Environmental drivers of space-timespatio-temporal dynamics in floodplain vegetation: grasslands as habitat for megafauna in Bardia National Park (Nepal)

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Abstract: Disturbance-dependent grasslands, often associated with hydromorphological and fire dynamics, are threatened, especially in subtropical climates. In the Nepalese and Indian Terai Arc Landscape at the foot of the Himalayas, natural and cultural grasslands serve a viable role for rhinosgreater one-horned rhinoceros (Rhinoceros unicornis) and thefor grazers which form prey of the Royal Bengal tiger (Panthera tigris). The grasslands are vulnerable forto encroachment of forest. We aimed to establish the effects of environmental drivers, in particular river discharge, river channel dynamics, precipitation, and forest fires, on space timethe spatio-temporal dynamics of these grasslands. The study area is the floodplain of the eastern branch of the Karnali River and adjacent western part of Bardia National Park. We created two-annual time series (1993-2019) of land cover with the use of field data, remotely sensed LANDSAT imagery and a supervised classification model. Additionally, we analysed aerial photographs of 1964 and the pattern of grassland patches. From and aerial photographs of 1964. Between 1964 to and 2019, grasslands saw a transition to forest and grassland patches decreased in size and number abundance and size due to encroachment of forest. Outside the floodplain, successional setbacks conversion of grassland coincide to bare substrate coincides with extreme precipitation events. Within the floodplain, successional setbacksconversion of grassland correlateto bare substrate correlates with the magnitude of the annual peak discharge. However of the bifurcated Karnali River. Since 2009, however, this relationship correlation is absent after 2009 due to a westward shift of the main discharge channel to the western branch of the bifurcated Karnali River with a vast expansion of Consequently, alluvial tall grasslands (Saccharum spontaneum dominant) as consequence. Since 2009, hydromorphological processes in the floodplain have become more static. This is supported by an observed decrease in water coverage (53%) in the dry season, an absence of successional setbacks, vastly expanded between 2009 and decreased morphodynamics of river channels. For forest fires, the surface area that annually burns is observed to be more variable in recent years and the maximum extent affected by fires is in an increasing trend2019. Because the hydromorphological processes in the floodplain have become more static, other sources of disturbances - local flooding of ephemeral streams, anthropogenic maintenance, grazing and fires - are more paramount to prevent encroachment of grasslands. Altogether, our findings underscore that a change in the environmental drivers impact the surface area and heterogeneity of grassland patches in the landscape, which can lead to cascading effects for the grassland-dependent fauna megafauna.

1. Introduction

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Grasslands occur in diverse environmental conditions that can be dry, wet, hot, cold, productive, barren, dynamic and static (Brown and Makings, 2014). They have globally reduced in areal extent with 40% since the Industrial Era. A transition occurs from grassland to other types of land cover and land use due to cultivation, urbanisation and increased forest growth, which

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eausesTheir global areal extent has decreased by 40% since the Industrial Era (White et al., 2000) due to cultivation, urbanisation and increased forest growth, causing loss of biodiversity (White et al., 2000; Veldman et al., 2015). Grasslands are home to most of the extant mammalian megafauna on Earth and provide food for herbivores, which in turn sustain their predators. This is for example the case in the African (sub)tropical grasslands as well as in the higher and colder grassland habitats in Mongolia and Nepal. In these mountainous regions, grassland dependent Blue Sheep (Pseudois nayaur) and Siberian Ibex (Capra sibirica) serve as principle prey species of the snow leopard (Phantera unica) and wolf (Canis lupus) (Filla et al., 2021; Oli, 1996; Shehzad et al., 2012; Rovero et al., 2020). At the foot of mountain ranges, warmer and wetter conditions are present due to the lower altitude and the orographic lift of seasonal moist winds. The constraints for forest growth caused by temperature and water availability are mountainous and colder grassland habitats in Mongolia and Nepal. Here, grassland-dependent blue sheep (Pseudois nayaur), Siberian ibex (Capra sibirica) and Tibetan argali (Ovis ammon) serve as principle prey species of the snow leopard (Panthera unica) and Himalayan wolf (Canis lupus chanco) (Oli, 1996; Lyngdoh et al., 2014; Chetri et al., 2017; Rovero et al., 2020; Filla et al., 2021). At the foot of mountain ranges, warmer and wetter conditions are present due to lower altitudes and orographic lift of seasonal moist winds, causing the constraints for forest growth to be lifted (Staver et al., 2011; Hirota et al., 2011) (Staver et al., 2011; Hirota et al., 2011). The relationship between prey, predator and grassland habitat is also present in these regions more suited for forest growth: grassland patches in forests at the foot of mountainous regions in South Asia sustain Royal Bengal tigers (*Panthera tigris*) and their prey. This is for example the case in the Western Gaths, South India (Sankaran, 2009), in the Duars, Eastern India and Bhutan, and in the Nepalese and Indian Terai Arc Landscape (TAL), a subtropical belt at the foot of the Indian and Nepalese Himalayas (Harihar et al., 2014, Irengbam et al., 2017a). Also the TAL has seen a reduction of grassland in protected areas (Irengbam et al., 2017).

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The TAL encompasses a subtropical belt at the foot of the Indian and Nepalese Himalayas where it is also the northern border of the Ganges sedimentary basin. Settlements and agricultural fields alternate with subtropical forests with distinctive grasslands and dynamic rivers draining the Himalayas. The grasslands can broadly be subdivided into 1) the natural alluvial tall grassland, located near active and in abandoned river channels, and 2) a mosaic of mixed tall grasslands and short grasslands (*phantas*) located further away from the river channels, of which a number of grassland patches are thought to be remnants of anthropogenic interferences (Peet et al., 1999a). Across the TAL, these grasslands have an ecologically important role for indigenous flagship species. The tiger density is highest on grasslands and these grasslands provide feeding grounds for Chital deer (*Axis axis*), the most important prey (DNPWC and DFSC, 2018). Next to the areal extent, the spatial pattern of grassland in the landscape is of importance: habitat heterogeneity is positively correlated with the principle prey for tigers in the National Park of Chitwan (Nepal) (Bhattarai and Kindlmann, 2012). For rhinos (*Rhinoceros unicornis*) and hog deer (*Hyelaphus porcinus*) the alluvial tall grasslands serve as an essential habitat (Jnawali and Wegge, 2000; Thapa et al., 2013; DNPWC and DFSC, 2018). It is thought that if changes occur in the composition and area of tall grassland that the abundance and distribution of these animals is greatly affected (Odden et al., 2005; Peet, 1997). It is of ecological concern when megafauna is extirpated not only for the megafauna itself as the megafauna influences the community structure of their habitat, too (Wikramanayake et al., 1998).

Grasslands in the TAL can broadly be subdivided into (1) natural alluvial tall grassland, located near active and in abandoned river channels, and (2) a mosaic of short grassland (phantas) and mixed tall grassland, located further away from the river channels. The alluvial tall grasslands serve as essential habitat for the greater one-horned rhino (*Rhinoceros unicornis*) and hog deer (*Hyelaphus porcinus*) (Jnawali and Wegge, 2000; Odden et al., 2005; Pradhan et al., 2008; Thapa et al., 2013; DNPWC and DFSC, 2018) and their numbers would be greatly affected if changes occur in composition and areal extent of

these grasslands (Odden et al., 2005; Peet, 1997). Tiger density is highest near grasslands and particularly the short grasslands provide nutritious feeding grounds for Chital deer (*Axis axis*), the most important prey (Moe and Wegge, 1994; DNPWC and DFSC, 2018). In addition to the surface area, the spatial pattern of grassland in the landscape is of importance: habitat heterogeneity is positively correlated with the abundance and habitat preference of Chital deer in the National Park of Chitwan (Nepal) (Bhattarai and Kindlmann, 2012).

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A number of grasslands in the TAL are considered to be maintained via physical disturbances, preventing that prevent succession to mature vegetational stages (Peet et al., 1999b; Lehmkuhl, 1994). (Seidensticker, 1976; Dinerstein, 1979a; Lehmkuhl, 1989, 1994; Peet et al., 1999b). In general, the pattern of vegetation is regulated by autogenic and allogenic factors (Tilman, 1988), of which the latter are either environmental or anthropogenic. The Allogenically, the community dynamics can be controlled allogenically via disturbances, of which the intensity and spatial extent arehave been identified as principal factors (Turner et al., 1998). The natural disturbances recorded in the TAL include hydromorphological processes, fires, climatic variables (precipitation, temperature, droughts) and herbivory, whereas the anthropogenic disturbances include cutting of grasses and trees for vegetation maintenance and resource collection, clearance by fires, eattlelivestock grazing, and, in the past, cultivation and creation of settlements (Seidensticker, 1976; Lehmkuhl, 1994, 1989; Dinerstein, 1979a). In disturbancedependent grasslands, changes in rainfall, fire and herbivory can cause forestation of grassland within several years (Seidensticker, 1976; Dinerstein, 1979a; Lehmkuhl, 1989, 1994). Changes in disturbance regimes impact other disturbance regimes (Van Langevelde et al., 2003; Trauernicht et al., 2013; Orem and Pelletier, 2015), the vegetation pattern, and thus the ecosystem as a whole (Parr et al., 2012). Conversion of disturbance-dependent grassland to forest can occur within several years when changes occur in rainfall, fire or herbivory (Bond, 2008). Changes in disturbance regimes have impact on other disturbance regimes and on the vegetation pattern, and thus the ecosystem as a whole (Parr et al., 2012). This susceptibility of land cover to changes in disturbance regimes not only holds for the grasslands in the TAL but also for other ecosystems where climate is conductive for forest growth. In the Brazilian Pantanal, where similar grasslands are found at the foot of the Brazilian Highlands, extreme hydrologic conditions are explicitly analysed, with the use of remote sensing, In Bardia National Park (BNP), a conservation area in the TAL located near the Karnali megafan in Western Nepal, grasslands have been declining due to encroachment of forest (Peet et al., 1999a; Jnawali and Wegge, 2000; Odden, 2007). From a hydrological perspective, there are indications that a major change occurred during the monsoon in 2009 (Sinclair et al., 2017): near the apex of the highly dynamic braided river system of the Karnali River, the dominant discharge channel relocated from its eastern branch (the Geruwa River that borders BNP) to the western branch (Kauriala River). The susceptibility of land cover to changes in disturbance regimes not only holds for the grasslands in the TAL but also for other ecosystems where climate is conductive for forest growth. In the Brazilian Pantanal, where similar grasslands are found at the foot of the Brazilian Highlands, extreme hydrologic conditions have been explicitly identified with remote sensing as a driving factor of the pattern of vegetation communities (Arieira et al., 2011). The Taquari megafan in the Pantanal experienced a shift in hydrological conditions due to an avulsion, increasing susceptibility to deforestation and fire (Louzada et al., 2020).

One of the conservation areas in the TAL where the grasslands are under threat is Bardia National Park (BNP), located near the Karnali megafan in Western Nepal. In earlier times, grasslands were more widespread in BNP and the area of these early successional habitats in the region has been declining by encroachment of wooden plants (Peet et al., 1999a; Jnawali and Wegge, 2000; Odden, 2007). From a hydrological perspective, there are indications that a major change occurred during monsoonal floods in 2009 (Sinclair et al., 2017): near the apex of the highly dynamic braided river system of the Karnali River, the dominant discharge channel relocated from its eastern branch, the Geruwa River that borders BNP, to the western branch

(Kauriala River). The current distribution of discharge is considered to be about 80% for the Kauriala river and 20% for the Geruwa River during low discharges (Sinclair et al., 2017; Dingle et al., 2020a) and the distribution is thought to be higher in the western branch (55 65%) during peak flow for monsoonal discharges (Dingle et al., 2017, 2020a). Reduced fluvial dynamics in the Geruwa River could possibly favour higher successional stages of vegetation in and near the floodplains at the western boundary of BNP.

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Wide ranging progress has been made in understanding the relation between drivers and space time dynamics of vegetation, including on the dynamics and pattern of riparian vegetation (e.g. Hupp and Osterkamp, 1996; Lorenz et al., 1997; Corenblit et al., 2007; Vesipa et al., 2017). The composition and spatial distribution of riparian vegetation communities is heavily influenced by the flood frequency, duration, intensity (Hupp and Osterkamp, 1996; Poff et al., 1997) and timing (Newbold and Mountford, 1997). Consequences of high discharges for vegetation can be physical damage and uprooting, anoxia due to prolonged inundation and burial by fresh alluvium. Low discharges can adversely affect vegetation with drought stress when the ground water table becomes too deep. These processes, where the three main elements of interaction are vegetation, water, and sediment flow, are integrated in concepts such as 'fluvial biogeomorphic succession', describing the reciprocal interactions between fluvial landforms and vegetation (Gunderson, 2002; Corenblit et al., 2007), 'flood pulse concept', describing lateral connectivity recognizing that flow and its variability are the main drivers of ecological processes in floodplains (Junk et al., 1989) and 'shifting mosaic steady state' for vegetation, which specifies the proportion of distinct successional land cover classes that remains relatively constant in a river reach or corridor (Arscott et al., 2002; Kollmann et al., 1999).

Additionally, advances have been made in understanding other environmental drivers, such fire dynamics in grasslands (Leys et al., 2018; Buisson et al., 2019; Hoetzel et al., 2013; Iglesias et al., 2014; Flannigan and Wotton, 2001), as well as the role of anthropogenic interferences and restoration pathways of (sub)tropical grasslands (Buisson et al., 2019). Nonetheless, the interrelations of environmental drivers are complex (Lehmann et al., 2014). There is not yet a complete understanding of the relations between the various environmental drivers and the effect on the vegetation pattern in space and time. It is, therefore, of interest to study vegetation and its drivers at a high temporal resolution over a long time span and in its historic context.

More region and context specific, linking annual land cover dynamics to their drivers in combination with long term data on environmental drivers was tackled less frequently and literature is particularly limited to a small number of geographic regions (Dufour et al., 2019). This includes the subtropical grasslands in the TAL such as in BNP, for which specifically no published work is available that evaluated or monitored grasslands with (annual) temporal series and in relation to environmental drivers and variation therein. It is especially vital to understand the functioning of such a valuable ecosystem in relation to possible changes in drivers (such as the indicated redistribution of discharge in 2009) in order to provide insight for nature management strategies aiming to conserve grasslands and associated fauna, such as the abundance of tiger's prey and the endangered tiger itself (Harihar et al., 2014), and also to expand on the knowledge for similar systems where grasslands sustain prey predator relationships.

Wide-ranging progress has been made in understanding the drivers of vegetation dynamics through space and time (Hupp and Osterkamp, 1996; Lorenz et al., 1997; Vesipa et al., 2017; Corenblit et al., 2007; Turner et al., 1998). The composition and spatial distribution of riparian vegetation communities are heavily influenced by flood frequency, duration, intensity (Hupp and Osterkamp, 1996; Poff et al., 1997) and timing (Newbold and Mountford, 1997). Consequences of high discharges for vegetation can be physical damage and uprooting, anoxia due to prolonged inundation, and burial by fresh alluvium. Low

160 discharges affect vegetation with drought stress when the ground water table becomes too deep. These processes are integrated in concepts such as 'fluvial biogeomorphic succession', describing the reciprocal interactions between fluvial landforms and vegetation (Gunderson, 2002; Corenblit et al., 2007), 'flood pulse concept', describing lateral connectivity and recognizing that flow and its variability are the main drivers of ecological processes in floodplains (Junk et al., 1989), and 'shifting-mosaic steady state' for vegetation, which specifies on the proportion of distinct successional land cover in floodplains (Arscott et al., 165 2002; Kollmann et al., 1999). Additionally, fire dynamics (Flannigan and Wotton, 2001; Hoetzel et al., 2013; Iglesias et al., 2014), herbivory (Owen-Smith, 1988; Van Langevelde et al., 2003; Allred et al., 2011) as well as anthropogenic interferences (Wang et al., 2015) have been studied. The interrelations of environmental drivers are complex (Lehmann et al., 2014) and there is not yet a complete understanding. It is, therefore, of interest to study vegetation and its drivers at a high temporal resolution over a long time span and in its historic context. Literature in this domain is particularly limited to a small number 170 of geographic regions (Dufour et al., 2019) and the subtropical grasslands in conservation areas in the TAL such as BNP are underrepresented. It is especially vital to understand the functioning of such a valuable ecosystem in relation to possible changes in drivers (such as the indicated redistribution of discharge in 2009) in order to provide insight for nature management strategies.

175 A promising method for studying land cover dynamics and its drivers is the use of earth observation techniques, which enable diachronic analysis-contributing valuable data and insights on the development of land cover (Lallias-Tacon et al., 2017; Dufour et al., 2019; Harezlak et al., 2020; Solins et al., 2017; Basumatary et al., 2021; Louzada et al., 2020; Van Iersel, 2020; Corenblit et al., 2010; Dufour et al., 2015). With remotely sensed imagery from satellites, vegetation vegetation in and near floodplains can be mapped at the following scales and characteristics: vegetation typestype (Alaibakhsh et al., 2017), species 180 composition (Rapinel et al., 2019; Plakman et al., 2020)(Rapinel et al., 2019; Plakman et al., 2020), physiological processesphysiology (Wagner-Lücker et al., 2013) and vegetation-structure (Straatsma and Baptist, 2008; Jalonen et al., 2015). Regarding delineation of vegetation types in conservational areas similar to BNPRegionally, remote sensing washas been successfully used for monitoring (moist) to monitor similar grasslands, with either single moments in time to map the distribution of land cover for habitat evaluation (Thapa, 2011) or diachronous maps to be able to detect changes in vegetation 185 (Biswas, 2010; Acharya, 2002; Sarma et al., 2008). Regionally, remote sensing was also used to study fire dynamics and occurrences(Thapa, 2011; Biswas, 2010; Acharya, 2002; Sarma et al., 2008), fires (Takahata et al., 2010; Ghimire et al., 2014) and hydromorphological processes such as shifting of the channels channel migration of the Karnali River (Rakhal et al., 2021).

We aim to quantify spatio-temporal land cover change and establish relationships with the environmental drivers with focus on the ecologically important grasslands. We address both grasslands within the floodplain of the Geruwa River and away from the Karnali River where hydromorphological processes occur on a smaller scale due to ephemeral streams. Specifically, we give answers to the subquestions: (a) what is the spatio temporal pattern of land cover; (b) what is the temporal (and spatial) variation in environmental drivers of the last three decades; and (c) what are the effects of environmental drivers on land cover change and what are possible mechanisms that explain these effects. We do so by mapping the historic development including recent annual dynamics of land cover with the use of field and remotely sensed data (LANDSAT) combined in a supervised land cover classification algorithm (*Random Forest*). We then relate the land cover dynamics to extremes in hydrological, meteorological and forest fire variables and discuss the mechanisms. The hypothesis is that from 1964 to 2019 a shift occurred from grasslands to later successional stages such as riverine and Sal forest due to the establishment as conservation area and the more recent avulsion in 2009. This is based on literature (Dinerstein, 1979a; Peet et al., 1999a; Odden, 2007), the consequences of avulsions in similar nature reserves (Louzada et al., 2020; Biswas, 2010; Sarma et al., 2008) and experiences

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of staff of BNP. On small time scales, years that experienced extreme hydrologic conditions are regarded as important incidents for removal of grass and riverine forest in the floodplain, and in the years thereafter an increase in alluvial grasslands is expected following the proposed successional trajectory of vegetation for the park (Dinerstein, 1979b; Lehmkuhl, 1989; Peet, 1997). Therefore, the discharge distribution We address both grasslands within the floodplain of the Geruwa River (the eastern branch of the Karnali River) and the grassland patches further away. Specifically, we give answers to the subquestions: (a) what is the spatio-temporal pattern of land cover; (b) what is the spatio-temporal variation in environmental drivers during the last three decades; and (c) what are the effects of environmental drivers on land cover change and what are possible mechanisms that explain these effects? We do so by mapping the historic development including recent annual dynamics of land cover with the use of field and remotely sensed data (LANDSAT). We then relate the land cover dynamics to abrupt changes in hydrological, meteorological and forest fire variables. The hypothesis is that between 1964 and 2019 a transition occurred from grasslands to later successional stages. The discharge distribution during peak discharges at the upstream bifurcation of the Karnali River would be a relevant driver for the vegetation pattern and associated habitats of megafauna.

2. Study area

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2.1 Bardia National Park and Karnali River

Bardia National Park, located in the southwest of Nepal, contains the only protected floodplain of the Karnali river in Nepal (Fig. 1). The park, a former hunting reserve, was established as the Royal Karnali Wildlife Reserve on 8 March 1976, holding an area of 368 km² at that time. It was expanded with the Babai River valley to the east in 1984, and received the status of National Park in 1988 (Brown, 1997). Bardia National Park has a core zone area of 968 km², surrounded by a buffer zone of 507 km² falling in IUCN category II (DNPWC and DFSC, 2018). Here, the western part of the park is studied with an area of 135 km², for its accessibility, presence of grasslands and location of the Karnali river. The park holds 82 97 tigers (DNPWC and DFSC, 2018), more than 100 elephants (Shrestha and Shrestha, 2021) and 38 rhinos according to the rhino consensus in 2021.

The study area is delimited by: Bardia National Park is located in the southwest of Nepal (Fig. 1) and was established as conservation area on 8 March 1976. The study area (135 km² of the now in total 968 km² of BNP) is delimited by (1) the Siwalik Hills in the north, where ephemeral streams originate that are tributaries to the Karnali River; (2) buffer zone and settlements in the southwest, where the park office and a field station of the Nepalese National Trust for Nature Conservation (NTNC) are located southeast; and (3) the Geruwa branch Branch of the Karnali River in the west; a highly dynamic braided river system (Sinclair et al., 2017) located on the Karnali megafan (USAID, 2018). The Karnali River rises from Mt. Kailash on the Tibetan Plateau and isas one of the three biggest rivers draining Nepal. It bifurcates a few kilometres south of the Siwalik Hills near the town of Chisapani (Fig. 1). The eastern branch is 1) the Karnali River bifurcates into the Geruwa River and the (eastern branch) and Kauriala River (western branch is the Kauriala River,), joining together in India as a tributary to the Ganges riverRiver. The subtropical climate in the region encompasses three distinct seasons: the monsoon, cool and dry post-monsoon, and a hot and dry pre-monsoon (USAID, 2018). Temperatures reach a maximum of 45°C in the hot season and fall to 10°C in January (Bolton, 1976). The mean annual precipitation is 1560 mm (DHM, 2017). The mean annual precipitation is 1560 mm (DHM, 2017).

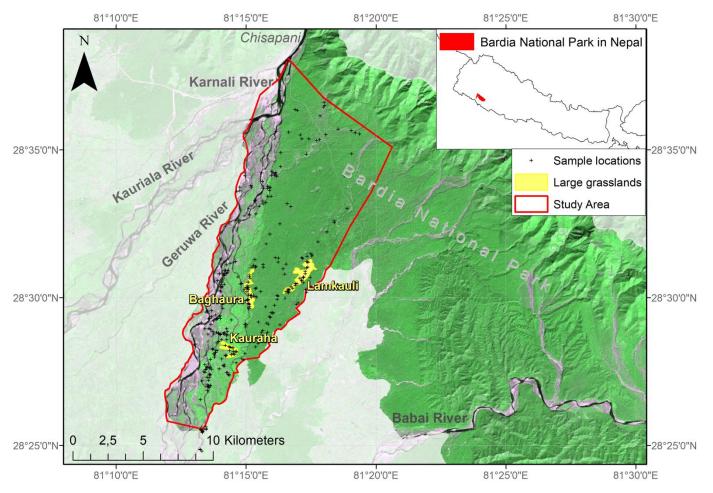


Figure 1: Delineation of Bardia National Park (dark green) and the study area (red) in and near the floodplain of the Karnali River.

The megafan of the Karnali River lies between the two branches (the Kauriala and the Geruwa River) at the west side of Bardia National Park. This is a false colour composite (RGB = 656 of LANDSAT 8 imagery, 2020).

2.2 Drivers of grasslands in the Terai Arc Landscape and Bardia NP

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In the TAL, natural and anthropogenic disturbances are regarded as important drivers for grasslands (Seidensticker, 1976; Lehmkuhl, 1994, 1989; Dinerstein, 1979a; Peet et al., 1999a). Drivers in the form of physical disturbances vary in type, temporal components (frequency, duration, and timing), their spatial component (location, extent) and their intensity. Fluvial processes are regarded to be important in shaping the spatial and temporal patterns of vegetation in Terai grasslands at the landscape level (Seidensticker, 1976; Lehmkuhl, 1989). Flooding of the Karnali River and associated clearance of vegetation together with deposition of fresh alluvium have been observed in Bardia National Park (Dinerstein, 1979a; Peet et al., 1999b; Lehmkuhl, 1994). Another type of disturbance that (sub)tropical grasslands often depend on are herbivory and trampling (Buisson et al. 2019), which are also relevant for BNP. A combination of fire and grazing pressure results in nutritious, very short grazing lawns on the *phantas*, and are thought to be a consequence of patch selective grazing after burning or cutting (Thapa et al., 2021). Larger herbivores such as rhinos and elephants are considered not to be of sufficient density to be deemed responsible for creating and maintaining grassland patches due to their low numbers in the park (Thapa et al., 2021). Fires, an integral ecological process (Flannigan and Wotton, 2001), also for subtropical grasslands (Ratnam et al., 2019; Sankaran, 2005), occur in BNP on a yearly basis, and can be natural as well as anthropogenic of origin. Forest fires are most frequently

observed in the Sal forest (Ghimire et al., 2014), while for grasslands fire is used annually to prevent encroachment, especially for the phantas (Peet, 1997; Dinerstein, 1979b).

Alongside man lit fires, BNP experienced anthropogenic influences for a long time (Bhatta, 2000; Bolton, 1976). Table 1 provides an overview of historic human activities in BNP as recorded in literature (see also Thing et al., 2017). Anthropogenic disturbances can be subdivided in disturbances by locals (cutting of thatch and grazing of cattle) and by park staff (maintenance purposes), detailed upon by Bhatta (2000). After establishment as conservation area in 1976, free access for locals, who extensively relied on BNP for resources, was refused. Later on active habitat management was initiated to compensate for the declined disturbances by locals, particularly carried out at the Khauraha, Baghaura and Lamkauli *phantas* (Fig. 1), which are short grasslands (< 2 meters) that were cultivated before the park was established and have been actively managed since to disrupt succession. Permits that allow resource collection are needed for locals to enter the park. Through the years, the number of permits at first increased threefold until 1999, whereas the number of days that the park is accessible for resource collection at first increased and after 1994 decreased (Bolton, 1976; Bhatta, 2000). At present, the number of permits drastically decreased and the park is accessible for 3 days (Thapa et al., 2021). Whether the amount of permits in a given year is proportional to the amount of removed biomass is not known (Bhatta, 2000).

Table 1: Overview of anthropogenic activities in BNP as recorded in literature

Year	Activity	Description	Source				
Drivers in	the form of physica	l disturbances vary in type, temporal c	omponents (frequency, duration, and timing), their spatial				
componen	t (location, extent) a	and their intensity (White, 1985; Turne	r et al., 1998). In the TAL, both natural and anthropogenic				
disturbanc	disturbances are regarded as important drivers for grasslands in shaping the spatial and temporal patterns of vegetation. In						
BNP, fluvial process are regarded to be important at the landscape level. Flooding of the Karnali River and associated clearance							
of vegetat	of vegetation together with deposition of fresh alluvium have been observed (Seidensticker, 1976; Dinerstein, 1979a;						
Lehmkuhl	, 1989, 1994; Peet e	et al., 1999b). Fires occur in BNP on a	yearly basis, and can be natural as well as anthropogenic.				
Forest fire	s occur most freque	ently in the Sal forest (Shorea robusta	dominated) and in the months April and May (Ghimire et				
al., 2014).	The phantas are ann	nually set on fire by park staff (Dinerste	in, 1979b; Peet, 1997). Appendix A provides an overview				
from the li	iterature of the activ	rities by locals and park staff (see also	Thing et al., 2017 and Bhatta, 2000). After establishment				
as conserv	ration area in 1976,	free access for locals, who extensively	relied on BNP for resources, was refused. Later on, active				
habitat ma	nagement was init	ated on grasslands to compensate for	the decreased disturbances. Permits that allow resource				
collection	are needed for loca	ls to enter the park for a certain number	er of days. Since 1983, the days of access and the number				
of permits	issued increased un	til 1994 and 1999, respectively, and he	reafter decreased drastically (Bolton, 1976; Bhatta, 2000).				
Whether t	he amount of permi	ts in a given year and the days of acce	ess are proportional to the amount of removed biomass is				
not knowr	(Bhatta, 2000). Be	fore and around 1976, extensive grazi	ng was recorded by cattle and wild ungulates in the south				
west of th	e study area. The	very short grazing lawns that are pres	ent in BNP are a result from patch-selective grazing by				
ungulates	after cutting or bur	ning (Thapa et al., 2021). Large herb	vores can strongly influence the pattern and community				
structure o	of vegetation (Owe	n-Smith, 1988; Dinerstein, 1992). In	the park, Asian elephants (Elephas maximus) (Ram and				
Acharya, 2	2020; Shrestha and	Shrestha, 2021) and rhinoceros (rhino	census in 2021) were virtually absent, but their numbers				
increased	to 113 and 38 in 20	021, respectively. Ungulates have been	n censused a number of times in BNP (Dinerstein, 1980;				
Wegge an	d Storaas, 2009; Dł	nakal et al., 2014; Karki et al., 2016; K	ral et al., 2017; DNPWC and DFSC, 2018, 2022). Some				
uncertaint	y is present around	the numbers of ungulates, which hav	e been debated in literature (van Lunenburg et al., 2017;				
Wegge et	al., 2019). In BNP,	the recorded densities of Chital deer (t	ne most abundant mid-sized grazer) are: 53.99 (SE 10.29)				

(Dhakal et al., 2014), 29.3 (SE 4.3) (Karki et al., 2016), 56.44 (SE 5.75) (DNPWC and DFSC, 2018), 44.05 (SE 6.3) (DNPWC and DFSC, 2022).

1925	Commercial forestry	5 years of extensive deforestation.	Bolton (1976)	
Since 1950	Increase of population	Increased deforestation and pressure on forest.	Brown (1997) and Bhatterai et al. (2017)	
1965-1975	Possible cultivation of Baghaura and Lamkauli phantas	Oral records deliver cultivation of these phantas.	Dinerstein (1979a) and Pokharel (1993)	
1970-1980	No deliberate management.	It was considered sufficient to exclude disturbances (except from fires and hunting expeditions) from protected areas in order to preserve biodiversity. In most protected area's the principle of 'nature balances itself' was strictly followed.	Bhattarai et al. (2017)	
1976	Free access for locals refused and livestock grazing was prohibited.	Establishment as conservation area.	Brown (1997)	
1979	Resource collection allowed, 7 days of access	Local communities were granted the rights to collect thatch grass and reeds from the reserves once a year, which was a pioneer step towards a people centered approach.	Brown (1997)	
1979-1983	Relocation of settlements	572 families were relocated from the Babai Valley of BNP.	Brown (1997)	
1983	Registered permits: 21,081; 15 days of access	Permits were issued for entering the part for resource collection.	Bhatta (2000)	
1994	Nr. of days for resource collection reduced to 10 days	=	Bhatta (2000)	
1995	Start of uprooting at the Khauraha and Baghaura phanta	Uprooting of unpalatable species (<i>Lantana sp.</i> and <i>Colebrookia sp.</i>) as part of the Bardia Integrated Conservation Project (1995-2001).	Bhatta (2000)	
1999	Registered permits: 57,255 Extensive uprooting of small bushes and trees at Khauraha phanta	Gradual increase of permits from 21k (1983) Bhatta (2000) o 57k (1999)		
2020	Nr. of days for resource collection is 3 days, drastical decrease of permits (number unknown)	-	Thapa et al. (2021)	

300 2.3 Vegetation types and successional stages

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Dinerstein (1979a) described six major vegetation types in the park, later modified by Pokharel (1993) into seven types: into seven types (Jnawali and Wegge, 1993; Pokheral, 1993): alluvial tall grasslands (*Saccharum spontaneum* dominant), mixed tall grasslands (wooded grasslands), short grasslands ((< 2 m, previously cultivated fields or 'phantas'), Khair-Sissoo forest, moist riverine forest, mixed hardwood forest and Sal forest. More detailed grassland associations have been delineated in literature and are clarified in Appendix A and are grouped as the three major grassland types used in this study (Lehmkuhl, 1989, 1994; Peet et al., 1999a; Dinerstein, 1979a).

The alluvial tall grasslands, which can be up to 4 metersm in height, containconsist of pioneer species and that quickly cover bare alluvium after disturbance events. With no disturbances present, the line of succession is from alluvial tall grasslands via Khair-Sissoo forest to other riverine forest types, mixed hardwood forest or the climax Sal forest. Khair-sissoo forest is prevalent along the river courses and on floodplain islands in the Terai due to its ability to withstand floods. Sal forest consists of predominantly *Shorea robusta* trees that are fire resistant and is the most dominant tree species in BNP. Disturbances prevent

succession of alluvial grasslands to forest, turning them into mixed tall grasslands. If disturbances continue, short and open grasslands may appear (locally called *phantas*) (Dinerstein, 1979a; Lehmkuhl, 2000). (< 2 m) and open grasslands may appear (locally called *phantas*) and their location is often more distal from stream channels (Dinerstein, 1979a; Lehmkuhl, 2000). A number of these short-grasslands (< 2 meters) are considered to owe their existence to human interferenceinterferences such as the Lamkauli, Bagaura and their location is more distal from the stream channels (Fig. 1) (Pokheral, 1993; Dinerstein, 1979b; Peet et al., 1999a). The largest phantas present are Lamkauli, Bagaura and Khauraha (Fig. 1). (Pokheral, 1993; Dinerstein, 1979b). Nowadays, they are predominantly vegetated with *Imperator Cylindrica* (Peet et al., 1999b). The largest phantas present are Lamkauli, Bagaura and Khauraha (Fig. 1). The origin is not clear for every grassland. The mixed tall grasslands (> 2 meters) are often dominated by *Narenga porphyrocoma* and *Erianthus ravennae* and are considered to lie on a continuum with short grasslands. Within the mosaic of short and mixed tall grasslands, very short grazing lawns are present (Thapa et al., 2021). In the mixed tall grasslands (> 2 m) the *Narenga porphyrocoma* and *Erianthus ravennae* species are abundant and these grasslands are considered to lie on a continuum with the short grasslands (Peet et al., 1999b).

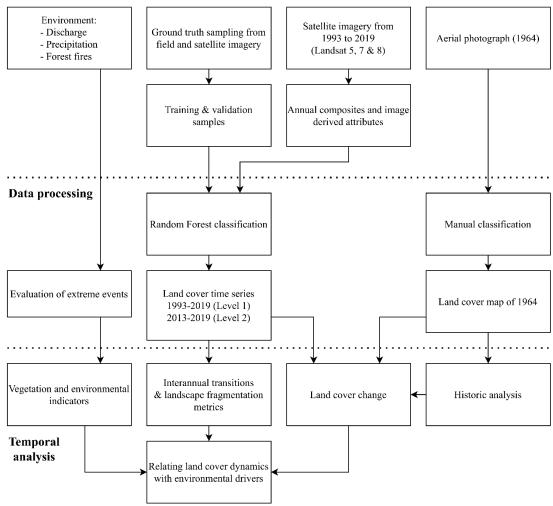
3. Data and Methods

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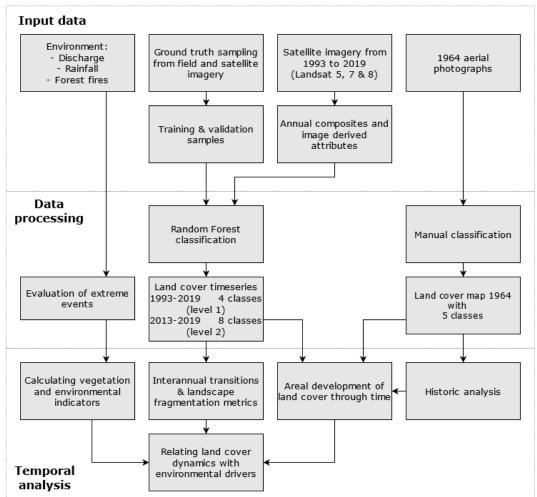
3.1 Outline of approach

Figure 2 provides a flow chart of the methodology. We used topographic maps of (1927, 1984) and aerial imagery of 1964 to assess the historic development of vegetation in the park and used LANDSAT imagery for reconstruction of an annual time series of land cover frommaps between 1993 to and 2019. During the post-monsoon of 2019, Bardia National ParkBNP was visited to collect ground truth data of vegetation types. This dataset was used for supervised classification of land cover using a Random Forest model for creating the land cover time series. From this time series, the development of the area, the areal extent, transitions and the pattern of the vegetation classes can be land cover were calculated. These are related to environmental variables with focus on extreme (precipitation and, discharge events and satellite observations from forest fires.

Input data



335 <u>fire occurrences</u>) to provide insight in the dynamics of land cover and the response to variation in environmental variables.



-Figure 2: Flowchart showing steps from data sources to analysis.

340 3.2 Satellite imagery and aerial photographs

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We selected imagery of the LANDSAT 5, 7 and 8 satellites from 1993 to 2019 to establish an annual time series of land cover. We used the Surface Reflectance dataset (Tier 1) of LANDSAT as proposed by Young et al. (2017) when the earth surface is compared at different moments in time and when multiple scenes are used for creation of composites. To establish a time series of yearly land cover, we selected imagery of the Surface Reflectance dataset (Tier 1) of LANDSAT 5, 7 and 8 satellites between 1993 and 2019. This imagery has a resolution of 30 m and the full pre-processing of the images is elaborated in Appendix B. The composites of 1995, 1997 and 2006 were not classified. Too few images of sufficient quality were available for 1995 and 1997 and a too large number of erroneous pixel values in the image composite was observed for 2006, caused by the LANDSAT 7 Scan-line error. For each year we calculated NDVI (Rouse et al., 1973; Rock et al., 1986; Myneni et al., 1995; Beeri et al., 2007), NDDI (Gu et al., 2007) This dataset of imagery is corrected for variation of the energy source (e.g. sun angle) and atmospheric effects (e.g. aerosol scattering and thin clouds) (Zanter, 2019). We accounted for the differences in sensors to provide continuity in data. Coefficients provided by Roy et al. (2016) were used to correct LANDSAT OLI imagery to LANDSAT 7 ETM+ imagery. Before computing the seasonal composites, imagery was selected based on the day of the year (between 240 366 of the post monsoon and 1-150 of the subsequent pre-monsoon), cloud cover (< 50 %), quality (> 9), geometric RMSE (Root Mean Square Error) of < 10 m and we performed cloud masking using the Fmask algorithm (Zhu et al., 2015).

For each annual land cover map, two separate seasonal image selections were made of the cool post monsoon and the dry premonsoon. The incorporation of seasonal information improves the classification results—as the spectral signatures of the
separate classes are more distinct due to phenological differences of vegetation throughout the year (Van Iersel et al., 2016;
Kelley et al., 2018). A composite was created from each seasonal selection. This was done by taking the median value of each
pixel to get rid of outliers. To enhance the available spectral information, NDVI (Rock et al., 1986; Beeri et al., 2007; Myneni
et al., 1995), NDDI (Rouse et al., 1973), and Tasseled Cap Transformations (Crist and Cicone, 1984) were calculated, which
have been shown to improve classification results of vegetation (Price et al., 2002; Biswas, 2010). To minimize noise in the
land cover maps caused by clouds, lower quality of certain LANDSAT images and the LANDSAT 7 Scan line error, the
seasonal selection of imagery and composite creation was repeated five times for each year with different selection dates for
the seasonal composites—and Tasseled Cap Transformations (Crist and Cicone, 1984) and included seasonal information by
forming seasonal composites and a harmonic trend analysis of NDVI. In addition, aerial photographs from 1964 were obtained
from of the Forest Research and Training Centre, Ministry of Forests and Environment, Kathmandu. The individual
photographs are combined and a point cloud is generated, followed by an orthomosaic with a resolution of 2-3 m (Fig. S1).
The resulting orthomosaic with a resolution of 2-3 m was resampled to a resolution of 30 m to compare it to the annual land
cover time series.

3.3 Field data

Vegetation types were classified according to two schemes based on vegetation assemblages in Bardia NP (Appendix C) (Dinerstein, 1979b; Jnawali and Wegge, 1993; Peet et al., 1999a)—After classification, the modus of the resulting five land cover maps was taken to obtain the most prevailing classification for each pixel for each year. The composites of 1995, 1997 and 2006 were not classified as too many clouds were present for 1995 and 1997 and a too large number of erroneous pixel values in the image composite was observed for 2006. This was caused by the LANDSAT 7 Scan-line error. The open access cloud platform software Google Earth Engine (Goreliek et al., 2017), earlier applied for mapping changes in floodplain ecosystems (Harezlak et al., 2020; Zurqani et al., 2018; Donehyts et al., 2016; Van Iersel, 2020), was used for selecting, processing and classifying LANDSAT imagery, whereas RStudio was used for post classification analysis.

3.3 Field data

Ground truth data was collected to (1) train the classification models and (2) validate the land cover maps of 2019. The study area was traversed with a handheld GPS and the vegetation type was noted at 352 sample locations (Appendix B). Seventy percent of the locations was used for training and thirty percent for validation (see section 3.4). The beforehand devised sampling scheme was not safe to execute in the field; therefore locations were selected based on accessibility and activities of staff at NTNC. Timing of surveying was in October and November 2019, coinciding with the first seasonal composite of LANDSAT imagery. For locations which were either inaccessible or unsafe to enter (water bodies, large patches of tall grasses, dense forest) the vegetation type was determined from a distance (elevated terrain or tower, Figure A1) and assigned coordinates via visual interpretation of high resolution satellite data (Planet Team, 2017) in combination with field knowledge. The Planet imagery has a spatial resolution of 5 m and its acquirement date coincides with the fieldwork dates.

Vegetation types were classified according to two schemes. The first land cover series (level 1) contains four classes and spans from 1993 to 2019. The second series (level 2), from 2013 to 2019, contains 8 classes and uses solely LANDSAT 8 imagery. Choice for which vegetation types to classify were done on the basis of literature on vegetation assemblages in Bardia NP (Dinerstein, 1979b; Jnawali and Wegge, 1993; Peet et al., 1999a) and on remote sensing studies in similar environments (Biswas et al., 2014; Sarma et al., 2008; Thapa, 2011; Biswas, 2010; Arieira et al., 2011; Sharma, 1999; Dinerstein, 4979a)(Biswas et al., 2014; Sarma et al., 2008; Thapa, 2011; Biswas, 2010; Arieira et al., 2011; Sharma, 1999; Dinerstein, 1979a). The first land cover series (level Level 1) uses the contains four classes (water, bare substrate, grassland (a combination of alluvial tall grassland, mixed tall grassland, short grassland) and forest (a combination of the Khair sissoo, riverine, mixedhardwood) and Sal forests), spans from 1993 to 2019. The more detailed dataset (level second series (Level 2) uses the solely <u>LANDSAT 8 imagery (2013 - 2019). The eight</u> classes consist of water, bare substrate, alluvial tall grasslands, short grasslands, mixed tall grassland, shrubland, riverine forest and Sal forest- (Appendix C). The grouping was based on literature, elaborated in Appendix A. Vegetation cover and height have been proven to discriminate and cover of vegetation, together with the dominant species present were used to assign the vegetation types (Appendix D). These criteria enable discrimination between riverine grassland from and non-riverine grasslands in the TeraiTAL (Biswas et al., 2014). Vegetation height, together with the dominant species present were used to assign the vegetation types (see Appendix C). For short grassesFor short grasslands the height criterion of < 2 meterm is used. Shrublands (< 5 m) entail no distinct vegetation assemblage recorded in BNP, but was used in the TeraiTAL by Biswas (2010)(2010) and Sarma et al. (2008), and provided useful successional information on impact of disturbances in Canada (Hermosilla et al., 2018).

3.4 Classification & validation

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The Random Forest model (Breiman, 2001) was used for classification of the remote sensing images, considering its earlier successful use in mapping of vegetation in and near floodplains (Van Iersel et al., 2018; Belgiu and Drăgu, 2016; Harezlak et al., 2020) and most satisfactory results during initial tests. Two time series of land cover maps (Level 1 and Level 2) were created based on the above described classification schemes. A manual classification was made of and the aerial photographs from 1964 according to the level 1 classification was classified manually (Level 1) (Figure A2). The composite of these photographs has a resolution of 2.3 meters. The Random ForestS1). For 2019, the classification model was trained on the image composites of 2019 with 70% of the field samples and the other 30% was used for validation—of this year. An error matrix was constructed for accuracy assessment and the user's and producer's accuracy was calculated (FAO, 2016). For classifying earlier years, all field samples of 2019—were used to train a classification model on the 2019 two seasonal image composite, as this sample dataset cannot be used for validating earlier years (1993–2018), but the composites of 2019 and image derived attributes. The trained model itself for 2019 is transferable through time (Gómez et al., 2016). Pixels of _each LANDSAT—composite were then labelled with the most voted class for these pixels (Lawrence et al., 2006; Pal, 2005).

To validate earlier years in the Level 1 time series, additional validation samples (106) were collected with publicly available imagery via Google Earth for the years that imagery is available that fully covered the study area (2000, 2010, 2011 and 2018). A grid was created digitally with an equal distance of 1200 meters between sample locations and to prevent underestimation of the classes grassland, bare substrate and water, a second grid was overlain for the floodplain area. This validation was done for the level 1 classification as the publicly available imagery is not detailed enough for delineation of the level 2 classes. Software used for the explained survey collection was CollectEarth (Bey et al., 2016).available years (2000, 2010, 2011 and

2018). The imagery was not detailed enough for delineation of the level 2 classes. A grid was created digitally with a distance of 1200 m between sample locations outside the floodplain and 600 m inside the floodplain.

435 **3.5 Environmental drivers**

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3.5.1 Meteorologic and hydrologic data

Hydrologic and meteorologic datasets of the Karnali River were purchasedobtained from the Department of Hydrology and Meteorology office in Kathmandu. The datasets include the yearly maximum and minimum discharges, daily discharges and monthly precipitation. They were This data was measured at the Chisapani weather station, nearnorth of the bifurcation of the Karnali River, and used to calculate hydrologic metrics (Table 1). Discharge ishas not been measured for the separate Karnali branches. From this data, hydrologic metrics were calculated (Table 2). Extreme discharges and their Based on the dataset of 54 years of yearly peak discharges (Q_{peak}), extreme discharges (Q_{ext}) with \geq 5-year recurrence times (\geq 12,500 m³ s⁻¹) were ealeulatedidentified using the Gumbel distribution (Gumbel, 1958) based on the dataset of 54 years. The 5-year recurrence time for discharges was used to discriminate extreme discharge years (> 12,500 m³ s⁺), as it. It has been postulated that discharges with these recurrences cause major changes in the river courses in the Ganga Plains (Thorne et al., 1993). For The extreme precipitation the extreme events (Pext) were identified by considering the peak rainfall months for each year. The third hydrologic indicator used was (Ppeak). Thirdly, the switch of the dominant discharge branch of the Karnali River system at the megafan was identified (Sinclair et al., 2017; Dingle et al., 2020b). Fourthly, the position of the water-filled channels during the dry period was extracted from the annual land cover maps (Table 2). The areal 1). We performed change detection of these water-filled channels was evaluated by comparing considering the land cover maps of time steps t and t + 1. Results were expressed in hectares, subdivided into, where t is 1 year. The results include newly covered area by water, disappeared water area and unchanged channels (Unchanged)..

3.5.2 Forest fires and anthropogenic disturbances

A forest fire dataset from 2000 to 2019 was obtained from NASA's Fire Information for Resource Management System (FIRMS). The dataset contains detections of fire, based on satellite imagery derived from the sensors on satellites (MODIS and VIIRS instruments.). Pixels that detected fire are flagged and attributes such as the brightness, confidence of occurrence and time of day are added. The MODIS imagery used has a 1 km resolution, it is 375 m for VIIRS. We interpret the number of pixels that detected a were flagged with presence of a fire were selected for each calendar year and we interpreted this number as the area that burned each year, but this is. This enables identification of years with high forest fire activity. It should be reminded that this method results in an overestimation of the actual area that actually burned. The coarse resolutions (1 km for MODIS, 375 m for VIRS) limit information on the exact extent of the fire. When a fire occurs at the border of two pixels, the two separate pixels are flagged, doubling its detected area. Also, when two fires occur at the same time on a spatial scale that is smaller than the resolution of the imagery, they are flagged as a single detection. Also, when one fire occurs at the border of two pixels, the two separate pixels are flagged doubling its detected area. The number of pixels that were flagged with presence of a forest fire for a calendar year was used as environmental indicator. The metric of forest fires is was expressed as the percentage of area that is affected by fires for both the area within and outside the floodplain. For these two zones, the surface area is corrected for the area that is covered by areal extent of water and bare substrate.

Table 21: Datasets of the environmental drivers (A, B) and land cover dynamics (C), the derived metrics and their timespan.

A. Hydrological and meteorological data	Period	Derived metric
Discharge dataset: Yearly maximum and minimum discharge.	1962-2015 1962-2015	Q _{ext} : Q _{peak} : Magnitude of yearly peak discharge (m³ s⁻¹) Q _{peak} : Yearly peak discharges and years with >Q _{ext} : Years where Q _{peak} ≥ 12,500 m³ s⁻¹ peak discharge
Precipitation dataset: sum of monthly precipitation (mm)	1964-2017 1964-2017	P _{ext.} : Yearly peak rainfall in a month (mm) P _{peak.} : Magnitude of the peak precipitation in a month for each year (mm) P _{ext.} : Years with extreme precipitation events
Change in dominant discharge branch of the Karnali River	2009	Moment in time
B. Forest fires		
Number of pixels that detected fires	2000-2019 (MODIS) 2012-2019 (VIIRS)	Surface area yearly affected by detected fires, in- and outside of floodplain (%)
C. Land cover dynamics		
Land cover area of classes, Level 1 and Level 2 classifications	Level 1: 1993—_2019 (excl. 1995, 1997, 2006) Level 2: 2013 - 2019	Yearly total area of each land cover class, within and outside floodplain (hectares)
Transitions of classes, Level 1 and Level 2 classifications	Level 1: 1993—2019 (excl. 1995, 1997, 2006) Level 2: 2013—2019	Interannual transitions of classes on pixel basis (hectares)
Channel dynamics: coverage of water class, based on Level 1 classification	Level 1: 1993—_2019 (excl. 1995, 1997, 2006)	Yearly coverage of new, disappeared and unchanged channels in floodplain (hectares)
Pattern of grassland class	Level 1: 1993—_2019 (excl. 1995, 1997, 2006)	PD, ED, AI and LSI (See Sect. 3.6 and Table 32)

3.6 Analysis of land cover dynamics

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From the Level 1 land cover timeseries time series, two metrics were derived representing land cover dynamics, with focus on the grasslands (Table 32). Firstly, land cover dynamics were expressed with interannual transitions between land cover classes-The transitions were calculated on a pixel basis with yearly time steps (Mas and Vega, 2012). A number of 16 combinations were calculated for the level 1 classification and 64 combinations for the level 2 classification assisted by existing scripts Intensity analysis (Pontius and Khallaghi, 2019) for crosstabulation and raster (Hijmans et al., 2011) packages in R (https://rproject.org, R Core Team 2016). A conversion factor of 0.0785 was used to transform the area of the LANDSAT pixels to hectares. For the analysis the yearly total removal of vegetation, which is the transition from both forest and grassland combined to bare substrate and water (Table 3). As second indicator of the land cover dynamics the yearly transition of grassland to bare substrate class was calculated, enabling insight in the loss of grassland. Secondly, land cover dynamics were characterized by their pattern with landscape fragmentation metrics, which quantify In total 16 combinations were calculated for the for the level 1 classification and 64 combinations for the level 2 classification. From these combinations we identified 'Removal of vegetation', which is the transition from forest and grassland towards bare substrate and water, and the 'Transition of grassland to bare substrate' (Table 2). Secondly, the land cover dynamics were characterized by their pattern with landscape fragmentation metrics. These metrics enable quantification of spatial patterns through time (McGarigal and Marks, 1995; McGarigal, 2002; Plexida et al., 2014). These metrics have been used for similar ecosystems in the Terai to quantify changes and enable comparison of the landscape through time and in different areas (Biswas, 2010; Thapa, 2011)(McGarigal and Marks, 1995; McGarigal, 2002; Plexida et al., 2014) and have been used for similar ecosystems in the Terai (Biswas, 2010; Thapa, 2011). Landscape fragmentation metrics of the aerial photograph of 1964 were not calculated, as its spatial resolution and classification method differ from the LANDSAT maps. This influences the delineation of grasslands, and with it, the outcome of the fragmentation calculations. To link the land cover dynamics to the environmental drivers, we analysed the coincidence of extreme events and performed a statistical test with the yearly magnitude of the metrics. First, we identified years where extremes in environmental drivers (Q_{ext} and P_{ext}) coincide with large interannual transitions of grassland to bare substrate (Table 2). This was substantiated with statistical tests (spearman tests in scatterplots) between the calculated metrics time series of the annual extremes of environmental drivers (Q_{peak}, P_{peak}) and the interannual transitions of land cover dynamics (yearly removal of vegetation cover yearly transition from grassland to bare substrate. (Table 2).

Table 32: The metrics that are used for establishing effects of drivers on land cover dynamics, including the method and whichthe part of the study area that is considered. Lower half: fragmentation metrics used for quantification of the grassland pattern (McGarigal and Marks, 1995; McGarigal, 2002; Plexida et al., 2014)(McGarigal and Marks, 1995; McGarigal, 2002; Plexida et al., 2014). This part of the table is partly adapted from Sertel et al. (2018).

Land cover metrics	Description	Symbol	Method	Area
Removal of vegetation	AreaSurface area (ha) changed from of forest and grassland to the water and bare substrate.	Qpeak	Spearman	Floodplain
		Qpeak	Spearman	Floodplain
Transition of grassland to bare substrate	AreaSurface area (ha) changed from grassland to bare substrate	Qext	Coincidence	Floodplain
to bare substrate		Pext	Coincidence	Outside FP
		Q_{peak}	Spearman	Floodplain
		P_{peak}	Spearman	Outside FP
Landscape				
fragmentation metrics				
Patch Density	Number of patches of corresponding patch type per unit area_	PD	Raster calculations	In- and outside FP
Edge Density	The sum of the lengths (m) of all edge segments in the landscape, divided by the total landscape area (m^2) .	ED	Raster calculations	In- and outside FP
Landscape Shape Index	A standardized measure of patch compactness that adjusts for the size of the patch.	LSI	Raster calculations	In- and outside FP
Aggregation Index	A percentage that describes the ratio between the observed number of like adjacencies and the maximum possible number of like adjacencies, with respect to the proportion of the landscape compromised of each patch type.	AI	Raster calculations	In- and outside FP

505 **4. Results**

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The spatial spatio-temporal dynamics of the 27-years land cover series are first presented along with the historic analysis of land cover in 1927 and 1964, whereafter and 1984. Hereafter, the datasets on environmental drivers are evaluated. The areal coverage through time and retrogressional presented and coupled to the yearly spatial extent of land cover classes and transitions from grassland to bare substrate are then presented and coupled to the extremes in environmental drivers to gain insight in the relation between the space-time dynamics of vegetation and environmental drivers.

4.1 Spatial dynamics of Spatio-temporal land cover dynamics in BNP

4.1.1 Classification of land Land cover maps

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For the reference year 2019, the overall accuracies ealculated with the set of ground truth data as selected for validation are 84.7% for the level 1 classification and 75.7% for the level 2 classification (Appendix E). The classes of drymixed tall grassland and shrubland underperform for the level 2 classification because of the low number of samples and their mixed presence in the field, often in patches smaller than the resolution of the LANDSAT pixels. For the level 1 classification, the calculated accuracies were 75.5% (2000), 86.8% (2010), 84.9% (2011), 78.3% (2018) as based on the data extracted from Google Earth (Appendix E). Classes typically underperforming in accuracy are bare substrate for the user's accuracy and water and grassland for the producer's accuracy. For water, seasonal differences in water level explain the interannual changes in coverage and lower accuracy, and not all available satellite imagery used for validation was available at the same date as the LANDSAT imagery for classification and the moment of sampling in the field.

A significant change in land cover (deforestation) around BNP can be seen when comparing the topographic maps of 1927 and 1984, particularly at the area in-between the Kauriala and Geruwa Rivers (Appendix F). This is in line with the general observation that the TAL experienced high rates of deforestation in the previous century. For BNP, the channel belt of the Geruwa river was located more eastwards in 1964 compared to the situation in 2019 (Fig 3a and 3b). For grasslands, the comparison of the classified LANDSAT data with the historic aerial imagery of 1964 indicates that several patches found on the land cover maps of SI I). A number of grassland patches observed between 1993 -and 2019 can be traced back to the classified aerial photograph composite of 1964 (Fig 3f). Oppositely, a large part of the grasslands has transitioned into forest (1082 ha from 1964 to 1993). Phantas However, grasslands have shrunk in size and surrounding smaller patches completely disappeared (Fig 3g3f). The phantas Baghaura, Lamkauli and Khauraha, located outside of the active channel belt, are part of thesethe remaining grasslands and consist of short and mixed tall grasslands (Fig 3). The Lamkauli grassland can also be tracedobserved on the topographic map of 1927 as the village, settlement or grassland of Lathwa. No signs of a village or agricultural fields are observed on the 1964 imagery, in line with the record that the Baghaura phanta was and Lamkauli phantas were cultivated after 1965 and Lamkauli possibly even later (Brown, 1995; Pokheral, 1993). Cutting patterns are visible (Appendix D) highlighting the impact of anthropogenic activities for this grassland as recorded in literature (Bhatta, 2000)(Pokheral, 1993; Brown, 1995).

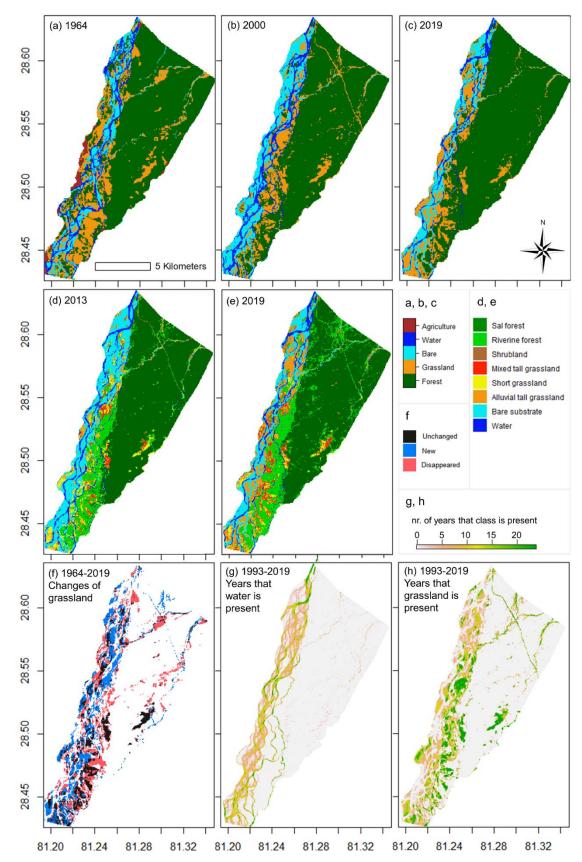


Figure 3: Overview of land cover maps. (a), the manual delineation of land cover classes from the aerial photograph composite of 1964; (b,e), level 1 data set; (d,e), level 2 data set; (f), the grasslands which are present in 1964 and during 2017-2019 and where they appeared and disappeared; (g,h), the number of years that water and grassland in total was present during the time period 1993-2019 (calculated from level 1 data set). The full series of land cover maps is presented in Appendix G.

When traversing the study area from west to east, the classes encountered most abundantly are in turn: bare substrate, alluvial tall grasslands, short and mixed tall grasslands, shrubland, riverine forest and lastly Sal forest (Fig. 3d and 3e). Shrubland is located close to the perennial streams and the national highway that crosses the park from Chisapani to the eastern border of the study area. Sal forest forms a sharp boundary with other vegetation types coinciding with a difference in elevation (Figure A3), which that decreases southwards. (Figure A2). Alluvial tall grasslands are dominant in and close to the active river channels and ephemeral streams, whereas shorter grasslands and mixed tall grasslands are located more distal from the active ehannel belt CHANNEL BELT (Fig. 3d and 3e). The changes are considerable active stream channels of the Geruwa River migrated westwards between years. Channel 1964 and 2019 (Fig 3a and 3b). Also migration of individual channels is observable observed, for example the change of the dominant discharge branch in the northwest (Fig. 3b and 3c) and the transition of abandonment a river channel southwest of the Baghaura phanta at 28.49 N°, 81.24 E°.

This channel transitions transitioned from an active river channel in 1964 (Fig. 3a) to grassland in 2000 (Fig. 3b) to forest in 2019 (Fig. 3c). Disappearance of grassland is observable adjacent to the active stream channel and throughout the forest. Between 1993 and 2019, grasslands and low-flow river channels practically have covered the entirety of theentire active channel beltCHANNEL BELT at least once, highlighting the dynamic nature of the braided river system (Fig. 3g). It indicates that the larger part of the present riverine forest located in and close to the stream channels is not older than 27 years. Compared to alluvial tall grassland in the active channel beltalong the Geruwa River and ephemeral streams, the more distal short and mixed tall grasslands are present for a larger time span. Also, grasslands along the highway (traversing the park from north to southeast) and ephemeral streams originating from the Siwalik hills in the northeast are present on a persistent basis (Fig. 3h).

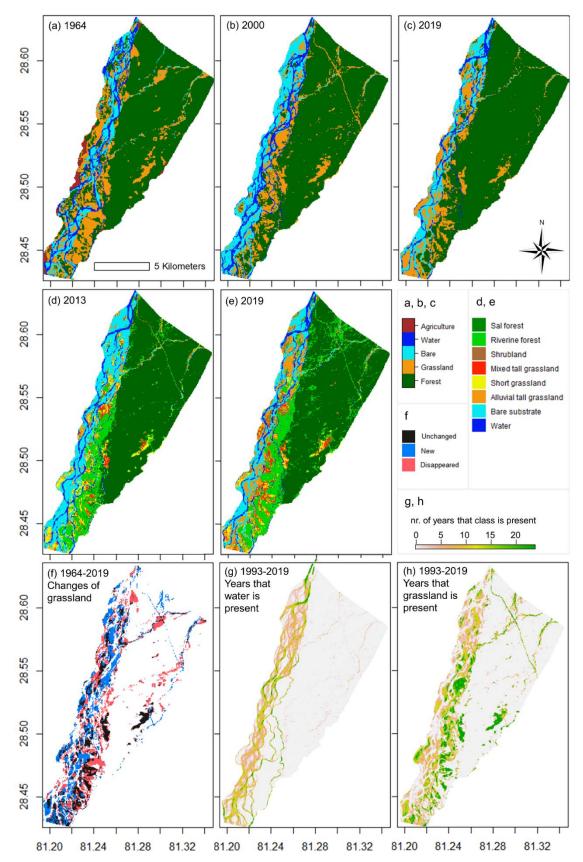


Figure 3: Overview of land cover maps. (a), the manual delineation of land cover classes from the aerial photograph composite of 1964; (b,c), level 1 data set; (d,e), level 2 data set; (f), the grasslands which are present in 1964 and during 2017-2019 and where they appeared and disappeared; (g,h), the number of years that water and grassland in total was present during the time period 1993-2019 (calculated from level 1 data set). The full series of land cover maps is presented in Figure A3.

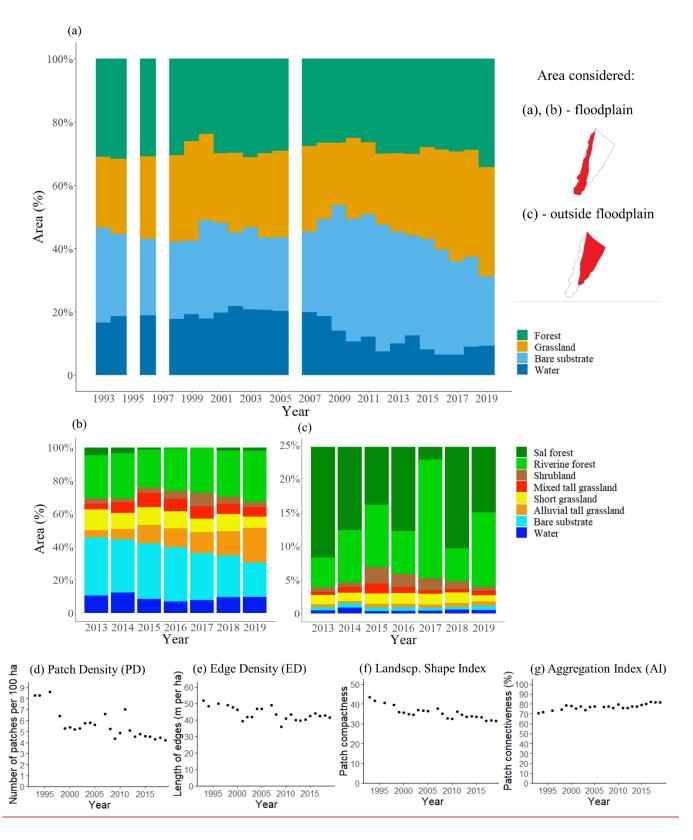


Figure 4: The temporal changes in land cover classes and pattern. (a), level 1 dataset; (b,c), level 2 dataset. White columns represent years with no data. Water can be interpreted as average water cover during the dry season. Note that a maximum of 25% was set in (c) for the surface area outside the floodplain for visualisation purposes; the remaining 75% of the area consists solely of Sal forest; (d,e,f,g), landscape fragmentation metrics 4.1.2 Areal development of land cover 1993 – 2019

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 $\frac{\text{The areal coverage of the land cover classes}(\underline{Table\ 2})}{\text{derived from the level 1 dataset}}.$

4-as predicted with the two classification sets. 1.2 Land cover dynamics 1993 – 2019

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Within the floodplain, succession and retrogression are observed in the levelLevel 1 classification (Fig. 4a) from which three distinctive-periods can be distinguished: 1993-1999, 2001-2008 and 2010-2019. These periods are characterised by a gradual decline in the areal extent of bare areasubstrate and an increase of the area covered by vegetation, particularly grassland. The three periods are separated by years wherein the surface area of bare substrate increased considerably as compared to the year before (2000; 2008 and 2009). In the first period, grassland is in an increasing trend until 1998 while the area covered by forest remains stable. In this period, water covers on average 18% of the floodplain, amounting to 932 ha. In the second period (2001-2008), the area covered by water (low flow channels) was slightly larger (1020 ha on average), while forest gradually decreased and grasslands saw an increase. In the third period (2010-2019), the area of bare substrate and water decreased to the lowest value since 1993. Bare substrate decreased gradually, whereas the coverage of water halved more abruptly to 478 ha on average (-53%).

%) as compared to the second period (2001-2008). The level 2 classification, spanning from 2013 to 2019 with 8 classes confirms this trendthe expansion of vegetation (Fig. 4b). Within the floodplain, increases are This expansion most evident for in the floodplain where the surface area of alluvial tall grasslands (grassland increased from 280 to 1185 ha, i.e., an increase of 151 ha per year). The short grasses and drymixed tall grasslands are more stable in terms of area covered not show such a trend. A notable decline of forest is observed in 2015, both on the floodplain (Fig 4b, to grassland) and outside of the

a trend. A notable decline of forest is observed in 2015, both on the floodplain (Fig 4b, to grassland) and outside of the floodplain (Fig. 4c, to mixed tall grassland and shrubland). Changes in shrubland coverage occur partly close to the perennial streams and the national highway that crosses the park from Chisapani to the eastern border of the study area. 4c). The grassland pattern shows a decrease of heterogeneity in the landscape: the number of patches decreased (Patch Density, Fig. 4d), the length of the edges of grassland with other classes decreased (Edge Density, Fig. 4e), the patches have become more compact

(Landscape Shape Index, Fig. 4f) and the patches have become more connected to each other (Aggregation Index, Fig. 4g).

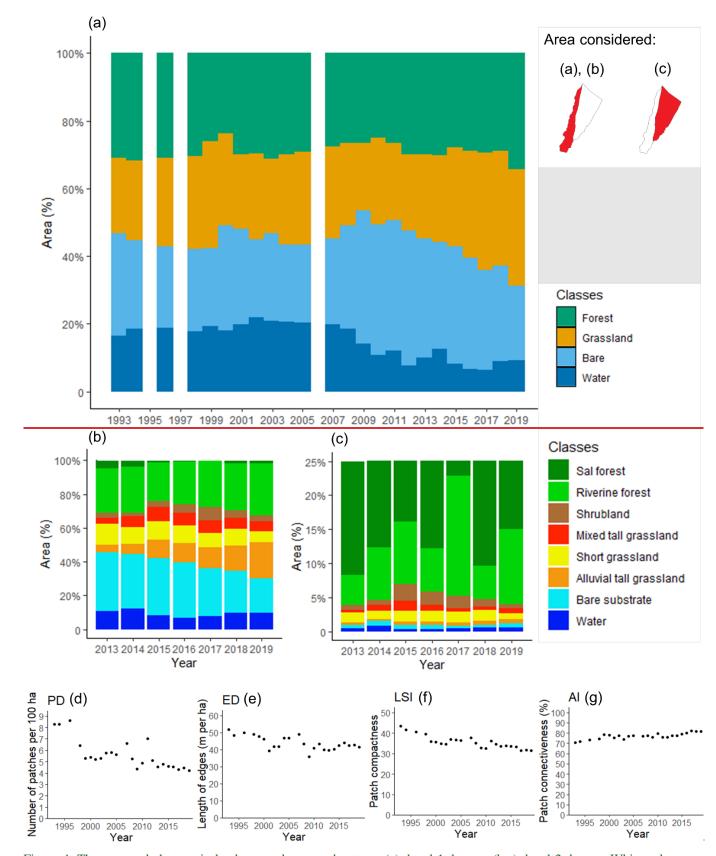


Figure 4: The temporal changes in land cover classes and pattern. (a), level 1 dataset; (b,c), level 2 dataset. White columns represent years with no data. Water can be interpreted as average water cover during the dry season. Note that a maximum of 25% was set in (c) for the surface area outside the floodplain for visualisation purposes; the remaining 75% of the area consists

605 solely of Sal forest; (d,e,f,g), landscape fragmentation metries-through time for the grassland class in the floodplain, derived from the level 1 dataset.

The vegetation follows a natural trajectory of succession after the abrupt increases of bare substrate in 2000 and 2009 (level 1), and decrease of riverine forest in 2015 (level 2). This natural trajectory is not observed for the areal extent of forest in 1999. In that year the areal extent of forest decreased and two years later, unnaturally fast, increased again. Other unrealistic interannual variations, that are not compliant with natural succession or retrogression, are observed for Sal forest and riverine forest in Fig. 4c. This indicates difficulty for the used level 2 classification model to discriminate between those two types of forest. Part of the inaccuracy of the level 2 classification is attributed to this.

The results for the temporal development of the grassland pattern are summarised as follows: the number of patches decreased (Patch Density, Fig. 4d), the length of the edges of the class grass with other classes decreased (Edge Density, Fig. 4e), the patches have become more connected to each other (Aggregation Index, Fig. 4e) and the patches of grass have become more compact (Landscape Shape Index, Fig. 4e). This indicates a decrease in heterogeneity in the landscape for the grassland pattern with respect to the bare substrate, water, and forest classes.

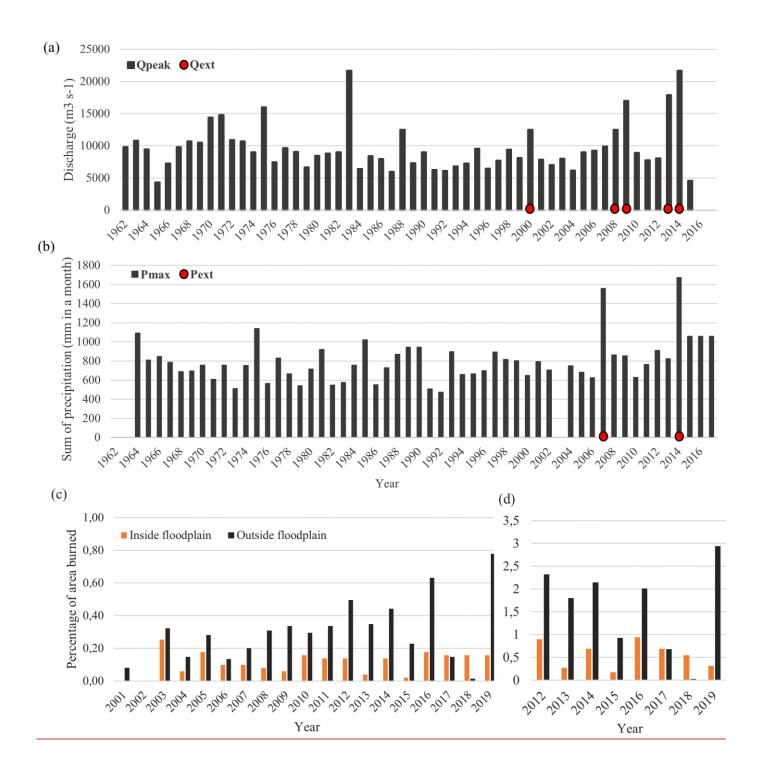
620 **4.2. Environmental drivers**

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The discharge regime of the Karnali near Chisapani is characterised from 1993 to 2015 by a large seasonal variation, with a mean discharge of 1,389 m³/s, a minimum of 173 m³/s and a maximum of 21,700 m³/s in August 2014. The median is 678 m³/s. Daily extremeExtreme discharges of the Karnali River can be up to three times the yearly average peak discharge. The monthly precipitation dataset shows a similar monsoonal variation, with a mean monthly rainfall of 196 mm and a maximum of 1,673 mm, a minimum of 0 mm and a median of 40 mm. For the studied period of 1993-2019, Fig. 5a shows the extreme peak discharges (Q_{ext}) with a recurrence time of more than 5 years as calculated with the Gumbel distribution and it also displays the periods with extreme precipitation (Pext in Fig. 5b). An overview of the daily discharge measurements is shown in Figure A4. The peak discharges of the daily and yearly dataset do not have the exact same values, but are in statistical correspondence to each other. We based our analysis on the yearly discharge dataset. Two months with extreme conditions were identified in 2007 and 2014, of which the second largest with 1560 mm for the monthly maximum (in 2007) is 47% larger than the third largest event- (Fig. 5b). On an annual basis, the highest cumulative precipitation was recorded for the same years: in 2007 and 2014 with 3,293 and 3,390 mm, respectively.

Forest fires were most abundant in the study area in the years 2012, 2013, 2014, 2016 and 2019, based on both the MODIS and VIIRS detection datasets (Fig. 5c and 5d). As of 2012 and onwards, the variance of the yearly affected area by fires in the floodplain increased: the maximum surface area affected increased, alternated with years wherein the surface area affected was very low. This is the case for the area outside of the floodplain. The channel dynamics of the channels waswere quantified by comparing the annual location of the stream channels in the dry period. The newly covered surface area of channels has decreased considerably after 2009 compared to the years before 2009 (Fig. 6). Most changes in surface area took place in the years 2007, 2008 and 2009, whereafter the average absolute area that changed decreased.



4.3. Relating environmental drivers to land cover dynamics

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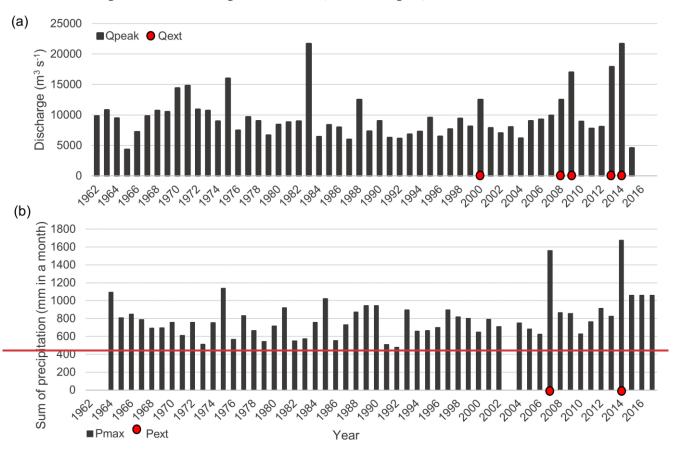
For the years 2000, 2008 and 2009, when extreme discharges occurred, a response in vegetation is observed in the floodplain in the form of increased transitions from grassland to bare substrate (Fig. 7a). These increased transitions are not observed for 2013 and 2014, whereas the recorded discharges were of a larger magnitude. The major change of the river course occurred in 2009, relocating the dominant discharge branch from the Geruwa River (in our study area) to the Kauriala River. The extreme discharges in 2009 most likely have markedly the changed the channel geometry, with a decrease in retrogression events after its relocation. This is also supported by the results of areal development of land cover (Fig. 4), where, before the channel relocation, the area of bare substrate only increases in the years that experienced extreme discharges, and not after the channel relocation. Outside the floodplain (Fig. 7b), the large transitions of grass to bare in 2007 and 2014 coincide with years that

experienced extreme precipitation events. These transitions occurred mostly near the ephemeral streams. Inside the floodplain, the surface area of grassland that transitioned is 20 times larger compared to outside the floodplain, and proportionally even larger as the floodplain area is smaller. This also reflects the more dynamic nature of the floodplain.

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For yearly peak discharges, a significant (p < 0.05) positive correlation is present between the maximum discharge in a year and the areal change from grassland to bare-within the floodplain in the period between 1993-2009 (Fig. 7e). This is not the case for 2010-2019 (p > 0.05), where the high peak discharges of 2013 and 2014 experience no increased transitions from grassland to bare compared to years with average peak discharges (Fig. 7d). Outside of the floodplain, extreme rainfall events and grassland to bare transitions show no correlation (p > 0.05), but the extreme precipitation months of 2007 and 2014 coincide with the highest transitions from grassland to bare (labelled in Fig. 7e).



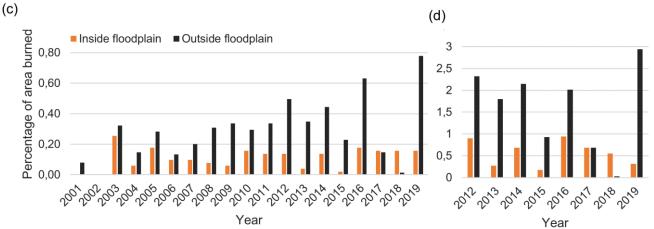


Figure 5: Environmental drivers. (a) Yearly maximum peak discharge (Q_{peak}) and identified extremes (Q_{ext}) are shown for the discharge dataset; (b) For each year the month with the most precipitation (P_{max}) is shown, along with the identified extreme years (P_{ext}) ; (c,d), surface area affected by fires in and outside the floodplain as detected from the MODIS (c) and VIIRS (d) datasets, respectively.

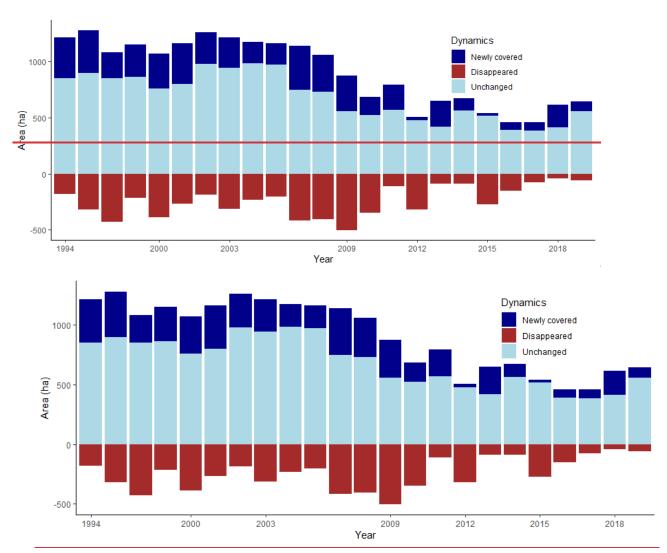


Figure 6: Change in surface area (hectares) covered by channels for each year compared to previous year; newly covered, area covered by new channels; disappeared, area abandoned by channels; unchanged, area covered by channels in current and previous year.

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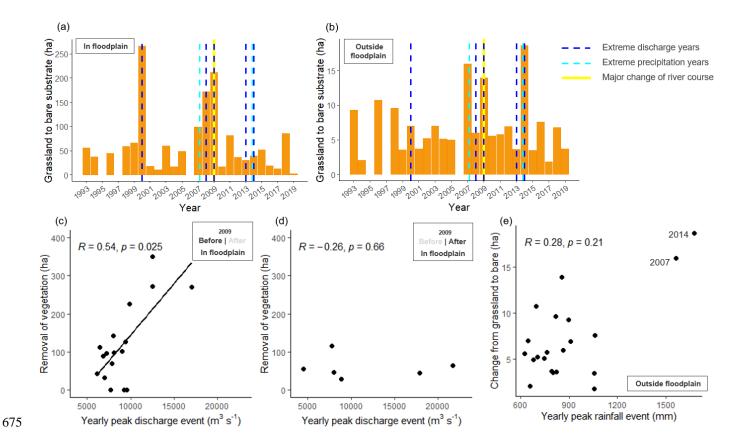


Figure 7: Environmental variables related to land cover dynamics derived from the Level 1 classification series. (a, b), annual transitions from grassland to bare substrate in hectares (orange bars), years that experienced extreme discharge events $(> 12,500 \, \text{m}^3/\text{s})$ and peak precipitation events within the floodplain (a) and outside the floodplain (b). Panels (c), (d) and (e) show correlation tests of environmental and land cover variables. (c) and (d) show the annual removal of vegetation (the transition of Sal forest and grassland towards water and bare substrate) as related to the magnitude of yearly peak discharge before the channel relocation (c) and after the channel relocation (d); (e) shows the yearly peak rainfall events (mm in a month) and the transition of grassland to bare substrate. Note the difference in scale on the y-axes.

4.3. Relating environmental drivers to land cover dynamics

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For the years 2000, 2008 and 2009, when extreme peak discharges occurred, a response in vegetation is observed in the floodplain in the form of increased transitions from grassland to bare substrate (Fig. 7a). These increased transitions are not observed for 2013 and 2014, while in these years the recorded peak discharges were higher than in 2000, 2008 and 2009. The major change of the river course occurred in 2009, when the dominant discharge branch relocated from the Geruwa River (in our study area) to the Kauriala River. The extreme discharges in 2009 most likely have markedly changed the channel geometry, with a decrease in retrogression events after its relocation. This is also supported by the results of areal development of land cover (Fig. 4), where, before the channel relocation, the area of bare substrate only increases in the years that experienced extreme discharges, and surprisingly not after the channel relocation in 2013 and 2014. Outside the floodplain (Fig. 7b), close to ephemeral streams the large transitions of grassland to bare in 2007 and 2014 coincide with years that experienced extreme precipitation events. These transitions occurred mostly near the ephemeral streams, on a much smaller (20 times) as compared to the floodplain of the Karnali River.

For yearly peak discharges, a significant (p < 0.05) positive correlation is present between the maximum discharge in a year and the areal change from grassland to bare substrate within the floodplain in the period between 1993-2009 (Fig. 7c). This is not the case for 2010-2019 (p > 0.05), where the high peak discharges of 2013 and 2014 experience no increased transitions from grassland to bare compared to years with average peak discharges (Fig. 7d). Outside of the floodplain, extreme rainfall

events and grassland to bare transitions show no correlation (p > 0.05), but the extreme precipitation months of 2007 and 2014 coincide with the highest transitions from grassland to bare (labelled in Fig. 7e).

5. Discussion

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We analysed the space time dynamics of land cover in Bardia National Park and the effects of environmental drivers and variation therein on these dynamics. We discuss the results of land cover development and highlight the hydrological, meteorological, forest fire and anthropogenic variables. The effects of these agents of environmental disturbances on land cover dynamics are discussed in a synthesis and are compared to similar ecosystems at the foot of mountain ranges. As land cover acts as a biophysical indicator of the state of an ecosystem (Meyer and Turner, 1992; Henderson Sellers and Pitman, 1992; Hansen and Loveland, 2012), the implications for BNP and possible cascading consequences for the ecosystem and nature management are considered.

5.1 Development of land cover

Our results show that grasslands decreased in areal coverage in the 29 years period between the aerial photograph of 1964 and satellite imagery of 1993. Almost all of the phantas present nowadays were also present in 1964, but declined in size and surrounding grassland patches disappeared. The encroachment of later successional stages of vegetation in BNP at the end of the previous century is in accordance with the decline of grasslands observed in previous studies (Peet et al., 1999b; Jnawali and Wegge, 2000). Odden (2007) showed that alluvial tall grasslands (Saccharum spontaneum dominated) and wooded grasslands (Imperata cylindrica dominated) saw a decrease in areal surface of 3.9% and 4.1%, respectively, between 1976 and 1997. These values were derived for the southern part of the study area by comparing land cover data from Dinerstein (1979a) and Sharma (1999). The discrepancy between the numbers on the decline of grassland as determined by Odden (2007) and this study are further discussed in Sect 5.5. In addition, Dinerstein (1979b) recorded that in 1976 the vegetation west and southwest of the Baghaura phanta was grassland with dispersed trees, which is in line with our findings for 1964. Our observation is that in 2019 this location is almost entirely covered by riverine forest (Fig. 3). This is also the case for the Mansuri Phanta, located north of the Lamkauli phanta, which was a grassland up to 1976 but has converted to forest due to succession (Fig. 3f) (Bhatta, 2000).

The annual time series of land cover for 1993—2019 shows periods with gradual successional changes and years with abrupt retrogression events. In 2019, the area in the floodplain covered by alluvial tall grasslands had expanded to a degree not observed during the previous three decades (1184 ha), but the total area of grasslands in the study area in 2019 (2114 ha) did not resemble the total coverage of grassland in the study area in 1964 (2612 ha). Our interpretation is that the total area of grasslands in 2019 contains a larger portion of alluvial tall grasslands compared to the situation in 1964, based on descriptions of Dinnerstein (1979b). With respect to the land cover pattern, a decrease of heterogeneity of the grassland pattern is observed from 1993 to 2019 as indicated by landscape fragmentation metrics. Grassland shows a trend towards more rounded patches; both the number of patches and length of edges of the grassland patches decreased. Additionally, our results show that the closer grasslands are located to the active river channel, the shorter in time they are present at that location (Fig. 3h). This corresponds with the highly dynamic nature of the braided river channel which causes rapid renewal of the floodplain guided by erosion and sedimentation processes. Riverine forest is abundant in the floodplain of the Karnali River and is observed in the level 2 time series as land cover along ephemeral streams. Bolton (1976) also recorded riverine forest as an 'undescribed'

riverine forest type' close to the ephemeral streams. Bare substrate is colonized quickly by early successional tall grasslands.

Concurrently, a portion of these alluvial tall grasslands transition towards the later successional grasslands and forests, but at a slower and presumably different rates. This difference in rate was also pointed out by Lehmkuhl (1989) who modelled the dynamics of the grassland and forest communities for Chitwan National Park. It is of interest to which extent the alluvial grasslands and riverine forest eventually expand, as their combined area still seems to be in an increasing trend in 2019.

5.2 Environmental drivers

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5.2.1 Hydrologic and meteorological drivers

The datasets on discharge and precipitation show a large seasonal difference between minima and maxima in a given year. Also, the interannual differences in maxima are considerable, both conform the monsoonal climate of the study area (USAID, 2018). The large discharge and precipitation fluctuations are typically associated with the existence of alluvial megafans (Leier et al., 2005) such as the Karnali megafan. The rainfall dataset shows two extreme rainfall events in 2007 and 2014. In 2014 the extreme precipitation in BNP and extreme discharge of the Karnali River coincide. Extensive loss of landmass and destruction of houses was recorded during the 2014 floods (Sinclair et al., 2017), and extensive flooding of the nearby Babai river also occurred in that year (Chhetri et al., 2020). In 1983, when the second largest discharge was recorded at Chisapani, extreme flooding occurred (MacClune et al., 2014) covering the Baghaura phanta with sediment (Bhatta, 2000). No such sedimentation on this grassland is observed with the classifications in this study. Furthermore, we observed abandonment of river channels and a westward migration of the active river channel belt of the Geruwa River by comparing the 1964 and 2019 land cover maps. This westward migration was also recorded by Rahhal et al. (2021).

The indications that a major redistribution of discharge occurred in 2009 is substantiated with the results that, since 2009, in the Geruwa River the water coverage during the dry period (-53%) decreased and the rate of channel migration declined. These lower dynamics implicate that erosion and sedimentation rates in the channel belt of the Geruwa River reduced, decreasing lateral erosion and rejuvenation of the floodplain following the principles and concepts of fluvial processes in floodplains (Corenblit et al., 2007). The reduced channel dynamics were not recorded by Rakhal et al. (2021), who contrastingly found an increasing trend in channel shifting in the eastern branch of the Karnali River from 1977 to 2013. Possibly, the reason for this discrepancy can be found in the timespan studied: in their study the last decade (2010 2019) was not fully assessed and the large scale channel migrations during the floods in 2009 could dominate the signal.

The reduction of discharge in the Geruwa branch after 2009 is in line with earlier observations (Sinclair et al., 2017). The discharge required to replace gravel that is present on bar surfaces is estimated to be 5,100 m³-s¹+ (Dingle et al., 2020b), which is exceeded annually. For coarser gravel near the apex of the Karnali megafan, it is estimated that a discharge of 23,500 m³-s¹+ to 31,500 m³-s¹+ is required to entrain the sediment and change the channel geometry with it (Dingle et al., 2020b). This is, however, larger than the discharge measured in 2009 of 17,000 m³-s¹+ that caused the channel relocation, but indications are very strong that indeed a major change in channel geometry occurred and consequently a redistribution of discharge.

5.2.2 Forest fires and anthropogenic disturbances

The areas that were affected by forest fires were largest in the years 2012, 2013, 2014, 2016 and 2019 according to both the VIIRS and the MODIS dataset. These years are alternated with years that conversely experienced the lowest areal affected by

forest fires. In the first 10 years of the dataset, the area affected by fires was more stable, whereas for the years from 2012 to 2019, the variance between years increased, attributed to fluctuations in area outside of the floodplain that was burned. Ghimire et al. (2014) created a fire hazard zonation map of BNP and found that most fires were present in the Sal forest in absolute terms and that the surface area that burned increased after 2008. They suggested changing climatic conditions as an explanation. Another explanation could be that anthropogenic management with fires increased, which is done extensively for maintenance purposes of the *phantas*.

We base the anthropogenic disturbances on the days of access for locals, the yearly amount of permits issued for resource collection and the stated status of Bardia National Park from literature (Table 1)(Bolton, 1976; Bhattarai et al., 2017; Bhatta, 2000; Thapa et al., 2021; Brown, 1997). Additionally, we considered the aerial photograph of 1964, which shows anthropogenic activities (cutting patterns) in the Lamkauli grassland, and the topographic map of 1927, where the Lamkauli grassland it is indicated as village. Based on the overview in Table 1 we interpreted the anthropogenic disturbances to be strongly diminished since establishment as a protected area in 1976, thereafter increasing towards the passing of the millennium as permits for resource collection increased threefold and management activities increased, followed by a large decrease in permits and days of access for resource collection (Bolton, 1976; Bhatta, 2000; Dinerstein, 1979b; Thapa et al., 2021; Brown, 1995). The intensity of the present day maintenance activities as compared to earlier times is not known. Persistent presence of short grasslands is observed on the channel bars southwest in the park as confirmed by field observations on that location (Figure A4).

5.3 The effects of drivers on land cover dynamics

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For the period before 2009, the results show a correlation between the transitions of grassland to bare substrate and the magnitude of yearly peak discharge. This relation is absent after 2009, although discharges were of a greater magnitude. This shows that during extreme discharges of the Karnali River, the impact of hydromorphological disturbances on vegetation in the floodplains of the Geruwa River decreased considerably. The absence of retrogression events during extreme discharges of the Karnali River in 2013 and 2014 contrasts with the part of our hypothesis that in Bardia National Park extreme discharges of the Karnali River always cause a removal of vegetation, followed by a period of regrowth. The change in dominant discharge branch and westward migration of the river channels of the Geruwa River have as consequence that the floodplain in Bardia National Park experiences less rejuvenation, as supported with the decreased channel dynamics (Fig. 7). This favours growth of later successional stages of vegetation. In this study this becomes clear by the expansion of alluvial tall grasslands and to a lesser extent that of riverine forest.

For precipitation events, the two most extreme events coincide with the largest transitions from grassland to bare substrate in the area outside of the floodplain. There is no correlation present for the yearly maxima of precipitation and retrogression of grassland. This implicates that only the extreme precipitation events generate sufficient discharge to reset grassland to bare substrate on a larger scale. These locations are predominantly close to the ephemeral streams.

We detected no direct signal between the used land cover metrics and years with extensive forest fires, although their importance for maintaining the disturbance dependent phantas is recorded in literature (Peet et al., 1999b; Bhatta, 2000). Occurrence of fires is particularly dependent on the water availability, temperature and presence of fuel. The annual minima of discharges through the Geruwa River do not correlate or coincide with years that saw a larger area impacted by fires and is

not regarded as a driver for forest fires occurrences. We interpret the forest fires at present not to be a principle factor that retard forest towards grassland, but that fires do play an important role in maintaining grassland patches, which was also suggested by Lehmkuhl (2000). A considerable portion of the recorded forest fires is thought to be anthropogenically induced, so natural forest fires in BNP even affect less surface area than calculated in this study.

Anthropogenic disturbances are considered to be the main driver of the presence of the grasslands on the *phantas*. This study supports the importance of anthropogenic management of phantas such as Lamkauli and Baghaura (Brown, 1995; Dinerstein, 1979b) and historic human interference. The reason *phantas* and surrounding grassland patches have decreased in areal coverage compared to the situation in 1964 can be found, firstly, in the establishment as a protected area in 1976 and associated decline in anthropogenic disturbances. Secondly, the westward migration of channels in the Geruwa River is followed by succession to later successional stages in the abandoned channels of the eastern side of the Geruwa River. The reason why the main *phantas* experienced less encroachment then surrounding smaller patches is plausibly because anthropogenic maintenance was predominantly executed on these large *phantas*, whereas surrounding smaller patches were excluded of these activities.

We analysed the spatio-temporal dynamics of land cover in Bardia National Park and the effects of environmental drivers and variation therein on these dynamics. The results are discussed and compared to similar ecosystems at the foot of mountain ranges. As land cover acts as a biophysical indicator of the state of an ecosystem (Meyer and Turner, 1992; Henderson-Sellers and Pitman, 1992; Hansen and Loveland, 2012), the implications for BNP and possible cascading consequences for the ecosystem and nature management are also considered.

5.41 Evaluation of obtained accuracies and limitations

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The observed temporal trends in land cover echo natural trajectories of succession or retrogression, apart from the rapid forest recovery observed in the years 1999-2000 and the unrealistic fluctuations in riverine and Sal forest for the level 2 classification. The difference in the calculated decline of grasslands (8% during the period 1976-1997) of Odden (2007) and this research (35% during the period 1964-1996) could be due to a number of explanations: (1) the studies differ in timespan studied; (2) the studies differ in area studied; (3) the studied years are in another phase of the biogeomorphologic cycle. The accuracies are similar to contemporary (Corenblit et al., 2007), eausing the areal extent of the grasslands to have changed due to a prior flood or absence of it; (4) for each map a different method for classification of land cover mapping was used, resulting in 4 different methods. The discrepancy could indicate a slight overestimation of forest classification of the level 1 classification and/or a slight underestimation of grassland classification of Dinnerstein (1979a) and Sharma (1999). Uncertainties in our results could originate from the classification models used, their input data (field samples and satellite image composites), validation data (collected in the field and with Google Earth), the change detection algorithm, and the datasets of the environmental drivers. Although these factors pose limitations to optimal classification accuracies, the accuracies are in line with similar studies (Hermosilla et al., 2018; Harezlak et al., 2020; Van Iersel, 2020). Therefore, we argue that the series of land cover maps are of sufficient accuracy (76%-87% for Level 1, 75% for Level 2) to draw conclusions on the general trends in land cover dynamics in the Geruwa River floodplain and adjacent part of Bardia National Park. The decreased surface area of grasslands is in accordance with observations in previous studies (Peet et al., 1999b; Jnawali and Wegge, 2000; Odden and Wegge, 2005; Neupane et al., 2020). Between 1976 and 1997, Odden (2007) recorded a decrease of grassland (-8%) in the southwest of the study area. Between 1990 and 2013, Neupane et al. (2020) recorded a decrease in the surface area of grasslands from 31% of the total surface area in BNP to 17% (-45%). In this study we recorded a decrease of -35% between 1964 and 1996. The differences in numbers could be due to: (1) the studies differ in time span studied; (2) the studies differ in area studied; (3) the studied years are in another phase of the biogeomorphologic cycle (Corenblit et al., 2007), causing the areal extent of For a study in Canada a more extensive algorithm was used, including a Hidden Markov Model and logical transition rules. Accuracies obtained with LANDSAT imagery were 70.3% for that study (Hermosilla et al., 2018). In the Dutch floodplains, vegetation was mapped by Harezlak (2020) on a 5 years basis with a 35 years LANDSAT series comparable to the level 2 classification in this study and an accuracy of 77% was obtained. As the accuracies of our land cover mapping are similar to contemporary studies, we argue that the series of land cover maps are of sufficient accuracy (76%-87% for level 1, 75% for level 2) to draw conclusions on the general trends in land cover dynamics in the Geruwa River floodplain and adjacent part of Bardia National Park. Important to keep in mind, when analysing the results on grasslands, is that our study does not map the grasses that are covered by tree crowns, which are also present in the study area, and therefore we underestimate the extent of grasslands in the study area. floodplain grasslands to differ greatly prior to or after a flood. This cycle is also reflected in our results (Fig. 4); (4) the methods to construct land cover maps differ. Uncertainties in our results could originate from the classification models used, the input data (field samples and satellite image composites), validation data (collected in the field and with Google Earth), the change detection algorithm, and the datasets of the environmental drivers.

5.2 Land cover dynamics from 1964 to 2019

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The deforestation outside BNP is in line with the general observation that the TAL experienced high rates of deforestation in the previous century, but in contrast with the encroachment of forest within BNP in the past decades. The total areal extent of grassland was higher in 1964 (2614 ha) than in 2019 (2114 ha). Our interpretation is that short grasslands were proportionally more abundant in 1964 than in 2019 and vice versa for the alluvial tall grasslands, which expanded to 1184 ha in 2019 (Dinerstein, 1979b). Dinerstein (1979b) recorded that vegetation west and southwest of the Baghaura phanta was grassland with dispersed trees in 1976, which is in line with our findings for 1964. In 2019, this location is almost entirely covered by riverine forest (Fig. 3a and Fig. 3e). This is also the case for the Mansuri phanta, north of the Lamkauli phanta, which was a grassland up to 1976 but has converted to forest due to succession (Fig. 3f) (Bhatta, 2000).

The period between 1993 and 2019 shows gradual successional changes and years with abrupt retrogression events. A decrease of heterogeneity of the grassland pattern is observed from 1993 to 2019, because grassland patches decreased in size and abundance and became more rounded and compacted. Riverine forest is observed in both the Karnali floodplain as well as along ephemeral streams. Bolton (1976) recorded the latter as an 'undescribed riverine forest type'. Bare substrate is quickly colonized by early successional tall grassland. Concurrently, a portion of these grasslands transition towards the latter successional grasslands and forests, but at a slower rate. Lehmkuhl (1989) pointed out these different rates of succession for vegetation in Chitwan National Park. It is of interest to which extent the alluvial grasslands and riverine forest eventually expand, as in the last decade their combined area shows an increasing trend.

5.3 Environmental drivers

890 <u>5.3.1 Hydrologic and meteorological drivers</u>

The large discharge and precipitation fluctuations are associated with the monsoonal climate at the foot of mountain ranges (USAID, 2018) and alluvial megafans (Leier et al., 2005) such as the Karnali megafan. and Kosi megafan. In 2014, the extreme precipitation in BNP and extreme discharge of the Karnali River coincided. Extensive loss of landmass and destruction of houses was recorded (Sinclair et al., 2017), and extreme flooding was also recorded in the nearby Babai River in BNP (Chhetri

et al., 2020). During the peak discharge in 1983, in magnitude second to the discharge in 2014, extreme flooding (MacClune et al., 2014) resulted in coverage of the Baghaura phanta with sediment (Bhatta, 2000). Between 1993 and 2019, we observed no conversion from grassland to bare substrate on this phanta. The active stream channels of the Geruwa River migrated westwards, which was also recorded by Rahhal et al. (2021).

For 2009 there are indications that a major redistribution of discharge occurred at the apex of the Karnali megafan. Dingle et al. (2020) suggested, based on LANDSAT imagery, that the western branch was already dominant in 2001. This is not reflected in the results of the annual land cover time series of this study, which show a relocation of the main discharge in 2009. In addition, the rate of channel migration declined since 2009 and the water coverage in the Geruwa River decreased in the dry period (-53%). In contrast, Rakhal et al. (2021) found an increase in long term channel migration in the eastern branch of the Karnali River between 1977 and 2013. The cause for the discrepancy can possibly be found in the time span studied: the last decade (2010-2019) was not fully assessed by Rakhal et al. (2021) and the large-scale channel migrations in 2008 and 2009 could dominate the signal. During low discharges, the current distribution of discharge is considered to be about 80% for the Kauriala River and 20% for the Geruwa River (Sinclair et al., 2017; Dingle et al., 2020a). During peak flow of monsoonal discharges, the distribution is thought to be higher in the Kauriala River (55-65%) (Dingle et al., 2017, 2020a). Time scales associated with avulsions at the apex of the Karnali megafan are estimated to be between ~400 and 7,000 years, (Dingle et al., 2020b).

suggesting that the new discharge situation at the apex could hold for centuries. It is estimated that a discharge of 23,500 m³ s⁻¹ to 31,500 m³ s⁻¹ is required to entrain the coarse gravel near the apex of the Karnali megafan (Dingle et al., 2020b). This is higher than the discharge measured in 2009 (17,000 m³ s⁻¹) that caused the channel relocation. Altogether, indications are very strong that indeed a major change in channel geometry occurred and consequently a redistribution of discharge, resulting in more static hydromorphological processes in the Geruwa River. If erosion and sedimentation rates decrease, also rejuvenation of the floodplain decreases (Gunderson, 2002; Corenblit et al., 2007). The extreme peak discharges are not the sole hydrological driver for vegetation (Pickett et al., 1987; Corenblit et al., 2007). For example, prolonged inundation is common in a number of conservation areas in the TAL that are more characterized as wetland complexes, such as in Chitwan National Park (Dinerstein, 1979a). Our interpretation is that that prolonged inundation is less of an important driver in BNP than in Chitwan National Park.

5.3.2 Forest fires, anthropogenic disturbances and herbivory.

The variance of area affected by forest fires between years increased: the surface area affected (predominantly outside the floodplain) was large in the years 2012, 2013, 2014, 2016 and 2019, with increasing maxima. Ghimire et al. (2014) found that the surface area that burned increased after 2008, and increase that we recorded since 2012. Ghimire et al. (2014) suggested changing climatic conditions as an explanation. We interpret the anthropogenic disturbances in relative terms, based on the days of access for locals, the yearly amount of permits issued for resource collection and what is known from literature on management activities (Appendix A) (Bolton, 1976; Brown, 1998; Bhatta, 2000; Bhattarai et al., 2017; Thapa et al., 2021). Anthropogenic disturbances by locals have strongly diminished since the establishment as a protected area in 1976, thereafter increasing towards the passing of the millennium, again followed by a considerable decrease (Bolton, 1976; Dinerstein, 1979b; Brown, 1995; Bhatta, 2000; Thapa et al., 2021). Persistent presence of (grazed) short grasslands is still observed on the channel bars southwest as confirmed by field observations on that location (Fig. A5). Maintenance by park staff also seem to have increased towards the passing of the millennium (Bhatta, 2000) and was extensive in recent years.

Herbivory effects land cover in a more gradual way than the pulse disturbances considered in this study, and therefore we do not attribute this driver to the abrupt vegetational changes. On the long term, interpretation of grazing pressure on the vegetation pattern is hindered by uncertainty around the development of numbers of livestock and wild herbivores before and after the establishment as conservation area. In 1975, biomass of livestock was 15-17 times higher 47,959 kg km⁻² than biomass of wild ungulates in 1977 (2842–3120 kg km⁻²) in the southwestern part of the study area (Dinerstein, 1980). Between 1976 and 1999, grazing by livestock decreased and ungulates are thought to have increased fourfold in numbers (Wegge and Storaas, 2009). It is of interest whether wild herbivores could have compensated for the decline in grazing pressure by livestock, on which no consensus exists in literature (Karki et al., 2016; van Lunenburg et al., 2017; Wegge et al., 2019).

5.4 The effects of drivers on land cover dynamics

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For the period before 2009, the results show a correlation between the transitions of grassland to bare substrate and the magnitude of yearly peak discharge. This relation is absent after 2009, although in 2013 and 2014 the peak discharges were of a higher magnitude than in 2009. This indicates that the impact of hydromorphological disturbances on vegetation in the floodplains of the Geruwa River decreased considerably during extreme discharges of the Karnali River. The absence of retrogression events during extreme discharges of the Karnali River in 2013 and 2014 contrasts with our hypothesis that in Bardia National Park extreme discharges of the Karnali River always cause a removal of vegetation, followed by a period of regrowth. The reduced hydromorphological dynamics favour transition to later successional stages of vegetation in the floodplain (Corenblit et al., 2007). This is in accordance with our observed expansions of alluvial tall grasslands and to a lesser extent that of riverine forest after 2009. The two most extreme precipitation events coincide with the largest transitions from grassland to bare substrate in the area outside of the floodplain along the ephemeral streams. Only during extreme precipitation events the magnitude of erosion and sedimentation processes seem large enough to convert vegetation to bare land, because no correlation was found between the yearly maxima of precipitation and the retrogression of grassland. On the spatial and temporal scale of this study, no signal was detected between the transitions of land cover classes and years with extensive forest fires. We interpret the forest fires not to be a principle factor that converts forest towards grassland, which was also suggested by (Lehmkuhl, 2000). However, fires do play an important role in maintaining grassland (Peet et al., 1999b; Bhatta, 2000; Lehmkuhl, 2000).

Altogether, we attribute the decrease in surface area and abundance of grassland patches to, firstly, the reduced anthropogenic disturbances by locals since the establishment as protected area in 1976. Secondly, the westward migration of channels in the Geruwa River is followed by succession to later successional stages in the abandoned channels of the Geruwa River. The contrasting recent expansion of alluvial tall grasslands is attributed to a decline in hydromorphological dynamics in the Geruwa floodplain. Phantas such as Lamkauli, Baghaura and Khauraha experienced significantly less encroachment than surrounding smaller patches. This is attributed to anthropogenic maintenance practices that are predominantly executed on these phantas, in contrast with the surrounding smaller grassland patches. If the grasslands would have been left unmanaged, encroachment of the forest would have been more severe.

5.5 Global comparison

Disturbance-dependent grasslands, especially those located in high rainfall regions, are threatened by indigenous forest expansion, which has been recorded in Australia, Africa, India and North and South America (Puyravaud et al., 2003; Banfai and Bowman, 2006; Silva et al., 2008; Wigley et al., 2009). Grasslands in the Nepalese TAL can be added to this list. This is

especially the case when changes occur in the environmental drivers which can cause cascading effects, impacting biodiversity (Parr et al., 2012) and the habitat of prey and their predators.

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The redistribution of discharge through the two branches is thought to have had a profound impact on land cover. The observed behaviour of this system can be compared to other nature reserves near dynamic fluvial systems to gain insight in the possible effects of a change in discharge regime. The settings of BNP are a parable to other nature reserves near lowland alluvial megafan systems, such as the Kosi megafan, located eastwards of the TAL, and the Taquari megafan in the Pantanal (Brasil). For the Pantanal, A comparison is made to parable conservation areas near lowland alluvial megafans. In the Pantanal (Brazil), extreme hydrologic conditions are explicitly analysed as a driving force in the pattern of vegetation communities (Arieira et al., 2011). The Taquari fanmegafan in the Pantanal experienced a shift in external conditions environmental drivers due to an avulsion, increasing susceptibility to deforestation and fire (Louzada et al., 2020). Furthermore At the foot of the Himalayas in northeastern India, extensive vegetational changes towards dryer vegetation types occurred after a shiftshifts in the river course in the Manas National Park (Sarma et al., 2008) and the courses in Jaldapara Wildlife Sanctuary (Biswas et al., 2014), both in northeastern India. For the Manas National Park, Sarma et al. (2008) stated that the change to forest could have been partly prevented by more active management of the grasslands after the relocation of the river. In the Jaldapara Wildlife Sanctuary, a shift in the river course of the Torsa River during floods in 1968 is seen as one of the primary factors for decreased coverage of grassland and increase in and in Manas National Park (Sarma et al., 2008). In Jaldapara Wildlife Sanctuary, floods in 1968 caused a relocation of the Torsa River. This is regarded as one of the primary drivers that caused a transition of grassland towards woodland (Biswas et al., 2014). For the Manas National Park, Sarma et al. (2008) stated that after the relocation of the river, the change of grassland to forest could have been partly prevented by more active management of the grasslands

These analogous scenarios demonstrate the possible consequences for BNP, namely a change in the (relative) dominance of environmental drivers, and possibly a shift towards dryer vegetation types and an increased susceptibility to droughts and fires. Without increased disturbances from another source, one expects that the now extensive alluvial tall grasslands will transition to ultimately a homogenic riverine forest or Sal forest as observed in comparable ecosystems and also considering the decreasing trend of heterogeneity of grassland in the landscape. The increased threat of encroachment makes anthropogenic interventions a greater necessity to maintain a large enough degree of disturbances for the early—to mid-successional stages of vegetation. However, if anthropogenic disturbances are extensive, the still present phantas could be maintained and the land cover could transition into short grassland (more *Imperata cylindrica* dominated), such as the *phantas* Lamkauli and Baghaura, based on successional relationships from literature (Dinerstein, 1979a; Lehmkuhl, 2000).

driver, for example anthropogenic disturbances, one could expect that a proportion of the now extensive alluvial tall grassland transition to a more homogenic (riverine) forest, especially considering the decreasing trend of heterogeneity of grassland in the landscape. This evokes questions on the degree that humans should intervene in a (semi-)natural system. The importance of 'human-dependent' habitat is also interesting in the light of how grasslands are described in literature. Classically, grasslands are seen as a transitional land cover from a successional point of view (Clements, 1916), while recent studies highlight the importance to subdivide grasslands from a conservation perspective (Buisson et al., 2019; Bond and Parr, 2010). Grasslands can be subdivided in into old-growth (or ancient) grasslands (Bond, 2008; Veldman et al., 2015) and cultural (or derived) grasslands, which are a result of anthropogenic activities. The natural. The old-growth grasslands are regarded to harbour a higher biodiversity and endemism when compared to the cultural grasslands. In contrast withto this subdivision for

conservation goals, <u>subtropical</u> grasslands in the TAL that are remnant of anthropogenic occupation and cultivation (the phantas) serve an important ecosystem role for tiger conservation goals in Bardia National Park by sustaining prey populations.

5.6 Relationship with nature management

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We recorded a shift of grasslands to forest from 1964 to 1993, most likely due to decreased anthropogenic disturbances after the establishment as protected area. We also recorded a second, more recent increase of vegetation in the floodplain due to a redistribution of discharge of the Karnali River. The state of the floodplain is regarded to be in transition towards a more static nature as the hydromorphological processes of the Geruwa River reduced in strength. It is of interest for how long the current distribution of discharge holds. Time scales associated with avulsions at the apex of the Karnali megafan are estimated to be between ~400 and 7,000 years, derived from SL dating and the amount of discharge needed to replace the coarser gravel at the apex (Dingle et al., 2020b). This means that the new discharge situation at the apex could hold for centuries. However, the peak discharge in 2009 that caused the change in dominant river branch was not very extreme as earlier highlighted and is more considered to have a 1 in 20 year chance of occurrence.

The channel geometries (and changes therein) and associated redistribution of discharge at the apex of the Karnali megafan are connected to the land cover distribution in the western part of BNP, and with it, the habitat of threatened megafaunal species. With the recent increase of coverage of alluvial tall grasslands, the areal extent of favourable habitat of megafauna as rhinoceros (Rhinoceros unicornis) and hog deer (Hyelaphus porcinus) (Odden et al., 2005; Jnawali and Wegge, 2000) increased, but the mechanism that maintains this habitat has reduced in strength. Not only the surface area, but also the spatial pattern of land cover is of importance. In Chitwan, habitat heterogeneity was positively correlated with the occurrence of the prey of tigers (Bhattarai and Kindlmann, 2012). Moe and Wegge (1994) supposed that a high density of ungulates in BNP is be possible due to the fine mosaic pattern of habitats in the Karnali floodplain. However, we observed a decreasing trend of heterogeneity for grasslands in BNP with the metrics calculated in this study. This ongoing trend to a more homogenic grassland pattern is unfavourable for the ungulate population (especially Chital deer) and tiger, although this is not directly observable in their numbers. Tigers have seen an impressive increase in numbers, particularly due to conservational effort, and there is a debate whether the number of ungulates declined or not (Wegge et al., 2019; Kral et al., 2017). More research is needed on the numbers and the explanation behind.

5.7 Future Research

That environmental drivers are important controls on the grasslands in valuable floodplain ecosystems in the TAL is clear in this study, but conservational effort and management decisions would benefit from a more detailed understanding of these relations. In the next decades, changes in the environmental drivers are expected, also due to the changing climatic conditions. For example, Rakhal et al. (2021) suggested that for the Karnali River basin, river dynamics could increase due to climate change. 5.6 In perspective of megafauna

Migrating channels at the apex of the Karnali megafan and associated redistribution of discharge are connected to the land cover distribution in the western part of BNP, and with it, habitat of threatened megafaunal species. With the recent ongoing increase of grassland near the Geruwa River, the spatial extent of favourable habitat for species that depend on alluvial tall grasslands, such as greater one-horned rhinoceros and hog deer (Jnawali and Wegge, 2000; Odden et al., 2005; Pradhan et al., 2008), increased, but the natural mechanism that maintains these early successional tall grasses has reduced in strength. In Chitwan, habitat heterogeneity was positively correlated with the abundance and habitat preference of prey species of tigers

(Bhattarai and Kindlmann, 2012). Moe and Wegge (1994) suggested that a high density of ungulates in BNP is possible due to the fine mosaic pattern of habitats in the Karnali floodplain. However, heterogeneity of grasslands in BNP shows a decreasing trend. This could be unfavourable for Chital deer and tiger, although their numbers suggest otherwise. Tiger population showed the largest increase in BNP among the conservation areas in Nepal (DNPWC and DFSC, 2022), particularly due to conservational efforts.

5.7 Future Research

In the next decades, changes in the environmental drivers are projected because of, amongst others, changing climatic conditions. Precipitation, temperature and mean annual discharge are projected to increase in the Karnali River basin except for the post-monsoon, when smaller amounts of precipitation and discharge could be expected (Dahal et al., 2020). Besides, mapping the other branch of the Karnali River, the Kauriala River, can shed additional light on the relation between the fluvial processes and land cover in this system. The two branches could possibly be seen as an antagonistic couple, wherein the areal coverage of the bare, grassland and forest could inversely mirror each other. The recorded reduced fluvial disturbances of the Geruwa River are likely to be of impact to the downstream Katerniaghat Wildlife Sanctuary in India as well.

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The combination of earth observation with data on environmental drivers proves to be useful for adding to the understanding of this ecosystem. The algorithm used in this study can be improved upon and could be used for the years to come to monitor the land cover development in BNP and other conservational areas, while monetary costs are minimal for this method. Future studies that use remotely sensed imagery are preferably designed with the advised optimizations incorporated, such as addition of Sentinel satellite imagery (ESA, 2015) and enhanced classification and change detection algorithms (e.g. Hermosilla et al., 2018; Van Iersel, 2020). Besides, imagery of a higher spatial and spectral resolution may aid in mapping of invasive species such as *Lantana Camara*, which pose a problem in BNP(Dahal et al., 2020; Rakhal et al., 2021). The combination of earth observation with time series of environmental drivers adds to the understanding of spatio-temporal dynamics of land cover and proves to be useful for monitoring with minimal monetary costs. With future technological advances in mind, the algorithm could be improved upon (e.g. Hermosilla et al., 2018; Van Iersel, 2020) to enable detailed analysis of vegetation dynamics on the community or species level, including invasive species (Bhatta et al., 2020). As we were only engaged in the descriptive and diagnostical fields of modelling, advances could also be made in predictive and prescriptive fields. This will expandIn addition, further research on the knowledge of such a systemdrivers and detail on the changing relative dominance of drivers, useful for will aid conservational effort and management purposes and decision making decisions in BNP and similar nature reserves where disturbance-dependent grasslands—mosaics are ofunder threat of forest intrusion.

6. Conclusion

We aimed to quantify the spatio-temporal land cover change and establish relationships with the environmental drivers. This was done in Bardia National Park, (BNP), located along the highly dynamic Karnali River at the foot of the Himalayas in a climate conductive of forest growth. We focussed on disturbance-dependent grasslands that are vulnerable forto encroachment. Grasslands in the floodplain of the Geruwa branch of the Karnali River and the adjacent part of Bardia National Park had a larger surface areaBNP were more abundant and were located more eastwards in 1964 than from 1993 to 2019. Grasslands at so-called phantas away from the floodplains are present nowadays and were also present in 1964. However, they declined in size and surrounding grassland patches disappeared. Landscape fragmentation metrics for 1993—2019 revealbetween 1993 and 2019. Between 1993 and 2019, grasslands showed a steady decline in heterogeneity in the landscape with respect to the

of the grasslands, pinpointing the encroachment decrease in size and abundance of the grassland patches in the study area. The more recent ongoing increase of surface area of grasslands can be attributed to expansion theof alluvial tall grasslands (Saccharum spontaneum dominated, a pioneer species) in the floodplain. Since 2009, the amount of water flowing through the Geruwa River decreased as a consequence of a redistribution of discharge at the apex of the Karnali Megafan to the Kauriala branch. Precipitation and discharge datasets show years with extreme circumstances. For forest fires the The maximum surface area that yearly burns affected by forest fires increased, but at the same time theaffected surface area affected by fires between years became more variable.

Before the shift in river course in between years. Since 2009, the Geruwa branch of the Karnali River was partially abandoned. Before 2009, the magnitude of discharge events shows a correlation with transitions from grassland to bare substrate. This correlation is not present anymore after 2009. In and no increased transitions are observed for 2013 and 2014 discharges measured upstream of, while the bifurcation pointrecorded peak discharges were of a greaterhigher magnitude than earlier measured peak discharges, but successional setbacks were absent during in these years. On top of that Also, the results show decreased channel dynamics of the eastern branch of the Karnali River- decreased since 2009. Hence, we confirm our hypothesis that from 1964 to 2019 a shift occurred from grasslands to later successional stages such as riverine and Sal forest. This shift is predominantly attributed to the a westward shift of the active stream channels of the Geruwa River and the establishment as conservation area in 1976 and with associated reduced anthropogenic disturbances by locals and the. The more recent expansion of alluvial tall grasslands is attributed to the redistribution of discharge in 2009 of the Karnali River. In 2019, land cover in the floodplain is more favourable for faunal species that dependent depend on tall alluvial tall grasslands such as the hog deer (Hyelaphus porcinus) and greater one-horned rhinoceros (Rhinoceros unicornis), although the mechanism to maintain these grassesgrasslands reduced in strength-since 2009. For faunal species dependent on short grasslands (Imperata eylindrica dominated)that benefit from a heterogenic vegetation pattern, land cover is thought to have become less favourable between 1964 and 2019 as the heterogeneity and surface area and the amount of the grassland patches decreased. As the Because hydromorphological disturbances processes in the floodplain became have become more static, other agents of disturbances are more paramount to prevent encroachment of grasslands.

1115 Data availability

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Data is available via the contact person by sending an email to jitsebijl@xs4all.nl.

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Competing interests

The authors declare that they have no conflict of interest that could influence the work reported in this paper.

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Appendices

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A: Anthropogenic activities in BNP
B: Pre-processing of imagery

505 C: Grouping of vegetation types
B:D: Criteria for classification
C:E: Error matrices

Figure A1: Photograph of grasslands

Figure A2: Photograph of incision of the Karnali River
Figure A3: Land cover maps, level 1 and level 2 series
Figure A3: Photograph of incision of the Karnali River

Figure A4: Hydrological year

Figure A5: Photograph of cattle grazing on the border of BNP

Supporting Information

EXECUTE Composite of aerial photographs of 1964

H:S2: Topographic maps:

III: S3: Image availability of LANDSAT

Appendix A: Overview of anthropogenic activities in BNP as recorded in literature

<u>Year</u>	<u>Activity</u>	Description	Source
<u>1925</u>	Commercial forestry	5 years of extensive deforestation.	Bolton (1976)
<u>Since 1950</u>	Increase of population	Increased deforestation and pressure on forest.	Brown (1997) and Bhatterai et al. (2017)

<u>1965-1975</u>	Possible cultivation of	Oral records deliver cultivation of these	Dinerstein (1979a) and Pokharel (1993)
	Baghaura and Lamkauli phantas	<u>phantas.</u>	
<u>1976</u>	Free access for locals refused and livestock grazing was prohibited.	Establishment as conservation area: The Royal Bardia Wildlife Reserve (348 km²)	Bhatta (2000)
1978-1994	Management practices	Controlled burning in phantas. Ploughing, introduction of grass species and uprooting of stumps from 1978 to 1994 on the Baghaura phanta for the reintroduced black buck (Antilope cervicapra)	Bhatta (2000)
<u>1979</u>	Resource collection allowed, 7 days of access	Local communities were granted the rights to collect thatch grass and reeds from the reserves once a year, which was a pioneer step towards a people-centered approach.	<u>Brown (1997)</u>
<u>1979-1983</u>	Relocation of settlements	572 families were relocated from the Babai Valley of BNP.	<u>Brown (1997)</u>
<u>1983</u>	Registered permits: 21,081; 15 days of access	Permits were issued for entering the part for resource collection.	Bhatta (2000)
<u>1994</u>	Nr. of days for resource collection reduced to 10 days	=	Bhatta (2000)
<u>1995</u>	Uprooting at the Khauraha and Baghaura phanta	Uprooting of unpalatable species (<i>Lantana sp.</i> and <i>Colebrookia sp.</i>) as part of the Bardia Integrated Conservation Project (1995-2001). Extensive uprooting was executed on the Khauraha phanta in 1999.	Bhatta (2000)
<u>1999</u>	Registered permits: 57,255 Extensive uprooting of small bushes and trees at Khauraha phanta	Gradual increase of permits from 21k (1983) to 57k (1999)	Bhatta (2000)
2020	Nr. of days for resource collection is 3 days. drastical decrease of permits (number unknown)	<u>=</u>	<u>Thapa et al. (2021)</u>

Appendix B: Pre-processing of imagery

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To establish an annual time series of land cover, we selected imagery of the Surface Reflectance dataset (Tier 1) of LANDSAT 5, 7 and 8 satellites between 1993 and 2019. This dataset of imagery is corrected for variation of the energy source (e.g. sun angle) and atmospheric effects (e.g. aerosol scattering and thin clouds) and is therefore best suited to use for diachronic analysis (Zanter, 2019; Young et al., 2017). We accounted for the differences in sensors by using coefficients (Roy et al., 2016) to correct LANDSAT OLI imagery to LANDSAT 7 ETM+ imagery. The incorporation of seasonal information improves the classification results, especially if height data is absent, as the spectral signatures of the separate classes are more distinct due to phenological differences of vegetation throughout the year (Van Iersel et al., 2016; Kelley et al., 2018). For each year, two separate collections of images were selected based on the day of the year (between 240-366 of the post-monsoon and 1-150 of the subsequent pre-monsoon). In addition, images were selected based on cloud cover (< 50 %), quality (> 9), geometric RMSE (Root Mean Square Error) of < 10 m and we performed cloud masking using the *Fmask* algorithm (Zhu et al., 2015). To enhance the available spectral information, NDVI (Rouse et al., 1973; Rock et al., 1986; Myneni et al., 1995; Beeri et al., 2007), NDDI (Gu et al., 2007), and Tasseled Cap Transformations (Crist and Cicone, 1984) were calculated, which have been shown to improve classification results of vegetation (Price et al., 2002; Biswas, 2010). From each seasonal selection the images were combined into a composite image by taking the median value of each seasonal collection. In addition, for each year a separate annual collection of images was created, which was used to calculate the seasonal variation of the NDVI metric with a harmonic trend for each pixel in terms of magnitude and timing of the annual change. This was added as indicator for the supervised classification model. To minimized noise caused by clouds, artifacts and the LANDSAT 7 Scan-line error we repeated the seasonal selection of imagery five times for each year with slightly different selection dates. After classification, the modus of the resulting five land cover maps was taken to obtain the most prevailing classification for each pixel for each year. The composites of 1995, 1997 and 2006 were not classified. Too many clouds were present for 1995 and 1997 and a too large number of erroneous pixel values in the image composite was observed for 2006, caused by the LANDSAT 7 Scan-line error. We used *Google Earth Engine* (Gorelick et al., 2017), earlier applied for mapping changes in floodplain ecosystems (Harezlak et al., 2020; Zurqani et al., 2018; Donchyts et al., 2016; Van Iersel, 2020),

550 Appendix A: Grouping of vegetation types: Vegetation assemblages present in Bardia National Park and its associated vegetation class as used in the level 2 classification model in this study. Grass assemblages from Peet et al. (1999a). Forest assemblages are from Dinerstein (1979a)

for selecting, processing and classifying LANDSAT imagery, whereas RStudio was used for post-classification analysis.

Appendix C: Grouping of vegetation types: Vegetation assemblages present in Bardia National Park and its associated vegetation class as used in the Level 2 classification model in this study. Grass assemblages from Peet et al. (1999a). Forest assemblages are from Dinerstein (1979a) and modification of (Jnawali and Wegge, 1993; Pokheral, 1993).

Assemblage	Vegetation class (level 2)
Typha elephantina assemblage; permanently waterlogged sites	Alluvial tall grassland
Phragmites karka-Saccharum spontaneum assemblage; seasonally inundated, heavily grazed	Alluvial tall grassland
Phragmites karka-Saccharum spontaneum-Saccharum arundinaceum assemblage	Alluvial tall grassland
Phragmites karka assemblage; Tall, dense riverine grassland, seasonal and permanent marsh	Alluvial tall grassland
Saccharum spontaneum assemblage; Mixed Saccharum spontaneum phase, Saccharum spontaneum phase, Saccharum spontaneum-Dalbergia sissoo phase, floodplain grasslands, alluvial soils, often inundated	Alluvial tall grassland
Imperata cylindrica-Narenga porphyrocoma assemblage; (1) Saccharum spontaneum-Saccharum bengalense phase, edges of wet sites, newer river terraces. (2) Imperata cylindrica phase, sites where tall grasses invading an Imperata cylindrica dominated sward.	Alluvial tall grassland if (1) Mixed tall grassland if (2) and > 50 % is higher than 2 m Short grassland: if > 50 % is lower than 2 m
Imperata cylindrica assemblage; Imperata cylindrica phase, Erianthus ravennae phase; Imperata-Saccharum phase; dry sites, well developed soils, previously cultivated	Short $\frac{\text{grasses grassland}}{\text{grassland}}$: if > 50% is lower than 2 m, or Mixed tall grassland: if > 50 % is higher than 2 m
Narenga porphyrocoma assemblage; Tall, dense grassland, older river terraces and wetter sites, influenced by fire	Mixed tall grassland
Themeda arundinacea assemblage; Tall, dense grassland, often at forest edge, well developed soils, influenced by fire	Mixed tall grassland
Sal forest	Sal forest
Dry sal forest	Sal forest
Hill sal forest	Sal forest
Khair-Sissoo forest (Dalbergia sissoo-Acacia catechu)	Riverine forest
Mixed hardwood forest (Ficus glomerata-Mallotus phillippinenis-Eugenia jambolana	Riverine forest
Moist riverine forest (Mallotus phillipinensis and Syzigium cumini)	Riverine forest

Appendix BD: Dominant species per vegetation type used for assigning classes. Cover and height are used for further discrimination of classes.

Criteria	Dry <u>Mixed</u> tall grasslands	WetAlluvial tall grasslands	Short grasslands	Sal forest	Riverine forest	Shrubland
	Narenga	Saccharum	Imperata	Shorea	Dalbergia sissoo,	
	porphyrocoma,	spontaneum,	cylindrica, Vitiveria	robusta,	Acacia catechu,	
Dominant	Themeda	Phragmites	zizanioides,	Terminalia	Mallotus	
species	arundinacea,	karka,	Desmostachyia	tomentosa	phillippinensis,	-
	Erianthus	Saccharum	bipinnata.		Syzigium cumini,	
	ravennae, Bombax	arundinaceum			Bombax ceiba,	

	ceiba	Lantana Camara				
Cover	> 50% grass	> 50% grass	> 50% short grass	> 50% trees	> 50% trees	> 50% shrubs
Height	> 2 m	> 2 m	< 2 m	> 5 m	> 5 m	< 5 m

Appendix CE: Confusion matrices of the level 1 and level 2 classifications for 2019, as calculated with 30% of the ground truth data and confusion matrices for 2000, 2010, 2011 and 2018 based on separate validation set (106 samples).

2019						
level Level 1	Forest	Grassland	Bare	Water	Sum	User's accuracy
Forest	38	5	0	0	43	88.4
Grassland	3	34	0	0	37	91.9
Bare	0	4	6	0	10	60.0
Water	1	1	1	5	8	62.5
Sum	42	44	7	5	98	
Producer's accuracy	90.5	77.3	85.7	100.0		84.7%

2019 level <u>Level</u> 2	Sal forest	Alluvial tall grassland	Short grassland	Bare	Water	Shrubland	Riverine forest	Mixed tall grassland	Sum	User's accuracy
Sal forest	24	0	0	0	0	0	0	0	24	100.0
Wet tall grassland	0	8	1	1	0	0	0	0	10	80.0
Short grassland	0	1	14	1	0	0	0	1	17	82.4
Bare	0	3	1	11	0	0	0	0	15	73.3
Water	0	1	0	0	11	0	0	0	12	91.7
Shrubland	1	0	1	0	0	2	1	3	8	25.0
Riverine forest	1	0	1	1	0	0	13	0	16	81. 3
DryMixed tall grassland	0	1	6	0	0	1	1	5	14	35.7
Sum	26	14	24	14	11	3	15	9	116	
Producer's accuracy	92.3	57.1	78.6	78.6	100.0	66.7	86.7	55.6		75.7%

2000 Level 1	Forest	Grassland	Bare	Water	Sum	User's accuracy
Forest	45	4	0	0	49	91.8

Grassland	3	20	0	0	23	87.0
Bare	0	7	10	5	22	45.5
Water	1	4	2	5	12	41.7
Sum	49	35	12	10	106	
Producer's accuracy	91.8	57.1	83.3	50.0		75.5%
p-value	1.43E-07	CI95	0.66	0.83	kappa	0.64

2010 Level 1	Forest	Grassland	Bare	Water	Sum	User's accuracy
Forest	45	5	0	0	50	90.0
Grassland	1	14	0	0	15	93.3
Bare	1	1	22	5	29	75.9
Water	1	0	0	11	12	91.7
Sum	48	20	22	16	106	
Producer's accuracy	93.8	70.0	100.0	68.8		86.8%
p-value	3.06E-15	CI95	0.79	0.93	kappa	0.81

2011 Level 1	Forest	Grassland	Bare	Water	Sum	User's accuracy
Forest	45	6	0	0	51	88.2
Grassland	1	13	0	0	14	92.9
Bare	0	1	20	7	28	71.4
Water	0	0	1	12	13	92.3
Sum	46	20	21	19	106	
Producer's accuracy	97.8	65.0	95.2	63.2		84.9%
p-value	1.10E-13	CI95	0.77	0.91	kappa	0.78

2018 Level 1	Forest	Grassland	Bare	Water	Sum	User's accuracy
Forest	46	4	0	1	51	90.0
Grassland	5	20	1	1	27	74.0
Bare	0	5	10	4	19	52.6
Water	0	2	0	7	9	77.8
Sum	51	31	11	13	106	
Producer's accuracy	90.2	64.5	90.9	53.8		78.3%
p-value	3.78-e09	CI95	0.69	0.86	kappa	0.67

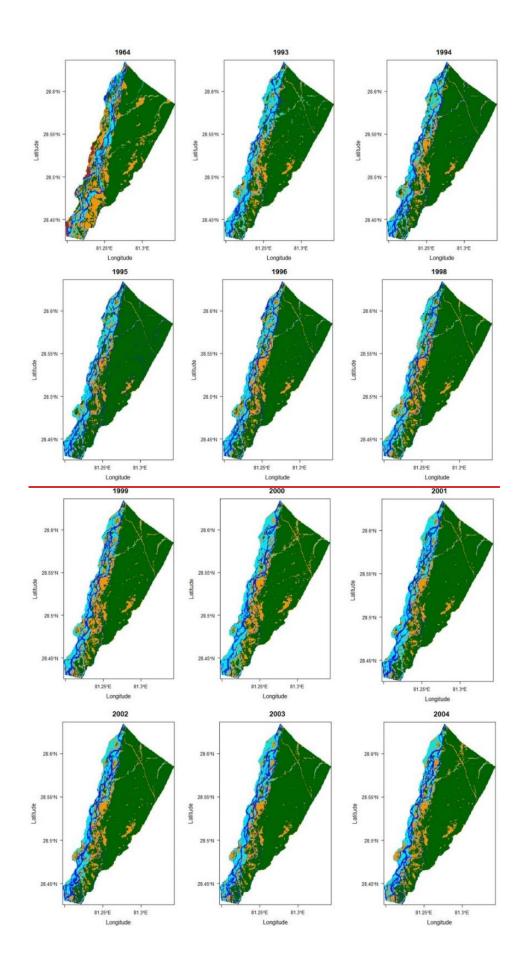


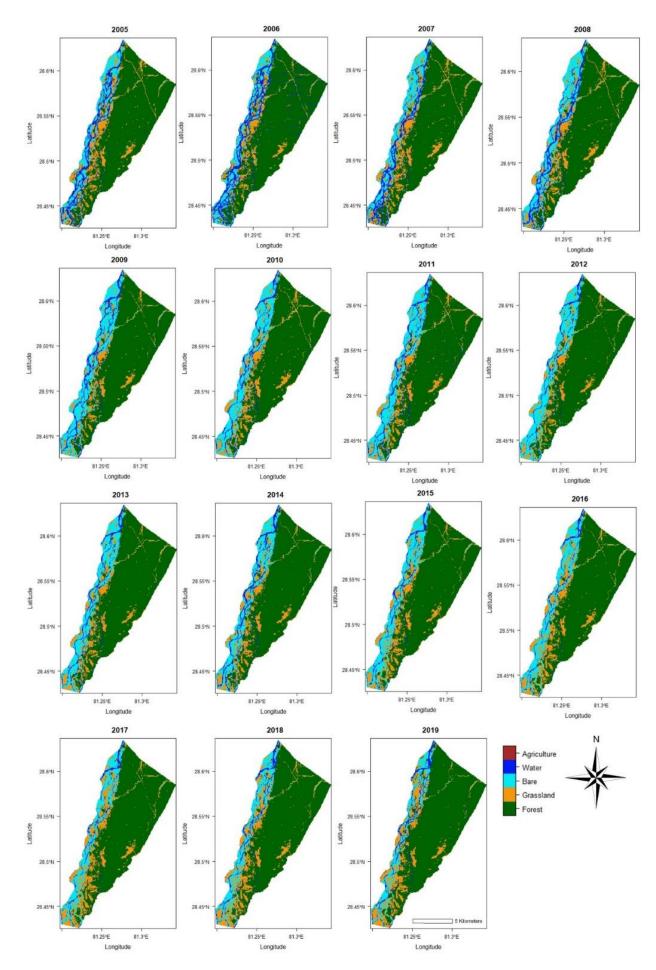
Figure A1: View from tower (Bagh Machan). In the front Saccharum spontaneum dominated tall grasslands grow, behind which.

Behind dryer types of grassland are present. At the right a patch of grassland is located present that was cut short. Behind the grasses, grassland riverine forest is present. Location The location is northwest from the Khauraha phanta. Such an As entering these grasslands is not recommended, this view is used to assign samples to map the ground truth locations for short grassland, alluvial tall grassland (in front) and on the far end mixed tall grasslands, as entering these grasslands is not recommended and riverine forest. This was done in combination with satellite imagery to obtain the exact coordinates. Photograph taken by authors.



Figure A2: Series of land cover maps of Level 1 and Level 2 classifications.





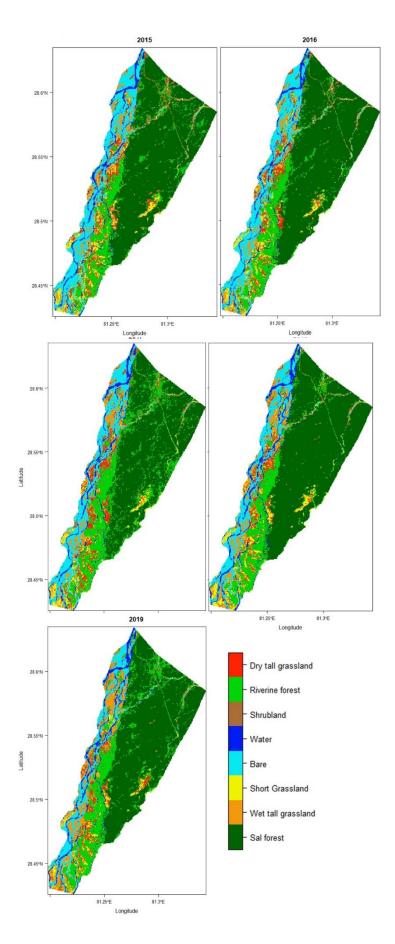
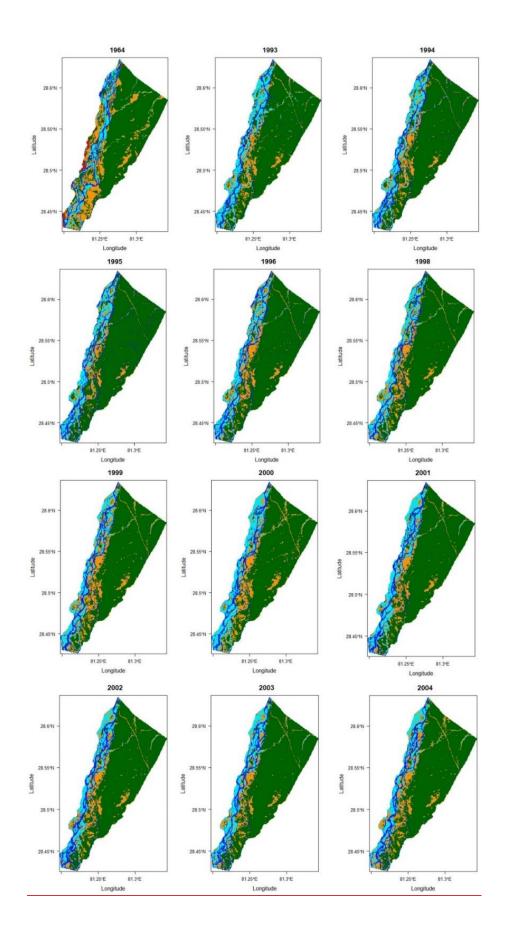
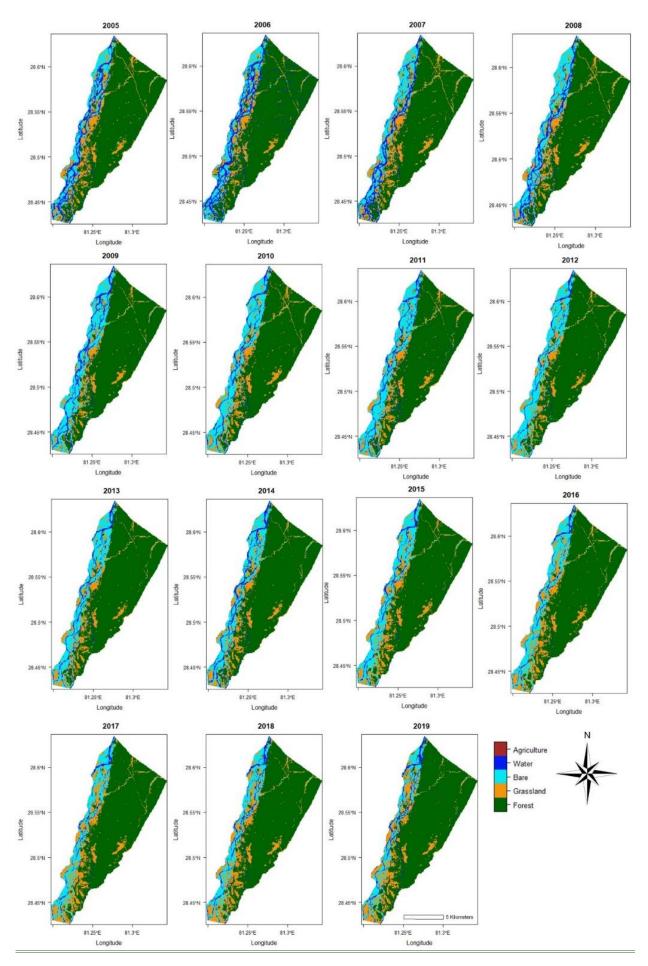


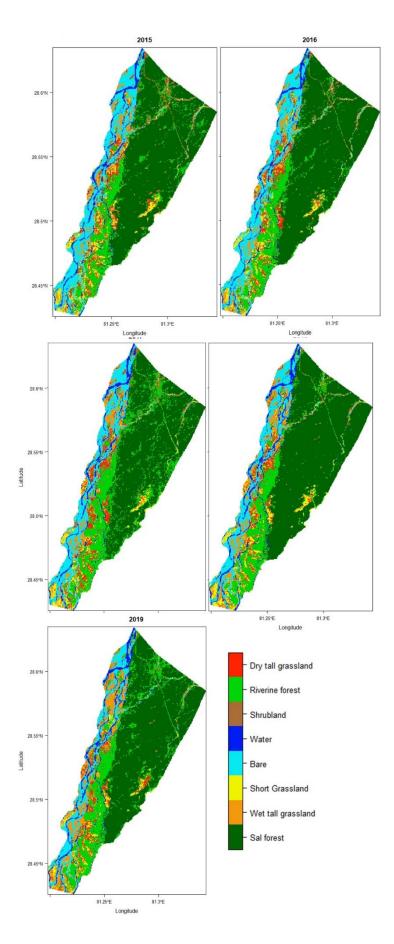


Figure A3: Sal forest on top and grasses below. On At the right hand, tall grasses grow, amongst the grasses classified as short grass. This grassland. The cliff marks the spatial extent of fluvial disturbances of the Karnali river. More southwards this boundary is lower The difference in height. This height differences elevation is largest at the boundary between Sal and the floodplain in the northwest of the study area, and gradually decreased to the southeast decreases southwards along the interface between of riverine forest and Sal forest. In the outcrop a loamy layer is positioned on top of a more sandy layer containing rounded pebbles, cobbles and boulders. Photograph taken by authors.

Figure A3: Series of land cover maps of Level 1 and Level 2 classifications.







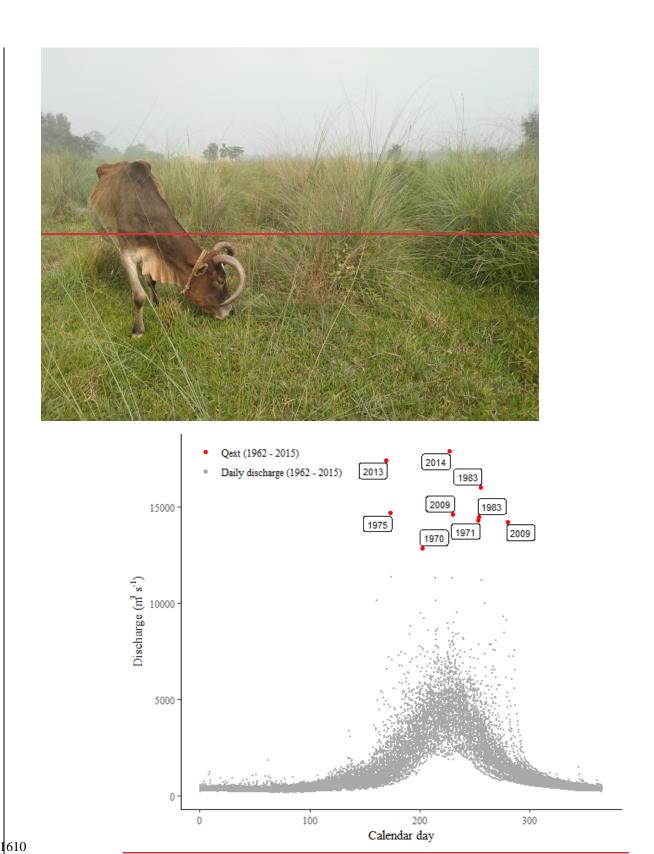


Figure A4: Discharge measurements and their calendar day. Red dots represent the identified extreme discharges (Q_{ext}) that are larger than $12,500 \text{ m}^3 \text{ s}^{-1}$.



Figure A5: Cattle grazing along southwestern border of study area on the western bank of the Geruwa River, along the southwestern border of study area. It demonstrates the grazing pressure herbivores pose on the vegetation pattern at the border of the park, nowadays, and, at earlier times, within the park when cattle grazing was allowed, within the park. Photograph taken by authors.