

# 1 Exploring environmental and physiological drivers of the annual 2 carbon budget of biocrusts from various climatic zones with a 3 mechanistic data-driven model

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15 **Abstract.** Biocrusts are a worldwide phenomenon, contributing substantially to ecosystem functioning. Their growth and  
16 survival depend on multiple environmental factors, including climatic ones, and the relations of these factors to physiological  
17 processes. Responses of biocrusts to individual environmental factors have been examined in a large number of field and  
18 laboratory experiments. These observational data, however, rarely have been assembled into a comprehensive, consistent  
19 framework that allows quantitative exploration of the roles of multiple environmental factors and physiological properties for  
20 the performance of biocrusts, in particular across climatic regions. Here we used a data-driven mechanistic modeling  
21 framework to simulate the carbon balance of biocrusts, a key measure of their growth and survival. We thereby assessed the  
22 relative importance of physiological and environmental factors for the carbon balance at six study sites that differ in climatic  
23 conditions. Moreover, we examined the role of seasonal acclimation of physiological properties using our framework, since  
24 the effects of this process on the carbon balance of biocrusts are poorly constrained so far. We found substantial effects of air  
25 temperature, CO<sub>2</sub> concentration, and physiological parameters that are related to respiration on biocrust carbon balance, which  
26 differ, however, in their patterns across regions. The ambient CO<sub>2</sub> concentration is the most important factor for biocrusts from  
27 drylands while air temperature has the strongest impact at alpine and temperate sites. Metabolic respiration cost plays a more  
28 important role than optimum temperature for gross photosynthesis at the alpine site; this is not the case, however, in drylands  
29 and temperate regions. Moreover, we estimated a small annual carbon gain of 1.5 g m<sup>-2</sup> yr<sup>-1</sup> by lichen-dominated biocrust and  
30 1.9 g m<sup>-2</sup> yr<sup>-1</sup> by moss-dominated biocrust at a dryland site, while the biocrusts lost a large amount of carbon at some of the  
31 temperate sites (e.g., -92.1 for lichen- and -74.7 g m<sup>-2</sup> yr<sup>-1</sup> for moss-dominated biocrust). These strongly negative values  
32 contradict the observed survival of the organisms at the sites and may be caused by the uncertainty in environmental conditions  
33 and physiological parameters, which we assessed in a sensitivity analysis. Another potential explanation for this result may be

34 the lack of acclimation in the modeling approach since the carbon balance can increase substantially when testing for seasonally  
35 varying parameters in the sensitivity analysis. We conclude that the uncertainties in air temperature, CO<sub>2</sub> concentration,  
36 respiration-related physiological parameters, and the absence of seasonal acclimation in the model for humid temperate and  
37 alpine regions may be a relevant source of error and should be taken into account in future approaches that aim at estimating  
38 the long-term biocrust carbon balance based on ecophysiological data.

## 39 **1 Introduction**

40 Non-vascular photoautotrophs, such as lichens, mosses, eukaryotic algae and cyanobacteria, together with heterotrophic  
41 microorganisms, form biological soil crusts (biocrusts) which occur in various environments across the globe and provide a  
42 wide range of important ecosystem functions, such as build-up of soil organic carbon and nutrients (Belnap et al., 2016;  
43 Chamizo et al., 2012; Dümig et al., 2014; Ferrenberg et al., 2018). Due to the importance of biocrusts in ecosystem functioning,  
44 their growth and survival have been extensively studied, through different methodological approaches (e.g. Ladrón de Guevara  
45 et al., 2018; Lange et al., 2006; Porada et al., 2019). An established measure to quantify the growth of biocrusts is their long-  
46 term carbon balance (hereafter, C balance), which corresponds to the (accumulated) net carbon flux across the system  
47 boundaries including all relevant carbon gains and losses.

48 In order to ensure survival, any species needs to achieve a positive C balance in the long-term, while negative values may  
49 occur for short periods. Acknowledging the importance of C balance, an increasing number of studies have investigated the  
50 long-term C balance of individual non-vascular organisms as well as biocrust communities, and their environmental drivers.  
51 An annual carbon budget of 21.49 g C m<sup>-2</sup> based on measured field data was reported in the study of Lange (2003b) on the  
52 crustose lichen *Lecanora muralis* growing on a rock surface in a temperate climate of southern Germany. Furthermore,  
53 Büdel et al. (2018) estimated an annual C balance of 1.7 g C m<sup>-2</sup> based on measurements on cyanobacteria-dominated  
54 biocrust in an Australian dry savannah ecosystem. Several other studies obtained long-term, large-scale values of the C  
55 balance by scaling up short-term, local measurements of CO<sub>2</sub> exchange rate under natural field conditions (Lange et al.,  
56 1994; Zotz et al., 2003). For an estimation of the global C balance of cryptogamic covers, which include biocrusts,  
57 conversion factors based on the maximum photosynthesis rate have been suggested as a best-guess solution (Elbert et al.,  
58 2012). However, there are some drawbacks to these approaches for acquiring a C balance at both organism and community  
59 scale. First, the measurement of the long-term continuous CO<sub>2</sub> exchange rate of an individual organism or biocrust  
60 community has technical limitations and is highly time- and resource-consuming. Second, upscaling via extrapolation may  
61 result in bias in annual C balance estimation if the length and the frequency of sampling cannot capture the full variability of  
62 CO<sub>2</sub> exchange throughout the year (Bader et al., 2010). Moreover, using empirical approaches alone, it is difficult to  
63 understand the underlying mechanisms of how climatic conditions affect individual physiological processes, and  
64 consequently which role these processes play in the observed changes in C balance and growth at the individual as well as

65 community level. Such approaches are thus subject to large uncertainty when used for projections of C balance under climate  
66 change.

67 Most studies on the relationships between C balance and environmental factors for biocrusts are based on laboratory  
68 experiments (e.g. Coe et al., 2012b; Cowan et al., 1992; Lange et al., 1998) or direct field measurements in situ over short  
69 periods of time (e.g. Brostoff et al., 2005; Lange et al., 1994). From this work cited above, it has been recognized that the C  
70 balance of biocrusts is strongly influenced by factors such as water supply, temperature, radiation, and CO<sub>2</sub> concentration and  
71 the complex relations of these factors to physiological processes such as photosynthesis and respiration. While the highest  
72 values of productivity under field conditions are achieved when the environmental factors are in the range that is optimal for  
73 the specific biocrust, it has been found that biocrusts are also able to achieve metabolic activity and thus, potential productivity,  
74 under sub-optimal conditions of temperature and light (Colesie et al., 2016; Raggio et al., 2017, 2014). It is largely unknown,  
75 however, which relative importance each of these environmental factors and physiological parameters has for the long-term C  
76 balance of biocrusts under natural field conditions, and if the importance of factors/parameters shows a spatial and temporal  
77 pattern. In addition, seasonal acclimation of photosynthetic and respiratory properties of species to intra-annually varying  
78 climate factors found by several studies (e.g. Gauslaa et al., 2006; Lange and Green, 2005; Wagner et al., 2014) may  
79 substantially affect biocrust C balance, thus leading to further complexity in the spatio-temporal relations between C balance  
80 and environmental factors/physiological parameters. One of the few experimental studies investigating biocrust acclimation  
81 potential to changing temperatures has found threshold temperatures for the survival of lichen species (Colesie et al., 2018)  
82 but the overall extent of the impact is poorly understood.

83 Here, we applied a mechanistic data-driven model to (a) complement empirical estimates of the annual C balance of biocrusts  
84 and (b) to address the knowledge gaps concerning the relative importance of different environmental factors and physiological  
85 parameters for the C balance in contrasting climates, thereby accounting for the role of seasonal acclimation. The advantage  
86 of this modeling approach is that it can predict at high temporal resolution the dynamic C balance of biocrust organisms for  
87 given locations by simulating the physiological processes driven by environmental factors. The model allows for a deeper  
88 mechanistic understanding of the C balance of biocrusts through factorial experiments and sensitivity analyses regarding  
89 physiological parameters and individual environmental factors, which would be impractical to realize in field or laboratory  
90 experiments. To complement our analyses using the data-driven model, a process-based dynamic non-vascular vegetation  
91 model, called LiBry, was employed as an supporting tool (Porada et al., 2013).

## 92 **2 Material and Methods**

93 We simulated the C balance of biocrusts from six climatically different study sites in a semi-empirical way using a data-driven  
94 model. The model simulates photosynthetic rate based on the Farquhar photosynthesis model developed by Farquhar and von  
95 Caemmerer (1982) and respiration rate based on a Q10 relationship. The C balance is computed as the difference of  
96 photosynthesis and respiration accumulated over a given time period. In the model, both photosynthesis and respiration depend

97 on surface temperature, relative water saturation, and the activity of the biocrust, which are all simulated in a coupled way via  
98 the surface energy balance as a function of climate input data. Photosynthesis additionally depends on ambient CO<sub>2</sub>  
99 concentration.

100 To calibrate the model, we first determined soil and land surface properties that are required for the coupled energy and water  
101 balance in the model through fitting simulated to measured surface temperature patterns. Then, we parameterized the  
102 physiological properties of the organisms using measured relationships between net photosynthesis and light intensity, water  
103 content, and temperature. Finally, we validated the model with regard to the water content or activity patterns of biocrusts and  
104 compared the modelled CO<sub>2</sub> assimilation rate to measured values. The data sets used for calibration and validation of the model  
105 as well as basic climate conditions of each site are described in Table 1. Sites were listed in ascending order of total annual  
106 precipitation based on measured data.

## 107 **2.1 Study sites**

108 In our study we considered six sites, namely two dryland sites at Almeria (Spain) and Soebatsfontein (South Africa; hereafter  
109 D1 and D2); three temperate sites at Gössenheim (Germany), Öland (Sweden), and Linde (Germany; Hereafter T1, T2, and  
110 T3); and one alpine site at Hochtorn (Austria; Hereafter A1) (Table 1). These sites were chosen based on data availability for C  
111 balance estimation, and because they cover a broad range of climatic conditions. The field and laboratory measurements  
112 conducted at all sites were following a similar protocol, which allows comparing the simulation results among sites. The  
113 necessary empirical data for C balance estimation regarding climatic conditions, species physiological characteristics, and  
114 status especially in terms of moisture such as water content or activity, have been monitored in a relatively small number of  
115 experiments, so far, and the six study sites chosen here to provide a good opportunity to utilize these data for an extended  
116 modeling approach. In this context, activity measurements are more suitable than soil moisture records since they are direct,  
117 non-invasive and they do not show deviations in the temporal patterns at high resolution, which may occur with soil moisture  
118 time series.

119 Sites D1 and D2 are characterized by an arid climate with mean annual precipitation of less than 250 mm, but a wet winter  
120 season (Büdel et al., 2014; Haarmeyer et al., 2010). Sites T1, T2, and T3 have a temperate climate. The mean precipitation in  
121 these three sites is around 550 mm (Büdel et al., 2014; Diez et al., 2019). Site A1 is located in a humid alpine region with a  
122 mean annual precipitation between 1750 and 2000 mm, of which more than 70% are snowfall; the A1 site is covered by snow  
123 for at least 200 days per year (Büdel et al., 2014). More detailed site descriptions are provided in the corresponding studies  
124 cited above.

## 125 **2.2 Observational Data**

### 126 **2.2.1 Climatic variables**

127 The proposed data-driven model for estimating the annual C balance of dominant biocrust types at each site was forced by

128 hourly microclimatic variables. The forcing data sets of the data-driven model include photosynthetically active radiation  
129 (PAR), long-wave radiation (near-infrared), relative air humidity, air temperature, wind speed, rainfall, and snowfall. All the  
130 microclimatic variables were measured on-site by climate stations with a temporal resolution of 10 min (5 min in A1 and D1;  
131 data available at <http://www.biota-africa.org>; Raggio et al., 2017; M. Veste, unpublished data), except for long-wave radiation  
132 and snowfall, which were taken from ERA5 dataset (<https://www.ecmwf.int/en/forecasts/datasets/reanalysis-datasets/era5>).  
133 Although directly measured surface temperature data are available for all sites, we use ERA5-based down-welling long-wave  
134 radiation instead to simulate surface temperature on biocrusts. This is necessary since, in our model, calculations of  
135 photosynthesis and respiration require not only surface temperature, but also depend on water saturation of biocrusts (affecting  
136 activity). However, we do not have water saturation data available at most sites. Therefore, we instead simulate the dynamic  
137 water saturation of biocrusts based on climate, via processes such as evaporation, rainfall, and dew. The calculation of  
138 evaporation and dew automatically includes the computation of a surface temperature that emerges from solving the surface  
139 energy balance, thereby including down-welling long-wave radiation. Since the simulated surface temperature that is  
140 connected to simulated water saturation slightly deviates from the observed surface temperature (see Fig. 1 and S1), we do not  
141 directly use the observed surface temperature as input in the modeling approach, to avoid inconsistencies. Then all these  
142 microclimate data were aggregated to data with hourly temporal resolution.

### 143 **2.2.2 Dynamic biocrust variables**

144 Besides the surface temperature, biocrust activity was either monitored directly (binary: active or not active) using a continuous  
145 chlorophyll fluorescence monitoring system (Raggio et al., 2014, 2017), or indirectly via the electrical conductivity of the  
146 substrate (BWP, Umweltanalytische Produkte GmbH, Cottbus, Germany; Weber et al., 2016; M. Veste, unpublished data).  
147 For site D2, the biocrust water content was calculated instead of activity based on electrical conductivity. Due to snow covering  
148 the measuring instruments, data of site A1 only covers the time from August to October. Samples from both lichen- and also  
149 moss-dominated biocrusts were measured at all sites, except for site T3 where four BWPs were mostly located in moss-  
150 dominated biocrusts. At site D2, additionally cyanolichen- and cyanobacteria-dominated biocrusts were monitored. The  
151 measured surface temperature, water content, and activity data at all sites were then aggregated to data with a temporal  
152 resolution of one hour.

153 As explained in Sect. 2.2.1, we did not directly use the observed surface temperature and activity (or water content) as forcing  
154 data for the model, but estimated the time-series of surface temperature and water saturation data at all sites based on a  
155 simulation of the energy and water balance. The activity of the organisms was then approximated via the empirical equations  
156 describing the link between water saturation and metabolic activity (see Porada et al., 2013). Furthermore, ambient CO<sub>2</sub>  
157 concentration was assumed to be constant at 400 ppm. The CO<sub>2</sub> concentration at the soil surface may be higher than 400 ppm  
158 due to the flux of respired CO<sub>2</sub> from the soil. Since our study sites are on open ground, we do not assume substantial  
159 accumulation of CO<sub>2</sub> in the near-surface boundary layer. We discuss the effect of uncertainties in CO<sub>2</sub> concentration below in  
160 Sect. 4.2.

161 For validation of C balance, we used data of the on-site CO<sub>2</sub>-exchange rate of different biocrust types (lichen- and moss- and  
 162 also cyano-dominated biocrusts removed from surplus soil; the latter composed of cyanolichen and cyanobacteria) that were  
 163 measured by a portable gas exchange system at several time intervals from November 4<sup>th</sup> to 8<sup>th</sup> at site D2 (Tamm et al., 2018).  
 164 For the other sites, additional field measurements of CO<sub>2</sub>-exchange were not available.

### 165 2.2.3 Photosynthesis response and water storage

166 For all sites, CO<sub>2</sub> exchange measurements under controlled conditions in the laboratory or in the field (site T3) were conducted  
 167 using a mobile gas exchange system GFS 3000 (Walz GmbH, Effeltrich, Germany) with an infrared-gas analyzer to explore  
 168 the physiological characteristics of samples of different biocrust types (same as those measured for validation; main species  
 169 see Table 1; Diez et al., 2019; Raggio et al., 2018; Tamm et al., 2018). Before measurements, the soil underneath these biocrust  
 170 samples was removed up to the amount necessary to preserve the physical structure of the biocrusts. And the samples were  
 171 subjected to reactivation for at least two days (D2) or three days (T1, T2, D1, A1). At T1, T2, A1 and D1, for instance, samples  
 172 were kept at 12°C under 12 h dark and 12 h light (100 μmol m<sup>-2</sup> s<sup>-1</sup>) conditions for three days and wetted once a day. Net  
 173 photosynthesis was measured at different ranges of light intensity, water content, and temperature. Light response curves, for  
 174 instance, were determined at optimum water saturation and 15 °C, water response curves were measured at 400 μmol m<sup>-2</sup> s<sup>-1</sup>  
 175 and 15 °C at sites D1, T1, T2, and A1 (Raggio et al., 2018). Moreover, the maximum water storage capacity (MWC) of the  
 176 samples was quantified in the laboratory for samples from sites D1, T1, T2, and A1 (Raggio et al., 2018), whereas the MWC  
 177 at site D2 was approximated as the maximum value when measuring water response curves (Tamm et al., 2018; Weber et al.,  
 178 2012). MWC at site T3 was estimated as the value of the same genus measured in Hamburg, Germany (*Cladonia portentosa*  
 179 and *Polytrichum formosum*, Petersen et al., in prep.). MWC was acquired since it is one of the essential parameters of the  
 180 model to convert the specific water content in mm to relative water saturation required by the model used here.

181 **Table 1:** Properties of the study sites and data which are available (+ sign) for calibration and validation of the data-driven  
 182 model.

Site	Almeria, Spain	Soebatsfontein, South Africa	Gössenheim, Bavaria, Germany	Öland, Sweden	Linde, Brandenburg, Germany	Hochtor, Austria
Code	D1	D2	T1	T2	T3	A1
Climate	arid	arid	Temperate	Temperate	Temperate	Alpine
Measured annual rainfall [mm]	110	141	424	441	449	744

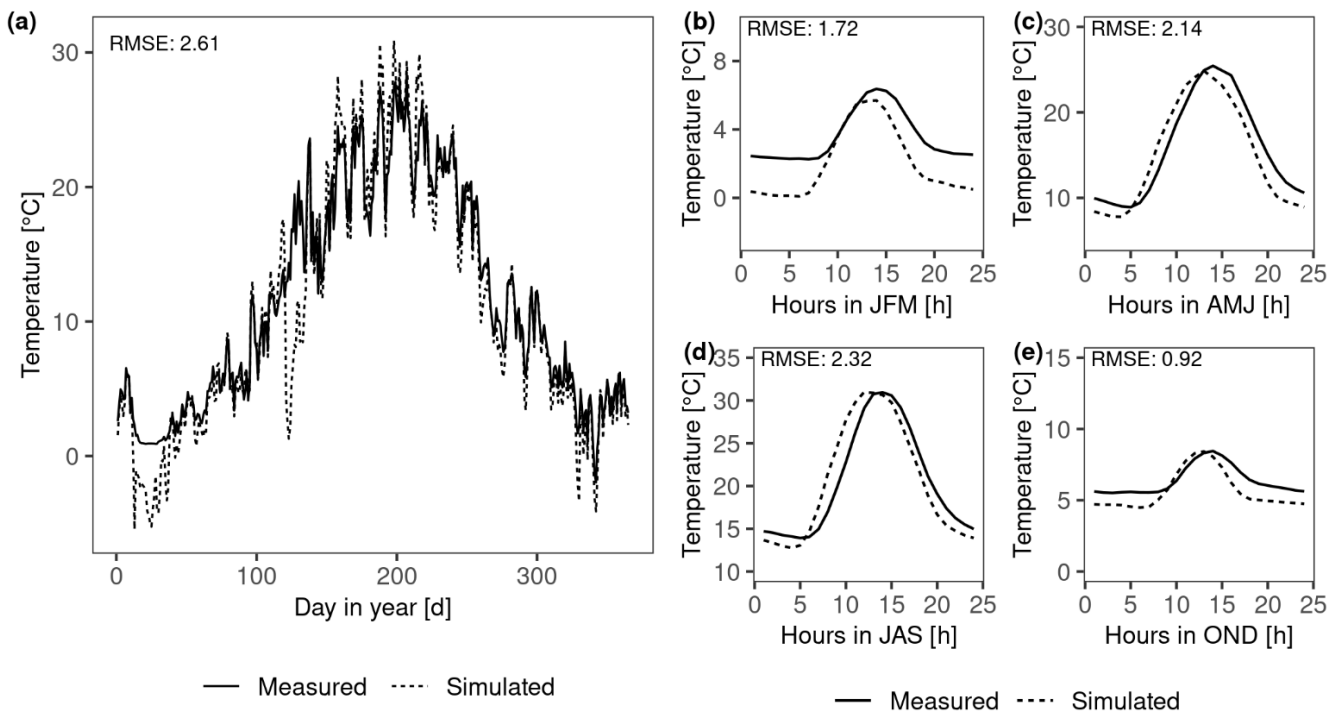
Dominant species at the site	<i>Psora decipiens</i> , <i>Didymodon rigidulus</i>	<i>Psora decipiens</i> , <i>Psora crenata</i> , <i>Ceratodon purpureus</i> , <i>Collema coccophorum</i>	<i>Psora decipiens</i> , <i>Trichostomum crispulum</i>	<i>Psora decipiens</i> , <i>Tortella tortuosa</i>	<i>Cladonia furcata</i> , <i>Polytrichum piliferum</i>	<i>Psora decipiens</i> , <i>Tortella rigens</i>
Data for Calibration	Laboratory CO <sub>2</sub> exchange response curves	Light, water, temperature	Light, water, temperature	Light, water, temperature	Light, water, temperature	Light, water, temperature
	Surface temperature	+	+	+	+	+
Data for Validation	Water content	-	+	-	-	-
	Activity	+	-	+	+	+
	CO <sub>2</sub> exchange on site	-	+	-	-	-
References	Raggio et al., 2018	Tamm et al., 2018; Weber et al., 2012	Raggio et al., 2018	Raggio et al., 2018	Veste, unpublished data; Diez et al. 2019	Raggio et al., 2018

183 **2.3 Parameterization of the data-driven model**

184 **2.3.1 Abiotic surface properties**

185 Several abiotic parameters of the data-driven model describing soil and land surface properties, such as roughness length or  
 186 soil thermal conductivity, were required to simulate the energy and water balance. These parameters were obtained by fitting  
 187 the daily and diurnal surface temperature patterns of lichen-dominated biocrust at all sites except for site T3. At site T3, we  
 188 compared the surface temperature patterns of simulated moss-dominated biocrusts to data collected by sensors in four  
 189 locations.

190 The set of parameters that corresponded to minimum differences between simulated and measured values (visual assessment)  
 191 was used in the data-driven model. The calibration results of surface temperature and the photosynthesis response curves at  
 192 site T2 are shown in Fig. 1 and Fig. 2, respectively. The results of dominant species at other sites are shown in Fig. S1 and S2.  
 193 The daily surface temperature was simulated accurately (the maximum value of root mean square error (RMSE) at all sites is  
 194 3.78) except for site T3 where the temperature during cold seasons was underestimated, and at site D1 the peak temperature  
 195 within a day in hot seasons was underestimated (Fig. S1). The peak in surface temperature occurred too early by around 1-2  
 196 hours at site D1, A1, T1 and T2, but the magnitude of the peak corresponded well to the measured data (Fig. 1 and S1).  
 197 Therefore, in general, the fitting of the surface temperature patterns was satisfactory. The mismatches may result from the  
 198 measured climate variables such as PAR or air temperature at 2 m being inconsistent with the measured surface temperature  
 199 (detailed descriptions are in Supplement) and bias in the calibrated soil properties such as soil thermal conductivity since it  
 200 may affect the surface temperature difference between morning and evening. A sensitivity analysis of soil thermal conductivity  
 201 to C balance is shown in the Supplement.



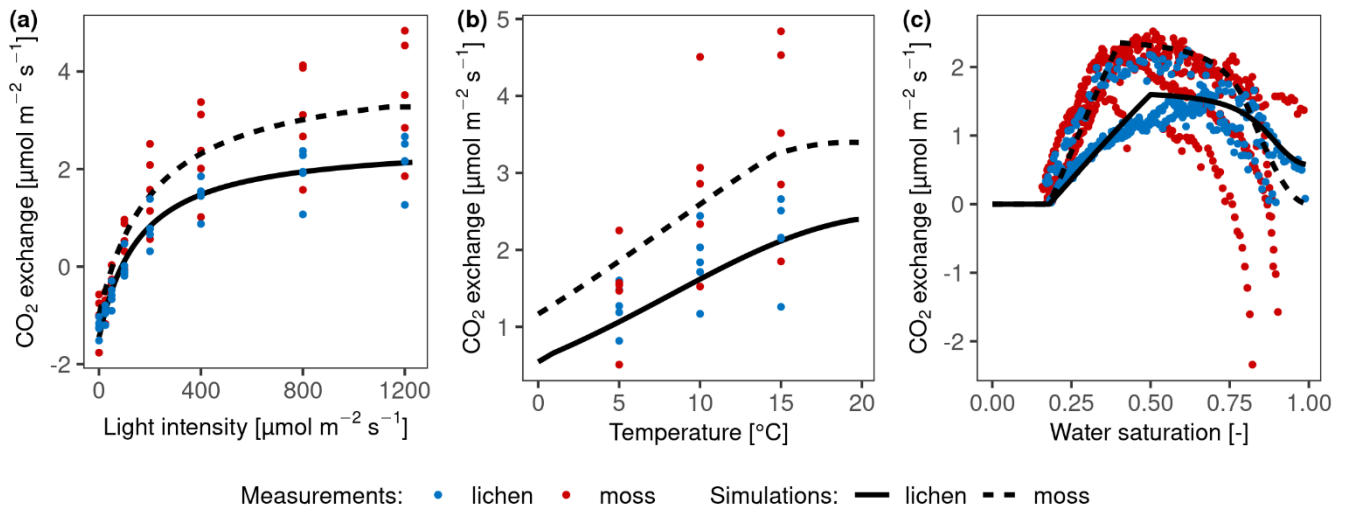
202

203 **Figure 1:** Calibration results of abiotic parameters of the data-driven model by fitting the daily (a) and diurnal (b-e) patterns  
 204 of surface temperature at site T2. (b) to (e) represent the patterns of average hourly surface temperature from January to March  
 205 (JFM), April to June (AMJ), July to September (JAS), and October to December (OND), respectively.

206 **2.3.2 Biocrust physiological properties**



207 Furthermore, several parameters required by the Farquhar photosynthesis scheme and the respiration scheme were directly  
 208 measured or calculated from gas exchange data sets, such as MWC of the thallus, optimum water content, the optimum  
 209 temperature for gross photosynthesis, metabolic respiration cost, and the Q10 value of respiration. Since the ambient  
 210 temperature range that was applied in the laboratory for samples from all sites except D2 was too small to capture the optimum  
 211 temperature of photosynthesis reliably, we approximated the optimum temperature from the measured data set as the average  
 212 surface temperature during active periods. In addition, the optimum temperature was also constrained by fitting the Farquhar  
 213 equations to photosynthesis curves, as related to environmental factors light, water content, and temperature. Such fitting  
 214 method was also used to obtain some other photosynthesis-related parameters of organisms, such as molar carboxylation and  
 215 oxygenation rate of RuBisCO ( $V_{cmax}$ ,  $V_{omax}$ ), respiration cost of RuBisCO enzyme ( $Rub\_ratio$ ), and minimum saturation  
 216 for activation ( $Sat\_act0$ ).



217

218 **Figure 2:** Calibration of photosynthesis parameters of the model by fitting photosynthesis response curves of moss- and lichen-  
 219 dominated biocrust samples to measurements at site T2. (a): net photosynthesis rate in response to light at optimum water  
 220 content and 15 °C. (b): net photosynthesis rate in response to temperature at 1200  $\mu\text{mol m}^{-2} \text{s}^{-1}$  light and optimal water content.  
 221 (c): net photosynthesis rate in response to relative water saturation at 400  $\mu\text{mol m}^{-2} \text{s}^{-1}$  light and 15 °C.

222 Since the measured values between replicates showed large variation, in particular the water and temperature-dependent data,  
 223 as illustrated by the coloured points in Fig. 2, we fitted the curves to the average values of all replicates. The calibration results  
 224 showed that visually and overall, the photosynthetic curves could be parameterized to fit the measurements well, given that  
 225 different samples were used for measuring responses to different driving factors, and considering the methodological  
 226 differences between light and water response curve measurements. However, the water responses were least well-fitted,  
 227 especially at high water contents. The measured net photosynthesis response rate was negative in some cases at high water  
 228 saturation, but it is not possible to reproduce this negative net photosynthesis rates with our adapted Farquhar photosynthetic

229 model for the light and CO<sub>2</sub> conditions of the laboratory setup. Under these conditions, gross photosynthesis is larger than  
230 respiration and thus CO<sub>2</sub> is required to diffuse from the atmosphere into the thallus, not out of it. Even under low diffusivity,  
231 caused by high water saturation, there will be no net diffusion of CO<sub>2</sub> from the thallus into the ambient air assuming that  
232 inward and outward flows of CO<sub>2</sub> share the same pathway and that diffusion of CO<sub>2</sub> between atmosphere and thallus is in  
233 steady-state with the flux (respiration minus gross photosynthesis). For details please see the Supplement. Furthermore,  
234 matching the simulated temperature-response of net photosynthesis to the measured data under cool conditions turned out to  
235 be difficult for samples from site T1 and A1 (Fig. S2). There were too few data points in the measured temperature response  
236 data set to constrain the optimum temperature and temperature relation (see Fig. 2), but the fitting turned out well because the  
237 simulated optimum temperature for net photosynthesis at site T1 was coincidentally close to the value of 17 °C reported by  
238 Colesie et al. for this site. (2014; see Fig. S2).

## 239 **2.4 Validation of the data-driven model**

240 The water saturation and activity estimated by the data-driven model were validated by comparing with the daily and diurnal  
241 patterns of measured activity (5 sites, not at site D2) and water content data (only at D2). Furthermore, the C balance estimated  
242 by the data-driven model was validated by comparison to the in situ measured net CO<sub>2</sub> exchange rate of moss, lichen, and  
243 cyanocrust-dominated biocrusts. These data were sampled at site D2 by removing the soil respiration rate, predicted by means  
244 of a fitted regression (Weber et al., 2012). Since data on water saturation were available, measured PAR, surface temperature,  
245 and water content were used to simulate the C balance using the data-driven model, in contrast to the setup described above.  
246 The activity, however, was calculated in the same way as described in the setup. Moreover, the parameters of the model were  
247 the same as the calibrated ones of the corresponding biocrust types.

## 248 **2.5 Sensitivity analysis**

249 To investigate the role of environmental factors, physiological properties, and also seasonal acclimation for the simulated  
250 annual C balance of biocrusts, we conducted three sensitivity analyses using our data-driven model. With this setup, we intend  
251 to put into context the effects of environmental conditions and the uncertainty associated with the physiological properties that  
252 were used to parameterize the model. We additionally explore the impact of seasonally acclimatized physiological properties  
253 on carbon assimilation at site T1, since variation between seasons represents additional uncertainty in the estimation of the C  
254 balance.

### 255 **2.5.1 Effects of environmental factors**

256 To investigate the role that environmental factors, namely air temperature (T<sub>air</sub>), light intensity (Light), ambient CO<sub>2</sub>  
257 concentration (CO<sub>2</sub>), and different types of water sources play in regulating the C balance of biocrusts, sensitivity analyses  
258 were conducted for lichen-dominated biocrusts from all study sites. The different types of water sources include rainfall (Rain)  
259 and non-rainfall water inputs such as dew and water vapor, which are determined by relative air humidity (R<sub>hum</sub>). All the

260 environmental factors were reduced and increased by half (+/- 50%), except for T<sub>air</sub> and R<sub>hum</sub>. The T<sub>air</sub> differences varied  
261 by 5 K and R<sub>hum</sub> by 20%. Moreover, relative humidity was constrained between 0 and 100% when the varied relative  
262 humidity exceeded this range.

263 The annual C balance for each modified environmental factors was then normalized following Eq. (1), and normalized again  
264 among different environmental factors within each climatic zone for comparing the relative importance of environmental  
265 factors:

$$266 \text{ Normalized } C \text{ balance} = \frac{C_{ij} - C_j}{|C_j|}, \quad (1)$$

267 where C<sub>ij</sub> is the C balance of factor j under operation i, and C<sub>j</sub> is the original C balance of factor j.

268 A positive normalized C balance demonstrates an increase in annual C balance when certain environmental factors change,  
269 and a larger magnitude of the normalized C balance number demonstrates a larger effect of this environmental factor compared  
270 to a factor with a smaller value.

### 271 **2.5.2 Effect of physiological parameters**

272 The sensitivity analysis of physiological parameters was conducted for lichen-dominated biocrust at all study sites. The original  
273 parameter values were obtained by calibration to measured net photosynthesis response curves. We then varied the values of  
274 the following physiological parameters by a consistent range for all sites: metabolic respiration cost per surface area  
275 (Resp<sub>main</sub>), Q10 value of respiration (q10), the optimum temperature for gross photosynthesis (Topt), respiration cost of  
276 RuBisCO enzyme (Rub<sub>ratio</sub>), and light absorption fraction in cells (ExtL), minimum saturation for activation (Sat<sub>act0</sub>), and  
277 minimum saturation for full activation (Sat<sub>act1</sub>). Specifically, we increased or decreased Resp<sub>main</sub>, ExtL, q10, Sat<sub>act0</sub> by  
278 30%, Rub<sub>ratio</sub> and Sat<sub>act1</sub> by 20%, and Topt by 5 K. These parameters are chosen since they are closely related to the  
279 response of photosynthesis and respiration to water, light, and temperature. These ranges of different parameters were  
280 determined based on the observed bounds of the photosynthetic response curves of all replicates (see calibration results with  
281 varied parameters at all sites in Fig. S5-S10), which have large deviations between each other at most sites as shown in Fig. 2  
282 and Fig. S2. The effects of the varied physiological parameters on the C balance were then normalized using the same  
283 normalization method as for the environmental factors (in Sect. 2.5.1) for comparison among parameters and climatic zones.

### 284 **2.5.3 Effect of seasonal acclimation**

285 Another sensitivity analysis was performed for site T1 to investigate the impact of seasonally acclimatized physiological  
286 properties on the C balance. We analysed the lichen- and moss-dominated biocrusts at site T1 as an example, because the  
287 measured time-series of activity showed that in temperate sites such as T1, the organisms were active most of the time, and  
288 thus the C balance would be more sensitive to seasonally varying properties.

289 In the analysis, rather than keeping all calibrated parameters fixed throughout the simulation period of the data-driven model,  
290 the physiological parameters metabolic respiration cost per surface area (Resp\_main), light absorption fraction in cells (ExtL),  
291 and the ratio of Jmax to Vcmax (jvratio) were set to another set of values in the winter months in order to adapt to the climatic  
292 conditions, since biocrusts at sites T1 were collected in summer months. These new, “dynamic” parameters were applied in an  
293 additional simulation and the resulting C balance was compared to the original simulation based on the “fixed” parameters.  
294 The dynamic parameters were chosen and varied based on the literature: Respiration of lichens was found to acclimate to  
295 seasonal changes in temperature (Lange and Green, 2005). Moreover, under low light, organisms showed shade-adapted  
296 physiological characteristics with low PAR compensation and saturation points (LCP and LSP; Green and Lange, 1991). These  
297 properties can be expressed by certain parameters of the data-driven model. For instance, the respiration rate is determined by  
298 the parameter Resp\_main; LCP and LSP can be affected by changing the slope of the photosynthesis-light relations through  
299 the parameter ExtL; LCP and LSP can also be modified via the parameter jvratio as it influences the value of light use efficiency  
300 at unsaturated light.

301 Accordingly, in an hourly simulation during September and December, January, and February, the parameter Resp\_main was  
302 reduced to half to lighten the respiratory cost for the samples collected at site T1. The size of ExtL was doubled to increase the  
303 slope of photosynthesis-light relations. In addition, the parameter jvratio was doubled as well to enhance the light use  
304 efficiency.

## 305 **2.6 LiBry Model**

306 LiBry is a process-based dynamic global vegetation model (DGVM) specific to non-vascular vegetation. The model mimics  
307 environmental filtering in the real world by simulating many different functional strategies and selecting those which maintain  
308 a positive C balance under the respective climatic conditions. The strategies are characterized by a combination of 11  
309 physiological and morphological parameters. More information about the model is briefly described in the Supplement, and a  
310 full detailed description can be found in Porada et al. (2013, 2019).

311 For this study, the LiBry model was run for 300 years with repeated microclimate forcing data of one year from the six study  
312 sites, calibrated abiotic parameters same as the data-driven model, and initially generated 1000 strategies. C balance and  
313 dynamics of the surface cover of the strategies were simulated until a steady state was reached, so that the final successful  
314 strategies were those where long-term biomass values were positive. Moreover, at the end of the simulation, the average values  
315 of functional traits were estimated by weighting all surviving strategies based on their relative cover. The (hypothetical)  
316 strategy characterized by these average values is called average strategy. The strategy with the largest cover area is called  
317 dominant strategy.

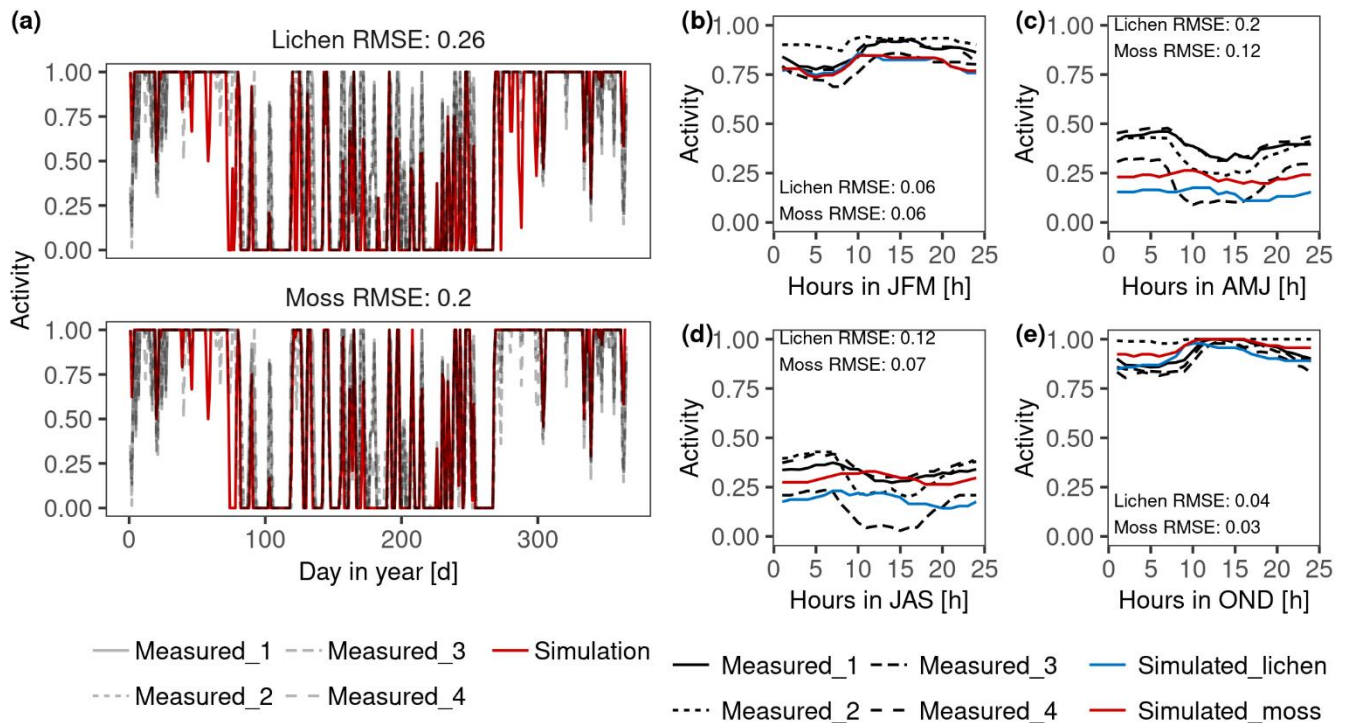
318 Furthermore, we compared the physiological parameters of the average strategy and the selected dominant strategies to the  
319 ones of organisms in the field by means of their respective photosynthesis response curves. This comparison can verify the C  
320 balance estimated by the data-driven model from a reversed perspective as LiBry model is based on the same processes as the

321 data-driven model, but the strategies were freely selected by the LiBry model based on their C balance, without prescribing  
 322 values based on site level observations.

### 323 3 Results

#### 324 3.1 Data-driven model

##### 325 3.1.1 Validation of the data-driven model

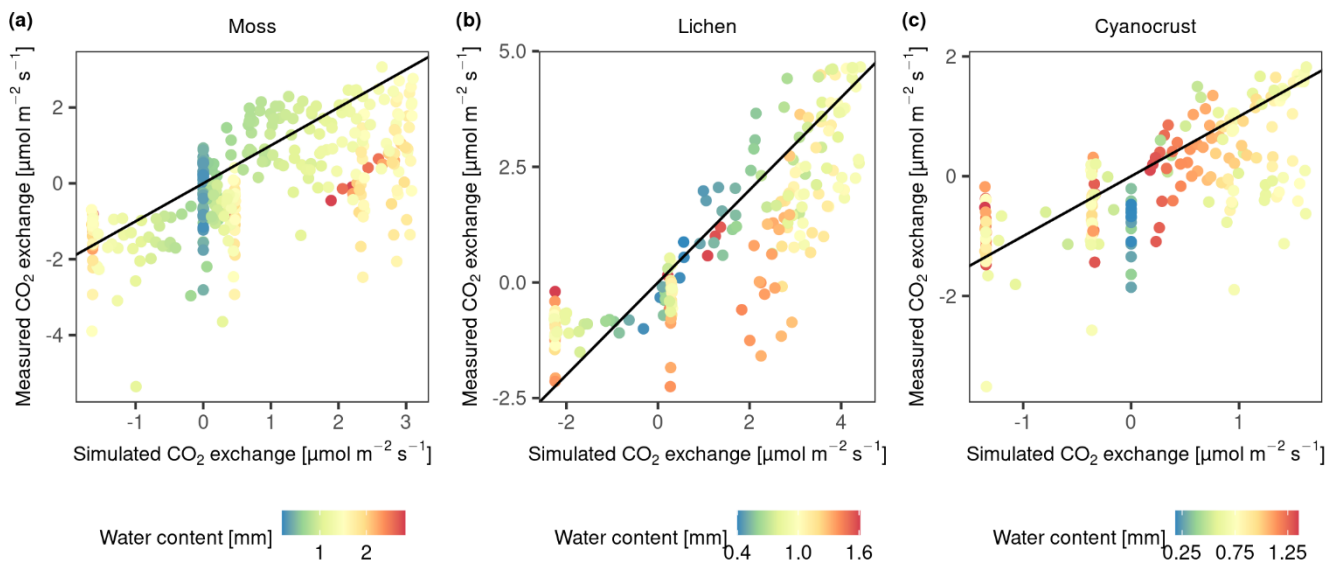


326

327 **Figure 3:** Validation of the estimated daily (a) and diurnal (b-e) patterns of activity of lichen- and moss-dominated biocrusts  
 328 at site T3. The simulated patterns of activity were compared to measured data by four sensors at different locations. (b) to (e)  
 329 represent the patterns of average hourly activity from January to March (JFM), April to June (AMJ), July to September (JAS),  
 330 and October to December (OND), respectively.

331 In general, the simulated daily and diurnal patterns of activity (water content at site D2) fit the measurements reasonably well  
 332 in magnitude (Fig. 3 and Fig. S3). However, our fitting resulted in a more dampened diurnal activity pattern simulated by the  
 333 model, and the activity at night and in the morning was underestimated during several seasons at sites D1, T1, T2 and A1. In  
 334 addition, both the daily and diurnal activity during April and June at site T1 were underestimated. Furthermore, water content  
 335 was overestimated for moss-dominated biocrust, especially when there was a large amount of water input at D2, although the

336 patterns corresponded well to the measured data for all three biocrust types (Fig. S3). This overestimation may have resulted  
 337 from the bias in measured MWC of samples used for constraining the water content in the model.  
 338 The comparison of simulated and on-site measured CO<sub>2</sub> exchange rates of three biocrust types (moss, lichen, and cyanocrust  
 339 composed of cyanolichens and cyanobacteria) at site D2 showed mismatches, especially when water saturation was at both  
 340 ends of the gradient (Fig. 4). The CO<sub>2</sub> exchange rate at high water content was overestimated compared to the measurements.  
 341 Moreover, there were large variations in measurements of respiration and CO<sub>2</sub> exchange rate as water content was low and  
 342 thus simulated CO<sub>2</sub> exchange rate was zero. Excluding the values at both ends of water content (0.58 and 1.74 mm for moss;  
 343 0.22 and 0.68 mm for lichen; 0.65 and 1.24 mm for cyanocrust), the accuracy of the model predictions was improved (root  
 344 mean square error (RMSE) decreased from 1.44 to 1.36 for moss, 1.27 to 0.65 for lichen, and 0.79 to 0.77 for cyanocrust).  
 345 Furthermore, the simulations were similar to measurements in magnitude. Therefore, despite the large variation, we are  
 346 confident about the general validity of the model.



347  
 348 **Figure 4:** Validation of the photosynthesis and respiration scheme of the data-driven model through comparison of modelled  
 349 and measured CO<sub>2</sub> exchange rate of moss, lichen, and cyanocrust given the measured water content, surface temperature, PAR,  
 350 and calculated activity. Observational data were collected in November in 2013 at site D2. The 1:1 line is shown in black.

### 351 3.1.2 Estimated C balance by data-driven model

352 The simulated annual C balance of each collected biocrust type at each site is listed in Table. 2. The annual C balance of lichen-  
 353 and moss-dominated biocrusts at dryland D1 showed a small positive value. Moreover, the moss-dominated biocrust at dryland  
 354 D2 gained small amount of carbon while lichen-dominated biocrust and a cyanocrust additionally measured at site D2 showed  
 355 a small net release of carbon in the model.

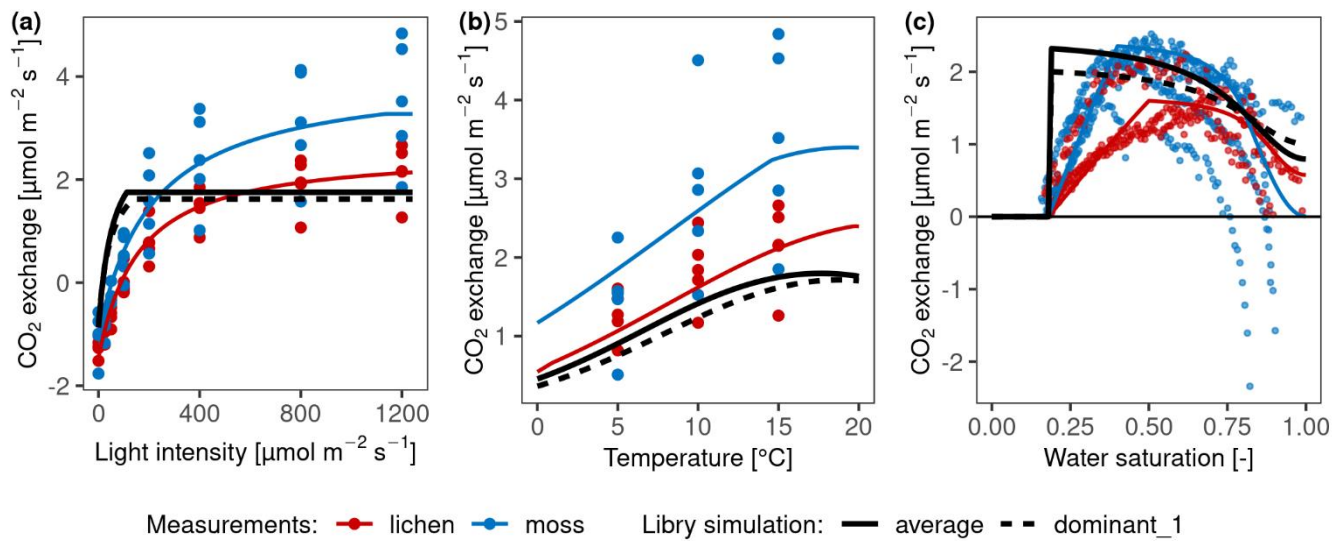
356 **Table 2:** Simulated annual carbon budgets of each biocrust type at all sites.

	Lichen	Moss	Cyanocrust
	g C m <sup>-2</sup> yr <sup>-1</sup>	g C m <sup>-2</sup> yr <sup>-1</sup>	g C m <sup>-2</sup> yr <sup>-1</sup>
D1 (Almería)	1.5	1.9	
D2 (Soebatsfontein)	-1.7	3.1	-8.3
T1 (Gössenheim)	-42.8	-39.4	
T2 (Öland)	-92.1	-74.7	
T3 (Linde)	9.4	18.7	
A1 (Hochtor)	-17.9	-6.8	

357 Furthermore, according to these data-driven model simulations, despite the C balance of two biocrust types being positive at  
358 site T3, a large amount of carbon was lost at the sites T1 and T2 in temperate humid regions. These results imply that according  
359 to the data-driven model, the biocrusts would not survive in the long-term at most of the temperate humid research sites. At  
360 the alpine site A1, both lichen- and moss-dominated biocrust lost carbon in a year with long periods of ice cover.

### 361 **3.2 Dominant strategies selected by the LiBry Model**

362 In general, the photosynthesis response curves of dominant and average strategies selected by the LiBry model did not fit well  
363 to the measurements, especially at temperate site T2 (Fig. 5; the results for the other sites with negative C balance are shown  
364 in Fig. S4). Specifically, the selected physiological traits which determine water and light acquisition of the dominant and  
365 average strategies in LiBry differed from those of the collected samples at all sites. Compared to the measured samples, the  
366 LiBry strategies showed markedly higher efficiency at low light intensity and faster activation. By design, the LiBry model  
367 selected strategies with a positive C balance in the long-term run, and thus the mismatches are consistent with the fact that the  
368 data-driven model simulated negative C balances.



369

370 **Figure 5:** Comparison of net photosynthesis response of measured samples with simulated selected dominant and average  
 371 strategies of LiBry at site T2. (a): light response curve; (b): temperature response curve; (c): water response relation. The  
 372 colored points represent the measured CO<sub>2</sub> exchange rates of moss and lichen, and the coloured lines correspond to the data-  
 373 driven model. The black lines show the photosynthesis response of the dominant strategy selected by the LiBry model (dashed)  
 374 and the average strategy (solid). The parameter values of the average strategy correspond to the average of all surviving  
 375 strategies.

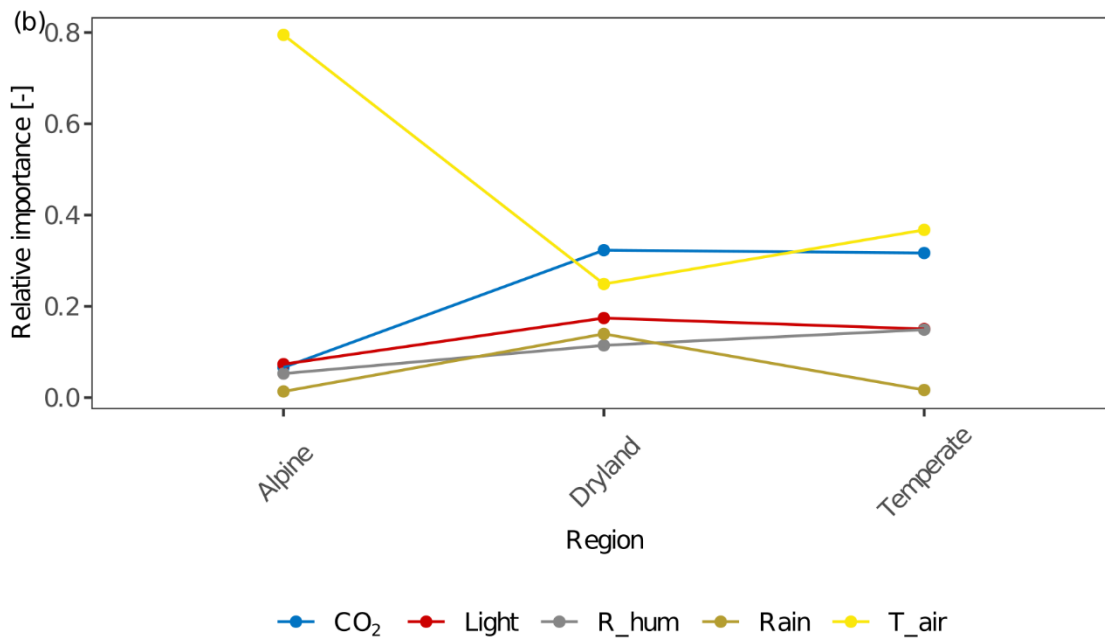
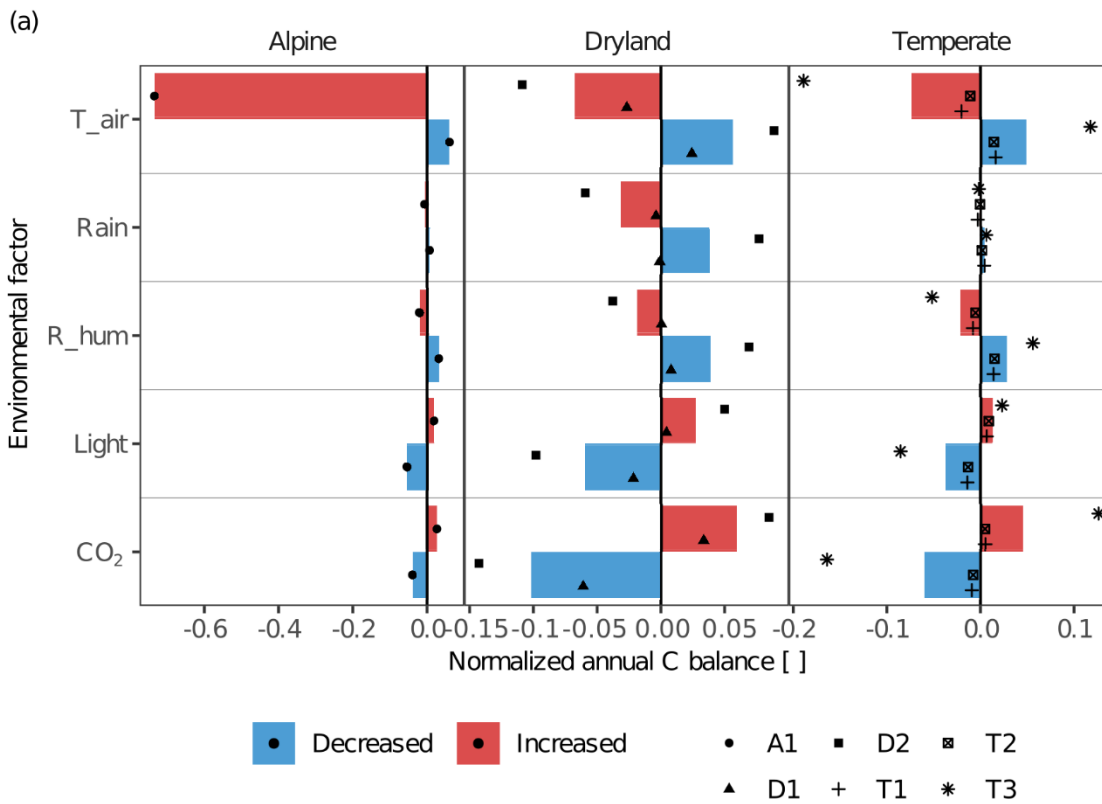
### 376 3.3 Driving factors of variation of the C balance

#### 377 3.3.1 Environmental factors

378 The environmental factors light intensity, CO<sub>2</sub> concentration, air temperature, and various water sources, had different effects  
 379 on the C balance of lichens in different climate zones (Fig. 6). For all sites within a given climate zone, the effects of different  
 380 environmental factors on C balance were overall similar but showed an apparent larger variation at the temperate site T3 in  
 381 contrast to the other two temperate sites, and at site D2 compared to D1 (Fig. 6a). This may be due to physiological differences  
 382 of the investigated biocrust species between these sites and consequently variations in the responses of net photosynthesis rate  
 383 to temperature, water, and light between them (Fig. 2 and S2).

384 Furthermore, the spatial patterns of the relative importance of different environmental factors show that the factors which have  
 385 the strongest effects differ between climatic regions (Fig. 6b).





387 **Figure 6:** (a) The effects of environmental factors - CO<sub>2</sub> concentration (CO<sub>2</sub>), relative air humidity (R\_hum), rainfall amount  
388 (Rain), air temperature (T\_air) and light intensity (Light) on the annual C balance of lichens in different climate regions. The  
389 altered annual C balance resulting from increasing or decreasing environmental factors is normalized by the C balance under  
390 original environmental conditions. The colored columns indicate the average value of the normalized C balance at sites with  
391 similar climate conditions. Various styles of black points indicate different sites. Positive normalized C balance implies that  
392 the annual C balance increases with varying environmental factors and more carbon was accumulated in a year at the site, and  
393 vice versa. A larger normalized C balance reflects that the C balance is more sensitive to the altering environmental factor,  
394 and thus the environmental factor has a larger effect on C balance. (b) Relative importance of each environmental factor  
395 compared to other factors across the climatic regions. Larger relative importance implies a more important effect the factor  
396 has on the C balance compared to other factors in the given climatic region, and vice versa.

397 In general, air temperature and CO<sub>2</sub> concentration were the most important drivers for C balance of biocrust organisms between  
398 climate zones. Light intensity and relative humidity played a relevant role in impacting the C balance as well. Rainfall amount  
399 had lower relative importance at all sites except dryland D2, where the effect of rainfall on C balance was similar to other  
400 factors (Fig. 6a). Therefore, rainfall amount showed a maximum in relative importance in drylands, compared to other regions.  
401 In general, the effect of the other water source, relative air humidity, was moderate but notable at all climate zones, and is  
402 slightly larger in temperate region in comparison to other climate zones. Furthermore, the humidity had a slightly larger impact  
403 on C balance in comparison to rainfall amount at all temperate and alpine sites (e.g., change amplitude was 0.007 for rainfall  
404 and 0.021 for humidity at T1). In drylands, however, the impacts of water sources on C balance varied between sites. The  
405 results showed that relative humidity had a slightly larger impact than rainfall amount at D1 while smaller at D2 (the change  
406 in amplitude at D2 was 0.108 for humidity and 0.137 for rainfall).

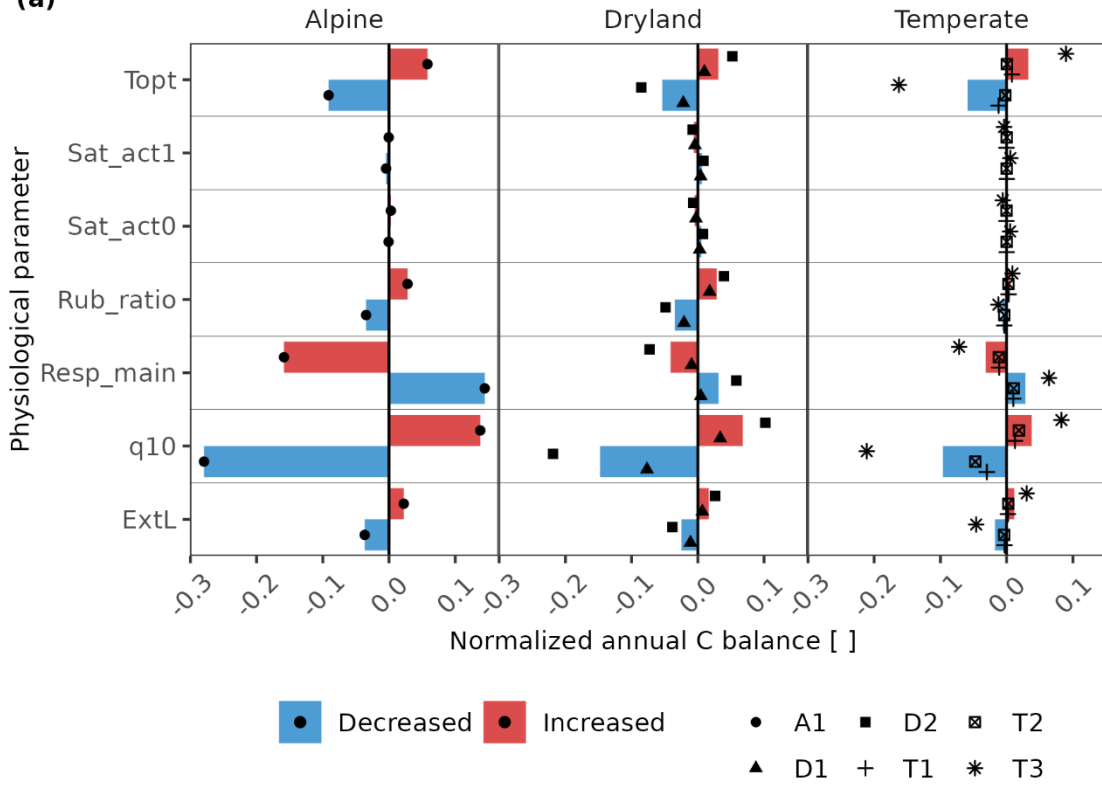
407 The ambient CO<sub>2</sub> concentration was an essential factor for the C balance at all sites especially in drylands, resulting in positive  
408 effects on C balance with increasing CO<sub>2</sub>. Furthermore, light intensity had a marked impact on the C budget at all sites, and it  
409 was relatively more important in drylands and temperate regions than the alpine site. At all sites, the normalized C balance  
410 increased with enhanced light intensity and vice versa (e.g., normalized C balance at T2 are -0.013 and 0.008 for half and  
411 doubled light intensity). Air temperature had a large impact on C balance at all sites. Especially at the alpine site A1, C balance  
412 decreased strongly as air temperature raised by 5 K (normalized C balance of -0.735), and at all sites, the direction of the effect  
413 remained constant, namely, warming decreased the C balance and vice versa.

### 414 **3.3.2 Physiological parameters**

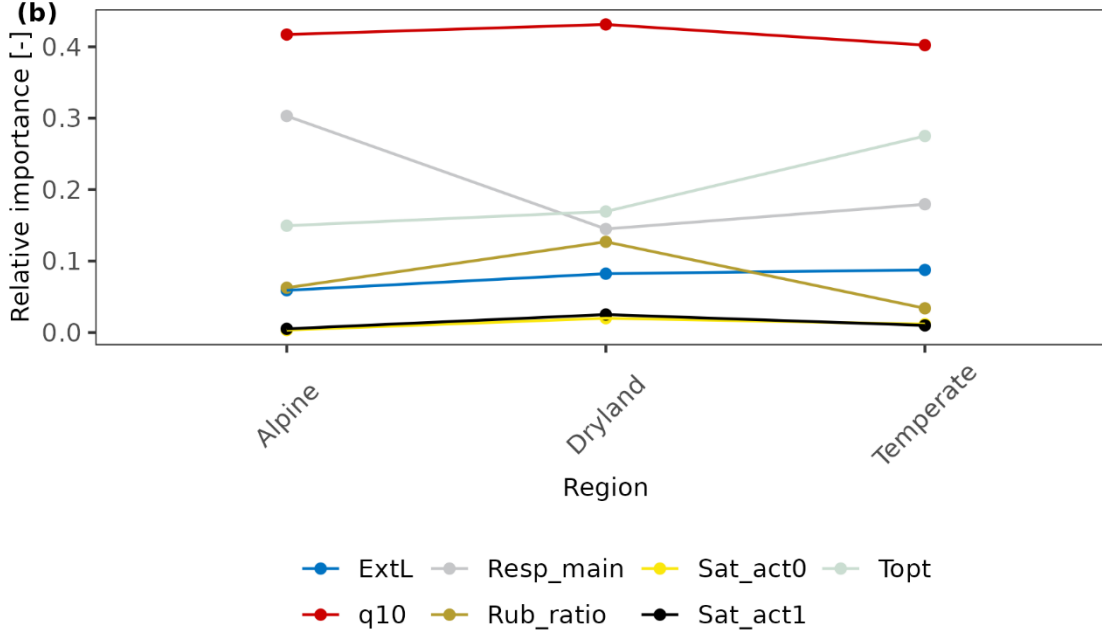
415 We found that physiology plays an important role in all regions. In particular, the respiration-related parameters such as q<sub>10</sub>,  
416 Resp<sub>main</sub>, and T<sub>opt</sub> have a notably higher impact on C balance estimation (Fig. 7). Furthermore, the relative importance of  
417 several physiological parameters showed similar patterns across climatic zones: in all regions, q<sub>10</sub> is the most essential  
418 parameter, Sat<sub>act0</sub> and Sat<sub>act1</sub> play little roles in affecting C balance. Other parameters showed slightly different patterns

419 among regions. Metabolic respiration cost (Resp\_main), for instance, plays a more important role than optimum temperature  
420 for gross photosynthesis (Topt) at the alpine site, while the optimum temperature is more essential in drylands and temperate  
421 regions (Fig. 7b).

(a)



(b)

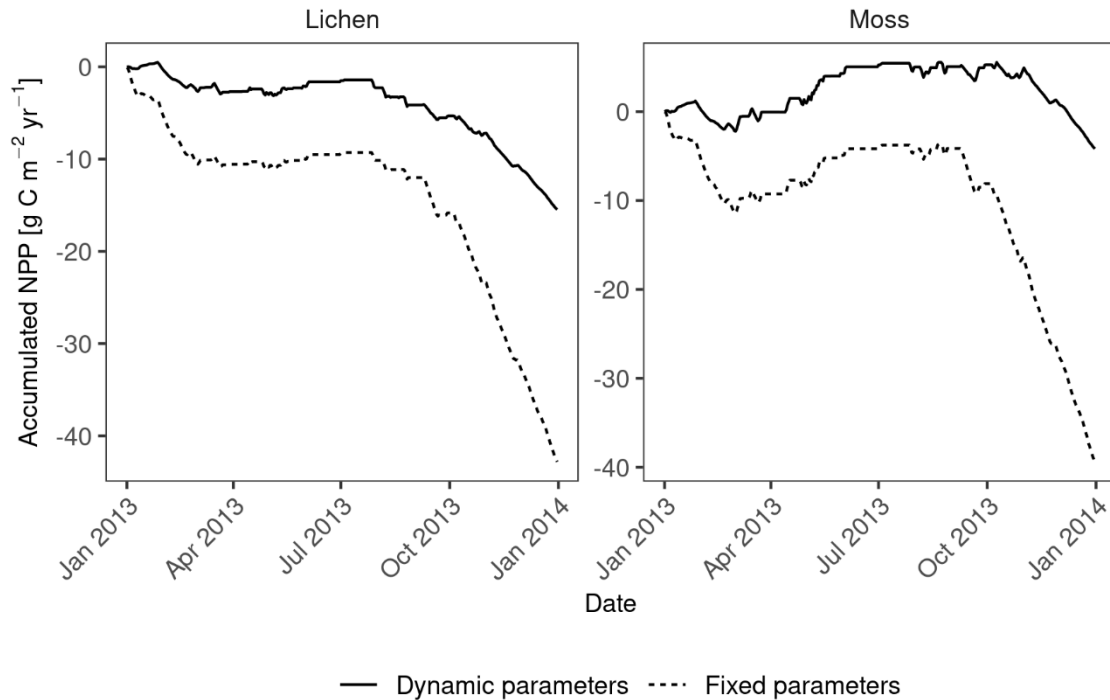


423 Figure 7: (a) The effects of physiological parameters – metabolic respiration cost per surface area (Resp\_main), Q10 value of  
424 respiration (q10), the optimum temperature for gross photosynthesis (Topt), respiration cost of RuBisCO enzyme (Rub\_ratio),  
425 and light absorption fraction in cells (ExtL), minimum saturation for activation (Sat\_act0), and minimum saturation for full  
426 activation (Sat\_act1) – on the annual C balance of lichen-dominated biocrusts in different climate regions. The parameters  
427 decreased or increased by a consistent range for all sites based on the measured deviation in photosynthesis response curves  
428 of replicates. The altered annual C balance resulting from increasing or decreasing parameters is normalized by the original C  
429 balance. The coloured columns indicate the average value of the normalized C balance at sites with similar climate conditions.  
430 Various styles of black points indicate different sites. (b) Relative importance of each physiological parameter compared to  
431 other parameters across the climatic regions. Larger relative importance implies a more important effect the parameter has on  
432 the C balance compared to other parameters in the given climatic region, and vice versa.

433 However, even though physiology parameters play an important role in all regions, the C balance at T1, T2 and A1 did not  
434 become positive when the physiological parameters were varied reasonably, that is the parameters were varied to relatively  
435 cover the deviation of response curves of replicates. Furthermore, the change in C balance value is much smaller in drylands  
436 compared to other regions (as shown in Fig. S11 in the Supplement).

### 437 **3.3.3 Acclimation of physiological properties**

438 The sensitivity analysis for acclimation showed a marked increase in annual productivity of lichen- and moss-dominated  
439 biocrusts at site T1 (Fig. 8) when the seasonal acclimation of several physiological parameters was included in the model (from  
440  $-42.8$  to  $-15.5$  g C m<sup>-2</sup> yr<sup>-1</sup> and from  $-39.4$  to  $-4.2$  g C m<sup>-2</sup> yr<sup>-1</sup>).



441

442 **Figure 8:** Comparison of accumulated annual C balance between two simulations in the sensitivity analysis of seasonal  
 443 acclimation of physiological properties. In the simulation “fixed parameters”, all parameters that have been calibrated or  
 444 measured remained constant throughout the simulation year. For the simulation “dynamic parameters” at site T1,, parameter  
 445 metabolic respiration cost per surface area (Resp\_main) was reduced by half, light absorption fraction in cells (ExtL) was  
 446 doubled but restricted to one, the ratio of Jmax to Vcmax (jvratio) was increased by two times from September to February to  
 447 adapt to the winter climates. For the other months, the “fixed” values were used.

#### 448 4 Discussion

##### 449 4.1 Simulated C balance of data-driven model

450 The data-driven model aims to provide observation-based estimates of the carbon fluxes of non-vascular photoautotrophs  
 451 which may serve as approximation for the C balance of vegetation in biocrust-dominated ecosystems. At the two dryland sites,  
 452 the moss-dominated biocrusts were estimated to be carbon sinks on an annual basis, and lichen-dominated biocrust can also  
 453 be a carbon sink at one of the dryland sites. As shown in the results, mosses accumulated 1.9 and 3.1 g C m<sup>-2</sup> yr<sup>-1</sup> at site D1  
 454 and D2, respectively, and lichens accumulated 1.5 g C m<sup>-2</sup> yr<sup>-1</sup> at site D1.

455 The estimated C balance at the two dryland sites is consistent with the magnitude of the annual C balance of different biocrust  
 456 types reported by various studies in arid habitats. Feng et al. (2014) recorded that the biocrusts composed of lichens, mosses,

457 and cyanobacteria of the Mu Us Desert in China took up 3.46 to 6.05 g C m<sup>-2</sup> yr<sup>-1</sup>. Brostoff et al. (2005) estimated a larger  
458 carbon gain by lichen biocrust of 11.7 g C m<sup>-2</sup> yr<sup>-1</sup> in the Mojave Desert, USA. For cyanobacteria, an annual carbon uptake of  
459 0.02 to 2.3 g C m<sup>-2</sup> was reported for deserts (Jeffries et al., 1993). The estimated C balance values in drylands fluctuate  
460 relatively largely, but the magnitude is consistent with the simulated results by the data-driven model at D1 and D2.

461 For biocrust lichens growing on rock surface without soil attached underneath in a temperate grassland, Lange (2003b)  
462 measured an annual carbon gain of 21.49 g C m<sup>-2</sup>. Additionally, several studies estimated the carbon budget in humid tundra  
463 habitats. An amount of ~12–70 g C m<sup>-2</sup> yr<sup>-1</sup> carbon was fixed by moss-dominated biocrust, for instance (Schuur et al., 2007).  
464 The magnitude of these values corresponds to the estimation of the C balance at T3. However, the estimated annual carbon  
465 losses of lichens and mosses by the data-driven model in temperate regions T1 and T2 should actually lead to the death of  
466 these organisms, which is not consistent with their dominant abundance in the field and is much lower than published by  
467 previous studies.

468 The mismatches of trait values between strategies predicted via selection by the LiBry model for the sites and the collected  
469 species with regard to their net photosynthesis response curves indicate that the physiological parameter values that would be  
470 necessary to maintain a positive C balance in LiBry are not compatible with those of the sampled biocrusts. This is in line with  
471 the results of the data-driven model, which also simulates a negative C balance and is based on the same physiological  
472 processes as LiBry. This also applies to the lack of seasonal acclimation in both modeling approaches, since the strategies in  
473 LiBry are assumed to have constant functional properties throughout the simulation.

#### 474 **4.2 Uncertainties of long-term C balance simulated by the data-driven model.**

475 The data-driven model simulated relatively reasonable C balance values in drylands but unrealistic, negative values at  
476 temperate sites T1 and T2. Since the same or similar gas exchange methodology has been used for all sites, differences in the  
477 simulated C balance among these regions likely result from variation in the species-specific interactions between climate and  
478 physiological processes, including seasonal variation in physiological properties due to acclimation.

479 As the results (Fig. 6) show, CO<sub>2</sub> concentration is an essential factor for the annual C balance of biocrusts, especially at dryland  
480 and some temperate sites. Therefore, uncertainty in the CO<sub>2</sub>-value prescribed in the model may be a source of error. The CO<sub>2</sub>  
481 concentration at the surface boundary might exceed the value of 400 ppm that was prescribed in the model because of CO<sub>2</sub>  
482 diffusion from the soil, which may lead to an underestimated C balance (Fig. 6a). However, with enhanced CO<sub>2</sub> concentration  
483 in the sensitivity analysis (600 ppm) at site T1, for instance, the estimated C balance increased only slightly, and is still strongly  
484 negative (-37.0 g C m<sup>-2</sup> yr<sup>-1</sup> for lichen and -30.2 g C m<sup>-2</sup> yr<sup>-1</sup> for moss). Hence, the lower CO<sub>2</sub> concentration can partially  
485 contribute to the strongly negative C balance at T1 and T2, but is not a major factor.

486 Furthermore, the negative C balance at temperate and alpine sites may result from the uncertainties in physiology, which were  
487 also observed between replicates (see Fig. 2 and Fig. S2). An overestimation of dark respiration rates of the photoautotrophs  
488 in the biocrust may result from including a small amount of heterotrophic respiration. The overestimated respiration rate then  
489 leads to an overestimation of the parameter metabolic respiration cost per surface area (Resp\_main) and might also cause an

490 underestimated Q10 value ( $q_{10}$ ) calculated from the respiration rates. The uncertainties of these two parameters reduce the  
491 estimated C balance largely (Fig. 7). Additionally, the optimum temperature ( $T_{opt}$ ), which is also the reference temperature  
492 for calculating the respiration rate, cannot be well constrained by the limited measured temperature response data set. Thus,  
493  $T_{opt}$  may be underestimated. The larger difference in surface temperature to  $T_{opt}$  results in a larger respiration rate, and lower  
494 gross photosynthesis, which leads to a lower C balance.

495 Although the uncertainty in individual physiological parameters may not lead to the markedly negative C balance estimates,  
496 as indicated by still negative values upon variation of these parameters (Fig. S11), additive effects of all parameters combined  
497 with long-term unfavourable environmental conditions may cause a large amount of carbon lost over a year. The optimal  
498 conditions are rare within a year, which was also described by Lange (2003b). Thus, the overestimated respiration rate leads  
499 to a lower carbon gain during the relatively optimal conditions, which may not be sufficient to compensate for exaggerated C  
500 losses under long-term harsh conditions, such as autumn and winter at site T1, for instance. For this reason, the simulated C  
501 balance of mosses and lichens in temperate humid regions was mostly negative.

502 In addition to the uncertainty in the values of physiological parameters, seasonal acclimation of these physiological traits to  
503 the current climatic conditions may play an important role in regulating the C balance at humid sites where the organisms are  
504 active throughout the year, such as site T1 (Fig. 8). It was observed, for instance, that the respiration of lichens shows  
505 acclimation to seasonal changes in temperature, and the maximum  $CO_2$  exchange rate of the organisms remains steady  
506 throughout the year (Lange and Green, 2005). Gauslaa (2006) found a higher chlorophyll a/b ratio in forest lichens with  
507 increasing light. Moreover, depression in quantum efficiency in summer under extremely dry conditions has been observed  
508 (Vivas et al., 2017). These varied physiological properties of organisms within a year may result in different photosynthesis  
509 and respiration rates, and thus different C balances in comparison to the ones that cannot acclimate to the seasonal climate.  
510 The missing seasonal acclimation of physiological traits may explain why the data-driven model estimated a negative C  
511 balance for biocrusts in humid regions.

512 Another limitation of the modeling approach may be the lack of separate responses of respiration and photosynthesis to  
513 metabolic activity. Both photosynthetic activity and respiration reach their maximum in the model once the water saturation  
514 reaches the optimum value for net photosynthesis ( $W_{opt\_np}$ ). In some cases, however, respiration rate may reach the maximum  
515 value only at a higher saturation than  $W_{opt\_np}$  (Lange, 1980), indicating that the model may overestimate respiration in the  
516 long-term.

517 In comparison to the unrealistic C balance numbers at T1 and T2, we estimated more reasonable values in drylands and at T3.  
518 However, we do not make a definitive statement about whether or not the model predicts an accurate C balance in drylands.  
519 since the measured climate data and photosynthesis response curves that were used for calibrating land surface properties and  
520 various physiological parameters represent only samples of the large physiological and climatic variation. A higher accuracy  
521 would be more likely to be expected in drylands as these regions have a more uniform climate throughout the year than  
522 temperate regions that show substantial seasonality. Additionally, variation in light conditions is slightly more relevant for the  
523 simulated C balance than variation in moisture (see Fig. 6) because the organisms are able to become inactive, meaning that



524 the dry season in drylands does not have a decisive effect on the C balance, while low light in winter in temperate climate does  
525 since organisms have to be active then. Furthermore, the longer total inactive period in drylands could reduce the bias in the  
526 magnitude of the simulated C balance caused by incorrectly estimated physiological parameter values. We estimated a smaller  
527 absolute change in annual C balance in drylands with varied physiological parameters in the sensitivity analysis (for instance,  
528 the C balance of lichens changed by 34.6 g C m<sup>-2</sup> yr<sup>-1</sup> for parameter T<sub>opt</sub> at T1, while it changed only by 1.5 g C m<sup>-2</sup> yr<sup>-1</sup> at  
529 D1).

530 Furthermore, the estimated C balance may be inaccurate due to the potential bias in estimated relative water saturation, which  
531 partly depends on prescribed MWC obtained by measurements. However, the outcome of the sensitivity analysis of MWC at  
532 T1 revealed that the annual carbon estimation is robust to the uncertainties with regard to the prescribed MWC (details in the  
533 Supplement). Another factor that potentially affects the accuracy of C balance estimates is interannual climatic variability.  
534 While the model estimated unrealistic C balance values of lichen-dominated biocrusts at T1 and T2 for current conditions, the  
535 C balance may have been different in other years. Therefore, the simulation of annual C balance based on multi-year climate  
536 data is worthy of future study to understand the long-term C balance better. Moreover, the estimated negative C balance of  
537 certain lichen and moss species may not be generalizable and representative for the overall situation in the field due to the  
538 large variation in physiological adaptation strategies to climate. There could be other organisms that form cryptogamic covers,  
539 for instance, that show a different degree of depression in net photosynthesis at high water content (Lange et al., 1995), and  
540 thus have more reasonable C balance values.

#### 541 **4.3 Potential factors influencing the C balance**

542 Despite diverse climatic conditions, we found similarities regarding the dominant environmental factors and physiological  
543 parameters controlling the C balance. Thereby, CO<sub>2</sub> and air temperature were the two most important environmental factors at  
544 all sites. Relative air humidity, partly precipitation, and light intensity were also relevant for the estimation of the C balance.  
545 In terms of physiological parameters, the respiration-related parameters were the most important drivers, while parameters that  
546 affect V<sub>cmax</sub> and thus the light-independent CO<sub>2</sub> assimilation rate were relevant, too.

547 The relative importance of these factors/parameters varied slightly among climatic regions. Regarding the comparison between  
548 environmental factors, we cannot rule out that the magnitudes of changes in environmental factors that we applied in the  
549 sensitivity analysis were not balanced, which may have led to an overestimation of the relative importance of certain factors,  
550 such as air temperature, for instance, compared to the others. The spatial patterns across climate regions of a given  
551 environmental factor, however, are not affected by this, which means that differences between climatic regions for a given  
552 factor are most likely robust. Hence, air temperature is more relevant at the alpine site and relative air humidity has a higher  
553 impact in temperate than in other regions, rainfall and CO<sub>2</sub> are likely to have the largest effect on C balance in drylands. Even  
554 though the data-driven model failed to estimate reasonable C balance at some sites, the comparison of the relative importance  
555 of the environmental factors across climatic regions may be valid since the measurement procedure is consistent. Moreover,  
556 the patterns of relative importance remain similar when excluding the sites with strongly negative C balance (T1, T2, and A1;

557 as shown in the Fig. S12 in the Supplement). Nevertheless, we only studied the sensitivity of the C balance of biocrusts  
558 dominated by the lichen *Psora decipiens* and *Cladonia furcata* (at T3), and there are variations between lichens of different  
559 growth forms and between biocrust types. For example, cyanolichens increase in abundance with increasing rainfall, but  
560 trebouxoid lichens have their physiological optimum in drier conditions (Phinney et al., 2021). Moreover, the impact of  
561 precipitation on isidiate lichens is weaker than that of temperature (Phinney et al., 2021).

#### 562 **4.3.1 Environmental factors**

563 Our results suggest that warming can result in a large amount of carbon loss at all sites, with a particular large effect in the  
564 alpine region. The consistent effects of warming on C balance of biocrusts are found in various field studies (e.g., Darrouzet-  
565 Nardi et al., 2015; Ladrón de Guevara et al., 2014; Li et al., 2021; Maestre et al., 2013). This can be explained by the overall  
566 less optimal water and temperature conditions associated with warming. The simulated increasing respiratory costs with  
567 warming overcompensate gains in gross photosynthesis.

568 Ambient CO<sub>2</sub> concentration affects the gross photosynthesis rate to a large extent in the model. Although the intra-annual  
569 change in air CO<sub>2</sub> concentration may be small in the field compared to other environmental factors, the increase of CO<sub>2</sub> in the  
570 atmosphere in recent decades (IPCC 2021) may alter the long-term C balance substantially. However, this beneficial effect of  
571 elevated CO<sub>2</sub> on photosynthesis and C balance may be reduced in reality due to future limitation of growth by nitrogen (Coe  
572 et al., 2012a), which is not considered in the model, or also due to shortened activity periods resulting from warmer and drier  
573 future climatic conditions.

574 Light intensity has the third largest effect on C balance, slightly larger than moisture. Light is one of the essential factors for  
575 photosynthesis as simulated by our model, and it is a limiting factor of photosynthetic carbon assimilation, in particular in  
576 winter at temperate and alpine sites (the mean value of radiation maxima in January is 244  $\mu\text{mol m}^{-2} \text{s}^{-1}$  at T1 and 245  $\mu\text{mol}$   
577  $\text{m}^{-2} \text{s}^{-1}$  at the alpine site). Hence, increasing light intensity can promote carbon accumulation.

578 Factors that determine water supply are rainfall and non-rainfall inputs such as dew and water vapor that are related to relative  
579 humidity. The relative importance of different moisture factors in mediating C balance varies in the model. Relative humidity  
580 plays a more important role in mediating the C balance than rainfall amount. This may be due to the timing of dew or water  
581 vapor uptake, which is greatest before sunrise (Chamizo et al., 2021; Ouyang et al., 2017) and prolongs the activated periods  
582 in the early morning when the organisms start assimilating carbon (Veste and Littmann, 2006). This may result in a markedly  
583 increased annual C balance in the model. Rainfall amount was not a key factor affecting the simulated biocrust performance  
584 at one of the arid sites, which is consistent with another study (Baldauf et al., 2021). At the other dryland site (D2), however,  
585 this was not the case. Moreover, we found that the effect of the amount of rainfall is small in humid temperate and alpine  
586 regions as well. The differing effects of rainfall on the C balance depend on the change in relative water saturation that follows  
587 from rainfall event sizes and patterns throughout the year (Reed et al., 2012). In some cases, decreased rainfall leading to lower  
588 water saturation of biocrusts may facilitate photosynthetic carbon gain via increasing the CO<sub>2</sub> diffusivity from the atmosphere  
589 into the chloroplast (Lange et al., 1997). Nevertheless, reducing water saturation below a certain value can cause a decline in

590 the duration of activity (Proctor, 2001; Veste et al., 2008) which thus reduces carbon accumulation. Thus, there may be a rain  
591 threshold below which decreasing rain may start having a negative effect on biocrust C balances. The threshold is likely  
592 species-specific as it is associated with the water-holding capacity of the organism. Our simulation results thus highlight the  
593 need for the combined application of field experiments and data-driven modeling to improve our understanding of differential  
594 responses to variation in precipitation.

#### 595 **4.3.2 Physiological parameters**

596 The parameter  $q_{10}$  is a key parameter that substantially affects respiration.  $Resp_{main}$  is the dark respiration rate at a reference  
597 temperature that is linked in the model to  $V_{cmax}$ , the maximum rate of carboxylation of RuBisCO in the Calvin Cycle of  
598 photosynthesis (Walker et al., 2014).  $Topt$  is a parameter that controls gross photosynthesis as well as respiration as it is also  
599 the reference temperature for calculating respiration rates.  $Rub_{ratio}$  can affect  $V_{cmax}$  and hence the maximum  $CO_2$   
600 assimilation rate, while  $ExtL$  regulates the light using efficiency under limited light conditions.  $Sat_{act0}$  and  $Sat_{act1}$  are two  
601 parameters that determine the range of water saturation for initial activation and full metabolic activity. They have the smallest  
602 effects on the C balance of lichen-dominated biocrusts at all sites.

603 Our modeling results give insights into the relative effects of individual physiological parameters on annual C balance across  
604 different climatic zones. However, the impacts of physiology on biocrust C balance are complex since they always arise from  
605 combinations of these physiological parameters. Thereby, different parameter combinations that correspond to different  
606 relative impacts on the C balance may lead to the same response curves. Hence, we cannot directly link individual physiological  
607 parameters to the underlying mechanisms since we do not have enough data to distinguish multiple possible parameter  
608 combinations from each other, in case they produce the same response curves.

#### 609 **4.4 Validation of the data-driven model**

610 The validation results of the model showed an overall good fit of daily and diurnal patterns of water content and activity (Fig.  
611 3 and S3), and C balance at D2 (shown in Fig. 4) given the uncertainties in the data used to parameterize and evaluate the  
612 model. This indicates that the data-driven model may be a reliable tool for C balance estimation, provided that a sufficient  
613 amount of suitable forcing data is available.

614 A potential explanation for the general underestimation of activity at night and morning during several periods in a year is the  
615 larger prescribed MWC and  $Sat_{act0}$  of organisms in the model compared to those of the samples from the activity  
616 measurements. Consequently, simulated saturation was lower, but minimal saturation for being active was higher than the  
617 samples. Thus, the activity may have been underestimated at small water inputs such as dew and water vapor, which occur  
618 mainly during the night and in the morning hours (Fig. S13). Moreover, underestimated activity in April and June at site T1  
619 (Fig. S3 F(b)) may have resulted from a gap in rainfall measurements during this period. Not only rainfall amount but also  
620 timing and frequency of rainfall events are essential for the physiological responses of biocrust communities (Belnap et al.,  
621 2004; Coe et al., 2012b; Reed et al., 2012). Therefore, although the measured annual total amount of rainfall is reasonable (424

622 mm at site T1), the missing rainfall during a series of days in summer at site T1 would lead us to incorrectly predict that the  
623 biocrusts remain inactive on these days.

624 Moreover, the mismatch between modelled and observed CO<sub>2</sub> assimilation rates at low or high water contents at site D2 (Fig.  
625 4) may have partly resulted from the calibration procedure. In the calibration the simulated CO<sub>2</sub> exchange rates were higher  
626 than measurements when the saturation exceeded the optimum saturation and hardly showed any negative values at high  
627 saturation (Fig. S2 f). In turn, the simulated CO<sub>2</sub> exchange rates of biocrusts with an extremely low water content were zero  
628 while the measurements showed negative values (see Fig. S2 f), pointing at a certain degree of metabolic activity in natural  
629 conditions. Furthermore, the samples used for validation were different from the ones for calibration, which can also lead to  
630 inaccuracies.

631 Additionally, the ability of the model to capture seasonality variations of C balance, which have been shown by other studies  
632 (Büdel et al., 2018; Lange, 2003a; Zhao et al., 2016), could not be evaluated here since the monitoring of C balance in the field  
633 and collection of samples used for photosynthesis performance measurements were conducted only during October and early  
634 November.

635 **5 Conclusions**

636 Our data-driven model provides possibilities to predict the long-term C balance of biocrusts in the field across various climate  
637 zones, and it enables us to analyse mechanisms that drive the C balance, despite marked uncertainties in the parametrization.  
638 We simulated reasonable C balance values in drylands but unrealistic ones at temperate sites with substantial seasonality.  
639 Uncertainties in environmental factors and respiration rate are likely to be the source of error for the C balance estimation  
640 since (1) all environmental factors that were examined in our study may act as relevant drivers for the C balance of biocrusts  
641 and (2) respiration-related parameters had the largest impacts compared to other physiological parameters, such as water  
642 relations or parameters solely related to V<sub>c</sub>max. CO<sub>2</sub> and air temperature showed the strongest effects among environmental  
643 factors and at the alpine site, the air temperature was most relevant. Compared to environmental factors, the relative impacts  
644 of physiological parameters are rather equal across climate regions. The optimum temperature may be slightly more relevant  
645 in temperate regions, while metabolic respiration cost is most important at the alpine site. Due to the importance of respiration-  
646 related physiological parameters, more studies to improve their accuracy are warranted in the future application of C balance  
647 modeling approaches.

648 Our study suggests that a better, more detailed understanding of the seasonal variation of physiological traits is necessary, as  
649 the more realistic estimations in drylands compared to temperate sites could be due to the weaker climate seasonality. The  
650 model needs to be calibrated with a larger number of samples collected and measured in various seasons to take the acclimation  
651 of physiological properties into account. Additionally, the integration of acclimation of physiological traits in process-based  
652 models may improve their accuracy in C balance estimation.

653 *Code and Data Availability Statement.* Source code of the data-driven model, LiBry modelling results, R-scripts to run the  
654 analysis in this manuscript are available in Zenodo repository at <https://doi.org/10.5281/zenodo.7756960> (Ma, 2023). Field  
655 and laboratory data are available in the corresponding publications cited in the manuscript and also from the corresponding  
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657 *Author Contributions.* YM and PP designed the study, BW, AK, MV and JR provided the observational data. YM did the  
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659 and all authors revised it.

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