



## Methane emissions from Arctic landscapes during 2000-2015: An analysis with land and lake biogeochemistry models

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**Abstract.** Wetlands and freshwater bodies (mainly lakes) are the largest natural source of greenhouse  
10 gas CH<sub>4</sub> to the atmosphere. Great efforts have been made to quantify these source emissions and their  
uncertainties. Previous research suggests that there might be significant uncertainties coming from  
“double accounting” emissions from freshwater bodies and wetlands. Here we quantify the methane  
emissions from both land and freshwater bodies in the pan-Arctic with two process-based  
15 biogeochemistry models by minimizing the double accounting at the landscape scale. Two non-  
overlapping dynamic areal change datasets are used to drive the models. We estimate that the total  
methane emissions from pan-Arctic are 35.81 Tg CH<sub>4</sub> yr<sup>-1</sup> during 2000-2015, of which wetlands and  
freshwater bodies are 21.38 CH<sub>4</sub> Tg yr<sup>-1</sup> and 14.45 Tg CH<sub>4</sub> yr<sup>-1</sup>, respectively. The emissions are  
significantly affected by humidity and vapor pressure, followed by temperature and landscape areal  
20 changes. Emissions from wetlands are more sensitive to landscape areal changes than from freshwater  
bodies while the latter emissions are more influenced by temperature than precipitation. Our sensitivity  
analysis indicates the pan-Arctic CH<sub>4</sub> emissions from freshwater bodies were highly influenced by  
temperature, but less by lake sediment carbon content from permafrost thaw.

### 1. Introduction

25 Atmospheric methane (CH<sub>4</sub>) is one of the major greenhouse gasses which contributes to about 20%  
of the warming effect, second only to carbon dioxide (CO<sub>2</sub>). Atmospheric methane concentrations have  
risen 2.5 times since the beginning of the industrial age (Hamdan and Wickland, 2016). However, its  
global warming potential is 25 times higher than CO<sub>2</sub> (IPCC 2014). Previous studies have suggested that  
inland water systems (wetlands and freshwater bodies) are the single largest natural source of the



greenhouse gas CH<sub>4</sub> (Saunois et al., 2020), in which the global wetland contributes 60–80% of natural  
30 CH<sub>4</sub> emissions (Quiquet et al., 2015; Hopcroft et al., 2017) and lakes account for approximately 30% of  
biogenic methane emissions (Guo et al., 2020), both have been found to increase under changing climate.

Wetland CH<sub>4</sub> emissions are the largest natural source in the global CH<sub>4</sub> budget, contributing to  
roughly one-third of total natural and anthropogenic emissions. Under the RCP 2.6 scenario, climate  
change-induced increases in boreal wetland extent and temperature-driven increases in tropical CH<sub>4</sub>  
35 emissions will dominate anthropogenic CH<sub>4</sub> emissions by 38 to 56% toward the end of the 21st century  
(Zhang et al., 2017). Likewise, lakes are the second largest CH<sub>4</sub> source of all inland water emissions after  
wetlands (Kyzivat, et al., 2022). They are especially common in high latitudes and account for about 10%  
of the boreal landscape (Guo et al., 2020). This high coverage of lakes especially the extensive shallow  
seasonally ice-covered ones in the subarctic landscapes has been considered as a major source of  
40 atmospheric methane in northern high latitudes (Bastviken et al., 2011). Unlike wetlands, shallow lakes  
have the highest methane emission potential in the cold season which dominate the spring methane  
release in the pan-Arctic area (Jammet et al., 2015). Because of the considerable total lake area and the  
substantial shallow lakes in the area of 40–70° N, this region was also found to be the dominant  
contributor (~30%) of global lake diffusive CH<sub>4</sub> emissions (Li et al., 2020). However, in comparison  
45 with land methane emission studies, less work has been done on studying lake CH<sub>4</sub> emissions through  
process-based modeling (Saunois et al., 2020), especially for the pan-Arctic region.

To date, although great efforts have been made to quantifying the uncertainties of global wetland  
and lake methane emissions separately (Liu et al., 2020; Guo et al., 2021), there are still significant  
differences between the estimates of the Arctic CH<sub>4</sub> natural sources using ‘bottom-up’ method which  
50 aggregated lakes, wetlands and coastal waters as CH<sub>4</sub> sources (32–112 Tg CH<sub>4</sub> yr<sup>-1</sup>; McGuire et al., 2009;  
Saunois et al., 2020) and ‘top-down’ method which determines the emissions based on the spatial and  
temporal variability of atmospheric CH<sub>4</sub> concentration measurements (15–50 Tg CH<sub>4</sub> yr<sup>-1</sup>; AMAP, 2015).  
In those studies, there are potential “double accounting” issues for certain areas of wetlands and lakes  
using low-resolution wetland and lake distribution data (Thornton et al., 2016). Specifically, some small  
55 lakes and ponds might have been considered as lakes using lake models while wetland modeling might  
have also treated those as wetlands, therefore being accounted for twice in the regional methane emission  
estimation.



Here we use two process-based biogeochemical models, the Terrestrial Ecosystem Model (TEM-MDM, Liu et al, 2020) and the Arctic Lake Biogeochemistry Model (ALBM, Guo et al., 2020), along with two dynamic area datasets for both wetlands (WAD2M, Version 2.0, unpublished data) and lakes (GLCP; Meyer et al., 2020) ecosystems, to quantify the methane emissions considering the impact of the landscape changes in both land ecosystems and freshwater bodies in the pan-Arctic region for the period 2000-2015.

## 2. Method

### 2.1. Model description

The Terrestrial Ecosystem Model (TEM) is a process-based biogeochemistry model which considers carbon, nitrogen, water, and heat processes in terrestrial ecosystems and was originally used to simulate ecosystem carbon and nitrogen dynamics (Melillo et al., 1993; Zhuang et al., 2001, 2002, 2003, 2004, 2007, 2013). The model considers important freeze-thaw processes and explicitly integrates soil thermodynamics in permafrost and non-permafrost region biogeochemical processes. It is also coupled with a complex hydrological module that enables the modeling of soil moisture profiles and water table depths in upland and wetland ecosystems. Zhuang et al. (2004) also developed a Methane Dynamics Module (MDM), which was integrated into TEM to estimate CH<sub>4</sub> emissions from northern high-latitude regions and further revised and extrapolated to the global scale to quantify soil methane consumption (Zhuang et al., 2013). Recently, Liu et al. (2020) revised the model to the version we used in this study by taking into account several more detailed land methane cycling processes, including various types of wetlands in different regions based on plant functional types, the impact of above-soil surface water on methane transport, and cumulative vertical methane concentrations in soil, such that it can give a more precise methane estimate on the global scale.

The Arctic Lake Biogeochemistry Model (ALBM) is a 1-D process-based climate-sensitive lake biogeochemical model originally developed for simulating CH<sub>4</sub> production, oxidation, and emission in Arctic lakes (Tan et al., 2015; Tan and Zhuang 2015a, 2015b) and later revised to predict both thermal and carbon dynamics of aquatic ecosystems in boreal lakes (Tan et al., 2017; Guo et al., 2020), and it was then successfully applied to temperate lakes (Tan et al., 2018; Guseva et al., 2020). Recently, the ALBM is also shown to be capable of simulating global lake thermal dynamics (Guo et al., 2021). The



model consists of several modules, including those for the water/sediment thermal circulation, conceptualized as the water thermal module (WTM) and the sediment thermal module (STM), and those for the gas diffusive and ebullition transportation, conceptualized as the bubble transport module (BTM) and the dissolved gas transport module (GTM) (Tan et al., 2015). The model also covers the radiative  
90 transfer processes and the water/sediment biogeochemistry, including the terrestrial ecosystems' organic carbon loading, the microbial and photochemical organic carbon degradation, the photosynthesis for inorganic carbon fixation, and phytoplankton biomass loss through respiration for further simulation of CO<sub>2</sub> dynamics. The ability of ALBM to simulate and represent the thawing and freezing cycles of sediments in thermokarst lakes and the organic carbon inputs induced by thermokarst activities, the  
95 degradation of dissolved organic carbon through photochemical mineralization, and the mobilization and mineralization of labile organic carbon in the deep sediments of yedoma lakes is crucial for understanding the carbon dynamics in Arctic lakes which makes it a better choice for simulating Arctic lake methane emission than other lake models that are usually lacking these processes (Tan et al., 2017).

## 2.2. Input Data

100 Here we use two global dynamic area changing datasets for both wetland and lake ecosystems. For wetlands, the Wetland Area and Dynamics for Methane Modeling (WAD2M) Version 2.0 was used as the TEM-MDM model input as transient wetland inundation fraction data. The dataset following the same processing method as Version 1.0 (Zhang et al., 2021), which was used for quantifying the global methane budget for 2000-2017 (Saunois et al., 2020), but included a few updates on the static inventories applied  
105 in WAD2M and used the same monthly SWAMPS version 3.2 (Jensen and McDonald, 2019), was provided for the Global Carbon Project wetland CH<sub>4</sub> (GCP-CH<sub>4</sub>) model intercomparison. Compared to the previous one, the new version applied Global River Width from Landsat (GRWL) Database (<https://zenodo.org/record/1297434>; Allen and Pavelsky, 2018) and HydroLAKES ([https://www.hydrosheds.org/images/inpages/HydroLAKES\\_TechDoc\\_v10.pdf](https://www.hydrosheds.org/images/inpages/HydroLAKES_TechDoc_v10.pdf); Messenger et al., 2016)  
110 instead of the Joint Research Center Global Surface Water (GSW) dataset (Pekel et al., 2016) to remove inland freshwater systems, defined as lakes, ponds, and rivers and the time period was extended to 2000-2020.

For lake simulation, we used the Global Lake area, Climate, and Population dataset (GLCP; Meyer et al., 2020) as the dynamic input for ALBM model. Using the HydroLAKES database version 1.0 for



115 the locations and numbers of lakes, the GLCP contains over 1.4 million lakes of at least 10 ha in surface  
area, with annual surface area (identified as permanent or seasonal water) from 1995 to 2015, paired with  
annual basin-level temperature, precipitation, and population values. Since it uses the same database that  
WAD2M uses to remove inland freshwater bodies (HydroLAKES version 1.0), which ensures that these  
two datasets won't overlap with each other and hence avoiding "double accounting" induced  
120 uncertainties in simulating methane emissions at the landscape scale. We further classified the lakes into  
four types based on their location and permafrost thawing type in the pan-Arctic area (above 45° north),  
including yedoma thermokarst lakes (yedoma/YDM), non-yedoma thermokarst lakes  
(thermokarst/TMK), non-thermo boreal lakes (boreal/BRL), and temperate lakes (temperate/TMP). From  
which, yedoma and thermokarst lakes are classified based on circum-polar Yedoma map (Jens et al., 2022)  
125 and Arctic Circumpolar Distribution and Soil Carbon of Thermokarst Landscapes (Olefeldt et al., 2016),  
non-thermo boreal lakes and temperate lakes were defined on whether their location is above 60° north.  
At the end, there are total 1,248,478 lakes were simulated, including 101,852 yedoma lakes, 249,434  
non-yedoma-thermokarst lakes, 390,687 non-thermo-boreal lakes, and 506,505 temperate lakes. Because  
the time period is different for these two datasets (2000-2020 for WAD2M and 1995-2015 for GLCP),  
130 we chose the overlap years 2000-2015 as our simulation time period.

For the climate forcing data, we used GSWP3-W5E5 dataset (gswp3-w5e5\_obsclim\_global\_daily,  
<https://data.isimip.org/10.48364/ISIMIP.982724>; Lange et al., 2022), a factual climate input daily dataset  
with a resolution of  $0.5^\circ \times 0.5^\circ$  globally provided by the Inter-Sectoral Impact Model Intercomparison  
Project (ISIMIP). These forcing data were used for both models to ensure that no additional uncertainties  
135 are introduced. Air temperature, surface pressure, 10 m wind speed, relative humidity, precipitation,  
snowfall, downward short-wave radiation and downward long-wave radiation were used in ALBM model  
as input forcing. For TEM-MDM model simulation, we only used air temperature, relative humidity,  
precipitation, and downward short-wave radiation, where air temperature and relative humidity were  
used to calculate the vapor pressure as another input.

### 140 2.3. Model parameters

The model parameters are derived from previous studies, both of which did the parameter calibration  
and validation on a global scale (Liu et al., 2020; Guo, et al., 2021). For TEM-MDM, 15 key parameters  
involved in wetland methane oxidation and production processes were calibrated and validated at the site



level (15 sites for calibration and 14 sites for validation) using the Shuffled Complex Evolution Approach  
145 (SCE-UA). Other information, such as vegetation type, soil texture, and wetland type, were also set based  
on site observations. For ALBM, 58 freshwater lakes of varying shapes, locations, climates, and  
landscapes were used for the calibration of nine lake sediment property related parameters. The  
calibration process used the Sobol sequence sampling method to generate a perturbed parameter  
ensemble (PPE) of 10,000 samples from the parameters space and then the Monte Carlo method was  
150 applied to simulate this PPE for each lake. Six years of the observation data from each lake were used  
for calibration and the rest were used for validation.

#### 2.4. Simulation protocols

Model simulations followed different protocols for different models. In wetland simulation (TEM-  
MDM), the Terrestrial Ecosystem Model 5.0 (TEM5) was first run in the same simulation area and time  
155 period to get the net primary production (NPP) and leaf area index (LAI), the outputs were then fed to  
TEM-MDM as input to calculate methane emissions. For TEM5 simulation, we first did the spin-up run  
10 times with 40 years per spin before the transient simulation to let the model reach a steady state using  
the first 40-year input data, 120 years transient simulation was run in TEM-MDM while the first 100  
years simulation was used as spin up. For lake simulation using ALBM, as discussed in section 2.2, the  
160 lakes were classified into four types based on their location and permafrost thawing type. We further  
grouped each type of lakes based on their surface area ( $<1 \text{ km}^2$ ,  $1\text{-}10 \text{ km}^2$ ,  $>10 \text{ km}^2$ ) and depth ( $<3 \text{ m}$ ,  $>3 \text{ m}$ )  
and whether they are in the same  $0.5^\circ \times 0.5^\circ$  pixel so that lakes in the same groups will be driven  
by the same meteorology input data. Different types of lakes used different parameter sets derived from  
calibration. For all the simulations, a spin-up period of 10 years was run first.

#### 165 2.5. Sensitivity test

Sensitivity tests were conducted towards lake emissions simulation in three aspects. According to  
the previous studies, by the end of the 21<sup>st</sup> century, the temperature will increase roughly by  $5^\circ \text{C}$  and the  
precipitation exhibits a trend at a rate of  $+1.5 \text{ mm/yr}$ , corresponding to  $\sim 15\%$  increase in the RCP 8.5  
scenario (IPCC 2014; Chen et al., 2014). Therefore, we rerun the simulation by 1) increasing the  
170 temperature by  $5^\circ \text{C}$ ; 2) increasing the precipitation by 15%, where both rain and snowfall were  
considered; and 3) adding additional 15% carbon into lake sediments to simulate the influence of



permafrost thawing due to global warming. For temperature and precipitation, we directly modified them at the data input step. For lake sediment carbon, we assumed that the additional carbon transferred straightly from old organic matter in thawing permafrost (old  $^{14}\text{C}$ -depleted organic carbon pool) to new  
175 organic matter at the water-sediment interface (young  $^{14}\text{C}$ -enriched organic carbon pool) and changed it by altering the labile carbon density ( $C_{\text{labile}}$ ) (Tan et al., 2015). Because the old  $^{14}\text{C}$ -depleted organic carbon pool only contributes to  $\text{CH}_4$  production in the permafrost thaw bulb under yedoma and thermokarst lakes, we just altered the corresponding  $C_{\text{labile}}$ .

### 3. Results

#### 180 3.1. Temporal dynamics of methane emissions at the landscape scale

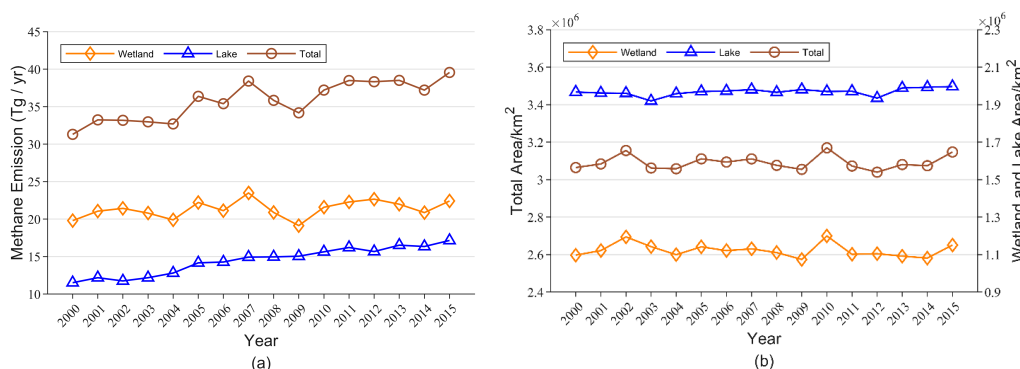
The ALBM model simulation driven with the GLCP dataset indicates that the methane emission from lakes in the pan-Arctic region ranges from  $11.51 \text{ Tg CH}_4 \text{ yr}^{-1}$  in the year 2000 to  $17.15 \text{ Tg CH}_4 \text{ yr}^{-1}$  in the year 2015 with a mean value of 14.45 and standard deviation of 1.84. For different types of lake, we estimate  $6.46 \pm 0.61 \text{ Tg CH}_4 \text{ yr}^{-1}$  (mean  $\pm$  standard deviation) for temperate lakes,  $2.98 \pm 0.67 \text{ Tg CH}_4$   
185  $\text{yr}^{-1}$  for boreal lakes,  $2.08 \pm 0.17 \text{ Tg CH}_4 \text{ yr}^{-1}$  for thermokarst lakes, and  $2.93 \pm 0.43 \text{ Tg CH}_4 \text{ yr}^{-1}$  for yedoma lakes, respectively. The TEM-MDM model driven with WAD2M 2.0 inundation data estimates land ecosystem net emissions of  $21.35 \pm 1.15 \text{ Tg CH}_4 \text{ yr}^{-1}$ , ranging from 19.13 in 2009 to 23.49 in 2007. Combined the two model simulations along with two dynamic area change datasets, we estimate that the total annual methane emission from inland water systems in the region of  $45^\circ \text{ N}$  north during 2000-2015  
190 is  $35.80 \pm 2.59 \text{ Tg CH}_4 \text{ yr}^{-1}$ , with the lowest value of  $31.30 \text{ Tg CH}_4 \text{ yr}^{-1}$  in the year 2000 and the highest value of  $39.56 \text{ Tg CH}_4 \text{ yr}^{-1}$  in 2015 (Fig. 1a & Table 1).

Fig. 1b shows the landscape change over the 2000-2015 period. From which, the wetland area was calculated using inundation fraction data and the lake area was directly derived from the GLCP dataset. The total annual average area of the inland water system in the study region is  $3,090,690 \pm 38,203 \text{ km}^2$   
195 with a minimum value of  $3,039,565 \text{ km}^2$  in 2003 and a maximum of  $3,169,494 \text{ km}^2$  in 2015. The total wetland area is  $1,122,493 \pm 36,303 \text{ km}^2$  ranging from  $1,074,079 \text{ km}^2$  (2009) to  $1,199,428 \text{ km}^2$  (2010). For lakes, the total area ranges from  $1,919,652 \text{ km}^2$  in 2003 to  $1,996,625 \text{ km}^2$  ( $1,968,197 \pm 19,708 \text{ km}^2$ ).



Table 1. Annual methane emissions and landscape area data

Year	Wetland emission (Tg CH <sub>4</sub> yr <sup>-1</sup> )	Lake emission (Tg CH <sub>4</sub> yr <sup>-1</sup> )	Total emission (Tg CH <sub>4</sub> yr <sup>-1</sup> )	Total lake area/ km <sup>2</sup>	Total wetland area/ km <sup>2</sup>	Total area/ km <sup>2</sup>
2000	19.79	11.51	31.30	1967248	1096965	3064213
2001	21.06	12.18	33.23	1962073	1121324	3083397
2002	21.43	11.75	33.17	1960859	1194271	3155130
2003	20.80	12.18	32.97	1919652	1142169	3061821
2004	19.89	12.80	32.69	1958535	1099228	3057762
2005	22.21	14.16	36.37	1969582	1141135	3110717
2006	21.12	14.27	35.39	1972748	1120970	3093718
2007	23.49	14.93	38.41	1980094	1130346	3110439
2008	20.88	14.96	35.84	1965655	1110515	3076170
2009	19.13	15.04	34.17	1980370	1074079	3054450
2010	21.57	15.63	37.21	1970067	1199428	3169495
2011	22.28	16.21	38.48	1971288	1101222	3072510
2012	22.66	15.67	38.32	1934793	1104772	3039565
2013	21.98	16.52	38.50	1988811	1091506	3080317
2014	20.86	16.34	37.20	1992754	1081296	3074051
2015	22.41	17.15	39.56	1996625	1150670	3147294



200 Figure 1. Annual (a) methane emissions and (b) landscape change.

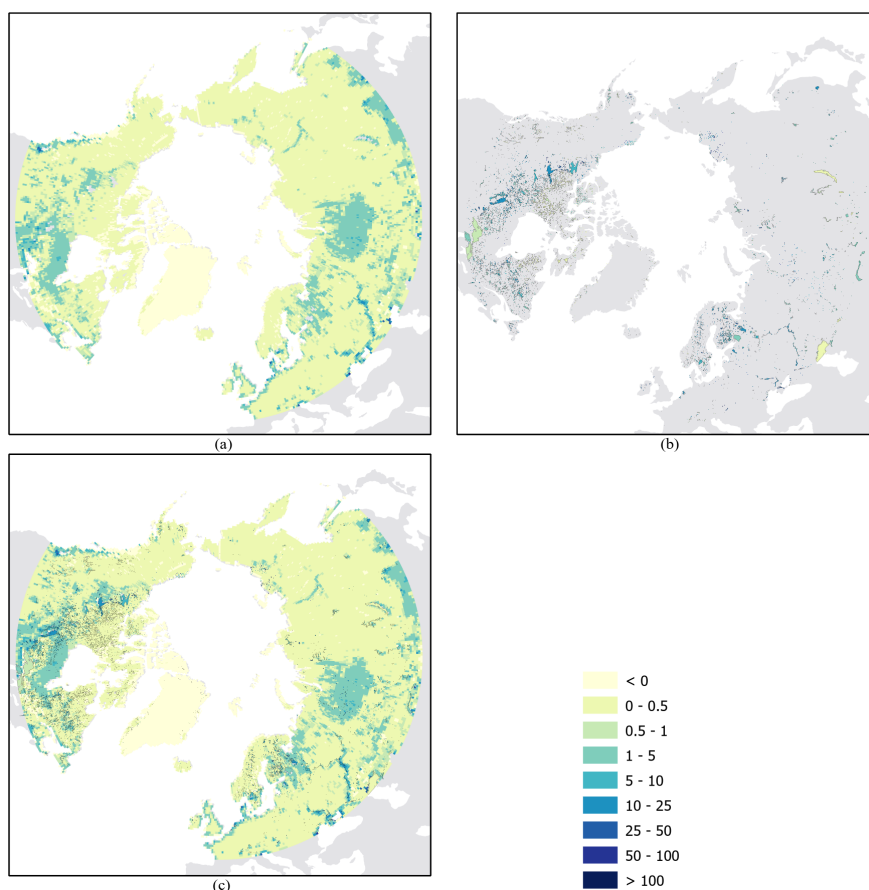
### 3.2. Spatial variations of landscape-level methane emissions

Spatial wetland and lake methane emissions are shown in Fig. 2a and b separately. West Siberia Lowland and the Hudson Bay Lowland were the two strong sources. There are many sporadic high emission sources in wet tundra and small wetlands in boreal forest regions, and river and coastal floodplains. Although a majority of lakes are located in the northern Hudson Bay area, they all have low emissions at around 1 g CH<sub>4</sub> m<sup>-2</sup> yr<sup>-1</sup>, compared to which, lakes near Mackenzie River delta of Canada and the Hudson Bay Lowland area have a relatively higher emission at 50 g CH<sub>4</sub> m<sup>-2</sup> yr<sup>-1</sup>, as well as lakes in northern Europe such as Sweden, Finland, and the northwest corner of Russia (around Lake Onega).





210 Fig. 2c shows the methane emission for inland water systems in the pan-Arctic area, from which wetland and lake emission maps overlapped very well, suggesting our two model simulation results agree with each other. It is worth noting that the average emissions of the lake are usually higher than the emission of the wetlands around the lake, indicating that lakes emit more methane than wetlands in same the region under the same conditions.



215 **Figure 2.** Spatial distribution of average annual methane emissions ( $\text{g CH}_4 \text{ m}^{-2} \text{ yr}^{-1}$ ) from (a) Wetlands, (b) Lakes, and (c) total inland water systems in the pan-Arctic region.

### 3.3. Correlation and sensitivity analysis results

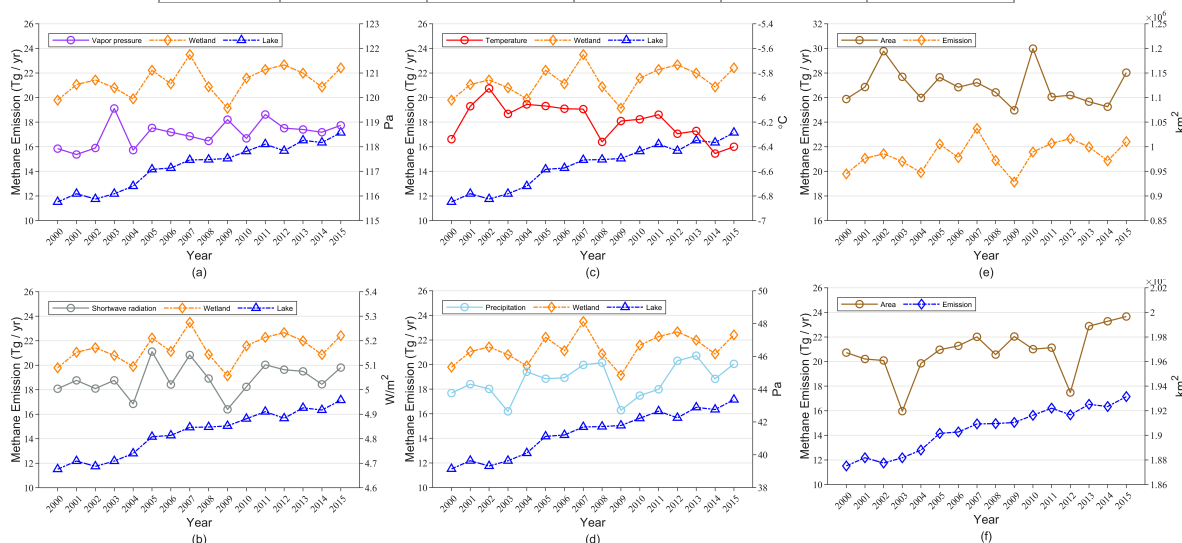
220 The relationship between annual methane emissions from inland water systems and climate drivers as well as landscape areal change are shown in Fig. 3. The studied climate drivers include vapor pressure (relative humidity), precipitation, temperature, and shortwave radiation. For areal changes, wetlands and lakes are shown separately. We also did a correlation analysis between annual methane emissions and



these drivers and the results are shown in Table 2. The shortwave radiation is fitted well with wetland emission with a high correlation of 0.88, which is the only one has a P-value less than 0.01. The precipitation captured the upward and downward trends of wetland emissions, with a relatively high and statistically significant correlation of 0.56. Compared to which, the vapor pressure and temperature have lower correlations with wetland emissions. For lake emissions, vapor pressure captured the most upward and downward trends, followed by precipitation with statistically significant correlations of 0.47 and 0.45, respectively. Although the annual average temperature shares a similar annual trend with lake methane emissions, the correlation analysis is low (Fig. 3c). In addition, the methane emissions from lakes (0.56) are more sensitive to landscape areal changes than to wetlands changes (0.27 with no statistical significance).

**Table 2. Correlations between annual methane emissions and climate drivers and landscape changes: (a) P-value less than 0.1; (b) P-value less than 0.05; (c) P-value less than 0.01**

	Vapor pressure	Precipitation	Temperature	Shortwave radiation	Areal change
Wetland Emission	0.20	0.56 <sup>b</sup>	0.31	0.88 <sup>c</sup>	0.35
Lake Emission	0.47 <sup>a</sup>	0.45 <sup>a</sup>	0.002	0.35	0.56 <sup>b</sup>



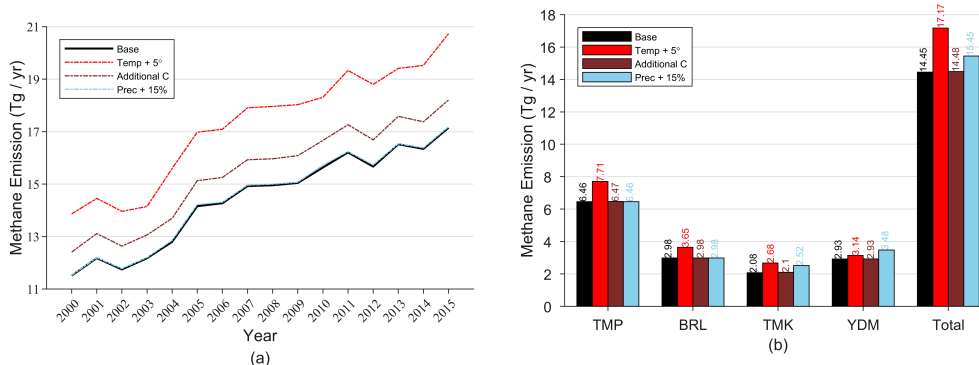
**Figure 3. Relationships between methane emissions from inland water systems and (a) vapor pressure, (b) shortwave radiation, (c) temperature, (d) precipitation, (e) wetland area, and (f) lake area changes.**



235 Different types of lakes have various sensitivities to increasing temperature, precipitation, and  
 additional lake sediment carbon (Fig. 4 and Table 3). Lake methane emission from above 45-degree  
 north is more sensitive to temperature changes than to precipitation or lake sediment carbon pool. When  
 temperature increases by 5° C, lake emissions increase by 19%, where thermokarst lakes are influenced  
 the most (28.5%) and yedoma lakes are influenced the least (7.35%). Precipitation has low impacts on  
 240 lake CH<sub>4</sub> emissions. The overall lake emissions only increase by 0.19% when the precipitation increased  
 by 15%. Thermokarst lakes remain relatively most sensitive to changes in precipitation (0.82), while the  
 other three types of lakes were all insensitive. For additional sediment carbon added due to permafrost  
 thaw, only thermokarst and yedoma lakes were impacted, with increasing by 15% carbon leading to a  
 similar increase for both types of lakes (20.85% and 18.98%), resulting in an overall CH<sub>4</sub> emission  
 245 increase by 6.85%.

**Table 3. Average increase for 4 types of lakes (temperate (TMP), boreal (BRL), thermokarst (TMK), and yedoma (YDM)) and total CH<sub>4</sub> emissions in 16-year period due to changes in temperature, precipitation, and lake sediment carbon.**

	TMP	BRL	TMK	YDM	Total
<b>Additional C</b>	0	0	20.85%	18.98%	6.85%
<b>Temperature</b>	19.24%	22.38%	28.49%	7.35%	18.81%
<b>Precipitation</b>	0.12%	0.05%	0.82%	0.06%	0.19%



250 **Figure 4. Sensitivity test for increasing temperature by 5° C, increasing precipitation by 15%, and adding additional 15% carbon into lake sediments (a); Average value of each type of lakes including temperate (TMP), boreal (BRL), thermokarst (TMK), and yedoma (YDM) (b).**



## 4. Discussion

### 4.1. Annual methane emissions from the landscape

From the previous studies, Wik et al. (2016) estimated 16.5 Tg CH<sub>4</sub> yr<sup>-1</sup> emissions from lakes and ponds north of 50° N while Bastviken et al. (2011) estimated 13.4 for the inland waters (lakes, reservoirs, streams, and rivers) >54° N, both of which are estimated using measurement data combined with inventories. Based on a new spatially-explicit dataset of lakes > 50° N which includes not only all the lakes that area greater than 0.1 km<sup>2</sup> but also 6.5 million smaller lakes (0.02–0.1 km<sup>2</sup>), Matthews et al. (2020) estimated the emissions are 13.8–17.7 Tg CH<sub>4</sub> yr<sup>-1</sup>. Using a process-based model (bLake4Me, a previous version of the ALBM model), Tan and Zhuang (2015a) estimated 11.86 Tg CH<sub>4</sub> yr<sup>-1</sup> in the year 2005–2008 ranging from 7.1 to 17.3 Tg CH<sub>4</sub> yr<sup>-1</sup> for north of 60° N. After this study, a coupled model of bLake4Me and a thermokarst lake-evolution model was used to estimate a total methane emission of 11.3 ± 2.1 Tg CH<sub>4</sub> yr<sup>-1</sup> from lakes >60° N in the year 2006 (Tan and Zhuang, 2015b). Compared to these estimates, our lake simulation results fall in a reasonable range.

For emissions from northern high latitude wetlands, Chen et al. (2015) estimated 36.1 ± 6.7 Tg CH<sub>4</sub> yr<sup>-1</sup> during 1997–2006 for the same pan-Arctic wetlands (north of 45° N) using an enhanced Variable Infiltration Capacity (VIC) model linked with the Walter and Heimann wetland CH<sub>4</sub> emissions model. Zhang et al. (2017) used a bottom-up approach with LPJ-wsl model, estimating methane emissions of 23.4 ± 0.76 Tg CH<sub>4</sub> yr<sup>-1</sup> from wetlands > 50° N over the period 1980–2000. Poulter et al. (2017) used an ensemble of biogeochemical models constrained with remote sensing surface inundation and inventory-based wetland area data (SWAMPS–GLWD, a previous version of WAD2M used in this study) estimating the boreal wetland emitted 44 ± 19 Tg CH<sub>4</sub> yr<sup>-1</sup> in 2012. Using TEM-MDM, but combined with different transient wetland inundation area fraction datasets, Liu et al. (2020) estimated the emissions are 38.90 Tg CH<sub>4</sub> yr<sup>-1</sup> from the region 45–90° N. Our estimates are at the lower end of these records. We attribute this to the change in inundation area data. The revision in WAD2M version 2.0 from version 1.0 slightly reduced vegetated wetland extent in the mid-latitudes especially for the region 45–70° N to further avoid the “double accounting” problem, which is the region with the most methane emissions in our study area. Compared to other model simulations that were also involved in the same project (Global Carbon Project wetland CH<sub>4</sub>, GCP-CH<sub>4</sub>) where 16 models give an annual average CH<sub>4</sub> emission of 28.8 ± 11.8 Tg CH<sub>4</sub> yr<sup>-1</sup> from northern wetlands >45° N in 2000–2020.



Besides biogeochemistry modeling approaches, atmospheric chemistry transport and inversion models have also been used to constrain the methane emission quantification from pan-Arctic wetlands and lakes. Bruhwiler et al. (2014) developed an assimilation system for atmospheric CH<sub>4</sub> and simulated the annual emissions from the wetland over the northern high latitudes (53–90° N) of about 23 Tg CH<sub>4</sub> yr<sup>-1</sup>. Tan et al. (2016) used a nested-grid high-resolution inverse model estimating methane emissions from north of 60° N in the range of 11.9–28.5 Tg CH<sub>4</sub> yr<sup>-1</sup>, of which wetlands and lakes accounted for 5.5–14.2 and 2.4–14.2 Tg CH<sub>4</sub> yr<sup>-1</sup>, respectively.

Based on the above comparisons with other estimates using either bottom-up or top-down methods for wetland and lake separately, we believe that our estimation narrows the gap between biogeochemistry modeling and atmospheric modeling since our landscape-level modeling avoids the uncertainty induced by the “double accounting” problem raised by Thornton et al. (2016).

#### 4.2. Climate drivers and sensitivity analysis

Considering that annual average values are not capable of capturing the true relationship (statistically insignificant), we then did another correlation analysis using monthly data. From which, monthly emissions, four climate drivers, and wetland inundation data were used while monthly lake area data are not available (Table 4). Except for the correlation between lake emission and areal change, each correlation in the table has a p-value lower than 0.01, which means they are all statistically significant. Vapor pressure, although not fully coincides with wetland emissions in interannual trends (Fig. 3), the monthly correlation is the highest among the five factors (0.96). Other studies also indicated that the impact of wet/dry cycles on regional methane emissions is evident, reflected in soil moisture effects (Watts et al., 2014). The second highest correlation with wetland emissions is temperature, followed by wetland area, shortwave radiation, and precipitation. Although the interannual variation of short-wave radiation most coincided with wetland emissions, their monthly correlation was only 0.77. In terms of the correlation of lake methane emissions, temperature has the highest value of 0.87, followed by relative humidity (vapor pressure) and precipitation. Since the correlation between lake methane emissions and area is still calculated using inter-annual data, the correlation of 0.56 can also be considered as high. We also did a correlation analysis between wetland area and climate drivers and found that temperature and vapor pressure are the climatic factors that have the greatest impact on wetland landscape areal changes.



310 **Table 4. Correlation between monthly methane emission and climate drivers and landscape changes**

	Vapor pressure	Precipitation	Temperature	Shortwave radiation	Areal change
Wetland Emission	0.96	0.76	0.89	0.77	0.79
Lake Emission	0.82	0.79	0.87	0.56	0.56
Wetland Area	0.88	0.75	0.93	0.69	

The sensitivity analysis suggests that a 5 °C increase in temperature increases the pan-Arctic lake methane emission by 20%. Compared to previous studies, Guo et al. (2020) estimated a 40% lake methane emission increase for the same study area by the end of the 21<sup>st</sup> century in the scenario that the temperature increases around 7.5 °C. Sepulveda-Jauregui et al. (2018) showed that for subarctic oligotrophic lakes, increasing lake water temperature by 2 °C leads to a net increase in CH<sub>4</sub> emissions by 47-56%. However, their work did not consider the ice cover season of high-latitude lakes, from which the methane fluxes can be blocked by a thick layer of ice for several months each year and then oxidized in the water column. In addition, the relatively low response of yedoma lakes (~7%) to the increasing temperature could be explained by their mobilized labile carbon is usually in deep sediments (Tan and Zhuang, 2015a), which means that the influence of the warming air temperature will take much longer to enhance methane production in the lake sediment. In contrast, when we directly increase the labile carbon density ( $C_{labile}$ ) at the water-sediment interface, the methane emission of yedoma lakes increased much higher (~19%), while the thermokarst lake were affected less (~20%) compared to its response to temperature change (~28%). For precipitation, although it was set in the model to bring the load of allochthonous carbon to the lake (Tan et al., 2017), increasing it by 15% only makes a negligible impact on methane emission. A plausible explanation is that the lakes are relatively saturated with extraneous carbon in sediments, so any increase brought by the additional precipitation tends to have small influences.

#### 4.3. Uncertainty analysis and future works

330 Although our simulation results more accurately estimate methane emissions from inland water systems in the pan-Arctic by avoiding the “double accounting” problem, there still exist some uncertainty



sources in this study. First, despite the use of two non-overlapping landscape change maps to avoid the uncertainty caused by “double accounting”, the precision of the two maps remains to be examined. The HydroLAKES database used in the GLCP and WAD2M datasets only contains lakes and reservoirs which  
335 area greater than 0.1 km<sup>2</sup> (Messenger et al., 2016), which means that lakes and ponds smaller than 0.1 km<sup>2</sup> are either not considered or misclassified as wetlands. Those small lakes and ponds cover in total about 1 × 10<sup>6</sup> km<sup>2</sup> which equals more than half the area of Alaska (Verpoorter et al., 2014). Also, some previous studies have found higher methane fluxes in small and shallow lakes (Holgerson et al., 2016; Sasaki et al., 2016), and lakes appear to emit more methane than wetlands, implying that lake methane emissions  
340 may still be underestimated. Secondly, during the simulation, although we classified the lakes based on their sediment type, size, and depth, we still assumed that all the same types of lakes to be homogeneous which were assigned to the same set of parameters. Nevertheless, lakes are highly heterogeneous across the globe (Guo et al., 2021), especially for those big lakes, such that regional lake simulation may introduce a high uncertainty.

345 Furthermore, recent studies have found that groundwater discharge could be an important pathway as lateral CH<sub>4</sub> inputs to Arctic lakes that links CH<sub>4</sub> production in thawing permafrost to atmospheric emissions via lakes (Olid et al., 2022). Jammet et al. (2015) also confirmed that spring is a crucial period for methane dynamics in subarctic shallow lakes while large methane emissions were observed during the spring thaw. These two important processes were not considered in our process-based ALBM model.  
350 Similarly, compared to other model results in GCP-CH<sub>4</sub> projects, our TEM-MDM modeled wetland CH<sub>4</sub> emissions are relatively low in subzero temperature months, while a field study found that substantial emissions occur during the “zero curtain” period, when subsurface soil temperatures are poised near 0 °C (Zona et al., 2016). Therefore, our next step will be modifying the TEM-MDM and ALBM models by taking those important processes into consideration. In addition, higher resolution maps of dynamic  
355 wetland inundation and lake landscape changes are highly needed.

## 5. Conclusions

By using two dynamic areal change datasets combined with process-based terrestrial and lake biogeochemical models, we are among the first to quantify methane emissions from both land and aquatic inland water systems, i.e., wetlands and freshwater bodies in the pan-Arctic, which avoids the uncertainty



360 caused by area “double accounting”. Our simulations indicate that the total methane emissions from pan-  
Arctic inland water system are 35.81 Tg CH<sub>4</sub> yr<sup>-1</sup> during 2000-2015, of which wetlands and lakes were  
21.38 Tg yr<sup>-1</sup> and 14.45 Tg yr<sup>-1</sup>, respectively. Our estimation narrows the gap between estimates with  
biogeochemistry models and atmospheric chemistry transport and inversion models. In the pan-Arctic,  
wetland methane emissions are most affected by vapor pressure (relative humidity), followed by  
365 temperature and landscape changes, while lake emissions are more sensitive to temperature than to  
relative humidity (vapor pressure) and precipitation. Furthermore, the methane emissions from wetlands  
are more sensitive to landscape areal changes than from lakes. West Siberia Lowland and the Hudson  
Bay Lowland were the two strong sources of wetlands and lakes have higher emissions around  
Mackenzie River delta of Canada and the Hudson Bay Lowland area. In addition, lakes emit more  
370 methane than wetlands under the same condition. Although the lack of understanding of the underlying  
methane cycle mechanisms in the lake makes the response of CH<sub>4</sub> emissions from Arctic lakes to climate  
change highly uncertain, our sensitivity test using the process-based model ALBM does indicate the pan-  
Arctic Lake CH<sub>4</sub> emissions were highly influenced by temperature, but less by lake sediment carbon  
increase.

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*Code and data availability.* The data used to reproduce figures, codes, model and samples of running  
directory can be accessed via Purdue University Research Repository:  
<https://purr.purdue.edu/publications/4166/1>.

380 *Author contributions.* QZ designed the study. XL conducted model simulation and analysis. XL and QZ  
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*Competing interests.* The authors declare that they have no conflict of interest.

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