a $n_{m_{\text{total}}}$ value of 8.2×10^7 g⁻¹ (2σ prediction bands: 2.8×10^7 - 2.4×10^8 g⁻¹)
615 at -27 °C, see above, from which we can derive the ~88 INP m⁻³ air (2σ : 3-250) at -27 °C in fall (May). This value can be
compared to *in situ* total INP measurements in the Southern Ocean south of Australia in fall (March-April) yielding values
between 34 and 207 INP m⁻³ air at -27 °C (McCluskey et al., 2018). Although the above calculations are order of magnitude
estimates, the comparison shows that it is not unreasonable that sea ice diatoms such as *F. cylindrus* and their fragments may
constitute a significant fraction of the INP in the Southern Ocean and Antarctic waters.

620 5 Summary and Conclusions

Cells and fragments of *F. cylindrus* diatoms can induce heterogeneous ice nucleation in artificial seawater by as much as up to
7.2°C higher than pure seawater temperature (for the highest diatom largest concentration investigated, (i.e., 5×10^7 cells per
mL)-) than the homogeneous ice nucleation temperature in pure seawater. We also observed an ice nucleating effect of
fragments smaller than 0.22 µm, in agreement with previous observations of the relevance of nanoscale biological fragments
625 for ice nucleation in clouds. We (O'Sullivan et al., 2015; Wilson et al., 2015; Irish et al., 2017; Irish et al., 2019; Hartmann et
al., 2021). For the ice-binding (antifreeze) protein *fcIBP11*, we did not observe any evidence for promoting ice nucleation at
low temperatures.

Using the information that *F. cylindrus* may serve as INPs, we can estimate their atmospheric relevance. Due to their smaller
630 size and, thus, longer atmospheric residence time, especially fragments of diatoms are expected to be relevant for atmospheric
ice nucleation because the atmospheric lifetime of entire *F. cylindrus* diatoms is estimated to be below one day due to
deposition (Hobbs, 2000; Seinfeld and Pandis, 2016). There are only a few studies that describe the aerosolization and
atmospheric transport processes of diatoms and diatom fragments as well as their atmospheric detection at different altitudes
(Brown et al., 1964; Gaudichet et al., 1989; Leck and Keith Bigg, 2008; Burrows et al., 2013). Based on order-of-magnitude
635 estimations comparing field observations of the Southern Oceans with our laboratory results, we suggest that diatoms like *F.*
cylindrus as well as their fragments may contribute to ice nucleation in marine environments of the polar regions at low
temperatures where sea ice diatoms become active for ice nucleation (Fig. 7). To improve these estimates, more observations
of the atmospheric abundance of diatoms and INPs in general and in the Antarctic marine environments are required and
modelling studies of the sea-to-air transfer of diatoms and their fragments are needed. In this respect, we observed a common
640 behaviour of the cumulative number of ice nucleating active sites per mass of diatom among three different types of sea ice
diatoms. This similarity may originate from a similar biological function of the ice nucleation ability in sea-ice diatoms, and

a corresponding parameterization developed thereof may simplify the representation of their properties in atmospheric biogeoscientific and biogeochemical models.

Data availability

645 The experimental data presented in this paper will be made freely available on a repository server of Bielefeld University upon final acceptance of the manuscript.

Author contribution

650 LE and TK designed the study. MBG provided the protein samples, LE performed the calibration and both the DSC and the microfluidic ice nucleation experiments, NR prepared the microfluidic devices. LE did the data analysis and the Poisson statistics calculations with input from TK. LE and TK prepared the figures, LE, TK and MBG wrote the manuscript with input from YR and NR. All authors contributed to the discussion of the data and text, and approved the final version of the manuscript.

Competing interests

The authors declare that they have no conflict of interest.

655 **Acknowledgements**

We thank Arika Allhusen and Klaus Valentin for providing the original *F. cylindrus* samples, and Luisa Ickes, Yu Xi and Allan Bertram for provision of original data sets on diatom ice nucleation and for helpful comments. We acknowledge support for the publication costs by the Open Access Publication Fund of Bielefeld University and the Deutsche Forschungsgemeinschaft (DFG).

660 **References**

- Ackley, S. F. and Sullivan, C. W.: Physical controls on the development and characteristics of Antarctic sea ice biological communities— a review and synthesis, *Deep Sea Research Part I: Oceanographic Research Papers*, 41, 1583–1604, doi:10.1016/0967-0637(94)90062-0, 1994.
- 665 Alpert, P. A., Aller, J. Y., and Knopf, D. A.: Ice nucleation from aqueous NaCl droplets with and without marine diatoms, *Atmos. Chem. Phys.*, 11, 5539–5555, doi:10.5194/acp-11-5539-2011, 2011.

- Aslam, S. N., Cresswell-Maynard, T., Thomas, D. N., and Underwood, G. J. C.: Production and Characterization of the Intra- and Extracellular Carbohydrates and Polymeric Substances (EPS) of Three Sea-Ice Diatom Species, and Evidence for a Cryoprotective Role for EPS, *Journal of phycology*, 48, 1494–1509, doi:10.1111/jpy.12004, 2012a.
- 670 Aslam, S. N., Strauss, J., Thomas, D. N., Mock, T., and Underwood, G. J. C.: Identifying metabolic pathways for production of extracellular polymeric substances by the diatom *Fragilariopsis cylindrus* inhabiting sea ice, *The ISME journal*, 12, 1237–1251, doi:10.1038/s41396-017-0039-z, 2018.
- Aslam, S. N., Underwood, G. J. C., Kaartokallio, H., Norman, L., Autio, R., Fischer, M., Kuosa, H., Dieckmann, G. S., and Thomas, D. N.: Dissolved extracellular polymeric substances (dEPS) dynamics and bacterial growth during sea ice formation in an ice tank study, *Polar Biol*, 35, 661–676, doi:10.1007/s00300-011-1112-0, 2012b.
- 675 Augustin, S., Wex, H., Niedermeier, D., Pummer, B., Grothe, H., Hartmann, S., Tomsche, L., Clauss, T., Voigtländer, J., Ignatius, K., and Stratmann, F.: Immersion freezing of birch pollen washing water, *Atmos. Chem. Phys.*, 13, 10989–11003, doi:10.5194/acp-13-10989-2013, 2013.
- Bar Dolev, M., Braslavsky, I., and Davies, P. L.: Ice-Binding Proteins and Their Function, *Annual review of biochemistry*, 85, 515–542, doi:10.1146/annurev-biochem-060815-014546, 2016.
- 680 Bartsch, A.: Sea Ice Algae of the Weddell Sea (Antarctica): Species Composition, Biomass, and Ecophysiology of Selected Species, *Ber. Polarforsch.*, 63, 1989.
- Bayer-Giraldi, M., Sazaki, G., Nagashima, K., Kipfstuhl, S., Vorontsov, D. A., and Furukawa, Y.: Growth suppression of ice crystal basal face in the presence of a moderate ice-binding protein does not confer hyperactivity, *Proceedings of the National Academy of Sciences of the United States of America*, 115, 7479–7484, doi:10.1073/pnas.1807461115, 2018.
- 685 Bayer-Giraldi, M., Uhlig, C., John, U., Mock, T., and Valentin, K.: Antifreeze proteins in polar sea ice diatoms: diversity and gene expression in the genus *Fragilariopsis*, *Environmental microbiology*, 12, 1041–1052, doi:10.1111/j.1462-2920.2009.02149.x, 2010.
- Bayer-Giraldi, M., Weikusat, I., Besir, H., and Dieckmann, G.: Characterization of an antifreeze protein from the polar diatom *Fragilariopsis cylindrus* and its relevance in sea ice, *Cryobiology*, 63, 210–219, doi:10.1016/j.cryobiol.2011.08.006, 2011.
- 690 Brown, R. M., Larson, D. A., and Bold, H. C.: Airborne Algae: Their Abundance and Heterogeneity, *Science (New York, N.Y.)*, 143, 583–585, doi:10.1126/science.143.3606.583, 1964.
- Budke, C. and Koop, T.: BINARY: an optical freezing array for assessing temperature and time dependence of heterogeneous ice nucleation, *Atmos. Meas. Tech.*, 8, 689–703, doi:10.5194/amt-8-689-2015, 2015.
- 695 Burrows, S. M., Hoose, C., Pöschl, U., and Lawrence, M. G.: Ice nuclei in marine air: biogenic particles or dust?, *Atmos. Chem. Phys.*, 13, 245–267, doi:10.5194/acp-13-245-2013, 2013.
- Cefarelli, A. O., Ferrario, M. E., Almandoz, G. O., Atencio, A. G., Akselman, R., and Vernet, M.: Diversity of the diatom genus *Fragilariopsis* in the Argentine Sea and Antarctic waters: morphology, distribution and abundance, *Polar Biol*, 33, 1463–1484, doi:10.1007/s00300-010-0794-z, 2010.

- 700 Collins, D. J., Neild, A., deMello, A., Liu, A.-Q., and Ai, Y.: The Poisson distribution and beyond: methods for microfluidic droplet production and single cell encapsulation, *Lab on a chip*, 15, 3439–3459, doi:10.1039/c5lc00614g, 2015.
- Creamean, J. M., Cenicerros, J. E., Newman, L., Pace, A. D., Hill, T. C. J., DeMott, P. J., and Rhodes, M. E.: Evaluating the potential for Haloarchaea to serve as ice nucleating particles, *Biogeosciences*, 18, 3751–3762, doi:10.5194/bg-18-3751-2021, 2021.
- 705 Creamean, J. M., Hill, T. C. J., DeMott, P. J., Uetake, J., Kreidenweis, S., and Douglas, T. A.: Thawing permafrost: an overlooked source of seeds for Arctic cloud formation, *Environ. Res. Lett.*, 15, 84022, doi:10.1088/1748-9326/ab87d3, 2020.
- Cui, S., Zhang, W., Shao, X., and Cai, W.: Do Antifreeze Proteins Generally Possess the Potential to Promote Ice Growth?, *Phys. Chem. Chem. Phys.*, doi:10.1039/D1CP05431G, 2022.
- 710 Davies, P. L.: Ice-binding proteins: a remarkable diversity of structures for stopping and starting ice growth, *Trends in Biochemical Sciences*, 39, 548–555, doi:10.1016/j.tibs.2014.09.005, 2014.
- DeMott, P. J., Hill, T. C. J., McCluskey, C. S., Prather, K. A., Collins, D. B., Sullivan, R. C., Ruppel, M. J., Mason, R. H., Irish, V. E., Lee, T., Hwang, C. Y., Rhee, T. S., Snider, J. R., McMeeking, G. R., Dhaniyala, S., Lewis, E. R., Wentzell, J. J. B., Abbatt, J., Lee, C., Sultana, C. M., Ault, A. P., Axson, J. L., Diaz Martinez, M., Venero, I., Santos-Figueroa, G.,
715 Stokes, M. D., Deane, G. B., Mayol-Bracero, O. L., Grassian, V. H., Bertram, T. H., Bertram, A. K., Moffett, B. F., and Franc, G. D.: Sea spray aerosol as a unique source of ice nucleating particles, *Proceedings of the National Academy of Sciences of the United States of America*, 113, 5797–5803, doi:10.1073/pnas.1514034112, 2016.
- DeMott, P. J., Möhler, O., Cziczo, D. J., Hiranuma, N., Petters, M. D., Petters, S. S., Belosi, F., Bingemer, H. G., Brooks, S. D., Budke, C., Burkert-Kohn, M., Collier, K. N., Danielczok, A., Eppers, O., Felgitsch, L., Garimella, S., Grothe, H.,
720 Herenz, P., Hill, T. C. J., Höhler, K., Kanji, Z. A., Kiselev, A., Koop, T., Kristensen, T. B., Krüger, K., Kulkarni, G., Levin, E. J. T., Murray, B. J., Nicosia, A., O'Sullivan, D., Peckhaus, A., Polen, M. J., Price, H. C., Reicher, N., Rothenberg, D. A., Rudich, Y., Santachiara, G., Schiebel, T., Schrod, J., Seifried, T. M., Stratmann, F., Sullivan, R. C., Suski, K. J., Szakáll, M., Taylor, H. P., Ullrich, R., Vergara-Temprado, J., Wagner, R., Whale, T. F., Weber, D., Welti, A., Wilson, T. W., Wolf, M. J., and Zenker, J.: The Fifth International Workshop on Ice Nucleation phase 2 (FIN-02):
725 laboratory intercomparison of ice nucleation measurements, *Atmos. Meas. Tech.*, 11, 6231–6257, doi:10.5194/amt-11-6231-2018, 2018.
- Dreischmeier, K., Budke, C., Wiehemeier, L., Kottke, T., and Koop, T.: Boreal pollen contain ice-nucleating as well as ice-binding 'antifreeze' polysaccharides, *Scientific reports*, 7, 41890, doi:10.1038/srep41890, 2017.
- Edd, J. F., Humphry, K. J., Irimia, D., Weitz, D. A., and Toner, M.: Nucleation and solidification in static arrays of
730 monodisperse drops, *Lab on a chip*, 9, 1859–1865, doi:10.1039/b821785h, 2009.
- Eicken, H.: The role of sea ice in structuring Antarctic ecosystems, *Polar Biol*, 12, doi:10.1007/BF00239960, 1992.

- Eickhoff, L., Dreischmeier, K., Zipori, A., Sirotinskaya, V., Adar, C., Reicher, N., Braslavsky, I., Rudich, Y., and Koop, T.: Contrasting Behavior of Antifreeze Proteins: Ice Growth Inhibitors and Ice Nucleation Promoters, *The journal of physical chemistry letters*, 10, 966–972, doi:10.1021/acs.jpcelett.8b03719, 2019.
- 735 Ekman, A. M. and Schmale, J.: Aerosol processes in high-latitude environments and the effects on climate, in: *Aerosols and climate*, Carslaw, K. S. (Ed.), Elsevier, Amsterdam, Kidlington, Cambridge, MA, 651–706, 2022.
- Garrison, D. and Buck, K.: The biota of Antarctic pack ice in the Weddell sea and Antarctic Peninsula regions, *Polar Biol*, 10, doi:10.1007/BF00238497, 1989.
- Gaudichet, A., Lefèvre, R., Gaudry, A., Ardouin, B., Lambert, G., and Miller, J.: Mineralogical composition of aerosols at
740 Amsterdam Island, *Tellus B*, 41B, 344–352, doi:10.1111/j.1600-0889.1989.tb00313.x, 1989.
- Gersonde, R. and Zielinski, U.: The reconstruction of late Quaternary Antarctic sea-ice distribution—the use of diatoms as a proxy for sea-ice, *Palaeogeography, Palaeoclimatology, Palaeoecology*, 162, 263–286, doi:10.1016/S0031-0182(00)00131-0, 2000.
- Govindarajan, A. G. and Lindow, S. E.: Size of Bacterial Ice-Nucleation Sites Measured in situ by Radiation Inactivation
745 Analysis, *Proceedings of the National Academy of Sciences of the United States of America*, 85, 1334–1338, 1988.
- Guillard, R. R. L. and Ryther, J. H.: Studies of marine planktonic diatoms: I. *Cyclotella nana* hustedt, and *Detonula* *convexa* (Cleve) Gran, *Can. J. Microbiol.*, 8, 229–239, doi:10.1139/m62-029, 1962.
- Günther, S. and Dieckmann, G. S.: Vertical zonation and community transition of sea-ice diatoms in fast ice and platelet layer, Weddell Sea, Antarctica, *Ann. Glaciol.*, 33, 287–296, doi:10.3189/172756401781818590, 2001.
- 750 Guo, S., Stevens, C. A., Vance, T. D. R., Olijve, L. L. C., Graham, L. A., Campbell, R. L., Yazdi, S. R., Escobedo, C., Bardolev, M., Yashunsky, V., Braslavsky, I., Langelaan, D. N., Smith, S. P., Allingham, J. S., Voets, I. K., and Davies, P. L.: Structure of a 1.5-MDa adhesin that binds its Antarctic bacterium to diatoms and ice, *Science advances*, 3, e1701440, doi:10.1126/sciadv.1701440, 2017.
- Hartmann, M., Gong, X., Kecorius, S., van Pinxteren, M., Vogl, T., Welti, A., Wex, H., Zeppenfeld, S., Herrmann, H.,
755 Wiedensohler, A., and Stratmann, F.: Terrestrial or marine – indications towards the origin of ice-nucleating particles during melt season in the European Arctic up to 83.7° N, *Atmos. Chem. Phys.*, 21, 11613–11636, doi:10.5194/acp-21-11613-2021, 2021.
- Herbert, R. J., Murray, B. J., Whale, T. F., Dobbie, S. J., and Atkinson, J. D.: Representing time-dependent freezing behaviour in immersion mode ice nucleation, *Atmos. Chem. Phys.*, 14, 8501–8520, doi:10.5194/acp-14-8501-2014,
760 2014.
- Hiranuma, N., Adachi, K., Bell, D. M., Belosi, F., Beydoun, H., Bhaduri, B., Bingemer, H., Budke, C., Clemen, H.-C., Conen, F., Cory, K. M., Curtius, J., DeMott, P. J., Eppers, O., Grawe, S., Hartmann, S., Hoffmann, N., Höhler, K., Jantsch, E., Kiselev, A., Koop, T., Kulkarni, G., Mayer, A., Murakami, M., Murray, B. J., Nicosia, A., Petters, M. D., Piazza, M., Polen, M., Reicher, N., Rudich, Y., Saito, A., Santachiara, G., Schiebel, T., Schill, G. P., Schneider, J.,
765 Segev, L., Stopelli, E., Sullivan, R. C., Suski, K., Szakáll, M., Tajiri, T., Taylor, H., Tobo, Y., Ullrich, R., Weber, D.,

- Wex, H., Whale, T. F., Whiteside, C. L., Yamashita, K., Zelenyuk, A., and Möhler, O.: A comprehensive characterization of ice nucleation by three different types of cellulose particles immersed in water, *Atmos. Chem. Phys.*, 19, 4823–4849, doi:10.5194/acp-19-4823-2019, 2019.
- 770 Hiranuma, N., Augustin-Bauditz, S., Bingemer, H., Budke, C., Curtius, J., Danielczok, A., Diehl, K., Dreischmeier, K., Ebert, M., Frank, F., Hoffmann, N., Kandler, K., Kiselev, A., Koop, T., Leisner, T., Möhler, O., Nillius, B., Peckhaus, A., Rose, D., Weinbruch, S., Wex, H., Boose, Y., DeMott, P. J., Hader, J. D., Hill, T. C. J., Kanji, Z. A., Kulkarni, G., Levin, E. J. T., McCluskey, C. S., Murakami, M., Murray, B. J., Niedermeier, D., Petters, M. D., O'Sullivan, D., Saito, A., Schill, G. P., Tajiri, T., Tolbert, M. A., Welti, A., Whale, T. F., Wright, T. P., and Yamashita, K.: A comprehensive laboratory study on the immersion freezing behavior of illite NX particles: a comparison of 17 ice nucleation measurement techniques, *Atmos. Chem. Phys.*, 15, 2489–2518, doi:10.5194/acp-15-2489-2015, 2015.
- 775 Hobbs, P. V. I.: Introduction to atmospheric chemistry a companion text to Basic physical chemistry for the atmospheric sciences: A companion text to Basic physical chemistry for the atmospheric sciences, Cambridge University Press, Cambridge, 262 pp., 2000.
- Hudait, A., Qiu, Y., Odendahl, N., and Molinero, V.: Hydrogen-Bonding and Hydrophobic Groups Contribute Equally to the Binding of Hyperactive Antifreeze and Ice-Nucleating Proteins to Ice, *Journal of the American Chemical Society*, 141, 7887–7898, doi:10.1021/jacs.9b02248, 2019.
- 780 Huebner, A., Srisa-Art, M., Holt, D., Abell, C., Hollfelder, F., deMello, A. J., and Edel, J. B.: Quantitative detection of protein expression in single cells using droplet microfluidics, *Chemical communications (Cambridge, England)*, 1218–1220, doi:10.1039/b618570c, 2007.
- 785 Ickes, L., Porter, G. C. E., Wagner, R., Adams, M. P., Bierbauer, S., Bertram, A. K., Bilde, M., Christiansen, S., Ekman, A. M. L., Gorokhova, E., Höhler, K., Kiselev, A. A., Leck, C., Möhler, O., Murray, B. J., Schiebel, T., Ullrich, R., and Salter, M. E.: The ice-nucleating activity of Arctic sea surface microlayer samples and marine algal cultures, *Atmos. Chem. Phys.*, 20, 11089–11117, doi:10.5194/acp-20-11089-2020, 2020.
- 790 Irish, V. E., Elizondo, P., Chen, J., Chou, C., Charette, J., Lizotte, M., Ladino, L. A., Wilson, T. W., Gosselin, M., Murray, B. J., Polishchuk, E., Abbatt, J. P. D., Miller, L. A., and Bertram, A. K.: Ice-nucleating particles in Canadian Arctic sea-surface microlayer and bulk seawater, *Atmos. Chem. Phys.*, 17, 10583–10595, doi:10.5194/acp-17-10583-2017, 2017.
- Irish, V. E., Hanna, S. J., Xi, Y., Boyer, M., Polishchuk, E., Ahmed, M., Chen, J., Abbatt, J. P. D., Gosselin, M., Chang, R., Miller, L. A., and Bertram, A. K.: Revisiting properties and concentrations of ice-nucleating particles in the sea surface microlayer and bulk seawater in the Canadian Arctic during summer, *Atmos. Chem. Phys.*, 19, 7775–7787, doi:10.5194/acp-19-7775-2019, 2019.
- 795 Kang, S.-H. and Fryxell, G.: *Fragilariopsis cylindrus* (Grunow) Krieger: The most abundant diatom in water column assemblages of Antarctic marginal ice-edge zones, *Polar Biol*, 12, doi:10.1007/BF00236984, 1992.
- Knopf, D. A., Alpert, P. A., Wang, B., and Aller, J. Y.: Stimulation of ice nucleation by marine diatoms, *Nature Geosci*, 4, 88–90, doi:10.1038/ngeo1037, 2011.

- 800 Kondo, H., Mochizuki, K., and Bayer-Giraldi, M.: Multiple binding modes of a moderate ice-binding protein from a polar microalga, *Physical chemistry chemical physics PCCP*, 20, 25295–25303, doi:10.1039/c8cp04727h, 2018.
- Koop, T.: Homogeneous Ice Nucleation in Water and Aqueous Solutions, *Zeitschrift für Physikalische Chemie*, 218, 1231–1258, doi:10.1524/zpch.218.11.1231.50812, 2004.
- Köster, S., Angilè, F. E., Duan, H., Agresti, J. J., Wintner, A., Schmitz, C., Rowat, A. C., Merten, C. A., Pisignano, D.,
805 Griffiths, A. D., and Weitz, D. A.: Drop-based microfluidic devices for encapsulation of single cells, *Lab on a chip*, 8, 1110–1115, doi:10.1039/b802941e, 2008.
- Krembs, C., Eicken, H., and Deming, J. W.: Exopolymer alteration of physical properties of sea ice and implications for ice habitability and biogeochemistry in a warmer Arctic, *Proceedings of the National Academy of Sciences of the United States of America*, 108, 3653–3658, doi:10.1073/pnas.1100701108, 2011.
- 810 Krembs, C., Eicken, H., Junge, K., and Deming, J.: High concentrations of exopolymeric substances in Arctic winter sea ice: implications for the polar ocean carbon cycle and cryoprotection of diatoms, *Deep Sea Research Part I: Oceanographic Research Papers*, 49, 2163–2181, doi:10.1016/S0967-0637(02)00122-X, 2002.
- Krembs, C. and Engel, A.: Abundance and variability of microorganisms and transparent exopolymer particles across the ice-water interface of melting first-year sea ice in the Laptev Sea (Arctic), *Marine Biology*, 138, 173–185,
815 doi:10.1007/s002270000396, 2001.
- Kunert, A. T., Pöhlker, M. L., Tang, K., Krevert, C. S., Wieder, C., Speth, K. R., Hanson, L. E., Morris, C. E., Schmale III, D. G., Pöschl, U., and Fröhlich-Nowoisky, J.: Macromolecular fungal ice nuclei in *Fusarium*: effects of physical and chemical processing, *Biogeosciences*, 16, 4647–4659, doi:10.5194/bg-16-4647-2019, 2019.
- Leck, C. and Bigg, E. K.: Biogenic particles in the surface microlayer and overlaying atmosphere in the central Arctic Ocean during summer, *Tellus B: Chemical and Physical Meteorology*, 57, 305–316, doi:10.3402/tellusb.v57i4.16546, 2005.
- 820 Leck, C. and Keith Bigg, E.: Comparison of sources and nature of the tropical aerosol with the summer high Arctic aerosol, *Tellus B*, 60, 118–126, doi:10.1111/j.1600-0889.2007.00315.x, 2008.
- Lizotte, M. P.: The Contributions of Sea Ice Algae to Antarctic Marine Primary Production, *Am Zool*, 41, 57–73, doi:10.1093/icb/41.1.57, 2001.
- 825 Lundholm, N. and Hasle, G. R.: Are *Fragilariopsis cylindrus* and *Fragilariopsis nana* bipolar diatoms? - Morphological and molecular analyses of two sympatric species, *Nova Hedwigia, Beiheft*, 133, 231–250, 2008.
- McCluskey, C. S., Hill, T. C. J., Humphries, R. S., Rauker, A. M., Moreau, S., Strutton, P. G., Chambers, S. D., Williams, A. G., McRobert, I., Ward, J., Keywood, M. D., Harnwell, J., Ponsonby, W., Loh, Z. M., Krummel, P. B., Protat, A., Kreidenweis, S. M., and DeMott, P. J.: Observations of Ice Nucleating Particles Over Southern Ocean Waters, *Geophys. Res. Lett.*, 45, 11,989–11,997, doi:10.1029/2018GL079981, 2018.
- 830 Mock, T., Otilar, R. P., Strauss, J., McMullan, M., Paaajanen, P., Schmutz, J., Salamov, A., Sanges, R., Toseland, A., Ward, B. J., Allen, A. E., Dupont, C. L., Frickenhaus, S., Maumus, F., Veluchamy, A., Wu, T., Barry, K. W., Falcioratore, A., Ferrante, M. I., Fortunato, A. E., Glöckner, G., Gruber, A., Hipkin, R., Janech, M. G., Kroth, P. G., Leese, F., Lindquist,

- 835 E. A., Lyon, B. R., Martin, J., Mayer, C., Parker, M., Quesneville, H., Raymond, J. A., Uhlig, C., Valas, R. E., Valentin, K. U., Worden, A. Z., Armbrust, E. V., Clark, M. D., Bowler, C., Green, B. R., Moulton, V., van Oosterhout, C., and Grigoriev, I. V.: Evolutionary genomics of the cold-adapted diatom *Fragilariopsis cylindrus*, *Nature*, 541, 536–540, doi:10.1038/nature20803, 2017.
- Murray, B. J., O'Sullivan, D., Atkinson, J. D., and Webb, M. E.: Ice nucleation by particles immersed in supercooled cloud droplets, *Chemical Society reviews*, 41, 6519–6554, doi:10.1039/c2cs35200a, 2012.
- 840 Olenina, I., Hajdu, S., Edler, L., Andersson, A., Wasmund, N., Busch, S., Göbel, J., Gromisz, S., Huseby, S., Huttunen, M., Jaanus, A., Kokkonen, P., Ledaine, I., and Niemkiewicz, E.: Biovolumes and size-classes of phytoplankton in the Baltic Sea, *HELCOM Balt.Sea Environ. Proc.*, 106, 2006.
- O'Sullivan, D., Murray, B. J., Ross, J. F., Whale, T. F., Price, H. C., Atkinson, J. D., Umo, N. S., and Webb, M. E.: The relevance of nanoscale biological fragments for ice nucleation in clouds, *Scientific reports*, 5, 8082, doi:10.1038/srep08082, 2015.
- 845 Pinti, V., Marcolli, C., Zobrist, B., Hoyle, C. R., and Peter, T.: Ice nucleation efficiency of clay minerals in the immersion mode, *Atmos. Chem. Phys.*, 12, 5859–5878, doi:10.5194/acp-12-5859-2012, 2012.
- Poulin, M., Daugbjerg, N., Gradinger, R., Ilyash, L., Ratkova, T., and Quillfeldt, C. von: The pan-Arctic biodiversity of marine pelagic and sea-ice unicellular eukaryotes: a first-attempt assessment, *Mar Biodiv*, 41, 13–28, doi:10.1007/s12526-010-0058-8, 2011.
- 850 Pummer, B. G., Bauer, H., Bernardi, J., Bleicher, S., and Grothe, H.: Suspendable macromolecules are responsible for ice nucleation activity of birch and conifer pollen, *Atmos. Chem. Phys.*, 12, 2541–2550, doi:10.5194/acp-12-2541-2012, 2012.
- Pummer, B. G., Budke, C., Augustin-Bauditz, S., Niedermeier, D., Felgitsch, L., Kampf, C. J., Huber, R. G., Liedl, K. R., 855 Loerting, T., Moschen, T., Schauerperl, M., Tollinger, M., Morris, C. E., Wex, H., Grothe, H., Pöschl, U., Koop, T., and Fröhlich-Nowoisky, J.: Ice nucleation by water-soluble macromolecules, *Atmos. Chem. Phys.*, 15, 4077–4091, doi:10.5194/acp-15-4077-2015, 2015.
- Rasmussen, D. H. and MacKenzie, A. P.: Effect of Solute on Ice-Solution Interfacial Free Energy; Calculation from Measured Homogeneous Nucleation Temperatures, in: *Water Structure at the Water-Polymer Interface*, Springer, 860 Boston, MA, 126–145, 1972.
- Raymond, J. A., Sullivan, C. W., and DeVries, A. L.: Release of an ice-active substance by Antarctic sea ice diatoms, *Polar Biol*, 14, doi:10.1007/BF00240276, 1994.
- Reicher, N., Budke, C., Eickhoff, L., Raveh-Rubin, S., Kaplan-Ashiri, I., Koop, T., and Rudich, Y.: Size-dependent ice nucleation by airborne particles during dust events in the eastern Mediterranean, *Atmos. Chem. Phys.*, 19, 11143–11158, 865 doi:10.5194/acp-19-11143-2019, 2019.
- Reicher, N., Segev, L., and Rudich, Y.: The Weizmann Supercooled Droplets Observation on a Microarray (WISDOM) and application for ambient dust, *Atmos. Meas. Tech.*, 11, 233–248, doi:10.5194/amt-11-233-2018, 2018.

- Riechers, B., Wittbracht, F., Hütten, A., and Koop, T.: The homogeneous ice nucleation rate of water droplets produced in a microfluidic device and the role of temperature uncertainty, *Physical chemistry chemical physics PCCP*, 15, 5873–5887, doi:10.1039/c3cp42437e, 2013.
- 870
- Roy, P., Mael, L. E., Hill, T. C. J., Mehndiratta, L., Peiker, G., House, M. L., DeMott, P. J., Grassian, V. H., and Dutcher, C. S.: Ice Nucleating Activity and Residual Particle Morphology of Bulk Seawater and Sea Surface Microlayer, *ACS Earth Space Chem.*, 5, 1916–1928, doi:10.1021/acsearthspacechem.1c00175, 2021.
- Roy-Barman, M. and Jeandel, C.: *Marine Geochemistry*, Oxford University Press, 2016.
- 875
- Šantl-Temkiv, T., Lange, R., Beddows, D., Rauter, U., Pilgaard, S., Dall'Osto, M., Gunde-Cimerman, N., Massling, A., and Wex, H.: Biogenic Sources of Ice Nucleating Particles at the High Arctic Site Villum Research Station, *Environmental science & technology*, 53, 10580–10590, doi:10.1021/acs.est.9b00991, 2019.
- Šantl-Temkiv, T., Sikoparija, B., Maki, T., Carotenuto, F., Amato, P., Yao, M., Morris, C. E., Schnell, R., Jaenicke, R., Pöhlker, C., DeMott, P. J., Hill, T. C. J., and Huffman, J. A.: Bioaerosol field measurements: Challenges and perspectives in outdoor studies, *Aerosol Science and Technology*, 54, 520–546, doi:10.1080/02786826.2019.1676395, 2020.
- 880
- Seinfeld, J. H. and Pandis, S. N.: *Atmospheric chemistry and physics: From air pollution to climate change*, Third edition, Wiley, Hoboken, New Jersey, 1120 pp., 2016.
- Steinke, I., DeMott, P. J., Deane, G. B., Hill, T. C. J., Maltrud, M., Raman, A., and Burrows, S. M.: A numerical framework for simulating the atmospheric variability of supermicron marine biogenic ice nucleating particles, *Atmos. Chem. Phys.*, 22, 847–859, doi:10.5194/acp-22-847-2022, 2022.
- 885
- Stohl, A. and Sodemann, H.: Characteristics of atmospheric transport into the Antarctic troposphere, *J. Geophys. Res.*, 115, doi:10.1029/2009JD012536, 2010.
- Tarn, M. D., Sikora, S. N. F., Porter, G. C. E., Shim, J.-U., and Murray, B. J.: Homogeneous Freezing of Water Using Microfluidics, *Micromachines*, 12, doi:10.3390/mi12020223, 2021.
- 890
- van Leeuwe, M. A., Tedesco, L., Arrigo, K. R., Assmy, P., Campbell, K., Meiners, K. M., Rintala, J.-M., Selz, V., Thomas, D. N., and Stefels, J.: Microalgal community structure and primary production in Arctic and Antarctic sea ice: A synthesis, *Elementa: Science of the Anthropocene*, 6, doi:10.1525/elementa.267, 2018.
- Vance, T. D. R., Bayer-Giraldi, M., Davies, P. L., and Mangiagalli, M.: Ice-binding proteins and the 'domain of unknown function' 3494 family, *The FEBS journal*, 286, 855–873, doi:10.1111/febs.14764, 2019.
- 895
- Wagner, R., Ickes, L., Bertram, A. K., Els, N., Gorokhova, E., Möhler, O., Murray, B. J., Umo, N. S., and Salter, M. E.: Heterogeneous ice nucleation ability of aerosol particles generated from Arctic sea surface microlayer and surface seawater samples at cirrus temperatures, *Atmos. Chem. Phys.*, 21, 13903–13930, doi:10.5194/acp-21-13903-2021, 2021.
- Welti, A., Bigg, E. K., DeMott, P. J., Gong, X., Hartmann, M., Harvey, M., Henning, S., Herenz, P., Hill, T. C. J., Hornblow, B., Leck, C., Löffler, M., McCluskey, C. S., Rauker, A. M., Schmale, J., Tatzelt, C., van Pinxteren, M., and Stratmann,
- 900

- F.: Ship-based measurements of ice nuclei concentrations over the Arctic, Atlantic, Pacific and Southern oceans, *Atmos. Chem. Phys.*, 20, 15191–15206, doi:10.5194/acp-20-15191-2020, 2020.
- 905 Wex, H., Augustin-Bauditz, S., Boose, Y., Budke, C., Curtius, J., Diehl, K., Dreyer, A., Frank, F., Hartmann, S., Hiranuma, N., Jantsch, E., Kanji, Z. A., Kiselev, A., Koop, T., Möhler, O., Niedermeier, D., Nillius, B., Rösch, M., Rose, D., Schmidt, C., Steinke, I., and Stratmann, F.: Intercomparing different devices for the investigation of ice nucleating particles using Snomax ® as test substance, *Atmos. Chem. Phys.*, 15, 1463–1485, doi:10.5194/acp-15-1463-2015, 2015.
- 910 Wilson, T. W., Ladino, L. A., Alpert, P. A., Breckels, M. N., Brooks, I. M., Browse, J., Burrows, S. M., Carslaw, K. S., Huffman, J. A., Judd, C., Kilhau, W. P., Mason, R. H., McFiggans, G., Miller, L. A., Nájera, J. J., Polishchuk, E., Rae, S., Schiller, C. L., Si, M., Temprado, J. V., Whale, T. F., Wong, J. P. S., Wurl, O., Yakobi-Hancock, J. D., Abbatt, J. P. D., Aller, J. Y., Bertram, A. K., Knopf, D. A., and Murray, B. J.: A marine biogenic source of atmospheric ice-nucleating particles, *Nature*, 525, 234–238, doi:10.1038/nature14986, 2015.
- Wolber, P. K., Deininger, C. A., Southworth, M. W., Vandekerckhove, J., van Montagu, M., and Warren, G. J.: Identification and Purification of a Bacterial Ice-Nucleation Protein, *Proceedings of the National Academy of Sciences of the United States of America*, 83, 7256–7260, 1986.
- 915 Xi, Y., Mercier, A., Kuang, C., Yun, J., Christy, A., Melo, L., Maldonado, M. T., Raymond, J. A., and Bertram, A. K.: Concentrations and properties of ice nucleating substances in exudates from Antarctic sea-ice diatoms, *Environmental science. Processes & impacts*, 23, 323–334, doi:10.1039/d0em00398k, 2021.

Revised supplemental information with changes indicated

Ice nucleating properties of the sea ice diatom *Fragilariopsis cylindrus* and its exudates

Lukas Eickhoff¹, Maddalena Bayer-Giraldi², Naama Reicher³, Yinon Rudich³, Thomas Koop¹

¹Faculty of Chemistry, Bielefeld University, Universitätsstraße 25, 33615 Bielefeld, Germany

²Alfred-Wegener-Institut Helmholtz-Zentrum für Polar- und Meeresforschung (AWI), Bremerhaven, Germany

³Department of Earth and Planetary Sciences, Weizmann Institute of Science, 76100 Rehovot, Israel

Correspondence to: Thomas Koop (thomas.koop@uni-bielefeld.de)

Supplemental Information

Differential scanning calorimetry experiments

The principle preparation procedure for the water-in-oil emulsion (w/o) samples was almost identical to the method described earlier (Dreischmeier et al., 2017). 1 mL of 7 wt% emulsifier Span[®]65 (Merck) dissolved in 93 wt% of a mixture of 50 vol% methylcyclopentane (Acros Organics, 99 %) and 50 vol% methylcyclohexane (Acros Organics, 95 %) was used as the organic phase. The aqueous phase consisted of 1 mL of an *F. cylindrus* suspension with a concentration of 1×10^7 cells per mL, see Sect. 2.2.2 in the main paper, or alternatively of 1 mL of pure artificial seawater for comparison. The mixtures of the organic and aqueous phase were subsequently emulsified by stirring with a high-speed disperser (IKA Ultra-Turrax T25 basic) for 10 min at 20'000 rpm. For a DSC measurement, about 10 mg of such an emulsion was filled into an aluminium pan that was sealed hermetically and then transferred into the calorimeter. The samples were cooled at a rate of -5 °C per min down to -60 °C, and subsequently reheated, first at 5 °C per min and then, in the temperature range between -20 °C and +5 °C, at 1 °C per min.

WISDOM microfluidic device

For the droplet generation, we used two syringe pumps (neMESYS NEM-B101-02 E), one filled with the aqueous sample and another with an organic phase consisting of 2 wt% Span[®]80 (Merck) dissolved in 98 wt% of a mineral oil (Sigma-Aldrich, mineral oil M3516). The microfluidic chip was connected to the pumps with PTFE tubes. The droplets generated within the chip had diameters of $90 \mu\text{m} \pm 5 \mu\text{m}$.

For the freezing experiments, we placed the microfluidic chip after the droplet production on a temperature-controlled cold-stage (Linkam, BCS 196) attached to an optical microscope (Olympus, BX51 TRF). The temperature of the droplets in the chip was calibrated with respect to the cooling (or heating) rate as well as to the absolute temperature, and is described in detail in a previous study (Eickhoff et al., 2019). The freezing of the droplets was observed using the transmission mode of the microscope and we recorded the images with a digital camera (Q-Imaging, MicroPublisher 5.0 RTV) for later analysis by a LabView routine that detects a freezing event from the change in grey values of a particular droplet upon freezing. Typical changes in the droplets' grey values during freezing experiments are depicted in Fig. S3. In each individual experiment, between about 45 to 70 droplets were observed simultaneously, depending upon the percentage of droplet-filled microcells within the droplet array of the chip.

For all *F. cylindrus* measurements, the chip was first cooled to a temperature of -20 °C at a rate of -5 °C per min, because no freezing events were detected in this temperature range. After equilibration at this temperature for 2 min, the samples were then cooled at a slower rate of -1 °C per min to -45 °C, at which all droplets were frozen. Thereafter, the chip was heated relatively quickly at a rate of 5 °C per min, until -10 °C, and after two minutes of equilibration, it was then heated to 5 °C at 1 °C per min. The detailed temperature profiles for each type of experiment are listed in Table S2.

Evaluation procedure for samples with small INP concentrations

Referring to section 2.3.3 in the main paper, the following Poisson distribution can be used to describe the probability $P_\lambda(k)$ that an individual droplet contains exactly k INPs when the average concentration is λ INPs per droplet:

$$P_\lambda(k) = \frac{\lambda^k}{k!} \exp(-\lambda). \quad (S1)$$

Note that the derivation of the Poisson distribution contains a simplification that require a larger number of droplets and hence Eq. (S1) becomes more accurate as the number of investigated droplets increases. For the microfluidic experiments performed in this work with more than a hundred droplets investigated for each sample the simplification applies.

The average number of INPs per droplet, λ , is easily calculated from the concentration c of INPs in the stock solution and the volume V of an individual microfluidic droplet:

$$\lambda = c \cdot V_{\text{drop}}. \quad (S2)$$

Furthermore, the droplet volume V_{drop} can be expressed by the droplet's radius r or alternatively by its diameter d :

$$V_{\text{drop}} = \frac{4}{3}\pi \cdot r^3 = \frac{1}{6}\pi \cdot d^3. \quad (\text{S3})$$

Figure S4 shows the calculated Poisson distributions of the number of cells per droplet for four different values of λ in a concentration range relevant to this study. For lower values of λ , the histograms exhibit the tilted shape typical of Poisson distributions, while for larger values of λ , the Poisson distribution approaches the more symmetrical shape of a normal distribution (Koop et al., 1997).

For the ice nucleation experiments considered here, only those droplets containing at least one INP and those without any INPs are relevant, as this determines whether they are subject to heterogeneous or homogeneous nucleation, respectively. Whether a droplet contains one, two or more INPs is of less importance, as long as every INP is identical and, thus, induces heterogeneous ice nucleation at the same temperature. The probability that a droplet does not contain any INPs can be calculated easily by inserting $k = 0$ into Eq. (S1):

$$P_{\lambda}(0) = \frac{\lambda^0}{0!} \exp(-\lambda) = \frac{1}{1} \exp(-\lambda) = \exp(-\lambda). \quad (\text{S4})$$

$P_{\lambda}(0)$ is shown as the black-coloured bar in each panel of Fig. S4. The probability that a droplet contains at least one INP, $P_{\lambda}(k \geq 1)$, is given by the combined probability of all red-coloured bars in each panel of Fig. S4, and it can be calculated using the fact that the sum of all probabilities $P_{\lambda}(k)$ for k from 0 to ∞ must become 1 (see Eq. (S5)):

$$P_{\lambda}(k) = \sum_{k=0}^{\infty} \frac{\lambda^k}{k!} \exp(-\lambda) = 1. \quad (\text{S5})$$

Hence, $P_{\lambda}(k \geq 1)$ can be calculated from the following difference:

$$P_{\lambda}(k \geq 1) = \sum_{k=1}^{\infty} \frac{\lambda^k}{k!} \exp(-\lambda) = \sum_{k=0}^{\infty} \frac{\lambda^k}{k!} \exp(-\lambda) - \sum_{k=0}^0 \frac{\lambda^k}{k!} \exp(-\lambda) = 1 - P_{\lambda}(0) = 1 - \exp(-\lambda). \quad (\text{S6})$$

Since λ can be expressed by the product of the droplets' diameter and the known concentration of INPs in the stock solution, c , (see Eq. (S2) and Eq. (S3)) this yields:

$$P_{\lambda}(k \geq 1) = 1 - \exp\left(-\frac{\pi}{6} \cdot c \cdot d^3\right). \quad (S7)$$

The equations above have been derived for applications where the average concentration c of INPs in solution is known. However, in ice nucleation experiments of natural samples, the concentration c of INPs per volume is often unknown a priori and other values such as the organic carbon content has to be used for comparison (Gute and Abbatt, 2020; Xi et al., 2021). In such cases, Eq. (S7) can be used to obtain a rough estimate of λ and, thus, c from ice nucleation experiments when a plateau in the experimental frozen fraction curve is observed. The frozen fraction is defined as the number of frozen droplets relative to the number of all droplets, at a given temperature (Budke and Koop, 2015). Here, we term the value of the frozen fraction at the plateau as f'_{ice} . If a sufficiently large number of droplets is investigated, then f'_{ice} corresponds to the fraction of droplets that froze heterogeneously and thus may be equated with that fraction of droplets containing at least one INP.

There are two underlying assumptions for experimentally obtaining f'_{ice} . First, every droplet containing at least one INP freezes heterogeneously, which appears entirely reasonable. Secondly, every droplet containing one or more INPs freezes at a higher temperature than those droplets without any INP, i.e. the difference between the heterogeneous and homogeneous ice nucleation temperature is large enough to be easily distinguished in the experiment. With these two assumptions a plateau in the frozen fraction curve can be interpreted as follows: the fraction of droplets below the plateau froze heterogeneously and contain at least one INP, and the fraction of droplets above the plateau froze homogeneously (when their freezing temperature is consistent with homogenous freezing) and, thus, do not contain INPs. In practice, this evaluation procedure does not work if none of the droplets froze heterogeneously or if all droplets froze heterogeneously at the same temperature without any obvious plateau, i.e. it is only applicable for intermediate average INP concentrations in what we term the “Poisson relevant range”.

We define this “Poisson relevant range” as the range of average INP concentrations, in which both droplets without any INP as well as droplets containing one or more INPs occur and, thus, both can be observed readily in the corresponding freezing experiments. For the experiments presented here, we establish the Poisson relevant range as the area between $P_{\lambda}(k \geq 1)$ values of 5.0 % and 99.5 %. The lower limit was set at 5.0% in order to avoid any influence of the freezing of a minor fraction of droplets induced by impurities and the upper limit corresponds to about one out of 200 droplets not containing any INP and thus freezing homogeneously, while the highest accuracy can be reached for a value of 50 % (see blue curve in Fig. S5a and Fig. S6). For higher concentrations, when every droplet contains at least one INP, the above Poisson evaluation is not needed and the classic method can be used, and so this upper limit sets an endpoint for the Poisson-based evaluation. The classic method indeed assumes that every observed droplet contains at least one INP and it has been described in detail previously (Murray et al., 2012; Budke and Koop, 2015).

To demonstrate the concentration range suitable for the Poisson method, i.e. the Poisson relevant range, the latter is indicated in Fig. S5a as the grey shaded area. The solid blue curve shows the values of $P_\lambda(k \geq 1)$ calculated using Eq. (S7) as a function of the average INP concentration c of the studied sample and a droplet diameter of 90 μm . The two dashed lines show the changes for a deviation of $\pm 5 \mu\text{m}$ in droplet diameter.

To verify the procedure, we investigated aqueous suspensions of the well-studied ice-nucleating bacterium *Pseudomonas syringae* in the form of the commercial product Snomax (Morris et al., 2011; Budke and Koop, 2015; Wex et al., 2015). The ice nucleation temperatures of each about 165 ± 15 droplets, from three single measurements with 45 to 70 droplets each, containing either pure double-distilled water or three different concentrations of *P. syringae* were investigated, see Table S3. These concentrations are also marked in Fig. S5a as vertical lines. A similar plot for the *F. cylindrus* diatoms can be found in Fig. S6. The resulting experimental frozen fraction curves of *P. syringae* are shown in Fig. S5b. Double-distilled water (black open symbols) shows a steep increase in frozen fraction below about $-34.0 \text{ }^\circ\text{C}$, in agreement with homogeneous ice nucleation rates of droplets of such diameter (Koop and Murray, 2016; Reicher et al., 2018; Eickhoff et al., 2019). Following this observation, all droplets of the *P. syringae* samples that froze at around or below this temperature are assumed to have nucleated homogeneously, i.e. they are considered to contain no INPs in the analysis below.

For all *P. syringae* samples, the first freezing events occur at much higher temperatures of about -8 to $-9 \text{ }^\circ\text{C}$, and the frozen fraction curve in each case initially increases strongly before reaching a plateau, and subsequently the remaining liquid droplets freeze only at very low temperatures. In each sample, the plateau occurs at a different value of the frozen fraction, e.g. f'_{ice} is higher the larger the *P. syringae* concentrations (pink > blue > orange). We determined the corresponding f'_{ice} values, as defined above, from the experimentally obtained frozen fraction curve as the value of the frozen fraction at $-34.0 \text{ }^\circ\text{C}$, i.e. at the threshold between heterogeneous and homogenous ice nucleation as defined above. The resulting f'_{ice} values for the three concentrations were 0.99, 0.61, and 0.39, respectively, indicated as the dashed horizontal lines in Fig. S5b. These f'_{ice} values correspond to $P_\lambda(k \geq 1)_{\text{measured}}$ and can be used to infer the average INP concentration from Eq. (S7). Because in the current experiments the INP concentrations are known (i.e., 1.4×10^7 , 2.8×10^6 , and $1.4 \times 10^6 \text{ mL}^{-1}$), these experimentally derived f'_{ice} values can be compared to the expected f_{ice} values, corresponding to $P_\lambda(k \geq 1)_{\text{calculated}}$ values calculated from Eq. (S7), yielding values of 1.00 ± 0.01 , 0.66 ± 0.06 and 0.41 ± 0.05 , respectively. These theoretical values are in good agreement (within experimental uncertainty) with the measured values and thus confirm our approach and the inferred INP concentrations of 1.2×10^7 , 2.5×10^6 and $1.3 \times 10^6 \text{ mL}^{-1}$ (see Table S3) deviate by about 14%, 11% and 7% from the prepared concentrations, which is very good given that INP concentrations can vary by orders of magnitude. For further validation that the Poisson distribution is necessary for a proper evaluation in the above-mentioned concentration range, the cumulative number of active ice-nucleating sites n_N per number of *P. syringae* bacteria was evaluated and discussed in the following section and the related Fig. S7.

Determination of INP concentration

In the main paper Above, we have defined f'_{ice} as the plateau region separating heterogeneous and homogenous freezing. Since f'_{ice} varies with the number of droplets containing at least one INP, an experimentally determined f'_{ice} value can be used to calculate the concentration of INPs for unknown samples using a variation of Eq. (7S7). Typically, a sample is investigated by means of a dilution series so that a different INP concentration is scanned in each experiment. If the INP concentration is too large, all droplets freeze heterogeneously, and if it is too low, no INP-induced heterogeneous nucleation occurs (apart from that induced by any impurity present) and, thus, all droplets freeze homogeneously. In both these cases, it is not possible to obtain the desired INP concentration. But if measurements are done in the Poisson relevant concentration range (see definition in the main paper), one can observe both heterogeneous as well as homogenous freezing of droplets, resulting in a plateau in the frozen fraction curve, as discussed above. With the frozen fraction value of this plateau, f'_{ice} , and the assumptions that, first, every INP induces heterogeneous freezing and that, secondly, all heterogeneously frozen droplets freeze before the first freezing of a homogenous frozen droplet, the following Equation can be obtained by rearranging Eq. (7S7):

$$c = -\frac{6 \ln(1 - P_\lambda(k \geq 1))}{\pi \cdot d^3} = -\frac{6 \ln(1 - f'_{ice})}{\pi \cdot d^3} \quad (S4S8)$$

A comparison with Fig. 3a in the original paper S5a implies that f'_{ice} values of about 0.5 will lead to more accurate results than values close to the limits of the Poisson relevant range, because of the larger slope of the curve at $P_\lambda(k \geq 1)$ at intermediate f'_{ice} values.

In order to verify this method, we determine the concentrations c_{measured} of the investigated *P. syringae* samples by using the already determined values of $P_\lambda(k \geq 1)_{\text{measured}}$, from Table S3, for calculating f'_{ice} . The resulting values for c_{measured} as well as the actually prepared concentrations of the samples c , are also listed in Table S4. The comparison shows that there are only minor differences between the prepared and measured concentrations, supporting the fact that this method provides a suitable and relatively accurate estimate of the INP concentration of an unknown sample. A similar treatment was performed for the *F. cylindrus* diatom samples and the related Fig. S4S6.

Figure S5S7 shows the cumulative number of ice-nucleating active sites n_N per *P. syringae* cell. As we will see below, the Poisson evaluation is required for this type of evaluation. The cumulative number of active ice-nucleating sites is given formally by Eq. (S2)-S9) (Budke and Koop, 2015):

$$n_N = \frac{-\ln(1-f_{ice})}{c \cdot V} \quad (S2S9)$$

Here, f_{ice} is the frozen fraction, c is the INP concentration in absolute number of INPs per volume unit (i.e., the number density), and V is the volume of each individual droplet.

Figure S5S7 shows that for temperatures lower than about -35 °C, n_N obtains values that are larger than one per bacterium cell, implying that one bacterium initiates ice nucleation in more than one droplet, which is of course unreasonable. Instead, these high n_N values result from homogenous ice nucleation in droplets that do not contain any *P. syringae* bacteria, which normally is not considered in the classical n_N evaluation. By applying Eq. (7S7) on these measurements, the threshold value between droplets that do contain INPs and those that do not can be determined. This treatment results in maximum values for n_N of about one per bacterial cell, as indicated by the filled and open circles in Fig. S5S7.

Results of the DSC experiments

For the DSC measurements, an inverse emulsion of pure artificial seawater as a reference was compared with an emulsion of artificial seawater containing 1×10^7 *F. cylindrus* cells per mL, see Fig. S8. First, the endothermic ice melting signals of the reference and the sample in Fig. S8a show almost the same signal, indicating that any colligative effect of the diatoms is negligible when compared to the amount of the dissolved ions in the artificial seawater. This similarity in the ice melting signals also implies no change in water activity of the artificial seawater upon the addition of the diatoms and, thus, no colligative effect on the homogeneous ice nucleation (freezing) signals is to be expected

The exothermic freezing signals for both emulsions are shown in Fig. S8b. For the seawater reference, one distinct nearly symmetrical freezing signal is revealed with a maximum at about -44 °C and an onset, which is defined as the freezing temperature of the sample, at about -40 °C. In contrast, the *F. cylindrus* sample shows the same maximum, but in addition a second exothermic signal in the form of a shoulder at about -42 °C, with an onset at a somewhat higher temperature of -39 °C when compared to the reference, and with small signals as high as -34 °C. Because of the colligative freezing point depression of the seawater, the freezing temperatures of the reference and the sample are shifted to lower temperatures, compared to pure water.

The larger broad signal in both emulsion samples corresponds to the homogeneous ice nucleation temperature of artificial seawater. This signal is also observed in the *F. cylindrus* sample because many of the emulsion droplets in that sample do not contain diatoms. The exothermic shoulder of the signal, which is not present in the reference, is most likely due to the freezing

of droplets containing a diatom cell or fragment, and the shift of the onset to higher temperature is a first indication for the heterogeneous ice nucleation activity of the diatoms.

Because of the fact that the diatoms are of similar size as the emulsion droplets and the potential of mechanical disruption of diatom cells during the fast stirring of the disperser during emulsion preparation, these emulsion experiments appear to us as not suitable for a quantitative analysis of the ice nucleation activity of *F. cylindrus*. Thus, we employed non-invasive methods in the experiments described below.

Parametrization of *F. cylindrus* ice nucleation efficiency

In Eq. (2) in the main paper, we provide a parametrization representing the ice nucleation of the different sea ice diatoms in shown in Fig. 7 of the main paper in terms of the number of ice active sites per total mass of diatom cells, n_{m_total} . We also derived a parametrization for the individual ice nucleation efficiency of the *F. cylindrus* diatoms (see Fig. S9), which is given in the following Eq. (S10):

$$\log_{10}(n_{m_total} \text{ g}^{-1}) = -0.521789^{\circ\text{C}^{-1}} \cdot T - 6.1761, \quad (\text{S10})$$

where T is temperature to be entered in units of °C. For numerical code verification, Eq. (S10) should result in a value for n_{m_total} of $8.2 \times 10^7 \text{ g}^{-1}$ at a temperature of -27.0 °C. This parametrization is valid over the temperature range between -24.5 °C to -34.5 °C.

Table S1: Salts used for the preparation of artificial seawater for the *F. cylindrus* ice nucleation experiments. The amounts of substances provided for each ion yield a mass of 500 g artificial seawater at a salinity of 34.5.

Salt	Supplier	<i>m</i> [g]	Na ⁺ [mmol]	K ⁺ [mmol]	Mg ²⁺ [mmol]	Ca ²⁺ [mmol]	Cl ⁻ [mmol]	SO ₄ ²⁻ [mmol]	H ₂ O [mmol]
NaCl	VWR Chemicals	11.8446	202.68				202.68		
KCl	VWR Chemicals	0.3758		5.04			5.04		
MgCl ₂ · 6H ₂ O	ITW Reagents	5.3280			26.21		52.42		157.25
Na ₂ SO ₄ · 10H ₂ O	Acros Organics	4.4902	27.87					13.94	139.36
CaCl ₂ · 2H ₂ O	ITW Reagents	0.7460				5.07	10.15		10.15
H ₂ O	double distilled water	477.23							26490.26
artificial seawater		500.01	230.55	5.04	26.21	5.07	270.28	13.94	26797.02

Table S2: Temperature parameters used in the microfluidic freezing experiments. The first number in each triplet is the final temperature of the respective step in °C, the second number indicates the rate of cooling or heating in °C per min, and the third number indicates the holding time at the final temperature in min. Reference samples were always investigated with the same parameters as those given for each sample.

Step	<i>F. cylindrus</i>	<i>F. cylindrus</i> (filtered)	<i>F. cylindrus</i> (pure cells)	<i>F. cylindrus</i> (Medium)	<i>fcIBP11</i>	<i>P. syringae</i>
1	-20/-5/2	-20/-5/2	-20/-5/2	-20/-5/2	-20/-5/2	-5/-5/2
2	-45/-1/0	-45/-1/0	-45/-1/0	-45/-1/0	-45/-1/0	-40/-1/0
3	-10/5/2	-10/5/2	-10/5/2	-10/5/2	-10/5/2	-10/5/2
4	5/1/0	5/1/0	5/1/0	5/1/0	5/1/0	5/1/0

Table S3: As prepared concentrations c of the *P. syringae* samples, calculated fractions of droplets with at least one bacterium $P_\lambda(k \geq 1)_{\text{calculated}}$, as well as measured fractions $P_\lambda(k \geq 1)_{\text{measured}}$ and experimentally determined concentrations c_{measured} based on the approach outlined above using Eq. (S+S8).

c / mL^{-1}	$P_\lambda(k \geq 1)_{\text{calculated}}$	$P_\lambda(k \geq 1)_{\text{measured}}$	$c_{\text{measured}} / \text{mL}^{-1}$
1.4×10^7	$1.00^{+0.00}_{-0.01}$	0.99	1.2×10^7
2.8×10^6	$0.66^{+0.06}_{-0.06}$	0.61	2.5×10^6
1.4×10^6	$0.41^{+0.05}_{-0.05}$	0.39	1.3×10^6

Table S4: Shifts in ice nucleation temperature relative to the ΔT_{50} of artificial seawater for the untreated *F. cylindrus* samples, as well as for the samples filtered with a 0.22 μm syringe filter.

c	unfiltered ΔT_{50}	filtered ΔT_{50}
$5 \times 10^7 \text{ mL}^{-1}$	7.2 °C	6.4 °C
$1 \times 10^7 \text{ mL}^{-1}$	6.0 °C	5.2 °C
$2 \times 10^6 \text{ mL}^{-1}$	5.4 °C	3.1 °C
$1 \times 10^6 \text{ mL}^{-1}$	4.8 °C	2.6 °C
$5 \times 10^5 \text{ mL}^{-1}$	2.8 °C	0.0 °C

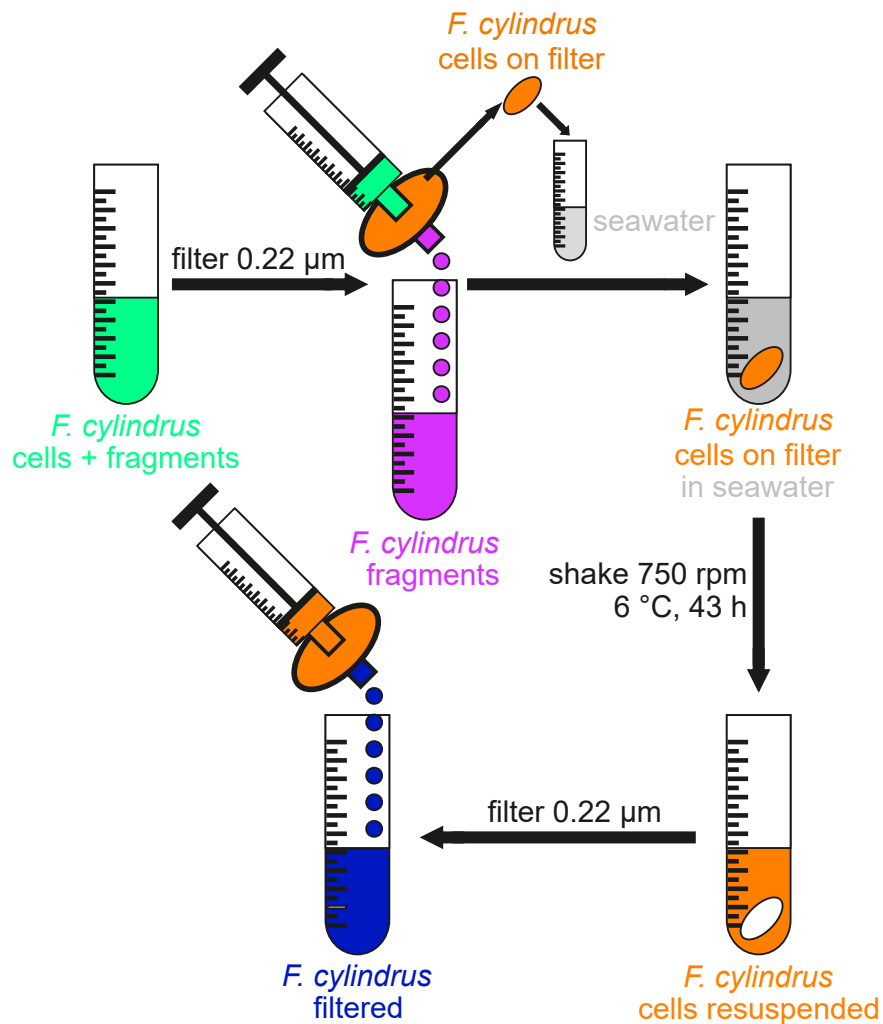


Figure S1: Extraction of the pure *F. cylindrus* cells by filtration of the stock solution (green). After filtration, the filtrate (purple) should only contain smaller cell fragments and soluble molecules such as *fcIBP*, while whole cells and larger fragments remain on the filter (orange filter). By shaking the filter in artificial seawater (grey), the cells were resuspended (orange solution). As a finally test, filtration of this suspension (blue) should not show any ice nucleation results different from those of pure artificial seawater.

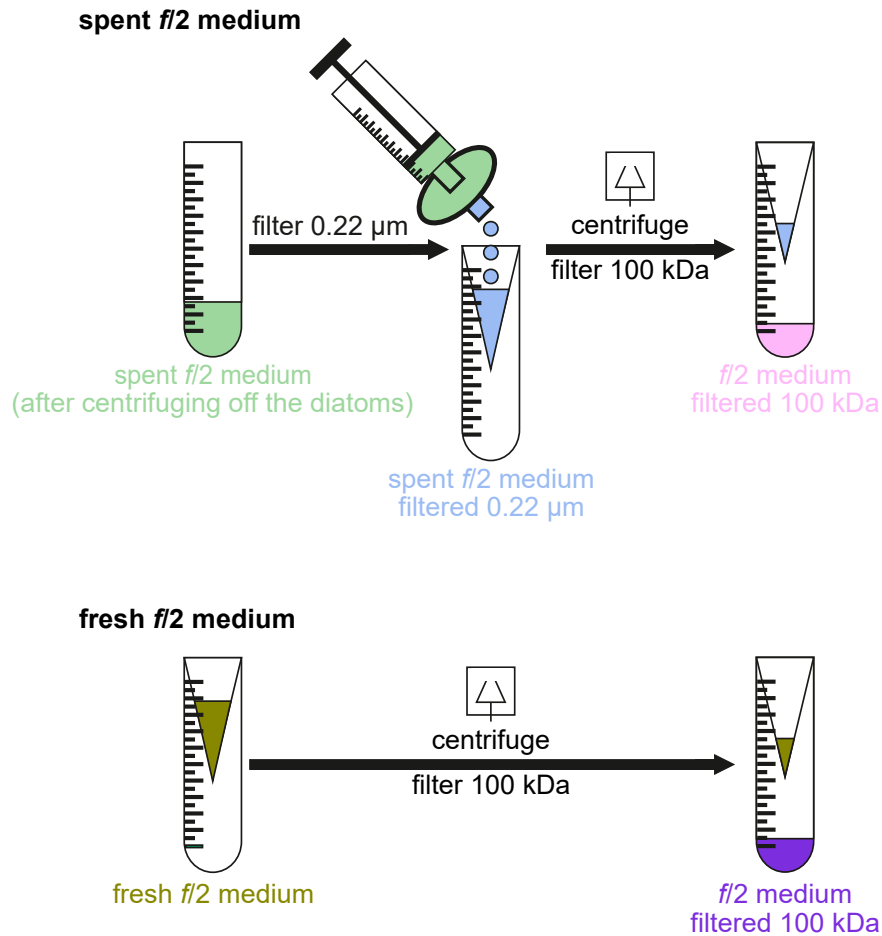


Figure S2: Sample preparation for the ice nucleation experiments with the *f/2* medium. The spent medium should only contain a few diatoms, because the diatoms were separated from the medium by centrifugation before (green vial). By filtration with a syringe-filter, we removed the remaining cells and retained smaller *F. cylindrus* fragments and the *fcIBP* in the filtrate (blue solution). The solution was filtered by centrifugation filtration and the resulting filtrate should only contain soluble macromolecules smaller than 100 kDa, e.g. *fcIBP* (pink vial). The fresh *f/2* medium (olive solution) does not contain any cells, fragments or *fcIBP* and was also filtered by centrifugation filtration as a reference (purple vial).

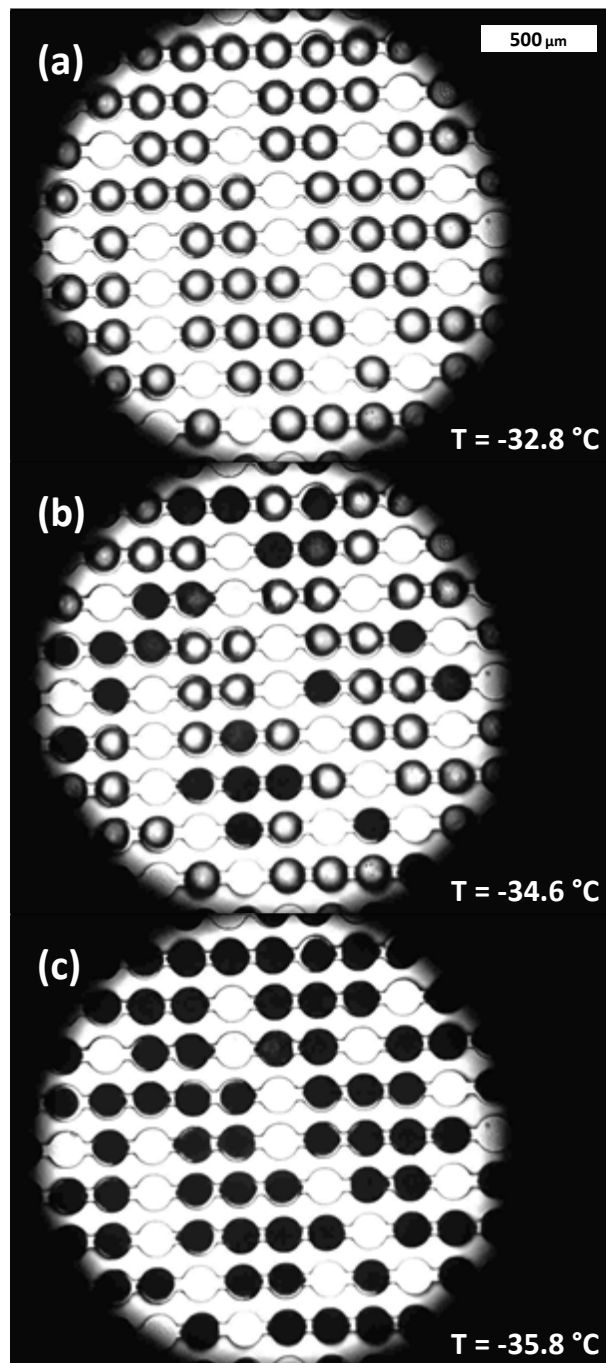


Figure S3: Optical photomicrographs of the freezing of microfluidic droplets during one of three freezing experiments with unfiltered *F. cylindrus* suspensions in artificial seawater (concentration of 2×10^6 cells per mL). The white scale bar in the top left indicates a length of 500 μm and is the same for all three images. The droplets' diameter is about (90 ± 5) μm. **a:** At a temperature of -32.8 °C all droplets are still liquid. This is the last picture before the freezing of the first droplet during this experiment. **b:** At a temperature of -34.6 °C some droplets are already frozen (black), while other droplets are still liquid (white). **c:** At a temperature of -35.8 °C all droplets are frozen. This is the first picture after the freezing of the last droplet in this experiment.

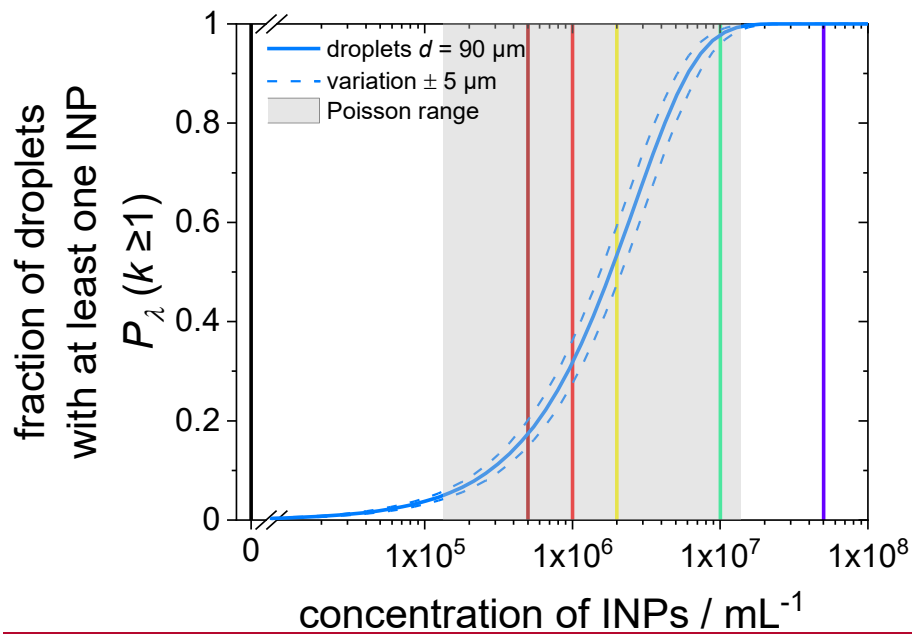


Figure S4

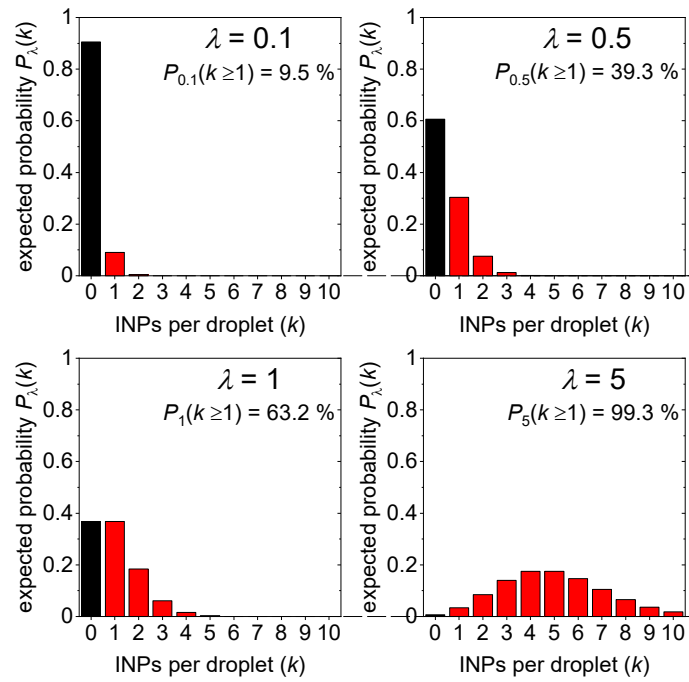


Figure S4: Calculated probability $P_\lambda(k)$ of the number k of INPs per droplet for different values λ of the average cell concentration per droplet. The black-coloured bars indicate the probability for the occurrence of droplets without any INPs, while the red-coloured bars indicate the combined probability $P_\lambda(k \geq 1)$ for droplets containing at least one INP. The corresponding values for $P_\lambda(k \geq 1)$ are annotated in each panel for different values of λ .

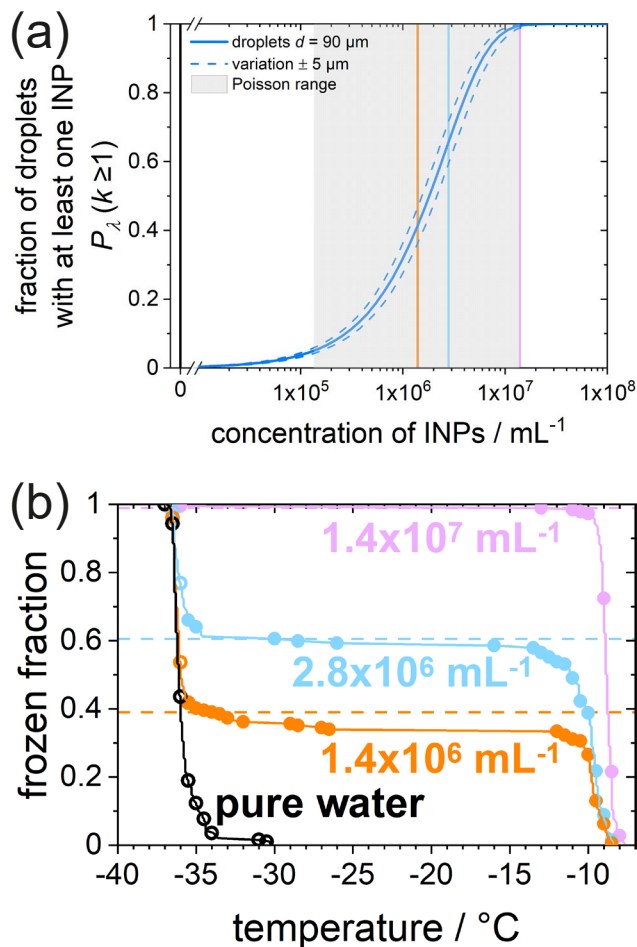


Figure S5: (a) Fraction of droplets containing at least one INP, $P_\lambda(k \geq 1)$, as a function of INP concentration in the investigated *P. syringae* sample. The solid blue curve represents the values of $P_\lambda(k \geq 1)$ for the droplets in the WISDOM experiment with a diameter of $90 \mu\text{m}$, calculated using Eq. (S7), the dashed curves indicate the uncertainty for a variation of $\pm 5 \mu\text{m}$ in droplet diameter. The grey shaded area shows the Poisson relevant range, with the lower and upper limits at the concentrations corresponding to $P_\lambda(k \geq 1) = 0.050$ and $P_\lambda(k \geq 1) = 0.995$, respectively. The coloured vertical bars mark the experimentally investigated concentrations of *P. syringae*: $1.4 \times 10^6 \text{ mL}^{-1}$ (orange), $2.8 \times 10^6 \text{ mL}^{-1}$ (blue), and $1.4 \times 10^7 \text{ mL}^{-1}$ (purple) and pure water (black). A comparable plot for the *F. cylindrus* diatoms can be found in Fig. S6. (b) Fraction of frozen droplets as a function of temperature for different concentrations of *P. syringae* bacteria in double-distilled water (coloured) and pure double-distilled water (black) for reference. The horizontal lines mark the values for $P_\lambda(k \geq 1)_{\text{measured}}$, see text. Data points of frozen fractions are binned in temperature intervals of 0.5°C (intervals without freezing events are not shown). Filled circles represent droplets containing *P. syringae* cells (based on calculations for $P_\lambda(k \geq 1)$ with Eq. (S7)), in which freezing was induced heterogeneously. Open circles represent droplets that should not contain *P. syringae* according to the calculations and, thus freeze homogeneously.

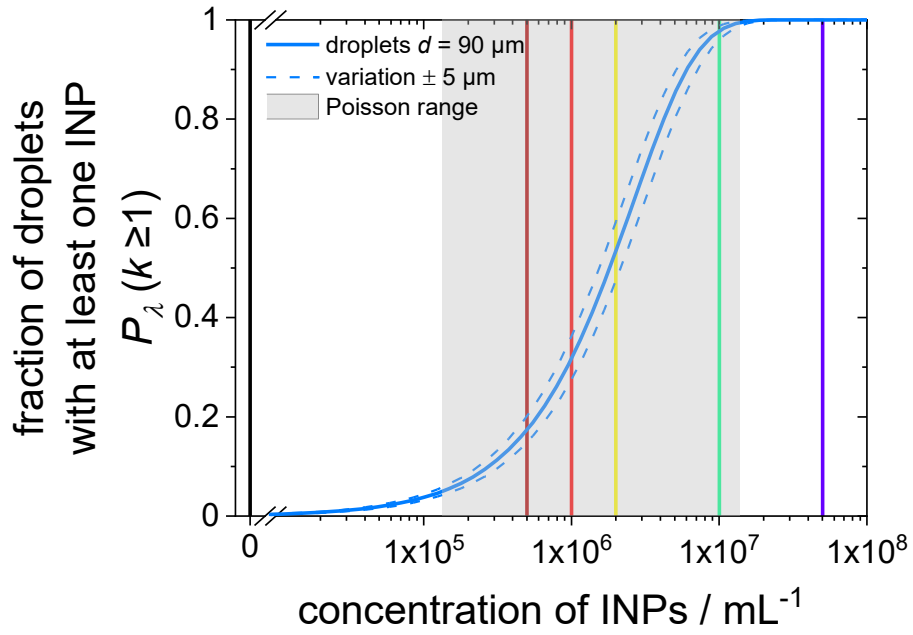


Figure S6: Fraction of droplets with at least one INP, $P_{\lambda}(k \geq 1)$, as a function of INP concentration in the investigated samples. The solid blue curve represents the values of $P_{\lambda}(k \geq 1)$ for the droplets in the microfluidic experiment with a diameter of $90 \mu\text{m}$. The dashed curves indicate the values for a variation of $\pm 5 \mu\text{m}$ in droplet diameter, i.e. $85\text{-}95 \mu\text{m}$. The calculations of these curves are based on Eq. (7S7). The grey shaded area shows the Poisson relevant range, see main text for definition, with the lower and the upper limits at the INP concentrations corresponding to $P_{\lambda}(k \geq 1) = 0.050$ and $P_{\lambda}(k \geq 1) = 0.995$. The vertical bars mark the concentration of the *F. cylindrus* diatom suspensions used in the experiments: $5 \times 10^5 \text{ mL}^{-1}$ (dark red), $1 \times 10^6 \text{ mL}^{-1}$ (bright red), $2 \times 10^6 \text{ mL}^{-1}$ (yellow), $1 \times 10^7 \text{ mL}^{-1}$ (green) and $5 \times 10^7 \text{ mL}^{-1}$ (purple) and pure seawater (black).

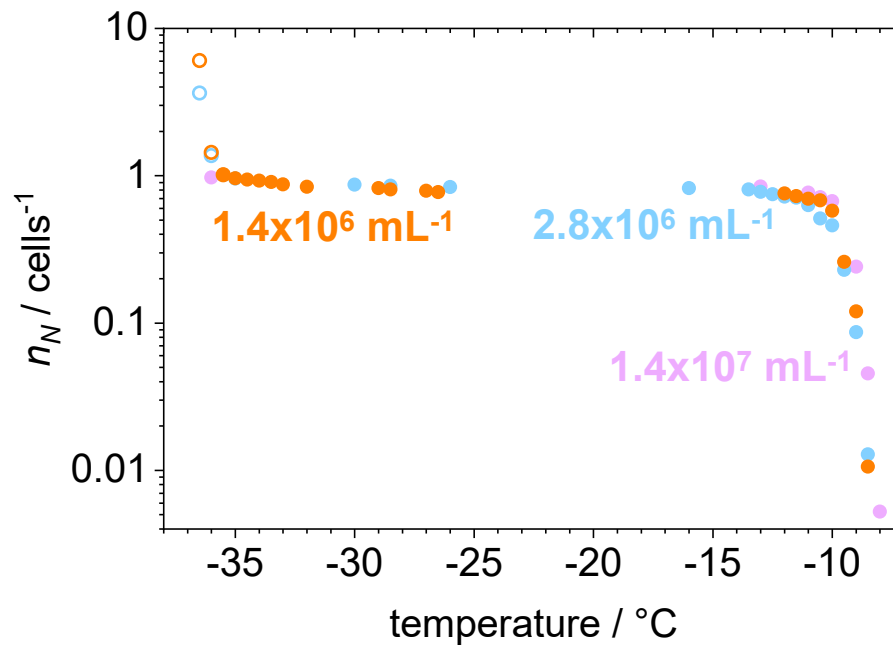
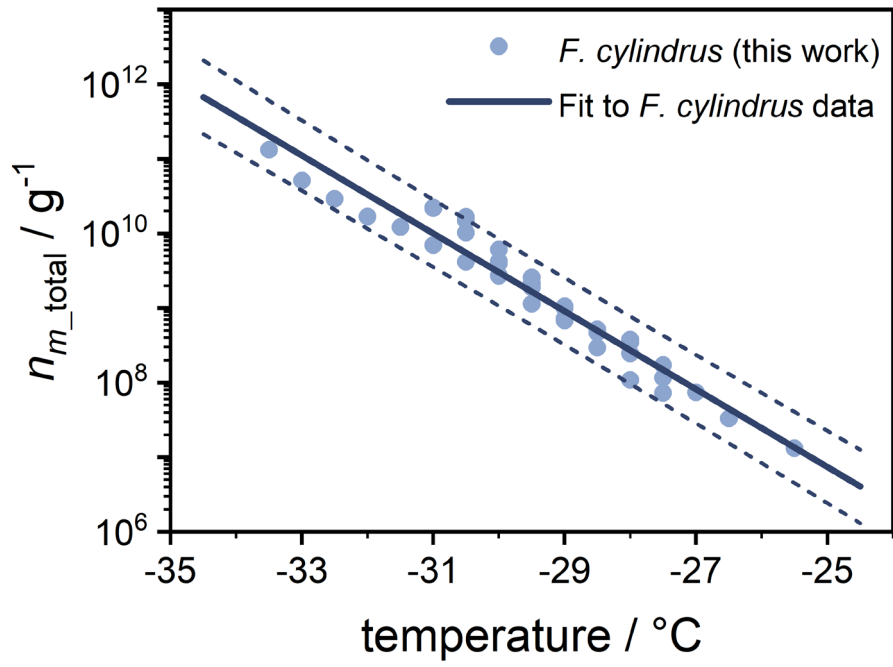


Figure S5S7: Cumulative number of ice nucleating active sites, n_N , for three different *P. syringae* bacteria suspensions with colours indicating the respective concentration. Filled circles represent droplets containing bacteria, as calculated from Eq. (7S7), while open circles represent droplets devoid of INP.



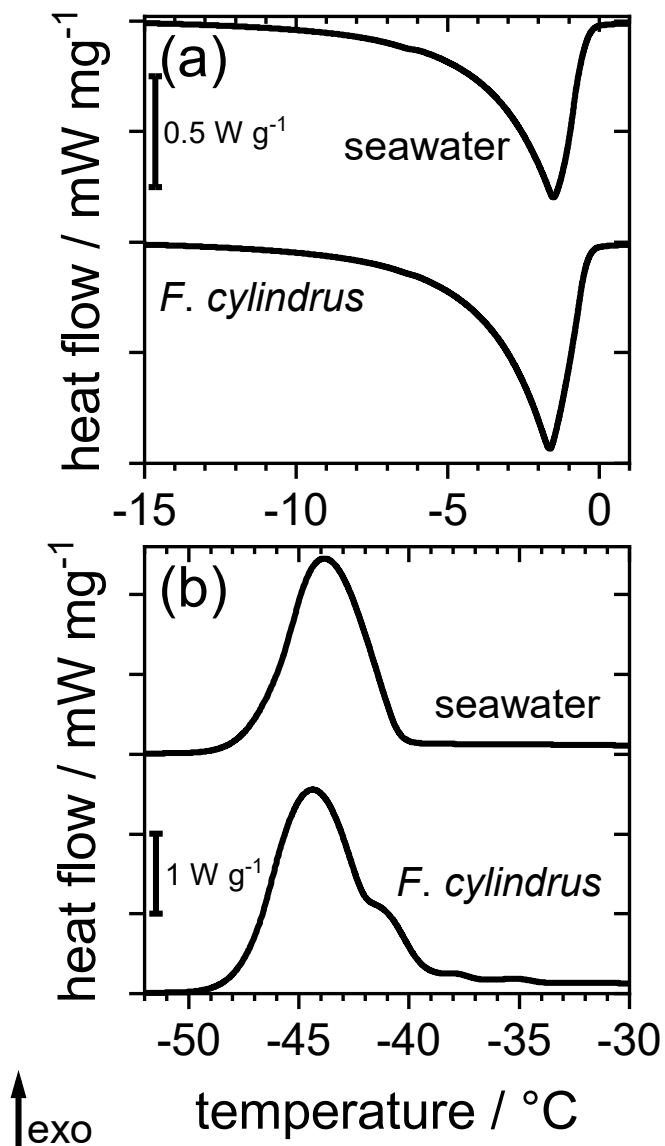


Figure S6

Figure S8: Comparison of DSC thermograms of water-in-oil emulsions containing pure artificial seawater and artificial seawater with *F. cylindrus* cells. **(a)** The endothermic melting-signals are almost identical for pure seawater and seawater containing diatoms. **(b)** Exothermic freezing signals for pure seawater and seawater containing diatoms. While the seawater emulsion shows only one freezing signal, the emulsion containing *F. cylindrus* shows the same signal but with a shoulder and smaller signals at higher temperature, indicative of diatom-induced heterogeneous ice nucleation.

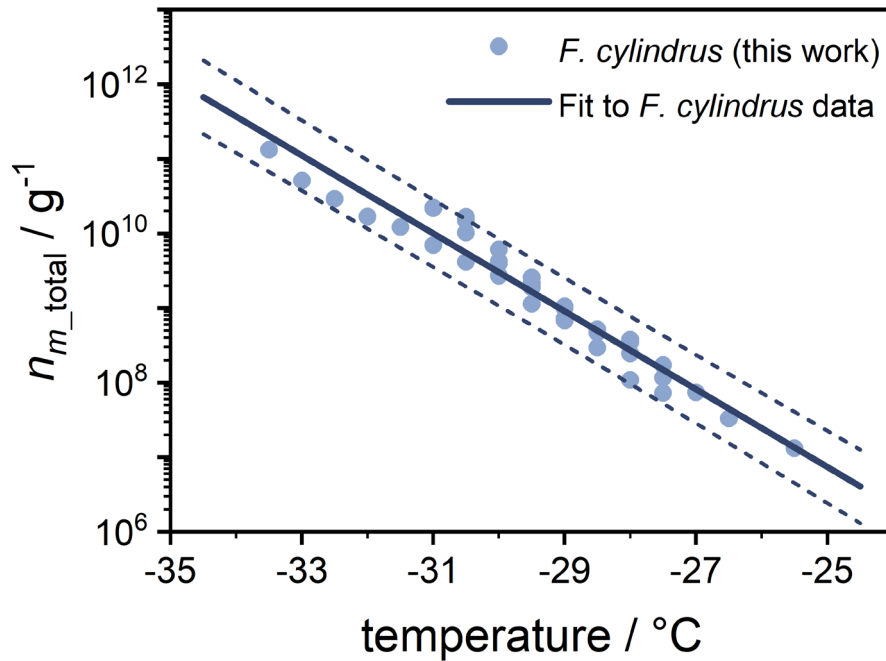


Figure S9: Measured values for n_{m_total} of *F. cylindrus* diatoms. The solid line represents a fit of the experimental n_{m_total} values (blue symbols) for the *F. cylindrus* diatoms. The parameterization of the fit is given in Eq. (S3S10). The dotted lines indicate to the upper and lower 2σ prediction bands of this fit. The temperatures are corrected for the freezing point depression of artificial seawater and, thus, represent ice nucleation temperatures in pure water.

References

- Budke, C. and Koop, T.: BINARY: an optical freezing array for assessing temperature and time dependence of heterogeneous ice nucleation, *Atmos. Meas. Tech.*, 8, 689–703, doi:10.5194/amt-8-689-2015, 2015.
- Dreischmeier, K., Budke, C., Wiehemeier, L., Kottke, T., and Koop, T.: Boreal pollen contain ice-nucleating as well as ice-binding 'antifreeze' polysaccharides, *Scientific reports*, 7, 41890, doi:10.1038/srep41890, 2017.
- Eickhoff, L., Dreischmeier, K., Zipori, A., Sirotinskaya, V., Adar, C., Reicher, N., Braslavsky, I., Rudich, Y., and Koop, T.: Contrasting Behavior of Antifreeze Proteins: Ice Growth Inhibitors and Ice Nucleation Promoters, *The journal of physical chemistry letters*, 10, 966–972, doi:10.1021/acs.jpcclett.8b03719, 2019.
- Gute, E. and Abbatt, J. P.: Ice nucleating behavior of different tree pollen in the immersion mode, *Atmospheric Environment*, 231, 117488, doi:10.1016/j.atmosenv.2020.117488, 2020.
- Koop, T., Luo, B., Biermann, U. M., Crutzen, P. J., and Peter, T.: Freezing of HNO₃/H₂SO₄/H₂O Solutions at Stratospheric Temperatures: Nucleation Statistics and Experiments, *J. Phys. Chem. A*, 101, 1117–1133, doi:10.1021/jp9626531, 1997.
- Koop, T. and Murray, B. J.: A physically constrained classical description of the homogeneous nucleation of ice in water, *The Journal of chemical physics*, 145, 211915, doi:10.1063/1.4962355, 2016.
- Morris, C. E., Sands, D. C., Bardin, M., Jaenicke, R., Vogel, B., Leyronas, C., Ariya, P. A., and Psenner, R.: Microbiology and atmospheric processes: research challenges concerning the impact of airborne micro-organisms on the atmosphere and climate, *Biogeosciences*, 8, 17–25, doi:10.5194/bg-8-17-2011, 2011.
- Murray, B. J., O'Sullivan, D., Atkinson, J. D., and Webb, M. E.: Ice nucleation by particles immersed in supercooled cloud droplets, *Chemical Society reviews*, 41, 6519–6554, doi:10.1039/c2cs35200a, 2012.
- Reicher, N., Segev, L., and Rudich, Y.: The Weizmann Supercooled Droplets Observation on a Microarray (WISDOM) and application for ambient dust, *Atmos. Meas. Tech.*, 11, 233–248, doi:10.5194/amt-11-233-2018, 2018.
- Wex, H., Augustin-Bauditz, S., Boose, Y., Budke, C., Curtius, J., Diehl, K., Dreyer, A., Frank, F., Hartmann, S., Hiranuma, N., Jantsch, E., Kanji, Z. A., Kiselev, A., Koop, T., Möhler, O., Niedermeier, D., Nillius, B., Rösch, M., Rose, D., Schmidt, C., Steinke, I., and Stratmann, F.: Intercomparing different devices for the investigation of ice nucleating particles using Snomax® as test substance, *Atmos. Chem. Phys.*, 15, 1463–1485, doi:10.5194/acp-15-1463-2015, 2015.
- Xi, Y., Mercier, A., Kuang, C., Yun, J., Christy, A., Melo, L., Maldonado, M. T., Raymond, J. A., and Bertram, A. K.: Concentrations and properties of ice nucleating substances in exudates from Antarctic sea-ice diatoms, *Environmental science. Processes & impacts*, 23, 323–334, doi:10.1039/d0em00398k, 2021.