



1 **Enabling a process-oriented hydro-biogeochemical model to simulate soil**
2 **erosion and nutrient losses**

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20 **Abstract**

21 Water-induced erosion and subsequent particulate carbon (C), nitrogen (N) and phosphorus
22 (P) nutrient losses were the vital parts of biogeochemical cycling. Identifying their intensity and
23 distribution characteristics is of great significance for the control of soil and water loss and N/P
24 nonpoint source pollution. This study incorporated the modules of physical soil erosion and the
25 particulate C, N and P losses into the process-oriented hydro-biogeochemical model
26 (CNMM-DNDC) to enable it to predict soil and water loss. The results indicated that the
27 upgraded CNMM-DNDC i) performed well in simulating the observed temporal dynamics and
28 magnitudes of surface runoff, sediment and particulate N/P losses in the lysimetric plot of the
29 Jieliu catchment in Sichuan Province; ii) successfully predicted the observed monthly dynamics
30 and magnitudes of stream flow, sediment yield and particulate N losses at the catchment outlet,
31 with significant zero-intercept univariate linear regressions and credible Nash–Sutcliffe indices
32 larger than 0.74. The particulate N accounted for 16.2%–26.6% of the TN components during
33 the period with larger precipitations. The intensities of soil erosion and particulate nutrient
34 losses in the Jieliu catchment was closely related to land use type in the order of sloping
35 cultivated cropland > residential area > forest land. The scenario analysis demonstrated that
36 high greenhouse gas (GHG) emissions scenarios provided a greater risk of soil erosion than did
37 low GHG emissions scenarios and land use change could help to mitigate soil and water loss
38 accelerated by climate change in the future. The upgraded model was demonstrated to have the
39 capability of predicting ecosystem productivity, hydrologic nitrogen loads, emissions of GHGs
40 and pollutant gases, soil erosion and particulate nutrient losses, which may become a decision



41 support tool for soil erosion and nonpoint source pollution control coordinated with increasing
42 production and reducing GHGs and pollutant gases emissions in a catchment.

43 **Keywords**

44 CNMM-DNDC, ROSE, soil erosion, particulate carbon/nitrogen/phosphorus loss

45 **1. Introduction**

46 Water-induced erosion and subsequent particulate carbon (C), nitrogen (N) and
47 phosphorus (P) nutrient losses are among the primary threats leading to the decline in soil
48 fertility and the increases in land degradation, channel sedimentation and eutrophication of
49 downstream rivers and lakes (Berhe et al., 2018; Ekholm and Lehtoranta, 2012; Garcia-Ruiz et
50 al., 2015). This global environmental issue has continued to deteriorate (Ma et al., 2021; Yang
51 et al., 2003). A previous study found that the vulnerability of water-induced erosion increased
52 over 51% of the global surface from 1982 to 2015 (Liu et al., 2019). Climate change and
53 anthropogenic activities (such as land use change) are the two principal driving forces that have
54 complicated and altered the hydrological cycle and water-induced erosion during recent decades
55 (Piao et al., 2007; Zeng et al., 2015).

56 Quantitative assessments of the water-induced soil erosion intensity and identification of its
57 temporal and spatial distribution characteristics are of great importance for preventing soil and
58 water loss and have attracted the attention of researchers (e.g., Jetten et al., 2003; Jiang et al.,
59 2017; Panagos et al., 2015). Lysimetric plot experiments have been developed as a direct field
60 measurement method for the accurate quantification of surface runoff and water-induced
61 erosion (e.g., Kosmas et al., 1997; Sumner et al., 1996; Zhu et al., 2009). However, the in situ
62 field measurements of water-reduced soil erosion with high cost of labor and money can only



63 cover a small piece of the sampling units. It is unrealistic to expect direct field measurements
64 to quantify water-induced erosion everywhere under various conditions.

65 Simulations of mathematical models are likely to compensate for the deficiency of direct
66 field measurements. The Universal Soil Loss Equation (USLE, Wischmeier and Smith, 1978)
67 and its revised version (RUSLE) (Renard et al., 1997) have been developed into widely used
68 empirical mathematical models to directly calculate soil erosion based on rainfall, soil property,
69 topography, cover and management data. The USLE or RUSLE quantify only the various
70 influencing factors that impact the soil loss associated with soil erosion, which is not directly
71 related to the process of surface runoff and does not involve the specific process of sediment
72 transport yet (Donovan, 2022; Meinen and Robinson, 2021). Fortunately, the physical
73 process-based ROSE model named after the name of developer (Rose et al., 1983)
74 conceptualizes the soil erosion process by conceiving three continuous and simultaneous
75 physical processes, including rainfall detachment, sediment entrainment and sediment
76 deposition, thus providing good performance in estimating sediment yield at the plot scale.
77 However, the ROSE model focuses only on the physical processes of water-induced erosion
78 without engaging the C and N cycles of the ecosystem. The Soil and Water Assessment Tool
79 (SWAT) (Arnold et al., 1998), a semi distributed hydrological model, incorporates the RUSLE
80 to predict soil erosion at the level of hydrological response units, in which the routing of
81 sediment transportation is not considered and the modeling of the biogeochemical element
82 cycle is relatively simple and empirical (Ferrant et al., 2011; Pohlert et al., 2007). However, the
83 transport of particulate C, N, and P nutrients accompanied by water-induced erosion crucially



84 depends on the C and N cycles of ecosystems in a catchment. Therefore, knowledge of the
85 coupling between the process-oriented hydro-biogeochemical model combined with the
86 complicated C and N cycles and the soil erosion model based on physical processes (such as
87 ROSE) is essential to accurately predict soil erosion and subsequent particulate C, N, and P
88 nutrient transport.

89 A recently developed hydro-biogeochemical model (CNMM-DNDC) by Zhang et al.
90 (2018) might become a realistic tool that can be used to address the abovementioned problem.
91 The CNMM-DNDC model introduces the complicated C and N biogeochemical modules
92 (including the modules of decomposition, nitrification, denitrification and fermentation) of a
93 widely used biogeochemical model (DeNitrification-DeComposition model, DNDC, Li et al.,
94 1992) into the distributed hydrological framework of the Catchment Nutrients Management
95 Model (CNMM, Li et al., 2017). The CNMM-DNDC model has been used to conduct a
96 comprehensive simulation of the complicated hydrological and biogeochemical processes
97 (such as ecosystem productivity, hydrologic N loads, gaseous N losses and greenhouse gas
98 emissions) of a subtropical catchment with various landscapes (Zhang et al., 2018), a model
99 evaluation of nitrous oxide and nitric oxide emissions from a subtropical tea plantation (Zhang
100 et al., 2020b), a model evaluation and regional simulation of nitrate leaching in the black soil
101 region of Northeast China (Zhang et al., 2021a) and a comprehensive model modification and
102 evaluation of NH₃ volatilization from fertilized croplands (Li et al., 2022b). However, the
103 CNMM-DNDC model still lacks the capacity to simulate the processes of soil erosion and
104 subsequent particulate C, N, and P nutrient transportation.



105 Therefore, we hypothesize that the accurate simulation of soil erosion and subsequent
106 particulate C, N, and P nutrient losses can be realized by incorporating the soil erosion physical
107 model and the element enrichment module into the process-oriented hydro-biogeochemical
108 model with complicated C and N cycles. Based upon the above hypothesis, the objectives of
109 this study were to i) introduce the ROSE model (a physical soil erosion model) and the
110 enrichment module of the particulate nutrients into the hydrological process of the
111 CNMM-DNDC model; ii) evaluate the performance of the CNMM-DNDC model in simulating
112 the temporal and spatial distributions of soil erosion and subsequent particulate C, N, and P
113 transportation at the plot and catchment scales; and iii) investigate the impact of climate change
114 and human activities (such as land use change) on the losses of soil and particulate nutrients.

115 **2. Materials and methods**

116 **2.1 Catchment description**

117 The Jieliu catchment (31°16'N, 105°28'E, 400–600 m a.s.l.), located in Sichuan Province
118 of Southwest China (Zhu et al., 2009), was used for the model calibration and validation. This
119 catchment is situated in the upper reaches of the Yangtze River and has a typical subtropical
120 monsoon climate. During the period from 2005 to 2018, the annual mean temperature was
121 16.7 °C, and the average annual precipitation was 720 mm, 75% of which occurred during the
122 period between June and September (<http://yga.cern.ac.cn>). The soil in the catchment is
123 dominated by Calcaric purple soil, classified as a Pup-Orthic Entisol in the Chinese Soil
124 Taxonomy or as an Entisol classified in the U.S. Soil Taxonomy (Zhu et al., 2009). The total
125 area of the Jieliu catchment is approximately 35 ha, and it is dominated by sloping croplands
126 (54%) and forest lands (31%). The primary crops cultivated in the sloping croplands are maize



127 (Zea mays L.), winter wheat (*Triticum aestivum* L.), rape (*Brassica napus* L.) and rice (*Oryza*
128 *sativa* L.). The N, P and potassium (K) fertilizers are applied at rates of 130–330 kg N ha⁻¹ yr⁻¹
129 (ammonium bicarbonate or urea), 72–162 kg P ha⁻¹ yr⁻¹ (calcium superphosphate) and 45–68
130 kg K ha⁻¹ yr⁻¹ (potassium chloride), respectively (Zhang et al., 2018).

131 **2.2 Model modifications**

132 The CNMM-DNDC model can simulate the lateral movements of water-soluble nutrients
133 (e.g., ammonium, nitrate, phosphate and dissolved organic matter) by surface and subsurface
134 runoff, whereas it lacks the capabilities of simulating soil erosion and sediment transport
135 caused by surface runoff and the subsequent transportation of particulate C, N, and P. To
136 address such a deficiency, this study incorporated the modules of soil erosion and element
137 enrichment into the lateral hydrological framework of the CNMM-DNDC model. Therefore,
138 the upgraded CNMM-DNDC model was equipped with the ability to estimate the movements
139 of soil particles and particulate nutrients transported with surface runoff in the lateral
140 dimension. The soil erosion module adopted the simplified ROSE model (Rose et al., 1983;
141 Stewart, 1985), which is a process-oriented soil erosion model. The ROSE model is based on
142 the dynamic equilibrium of three simultaneous processes, including rainfall detachment, runoff
143 detachment, and sediment deposition. In an individual erosion event, the process of runoff
144 detachment dominates, and the latter two processes of rainfall detachment and sediment
145 deposition can be generally neglected. Therefore, in the simplified ROSE module, as shown by
146 Eq. (1), the sediment yield (Y_s , kg dry soil ha⁻¹) resulting from soil erosion was driven by
147 surface runoff (R_s , m) and concomitantly regulated by the land's slope (S_i , dimensionless) and



148 the coverage fraction of vegetation (C_v , fraction).

$$Y_s = 27 \times 10^6 (1 - C_v) \eta S_1 R_s \quad (1)$$

149 Where R_s is calculated from the existing hydrological module of the CNMM-DNDC
150 model, in which R_s occurs in the following two cases. First, R_s is caused by the mechanism of
151 excess infiltration, in which the water input (i.e., precipitation and irrigation) is greater than the
152 maximum infiltration capacity of the soil. Second, R_s is derived from the mechanism of excess
153 storage, in which precipitation or irrigation still occurs when the soil surface water content
154 exceeds the corresponding saturated water content. The direction of the surface runoff conflux
155 is estimated by the distributed weights of four neighboring grids (i.e., in the upper, lower, left
156 and right directions), which are calculated based on the elevation of these grids. S_1 is the sine
157 value of the slope radian value. In this study, the value of C_v for the crop is approximately
158 equivalent to the growing index, which is estimated by the ratio of the accumulated
159 temperature from sowing to the present time to the accumulated thermal degree for maturity in
160 the plant growth module. Particularly, using the observed sediment yield in the catchment
161 outlet, the value of C_v for the natural vegetation (e.g., forest and grass) was addressed and
162 calibrated as half of the ratio of the real leaf area index (LAI) and the maximum LAI, which is
163 one of the model inputs. The value of C_v of the artificial lands (e.g., the urban or rural
164 residential areas) was calibrated and set to 0.1, which represented the effects of concrete roads
165 and residential buildings on the reduction of the soil area exposed to erosion. Usually, the
166 values of C_v need to be calibrated by the soil exposure ratio and sediment yield observations



167 for a given study area. η (dimensionless) is referred to as the efficiency of sediment entrained
168 by surface runoff, which depends on soil texture and C_v , as shown in Eq. (2). In Eq. (2), a_1 is
169 referred to as the rate of sediment carried by surface runoff on bare land, which differs for
170 various soil types and generally needs to be calibrated by the observed data of sediment loss for
171 a given study area. Loch and Donnollan (1983) reported that a_1 varies from 1.0% to 8.7% in
172 Middle Ridge clay loam and Irving clay soils. The soil mass balance was not considered in the
173 upgraded model.

$$\eta = a_1 e^{-0.15C_v} \quad (2)$$

174 It is known that the C, N, and P elements of the eroded sediments are usually richer than
175 those of the in situ soils from which the eroded sediments originate (Massey and Jackson, 1952;
176 Schiettecatte et al., 2008; Wan and El-Swaify, 1998). The above phenomenon is usually referred
177 to as sediment enrichment, which can be quantified by an empirically based enrichment ratio
178 (E). E is usually defined as the ratio of the concentration of C, N, and P elements in the eroded
179 sediment to that in the source soil (Sharpley, 1980; Teixeira and Misra, 2005). Generally, as
180 more eroded sediment is produced, the richness of the C, N, and P elements decreases. The
181 enrichment ratio of the C and N nutrients (E_{CN}) is estimated by Eq. (3), which was adapted
182 from McElroy et al. (1976) and Williams and Hann (1978). The pre-exponential factor (k_1) of
183 Eq. (3) was calibrated to 1.2 using the particulate N data observed at the lysimetric plot in this
184 study. The enrichment ratio of P nutrients (E_P) is calculated by Eq. (4) cited from Sharpley
185 (1980).



$$E_{CN} = k_1(Y_s \times 10^{-4})^{-0.2468} \quad (3)$$

$$E_P = e^{(2.46 - 0.21 \log Y_s)} \quad (4)$$

186 The yields of particulate C (P_C , kg C ha⁻¹), N (P_N , kg N ha⁻¹), and P (P_P , kg P ha⁻¹)
187 nutrients caused by soil erosion were calculated based on E , Y_s and the content of the
188 corresponding organic C (C_C , g C ha⁻¹), N (C_N , g N ha⁻¹), and P (C_P , g P ha⁻¹) pools in topsoil
189 using Eqs. (5–7), respectively. BD (g m⁻³) and D_s (m) refer to the soil bulk density and the
190 depth of topsoil, respectively. Eight of the soil organic C and N subpools participated in the
191 process of soil erosion, including the pools of very reliable, reliable, and resistant
192 decomposable litters, reliable and resistant active microbes, reliable and resistant humads and
193 passive humus, whereas five of the soil organic P subpools were involved in the process of soil
194 erosion, including the pools of active and passive organic P, active and dead microorganism P,
195 and inert stable P. The particulate C, N, and P losses calculated by the element enrichment
196 module were also deducted from the corresponding subpools of the topsoil. Subsequently, the
197 eroded soil and the particulate C, N, and P nutrients are transported with surface runoff and
198 eventually drain into streams.

$$P_C = \sum_{i=1}^8 \frac{10^{-7} E_{CN} C_{C_i} Y_s}{BD \cdot D_s} \quad (5)$$

$$P_N = \sum_{j=1}^8 \frac{10^{-7} E_{CN} C_{N_j} Y_s}{BD \cdot D_s} \quad (6)$$



$$P_P = \sum_{k=1}^5 \frac{10^{-7} E_P C_{P_k} Y_s}{BD \cdot D_s} \quad (7)$$

199 **2.3 Preparation for model simulation**

200 The input data for driving the model operation consisted of the meteorological data at the
201 3-hour scale (including average air temperature, solar radiation, long wave radiation, wind
202 speed, humidity and total precipitation), the spatialized soil properties (including soil texture,
203 soil organic carbon, bulk density and pH), the gridded digital elevation model (DEM, Fig. 1)
204 with a resolution of 5 m × 5 m, the spatial distribution of land use (Fig. 1) and cropping systems,
205 and the field management practices. Taking the efficiency of the model calculation and the
206 accuracy of the biogeochemical process description into consideration, the upgraded
207 CNMM-DNDC model conducted a simulation with a grid of 15 m × 15 m from 2004 to 2017,
208 with an initial spin-up period of ten years. The DEM, soil properties, land use, cropping
209 systems, field management practices and meteorological data from 2004 to 2014 were
210 primarily adapted from Zhang et al. (2018). The remaining meteorological data were adapted
211 from the hourly observations provided at the National Science & Technology Infrastructure
212 (<http://rs.cern.ac.cn>). The input data of soil properties, DEM, land use, cropping systems, and
213 field management practices were resampled to the ASCII grids with a resolution of 15 m × 15
214 m using the ArcGIS 10.0 software package (ESRI, Redland, CA, USA). The observation data
215 measured at the lysimetric plot and the catchment outlet (Fig. 1) contributed to model
216 calibration and validation. The surface runoff, subsequent sediment yield, and particulate and
217 total N losses from 2004 to 2006 with three replicates and the surface runoff, subsequent
218 sediment yield, and total P loss from 2017 to 2018 with three replicates measured at the



219 lysimetric plots were adapted from Deng et al. (2011) and Hu (2020), respectively. The
220 monthly stream flow, sediment yield, and particulate and total N losses from 2007 to 2008
221 measured at the catchment outlet were directly cited from Deng et al. (2011). Total N referred to
222 the total amount of NH_4^+ , NO_3^- , dissolved organic N and particulate N. Total P referred to the
223 total amount of dissolved organic and inorganic P and particulate P. Among them, the observed
224 data from the lysimetric plot in 2004 (with seven observation times) and 2016 (with four
225 observation times and a heavy precipitation event) and the observed data from the catchment
226 outlet in 2007 were used for model calibration, and the remaining observed data were used for
227 model validation.

228 **2.4 Climate and land use scenario settings**

229 Scenario analysis was adopted to assess the impact of climate change and land use change
230 on water-induced erosion and its accompanying nutrient losses. The annual accumulated yields
231 of sediment and particulate C, N, and P nutrient losses at the outlet in 2008 (the year for model
232 validation) simulated by the upgraded model were used as the baseline values of the scenario
233 analysis. Two groups of climate change and land use change scenarios were designed:
234 single-factor change and multifactor change scenarios. The single-factor change scenarios
235 altered only one factor while keeping the others constant. The single-factor change scenarios of
236 climate change consisted of two parts. One part for air temperature change was altered within
237 the range of $-4\text{ }^\circ\text{C}$ to $+4\text{ }^\circ\text{C}$ with an interval of $0.2\text{ }^\circ\text{C}$ (abbreviated as $T_{\text{air}+}$ the increase value
238 of air temperature or $T_{\text{air}-}$ the decrease value of air temperature). The other part for
239 precipitation change was altered by the range from -30% to $+30\%$ with an interval of 2%



240 (abbreviated as $P+$ the increase percentage of precipitation or $P-$ the decrease percentage of
241 precipitation). A single-factor change scenario of land use was designed as the sloping upland
242 changed into forest land with the lower soil erosion rate (i.e., FL scenario). The existing land
243 use conversion to another type, such as the change from cropland to forest land or some other
244 land use, is a kind of compromise and required a sensitivity analysis to the model simulation
245 rather than representing the conditions of the real natural system (Dey and Mishra, 2017). The
246 multifactor change group was designed to simultaneously consider climate and land use change.
247 The IPCC's Summary for Policy-makers (IPCC, 2021) points out that the average annual
248 global land precipitation is projected to increase by 10.5% and 30.2% at the 1.5 °C and 4 °C
249 warming levels, respectively. According to the correspondence between climate warming and
250 increasing precipitation in the IPCC's AR6, the multifactor change scenarios designed two
251 multiple climate change scenarios: the low and high greenhouse gas (GHG) emissions
252 scenarios. The low GHG emissions scenario represents air temperature and precipitation
253 increasing by 1.5 °C and 10%, respectively, while the high one represents air temperature and
254 precipitation increasing by 4 °C and 30%, respectively. For the sake of argument, we also
255 divided air temperature and precipitation single-factor scenarios into four sets. The scenarios
256 with air temperature increases greater than 2 °C were defined as the higher warming scenarios,
257 while the lower ones were defined as the scenarios with air temperature changes from 0 °C to
258 2 °C. The scenarios with air temperature reductions greater than 2 °C were defined as the
259 higher cooling scenarios, while the lower cooling scenarios were defined as the scenarios with
260 air temperature changes from -2 °C to 0 °C. Similarly, the scenarios with precipitation



261 increases greater than 20% were defined as the higher rain-enhanced scenarios, while the lower
262 ones were defined as the scenarios with precipitation changes from 0% to 20%. The scenarios
263 with precipitation decreasing more than 20% were defined as the higher rain-reduced scenarios,
264 while the lower rain-reduced scenarios were defined as the scenarios with precipitation change
265 from -20% to 0%. Furthermore, we also explored the effect of the low and high GHG
266 emissions scenarios in a combination of land use change scenarios (i.e., FL scenario) on
267 sediment yield and particulate nutrient yields. The tillage scenario analysis was involved in the
268 scenario analysis of alternative management practices, which were conducted as the no tillage
269 operations of the short term in 2008 and the long term from 2004 to 2008. The relative change
270 deviations of the simulated annual accumulated sediment and particulate nutrient losses of the
271 designed scenarios from the baseline were provided as the quantitative evaluation index for
272 scenario analysis (Abdalla et al., 2020; Dubache et al., 2019).

273 **2.5 Evaluation of model performance and statistical analysis**

274 The performance of the upgraded CNMM-DNDC model in simulating sediment and
275 particulate nutrient losses was evaluated using the normalized root mean square error (nRMSE),
276 the Nash–Sutcliffe index (NSI) and the slope, determination coefficient (R^2) and significance
277 level (p) of the zero-intercept univariate linear regression (ZIR) between the simulation and
278 observation. The nRMSE and NSI values are calculated by Eq. 8 and Eq. 9, respectively. O_i
279 and S_i are the observed and simulated values, respectively. \bar{O} is the mean value of the
280 observed data, and n is the number of paired samples. If the value of nRMSE is closer to 1, the
281 values simulated by the model are more coincident with the observed values (Cui et al., 2014;



282 Smith et al., 1997). The value of the NSI provides the discrepancy between the simulated
283 values and the mean of the observed values, with a positive value indicating an acceptable
284 simulation (Li et al., 2022a). The closer to 1 the slope and R^2 of the ZIR are, the better the
285 simulated values match the observed values. The Origin 8.0 (OriginLab Ltd., Guangzhou,
286 China) and ArcGIS 10.0 software packages were used for graph drawing.

$$\text{nRMSE} = \frac{100}{\bar{O}} \sqrt{\frac{\sum_{i=1}^n (S_i - O_i)^2}{n}} \quad (8)$$

$$\text{NSI} = 1 - \frac{\sum_{i=1}^n (S_i - O_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (9)$$

287 In addition, linear correlations were carried out to study the relationships between the
288 variables relevant to soil erosion and that related to the biogeochemistry process. The R project
289 was applied for the graph drawing of correlation matrix.

290 3. Results

291 3.1 Model performance in simulating soil erosion in the lysimetric plot

292 Among the only eight soil erosion observations conducted at the lysimetric plot from 2015
293 to 2017, four observations in 2016 were provided for model calibration. Given the limited size
294 of the samples, the performance of the upgraded CNMM-DNDC model was revealed using
295 only the graph of the predictions and observations (Fig. 2a–c), without a quantitative
296 evaluation with the above statistical criteria. The temporal dynamic patterns of the simulated
297 surface runoff, sediment and concomitant particulate P yields were in accordance with the
298 observed values when either model calibration or validation was performed (Fig. 2 a–c).
299 Nevertheless, on July 23, which was a heavy precipitation event (213 mm precipitation during



300 the seven days prior to the observation day) in 2016, the upgraded model overestimated the
301 observed sediment yield by approximately 6 times (3.6 versus 0.6 t ha⁻¹, Fig. 2b). However, the
302 simulated surface runoff and total P loss were only approximately 60% and 20% larger than the
303 observed values, respectively. Unfortunately, the simulated peaks of surface runoff and
304 sediment yield at the end of June 2015 lacked the support of the observations. Previous results
305 by the authors also demonstrated that the upgraded CNMM-DNDC performed well in
306 simulating surface runoff, sediment yield, and particulate N loss in 2004 (i.e., the verification
307 period) in the lysimetric plot, with significant ZIRs and credible nRMSE values of 15.2%,
308 32.0%, and 88.0%, respectively (Table S1 and Fig. 2d–g). Moreover, we conducted an
309 evaluation of the simulated and observed NH₄⁺ and NO₃⁻ losses accompanied by surface runoff
310 in the lysimetric plot (Fig. S1). The upgraded model generally captured the temporal variation
311 and magnitude of the observed NH₄⁺ and NO₃⁻ loss, although discrepancies existed in the
312 magnitude of the peak loss (i.e., the model underestimated NH₄⁺ loss caused by approximately
313 100 mm precipitation on September 4, 2006; Fig. S1).

314 **3.2 Model performance in simulating soil erosion at the catchment outlet**

315 The monthly observed and simulated stream flow, sediment yield, particulate and total N
316 losses at the outlet of the Jieliu catchment from 2007 to 2008 are illustrated in Fig. 3. The
317 observed stream flow and sediment yield began to increase dramatically with the concentrated
318 precipitation in summer and early autumn but rarely occurred in winter and spring (Fig. 3a–c).
319 The upgraded CNMM-DNDC model successfully predicted the above temporal pattern of the
320 stream flow and sediment yield at the catchment outlet with credible NSI values of 0.89 and



321 0.89 and significant ZIRs with R^2 values of 0.94 and 0.92 and slope values of 0.85 and 0.88 for
322 model validation, respectively (Table 1). Moreover, model validation of sediment yield resulted
323 in a larger nRMSE (38.23%) than that of stream flow simulation (34.57%).

324 The observed particulate and total N losses revealed a similar temporal pattern to that of
325 sediment yield (Fig. 3d–e) ranged from 0 to 56.3 kg mon⁻¹ and 0.9 to 283.1 kg N mon⁻¹ with a
326 mean value of 10.5 and 55.9 kg N mon⁻¹, respectively. The corresponding simulated
327 particulate and total N losses resulted in ranges of 0.5 to 50.4 kg N mon⁻¹ and 18.8 to 196.0 kg
328 N mon⁻¹ with averages of 12.0 and 65.1 kg N mon⁻¹, respectively. The upgraded model
329 provided an overestimation of the particulate N loss in August 2007 and September 2008 by
330 11.3 and 14.8 kg N mon⁻¹, respectively. The particulate N losses in February 2007, March 2007,
331 July 2007 and July 2008 and total N loss in summer were underestimated. However, in terms
332 of the validation, statistical comparisons between the simulated particulate and total N losses
333 yielded significant ZIRs with R^2 values of 0.85 and 0.86 and slope values of 0.80 and 0.96,
334 nRMSE values of 57.75% and 42.55%, and NSI values of 0.74 and 0.86, respectively ($n = 12$;
335 Table 1). Meanwhile, the upgraded CNMM-DNDC model successfully predicted the temporal
336 variation and magnitudes of NO₃⁻ loss at the catchment outlet, although the model slightly
337 underestimated the peak loss in July and August of 2007 and in September of 2008 (Fig. S2).
338 As the above results demonstrated, the simulated and observed particulate and total N losses at
339 the catchment outlet indicated good agreement despite the slight underestimation of the
340 individual large values when heavy precipitation occurred.



341 **3.3 Components of the simulated TN and PN at the catchment outlet**

342 The monthly components of TN and/or PN simulated from the original and upgraded
343 CNMM-DNDC model during the model validation of 2008 at the catchment outlet were
344 illustrated in Fig. 4. Among the TN components including PN, NH_4^+ , dissolved organic
345 nitrogen (DON) and NO_3^- , the simulation from both of the original and upgraded
346 CNMM-DNDC demonstrated that the proportion of NO_3^- at the catchment outlet was larger
347 than that of NH_4^+ during the period from May to September when the larger precipitations
348 appeared. Moreover, the upgraded CNMM-DNDC demonstrated that the PN accounted for up
349 to 16.2%–26.6% of the TN components during the period with larger precipitations.
350 Meanwhile, the humads N accounted for 11.3%–20.3% of the PN components, though the
351 humus N accounted for the largest of the PN components. In addition, compared with the
352 original model, the upgraded model simulated the observed TN with smaller nRMSE (42.55%
353 versus 51.67%), better NSI (0.86 versus 0.80) and improved r^2 of the ZIRs (0.98 versus 0.90)
354 though no significant difference was found of the ZIRs between the original and upgraded
355 model (Fig. 4).

356 **3.4 Spatial distributions of sediment yield and particulate C, N, and P losses**

357 Figure 5 illustrated the simulated spatial distributions of the sediment yield and particulate
358 C, N, and P losses and the effects of different land uses on those in the validation year 2008.
359 The annual accumulated sediment yield simulated by the upgraded model amounted to 0–106.6
360 $\text{t ha}^{-1} \text{ yr}^{-1}$ with an average of $5.0 \text{ t ha}^{-1} \text{ yr}^{-1}$ in 2008, which was a moderate rainfall year (952
361 mm) with eight large rainstorm events (exceeding 50 mm rainfall within 24 hours). The
362 simulated annual accumulated particulate C, N, and P losses yielded 0–595.7 $\text{kg C ha}^{-1} \text{ yr}^{-1}$,



363 0–56.0 kg N ha⁻¹ yr⁻¹, and 0–7.9 kg P ha⁻¹ yr⁻¹ with averages of 63.6 kg C ha⁻¹ yr⁻¹, 6.1 kg N
364 ha⁻¹ yr⁻¹ and 0.9 kg P ha⁻¹ yr⁻¹, respectively. The sloping cultivated cropland areas contributed
365 to the greatest losses of sediment and particulate C, N, and P nutrients, with 68%, 60%, 58%
366 and 57% of the total, respectively. Approximately 21% of sediment loss came from the
367 residential areas as the second largest contributor to sediment loss, while the forest areas were
368 the secondary sources to particulate C, N, and P losses, with 30%, 32%, and 32% of total losses,
369 respectively. Meanwhile, the highest rates of the particulate C, N, and P losses per unit area
370 occurred in the sloping cultivated cropland areas, with 84.1 kg C ha⁻¹ yr⁻¹, 7.7 kg N ha⁻¹ yr⁻¹
371 and 1.1 kg P ha⁻¹ yr⁻¹, respectively. However, the residential areas yielded to the highest rates
372 of sediment, i.e., 8.6 t ha⁻¹ yr⁻¹. The second largest loss rates per unit area of the particulate C,
373 N, and P appeared in the residential areas. These results demonstrated that sediment yield and
374 particulate C, N, and P losses caused by surface runoff in the Jieliu catchment were directly
375 relevant to the type of land use, and the sloping cultivated cropland area became the primary
376 source of sediment yield and particulate C, N, and P losses. Meanwhile, sediment and
377 particulate C, N, and P losses from the residential areas could not be neglected.

378 Moreover, the upgraded CNMM-DNDC model coupled the biogeochemical processes
379 with soil erosion, which was able to predict the crucial variables relevant to biogeochemical
380 processes, including the productivity, greenhouse gases, contaminated gases and NO₃⁻ loss and
381 the variables related to soil erosion, including the losses of sediment and particulate C, N and P
382 (Fig. S3, Text S1).



383 **3.5 Sediment yield and particulate C, N, and P losses under different scenarios**

384 The simulated results of the single-factor change scenarios of precipitation and
385 temperature were presented in Figure 6. The sediment yield and particulate C, N, and P losses
386 (i.e., the target variables) increased with precipitation or air temperature which was reflected by
387 the positive values. The more positive the slope value is, the greater the target variables
388 increase and vice versa. The slopes between the air temperature changes of the higher and
389 lower cooling and warming scenarios and the sediment yield changes yielded -1.25 , -1.00 ,
390 -0.38 and -0.40 , respectively. Compared to the slopes of the lower warming scenarios and the
391 lower cooling scenarios, the slopes of the higher warming scenarios and the higher cooling
392 scenarios provided 21% and 5% higher yields of sediment, respectively. Meanwhile, the
393 changes in particulate C, N, and P losses provided similar but stronger responses to the higher
394 cooling scenarios. However, the particulate nutrient losses showed a complicated response to
395 the warming scenarios. For the lower warming scenarios, the particulate nutrient losses
396 increased with air temperature. The changes in the particulate nutrient losses provided an
397 increasing tendency in response to the increase in air temperature. In terms of the higher
398 warming scenarios, the particulate nutrient losses were still increasing, but the rates of increase
399 rate decreased. These results proved that the increase in air temperature decreased the losses of
400 sediment but increased the particulate C, N, and P losses, although the promoting effect
401 became weaker for the higher warming scenarios.

402 The slopes between the precipitation changes of the higher and lower rain-reduced and
403 rain-enhanced scenarios and the sediment yield changes resulted in the values of 0.27, 0.37,



404 0.52 and 0.65, respectively. In comparison with the lower rain-enhanced and rain-reduced
405 scenarios, the slopes of the higher rain-enhanced and rain-reduced scenarios provided 24%
406 higher and 34% lower yields of sediment, respectively. Meanwhile, the changes in particulate
407 nutrient losses provided similar but weaker responses to the changes in precipitation. The
408 above results demonstrated that the losses of sediment and particulate nutrients increased with
409 the increasing precipitation. In addition, the contribution from such an elevation role of
410 precipitation tended to be stronger for the higher rain-enhanced scenarios. Furthermore, the
411 changes in sediment and particulate C, N, and P losses were more sensitive to precipitation
412 scenarios than to temperature scenarios.

413 Table 2 illustrated the results of the multifactor change scenarios and the land use change
414 single-factor scenario (FL scenario). Compared to the baseline scenario, the FL scenario
415 reduced stream flow, sediment yield, and particulate nutrient losses by -12.2% , -3.6% , -5.6% ,
416 -7.0% , and -7.2% , respectively. In comparison with the baseline scenario, the low GHG
417 emissions scenario with air temperature increasing by $1.5\text{ }^{\circ}\text{C}$ and precipitation increasing by 10%
418 increased the stream flow, sediment yield and particulate C, N, and P losses by 21.2% , 4.1% ,
419 5.3% , 5.3% and 5.3% , respectively. The increasing effects of the high GHG emissions
420 scenarios on the sediment and particulate nutrient losses were more than three times those of
421 the low GHG emissions scenarios. The low GHG emissions under the FL scenario increased
422 the stream flow and sediment yield by 5.2% and 0.2% , respectively, but decreased the
423 particulate C, N, and P losses by -0.8% , -2.3% , and -2.5% , respectively. Moreover, the high
424 GHG emissions under the FL scenario increased the stream flow, sediment yield, and



425 particulate C, N, and P losses by 47.9%, 9.2%, 9.3%, 7.8%, and 7.7%, respectively. The
426 short-term and long-term no-tillage scenarios decreased the losses of particulate nutrients by
427 approximately 1% and 20%, respectively, but provided no effect on sediment yield compared
428 with the baseline scenario (data not shown).

429 **3.6 Relationship among the variables relevant to soil erosion, productivity and C/N losses**

430 Figure 7 illustrated the relationships between the variables relevant to soil erosion and
431 biogeochemistry for different land use types. No soil erosion in the RF crop system because of
432 the year-round flooding regime. For the other three land use types, the significant positive
433 correlations ($r > 0.88$) between sediment yield and particulate nutrients were found, because
434 they were entrained by water and moved with water flow. With regard to the SU crop system,
435 the particulate nutrients were significantly correlated with NO_3^- losses through leaching ($r >$
436 0.6), though the correlation coefficient between sediment yields and NO_3^- losses through
437 leaching only yielded to 0.26 (insignificantly). For the SP crop system, the variables related to
438 soil erosion (including sediment yields and particulate nutrients) were negatively correlated
439 with NH_3 emissions ($r > 0.65$), while they were positively correlated with NO_3^- losses through
440 runoff ($r < -0.61$). As to FL, significantly positive correlations between the variables related to
441 soil erosion and NO_3^- losses through leaching/runoff were found ($r > 0.72$).

442 **4. Discussion**

443 **4.1 Effect of land use on soil erosion and particulate C, N, and P losses**

444 Land use change has been considered one of the most important factors affecting the
445 intensity and distribution of surface runoff and soil erosion (Dunjó et al., 2004; Kosmas et al.,
446 1997; Wei et al., 2007; Zhang et al., 2021b). Our study also provided consistent results, which



447 indicated that the intensity of soil erosion and the corresponding particulate C, N, and P losses
448 in the Jieliu catchment were closely related to land use, with the following order: sloping
449 cultivated cropland > residential area > forest land. There were three major reasons why forest
450 land contributed to the lowest losses of sediment and particulate nutrients among the above
451 three land uses. First, canopy interception reduced the amount of rainfall reaching the ground,
452 which directly decreased the occurrence of runoff and subsequent erosion (Greene and Hairsine,
453 2004; Hou et al., 2020; Vasquez-Mendez et al., 2010). Several previous studies also reported
454 that forest land with a thick canopy exhibited a lower amount of runoff than did other land uses
455 (Mehri et al., 2018; Mohammad and Adam, 2010; Nunes et al., 2011). Fortunately, the direct
456 protection mechanism by canopy interception was involved in the CNMM-DNDC model,
457 which was calculated using the leaf area index (Zhang et al., 2018). Second, the litter cover of
458 forest land protects the soil surface from the direct splash and detachment of raindrops, which
459 can decrease the formation of mechanical crusts and increase the infiltration capacity and
460 hence diminish the potential for surface runoff and soil erosion (Casermeiro et al., 2004;
461 Lemenih et al., 2005; Wainwright et al., 2002). However, the CNMM-DNDC did not take the
462 protection of litter cover on the soil surface into consideration. Further observation data and
463 studies are needed to introduce the mechanism of the effect of litter cover on surface runoff and
464 soil sediment into the CNMM-DNDC model. Last, forest land is equipped with higher soil
465 organic matter and hydraulic conductivity than other land uses, which can indirectly enhance
466 soil infiltration and reduce surface runoff (Abrishamkesh et al., 2011; Fu et al., 2000; Lemenih
467 et al., 2004). The excellent soil properties of forest land soil (e.g., higher soil organic matter



468 and vertical saturated hydraulic conductivity) have been involved in the CNMM-DNDC model
469 inputs. Moreover, as the forest litterfall returned to the soil and participated in further C and N
470 cycling, the content of soil organic matter was enhanced and accumulated. With regard to the
471 scenario analysis, we found that the scenarios related to FL contributed to greater decreases in
472 sediment yield than surface runoff (Table 2). The results of the lysimetric plot experiments by
473 Chen et al. (2012) also demonstrated that vegetation types and human interference had a
474 relatively small impact on surface runoff but had an appreciable effect on sediment yield.

475 The canopy of the cultivated cropland served as a weaker hindrance to rainfall than that of
476 the forest canopy, which suffered from more surface runoff. Furthermore, frequent agricultural
477 activities (i.e., tillage) loosen the subsurface soil and nutrients, which raises the risk of soil
478 erosion and the subsequent loss of particulate nutrients (Gregorich et al., 1998; Moldenhauer et
479 al., 1967; Muukkonen et al., 2009). The CNMM-DNDC model has taken the vertical mixing
480 effect of tillage on the upper and lower soils into consideration, and this process left the
481 subsurface soil organic nutrients unprotected and prone to erosion. This explained the
482 reduction in particulate C, N, and P nutrient losses under the no-tillage scenarios (data not
483 shown). However, several studies found that tillage disturbed the soil structure and pore size
484 distribution (Carof et al., 2007; Castellini and Ventrella, 2012; Kay and VandenBygaart, 2002;
485 Nunes et al., 2010), which made the effect of agricultural activities on surface runoff and soil
486 erosion difficult to model (Leitinger et al., 2010). Given that the effect of tillage on soil was not
487 considered in the CNMM-DNDC, the yields of surface runoff and sediment resulting from the
488 no-tillage scenario were not decreased compared with the baseline scenario with tillage.



489 **4.2 Effect of climate change on soil erosion and particulate C, N, and P losses**

490 In past decades, the frequent occurrence of warming and extreme weather events (e.g.,
491 extreme precipitation events) has been irrefutable (IPCC, 2019). From 1998 to 2021, the
492 observed annual average air temperature and annual precipitation in the Jieliu catchment also
493 presented an increasing trend but did not have a significant regression relationship (Fig. S4).
494 The reason of why the simulated soil erosion responded to the air temperature changes might
495 be that the climate-sensitive vegetation growth affected the C_v which was the effect factor of
496 the soil erosion in Eq. 1. The asymmetric response of sediment yield and particulate nutrient
497 losses to the cooling and warming scenarios might result from the different response of the
498 vegetation growth to the cooling and warming of air temperature. Our results of the scenario
499 analysis indicated that the losses of sediment slightly decreased with the scenarios treated with
500 climate warming alone, which lay in the higher C_v caused by the enhanced vegetation growth
501 (Ficklin et al., 2009; Zhang et al., 2020a; Zhou et al., 2003). We found that the decreasing
502 effect of increasing air temperature on sediment loss decreased (especially for the scenario with
503 an air temperature increase of 4 °C, Fig. 6), which might be because the enhanced effect of
504 increasing air temperature on vegetation growth is not unlimited. Once the air temperature
505 exceeds the threshold of the optimum temperature for photosynthesis and vegetation growth, it
506 would have a negative or even harmful impact on plant growth (Chapin, 1983; Schlenker and
507 Roberts, 2009). The complex response of the particulate C, N, and P losses to air temperature
508 increased, probably because they increased with ER and sediment yield, but the ER decreased
509 with sediment. Therefore, the slightly increasing sediment with increasing air temperature and



510 the corresponding decreasing ER might lead to upward or downward fluctuations in particulate
511 C, N, and P losses. However, we found that the rate of soil loss increased with increasing
512 precipitation amount and the corresponding increase in heavy rain events. Jiang et al. (2017)
513 also found that the increase in sediment loss was amplified by the increased precipitation,
514 which was directly accompanied by a dramatic and sustained increase in surface runoff.
515 Therefore, the higher GHG emissions scenarios, in which the soil erosion provided a higher
516 increase response to the rising precipitation and a lower and smaller decrease response to the
517 rising air temperature, might provide a greater risk of soil erosion than the low GHG emissions
518 scenario. Overall, our results indicated that the hydrology of the Jieliu catchment is very
519 sensitive to potential future climate changes, especially to the higher GHG emissions scenarios.

520 **4.3 Interactive effect of climate and land use change on soil and nutrient losses**

521 Changes in either climate or land use imply considerable influences on water and nutrient
522 cycles in a catchment or region (Labat et al., 2004; Milliman et al., 2008; Piao et al., 2007; Yin
523 et al., 2017). Our simulated results indicated that the reduction extent of the scenario with FL
524 on soil erosion, especially on sediment yield and subsequent nutrient losses, offset the
525 increasing extent caused by the low GHG emissions scenario. However, the scenario with FL
526 was insufficient to totally offset the sediment and particulate C, N, and P losses caused by the
527 high GHG emissions scenario. Nevertheless, vegetation restoration might still be able to slow
528 the soaring process of soil erosion caused by climate change in the future. Previous studies
529 primarily focused on the effects of human activity and climate change on the changes in
530 surface runoff or stream flow. Wang et al. (2016) demonstrated that human activity contributed



531 to slightly larger effects on stream flow changes than climate (59% versus 41%) by analyzing
532 the long-term records of hydrological data in the Luan River basin in North China. The results
533 in the Heihe River basin in Northwest China showed that human activities were the dominant
534 contributor to the variation in runoff in the upper and middle reaches when compared to
535 climate change (Qiu et al., 2015). However, other studies have shown that the influence of
536 climate change on soil and water loss was greater than that of human activities. Jiang et al.
537 (2017) pointed out that climate change, in comparison with anthropogenic activities, was the
538 primary factor causing the changes in either stream flow or sediment discharge in the Yellow
539 River basin and Yangtze River basin in China. The Huron River catchment in southeastern
540 Michigan in the U.S.A. was more sensitive to climate change than to land use change, as
541 demonstrated by Barlage et al. (2002).

542 Furthermore, we found that the promoting impacts of both high and low GHG emissions
543 scenarios on surface runoff were greater than those on sediment yield and subsequent
544 particulate nutrient losses. In contrast, the reduction effect of the scenarios with FL on sediment
545 yield and subsequent particulate nutrient losses was stronger than that on surface runoff (Table
546 2). These results demonstrated that human activity, e.g., the conversion from cropland with
547 intensive human disturbance to forest land, resulted in a greater mitigation effect on sediment
548 yield and subsequent particulate nutrient losses than on surface runoff. Therefore, further
549 studies should consider the effects of human activity and climate change on surface runoff and
550 on soil erosion as well as the subsequent nutrient losses. In summary, reasonable human
551 intervention, such as rational land use change, is expected to be a feasible practice to decelerate



552 soil erosion and subsequent particulate nutrient losses without altering and disturbing the
553 hydrological cycle of a catchment in the context of global warming.

554 **Conclusions**

555 The hydro-biogeochemical model (CNMM-DNDC) was improved by introducing the soil
556 erosion physical model (adopted from the simplified ROSE model) and the element (i.e.,
557 carbon, nitrogen and phosphorus) enrichment module to estimate soil erosion and the
558 movements of particulate nutrients. The comparability between the simulation and observation,
559 including surface runoff, sediment yield, and particulate nitrogen and phosphorus losses at the
560 lysimetric plot and the stream flow, sediment yield, and particulate N loss at the outlet of Jieliu
561 catchment, demonstrated that the upgraded CNMM-DNDC model could reliably simulate soil
562 erosion and the consequential particulate nutrient losses. The spatial distribution characteristics
563 of sediment yield and the consequential particulate carbon, nitrogen and phosphorus losses
564 were directly related to the spatial distribution of land use type, among which the sloping
565 cultivated cropland areas contributed to the greatest losses. The analysis of climate
566 single-factor change scenarios implied that the high GHG emissions scenarios provided a
567 greater potential risk of soil erosion, which resulted in the larger soil erosion rates than those in
568 the low GHG emissions scenarios. The scenarios with all non-forest land changes into forest
569 land decreased stream flow, sediment yield and particulate C, N, and P losses compared to the
570 baseline scenario. Anthropogenic activities (e.g., land use change) might be expected to help
571 mitigate the processes of soil and water losses accelerated by climate change in the future.

572 **Code and data availability**

573 The CNMM-DNDC model was originally developed by the Institute of Atmospheric



574 Physics using C++ language, which can be run on a standard PC. The upgraded model is
575 available on the FigShare (<https://doi.org/10.6084/m9.figshare.20210546>).

576 **Author contribution**

577 Siqi Li arranged data, improved model and implemented the simulation, prepared the
578 original draft. Yong Li, Xunhua Zheng, Wei Zhang developed the conceptualization and
579 methodology of this study. Bo Zhu, Pengcheng Hu, Jihui Fan, Tao Wang collected and arranged
580 data. Shenghui Han, Rui Wang, Kai Wang analyzed data and verified the results. Zhisheng Yao,
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590 **Competing interests**

591 The authors declare that they have no conflict of interest.

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862



863 Table 1 Performance of the revised CNMM-DNDC model in simulating the stream flow,
 864 sediment, and particulate and total nitrogen (N) losses at the Jieliu catchment outlet from 2007 to
 865 2008. Total N refers to the total amount of NH_4^+ , NO_3^- , dissolved organic N and particulate N.

Variables	Operation	Size	nRMSE	NSI	ZIR		
					Slope	R^2	p
Stream flow	Calibration	12	18.29	0.98	0.95	0.98	< 0.001
	Validation	12	34.57	0.89	0.85	0.94	< 0.001
Sediment loss	Calibration	12	34.02	0.94	0.91	0.96	< 0.001
	Validation	12	38.23	0.89	0.88	0.92	< 0.05
Particulate N loss	Calibration	12	49.45	0.87	0.93	0.88	< 0.001
	Validation	12	57.75	0.74	0.80	0.85	< 0.001
Total N loss	Calibration	12	56.98	0.86	1.21	0.90	< 0.001
	Validation	12	42.55	0.86	0.96	0.86	< 0.001

866 The statistical criteria used to quantify the discrepancy between observations and simulations
 867 include the normalized root mean square error (nRMSE), the Nash–Sutcliffe index (NSI) and the
 868 slope, determination coefficient (R^2) and significance level (p) of the zero-intercept univariate
 869 linear regression (ZIR). Size represents the sample size.

870

871



872 Table 2 Simulated comprehensive effects of precipitation, air temperature and land use change
873 on surface runoff, sediment yield, and particulate carbon (C), nitrogen (N) and phosphorus (P)
874 losses in the validation year of 2008. The low greenhouse gas (GHG) emission scenario
875 represents the scenario of air temperature increasing by 1.5 °C and precipitation increasing by
876 10%. The high GHG emission scenario represents the scenario of an air temperature increase of
877 4 °C and a precipitation increase of 30%. The FL scenario is the abbreviation of the scenario of
878 upland change into forest land.

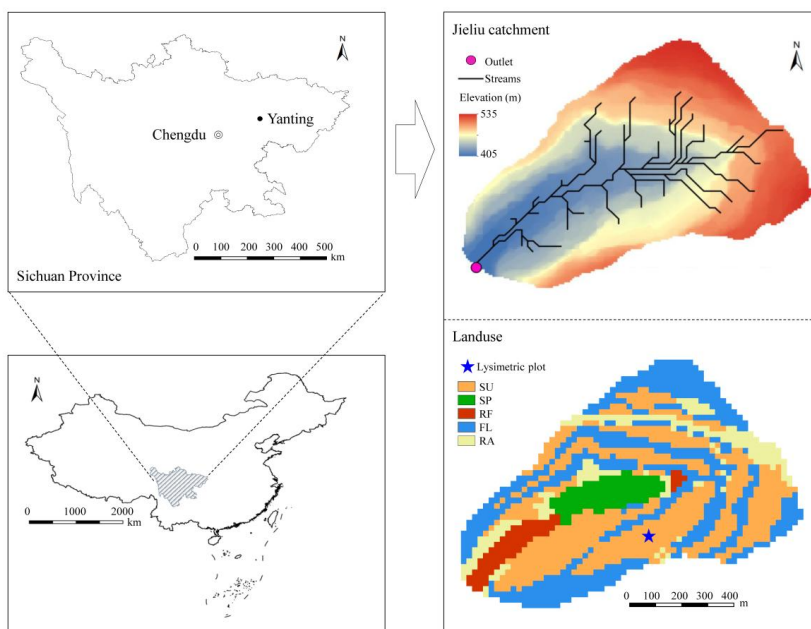
Scenario	Change between the scenario and the baseline (%)				
	Surface runoff	Sediment yield	Particulate C	Particulate N	Particulate P
Low GHG	21.2	4.1	5.3	5.3	5.3
High GHG	72.9	14.8	17.8	18.0	18.1
FL	-12.2	-3.6	-5.6	-7.0	-7.2
Low GHG with FL	5.2	0.2	-0.8	-2.3	-2.5
High GHG with FL	47.9	9.2	9.3	7.8	7.7

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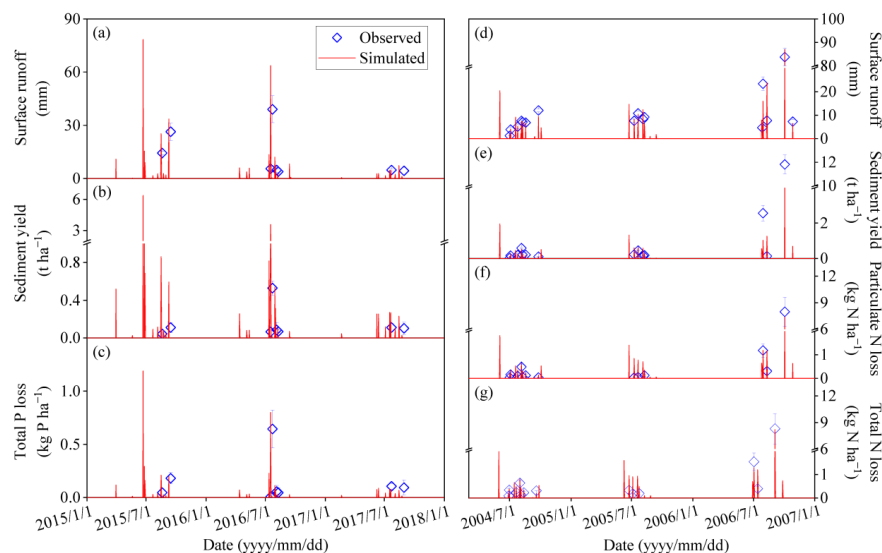
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883 Fig. 1 The location, digital elevation model and land use types of the Jieliu catchment. The land
884 use types are the sloping uplands with the summer maize–winter wheat rotation (SU), seasonally
885 waterlogged paddy with the paddy rice–winter wheat rotation or paddy rice–rape rotation (SP),
886 the winter-flooding paddy with the paddy rice-flooding fallow regime (RF) and the forest land
887 (FL).

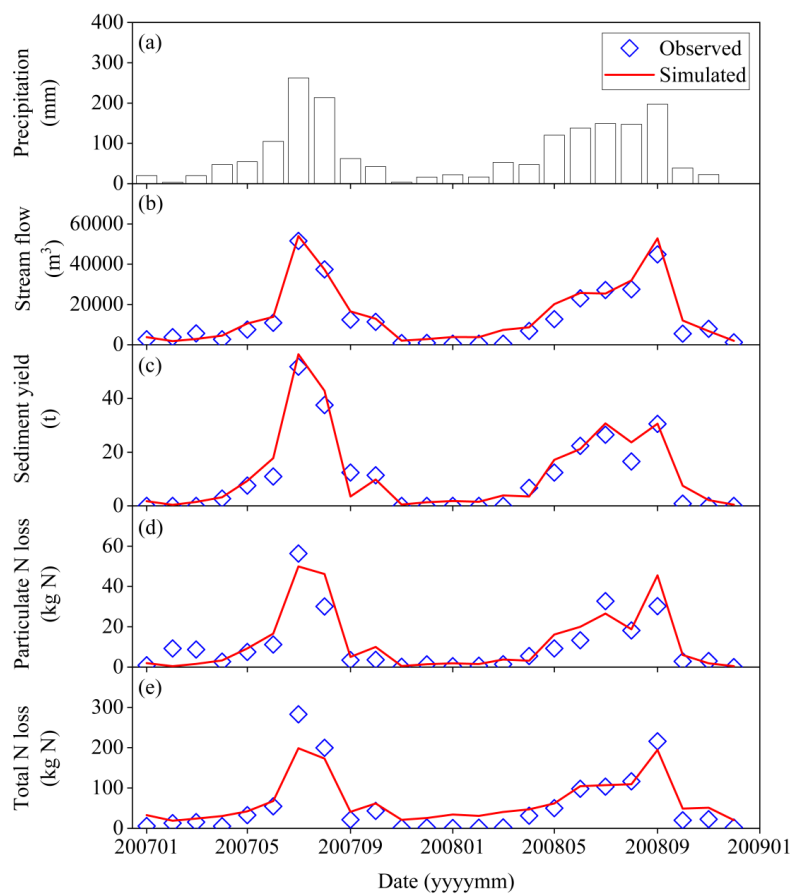
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890 Fig. 2 Observed and simulated surface runoff (a), sediment yield (b) and total phosphorus (P)
891 losses (c) from 2015 to 2017 and surface runoff (d), sediment yield (e), particulate nitrogen (N)
892 loss (f) and total N loss (g) from 2015 to 2017 in the lysimetric plot. Total P refers to the
893 dissolved and particulate P. Total N refers to the total amount of NH_4^+ , NO_3^- , dissolved organic
894 N and particulate N. The vertical bars indicate the standard error of three spatial replicates. The
895 observed data cited from Deng et al. (2011), Zhang et al. (2018), Li et al. (2022) and Hu (2020)
896 were provided by Bo Zhu.

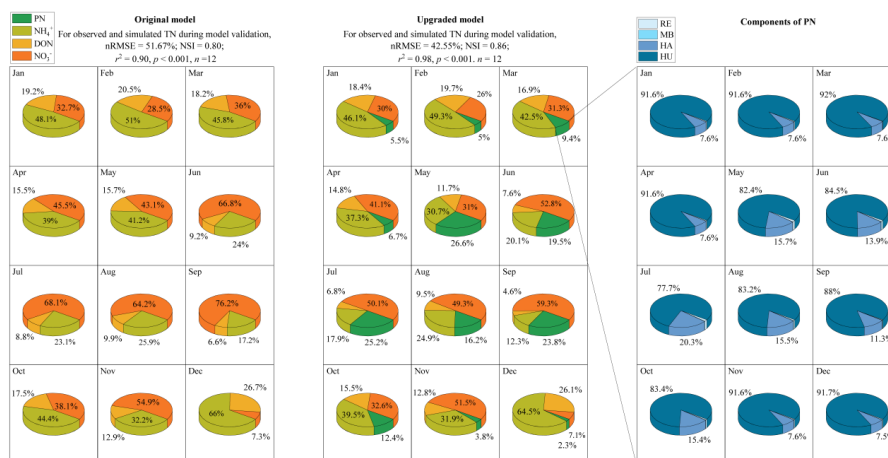
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899 Fig. 3 Monthly observed precipitation (a), observed and simulated stream flow (b), sediment
900 yield (c), particulate nitrogen (N) loss (d) and total N loss (e) at the outlet of the Jieliu catchment
901 from 2007 to 2008. Total N refers to the total amount of NH_4^+ , NO_3^- , dissolved organic N and
902 particulate N. The observed data cited from Deng et al. (2011) and Zhang et al. (2018) were
903 provided by Bo Zhu.

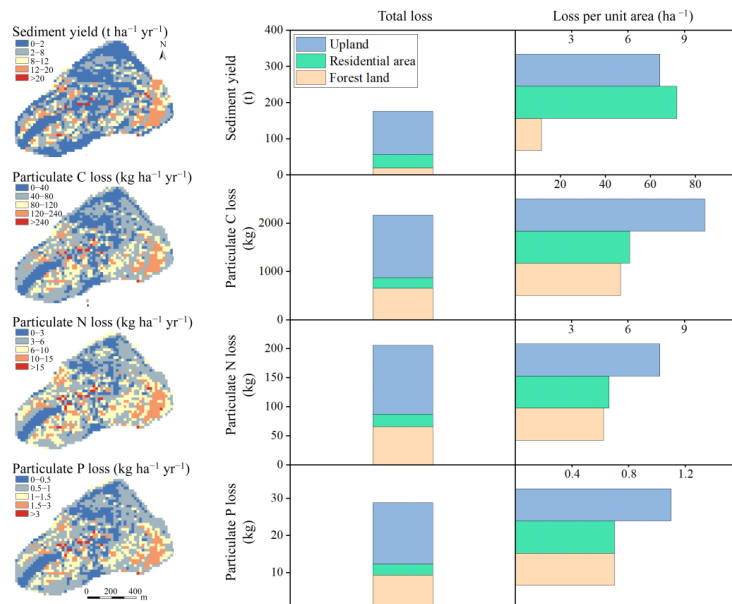
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906 Fig. 4 Components of the simulated total nitrogen (TN) and/or particulate N (PN) of the
 907 original and upgraded CNMM-DNDC model during the model validation. DON is the
 908 abbreviation of the dissolved organic nitrogen. The components of PN are the N from residue
 909 (RE), microbe (MB), humads (HA) and humus (HU).

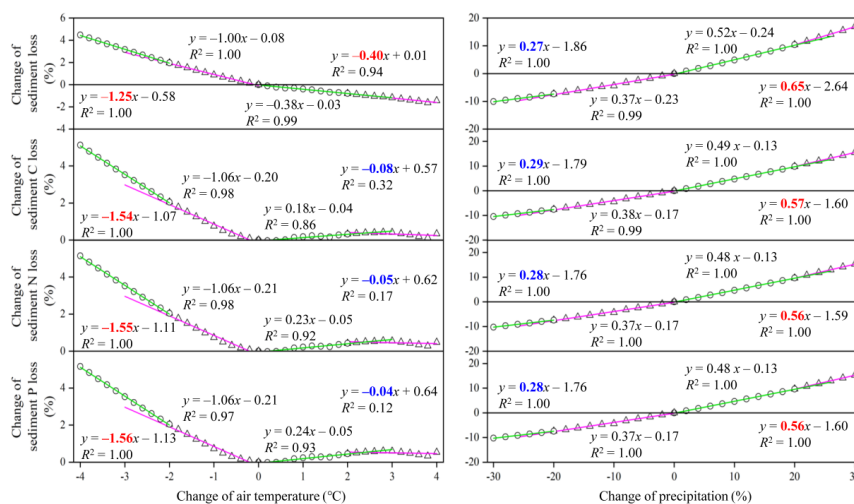
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911

912 Fig. 5 Simulated spatial distributions of sediment yield, particulate carbon (C), nitrogen (N) and
 913 phosphorus (P) losses and the effects of different land uses (i.e., cropland, residential area and
 914 forest land) in the validation year of 2008.

915



916

917 Fig. 6 Simulated effects of precipitation and air temperature change on sediment yield and

918 particulate carbon (C), nitrogen (N) and phosphorus (P) losses in the validation year of 2008.

919 The air temperature and precipitation single-factor scenarios were divided into four sets. The

920 scenarios with air temperature reductions and increases 0°C~2°C and greater than 2°C were

921 defined as the lower and higher cooling and warming scenarios, respectively. Similarly, the

922 scenarios with precipitation reductions and increases 0%~20% and greater than 20% were

923 defined as the lower and higher rain-reduced and rain-enhanced scenarios, respectively. The

924 numbers in blue and red in front of the letter x represent that the higher warming or cooling

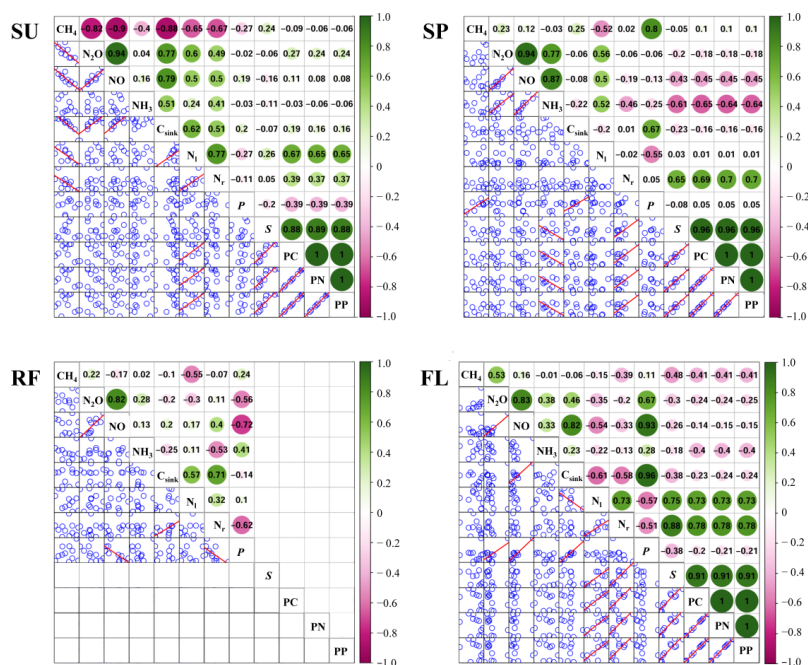
925 scenarios (or the higher rain-enhanced or rain-reduced scenarios) result in more and lower

926 effects on sediment yield and particulate C, N and P losses than the lower ones, respectively.

927



928



929

930 Fig. 7 Correlation analysis among the simulated sediment (S), particulate carbon (PC), nitrogen
 931 (PN) and phosphorus (PP) losses, productivity (P), C sink density (C_{sink}), methane (CH_4), nitrous
 932 oxide (N_2O), nitric oxide (NO) and ammonia (NH_3) emissions, losses of nitrate through leaching
 933 (N_i) and surface runoff (N_r) for different land use types. The land use types are the sloping
 934 uplands with the summer maize–winter wheat rotation (SU), seasonally waterlogged paddy with
 935 the paddy rice–winter wheat rotation or paddy rice–rape rotation (SP), the winter-flooding
 936 paddy with the paddy rice–flooding fallow regime (RF) and the forest land (FL). No losses of S,
 937 PC, PN, and PP in the RF crop system because of the year-round flooding regime. The figures in
 938 the circles stand for the correlation coefficients. The correlations with the level of $p < 0.05$ are
 939 considered as significant and the linear regression lines are exhibited.