

1 **Soil priming effects and involved microbial community along salt gradients**

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15 5 Figures

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18 27 text pages

19 **Abstract**

20 Soil salinity mediates microorganisms and soil process, like soil organic carbon (SOC) cycling.
21 Yet, how soil salinity affects SOC mineralization via shaping bacterial communities diversity and
22 composition remains elusive. Therefore, soils were sampled along a salt gradient (salinity at 0.25%,
23 0.58%, 0.75%, 1.00% and 2.64%) and incubated for 90 days to investigate i) SOC mineralization (i.e.,
24 soil priming effects induced by cottonseed meal, as substrate) and ii) responsible bacteria community,
25 by using high throughput sequencing and natural abundance ¹³C isotopes (to partition cottonseed meal
26 derived CO₂ and soil derived CO₂). We observed negative priming effect during first 28 days of
27 incubation ~~but and~~ turned to positive priming effect after day 56. Negative priming at the early stage
28 might be due to the preferential utilization of cottonseed meal. The followed positive priming
29 decreased with the increase of salinity, which might be caused by the decreased alpha diversity of
30 microbial community in soil with high salinity. Specifically, soil pH and EC along salinity gradient
31 were the dominant variables modulating the structure of microbial community and consequently SOC
32 priming (estimated by distance-based multivariate analysis and path analysis). By adopting O2PLS,
33 priming effects were linked with specific microbial taxa, e.g., *Proteobacteria (Luteimonas, Hoeflea*
34 *and Stenotrophomonas)* were the core microbial genus that attributed to the substrate induced priming
35 effects. Here, we highlight that the increase of salinity reduced the diversity of microbial community
36 and shifted dominant microorganisms(*Actinobacteria* and *Proteobacteria (Luteimonas, Hoeflea* and
37 *Stenotrophomonas)*) that determined SOC priming effects, which provides a theoretical basis for
38 understanding of SOC dynamics and microbial drivers under salinity gradient.

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40 **Keywords:** *Salt gradient, priming effects, bacterial community, core microorganisms*

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41 1. Introduction

42 Soil salinization is an increasing environmental problem caused by natural and
43 human activities in the arid and semi-arid area (Wichern et al., 2006). Salinization is
44 often a major threat to crop productivity in agricultural land. Soil microorganisms suffer
45 from osmotic stress. Soil salinity often cause microbial death or dormant. It was widely
46 reported that the increased salinity decrease microbial biomass, enzymatic activity, and
47 alpha diversity of microbial community (Laura, 1974; Pathak ~~et al. and Rao,~~ 1998;
48 Rietz ~~et al. and Haynes,~~ 2003). Soil salinity is reported to the major determinants of
49 composition, activity of microbial community (Kamble et al., 2014). Although salinity
50 is reported to be a vital factor in influencing microorganisms in the arid and semi-arid
51 area, limited studies investigated C processes (e.g., priming effect) driven by microbial
52 community in salinity soils (Sardinha et al., 2003).

53 Soil organic carbon (SOC) is the largest pool (1500 Pg C) in the terrestrial carbon
54 (C) cycle, and contains twice as much C as the atmosphere (Filley ~~et al., and Boutton,~~
55 2006; Wiesmeier et al., 2019). The input of substrate C can influence the output (i.e.,
56 CO₂ release) through a phenomenon called priming effect, which was firstly discovered
57 by LÖhnis (1926). Substrate additions accelerate or decrease soil organic C
58 mineralization, referred to positive or negative priming effects (Kuzyakov et al., 2000).
59 The intensity of the priming effect affects the turnover of SOC and thus storage pool
60 (Sullivan ~~et al., and Hart,~~ 2013). Soil priming effects are affected by many biotic and
61 abiotic factors (Lavelle, ~~et al., 1997;~~ Martin ~~et al., W,~~ 2019), to investigate abiotic and
62 biotic mechanisms underlying SOC priming enhance strong understanding of the SOC
63 cycling.

64 Soil priming effects is affected by soil fauna animals (Scheu ~~et al., and Parkinson,~~
65 1994), activities, diversity and composition of microbial community (~~Di Leonardo et al.,~~
66 ~~2017;~~ Fontaine et al., 2011). The microbial decomposers are the major player in the
67 decomposition process of added C sources. The addition of substrate, such as composts
68 (Xun et al., 2016), animal sludges (Hartmann et al., 2015), sewage sludges (Su et al.,

69 2017; Wagner ~~et al.,~~ ~~and Raquel,~~ 2011) and plant residues (Dai et al., 2017), generally
70 increases soil microbial biomass C and stimulates the microbial activities thus enhanced
71 the loss of SOC (positive priming effects) (Fontaine et al., 2003; Bird et al., 2011; Li et
72 al., 2018; Ali et al., 2019).

73 Concerning abiotic factors, the priming effect can be controlled by climate
74 variables (Hagemann, 2008), and soil properties, like pH, EC, TN, etc (Blagodatskaya
75 ~~et al., and Kuzyakov,~~ 2008; Luo et al., 2017). To understand how environmental and
76 edaphic factors affect the processes of SOC mineralization, is important to estimate
77 terrestrial C pool (Lehmann ~~et al.,~~ ~~and Kleber,~~ 2015). Although many studies have
78 tested the effects of soil pH, SOC content, and other edaphic variables on soil priming
79 effect, few study investigated soil priming effects in salinity soil (Asghar et al., 2012),
80 especially linked with soil microbial community structure and their functions in C
81 decomposition (Soina et al., 2018).

82 Thus, we sampled the soils along natural salinity gradients (0.25%, 0.58%, 0.75%,
83 1.00%, 2.64% apart from total water-soluble salt). Based on these soils, we conducted
84 a 90 days of indoor incubation applying C3 substrate of cottonseed meal ($\delta^{13}\text{C}=-$
85 23.47‰) to C4 soils with salt gradient ($\delta^{13}\text{C}$ between -14.21‰ and -16.01‰), to
86 investigate: 1) mineralization rate of cottonseed meal and induced soil priming effects
87 along salt gradients; 2) diversity of microbial community in the soils with increased
88 salinity, and 3) identify the bacteria taxa associated Soil priming. We hypothesized that
89 i) soil microbial community diversity and composition will be different with the
90 different in soil variables particularly pH and EC along salinity gradients, and ii) Soil
91 C processes like priming effects will be regulated mainly by microbial community and
92 especially the core microbial species. To clarify the priming effects and involved
93 microbial groups would help us better understanding C sequestration potential and
94 underlying mechanisms in saline soils.

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97 **2. Materials and methods**

98 **2.1. Soil sampling and cottonseed meal production.**

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99 The soil type was gray desert soil, which was collected from farmlands (82.90° E,
100 44.96° N) in Xiao Yinpan town(82.90° longitude, 44.96° latitude) , Bole City, Bortala,
101 northern Xinjiang Uygur Autonomous Region, northwest China. The farmlands soil is
102 naturally formed original saline-salinity soil and with a continuous 30 years planting of
103 maize (C4 crop) and maize straw returning to soil for 7-8 year. In September 2021, we
104 determining the sampling area, and use the five-point sampling method to collecting
105 non-rhizosphere soil. –The soil samples were indoor air drying and hand-picked to
106 remove visible other debris, animal and plant residues and then sieved at field moisture
107 (<2mm) and subsequently adjusted to 40% of water holding capacity (WHC).
108 Determination of five salinity gradients at 0.25%, 0.58%, 0.75%, 1.00% and 2.64%
109 through soil salinity measurements. Texture was determined by the pipette method
110 without carbonate in all soil samples. They were then incubated at 25 °C for 7 days
111 before starting the experiments, to allow any early sampling and sieving effects to
112 subside.

113 Cottonseed meal is a kind of reddish or yellow granular material obtained by
114 pressing, leaching and other cottonseed. The cottonseed meal was purchased from the
115 market and dried at 105 °C for 24 h indoor, then further pulverized by a ball mill and
116 passed through < 2 mm sieve.

117

118 **2.2. Soil and substrate analyses**

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119 EC and pH of soil and cottonseed meal were measured at a soil: water ratio of 1:5
120 (weight/weight) (Bao, 2000). Air-dry soil (5 g, <2 mm) and 25 ml of deionised water
121 were shaken together for 1 min and left to settle for 30 min, which was repeated once
122 more before pH was determined with a pH electrode. Soil water-soluble salt was
123 analyzed by weighted at a soil:water ratio of 1:5 (weight/weight). Air-dry soil (5 g, <1
124 mm) and 25 ml of deionised water were shaken together for 30 min, filtration to obtain

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125 clear filtrate, using thermostat water bath to evaporate and weigh (Bao, 2000). Soil total
126 carbon (TC), total nitrogen (TN) are collect soil to be tested was dried and ground
127 through a 0.15mm screen, and a certain amount of treated soil sample was wrapped in
128 tin foil and placed in an element analyzer for determinatio (air-dried, milled <150 μm)
129 were determined by dry combustion ~~Total soil C and N concentrations (air-dried, milled~~
130 ~~<150 μm) were determined by dry combustion~~ (LECO CNS 2000, LECO Corporation,
131 Michigan, USA). Soil microbial biomass C was determined by fumigation extraction
132 (Vance et al., 1987; Wu et al., 1990). The K₂SO₄ extractable organic C was determined
133 using an organic carbon autoanalyser (Shimadzu, Analytical Sciences, Kyoto, Japan).
134 Soil microbial biomass C (Bc) was calculated from: Bc = 2.22 Ec, where Ec = [(organic
135 C extracted from fumigated soil) minus (organic C extracted from non-fumigated soil)].
136 The natural δ¹³C (‰) abundance of the soils (air-dried, milled <200 μm) was
137 determined using an elemental analyser-isotope ratio mass spectrometer (Sercon Ltd,
138 Crewe, UK). All measurements are given on an oven-dry weight basis (o.d., 105 °C, 24
139 h).

140 The δ¹³C (‰) abundance of the cottonseed meal (air-dried, milled <200 μm) was
141 determined using an elemental analyser-isotope ratio mass spectrometer (Sercon Ltd,
142 Crewe, UK). The main elemental composition of the substrate was determined using
143 elemental analysis (Vario EL Cube, Hanau, Germany), with the samples combusted at
144 1200 °C. Natural δ¹³C (‰) abundance ,the total carbon, total nitrogen contents and C/N
145 of the cottonseed meal was presented in Table 1.

147 2.3. Experimental design

148 After pre-incubation, five soils with salinity gradient (salinity at 0.25%, 0.58%,
149 0.75%, 1.00% and 2.64%) were thoroughly mixed with cottonseed meal at 20 mg C
150 g⁻¹ soil (d.w. basis), and incubated over 90 days following moisture adjustment to 40%
151 of water-holding capacity (WHC) to investigate the substrate mineralization and

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152 priming effects. Each soil sample (40 g d.w. basis) was incubated in a 100 ml beaker
153 inside a 1 L brown glass jar. Three jars with only water and NaOH were set as blank.
154 All the jars were sealed with a rubber bung and incubated in a randomized block design
155 at 25 °C for the 90 days of incubation. The NaOH vials were changed after 1, 3, 5, 7,
156 14, 28, 56 and 90 days for determination of evolved CO₂ and ¹³C–CO₂ (‰). Meanwhile,
157 soil biomass C, NH₄⁺, NO₃⁻, pH, EC, TC, TN and DNA extraction were measured at
158 day 28.

159

160 2.4. Soil CO₂-C and its isotopic composition

161 Soil C evolved as CO₂-C in jars was measured by trapping CO₂ in 1 M NaOH
162 (20 ml) during soil incubation. After the NaOH (20 ml) trapping CO₂ at different
163 periods of soil incubation, 5 ml 1 M NaOH of each sample was mixed with 10 ml
164 deionised water and titrated with 0.05 M standardised HCl by the TIM840 autotitrator
165 (Radiometer Analytical, Villeurbanne Cedex, France). Meanwhile, the δ¹³C (‰) of
166 trapped CO₂-C was precipitated, with 8 ml of the 1 M NaOH (20 ml) mixed with 8 ml
167 1.5 M BaCl₂ in vials (Aoyama et al., 2000). The BaCO₃ precipitate was trapped on the
168 glass fibre the filter, rinsed with deionised water several times, and dried overnight
169 (80 °C), weighed (0.100-0.200 mg) into tin capsules, and analyzed for δ¹³C on an
170 elemental analyzer-isotope ratio mass spectrometer (Sercon Ltd, Crewe, UK).

171

172 2.5. DNA exaction and sequencing

173 The total soil DNA was extracted from 0.50 g of moist soil using a FastDNA Spin
174 Kit (MP Biomedicals, Santa Ana, CA, USA) according to the manufacturer's protocol.
175 The extracted DNA was dissolved in 50 µl of TE buffer, quantified using a
176 spectrophotometer and stored at -20 °C until sequencing.

177 V3-V4 hypervariable regions of the bacterial 16S rRNA gene were amplified with
178 primers 341F (5'-CCTAYGGRBGCASCAG-3') and 806R(5'-
179 GGACTACHVGGGTWTCTAAT-3'). The PCR reactions were conducted with a

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180 thermocycler PCR system (GeneAmp 9700, ABI, USA) by using the following
181 programs: 3 min of denaturation at 95 °C; followed by 27 cycles of 30 s at 95 °C, 30 s
182 at 55 °C, and 45 s at 72 °C; and a final extension at 72 °C for 10 min with a thermocycler
183 PCR system (GeneAmp9700, ABI, USA). PCR amplicons pooled from the triplicate
184 reactions were purified using a QIAquick PCR purification kit (Qiagen, Shenzhen,
185 China), and quantified using a NanoDrop ND-1000 spectrophotometer (Thermo
186 Scientific, Waltham, MA, USA). The PCR products were purified, mixed, and sent to
187 Majorbio, Inc. (Shanghai, China) for sequencing based on the Illumina MiSeq platform.

188

189 2.6. Calculations

190 2.6.1. CO₂-δ¹³C emission

191 The mineralisation of cottonseed meal was separated from SOC mineralisation
192 according to the change of stable isotopic composition ($\delta^{13}\text{C}$) with time. The standard
193 equation for determining $\delta^{13}\text{C}$ (‰) is derived from:

$$194 \delta^{13}\text{C} (\text{‰}) = [(R_{\text{sample}}/R_{\text{VPDB}}) - 1] \times 1000, \quad \text{Eqn. 1}$$

195 where R_{sample} is the mass ratio of ¹³C to ¹²C of each sample and R_{VPDB} is the
196 international PDB (Peedee Belemnite) limestone standard. The labeled ¹³C (‰) of
197 cottonseed meal was then estimated from:

$$198 \text{CO}_2\text{-}^{13}\text{C} (\text{‰}) = (\delta_{\text{treatment}} - \delta\text{C4}) / (\delta\text{C3} - \delta\text{C4}), \quad \text{Eqn. 2}$$

199 where $\text{CO}_2\text{-}^{13}\text{C}$ (‰) is the proportion of evolved CO₂ from C3 (cottonseed meal)
200 matter, $\delta_{\text{treatment}}$ is the $\delta^{13}\text{C}$ (‰) in treatments of soil with cottonseed meal, δC4 is the
201 $\delta^{13}\text{C}$ (‰) in control soil and δC3 is the $\delta^{13}\text{C}$ (‰) from cottonseed meal. Thus, the CO₂-
202 C produced from cottonseed meal during the incubation was calculated from:

$$203 \text{CO}_2\text{-}^{13}\text{C} (\mu\text{g g}^{-1} \text{soil}) = \text{CO}_2\text{-}^{13}\text{C} (\text{‰}) \times \text{total CO}_2\text{-C} (\mu\text{g g}^{-1} \text{soil})/100, \quad \text{Eqn. 3}$$

204 CO₂ from SOC was CO₂-¹³C subtracted from total evolved CO₂-C. The absolute
205 soil priming effect (or primed soil CO₂-C) with the addition of cottonseed meal was
206 calculated from:

$$207 \text{Primed soil CO}_2\text{-C} (\mu\text{g C g}^{-1} \text{soil}) = \text{CO}_2\text{-C}_{\text{treatment}} - \text{CO}_2\text{-C}_{\text{control}} \quad \text{Eqn. 4}$$

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208 where $\text{CO}_2\text{-C}_{\text{treatment}}$ is the non-isotopically labeled $\text{CO}_2\text{-C}$ evolved from
209 cottonseed meal amended soil, $\text{CO}_2\text{-C}_{\text{control}}$ is non-isotopically labeled $\text{CO}_2\text{-C}$ evolved
210 from soil without cottonseed meal.

211

212 2.7. Statistics

213 The data of ^{16}S gene sequencing were processed using the Quantitative Insights
214 Into Microbial Ecology (QIIME) 1.9.0-dev pipeline (Caporaso et al., 2010). In brief,
215 Reads with less than length 200 bp and ambiguous bases were discarded. The sequences
216 were then binned into operational taxonomic units (OTUs) by UCLUST (Edgar, 2010)
217 based on 97% pairwise identity. Chimeric OTUs identified by USEARCH (Edgar et al.,
218 2011) in QIIME were removed. The most abundant sequence from each OTU was
219 selected to represent that OTU. Taxonomy was assigned to 16S OTUs against a subset
220 of the Silva 104 database. The representative OTU sequences were aligned using
221 PyNAST (Caporaso et al., 2010). We obtained between 64,425 and 89,989 clean_reads
222 per sample for all experimental samples.

223 To avoid potential bias caused by sequencing depth, all sample datasets were
224 rarefied for the bacteria α -diversity and β -diversity analyses. Faith's phylogenetic
225 diversity was calculated to provide an integrated index of the phylogenetic breadth
226 across taxonomic levels (Faith, 1992). To compare β -diversity between samples,
227 principal coordinate analyses based on the unweighted and weighted UniFrac
228 (Lozupone et al., 2007a) distances were calculated using the function 'pcoa' in the R
229 package 'Ape'. Additionally, permutational multivariate analysis of variance
230 (PERMANOVA) was carried out using the function 'adonis' in the R 'vegan' to
231 measure effect size and significance on β -diversity. The variable influence projection
232 (VIP) value was processed using the way of O2PLS analysis by the SIMCAP 14
233 (Version 14.1.0.2047) (Wang et al., 2016). The y-matrix was defined as the
234 environmental factors datasets and the x-matrix was defined as the microbial
235 community on genus level dataset.

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236 Data were logarithmically transformed and analyzed by ANOVA. All analyses
237 were performed using SPSS software (13th edition). Pearson's correlation analyses were
238 performed to assess the linear correlation among soil physio-chemical properties and
239 microbial community. MULTIVARIATE analysis were operated to investigate
240 interaction of salinity treatments on bacteria community parameters.

241

242 3. Results

243 3.1. Soil physicochemical properties along salt gradients

244 The major soil physicochemical properties along salt gradients were presented
245 (Table 1) and all of soil physicochemical properties has significant difference ($P <$
246 0.05). The total soluble salinity content in the soils ranged from 0.25% to 2.64% of
247 salinity soils, soil salt gradients increasing gradually from salinity 1 samples to salinity
248 5 samples. The pH and EC in soils ranged from 8.45 to 8.85 and from 1.06 ms cm⁻¹ to
249 7.75 ms cm⁻¹. Soil total C and N were increased with salinity, ranging from 3.16% to
250 3.57%, and from 0.18% to 0.26%. The $\delta^{13}\text{C}$ value for soils are between -14.21‰ and -
251 16.01‰, which were relatively enriched compared to cottonseed meal (-23.47‰). This
252 allowed separation of soil derived CO₂ from total evolved CO₂, according to the classic
253 mixed modeling.

254

255 3.2. Total CO₂ evolution

256 During the whole 90 days of incubation, the cumulative CO₂ evolved had
257 similar trends, which the amount of CO₂ increased with the incubation times (Fig. S1).
258 The cumulative CO₂ evolved increased more rapidly with the addition of cottonseed
259 meal before 14 days, compared to non-amended soils. At 90 days of incubation. The
260 cumulative CO₂ evolved in the soil with the lowest salinity (Salinity 1) gave the lowest
261 CO₂ emission (597 $\mu\text{g C g}^{-1}$) in the non-amended soils (Fig. S1, $P < 0.001$).

262

263 3.3. Cottonseed derived ¹³CO₂ and soil priming effects

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264 The total cumulative CO₂-C was divided three parts based the δ¹³C value,
265 including basal soil-derived CO₂, cottonseed meal-derived CO₂ and primed soil CO₂
266 (Fig.1). The cottonseed meal-derived $\epsilon_{\text{CO}_2\text{-CO}_2}$ had a significant contribution to the total
267 CO₂ evolved during the early incubation period. The cottonseed meal-derived CO₂ was
268 significantly higher in Salinity 1, Salinity 2 and Salinity 3 than in Salinity 4 and Salinity
269 5 before 28 days incubation. Meanwhile, the soil priming effects was negative in all
270 amended soil treatments before 28 days incubation and the direction of priming effect
271 in most of soil samples turned into positive after 28 days. During the whole 90 days
272 incubation, there was a negative correlation between cottonseed meal-derived CO₂ and
273 primed soil CO₂ (Fig. 2).

274

275 3.4. Bacterial diversity and community structure

276 The number of sequences ranged from 64,425 to 91,261 for per sample (average
277 value of 80,602). About 27,990 OTUs in total were obtained under different five
278 treatments. Bacterial community diversity was measured by a series of OTU-based
279 analyses of alpha diversity including chao1 estimator, and observed_species in the
280 QIIME pipeline (Fig. 3). Chao1 diversity estimator and observed_species was
281 significantly different in treatments, being the highest in Salinity 1, followed by Salinity
282 3, Salinity 2, Salinity 4 and Salinity 5 ($P < 0.01$). In general, bacterial community
283 diversity decreased with increasing salinity (Fig. 3).

284 The most abundant phylum in the soils and their correlation with salinity were
285 shown in Fig. 4. Among them, *Actinobacteria* was the dominant taxa in all soils, with
286 the abundance ranging from 50.07 % (Salinity 3) to 68.99 % (Salinity 4). The relative
287 abundance of *Bacteroidetes*, *Firmicutes*, and *Deinococcus-Thermus* increased with
288 the salinity, while *Acidobacteria* decreased with salinity degree.

289 Based on OTUs of five gradient salt treatments, the PCA analysis showed that
290 treatments from Salinity 2 and Salinity 4 clustered together. Meanwhile, soil samples

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291 of Salinity 1, Salinity 3 and Salinity 5 distributed in the first, fourth and three quadrant,
292 which indicated that these treatments had large environmental heterogeneity (Fig. S4).

293 In order to visualize the relationship between environmental factors and microbial
294 community, *Canonical Correspondence Analysis* (CCA) was conducted, showing that
295 NO₃⁻-N, EC and TC had a more obvious impact than other factors for microbial
296 community (Fig. 3). Soil EC were positively correlated with pH, NH₄⁺-N, and
297 negatively correlated with TN, TC and MBC. Mantel test and Distance-based
298 multivariate analysis showed the contribution rate of different environmental factors
299 account for 78% of the variability of microbial communities (Table 2). The value of pH
300 (31%) and EC (12%) had a strong influence on microbial community.

301

302 3.5. Relation between soil microbial community and C dynamics

303 Based on the O2PLS analysis, the variable influence projection (VIP) values of
304 bacterial genus more than 1.00% were showed their contributions to C decomposition
305 of cottonseed meal-derived C, basal soil-derived C, and primed soil C (Table 3). There
306 were many microbial taxa positively correlating to soil primed CO₂, for insatnce, genera
307 of *Actinomarinales*, *Luteimonas*, *Nocardioides*, *Hoeflea*, *Intrasporangium*,
308 *Nitrolancea*, *Pseudarthrobacter* and *Stenotrophomonas* had a positive correlation with
309 primed CO₂. In order to further to evaluate the relationship between soil properties, soil
310 bacterial communities and C decomposition, we used the structural equation modeling
311 (SEM) to suggest the direct and indirect impacts of salinity and microbial community
312 on soil C decomposition (Fig. 7). The result showed that soil pH and EC had negative
313 contribution to bacterial diversity, while bacterial diversity had a strong positive
314 influence on the primed soil C (Fig. 5). For instance, salinity properties of EC had a
315 directly negative influence on the bacterial diversity but positive influence on the
316 primed soil C. Meanwhile, pH were negatively correlated with bacterial diversity and
317 positively correlated with substrate derived C.

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319 **4. Discussion**

320 4.1. Soil priming effects along salty gradients

321 Understanding soil C dynamics along salinity gradients is crucial to predict C
322 sequestration in salty soils. In the early stage of the incubation, we observed that the
323 cumulative substrate derived CO₂ in the soils with lower salinity was significantly
324 higher than soils with higher salinity (Fig. 1), which can be possibly explained by that
325 high salinity inhibited microbial activity. Many studies have reported the influence of
326 soil salinity on organic matter decomposition, mostly, the decomposition of organic
327 matter are decreased by salinity (Wichern et al., 2006; Ghollarata ~~et al.~~ ~~and Raiesi,~~
328 2007; Tripathi et al., 2007; Setia et al., 2012). Yet, the response of microbial community
329 to the increasing levels of salinity and consequent effects on soil priming effects
330 remains largely unknown.

331 Here, we found soil priming effects was gradually changed from negative to
332 positive priming effect (Fig. 1). The early pattern of the dynamics of the priming effect
333 in this study was similar to other studies showing preferential utilization of labile C
334 substance. The first phase of negative priming effects was likely to be caused by
335 microbial assimilation of substrate. The soil microbes turned to use the new added
336 substrate and thus used less of the original SOC. This was attributed to “preferential
337 substrate utilization” (Perelo et al., 2005).

338 Soil microbial biomass-related growth predominating in the first phase were most
339 likely to utilize SOC, leading to a positive priming effects after substrate was largely
340 vanished. The magnitude of priming effects depends on soil microbial biomass size
341 (Schneckenberger et al., 2008). It was found that the amount of added easily available
342 organic C is beyond 50% of microbial biomass C (Blagodatskaya ~~et al.~~ ~~and~~
343 ~~Kuzyakov,~~ 2008). Namely, the second phase of positive PEs probably was due to
344 increased biomass size and enhanced demand on SOC. Secondly, C that was
345 assimilated into microbial biomass in the first stage may also be mineralized in the

346 second stage due to the turnover of microbial biomass (Shahbaz et al., 2017; Perelo et
347 al., 2005).

348

349 4.2. Microbial community along salt gradients

350 Previous studies concerning the impact of salinity on soil microbial community
351 used different soils with a range of salt levels. In the present study we investigated the
352 influence of soil salinity on microbial communities in soils from the closed area
353 covering a range of salt content. Similarly, Rousk et al. (2011) also used agricultural
354 soils from the same area representing a range of soil salinity. Here, we found microbial
355 diversity (alpha diversity) decreased with increasing salinity (Fig. 3). The negative
356 impact on microbial diversity can be explained by that the accumulation of large
357 amounts of salt in the soil raised the extracellular osmotic concentration (Rath et al.,
358 2015; Rath and Rousk, 2015; Oren, 2011). The high osmotic pressures made it difficult
359 for many microorganisms to adapt to and thus reduce their biological activity. The
360 changes of soil microbial community structure were also explained by salinity
361 (Herlemann et al., 2011; Campbell et al., and Kirchman, 2013). We found that
362 *Bacteroidetes*, *Firmicutes*, *Acidobacteria* and *Deinococcus-Thermus* were dominant in
363 these soils (Fig. 4). These results are supported by previous findings that *Firmicutes*
364 possess the high salinity resistance. Other studies also found that *Bacteroidetes* is
365 dominant taxa in alkaline saline soil because of its resistant to salt (Valenzuela-Encinas
366 et al., 2009; Keshri et al., 2013). Other study shows that the dominant phyla are
367 *Bacteroidetes* and followed by *Proteobacteria* in the haloalkaline soil (Keshri et al.,
368 2013). These results are consistent with the esuarine or marine environments, despite
369 some studies suggest that soil salinity is not found to be a decisive factor for bacterial
370 community and their growth (Rousk et al., 2011).

371 The difference of microbial community structure is affected by many soil variables,
372 and pH and EC were the most important ones (Fig. 3; Table 2). Our results showed that
373 the value of soil pH and EC would significantly affect the microbial community

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374 structure and the combined contribution rate of these two variables to microbial
375 community was 43% (Table 2). At high levels of salt and alkaline arid condition, soil
376 pH has been also shown to have a very powerful influence on the soil bacterial
377 community structures (Bååth et al., ~~and Anderson~~, 2003; Fierer et al., ~~and Jackson~~,
378 2006; Rousk et al., 2010). Meanwhile, it is consequently unlikely that soil pH
379 differences between the studied soils obscured the influence of salt (Rousk et al., 2011).
380 Salinity has been identified as one of the most potent environmental factors that
381 determine assembly of microbiome. Salinity has been regarded to play the vital role in
382 shaping microbial community in different ecosystem. This, despite the clear evidence
383 from aquatic microbial ecology (Lozupone et al., ~~and Knight~~, 2007b), show a
384 potential for salt to affect soil microbial communities apart from that of pH (Rath et al.,
385 ~~and Rousk~~, 2015).

386

387 4.3. The core microbial taxa regulating C decomposition along salinity gradient

388 The correlation of microbial taxa and SOC decomposition (priming) were found
389 according to the results of O2PLS and SEM (Table 3; Fig. 5). Here we showed that
390 *Streptomyces* (*Actinobacteria*), *Glycomyces* (*branch of Actinobacteria*), *Agromyces*
391 (*branch of Actinobacteria*), and *Sphingomonas* (*branch of Proteobacteria*) at the genus
392 level were significantly correlated with the C process particularly primed soil-driven C.
393 Most of these functional taxa belonged to *Actinobacteria* and *Proteobacteria*. In a
394 recent study, Ren et al. (2018) found that *Actinobacteria* had negative impact on SOC
395 mineralization across land-use change (Fierer et al., 2007; Goldfarb et al., 2011) and
396 *Proteobacteria* drove the positive soil respiration (He et al., 2012; Stevenson et al.,
397 2004), indicating the balance of soil C dynamics were largely regulated by these two
398 phyla. We found similar result that *Streptomyces* (*branch of Actinobacteria*) had a
399 negative correlation with primed soil CO₂. *Actinobacteria* are able to grow
400 preferentially on the C-rich refractory materials and relatively easily decompose the

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401 cellulose, lignocellulose (Khodadad et al., 2011), indicating these microorganisms
402 preferentially use the C source that is used partially by others.

403 Although some studies suggest soil salinity may not be a vital factor for C
404 decomposers (Rousk et al., 2011), the composition of microbial community are
405 considered to play a decisive role in determining C dynamic processes in response to
406 salt stress (Ramsey et al., 2005; Schimel et al., 2007; Nottingham et al., 2009). Here,
407 SEM analysis showed that soil pH and EC in salted soils reduced microbial diversity
408 and thus limited the utilization of SOC by microbial community, It was reported that
409 high pH and salinity are the major determinants of soil microbial activity and
410 community structure (Kamble et al., 2014).

411

412 **5. Conclusion**

413 Cotton meal is a kind of organic material with high nitrogen content, adding cotton
414 meal in salinised soil can stimulate and promote the release of soil nutrients. The
415 microorganisms mainly use the organic matter in the cotton meal in the pre-culture
416 period, so the soil carbon excitation is negative excitation. Soil priming effect turned
417 from negative to positive at the later stage of incubation (day 28), because
418 microorganisms turned to decompose SOC from the labile substrate. With the increase
419 of salinity, the diversity of microbial community decreased. Soil microbial community
420 was mainly controlled by soil pH and EC. By O2PLS, we found *Actinobacteria* and
421 *Proteobacteria (Luteimonas, Hoeflea and Stenotrophomonas)* dominant in these soils
422 were the core microbial taxa that affecting the process of organic C mineralization,
423 particularly soil primed CO₂.

424

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429 **Data availability**

430 The datasets used and analysed during the current study available from the
431 corresponding author on reasonable request.

432

433 **Author contributions**

434 K.W. conceptualized and conducted the experiment. H.Z. and D.C. conducted the
435 data analysis and wrote the manuscript, conducted the indoor experiment. C.M. and
436 Z.Z. assisted in conducting the experiment. All authors reviewed the manuscript. All
437 authors contributed to the manuscript and approved the submitted version.

438

439 **Competing interests**

440 The authors declare no competing interests.

441

442 **Reference**

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Table 1. Soil samples and Cottonseed meal properties

	Salinity 1	Salinity 2	Salinity 3	Salinity 4	Salinity 5	Cottonseed meal
Total C (%)	3.38b	3.18c	3.16c	3.57a	3.35b	42.98
Total N (%)	0.18d	0.19d	0.20c	0.22b	0.26a	5.84
C/N ratio	18.32a	16.56b	15.71c	16.54b	12.94d	7.38
$\delta^{13}\text{C}$ value (%)	-14.21a	-14.79c	-14.60b	-14.55b	-16.01d	-23.47
pH (H₂O)	8.85a	8.45c	8.58b	8.59b	8.55b	7.63
EC (dS m⁻¹)	1.06e	1.96c	1.28d	2.64b	7.75a	2.56

Salinity (%)	0.25e	0.58d	0.75c	1.00b	2.64a	ND
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685 **Table 2.** Mantel test and Distance-based multivariate analysis relevance and
 686 contribution rate between soil properties and bacterial community compositions.

	pH	EC	NO ₃ -N	NH ₄ ⁺ -N	MBC	TN	TC
Correlation	0.74**	0.56**	0.36**	0.68**	0.31**	0.11	0.27
Contribution	0.31**	0.12**	0.05	0.04	0.16	0.03	0.07**

687 Note: * $P < 0.05$, ** $P < 0.01$

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690 **Table 3.** The variable influence projection (VIP) value and Spearman's correlation
 691 between the relative abundances of genera and C dynamic.

Phylum-Genus	VIP	Cottonseed meal CO ₂ -C(μg g ⁻¹)	Primed soil CO ₂ -C(μg g ⁻¹)	Basal soil CO ₂ -C(μg g ⁻¹)
<i>Actinobacteria-Actinomarinales</i>	1.36		0.63**	
<i>Proteobacteria-Luteimonas</i>	1.31		0.80**	
<i>Actinobacteria-Nocardioides</i>	1.30		0.54*	
<i>Proteobacteria-Hoeflea</i>	1.29		0.73**	
<i>Actinobacteria-Streptomyces</i>	1.27		-0.84**	
<i>Actinobacteria-Glycomyces</i>	1.26	0.63**		
<i>Actinobacteria-Marmoricola</i>	1.26	-0.52		
<i>Proteobacteria-Nitrosospira</i>	1.23		0.59	
<i>Actinobacteria-Intrasporangium</i>	1.22		0.60*	
<i>Actinobacteria-Agromyces</i>	1.19			0.58*
<i>Proteobacteria-Sphingomonas</i>	1.18			0.65**
<i>Actinobacteria-Myceligenans</i>	1.16			
<i>Chloroflexi-Nitrolancea</i>	1.15		0.65**	
<i>Actinobacteria-Pseudarthrobacter</i>	1.06		0.62**	
<i>Proteobacteria-Stenotrophomonas</i>	1.00	-0.50	0.72**	

692 Note: * $P < 0.05$, ** $P < 0.01$

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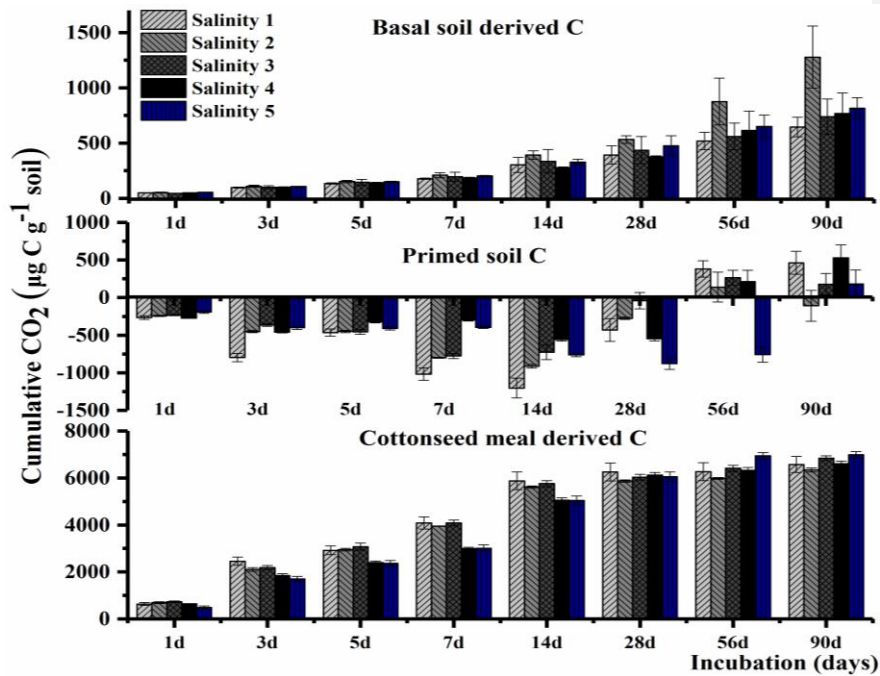
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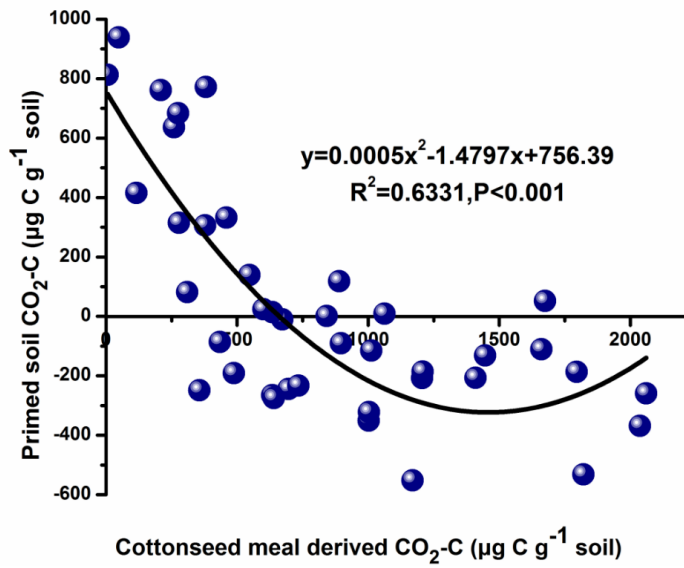
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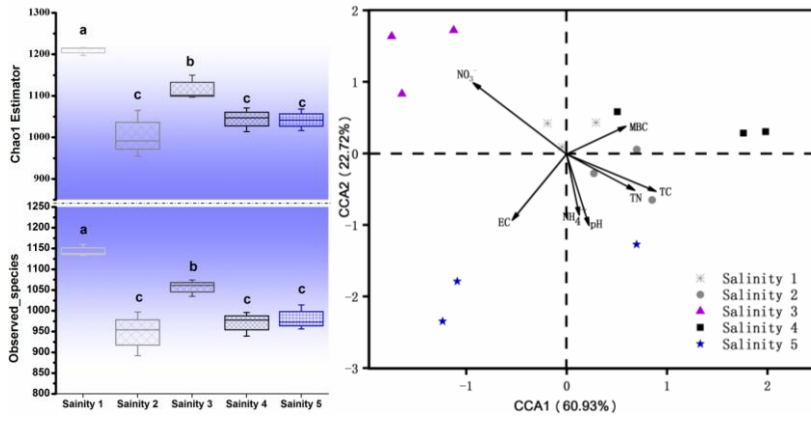


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700 **Fig. 1.** Partitioning of CO₂ evolution after addition of cottonseed meal in different five
701 salinity soils. Cumulative CO₂ evolved from salinity soil of 0.25 % (a) , 0.58 % (b) ,
702 0.75 % (c) ,1.00% (d) and 2.64%(e) . Error bars represent standard errors of the means
703 (n = 3).
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 709 **Fig. 2.** Correlation between primed soil mineralisation and cottonseed meal
 710 mineralisation following different five salinity soils during 90 days incubation
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723 **Fig. 3.** Microbial community alpha diversity (Chao1) observed_species and beta
 724 diversity. Within each panel, boxplot data refer to maximum date (top line), 99% (the
 725 second line), mean (the third line), 1% (the fourth line) and minimum date (bottom line)
 726 of the different treatments, with statistical significance ($P < 0.05$).

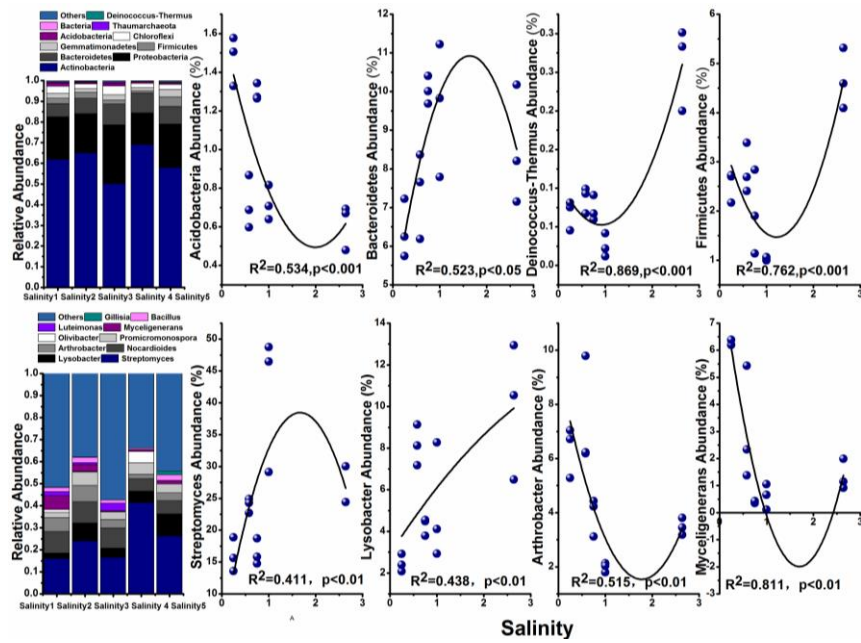
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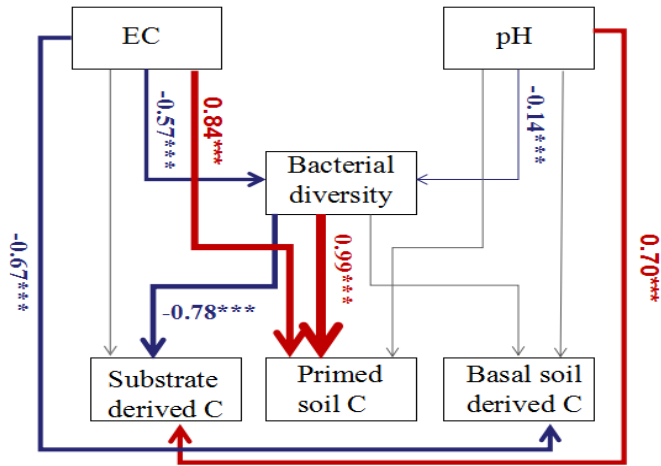
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 733 **Fig. 4.** The top 10 of phylums and genes in bacterial community in soils with a gradient
 734 of salinity
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$\chi^2 = 0.85, P = 0.65, GFI = 0.98, RMSEA < 0.001$

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Fig. 5. Path analysis detecting the underlying causal relationships between soil salinity physicochemical factors and microbial community composition of carbon dynamics in the soil system. Red lines indicate positive relationships, while blue lines indicate negative relationships. The width of arrows indicates the strength of significant standardized path coefficients ($P < 0.05$). Paths with non-significant coefficients are presented as gray lines. *** $P < 0.001$; ** $P < 0.01$; * $P < 0.05$

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