

1 **Title:** High metabolism and periodic hypoxia associated with drifting macrophyte detritus in the
2 shallow subtidal Baltic Sea

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13 **Abstract**

14 Macrophytes form highly productive habitats that export a substantial proportion of their primary
15 production as particulate organic matter. As the detritus drifts with currents and accumulates in
16 seafloor depressions, it constitutes organic enrichment and can deteriorate O₂ conditions on the
17 seafloor. In this study, we investigate the O₂ dynamics and macrobenthic biodiversity associated
18 with a shallow ~2300 m² macrophyte detritus field in the northern Baltic Sea. The detritus,
19 primarily *Fucus vesiculosus* fragments, had a biomass of ~1700 g dry weight m⁻², approximately
20 1.5-fold larger than nearby intact *F. vesiculosus* canopies. A vertical array of O₂ sensors placed
21 within the detritus documented that hypoxia ([O₂] < 63 μmol L⁻¹) occurred for 23% of the time and
22 terminated at the onset of wave-driven hydrodynamic mixing. Measurements in five other habitats
23 nearby spanning bare sediments, seagrass, and macroalgae indicate that hypoxic conditions were
24 unique to detritus canopies. Fast-response O₂ sensors placed above the detritus documented pulses
25 of hypoxic waters originating from within the canopy. These pulses triggered a rapid short-term (~5
26 min) deterioration of O₂ conditions within the water column. Eddy covariance measurements of O₂
27 fluxes indicated high metabolic rates with daily photosynthetic production offsetting up to 81 % of
28 the respiratory demands of the detritus canopy, prolonging its persistence within the coastal zone.

29 The detritus site had a low abundance of crustaceans, bivalves, and polychaetes when compared to
30 other habitats nearby, likely because their low-O₂ tolerance thresholds were often exceeded.

31 **1. Introduction**

32 Oxygen availability determines ecosystem health and the biogeochemical function of coastal waters
33 (Diaz and Rosenberg, 2008; Middelburg and Levin, 2009; Breitburg et al., 2018). When in gaseous
34 equilibrium with air, seawater typically contains an O₂ concentration ([O₂]) between 200-400 μmol
35 L⁻¹, depending on the water temperature and the salinity (Garcia and Gordon, 1992). However, both
36 abiotic and biotic processes cause significant departures from equilibrium. The main source of O₂ to
37 coastal waters is the atmosphere, where the diffusion of O₂ is governed by the air-to-sea gas
38 exchange rate (Berg and Pace, 2017; Long and Nicholson, 2018). In shallow waters and light-
39 exposed seafloor sediments, O₂ is produced by primary producers as a by-product of
40 photosynthesis, and it is consumed by consortia of microbes and fauna directly, through aerobic
41 respiration, and indirectly, through the oxidation of reduced substances (Glud, 2008). If O₂
42 consumption exceeds supply for a sufficiently long period, O₂ conditions deteriorate and become
43 hypoxic ([O₂] < 63 μmol L⁻¹). Hypoxia is becoming more common, more intense, and is affecting
44 larger areas of coastal waters, increasingly placing ecosystems and the services they provide at risk
45 (Breitburg et al., 2018). There are several well-known variants of coastal hypoxia (Diaz and
46 Rosenberg, 2008; Carstensen and Conley, 2019). Seasonal hypoxia, the most common form,
47 typically occurs in summer when warm waters, strong stratification, and high organic enrichment
48 combine to deplete O₂ until autumn (Robertson et al., 2016). Periodic hypoxia, in contrast, occurs
49 more often due to local weather dynamics and tidal cycles but individual events are shorter (Diaz
50 and Rosenberg, 1995), whereas diel cycles with large day-to-night [O₂] excursions trigger hypoxia
51 for a few hours daily (Davanzo and Kremer, 1994; Tyler et al., 2009). All events are expected to
52 affect biodiversity and biogeochemical cycling to varying degrees. Seasonal and periodic hypoxia
53 are associated with large-scale mortality of organisms and a switch between retention and removal
54 of bioavailable nutrients such as nitrate, ammonium, phosphate, and toxic hydrogen sulfide
55 (Middelburg and Levin, 2009; Carstensen and Conley, 2019). Short-term hypoxia can similarly
56 exceed lethal and non-lethal thresholds for many taxa (Vaquer-Sunyer and Duarte, 2008), although,
57 due to their sporadic nature, their occurrence and impacts are less understood.

58 Given the importance of O₂ in coastal waters, [O₂] is one of the most frequently measured
59 environmental parameters. Near-seabed [O₂] is typically measured using long-term stable O₂

60 sensors (e.g. optodes) (Bittig et al., 2018) that are moored ~0.3-1.0 m above the seafloor, or by
61 performing vertical profiles of water column [O₂] down to ~1.0 m above the seafloor using
62 multiparameter sondes. National monitoring programs such as those maintained by the Swedish
63 Meteorological and Hydrological Institute and the Finnish Environment Institute provide a wealth
64 of essential open-access data, enabling important analyses detailing the prevalence and intensity of
65 coastal hypoxia (Virtanen et al., 2019; Conley et al., 2011; Carstensen and Conley, 2019).
66 Notwithstanding the progress being made in coastal monitoring, it was demonstrated more than 40
67 years ago that the largest [O₂] gradients may occur just a few cm above the seafloor due to the high
68 reactivity of marine sediments and a strong benthic O₂ demand (Jorgensen, 1980). To date, records
69 of hypoxia in the shallow subtidal zone are still somewhat scarce. In a compilation of monitoring
70 data for the northern Baltic Sea (Gulf of Finland and Archipelago Sea), Virtanen et al. (2019) found
71 that just 11 out of 461 (or 2.4%) of the monitoring stations that registered hypoxia occurred in
72 waters < 5 m depth. While this may reflect a true signal that hypoxia is more widespread in deeper
73 coastal waters, it is also likely that hypoxic conditions go undetected if measurements are
74 performed away from the seafloor, as is common practice (Conley et al., 2011; Virtanen et al.,
75 2019).

76 Around two-thirds of the ocean's photosynthetic biomass is bound in macrophytes growing in
77 shallow waters along the world's coastline (Smith, 1981). Through seasonal decay, epiphyte
78 growth, grazing, and physical forcing (e.g. waves, currents, ice scouring), macrophytes export a
79 large proportion of their primary production (~40 %) to their surroundings as detritus (Attard et al.,
80 2019a; Krumhansl and Scheibling, 2012; Duarte and Cebrián, 1996). Macrophyte detritus drifts
81 with the currents and accumulates on the shoreline and in low-energy marine environments (e.g.
82 shallow seafloor depressions and in deeper waters), where it constitutes habitat structure and
83 organic enrichment to the receiving habitat (Norkko and Bonsdorff, 1996b). Given high enough
84 abundance, detritus suppresses the diffusion of O₂ from the water column to the sediment surface
85 and it exacerbates O₂ depletion on the seabed as it decays. Large accumulations of unattached
86 ephemeral macroalgae such as the brown algae *Ectocarpus siliculosus* and *Pylaiella littoralis* are
87 common in eutrophic coastal waters such as the Baltic Sea, forming thin mats above the seafloor
88 typically a few centimeters thick (Norkko and Bonsdorff, 1996a). While coastal hypoxia is most
89 commonly associated with eutrophic waters such as the Baltic Sea (Carstensen and Conley, 2019;
90 Conley et al., 2011), hypoxic (and even sulfidic) conditions have been reported in remote and more
91 pristine environments such as the high Arctic due to large accumulations of detritus produced from

92 perennial brown seaweeds (Glud et al., 2004). However, the O₂ dynamics within accumulations of
 93 drifting detritus and the potential implications for the associated fauna remain poorly understood.
 94 Understanding the ecological and biogeochemical implications of drifting macrophyte detritus is
 95 particularly important given the ambitions to vastly increase macroalgal farming (Broch et al.,
 96 2019), which would result in increased deposition of macrophyte detritus on the coastal seafloor
 97 (Broch et al., 2022).

98 In this study, we investigate the O₂ dynamics and macrobenthic biodiversity associated with a
 99 shallow ~2300 m² macrophyte detritus field composed of *Fucus vesiculosus* fragments in the
 100 northern Baltic Sea. To assess O₂ production versus consumption rates of the detritus canopy, we
 101 deployed an eddy covariance system on multiple occasions to extract benthic O₂ fluxes non-
 102 invasively. Using a vertical array of O₂ sensors and an acoustic velocimeter, we monitored O₂
 103 distribution within the canopy and the hydrodynamics above the canopy to assess the occurrence
 104 and intensity of hypoxic events and their links to local hydrodynamics. We performed biodiversity
 105 surveys to identify the prevailing taxa, and we compared hypoxic thresholds of these taxa to [O₂]
 106 measured *in situ* to identify potential stress. Measurements were also performed in five other
 107 habitats nearby spanning bare sediments, seagrass, and macroalgae for comparison.

108 2. Materials and Methods

109 2.1. Study location

110 The study was performed in the microtidal Baltic Sea nearby the Tvärminne Zoological Station in
 111 SW Finland. Although the focus of our study was to investigate drifting macrophyte detritus, we
 112 selected an additional five study sites within the shallow subtidal zone (2-4 m depth) for
 113 comparison, representing key habitats in the Baltic Sea: one site with bare sediments, two sites with
 114 seagrass (predominantly *Zostera marina*; sheltered and exposed), and two sites with intact
 115 macroalgae canopies (predominantly *Fucus vesiculosus*; sheltered and exposed) (Table 1).

Site	Location	Deployment start	Deployment duration (h)	Water depth (m)	Water temperature (°C)	Minimum O ₂ (μmol L ⁻¹)	Maximum O ₂ (μmol L ⁻¹)	Hypoxia duration (h)
Macrophyte detritus	59 811613 N 23 206624 E	29-05-2018	120	3.0	12	0.6	429	27
Bare sediments	59 841532 N 23 253370 E	20-05-2018	96	3.7	11	307	407	0
Sheltered <i>Z. marina</i>	59 841551 N 23 251203 E	27-05-2018	87	4.0	16	272	333	0

Exposed <i>Z. marina</i>	59 827008 N 23 151976 E	08-06-2018	120	2.9	10	281	437	0
Sheltered <i>F. vesiculosus</i>	59 826856 N 23 209721 E	08-06-2018	120	2.0	10	253	489	0
Exposed <i>F. vesiculosus</i>	59 811359 N 23 207281 E	31-05-2018	116	2.0	9	287	427	0

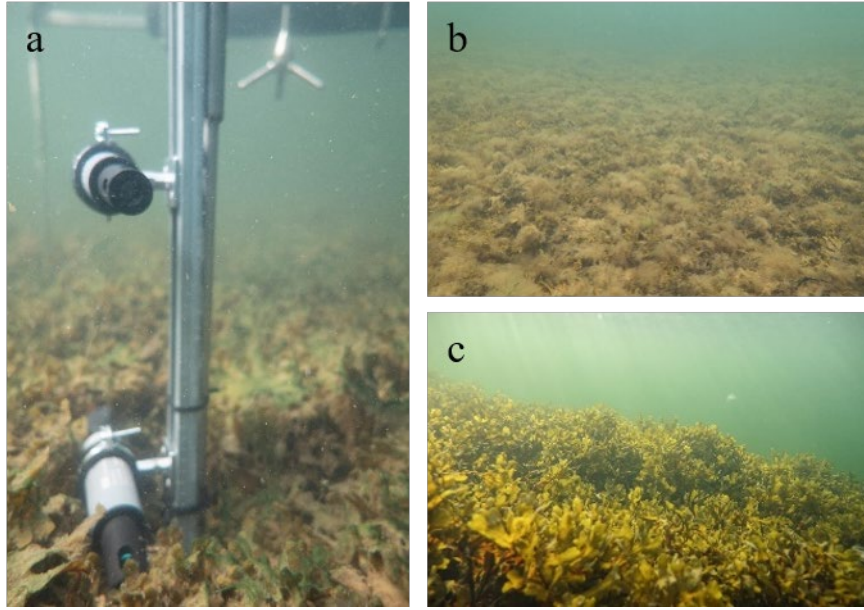
116 Table 1: Environmental conditions and low-oxygen events at the six study sites

117

118 2.2. $[O_2]$ dynamics in benthic habitats

119 To investigate the near-bed $[O_2]$ dynamics and its environmental controls, we equipped a tripod
120 frame with a suite of sensors consisting of three cross-calibrated dissolved $[O_2]$ loggers with inbuilt
121 temperature compensation (HOBO U26-001, Onset), a 6 MHz acoustic velocimeter (Vector,
122 Nortek), a photosynthetic active radiation (PAR) sensor (RBRsolo with Licor PAR Quantum
123 192SA), and a saltwater conductivity sensor (HOBO U24-002-C). The $[O_2]$ loggers have a factory-
124 specified accuracy of $\pm 6 \mu\text{mol L}^{-1}$ from 0 to $250 \mu\text{mol L}^{-1}$, $\pm 16 \mu\text{mol L}^{-1}$ from $250\text{-}625 \mu\text{mol L}^{-1}$, a
125 resolution of $0.6 \mu\text{mol L}^{-1}$ and a 90% response time (T_{90}) < 2 min. The $[O_2]$ and conductivity
126 sensors were mounted onto a 75 cm-long stainless-steel rail affixed to the tripod leg (Fig. 1). The
127 sensors were secured to the rail at various heights above the seabed using rail mount clamps. For the
128 study sites with canopies, two sensors were set inside the canopy; one sensor was ~ 5 cm above the
129 seafloor and one was close to the top of the canopy (15-25 cm). The third sensor was placed in the
130 water above the canopy (~ 35 cm above the seafloor). The tripod was deployed by divers from a
131 small boat and was carefully positioned on the seafloor using a lift bag. The exact sensor heights
132 were noted by the divers once the instrument was on the seafloor. The instrument was left to record
133 data for 3-5 days at each site. The velocimeter sampled three-dimensional flow velocity
134 continuously at 8 Hz, whereas the $[O_2]$, temperature, conductivity, and PAR sensors recorded data
135 every minute.

136 To investigate $[O_2]$ dynamics and its environmental drivers, all sensor time series were aligned in
137 time and analyses were performed to investigate vertical gradients in O_2 distribution, diel $[O_2]$
138 excursions, and boundary-layer hydrodynamics. We assessed the occurrence of hypoxia ($[O_2] < 63$
139 $\mu\text{mol L}^{-1}$) by quantifying the magnitude (lowest $[O_2]$ value) and the duration (in hours) of hypoxic
140 events. The high-frequency velocity data were used to calculate mean flow velocity magnitude (\bar{U})
141 as the sum of streamwise (u) and traverse (v) components, as $\bar{U} = \sqrt{u^2 + v^2}$.



142

143 Fig. 1: The study area showing (a) the instrument deployed within the detritus canopy, (b) a broad-
144 scale view of the detritus accumulation area, and (c) a nearby intact *Fucus vesiculosus* canopy.

145 2.3. Benthic O_2 fluxes

146 An aquatic eddy covariance system was deployed at the detritus site to quantify benthic O_2 fluxes at
147 the canopy-water interface on three occasions (June 2017, September 2017, and May 2018). Eddy
148 covariance integrates over a relatively large seafloor area (typically $\sim 30 \text{ m}^2$) (Berg et al., 2007), and
149 extracts fluxes without disturbing the hydrodynamics or the light, which is particularly useful when
150 trying to understand the mechanistic drivers of $[O_2]$ dynamics (Berg et al., 2022). The eddy
151 covariance setup was identical to the tripod frame described above, with the addition of a fast-
152 response ($T_{90} < 0.3 \text{ s}$) $[O_2]$ microsensor setup for covariance measurements (McGinnis et al., 2011).
153 The hardware and data processing techniques are described in detail in Attard et al. (2019b). This
154 instrument can capture the entire range of flux-contributing turbulent eddies within the benthic
155 boundary layer, and this information is used to approximate the benthic O_2 flux non-invasively
156 (Berg et al., 2003; Berg et al., 2022). The instrument recorded co-located measurements of the
157 vertical velocity (w) and the O_2 concentration (C) at 32 Hz, and the data were processed using a
158 multiple-step protocol detailed in Attard et al. (2019b) to extract and quality-check benthic fluxes.
159 In short, the data streams for w and C were decomposed into mean and fluctuating components
160 using Reynolds decomposition, as $w = \bar{w} + w'$ and $C = \bar{C} + C'$ (Berg et al., 2003). The turbulent
161 flux (J_{EC}) was then computed in units of $\text{mmol } O_2 \text{ m}^{-2} \text{ h}^{-1}$ as $J_{EC} = \overline{w'C'}$, where the overbar

162 represents a period of 15 min. The turbulent flux was then summed with a storage correction term to
163 calculate the total benthic flux ($J_{benthic}$, mmol O₂ m⁻² h⁻¹) (Rheuban et al., 2014), as:

164
$$J_{benthic} = J_{EC} + \int_0^h \frac{\partial C}{\partial t} dz$$

165 The storage correction term was defined using the three [O₂] optodes placed within and above the
166 canopy. For the correction, we defined a matrix with the number of rows n corresponding to the
167 sensor measurement height above the seafloor (1 row per cm) (Camillini et al., 2021). To do this,
168 the oxygen time series, consisting of [O₂] measurements performed at three heights within the
169 canopy, were converted to a matrix using the software package OriginPro 2022. Since the
170 measurement height of the three sensors were spaced nonlinearly, the data were first converted to
171 XYZ column format using the w2xyz function. Next, the three rows, representing the [O₂] time
172 series measurements at three heights, were expanded to n rows, with n representing the sensor
173 measurement height in cm (from 0 to n cm above seabed, 1 row per cm) using the XYZ Gridding
174 function. This generated a matrix of n rows consisting of linearly interpolated [O₂]. Interpolation
175 was performed using the Random (Renka Cline) gridding method. Next, a storage correction term
176 was calculated for each 1 cm cell as described by Rheuban et al. (2014), and the total storage
177 correction was subsequently computed for the water volume below the sensor measurement height
178 as the sum of the n rows. The high-frequency [O₂] time series from the fast-response microsensors
179 were also analyzed to identify any pulses of low [O₂] waters originating from within the canopy and
180 propagating upwards into the water column.

181 2.4. Benthic metabolic rates

182 The O₂ flux time series was separated into individual 24 h periods (midnight to midnight). The
183 daytime flux ($Flux_{day}$, mmol O₂ m⁻² h⁻¹) was computed as a bulk average of fluxes measured when
184 PAR > 1.0 μmol m⁻² s⁻¹. The nighttime flux ($Flux_{night}$, mmol O₂ m⁻² h⁻¹) was calculated as the
185 average of the remaining fluxes, when PAR < 1.0 μmol m⁻² s⁻¹. These two values and the number of
186 daylight hours (h_{day}) were used to estimate the daily photosynthetic rate, termed the gross primary
187 production (GPP , in mmol O₂ m⁻² d⁻¹), as $GPP = Flux_{day} + abs(Flux_{night}) * h_{day}$, and daily
188 respiration (R , in mmol O₂ m⁻² d⁻¹), as $R = abs(Flux_{night}) * 24$, assuming a light-independent
189 respiration rate. The latter is a common assumption, but it is known that it underestimates the true
190 metabolic activity (Fenchel and Glud, 2000; Juska and Berg, 2022). The daily balance between

191 *GPP* and *R*, termed the net ecosystem metabolism (*NEM*, in mmol O₂ m⁻² d⁻¹) was estimated as
192 $NEM = GPP - R$ (Attard et al., 2019b).

193 The relationship between seafloor PAR and the in situ benthic O₂ flux was investigated using light-
194 saturation curves. Hourly O₂ fluxes were plotted against the corresponding near-bed incident PAR
195 and the relationship between the two was investigated using a modified tangential hyperbolic
196 function by Platt et al. (1980), as $O_2 \text{ flux} = P_m * \tanh\left(\frac{\alpha I}{P_m}\right) - R$, where P_m is the maximum rate of
197 hourly gross primary production, α is the initial quasi-linear increase in O₂ flux with PAR, I is near-
198 bed irradiance (PAR), and R is the dark respiration rate. The photosaturation parameter, I_k (μmol
199 PAR m⁻² s⁻¹) was derived as P_m/α . Non-linear curve fitting was performed in OriginPro 2022 using
200 a Levenberg–Marquardt iteration algorithm, until a Chi-Squared tolerance value of 1E-9 was
201 reached (Attard and Glud, 2020).

202 2.5. Biodiversity sampling

203 At all six sites, we aimed to obtain a quantitative understanding of the abundance, biomass, and
204 species richness of macrophytes and macrofauna (infauna and epifauna). The different habitats
205 required different sampling strategies, since four sites were sedimentary (bare sediments site, two
206 seagrass sites, and the detritus sites) and two sites were rocky (two macroalgal sites) (Rodil et al.,
207 2019).

208 At the time of our study, the detritus site had a ~20-cm thick detritus mat covering the seabed
209 sediments. The detritus canopy was sampled using large stainless steel core liners (inner diameter =
210 19 cm; $n = 4$) capable of cutting through the mat, and the collected samples were transferred into a
211 fine-mesh bag. In the laboratory, the detritus was rinsed through a 0.5 mm sieve to collect the
212 associated epifauna. Samples of algal detritus were dried at 60°C for 48 hours and the biomass was
213 calculated as dry weight /m².

214 Macroinfauna at the four sedimentary habitats was sampled using six sediment cores (inner
215 diameter = 5.0 cm, depth = 15 cm). The samples were sieved through a 0.5 mm sieve and animals
216 were stored in alcohol for later identification. At the seagrass sites, representative macrophyte
217 samples were collected by divers from an area around the tripod frame at the end of the deployment
218 using four randomly-placed quadrats (20 x 20 cm). The seagrass within each quadrat was gently
219 uprooted and was transferred into a net-bag. In the laboratory, the samples were rinsed through a
220 0.5 mm sieve to collect all the associated epifauna. The animals were stored in alcohol for later

221 identification, and the seagrass was frozen in sealed bags for further processing. The seagrass
222 samples were later thawed, and individual shoots were counted to determine the canopy density in
223 m². The above- and below-ground macrophyte biomass was separated, dried at 60°C for 48 hours
224 and weighed.

225 At the rocky sites, *F. vesiculosus* individuals ($n = 4$) were randomly collected from around the
226 instrument in fine-mesh bags. Randomly-placed quadrats (1 m², $n = 4$) were used to quantify the
227 number of *F. vesiculosus* individuals per m². At the laboratory, the collected *F. vesiculosus* samples
228 were carefully rinsed through a 0.5 mm sieve to collect the epifauna. The height of the *F.*
229 *vesiculosus* canopy was determined from the average length of the sampled individuals. Both *F.*
230 *vesiculosus* and epiphytes were separated to the extent possible, dried at 60 °C for 48 h and
231 weighed. To collect any macrofauna on the bare rock beneath the *F. vesiculosus* canopy, Kautsky-
232 type samplers were placed on the seafloor and the 20 cm x 20 cm area was gently scraped using a
233 spoon into a fine-mesh sampling bag. In the laboratory, all the macrofauna from the four replicates
234 were sieved through a 0.5 mm sieve and stored in alcohol.

235 The fauna from all habitats was sorted, identified to species level, counted, and weighed. The wet
236 weight for each species was noted with 0.0001 g accuracy. In cases where the fauna occurred in
237 very high numbers, the sample was placed in a water-filled tray and divided into eight sectors. Four
238 sectors were randomly chosen to calculate abundance and biomass. The length of gastropods and
239 bivalves was measured from anterior to posterior axis using Vernier callipers (accuracy = 0.01 mm)
240 for conversion to ash-free dry mass (AFDM). The AFDM of bivalves and gastropods was
241 calculated using established relationships between length and weight for Baltic Sea fauna (Rumohr
242 et al., 1987).

243 The abundance (ind m⁻²) and biomass (AFDM/SFDM g m⁻²) of the invertebrates across sites were
244 calculated. Primer (v.7 and PERMANOVA+) software was used to perform the nonmetric
245 multidimensional scaling (nMDS, with fourth-root-transformed data) to visualize macrofauna
246 assemblages between sites. ANOSIM based on the Bray-Curtis similarity matrix was also
247 performed in Primer (site as a fixed factor, 4999 random sample permutations) to compare
248 differences in macrofauna abundance and biomass between sites.

249

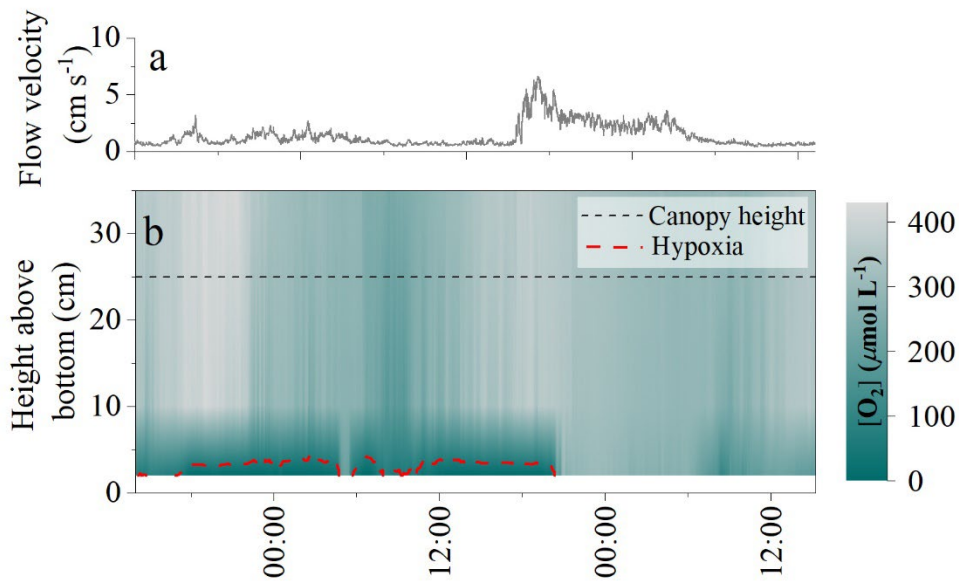
250 **3. Results**

251 *3.1. Environmental conditions*

252 Average water depth ranged from 2.0 m to 4.0 m at the six study sites, and average water
253 temperature ranged from 9 °C to 16 °C during the study period (Table 1). Hypoxic conditions were
254 only detected at the detritus site. Bottom-water [O₂] at the detritus site ranged from 1 μmol L⁻¹ to
255 429 μmol L⁻¹, with hypoxic conditions occurring for 27 h out of the 120 h long deployment (i.e. for
256 23 % of the time) (Table 1). At the five other measurement sites, [O₂] were well above hypoxic
257 conditions, with overall concentrations following diel patterns and ranging from 250 μmol L⁻¹ to
258 490 μmol L⁻¹ (Table 1).

259 3.2. [O₂] dynamics in detritus canopies

260 The [O₂] measurements within the detrital canopy document a highly dynamic [O₂] environment
261 driven by light availability and flow velocity (Fig. 2). Within the upper layers of the canopy (i.e.
262 ~10 to 25 cm above the seafloor), [O₂] and temporal dynamics largely follow diel patterns driven by
263 light availability, with large ~250 μmol L⁻¹ diel excursions in [O₂]. There, the [O₂] was lowest in
264 the morning (~160 μmol L⁻¹) and highest in the evening (~430 μmol L⁻¹). In all cases, [O₂] within
265 the upper canopy region was above hypoxic thresholds. However, under low average flow
266 velocities < 2 cm s⁻¹, [O₂] within the lower canopy region (< 10 cm) deviated substantially from the
267 conditions above. No diel variations in [O₂] were observed during these periods, and [O₂] rapidly
268 became hypoxic for sustained periods (> 24 h long), with [O₂] being very low (< 10 μmol L⁻¹)
269 during ~10 hr (~8 % of the time) (Fig. 2). As hypoxia persisted throughout the night under low flow
270 velocities, low [O₂] extended upwards into the canopy. Hypoxic conditions ended at the onset of
271 higher mean flow velocities of ~7 cm s⁻¹, which initiated a rapid (i.e. within 1.5 hr) oxygenation of
272 the entire canopy.

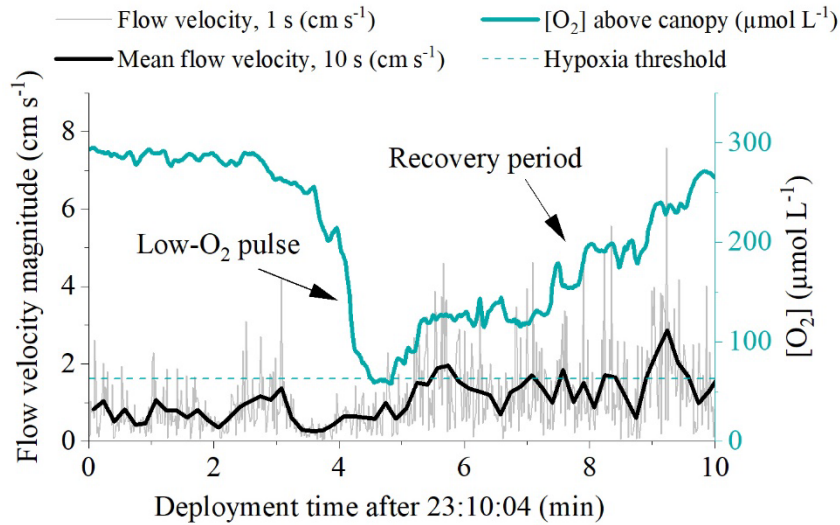


273

274 Fig. 2: (a) Flow velocity measured by the velocimeter 10 cm above the detritus canopy and (b) O_2
 275 distribution within the canopy as resolved by three O_2 sensors located at 3 cm, 10 cm, and 35 cm
 276 above the seafloor. Deployment starting from 29th May 2018.

277 3.3. Pulses of hypoxic waters

278 High-frequency $[\text{O}_2]$ measurements performed 10 cm above the detritus canopy document transient
 279 pulses of hypoxic water originating from within the canopy and propagating upwards into the water
 280 column (Fig 3). Such pulses typically followed quiescent weather and occurred at the onset of
 281 increased flow velocities. It took < 1 min to reduce $[\text{O}_2]$ in the water column from $220 \mu\text{mol L}^{-1}$ to
 282 $65 \mu\text{mol L}^{-1}$. Subsequently, a recovery period followed where $[\text{O}_2]$ gradually increased back to
 283 previous concentrations over a ~ 5 min period. These rapid variations in water column $[\text{O}_2]$ were not
 284 captured by the slow-response $[\text{O}_2]$ optode sampling at 1 min intervals.



285

286 Figure 3: High-frequency [O₂] measured 10 cm above the detrital canopy documented pulses of
 287 hypoxic water originating from within the canopy and propagating upwards into the water column.
 288 Data from 20th September 2017.

289

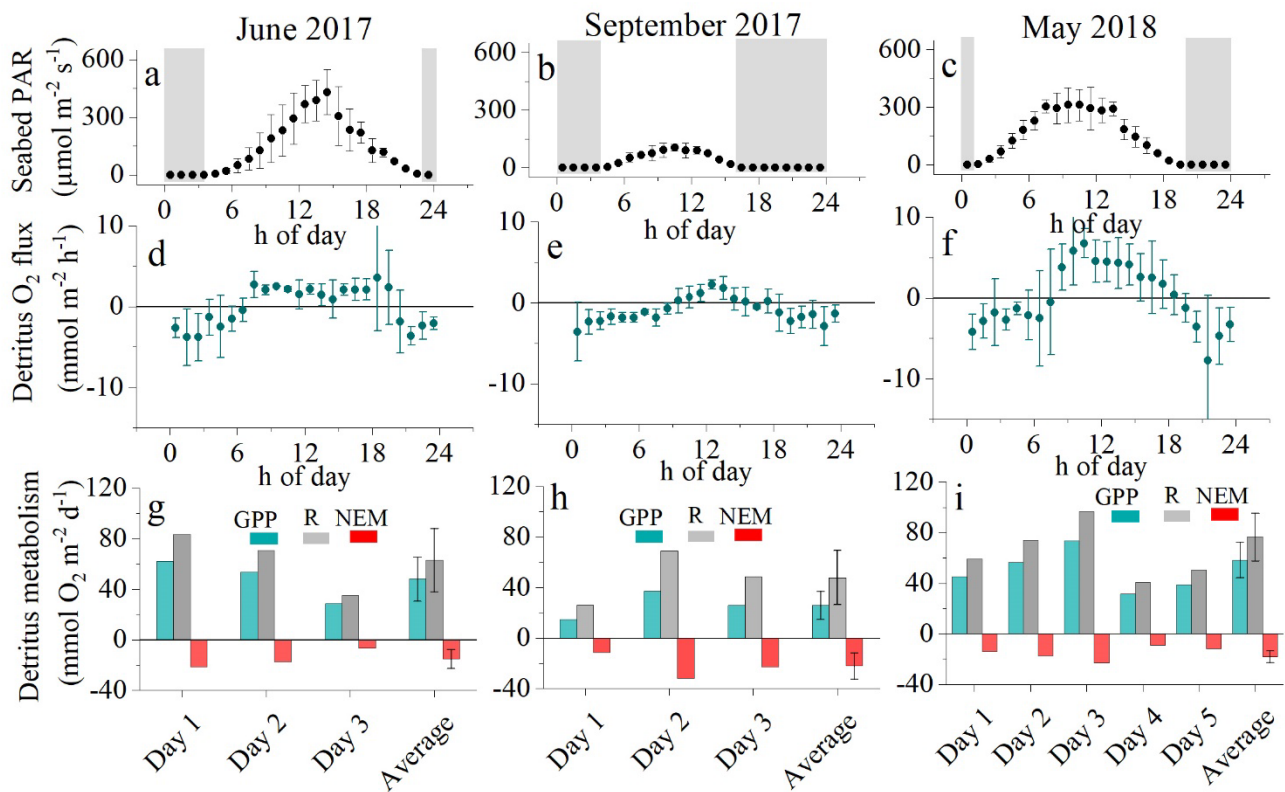
290 3.4. Benthic O₂ fluxes and detritus metabolic rates

291 The eddy covariance measurements at the detritus site produced three days of continuous flux data
 292 in June 2017, three days of data in September 2017, and five days of data in May 2018. Benthic O₂
 293 fluxes documented a dynamic O₂ exchange rate driven by light availability and flow velocity.
 294 During quiescent periods with low flow velocity < 2 cm s⁻¹, a clear diel signal in the O₂ flux was
 295 observed, indicating substantial primary production associated to the detritus canopy. Higher flow
 296 velocities stimulated O₂ uptake rates by up to 5-fold, indicating that canopy ventilation through
 297 mixing increased O₂ uptake.

298 Hourly O₂ fluxes ranged from -22 mmol O₂ m⁻² h⁻¹ at night to 13 mmol O₂ m⁻² h⁻¹ during the day
 299 and showed a distinct diel cycle in response to sunlight availability (Fig. 4). Daily R ranged from 26
 300 to 97 mmol O₂ m⁻² d⁻¹, and daily GPP was between 15 and 74 mmol O₂ m⁻² d⁻¹. Daily R exceeded
 301 GPP in all 11 measurement days (net heterotrophic), with NEM ranging from -7 to -32 mmol O₂ m⁻²
 302 d⁻¹ (Fig. 4, Table A1). The deployment average (± SD) GPP:R for the detritus canopy was 0.77 ±
 303 0.04 in June 2017 (n = 3), 0.55 ± 0.02 in September 2017 (n = 3), and 0.77 ± 0.00 in May 2018 (n
 304 = 5), and the global mean GPP:R was 0.71 ± 0.11 (n = 11).

305 There was a significant positive relationship between daily detritus GPP and R in all measurement
 306 campaigns, with the detritus canopy seemingly becoming more heterotrophic (i.e. R > GPP) as the

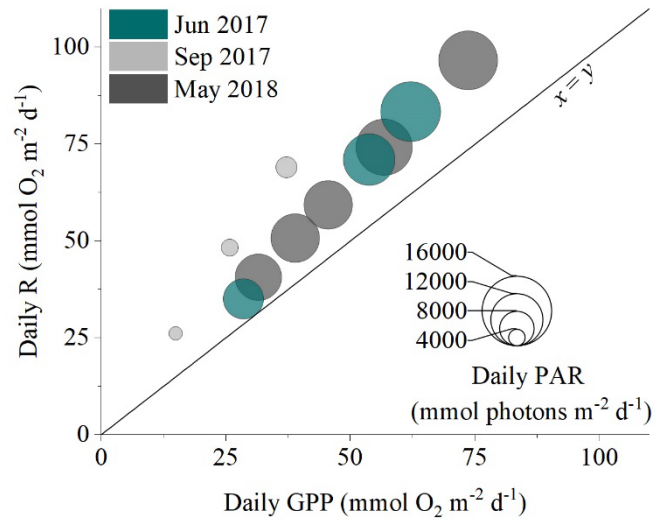
307 magnitude of the metabolic rates increased (Fig. 5, Table A1). Significant positive relationships
 308 were also observed between daily detritus GPP and daily seabed PAR (Table A1). Canopy light-use
 309 efficiency (LUE), estimated as the ratio between daily GPP and daily PAR (Attard and Glud, 2020),
 310 was 0.004 O₂ photon⁻¹ in June 2017, 0.006 O₂ photon⁻¹ in September 2017, and 0.004 O₂ photon⁻¹ in
 311 May 2018 (Table A1).



312
 313 Fig. 4: Hourly seabed PAR (a, b, c) with night-time periods grey shaded, hourly O₂ fluxes (d, e, f)
 314 and daily metabolism estimates of gross primary production (GPP), respiration (R), and net
 315 ecosystem metabolism (NEM) for the detritus canopy for the three measurement campaigns (g, h,
 316 i). Seabed PAR and O₂ fluxes are shown as mean ± 1 s.d. and are binned by the hour of day.

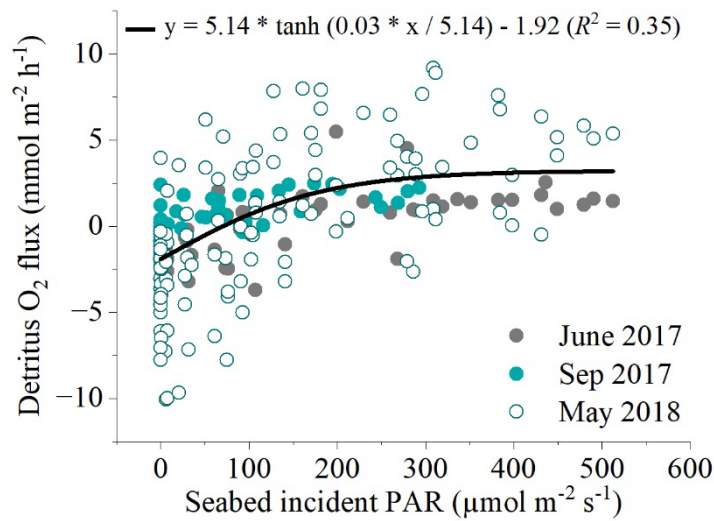
317 There was a significant positive relationship between near-bed incident PAR and the benthic O₂
 318 flux (Fig. 5). Light-saturation curves fitted to hourly data from all deployments indicated a
 319 maximum gross primary production rate (P_m) of $5.14 \pm 0.56 \text{ mmol O}_2 \text{m}^{-2} \text{h}^{-1}$, an α of 0.03 ± 0.01 ,
 320 and a R rate of $1.92 \pm 0.26 \text{ mmol O}_2 \text{m}^{-2} \text{h}^{-1}$. Light saturation (I_k) of the detritus canopy occurred at
 321 irradiances greater than $\sim 170 \mu\text{mol PAR m}^{-2} \text{s}^{-1}$.

322



323

324 Fig. 5: The daily balance between detritus gross primary production (GPP) and respirati®(R) for the
 325 three measurement campaigns. Symbol size corresponds to the daily integrated PAR reaching the
 326 seafloor.



327

328 Fig. 6: Relationship between all hourly in situ benthic O₂ fluxes at the detritus site and light
 329 availability from the three flux datasets measured. A modified photosynthesis-irradiance curve by
 330 Platt et al. (1980) is shown as the line-of-best-fit to the global dataset.

331

332 3.5. Macrobenthic diversity and abundance

333 The detritus site had a biomass of accumulated macrophyte (*F. vesiculosus*) detritus of 1666 ± 223
 334 g dry weight m⁻² (mean \pm SE, $n = 4$), approximately 1.5-fold larger than nearby intact *F. vesiculosus*

335 canopies (Table 2). Detritus accumulation in the five other habitats was around 100-fold smaller.
336 The area of the detritus site estimated using Google Earth was 2300 m², amounting to 3,800 kg dry
337 weight of *F. vesiculosus* fragments. Macrofauna abundance ranged from 2700 ± 900 ind. m⁻² at the
338 bare sediments site to 17300 ± 2400 ind. m⁻² at the sheltered *F. vesiculosus* site (mean ± SE, *n* = 4)
339 (Table 3). Macrofauna biomass ranged from 6 ± 2 g m⁻² at the bare site to 41 ± 9 g m⁻² at the
340 exposed seagrass site (mean ± SE, *n* = 4), and the number of species ranged from 6 to 23, with the
341 lowest values measured at the bare sediments and detritus sites, and the highest values at the
342 sheltered *F. vesiculosus* site (Table 3).

343 At the detritus site, there was a low abundance of epifaunal crustaceans when compared to other
344 habitats with canopies. Key species, such as the amphipod *Gammarus spp.* were notably absent, and
345 isopods such as *Idotea spp.* were present in low abundance (Table A3). Similarly, there was a
346 notable absence of bivalves such as the soft-shelled clam, *Mya arenaria*, and the cockle
347 *Cerastoderma glaucum*. Polychaetes such as *Hediste diversicolor* and *Marenzelleria spp.* were also
348 absent from the detritus site but present in other sedimentary habitats (Table A3). The nMDS
349 ordination of the macrofaunal assemblages indicated a clear separation of points representing the
350 different habitat sites (ANOSIM: R² = 0.865; *p* < 0.001). The assemblages from the bare sand and
351 the detritus sites formed separated site groupings compared to the vegetated sites ('*Fucus*' and
352 'seagrass', both exposed and sheltered). Within the vegetated sites, the assemblages of the 'seagrass
353 sheltered' and the '*Fucus* sheltered' sites were the most different (Fig. 7).

354

355 Table 2: Vegetation abundance and biomass (dry weight) at the six study sites. Abundance is shoots
 356 per m² for seagrass and individuals per m² for *F. vesiculosus*. Values are mean ± SE.

Site	Abundance per m ²	Above-ground biomass (g m ⁻²)	Belowground biomass (g m ⁻²)	Detritus (g m ⁻²)	Biomass other species (g m ⁻²)
Macrophyte detritus	-	-	-	1666 ± 223	-
Bare sediments	-	-	-	-	-
Sheltered <i>Z. marina</i>	768 ± 92	21 ± 2	8 ± 1	58 ± 13	0.1 ± 0.1
Exposed <i>Z. marina</i>	2565 ± 164	69 ± 7	25 ± 3	16 ± 2	0.2 ± 0.2
Sheltered <i>F. vesiculosus</i>	16 ± 2	1244 ± 58	-	55 ± 11	-
Exposed <i>F. vesiculosus</i>	16 ± 2	1112 ± 119	-	20 ± 2	-

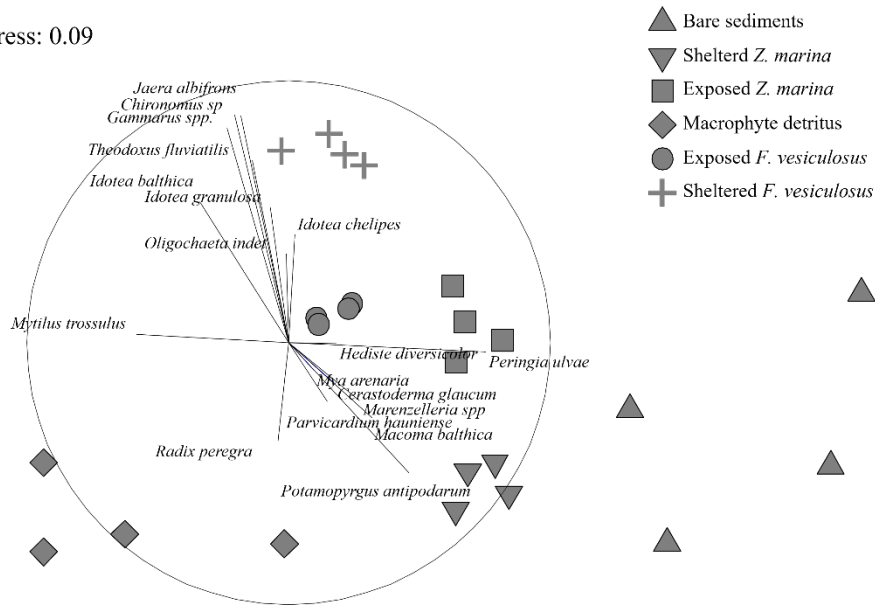
357

358 Table 3: Macrofauna abundance, biomass (ash-free dry weight), and number of species at the six
 359 study sites.

Site	Infauna abundance (ind. m ⁻²)	Epifauna abundance (ind. m ⁻²)	Total abundance (ind. m ⁻²)	Infauna biomass (g m ⁻²)	Epifauna biomass (g m ⁻²)	Total biomass (g m ⁻²)	Number of species
Macrophyte detritus	4175 ± 2885	493 ± 37	4668 ± 2885	5 ± 3	5 ± 0	9 ± 3	6
Bare sediments	2719 ± 854	-	2719 ± 854	6 ± 2	-	6 ± 2	6
Sheltered <i>Z. marina</i>	6110 ± 787	3020 ± 874	9130 ± 1176	30 ± 6	2 ± 0	33 ± 6	18
Exposed <i>Z. marina</i>	6959 ± 620	3316 ± 772	10275 ± 990	31 ± 8	10 ± 2	41 ± 9	16
Sheltered <i>F. vesiculosus</i>	-	17259 ± 2421	17259 ± 2421	-	11 ± 2	11 ± 2	23
Exposed <i>F. vesiculosus</i>	-	3551 ± 609	3551 ± 609	-	7 ± 2	7 ± 2	12

360

Stress: 0.09



361

362 Fig. 7: A non-metric multidimensional scaling (nMDS) ordination of the macrofaunal assemblages
363 indicated a clear separation of points representing the different habitat sites. The assemblages from
364 the bare sand and the detritus sites formed separate site groupings compared to the vegetated sites.
365 Data from May/June 2018 (see Table 1).

366 4. Discussion

367 4.1. Detritus metabolism

368 The eddy covariance measurements document a highly active detrital canopy that photosynthesized
369 as well as respired. High daily rates of GPP of up to $75 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$ and R of $100 \text{ mmol O}_2 \text{ m}^{-2}$
370 d^{-1} are comparable to some of the most productive habitats in the area, such as dense seagrass
371 meadows (*Zostera marina*) and intact canopies of bladder wrack (*Fucus vesiculosus*) (Attard et al.,
372 2019b). However, intact canopies of *F. vesiculosus* function very differently to detritus canopies
373 from a metabolic standpoint. In June 2017, two eddy covariance instruments were deployed in
374 parallel: one at the detritus site, and another at a nearby intact canopy. While the detritus was net
375 heterotrophic (NEM = $-15 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$; GPP:R = 0.76), the intact *F. vesiculosus* canopy was
376 strongly net autotrophic (NEM = $167 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$; GPP:R = 6.40) (Attard et al., 2019b). Daily
377 R at the detritus site was up to ~5-fold larger than that at a nearby (within 4 km) site with bare
378 sediments and up to twice as high as a neighbouring intact canopy of *F. vesiculosus* (Attard et al.,
379 2019b). Decaying (and respiring) fragments of *F. vesiculosus* could contribute substantially to the
380 O_2 uptake rate: laboratory incubations of *F. vesiculosus* fragments resolved respiration rates ~5

381 $\mu\text{mol O}_2 \text{ g dw}^{-1} \text{ h}^{-1}$, equivalent to $\sim 25 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$ when upscaled to *in situ* biomass observed at
382 the detritus site (data not shown).

383 Notwithstanding the key metabolic differences between detritus and other neighbouring sites, the
384 flux measurements (Figure 4) indicate that shallow detritus accumulation zones are not just regions
385 of organic matter remineralization, but rather they synthesize substantial amounts of organic matter
386 through primary production. The range in daily GPP:R from 0.53 to 0.81 indicates that primary
387 production can offset a substantial proportion of the respiratory demand, which extends the
388 persistence of detritus in the coastal zone. These observations are consistent with the laboratory
389 study by Frontier et al. (2021), who determined that following detachment, kelp (*Laminaria*
390 *hyperborea* and *L. ochroleuca*) fragments retain physiological and reproductive capabilities for up
391 to several months. Carbon retention within the coastal zone and export to deeper, sedimentary
392 accumulation regions would therefore be larger than would be predicted by decomposition theory
393 alone. Similarly, slow, and incomplete degradation of algae detritus under low $[\text{O}_2]$ conditions,
394 which could occur, for instance, in the bottom layers of detrital canopies or in the large anoxic
395 basins of the Baltic Sea (Conley et al., 2009), would increase carbon retention, transfer, and
396 sequestration potential (Pedersen et al., 2021).

397 4.2. Periodic benthic hypoxia

398 Our *in situ* measurements performed over a few days in late spring document that subtidal detritus
399 accumulation zones uniquely experience dynamic $[\text{O}_2]$ conditions driven by sunlight availability
400 and flow velocity, with rapid $[\text{O}_2]$ oscillations and frequent periods of hypoxia (Table 1). Hypoxic
401 conditions were largely restricted to the lower $\sim 5 \text{ cm}$ of the canopy and were only revealed by
402 sensors placed directly above the sediment surface ($< 5 \text{ cm}$ distance). At the onset of wave-driven
403 mixing, hypoxic waters from within the canopy propagated upwards into the water column and
404 were registered by fast-response $[\text{O}_2]$ sensors located 10 cm above the canopy ($\sim 35 \text{ cm}$ above the
405 seafloor). This observation suggests that the $[\text{O}_2]$ conditions inside the entire canopy and even in the
406 water column directly above can reach hypoxic conditions for a few minutes (Fig. 3). Such pulses,
407 however, were not registered by the slow-response $[\text{O}_2]$ optodes with a factory-specified $T_{90} < 2$
408 min. The minimum $[\text{O}_2]$ observed by these sensors placed at 10 cm and 35 cm above the seafloor
409 was 158 and $229 \mu\text{mol L}^{-1}$, respectively, and thus well above hypoxic conditions.

410 The importance of measuring $[\text{O}_2]$ close to the seafloor was demonstrated more than 40 years ago
411 by Jorgensen (1980), and since then, other researchers have investigated the distribution of

412 dissolved constituents such as O₂ and nutrients in the benthic boundary layer (Holtappels et al.,
413 2011). These studies document that solute gradients are largest near the seafloor. For practical
414 reasons, however, coastal monitoring programs measure [O₂] further away from the seafloor.
415 Models based on monitoring data suggest that hypoxia is prevalent in only small areas of the
416 shallow subtidal zone. For instance, models for the northern Baltic Sea, which cover a total seabed
417 area of 12435 km² of which 2211 km² is in shallow waters <5 m depth, indicate that just 16.5 km²
418 (or 0.75% of shallow waters) are prone to hypoxia (Virtanen et al., 2019). Given that large
419 quantities of drifting macrophytes are a common phenomenon in the shallow subtidal zone of the
420 northern Baltic Sea (Norkko and Bonsdorff, 1996a), it is likely that coastal hypoxia is currently
421 underestimated because large-scale models are largely based on measurements performed higher
422 above the seafloor (0.5-1.0 m) (Virtanen et al., 2019; Conley et al., 2011).

423 4.3. Biodiversity and [O₂] dynamics in detritus canopies

424 Despite being considered a temporary habitat, detritus was found in abundance at our study site on
425 all occasions in May, June, and September. This type of habitat is likely quite widespread in the
426 Baltic. Habitat distribution models for the area indicate a dominance of *F. vesiculosus* canopies in
427 shallow waters < 5 m depth (Virtanen et al., 2018), and these canopies are expected to export
428 substantial amounts of organic matter (~0.3 kg C m⁻² yr⁻¹) which can accumulate in topographical
429 depressions with limited water exchange (Attard et al., 2019a). Topographic depressions occupy
430 ~1350 km² or ~11% of the northern Baltic Sea (Virtanen et al., 2019). During a recent seasonal
431 study, we observed the highest abundance of detritus at our study site in summer and autumn,
432 coinciding with high southerly winds that erode intact canopies in shallower waters (Attard et al.,
433 2019a). However, we also observed significant canopy erosion in winter when a substantial biomass
434 of *F. vesiculosus* froze into sea ice and got dislodged once the ice broke up (Fig. 7). Therefore,
435 some degree of drifting detritus might be common throughout the year. Drifting detritus constitutes
436 a significant habitat structure. Given high enough biomass, however, detritus canopies can be a
437 challenging habitat for most species. At our study site, hypoxic conditions uniquely occurred at the
438 detritus site and for around a quarter of the deployment time (Table 1). We can expect these
439 conditions to be particularly challenging for crustaceans, the most hypoxia-sensitive
440 macroinvertebrate group (Vaquer-Sunyer and Duarte, 2008). Indeed, we only found one crustacean
441 species at this site- the isopod *Idotea balthica* (Table A3)- which is mobile and can tolerate hypoxic
442 conditions for a few hours (Vetter and Dayton, 1999). All other invertebrates observed at the
443 detritus site were mollusks (Table A3), the most hypoxia-tolerant marine invertebrate group

444 (Vaquer-Sunyer and Duarte, 2008). Other tolerant species include the blue mussel *Mytilus trossulus*
445 *x edulis* that can survive > 300 h of anoxia (Jorgensen, 1980), although the survival of larvae
446 depends on its developmental stage (Diaz and Rosenberg, 1995). Similarly, the mudsnail
447 *Peringia ulvae* is highly mobile and can survive > 150 h of anoxia (Jorgensen, 1980; Norkko et al.,
448 2000).

449 Overall, the dynamic [O₂] conditions in detrital canopies seem to be challenging for most species in
450 this region of the Baltic Sea, with lethal and non-lethal thresholds frequently being exceeded on
451 timescales of hours to days. We currently have a poor understanding of the extent of periodic
452 hypoxia in coastal waters, because [O₂] measurements are performed at some distance away from
453 the seabed. While this is a practical approach that is done to minimize sensor fouling and damage, it
454 does not reveal the full extent of coastal hypoxia. If implemented widely, sensor arrays, as
455 described herein, and sensor elevators (e.g. (Holtappels et al., 2011)) can fill in this knowledge gap
456 and provide important insights into the ecological status and biogeochemical cycling that is needed
457 for the sustainable management of coastal ecosystems.



458

459 Fig. 8: substantial detritus accumulation was observed in late winter (March 2021) when *F.*
460 *vesiculosus* froze into sea ice and got dislodged once the ice broke up. (Photo by Alf Norkko)

461 Table A1: A summary of the eddy covariance flux measurements performed on the detritus canopy
 462 during the three measurement campaigns. Daily integrated seabed PAR and detritus light-use
 463 efficiency (LUE, calculated as daily GPP/ daily PAR) are also presented.

Field campaign	Day	Daily GPP (mmol O ₂ m ⁻² d ⁻¹)	Daily R (mmol O ₂ m ⁻² d ⁻¹)	GPP:R	Daily PAR (mmol photons m ⁻² d ⁻¹)	LUE (O ₂ photon ⁻¹)
Jun 2017	1	62	83	0.74	13554	0.005
	2	54	71	0.76	11710	0.005
	3	29	35	0.81	9044	0.003
Sep 2017	1	15	26	0.57	3013	0.005
	2	37	69	0.54	4827	0.008
	3	26	48	0.53	3815	0.007
May 2018	1	46	59	0.77	10997	0.004
	2	57	74	0.76	12732	0.004
	3	74	97	0.76	13336	0.006
	4	32	41	0.78	10523	0.003
	5	39	51	0.77	10915	0.004

464

465

466 Table A2: Fit statistics for linear regressions performed between daily detritus GPP and R, and daily
 467 GPP and benthic PAR. Where relevant, values are presented \pm SE. The SE was scaled with the
 468 square root of the reduced Chi-Sqr. ANOVA was used to test slope significance. Asterisks indicate
 469 that the slope was significantly different from zero at the 0.05 level.

Relationship between daily GPP and daily R				
Field campaign	Slope of linear regression \pm SE	Intercept \pm SE	R^2	ANOVA Prob > F
Jun 2017	1.43 \pm 0.02	-5.91 \pm 0.77	0.99	0.01*
Sep 2017	1.93 \pm 0.06	-2.19 \pm 1.70	0.99	0.02*
May 2018	1.33 \pm 0.00	-1.09 \pm 0.17	0.99	0.00*
Global	1.16 \pm 0.13	9.90 \pm 5.92	0.89	0.00*
Relationship between daily GPP and daily PAR				
Field campaign	Slope of linear regression \pm SE	Intercept \pm SE	R^2	ANOVA Prob > F
Jun 2017	128 \pm 23	5293 \pm 1164	0.94	0.11
Sep 2017	82 \pm 4	1765 \pm 121	0.99	0.03*
May 2018	73 \pm 12	8103 \pm 609	0.90	0.01*
Global	182 \pm 40	1725 \pm 1852	0.66	0.00*

470

471

472 Table A3: Species list for the five studied sites. Presence is indicated by 'x'.

Group	Species	Macrophyte detritus	Bare sediments	Sheltered Z. marina	Exposed Z. marina	Sheltered F. vesiculosus	Exposed F. vesiculosus
Crustacea	<i>Amphibalanus improvisus</i>			x			
	<i>Asellus aquaticus</i>					x	
	<i>Corophium</i> spp.			x			
	<i>Gammarus</i> spp.			x	x	x	x
	<i>Idotea balthica</i>	x			x	x	x
	<i>Idotea chelipes</i>				x	x	x
	<i>Idotea granulosa</i>			x	x	x	x
	<i>Jaera albifrons</i>			x	x	x	x
	Cladocera					x	
	Copepoda					x	
	Ostracoda sp.					x	
	Mysid					x	x
Bivalvia	<i>Cerastoderma glaucum</i>			x	x		
	<i>Parvicardium hauniense</i>			x	x		
	<i>Macoma balthica</i>	x	x	x	x	x	x
	<i>Mya arenaria</i>			x	x		
	<i>Mytilus trossulus x edulis</i>	x		x	x	x	x
Gastropoda	<i>Peringia ulvae</i>	x	x	x	x	x	x
	<i>Radix</i> sp.	x		x			x
	<i>Potamopyrgus antipodarum</i>		x	x			
	<i>Theodoxus fluviatilis</i>	x	x	x	x	x	x

Polychaeta	<i>Hediste diversicolor</i>			X	X		
	<i>Halicryptus spinulosus</i>					X	
	<i>Maranzelleria</i> spp.		X	X	X	X	
	Nematoda					X	
	Oligochaeta			X	X	X	
	<i>Pygospio elegans</i>					X	
Others	<i>Chironomus</i> sp			X	X	X	X
	Coleoptera larvae						X
	Odonata						X
	<i>Cyanophthasma obscura</i>						X
	Hydrachnidae		X				X

473

474 **Author contribution**

475 All authors contributed significantly to designing the research, funding the study, collecting the
476 data, analyzing samples and data, and interpreting the results. KMA wrote the paper with input from
477 all authors.

478 **Competing interests**

479 The authors declare that they have no conflict of interest

480 **Data availability**

481 All data presented in this paper will be made available in a FAIR-aligned data repository upon
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