Palaeoecology of ungulates in northern Iberia during the Late Pleistocene through isotopic analysis of teeth

~~Evolutionary ecology of ungulates in northern Iberia during the Late Pleistocene through isotopic analysis on teeth~~

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Abstract

During the Late Pleistocene, stadial and interstadial fluctuations affected vegetation, fauna, and human groups that were forced to cope with these pronounced spatial-temporal climatic and environmental changes. These changes were especially abrupt during the Marine Isotopic Stage (MIS) 3. Here, we reconstruct the climatic trends in northern Iberia considering the stable isotopic composition of ungulate skeletal tissues found in archaeological deposits dated between 80 to 15 ka cal BP. The carbon and oxygen isotopic composition preserved in the carbonate fraction of tooth enamel provides a reliable and high-resolution proxy of the food and water consumed by these animals, which is indirectly related to the local vegetation, environment, and climate, allowing us to estimate paleotemperatures and rainfall intensity. This study presents new isotope data from 44 bovine, equid, and cervid teeth from five archaeological sites in the Vasco-Cantabrian region (El Castillo, Axlor, Labeko Koba, Aitzbitarte III interior and El Otero,) and one in northeastern Iberia (Canyars), where human evidence is attested from the Mousterian to the Magdalenian. The carbon isotope values reflect animals feeding on diverse C3 plants in open environments, and point to differentiated ecological niches for equids and bovines, especially during the Aurignacian in the Vasco-Cantabrian region. Temperature estimations based on oxygen isotopic compositions and rainfall obtained from carbon isotopic compositions indicate colder and more arid conditions than nowadays for the human occupations from the Late Mousterian to the Aurignacian. The contemporary northeastern Iberia site shows slightly lower temperatures related to an arid period when animals mainly graze in open landscapes. In the Vasco-Cantabrian region, during the MIS2, the Gravettian data reflect a landscape opening, whereas the Magdalenian points to warmer (but still arid) conditions.

Keywords: Middle and Upper Palaeolithic; Neanderthal; Homo sapiens, palaeoecology; geochemistry

1. Introduction

Understanding local and regional climatic variability during the Late Pleistocene in southern Europe is crucial for assessing the potential impact of climate on the adaptation and decline of Neanderthals and the subsequent expansion and resilience of Anatomically Modern Humans during the Upper Paleolithic (e.g., D'Errico and Sánchez Goñi, 2003; Finlayson and Carrión, 2007; Sepulchre et al., 2007; Staubwasser et al., 2018). During the Late Pleistocene, the climatic records demonstrate stadial and interstadial continuous fluctuations during the Marine Isotope Stage 3 (MIS 3, ca. 60-27 ka) and MIS 2 (ca. 27-11 ka). Human groups had to face those episodes, which affected vegetation and fauna to different extents, depending on the region. Northern Iberia is a key study area due to the abundance of well-preserved archaeological caves and rock shelters where, in the last decade, an updated and multidisciplinary approach has been applied to disentangle how changing environmental conditions affected the subsistence dynamics of Middle and Upper Paleolithic hominins. Recent chronological, technological, subsistence studies and ecological reconstructions are revealing a more complex regional panorama than previously known (e.g., Sánchez Goñi, 2020; Vidal-Cordasco et al., 2022; 2023; Timmermann, 2020; Klein et al., 2023).

The Vasco-Cantabrian region, located in northwestern Iberia, is subject to the influence of Atlantic climatic conditions, where recently has been evaluated the impact of the glacial-interglacial oscillations during MIS3 (Vidal-Cordasco et al., 2022). Modelling of traditional environmental proxies (small vertebrates and pollen) associated to archaeo-paleontological deposits show a progressive shift in the climatic conditions with decreasing temperatures and rainfall levels detected during the late Mousterian (Fernández-García et al., 2023). Ecological alterations have been observed in large mammals, such as niche partitioning between horses and cervids (Jones et al., 2018), a decrease in the available biomass for secondary consumers, and consequently, a reduction in the ungulate carrying capacity ((Jones et al., 2018; Vidal-Cordasco et al., 2022). Cold and arid conditions are observed during the Aurignacian and the Gravettian until the onset of MIS2. Afterwards, during the Last Glacial Maximum (LGM, 23-19 ka), the global climatic deterioration associated with this glacial phase results in colder and more arid conditions in the region, with a predominance of open landscapes. However, this region still provided resources for human survival acting as a refugia with more humid conditions in comparison to the Mediterranean area (Cascalheira et al., 2021; Fagoaga, 2014; Fernández-García et al., 2023; Garcia-Ibaibarriaga et al., 2019a; Lécuyer et al., 2021; Posth et al., 2023). By the end of the LGM, a climate amelioration and a moderate expansion of the deciduous forest are documented from the late Solutrean through the Magdalenian (Garcia-Ibaibarriaga et al., 2019a; Jones et al., 2021).

In contrast, northeastern Iberia is influenced by the Mediterranean climate. The MIS 3 human settlement in this region have been linked to cooler temperatures and with higher rainfall, compared to the present, but with climatic fluctuations less pronounced compared to the Vasco-Cantabrian region (López-García et al., 2014; Fernández-García et al., 2020; Vidal-Cordasco et al., 2022). Archaeobotanical and small vertebrate evidence indicate relatively stable climatic conditions, but also suggest the persistence of open forests during the Middle to Upper Paleolithic transition, as found in northwestern Iberia (Allué et al., 2018; Ochando et al., 2021). However, certain archaeological records indicate specific climatic episodes, such as increased aridity and landscape opening during Heinrich ~~Events~~ Stadials 4 and 5 (e.g., Álvarez-Lao et al., 2017; Daura et al., 2013; López-García et al., 2022; Rufí et al., 2018).

These multi-proxy studies have significantly expanded our understanding of the environmental evolution in Iberia, alongside proxies derived from marine core records in Iberia margins (Fourcade et al., 2022; Martrat et al., 2004; Naughton et al., 2007; Roucoux et al., 2001; Sánchez-Goñi et al., 1999, 2009) and other regional paleoclimatic records sourced from local natural deposits (e.g., Pérez-Mejías et al., 2019; Moreno et al., 2010, 2012; González-Sampériz et al., 2020; Ballesteros et al., 2020). However, the availability of proxies enabling the direct connections between these environmental shifts and human activities remains limited.

In this study, we investigate the palaeoecological and palaeoenvironmental dynamics in northern Iberia during the late Middle and Upper Paleolithic by measuring the carbon and oxygen isotopic composition of bioapatite carbonates ($\delta^{13}C_{carb}/\delta^{18}O_{carb}$) preserved in archaeological mammal teeth. These analyses provide high-resolution snapshots of ecological information from animals accumulated during human occupations at

the caves. Tooth enamel forms incrementally and does not biologically remodel (Kohn, 2004; Passey and Cerling, 2002), in contrast to other bodily tissues such as bone, which implies that the isotope values measured on them reflect the animal diet and water sources consumed during its mineralisation, around one to two years of life for the species included in our study (bovids, equids, cervids)(e.g., Hoppe et al., 2004; Pederzani and Britton, 2019; Ambrose and Norr, 1993; Luz et al., 1984). The preserved carbon isotope composition relies on animal dietary choices reflecting mainly the type of plant consumed (C3/C4), exposition to light and humidity levels. Otherwise, the oxygen isotope composition reflects mainly the environmental water consumed by animals, directly by drinking or through diet, which reflects isotopic information derived from water sources as well as changes in climatic conditions. Both indirectly provide information on the vegetation and climate that allows estimating past temperatures, rainfall, and moisture on a sub-annual scale, returning isotopic data of the foraging areas where animals were feeding during teeth formation.

By analysing the stable isotopic composition of 44 ungulate teeth obtained from 15 archaeological levels directly associated with human occupation, including El Castillo, Axlor, Labeko Koba, Aitzbitarte III interior and El Otero in northwestern Iberia, and Terrasses de la Riera dels Canyars in northeastern Iberia, this study presents novel insights into local and regional environmental and climatic trends associated to human presence during the Late Pleistocene (Fig.1; Fig.2; Appendix A). Specifically, it focuses on the Middle to Upper Paleolithic transition in both areas and the post-LGM period in the Vasco-Cantabrian region.

The main objectives of this work are: 1) to assess how regional environmental conditions, including changes in moisture and vegetation cover, but also temperatures and rainfall, are recorded in the stable isotopic composition of tooth enamel; 2) to characterize animal diet and their ecological niches; 3) to obtain quantitative temperature data to compare with available proxies; 4) to characterise seasonal patterns of animals found in the archaeological sites by identifying winter and summer fluctuations.

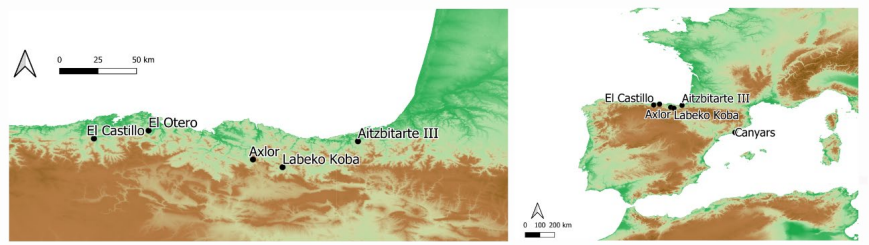

Figure 1. Location of the archaeological sites included in this study. From west to east, in the autonomous community of Cantabria, El Castillo, and El Otero; in the Basque Country, Axlor and Aitzbitarte III interior; in Catalonia, Canyars.

2. Archaeological sites and sampled material

This study selected a total of 44 ungulate teeth including 25 bovines (*Bos primigenius, Bison priscus, Bos/Bison* sp.), 14 equids (*Equus* sp. and *Equus ferus*), and five cervids (*Cervus elaphus*) originating from five archaeological sites in the Vasco-Cantabrian region (El Castillo, El Otero, Axlor, Labeko Koba, Aitzbitarte III interior) and one in the northeastern area (Terrasses de la Riera dels Canyars, henceforth Canyars). These teeth were recovered from 15 archaeological levels attributed to the following technocomplexes: Mousterian (n=14), Transitional Aurignacian (n=10), Châtelperronian (n=2), Aurignacian (n=12), Gravettian (n=1) and Magdalenian (n=5) (Table 1 and 2). Archaeozoological studies of the archaeological sites are available (synthesis in Marín-Arroyo and Sanz-Royo, 2022; Daura et al., 2013) and most prove that faunal remains were accumulated by human acquisition during the different cultural phases. The isotopic results of equids teeth and other ungulates bone collagen from El Castillo were previously published by Jones et al. (2019) in combination with the stable isotopes of ungulates from the site, as well as the combined bioapatite carbonate and phosphate analyses of bovines from Axlor (Pederzani et al., 2023). A comprehensive description of each archaeological site is provided in Appendix A.

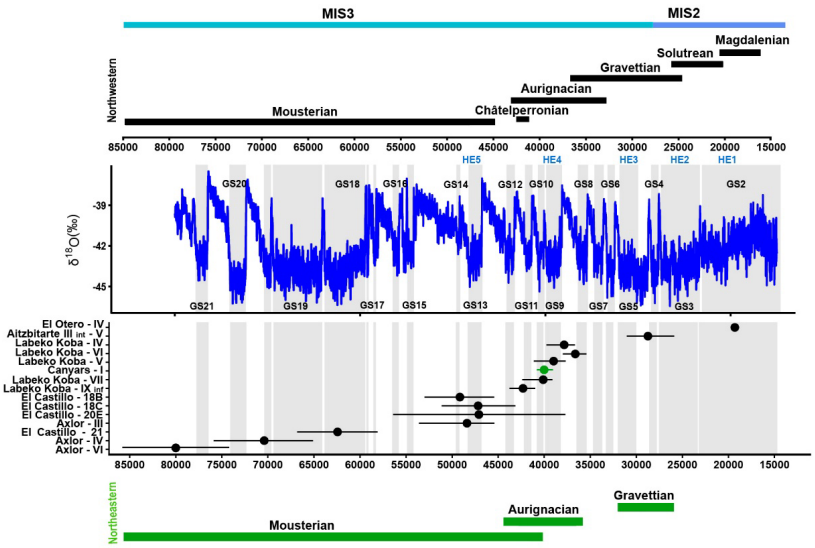

Figure 2. Representation of the duration each archaeological level (dots represent the median values, bars represent 95% confidence intervals for -14C dates and 68% for ESR and OSL dates) related to techno-complexes in both northwestern (in black) and northeastern Iberia (in green) and the δ18O record from the NGRIP (North Greenland Ice Core Project members, 2004; Rasmussen et al., 2014). Grey bands indicate Greenland Stadials (GS). Dates from EL Castillo (C14 UF, ESR), El Otero (C14 UF), Axlor (C14 UF, OSL), Labeko Koba (C14 UF), Aitzbitarte III-interior (C14 AMS) and Canyars (C14 UF, ABA, ABOx-SC) are shown in Appendix B and C.

~~**Figure 2.** Representation of the duration each archaeological level (dots represent the median values, bars represent 95% confidence intervals for ^{14}C dates and 68% for ESR and OSL dates) related to techno-complexes in both northwestern (in black) and northeastern Iberia (in green) and the δ^{18}O record from the NGRIP (North Greenland Ice Core Project members, 2004; Rasmussen et al., 2014). Grey bands indicate Greenland Stadials (GS). Detailed information on OSL, ESR and ^{14}C dates, along with ^{14}C calibration, are shown in Appendix B and C.~~

3. Methods

3.1 Methods: Dating methods

Individual Bayesian age models were built for Canyars, El Castillo, Labeko Koba and Aitzbitarte III interior based on radiocarbon dates (AMS UF and non-UF, ABOx-SC and ABA pretreatments on bones and charcoal remains) using OxCal4.4 software (Ramsey, 2009), considering the INTCAL20 calibration curve (Reimer et al., 2020) (Appendix C). The Bayesian model enables the modification of the calibrated Probability Distribution Function (PDF) of individual dates based on the existing relative stratigraphic and other relative age information. A resolution of 20 years was assumed, being a reasonable balance between required accuracy and computational costs. An order function in the OxCal was used to calculate the probability that one PDF predated another, providing information to assess synchronicity and temporal overlap of individual archaeological levels and cultural phases in each of the four separate sites modelled. Dates were organised into a 'Sequence,' and chronological information for each level was grouped into a single 'Phase' with start and end 'boundaries' to bracket each archaeological level. The interval between the start of each level and its end provided the duration of each level. In all cases, convergence was greater than 95%. CQL codes, individual Bayesian models and modelled dates per site are reported in Appendix C.

No chronological models were built for El Otero because only a single date was obtained for level IV and El Castillo levels 20E and 21 (ESR dated) and Axlor levels III, IV and VI (OSL dated) because dates go beyond the limit of the radiocarbon. To show the duration of these levels in combination with the other sites and levels, each of these dates was estimated by adding and subtracting the sigma (68% Confidence Interval) from the uncalibrated date. In this way, we estimated the duration of these levels to be beyond 55 ka cal BP.

3.2 Tooth sampling

All teeth included were sequentially sampled to reconstruct the complete $\delta^{18}O_{carb}$ and $\delta^{13}C_{carb}$ intratooth profiles based on enamel carbonate bioapatite. Intratooth sequential sampling was applied to the second and third molars and third and fourth premolars. Bovine and horse teeth sampled exceeded 3-4 cm of crown height to ensure that at least a one-year isotopic record of animal life was obtained (Britton et al., 2019; Hoppe et al., 2004). Samples were taken perpendicular to the growth axis on the tooth where the enamel was best preserved, avoiding, whenever possible taphonomic alterations such as cracks or postdepositional damages. Samples were performed in the buccal face for the lower teeth and the lingual part for the upper ones. The outermost enamel surface was abraded to remove the superficial enamel, calculus, cementum, or concretions adhering to the surface to avoid contaminations. The sequential sampling consisted of straight strips (ca. 8 x 1.5 x 1 mm) covering the width of the selected lobe, approximately every 2-3 mm, from the crown to the Enamel-Root-Junction (ERJ). The sample depth covered around 75% of the enamel depth, and dentine inclusion was avoided. A low-revolution variable-speed manual drill was used, equipped with 1 mm diamond-coated drill bits of conical and cylindrical shape. About 10-15mg of enamel powder was collected in each subsample, generating 693 subsamples for IRMS measurements (see complete intratooth profiles in Appendix D).

Site	Level - Cultural period	Bovines	Horses	Red deer	Teeth	Subsamples
Axlor	VI - Mousterian	2			2	32
	IV - Mousterian	1			1	12
	III - Mousterian	4			4	62
El Castillo	21A - Mousterian	2	1		3	47
	20E - Mousterian	2	2		4	56
	18C - Trans. Aurignacian	4			4	66
	18B - Trans. Aurignacian	3	2	1	6	93
Labeko Koba	IX inf - Châtelperronian		1	1	2	24
	VII - ProtoAurignacian	3			3	68
	VI - Aurignacian		1		1	16
	V - Aurignacian	1	1		2	39
	IV - Aurignacian		1		1	16
Canyars	I - Aurignacian	2	3		5	76
Aitzbitarte III interior	V - Gravettian	1			1	18
El Otero	IV - Magdalenian		2	3	5	68
TOTAL		**25**	**14**	**5**	**44**	**693**

Table 1. Number of teeth sampled by species, archaeological sites and cultural periods.

3.3 Sample treatment and stable isotope mass spectrometry

Several authors have debated the necessity of chemical pre-treatments to remove organic matter and secondary carbonates from bioapatite carbonates before stable isotopic analysis. Some chemical treatments can introduce secondary carbonates, increase carbonate content, and alter the original isotopic

signal (Pellegrini and Snoeck, 2016; Snoeck and Pellegrini, 2015). For this reason, in this work, most of the samples were not pretreated except for the equids and cervids samples from Labeko Koba, El Otero and El Castillo that were sampled and pretreated in an earlier phase of the project. The absence of pretreatment can elevate the risk of secondary carbonates (Chesson et al., 2021; France et al., 2020). Nonetheless, any pretreatment method cannot guarantee their complete removal, and the 'side effects' may compromise the final isotopic signal to a greater extent. While variations in pretreatment methods exist among samples in this study, the lack of a universally accepted protocol necessitates careful consideration of any potential isotopic effects resulting from these differences.

Pretreatment was followed for above-mentioned samples from fourteen teeth, where around 7 mg of powdered enamel was prepared and pretreated with 3% of sodium hypochlorite (NaOCl) at room temperature for 24 h (0.1 ml/mg sample) and thoroughly rinsed with deionised water, before a reaction with 0.1M acetic acid for 4 h (0.1 ml/mg sample) (Balasse et al., 2002; equivalent protocol in Jones et al., 2019). Samples were then thoroughly rinsed, frozen, and freeze-dried. NaOCl is one of the most common agents used for pretreating carbonates and works as a base that removes organic matter by oxidation. Although it is considered one of the most efficient agents for removing organic matter, it can induce the absorption of exogenous carbonates, such as atmospheric CO_2 and secondary carbonates (Pellegrini and Snoeck, 2016; Snoeck and Pellegrini, 2015). It is argued that acetic acid after NaOCl pretreatment can remove exogenous carbonates absorbed during NaOCl application. However, it is unclear if all newly introduced carbonates are finally released and which effect they produce on the original isotopic composition. These samples were analysed in the Godwin Laboratory (Department of Earth Sciences, University of Cambridge). Enamel powder samples were reacted with 100% orthophosphoric acid for 2 h at 70°C in individual vessels in an automated Gasbench interfaced with a Thermo Finnigan MAT253 isotope ratio mass spectrometer. Results were reported in reference to the international standard VPDB and calibrated using the NBS-19 standard (limestone, $\delta^{13}C$ = +1.95‰ and $\delta^{18}O$ = -2.2‰; Coplen, 2011) for which the precision is better than 0.08‰ for $\delta^{13}C$ and 0.11‰ for $\delta^{18}O$.

For the non-pre-treated samples, carbon and oxygen stable isotopic ratios were measured using continuous flow-isotope ratio mass spectrometry, specifically a Europa Scientific 20-20 IRMS coupled to a chromatograph, at the Iso-Analytical laboratory in Cheshire, UK. The samples were weighed into clean exetainer tubes after being flushed with 99.995% helium. Phosphoric acid was then added to the samples, and they were allowed to react overnight to ensure the complete conversion of carbonate to CO_2, following the method outlined by Coplen et al. (1983). The reference materials used for VPDB calibration and quality control of the analysis included IA-R022 (calcium carbonate, $\delta^{13}C$ = -28.63‰, $\delta^{18}O$ = -22.69‰), NBS-18 (carbonatite, $\delta^{13}C$ = -5.01‰, $\delta^{18}O$ = -23.2‰), IA-R066 (chalk, $\delta^{13}C$ = +2.33‰; $\delta^{18}O$ = -1.52). The accepted values of the in-house standards IA-R022 and IA-R066 were obtained by calibrating against IAEA international reference materials, NBS-18 and NBS-19, and NBS-18 and IAEA-CO-1 (Carrara marble, $\delta^{13}C$ = 2.5‰, and $\delta^{18}O$ = -2.4‰), respectively. Additionally, in-house standards long-term measured were used: ILC1 (calcite, $\delta^{13}C$ = 2.13, $\delta^{18}O$ = -3.99‰), and Y-02 (calcite, $\delta^{13}C$ = 1.48, $\delta^{18}O$ = -9.59‰). The analytical precision of quality control standard replicates was better than 0.09‰ for $\delta^{13}C$ and better than 0.12‰ for $\delta^{18}O$. The calcium carbonate content test of these samples, ranging between 3.9% and 8.9%, does not indicate a substantial presence of secondary carbonates, considering Chesson et al. (2021). Additionally, phosphate results on samples from Axlor showed $\delta^{18}O_{carb}$-$\delta^{18}O_{phos}$ offsets within the expected range for well-preserved samples (Pederzani et al., 2023).

3.4 Carbon stable isotopic compositions as environmental and ecological tracers

To unravel animal diet and compare the different species, in standardised terms, it is necessary to consider the enrichment factor (ε^*) between $\delta^{13}C$ obtained by the animal on its diet ($\delta^{13}C_{diet}$) and $\delta^{13}C$ recorded on

enamel carbonates ($\delta^{13}C_{carb}$) (Bocherens, 2003; Cerling and Harris, 1999). The ε* estimated for large ruminant mammals results in an offset of around 14.1‰ between diet and dental enamel, commonly applied to medium-sized herbivores. However, it is well-known that this offset varies between species, considering animals' different physiological parameters. Recently, a formal model to predict species-specific diet-consumer isotopic offsets has been proposed, which uses body mass (BM) and digestive physiology as the main factors that regulate the ε* (Tejada-Lara et al., 2018). This model proposes the following prediction equations for ruminant or foregut fermenters (Equation 1: Eq.1) and hindgut fermenters (Eq. 2):

(Eq. 1) $\varepsilon^* = 2.34+0.05\ (BM)$ [r^2=0.78; p-value=0.008]

(Eq. 2) $\varepsilon^* = 2.42+0.032\ (BM)$ [r^2=0.74; p-value=0.003]

This work compares species with different digestive physiology, ruminants for bovines and cervids, and non-ruminants for equids. The ε* value was adjusted for each animal to avoid bias from digestive physiology when comparing these species. The following enrichment factors have been used: 14.6‰ for *Bos taurus* (Passey et al., 2005a), 13.7‰ for *Equus caballus* (Cerling and Harris, 1999), and 13.2‰ for *Cervus elaphus* (Merceron et al. (2021) following (Eq. 1) for ruminants with a mean body mass of 125 kg.

In body tissues, carbon isotopic composition is considered a combination of diet (understood as consumed food), environment openness (and associated exposure to light), and the amount of precipitation. Assuming that $\delta^{13}C$ of past vegetation is close to $\delta^{13}C_{diet}$ of ungulates, Lécuyer et al. (2021) proposed to estimate Mean Annual Precipitations (MAP) from $\delta^{13}C_{carb}$, derived from diets based on C3 plants. After transforming $\delta^{13}C_{carb}$ to $\delta^{13}C_{diet}$ using the enrichment factors established above, this work suggested transforming this value to $\delta^{13}C$ from vegetation ($\delta^{13}C_{leaf}$). However, the isotopic composition of animals' diet may not directly reflect vegetation cover, but rather the food preference of the animal and this approach should be discussed alongside other environmental data.

The MAP estimation is based on least square regression developed by Rey et al. (2013) and based on Kohn (2010) dataset (Eq.4), which requires first to estimate the $\delta^{13}C_{leaf}$ (Eq. 3). The $\delta^{13}C$ values of atmospheric CO_2 ($\delta^{13}C_{atm}$) are fixed in -7‰ (Lécuyer et al., 2021; Leuenberger et al., 1992; Schmitt et al., 2012). Atmospheric CO_2 levels have varied throughout the Late Pleistocene, with $\delta^{13}C_{atm}$ range between -7 to -6.4‰ (Eggleston et al., 2016), favouring an age-specific correction approach. However, maintaining general corrections is preferred considering the chronological uncertainty of the studied levels.

(Eq.3) $\delta^{13}C_{leaf}\ (VPDB) = (\delta^{13}C_{atm} - \delta^{13}C_{diet}) / [1+(\delta^{13}C_{diet} / 1000)]$

(Eq.4) $Log1(MAP+300) = 0.092(\pm0.004) \times \delta^{13}C_{leaf} + 1.148(\pm0.074)$

Additionally, Lécuyer et al. (2021) equation also accounts for the pCO_2 effect on $\delta^{13}C_{leaf}$ estimation, which is expected to result in an offset of +1‰ from current levels (considering that pCO_2 was lower than that experienced after the deglaciation period). If this correction was not applied, MAP results could be underestimated by -150mm. In agreement with Lécuyer et al. (2021) appreciation, these MAP estimations are a preliminary approximation and should be cross-validated with other environmental proxies. The associated uncertainties range from ±100 to 200 mm, influencing the interpretation of the final values.

3.5 Oxygen stable isotope compositions as environmental tracers

Stable oxygen isotopes from meteoric water (mainly derived from rainfall) strongly correlate with mean air temperatures in mid to high latitudes (Dansgaard, 1964; Rozanski et al., 1992) on a regional-to-local scale. Obligate drinkers, like bovines and horses, acquire this water and record its isotopic composition in their teeth and bones with a fixed but species-specific offset (Pederzani and Britton, 2019). Considering this two-

step relationship, past climatic conditions can be estimated. However, most of the temperature reconstructions based on $\delta^{18}O$ have considered the $\delta^{18}O$ from the phosphate fraction of bioapatite enamel ($\delta^{18}O_{phos}$) to build linear correlations between tooth enamel and drinking water $\delta^{18}O$ and obtain climatic information. For this reason, the $\delta^{18}O_{carb}$ values obtained in this work were converted into $\delta^{18}O_{phos}$. To do so, first, to express in VSMOW notation, the $\delta^{18}O_{carb}$ was corrected using the following correlation (Brand et al., 2014; Coplen et al., 1983):

(Eq.5) $\delta^{18}O_{carb}$ (VSMOW)= 1.0309 x $\delta^{18}O_{carb}$ (VPDB) + 30.91

Second, considering the relationship existent in tooth enamel between the carbonate and phosphate fraction (Iacumin et al., 1996; Pellegrini et al., 2011), from a compilation of the existent bibliography of modern animals measurements (Bryant et al., 1996; Pellegrini et al., 2011; Trayler and Kohn, 2017), Pederzani et al. (2023) proposed the following correlation:

(Eq.6) $\delta^{18}O_{phos}$ (VSMOW) = 0.941 x c (VSMOW) - 7.16

Once the isotopic information is expressed in $\delta^{18}O_{phos}$ (VSMOW), we can estimate the $\delta^{18}O$ on meteoric waters ($\delta^{18}O_{mw}$). It is known that different physiological factors will condition how oxygen isotope composition is fixed in each mammalian group. Thus, the correlations are usually species-specific and developed considering the physiology of each animal group. The obligate drinkers heavily rely on consuming large amounts of liquid drinking water, being the relative contribution of water from plants negligible and then minimizing the possible impact of isotopic enrichment through evapotranspiration in plants (Hoppe, 2006; Maloiy, 1973, Pederzani and Britton, 2019). However, certain types of drinking behaviours can impact $\delta^{18}O$, such as systematic consumption of certain highly buffered water sources (rivers or lakes), can significantly attenuate the final signal recorded. The correlation employed by this work relies on recent data compilations (Pederzani et al., 2021b, 2023). In the case of horses (Eq. 7), it has been considered the data combination of Blumenthal et al. (2019); Chillón et al. (1994); Bryant et al., 1994; Delgado Huertas et al., 1995), whereas for bovines (Eq. 8) the data from D'Angela and Longinelli (1990) and Hoppe (2006) have been put together in Eq. 4. To estimate $\delta^{18}O_{mw}$ from red deer remains, we selected D'Angela and Longinelli (1990) correlation (Eq. 9):

(Eq.7) $\delta^{18}O_{mw}$ (VSMOW)= ($\delta^{18}O_{phos}$ (VSMOW) - 22.14) / 0.62

(Eq.8) $\delta^{18}O_{mw}$ (VSMOW)= ($\delta^{18}O_{phos}$ (VSMOW) - 22.36) / 0.78

(Eq.9) $\delta^{18}O_{mw}$ (VSMOW)= ($\delta^{18}O_{phos}$ (VSMOW) - 24.39) / 0.91

Finally, paleotemperatures estimations from $\delta^{18}O_{mw}$ are typically approached using a geographically adjusted linear regression, which can vary from precise adjustments (aimed at reducing errors) to broader geographical adjustments that encompass more variability but are less precise (e.g., Pryor et al., 2014; Skrzypek et al., 2011; Tütken et al., 2007). In this work, temperatures were calculated considering the linear regression model relating $\delta^{18}O_{mw}$ and air temperatures proposed by Pederzani et al. (2021) based on monthly climatic records (monthly mean $\delta^{18}O_{mw}$ and monthly mean air temperatures), from Western, Southern and Central Europe stations from the Global Network of Isotopes in Precipitation (IAEA/ WMO, 2020). Considering current IAEA data sets from northern Iberia, there is a strong positive relationship between $\delta^{18}O_{mw}$ and annual or monthly temperatures (Moreno et al., 2021). However, it is known that Iberia is under a mixed influence between Atlantic and Mediterranean moisture sources that affects the isotopic composition of rainfall (Araguas-Araguas and Diaz Teijeiro, 2005; García-Alix et al., 2021; Moreno et al., 2021). Given uncertainties in past atmospheric circulation patterns and the limited availability of reference stations, it was deemed most appropriate to select an equation that extends beyond the borders of Iberia

and incorporates higher variability. Different correlations were for mean annual temperature (Eq. 10), summer (Eq. 11), and winter (Eq. 12) temperatures (T):

(Eq.10) $\delta^{18}O_{mw}$ (VSMOW)= (0.50 x T) - 13.64

(Eq.11) $\delta^{18}O_{mw}$ (VSMOW) = (0.46 x T) - 14.70

(Eq.12) $\delta^{18}O_{mw}$ (VSMOW)= (0.52 x T) - 11.26

Nonetheless, oscillations between glacial and interglacial conditions in the past have influenced global ice volume and sea level fluctuations (Dansgaard, 1964; Shackleton, 1987), impacting seawater oxygen isotope composition and the surface hydrological cycle on a worldwide scale, including $\delta^{18}O_{mw}$ (Schrag et al., 2002). Prior studies have used sea level information to correct $\delta^{18}O_{mw}$ (e.g., Fernández-García et al., 2019; Schrag et al., 2002). Given the chronological uncertainty in the studied levels, a general correction was applied to $\delta^{18}O_{mw}$ before temperature estimations, following Fernández-García et al. (2020) approach. Considering the mean sea level descent for the MIS 3 period (50 meters below present-day sea level)(Chappell and Shackleton, 1986), this may have contributed to a potential increase in the global $\delta^{18}O_{mw}$ value by ≈0.5‰, inferring a bias in calculated air temperatures of ≈1ºC.

Due to the uncertainties incurred from converting stable isotope measurements to palaeotemperature, the final estimations in this work should be considered exploratory and as a method of standardisation to make results comparable among different sites, species, and other non-isotopic palaeoclimatic records. In these estimations, the associated error from converting $\delta^{18}O_{phos}$ to MAT is enlarged by the uncertainty derived from the transformation of $\delta^{18}O_{carb}$ (VPDB) to $\delta^{18}O_{phos}$ (VSMOW) (see Pryor et al., 2014; Skrzypek et al., 2016 for further discussion). However, Pryor et al. (2014) and Pederzani et al. (2023) concluded that the impact of this conversion is negligible compared to the error propagation in subsequent calibrations used for temperature estimations from $\delta^{18}O_{phos}$. These associated errors were quantified following the methodology outlined by Pryor et al. (2014) (Appendix B).

3.6 Inverse modelling applied to intratooth profiles

Intratooth profiles frequently provide a time-averaged signal compared to the input isotopic signal ($\delta^{13}C$/ $\delta^{18}O_{carb}$) during enamel formation (Passey et al., 2005b). This signal attenuation is caused by time-averaging effects incurred through the extended nature of amelogenesis and tooth formation, and through the sampling strategy. During mineralisation, the maturation zone, which is time-averaged, often affects a large portion of the crown height and might affect the temporal resolution of the input signal of the sample taken. To obtain climatically informative seasonal information on the analysed teeth, the inverse modelling method proposed by (Passey et al. (2005b) is applied in this work. This method computationally estimates the time-averaging effects of sampling and tooth formation to obtain the original amplitude of the isotopic input signal more accurately, thus, to summer and winter extremes (Appendix E). This method considers parameters based on the amelogenesis trends of each species and sampling geometry, which are critical for a meaningful interpretation of intratooth isotope profiles. The model also estimates the error derived from the sampling uncertainty and the mass spectrometer measurements to evaluate the data's reproducibility and precision. This method was initially developed for continuously growing teeth, taking into account a constant growth rate within a linear maturation model, with a progressive time-average increment as sampling advances along the teeth profile. The species studied in this research exhibit non-linear tooth enamel formation, particularly in later-forming molars (Bendrey et al., 2015; Blumenthal et al., 2014; Kohn, 2004; Passey and Cerling, 2002; Zazzo et al., 2012). Although the model mentioned above is not ideal, as it does not take into account non-linear enamel formation and specific growth parameters for the species included are unknown, it is the best estimation based on the current state of the field and remains widely used (Pederzani et al., 2021a, b, 2023). Flat and less sinusoidal profiles are less suitable for the application of the model, given its

inherent assumption of an approximately sinusoidal form. Therefore, we chose not to apply this methodology in the analysis of intratooth $\delta^{13}C$ profiles, and it is recommended to approach the interpretation of model outcomes for non-sinusoidal $\delta^{18}O$ curves with caution. Further details on the application of this method can be found in Appendix E.

Following Pederzani et al. (2021b), mean annual temperatures (MAT) were deduced from the average of $\delta^{18}O_{carb}$ values between summer and winter detected in original sinusoidal intratooth profiles (Appendix D). This work shows that comparable results for annual means can be obtained before and after model application, but doing it beforehand avoids the associated errors induced by the inverse model. To maximize data, in non-sinusoidal teeth profiles, MAT was deduced from the average of all points within a tooth. However, this approach is less reliable when complete annual cycles are not recorded. When possible, summer and winter temperature estimations were derived from the obtained $\delta^{18}O_{carb}$ values after inverse modelling application, aiming to identify the corrected seasonal amplitude, which is dampened in the original $\delta^{18}O_{carb}$ signal.

3.7 Present-day isotopic and climatic data

Present-day climatic conditions surrounding each site have been considered, allowing an inter-site comparison, essential for compare this study with other regional and global data. Considering current MATs and MAPs, estimated climatic data is expressed in relative terms as MAT and MAP anomalies. Present-day summer and winter temperatures were also considered. Present-day temperatures and precipitation values were obtained from the WorldClim Dataset v2 (Fick and Hijmans, 2017) (Appendix B). This dataset includes the average of bioclimatic variables between 1970-2000 in a set of raster files with a spatial resolution every 2.5 minutes. The exact location of the selected archeo-palaeontological sites was used, using geographical coordinates in the projection on modern climatic maps with QGIS software.

Present-day $\delta^{18}O_{mw}$ values from the analysed sites' areas were obtained using the Online Isotopes in Precipitation Calculator (OIPC Version 3.1 (4/2017); Bowen, 2022) based on datasets collected by the Global Network for Isotopes in Precipitation from the IAEA/WMO (Appendix B).

Site	Level	Culture	Species	Tooth type	Code	CCE (%)	n	δ13Ccarb VPDB (‰)	min	max	SD	Range	δ18Ocarb VPDB (‰)	min	max	SD	Range
Axlor	III	Mousterian	*Bos/Bison* sp.	LRM3	AXL59	5.6	14	-8.9	-9.6	-8.2	1.4	0.4	-6.0	-7.3	-5.2	0.7	2.1
Axlor	III	Mousterian	*Bos/Bison* sp.	LRM2	AXL60	5.5	18	-9.7	-10.0	-8.9	1.1	0.3	-5.7	-6.8	-4.6	0.7	2.2
Axlor	III	Mousterian	*Bos/Bison* sp.	LRM3	AXL65	6.2	13	-8.9	-9.3	-8.1	1.2	0.4	-6.0	-7.2	-4.6	0.8	2.6
Axlor	III	Mousterian	*Bos/Bison* sp.	LRM2	AXL66	5.6	16	-8.9	-9.8	-8.3	1.5	0.5	-4.8	-6.1	-3.8	0.7	2.3
Axlor	IV	Mousterian	*Bos/Bison* sp.	LRM2	AXL70	5.7	12	-9.1	-9.4	-8.6	0.7	0.3	-5.3	-7.3	-3.9	1.2	3.4
Axlor	VI	Mousterian	*Bos/Bison* sp.	LLM3	AXL77	5.9	14	-9.7	-10.2	-9.2	1.0	0.4	-6.2	-7.9	-5.0	0.9	2.9
Axlor	VI	Mousterian	*Bos/Bison* sp.	LLM3	AXL86	5.5	18	-9.9	-10.2	-9.3	0.9	0.3	-5.4	-6.5	-3.8	0.7	2.6
El Castillo	20E	Mousterian	*Equus* sp.	LRP3/LRP4	CAS60		14	-11.9	-12.5	-11.5	1.0	0.3	-3.3	-4.1	-2.4	0.4	1.6
El Castillo	20E	Mousterian	*Equus* sp.	LRP3/LRP4	CAS61		14	-12.2	-12.4	-12.1	0.3	0.1	-4.9	-5.8	-4.3	0.4	1.5
El Castillo	20E	Mousterian	*Bos/Bison* sp.	LLM2	CAS139	6.7	16	-11.6	-12.2	-11.2	0.9	0.3	-5.6	-6.3	-4.9	0.5	1.4
El Castillo	20E	Mousterian	*Bos/Bison* sp.	LLM2	CAS140	5.7	12	-11.5	-11.9	-11.1	0.8	0.3	-5.5	-6.3	-4.6	0.6	1.7
El Castillo	21A	Mousterian	*Bos/Bison* sp.	LLM3	CAS141	5.7	15	-11.2	-11.5	-10.9	0.6	0.2	-5.4	-6.5	-4.3	0.6	2.2
El Castillo	21A	Mousterian	*Bison priscus*	LLM3	CAS142	6.1	15	-11.2	-11.7	-10.9	0.7	0.2	-5.0	-5.7	-4.4	0.4	1.3
El Castillo	21A	Mousterian	*Equus* sp.	LLM3	CAS143	6.5	17	-12.6	-12.9	-12.5	0.4	0.1	-6.2	-7.2	-5.4	0.5	1.8
El Castillo	18B	Transitional Aurignacian	*Bos/Bison* sp.	ULM2	CAS132	6.2	13	-11.3	-11.5	-10.9	0.6	0.2	-6.2	-7.4	-4.9	0.7	2.6
El Castillo	18B	Transitional Aurignacian	*Bos/Bison* sp.	ULM2	CAS133	6.8	18	-10.9	-11.6	-10.5	1.1	0.3	-5.4	-6.5	-4.2	0.7	2.2
El Castillo	18B	Transitional Aurignacian	*Bos/Bison* sp.	ULM2	CAS134	6.6	18	-12.4	-12.8	-11.6	1.2	0.3	-5.4	-6.3	-4.5	0.5	1.8
El Castillo	18C	Transitional Aurignacian	*Bos/Bison* sp.	LLM3	CAS135	6	17	-11.3	-11.5	-11.0	0.5	0.2	-6.1	-6.6	-5.5	0.3	1.1
El Castillo	18C	Transitional Aurignacian	*Bos/Bison* sp.	LLM3	CAS136	5.8	17	-12.0	-12.5	-11.7	0.9	0.2	-5.8	-6.7	-5.0	0.6	1.7
El Castillo	18C	Transitional Aurignacian	*Bos/Bison* sp.	LLM3	CAS137	6.6	14	-10.2	-10.6	-9.9	0.7	0.2	-5.8	-6.5	-4.1	0.7	2.4
El Castillo	18C	Transitional Aurignacian	*Bos/Bison* sp.	LLM3	CAS138	6.1	18	-11.6	-11.8	-11.4	0.4	0.1	-5.3	-5.9	-4.8	0.3	1.2
El Castillo	18B	Transitional Aurignacian	*Cervus elaphus*	ULM2+ULM3	CAS8		11	-13.0	-14.9	-12.1	2.8	1.0	-6.8	-10.4	-4.1	2.1	6.3
El Castillo	18B	Transitional Aurignacian	*Equus* sp.	ULP3/ULP4	CAS58		19	-11.7	-11.8	-11.5	0.3	0.1	-6.6	-7.5	-5.6	0.5	1.8
El Castillo	18B	Transitional Aurignacian	*Equus* sp.	LLP3/LLP3	CAS59		14	-11.5	-11.7	-11.0	0.7	0.2	-4.0	-4.7	-3.5	0.4	1.2
Labeko Koba	IX inf	Chatelperronian	*Equus* sp.	URM3	LAB38		17	-12.0	-12.2	-11.9	0.3	0.1	-6.6	-7.7	-5.9	0.5	1.9
Labeko Koba	IX inf	Chatelperronian	*Cervus elaphus*	LLM2	LAB02		7	-12.3	-12.4	-12.1	0.3	0.1	-4.7	-6.0	-3.7	1.0	2.3
Labeko Koba	VI	Aurignacian	*Equus* sp.	URM2	LAB20		16	-12.0	-12.2	-11.8	0.4	0.1	-5.3	-6.1	-4.4	0.6	1.7
Labeko Koba	V	Aurignacian	*Equus* sp.	LRM3	LAB42		17	-11.9	-12.3	-11.5	0.2	0.7	-5.7	-6.6	-5.0	0.5	1.6
Labeko Koba	IV	Aurignacian	*Equus* sp.	LRM2	LAB36		17	-11.6	-11.8	-11.3	0.6	0.2	-5.9	-6.2	-5.5	0.2	0.7
Canyars	I	Aurignacian	*Equus* sp.	URM3	CAN01	7.8	12	-10.0	-10.4	-9.5	0.9	0.3	-4.8	-5.3	-4.3	0.3	1.1
Canyars	I	Aurignacian	*Equus ferus*	URM3	CAN02	6.2	17	-10.5	-10.7	-10.3	0.4	0.1	-4.4	-5.0	-3.6	0.5	1.4
Canyars	I	Aurignacian	*Equus ferus*	URP3/URP4	CAN03	6.4	17	-10.7	-11.2	-10.4	0.8	0.2	-4.8	-5.3	-4.0	0.4	1.4
Labeko Koba	VII	Aurignacian	*Bos primigenius*	LRM3	LAB53	5.2	23	-9.5	-10.1	-8.7	1.4	0.3	-5.7	-7.0	-4.2	0.9	2.8
Labeko Koba	VII	Aurignacian	*Bos primigenius*	LRM3	LAB55	5.6	23	-10.4	-11.5	-9.8	1.6	0.3	-5.1	-7.0	-2.7	1.2	4.3
Labeko Koba	VII	Aurignacian	*Bos/Bison* sp.	LRM3	LAB62	6.5	21	-9.7	-10.2	-9.1	1.2	0.3	-7.2	-8.1	-6.2	0.6	2.0
Labeko Koba	V	Aurignacian	*Bos primigenius*	LRM3	LAB69	5.5	21	-9.3	-10.3	-7.3	3.0	0.9	-7.2	-8.8	-5.5	0.9	3.3
Canyars	I	Aurignacian	*Bos primigenius*	ULM3	CAN04	6.8	14	-9.3	-9.8	-8.7	1.1	0.3	-3.6	-4.2	-2.6	0.5	1.6
Canyars	I	Aurignacian	Bos primigenius	ULM3	CAN05	6.6	14	-9.0	-9.5	-8.5	0.9	0.3	-5.5	-6.2	-5.0	0.4	1.2
Aitzbitarte III	V (int)	Gravettian	*Bos/Bison* sp.	LLM3	AITI10	5.5	17	-9.2	-9.6	-8.7	0.9	0.3	-5.5	-6.5	-4.3	0.5	2.2
El Otero	IV	Magdalenian	*Cervus elaphus*	LLM2+LLM3	OTE1		11	-11.4	-11.6	-11.2	0.4	0.1	-4.4	-5.8	-2.9	1.0	2.9
El Otero	IV	Magdalenian	*Cervus elaphus*	LLM2+LLM3	OTE5		10	-11.3	-11.5	-11.0	0.5	0.2	-5.1	-5.7	-3.8	0.6	1.9
El Otero	IV	Magdalenian	*Cervus elaphus*	LLM2+LLM3	OTE6		14	-11.4	-11.8	-10.6	1.2	0.3	-4.6	-5.4	-4.0	0.4	1.4
El Otero	IV	Magdalenian	*Equus* sp.	LLP3/LLP4	OTE11		17	-11.6	-11.8	-11.4	0.5	0.1	-5.0	-6.3	-3.9	0.7	2.4
El Otero	IV	Magdalenian	*Equus* sp.	LLP3/LLP4	OTE12		16	-11.3	-11.5	-10.9	0.6	0.1	-3.9	-4.9	-3.3	0.6	1.6

Table 2. Mean, maximum value (Max), minimum value (Min), and standard deviation (SD) of $\delta^{13}C$ and $\delta^{18}O$ values per archaeological site and level organised by cultural periods. CCE, calcium carbonate equivalent; n, number of intratooth subsamples measured. In tooth type: position (U, upper; L, lower); laterality (R, right; L, left); tooth (M, molar; P, premolar).

4. Results

In northwestern Iberia, specifically in the Vasco-Cantabrian region, the mean $\delta^{13}C_{carb}$ values range from -13‰ to -8.9‰, with a mean value of -11‰ (SD = 1.2‰) (Table 2; Table 3). Considering species' different enrichment factors, the $\delta^{13}C_{carb}$ were transformed in $\delta^{13}C_{diet}$, resulting in mean values that extend from -27‰ to -23.5‰ (Fig. 4). It must be considered that average values may reflect slightly different periods or be affected by seasonal bias because different teeth encompass diverse periods, but it has been verified in our teeth that the variations are limited when the seasonal information of the sequential sampling is incorporated (±0.2; Appendix B). The carbon isotopic composition varies between species. The bovines have generally higher mean $\delta^{13}C_{carb}$ (from -12.4‰ to -8.9‰) than the horses (from -12.6‰ to -11.3‰), whereas the red deer fall within the horses' range (from -13‰ to -11.3‰). Average values of $\delta^{18}O_{carb}$ in all Vasco-Cantabrian individuals extend between -7.2‰ and -3.3‰ (mean = -5.5‰; SD = 0.8‰). When transformed to $\delta^{18}O$ expected from meteoric waters ($\delta^{18}O_{mw}$), with species-adapted correlations, the $\delta^{18}O_{mw}$ values range from -10.6‰ to -5.5‰. Less clear patterns in $\delta^{18}O_{carb}$ are observed between bovines and horses, with mean values of -5.7‰ and -5.2‰, respectively. In northeastern Iberia, the site of Canyars, both species have relatively high $\delta^{18}O_{carb}$ values that fall inside the range of variation observed in the Cantabria region, between -5.5‰ and -3.6‰ in bovines and between -4.8‰ and -4.4‰ in horses.

		Vasco-Cantabrian region (NW Iberia)				Northeastern Iberia			
		δ^{13}Ccarb VPDB (‰)	δ^{13}Cdiet VPDB (‰)	δ^{18}Ocarb VPDB (‰)	δ^{18}Omw VSMOW (‰)	δ^{13}Ccarb VPDB (‰)	δ^{13}Cdiet VPDB (‰)	δ^{18}Ocarb VPDB (‰)	δ^{18}Omw VSMOW (‰)
Total	Mean	-11.0	-25.1	-5.5	-8.0	-9.9	-24.0	-4.6	-7.1
	Max	-8.9	-23.5	-3.3	-5.5	-9.0	-23.6	-3.6	-5.0
	Min	-13.0	-27.0	-7.2	-10.6	-10.7	-24.4	-5.5	-7.9
	Range	4.1	3.5	3.9	5.1	1.7	0.8	1.9	2.9
	SD	1.2	0.9	0.8	1.2	0.8	0.3	0.7	1.2
Bovines	Mean	-10.4	-25.0	-5.7	-7.7	-9.1	-23.7	-4.5	-6.2
	Max	-8.9	-23.5	-4.8	-6.5	-9.0	-23.6	-3.6	-5.0
	Min	-12.4	-27.0	-7.2	-9.5	-9.3	-23.9	-5.5	-7.4
	Range	3.5	3.5	2.4	3.0	0.3	0.3	1.9	2.4
	SD	1.1	1.1	0.6	0.7	0.2	0.2	1.4	1.7
Horses	Mean	-11.8	-25.5	-5.2	-8.5	-10.4	-24.1	-4.7	-7.6
	Max	-11.3	-25.0	-3.3	-5.5	-10.0	-23.7	-4.4	-7.2
	Min	-12.6	-26.3	-6.6	-10.6	-10.7	-24.4	-4.8	-7.9
	Range	1.4	1.4	3.3	5.1	0.7	0.7	0.5	0.7
	SD	0.4	0.4	1.1	1.8	0.3	0.3	0.3	0.4

Table 3. Mean δ^{13}C from enamel carbonate ($\delta^{13}C_{carb}$) and diet ($\delta^{13}C_{diet}$), and δ^{18}O from enamel carbonate ($\delta^{18}O_{carb}$) and meteoric waters ($\delta^{18}O_{mw}$), by species on the Vasco-Cantabrian and northeastern Iberia areas. Max: maximum value; Min: minimum value; SD: standard deviation.

4.1 Axlor (Mousterian, ca. 80 ka BP - 50 ka cal BP)

A total of seven bovine teeth were included from levels III (n = 4), IV (n = 1), and VI (n = 2) of Axlor cave (Pederzani et al., 2023). The mean $\delta^{13}C_{carb}$ range from -9.9‰ to -8.9‰ ($\delta^{13}C_{diet}$= -24.5‰ to -23.5‰); whereas mean $\delta^{18}O_{carb}$ values are between -6.2‰ and -4.8‰ ($\delta^{18}O_{mw}$ = -8.3‰ and -6.5‰), indicating a range of variation around 1‰ and 1.4‰, respectively (Fig. 3; 4). Considering isotopic compositions by levels, mean $\delta^{13}C_{carb}$ decreases from level III to level IV, whereas mean $\delta^{18}O_{carb}$ remains stable through the sequence (Table 2; Appendix B). A range between 0.3‰ and 0.5‰ is observed in $\delta^{13}C_{carb}$ variation within tooth profiles. Individuals show clear δ^{18}O sinusoidal profiles, with peaks and troughs and intratooth ranges from 2.1‰ to 3.4‰. The $\delta^{18}O_{mw}$ after inverse modelling intratooth profiles range from -9.1‰ to -7.35‰ (Appendix D; E). Mean Annual Temperatures (MATs) oscillated between 9.1ºC and 12.6ºC (MATAs = -3.1/+0.4ºC) (Table 4). From sinusoidal profiles, summer temperatures were extracted from peaks, resulting from 15.4ºC to 23.7ºC, and winter temperatures from troughs provided values ranging from -7ºC to 10.8ºC. Mean Annual Precipitation (MAPs), extracted from $\delta^{13}C_{carb}$, extend between 204mm and 326mm (MAPAs = -843/-721mm). Based on these estimations, a non-clear climatic trend is observed through these levels.

4.2 El Castillo (Mousterian and Transitional Aurignacian, 62.5 ka BP – 46.4 ka cal BP)

From El Castillo, this work includes bovines (n = 11), horses (n = 5), and red deer (n = 1) teeth from the Mousterian (21 and 20E) and the Transitional Aurignacian levels (18B and 18C). The mean $\delta^{13}C_{carb}$ values are lower for horses, bovines, and red deer (-13‰ to -10.2‰) than other sites. Between -12.4‰ and -10.2‰ for bovines ($\delta^{13}C_{diet}$ = -24.6‰ to -25.8‰) and between -12.6‰ and -11.5‰ for horses ($\delta^{13}C_{diet}$ = -26.3‰ to -25.2‰) (Fig. 3). The mean $\delta^{18}O_{carb}$ values extend from -6.8‰ and -3.3‰. Horses and bovines overlap in their isotopic niche (Fig. 4), mainly due to the notably lower $\delta^{13}C_{carb}$ reported by bovines. The mean $\delta^{13}C_{carb}$ (-13‰) of the single red deer tooth is inside the variation range of bovines and horses but with a lower $\delta^{18}O_{carb}$ mean value (-6.8‰). Considering these isotopic compositions by levels, bovine mean $\delta^{13}C_{diet}$ values highly increase the variation range from Mousterian levels (20E and 21A) to Transitional Aurignacian levels (18C and 18B). In contrast, horses increase mean $\delta^{13}C_{diet}$ values (Fig. 5). Bovine mean $\delta^{18}O_{mw}$ values decrease from level 21A to level 18B, while horses from 18B have a large intra-level amplitude.

The mean $\delta^{18}O_{carb}$ values from horses have a more significant variation (range = 3.3‰) than bovines (range = 2.2‰). All individuals show flat $\delta^{13}C_{carb}$ intratooth profiles (<0.4‰), except for red deer (1‰) (Appendix D). Intratooth $\delta^{18}O_{carb}$ ranges of individuals are around 1-2‰ for horses and 1-3‰ for bovines. Some of the individuals analyzed do not show non-complete annual cycles. No precise $\delta^{18}O_{carb}$ sinusoidal profiles are

detected in three teeth; the other six have particularly unclear profiles. After modelling, individual $\delta^{18}O_{carb}$ ranges oscillated between 2.7‰ and 7.4‰ (Appendix E). MATs oscillated between 4.6ºC and 12.6ºC (MATAs = -8.8ºC/-0.9ºC), with mean summer temperatures from around 20.5ºC and mean winter temperatures around -1.1ºC. MAPs extend between 376mm and 784mm (MAPAs = -656/-248mm) (Table 4). Non-important differences in rainfall estimations based on bovines and equids are noticed, probably because they feed on similar ecological resources. Diachronic trends are unclear along the sequence but mean annual and winter temperatures from levels 18C and 18C seem slightly lower. MAPs estimations oscillated more in the upper levels.

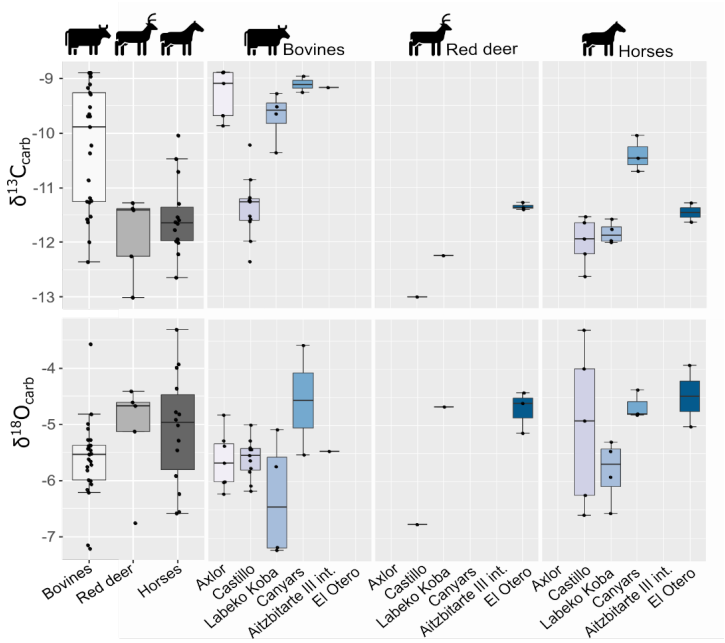

Figure 3. Distribution of mean carbon ($\delta^{13}C_{carb}$) and oxygen ($\delta^{18}O_{carb}$) isotopic values of enamel carbonate by species and archaeological site.

Figure 4. Biplot crossing $\delta^{13}C$ from diet ($\delta^{13}C_{diet}$) and $\delta^{18}O$ from meteoric waters ($\delta^{18}O_{mw}$) by species and archaeological site.

4.3 Labeko Koba (Châtelperronian and Aurignacian, 45.1-36.3 ka cal BP)

This work includes bovines (n = 4), horses (n = 4), and red deer (n = 1) teeth from levels related to Châtelperronian (IXb inf), ProtoAurignacian (VII), and Aurignacian (VI, V, and IV). Significant differentiation in mean $\delta^{13}C_{carb}$ between bovines and horses is observed, with higher values between -9.3‰ and -10.4‰ in bovines ($\delta^{13}C_{diet}$ = -25‰ to -23.8‰) than equids, whose values extend from -12‰ to -11.6‰ ($\delta^{13}C_{diet}$ = -25.8‰ to -25.2‰) (Fig. 3;). These horses' values are within the ranges observed from this species in the region. Red deer have similar $\delta^{13}C_{carb}$ values to those of horses ($\delta^{13}C_{carb}$ = -12.3‰; $\delta^{13}C_{diet}$ = -25.5‰). Mean $\delta^{18}O_{carb}$ values are similar between species from -7.2‰ to -4.7‰ ($\delta^{18}O_{mw}$ = -8.5‰ to -6.1‰). However, bovines have a very high variation within mean $\delta^{18}O_{carb}$ values (2.1‰), also reflected in the intratooth profiles. These $\delta^{18}O$ values are lower than in other Vasco-Cantabrian sites, especially for two individuals in levels VII and V (Table 3). Differences in $\delta^{13}C_{diet}$ values between bovines and horses result in isotopic niche differentiation between both species (Fig. 4). The red deer niche is placed within the horses' niche. The evolution of niche over time cannot be evaluated by levels due to the limited sample. Considering the isotopic compositions by levels (Fig. 5), both bovines and horses experienced a slight increase in mean $\delta^{13}C_{diet}$ from levels IX inf to IV, from Châtelperronian to Aurignacian. Mean $\delta^{18}O_{mw}$ values of bovines decrease from VII to V, whereas horses increase from IXb ~~i~~inf to VI to decrease from VI to IV.

Variability of $\delta^{13}C_{carb}$ values in intratooth profiles is slightly higher (0.1-0.7‰), especially in bovines (0.3-0.9‰), with more oscillating profiles than generally flat profiles observed in horses and red deer (Appendix D; E). Intratooth profiles ranges of $\delta^{18}O_{carb}$ are also larger within bovines (2-4‰) than in horses (1-2‰). Inverse-modelled individual $\delta^{18}O_{carb}$ ranges oscillated between 5-8‰ and 2-4‰, respectively. Sinusoidal curves are observed in horses and bovines, but bovine profiles are noisier. The red deer has an extensive $\delta^{18}O_{carb}$ range (6.3‰) from summer peak to an incomplete winter thought. We detect an inverse relation between $\delta^{13}C_{carb}$ and $\delta^{18}O_{carb}$ in some points of these individual profiles. MATs oscillated between 5.2ºC and 11.4ºC (MATAs = -5.6/+1.1ºC), with summer temperatures from 14.5ºC to 27.3ºC and winter temperatures from 1.9ºC to -4.9ºC. MAPs extend between 248mm and 521mm, notably drier than nowadays (MAPAs = -

798/-525mm) (Table 4). Lower rainfall levels and higher seasonal amplitudes are recorded along the sequence, especially in samples from the ProtoAurignacian level VII. Relevant differences are noticed between MAPs estimated from bovines and equids, the first providing more arid conditions.

4.4 Aitzbitarte III interior (Gravettian, 27.9 ka cal BP)

A single bovine individual was analysed from Gravettian level V located in the inner part of the cave. It has a high mean $\delta^{13}C_{carb}$ (-9.2‰) considering the observed range in bovines from the Vasco-Cantabrian region, whereas the $\delta^{18}O_{carb}$ mean value (-5.5‰) is inside the common $\delta^{18}O_{carb}$ variation observed (Fig. 3). The mean $\delta^{13}C_{diet}$ value of -23.8‰ is comparable with Canyars and some individuals from Axlor but different from Labeko Koba and El Castillo individuals. The individual δ^{13}Ccarb fluctuation is slight (0.3‰) (Appendix D; E). These teeth show not quite sinusoidal profile shape in $\delta^{18}O_{carb}$, with an intratooth range of around 2.2‰. Climatic information is extracted but may be considered cautiously due to the profile shape and the limited sample size. From the inverse modelled mean $\delta^{18}O_{mw}$ value (-5.4‰), we estimate a MAT of 13ºC (MATA = -0.4ºC) with a summer temperature of 19.7ºC and winter temperature of -2.9ºC. The MAP estimation reached 235mm (-1127mm to nowadays) (Table 4).

4.5 El Otero (Magdalenian, ca. 17.3 ka cal BP)

Two equids and three cervids are included from level IV from El Otero, recently redate and chronologically related to the Magdalenian (Marín-Arroyo et al., 2018). The mean $\delta^{13}C_{carb}$ values are close, between -11.4‰ and -11.3‰ for red deer ($\delta^{13}C_{diet}$ = -24.4‰ and -24.6‰) and -11.6‰ and -11.3‰ for horse ($\delta^{13}C_{diet}$ = -25.3‰ and -25.3‰) (Fig. 3). These $\delta^{13}C$ values for both species are relatively high concerning other studied samples, especially for cervids (around +1-2‰). Both species have higher $\delta^{18}O_{carb}$ values concerning the common range of variation observed in the Vasco-Cantabria region, between -5‰ and -3.9‰ for horses and between -5.1‰ and -4.4‰ for red deer. When values are transformed to $\delta^{13}C_{diet}$ and $\delta^{18}O_{mw}$, equids and cervids isotopic niches are separated (Fig. 4). All individuals show low amplitude $\delta^{13}C_{carb}$ intratooth profiles (<0.3‰), but especially equids with an intratooth variation around 0.1‰ (Appendix D; E). Equids and cervids show $\delta^{18}O_{carb}$ sinusoidal profiles, with intratooth ranges between 1.4‰ and 2.4‰. Climatic estimations are proposed only for equids, providing MATs estimations from 8.8ºC to 12.6ºC (MATAs = -4.9/-1ºC) and MAP between 400mm and 456mm (MAPAs = -755/-699mm) (Table 4). A high-temperature seasonality can be seen, with summer temperatures between 19.7ºC and 23.8ºC and winter temperatures from -10.4ºC to -3.1ºC.

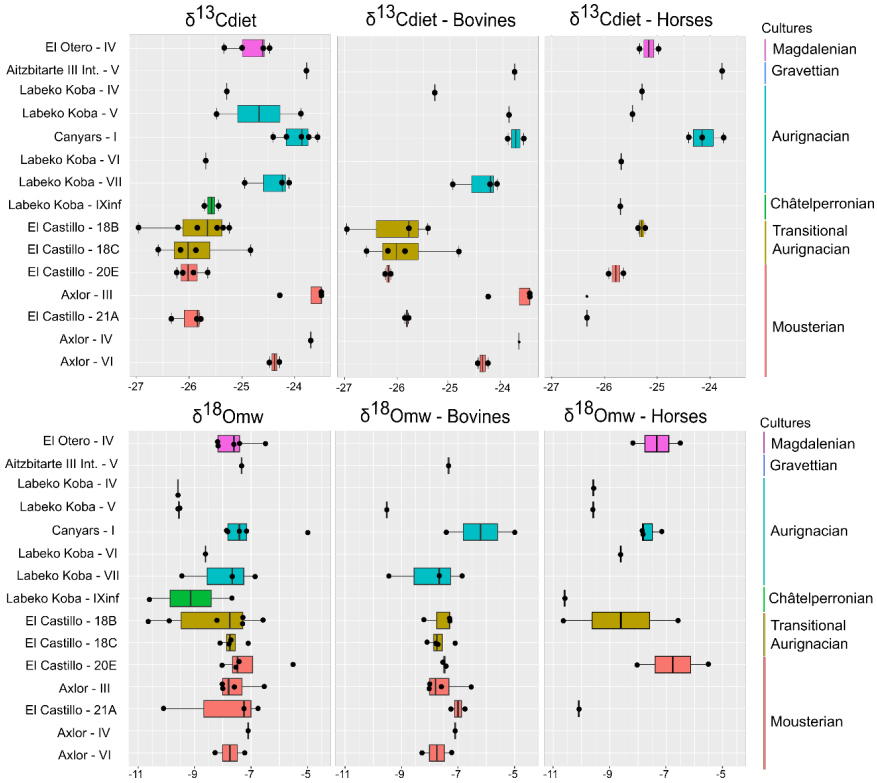

Figure 5. Evolution of $\delta^{13}C$ in diet ($\delta^{13}C_{diet}$) and $\delta^{18}O$ in meteoric waters ($\delta^{18}O_{mw}$) by archaeological levels in a diachronic order. From right to left: all species, including cervids, bovines and horses. Colours correspond to different chrono-cultures.

4.6 Canyars (Aurignacian, 39.7 ka cal BP)

From the archaeological level I at Canyars, corresponding to the Aurignacian, this work includes bovines (n = 2) and equids (n = 3) teeth. The mean $\delta^{13}C_{carb}$ values for bovines are between -9‰ to -9.3‰ ($\delta^{13}C_{diet}$ = -23.6‰ and -23.8‰), and for horses between -10‰ and -10.7‰ ($\delta^{13}C_{diet}$ = -23.7‰ and -24.4‰) (Fig.3). In this site, the $\delta^{13}C_{carb}$ values for horses are notably higher than in the Vasco-Cantabrian region (around +1-2‰) (Table 3). Both species have relatively high $\delta^{18}O_{carb}$ values, but they fall inside the range of variation observed in the Vasco-Cantabrian region, between -5.5‰ and -3.6‰ in bovines and between -4.8‰ and -4.4‰ in horses. Bovine and equid isotopic niches overlap (Fig. 4), but different responses are seen in mean $\delta^{18}O_{mw}$ values between the two bovines, with one high mean value but close $\delta^{13}C_{diet}$ mean values.

All individuals show flat $\delta^{13}C_{carb}$ intratooth profiles (<0.3‰ variation). Some individuals analysed do not show $\delta^{18}O_{carb}$ sinusoidal profiles, with intratooth profiles moderately flat and ranging from 1.1‰ to 1.6‰. We detect an inverse relation between $\delta^{13}C_{carb}$ and $\delta^{18}O_{carb}$ in some points of bovine individual isotopic profiles. MATs oscillated between 9.8ºC and 11.9ºC (MATAs = -5.4ºC/-3.3ºC), with summer temperatures from 16.3ºC to 27.5ºC and winter temperatures from -0.5ºC to 1.8ºC (Table 4). MAPs extend between 211mm and 316mm (MAPAs = -431/-326mm). No substantial differences are noticed in the estimations based on bovines and equids because mean $\delta^{13}C$ diet values differed relatively little.

				MAT (ºC)		Summer (ºC)		Winter (ºC)		Seasonality	MAP (mm)	
Site	**Sample**	**Level**	**Species**	**Estimated**	**Relative**	**Estimated**	**Relative**	**Estimated**	**Relative**	**(ºC)**	**Estimated**	**Relative**
Axlor	AXL59	III	*Bos/Bison* sp.	9.4	-2.8	17.6	-0.3	-3.9	-11.0	21.5	204	-843
	AXL60	III	*Bos/Bison* sp.	10.8	-1.4	22.7	4.7	4.8	-2.3	17.9	300	-747
	AXL65	III	*Bos/Bison* sp.	9.7	-2.5	22.7	4.8	-2.5	-9.6	25.2	204	-843
	AXL66	III	*Bos/Bison* sp.	12.6	0.4	22.8	4.8	-3.2	-10.3	26.0	204	-843
	AXL70	IV	*Bos/Bison* sp.	11.1	-1.1	21.9	3.9	-8.0	-15.1	29.9	227	-820
	AXL77	VI	*Bos/Bison* sp.	9.1	-3.1	20.4	2.5	-10.9	-17.9	31.3	300	-747
	AXL86	VI	*Bos/Bison* sp.	11.1	-1.1	25.9	8.0	3.1	-4.0	22.8	326	-721
El Castillo	CAS141	21A	*Bos/Bison* sp.	11.7	-1.7	24.2	5.6	-0.8	-9.9	25.1	546	-486
	CAS142	21A	*Bison priscus*	12.6	-0.9	19.6	1.0	3.1	-5.9	16.5	536	-496
	CAS143	21A	*Equus* sp.	5.7	-7.8	20.7	2.1	-5.6	-14.7	26.3	645	-387
	CAS60	20E	*Equus* sp.					1.6	-7.5		510	-522
	CAS61	20E	*Equus* sp.	9.7	-3.8	25.9	7.3	-4.1	-13.2	30.1	561	-471
	CAS139	20E	*Bos/Bison* sp.	11.2	-2.3	18.8	0.2	1.8	-7.3	17.0	622	-410
	CAS140	20E	*Bos/Bison* sp.	*11.3*	*-2.1*						602	-430
	CAS135	18C	*Bos/Bison* sp.			17.0	-1.6				551	-481
	CAS136	18C	*Bos/Bison* sp.	*10.6*	*-2.9*						699	-333
	CAS137	18C	*Bos/Bison* sp.					0.0	-9.1		376	-656
	CAS138	18C	*Bos/Bison* sp.	11.8	-1.7	18.3	-0.3	3.1	-6.0	15.3	612	-420
	CAS132	18B	*Bos/Bison* sp.	9.8	-3.6	26.3	7.6	-1.2	-10.3	27.5	548	-484
	CAS133	18B	*Bos/Bison* sp.					-0.1	-9.2		477	-555
	CAS134	18B	*Bos/Bison* sp.					0.8	-8.3		784	-248
	CAS58	18B	*Equus* sp.	4.6	-8.8	13.5	-5.1	-11.2	-20.3	24.7	460	-572
	CAS59	18B	*Equus* sp.	*13.0*	*-0.5*						440	-592
Labeko Koba	LAB38	IX inf	*Equus* sp.	5.2	-7.4	14.5	-4.1	-1.8	-9.1	16.2	521	-526
	LAB36	IV	*Equus* sp.	7.0	-5.6	16.3	-2.3	-2.4	-9.7	18.7	448	-599
	LAB42	V	*Equus* sp.	*7.6*	*-5.0*				-7.3		501	-546
	LAB69	V	*Bos primigenius*	6.3	-6.3	17.3	-1.2	-4.9	-12.2	22.2	248	-799
	LAB20	VI	*Equus* sp.	9.1	-3.5	15.7	-2.9	-0.9	-8.2	16.6	517	-530
	LAB53	VII	*Bos primigenius*	11.3	-1.3	27.3	8.7	-2.4	-9.7	29.7	278	-769
	LAB55	VII	*Bos primigenius*	11.4	-1.2	26.3	7.8	1.9	-5.4	24.4	397	-650
	LAB62	VII	*Bos/Bison* sp.	7.2	-5.4	20.6	2.1	-2.9	-10.2	23.5	295	-752
Canyars	CAN01	I	*Equus* sp.	9.8	-5.4	16.3	-5.9	1.7	-7.5	14.6	232	-410
	CAN02	I	*Equus ferus*	*11.9*	*-3.3*						284	-358
	CAN03	I	*Equus ferus*	10.4	-4.7	18.6	-3.6	-0.5	-9.7	19.1	316	-326
	CAN04	I	*Bos primigenius*	17.2	2.1	27.5	5.3				247	-395
	CAN05	I	*Bos primigenius*	11.3	-3.9	17.5	-4.7	1.8	-7.4	15.7	211	-431
Aitzbitarte III int	AITI10	V	*Bos/Bison* sp.	13.0	-0.4	19.7	0.7	-2.9	-11.4	22.6	235	-1127
Otero	OTE11	IV	*Equus* sp.	8.8	-4.9	19.7	0.9	-10.4	-19.8	30.1	456	-699
	OTE12	IV	*Equus* sp.	12.6	-1.0	23.8	5.0	-3.1	-12.5	26.8	400	-755

Table 4. Summary of paleoclimatic estimations, based on $\delta^{18}O$ for temperatures (Mean Annual Temperatures, MAT; summer; winter) and in $\delta^{13}C$ for precipitation (Mean Annual Precipitations, MAP). Summer and winter temperature estimations were obtained from teeth with clear seasonal profiles after modelling, while MAT was averaged between summer and winter before modelling. In profiles with an unclear seasonal shape, MAT was deduced from the original average of all teeth points (values marked in italics). Mean error associated to temperature estimations is 5.1±0.6 (see details in Appendix B). Seasonality is calculated as the temperature difference between summer and winter.

5. Discussion

5.1 Diet and ecological niches: carbon ratios

Carbon isotopic ratios are valuable indicators for discerning past animal diets, partially influenced by the physiology of the animal. Considering species trends in the studied sites, bovines have generally higher mean $\delta^{13}C_{carb}$ values (from -12.4‰ to-8.9‰) than horses (from -12.6‰ to -11.3‰), whereas the red deer fall within the horses' range (from -13‰-to 11.3‰). In the notheastern site of Canyars, bovines also show higher mean $\delta^{13}C_{carb}$ values (-9‰ to -9.3‰) compared to horses (-10.7‰ to -10‰). These differentiated isotopic ranges for equids and bovines can be potentially linked to feeding behaviour. Still, these species are expected to present different basal $\delta^{13}C_{carb}$ driven by their feeding behaviour and distinct physiological characteristics. Bovines, being ruminants, have been suggested in previous studies to exhibit higher $\delta^{13}C_{carb}$ values due to increased methane production (Cerling and Harris, 1999; Tejada-Lara et al., 2018). Therefore, transforming $\delta^{13}C_{carb}$ to $\delta^{13}C_{diet}$ values using species-specific equations is crucial to mitigate the species-specific impact, particularly when comparing ruminants and non-ruminants. Bovines report $\delta^{13}C_{diet}$ values between -27.5‰- and -23.5‰ and horses between -26‰ and -25‰. These carbon compositions are typical of animals feeding on C3 plants (commonly accepted range between -34‰ and -23‰), as can be expected

from high-latitude ecosystems during the Pleistocene (Bocherens, 2003; Cerling and Harris, 1999; Drucker, 2022).

Environmental factors such as light exposure, water stress, temperature fluctuations, salinity, and atmospheric CO_2 changes can influence variations in $\delta^{13}C$ values in a diet primarily based on C3 plants (Bocherens, 2003; Kohn, 2010). Typically, $\delta^{13}C_{diet}$ values below -27‰ ($\delta^{13}C_{carb}$ = -13‰) are associated with animals feeding on C3 vegetation found in closed forested environments, whereas $\delta^{13}C_{diet}$ values between -27‰ and -23‰ are linked to C3 open landscapes, which could include grasslands and steppe areas (Bocherens, 2003). The relatively high $\delta^{13}C_{diet}$ observed here points to animals predominantly feeding in open environments. The canopy effect, characterised by a depletion in ^{13}C isotopes due to dense tree cover, seems unlikely among the analysed samples since none of the individuals reported $\delta^{13}C_{diet}$ below the standard cut-off of -27‰ (Drucker et al., 2008; Kohn, 2010; van der Merwe, 1991). Therefore, in general terms, open mosaic landscapes, ranging from light forests to meadows and grasslands, can be inferred for northwestern Iberia. Given the generally higher $\delta^{13}C_{diet}$ values reported by bovines, it is likely that they were foraging in more open environments than horses and can be considered predominantly grazers. Particularly, bovines from El Castillo exhibit distinct feeding behaviour compared to other Vasco-Cantabrian sites, as evidenced by their lower $\delta^{13}C_{diet}$ values, indicating a potential preference for browsing and feeding in closer environments, possibly in lightly forested areas. Both extinct aurochs (*Bos primigenius*) and steppe bison (*Bison priscus*) are usually classified as grass-dominant mix-feeders during the Pleistocene, although it should be noted that modern European bison (*Bison bonasus*) could include browsing in their diet (Rivals et al., 2022). For aurochs, a browse-dominated mixed feeding behaviour is also frequently described.

The $\delta^{13}C_{diet}$ range in equids also indicates feeding in open environments, suggesting a general mixed-feeding pattern for the Vasco-Cantabrian region. However, individuals from northeastern Iberia are likely grazing in more open environments, as evidenced by their notably higher $\delta^{13}C_{diet}$ values compared to the Vasco-Cantabrian region (+1-2‰). Evaluating if other factors contribute to lower $\delta^{13}C_{diet}$ values in horses is critical. In the case of equid from the Vasco-Cantabrian region, it should be considered that they have been pretreated with a combination of NaClO and acetic acid, which could potentially affect the isotopic values. Samples after organic removal pretreatment can potentially show either higher or lower $\delta^{13}C$ values and higher $\delta^{18}O$ values based on previous experiments (Pellegrini and Snoeck, 2016; Snoeck and Pellegrini, 2015), with $\delta^{13}C$ values generally varying below 0.3‰. Based on the observation that horses in the Vasco-Cantabrian region present lower $\delta^{13}C_{carb}$ values compared to bovines but similar mean $\delta^{18}O_{carb}$ value ranges, the influence of the pre-treatment on our samples is deemed to be limited.

Furthermore, the high variability in $\delta^{18}O_{carb}$ values at El Castillo and Labeko Koba does not correlate with a significant variation in $\delta^{13}C_{carb}$ values. Based on dental wear and stable isotopes analysis, Middle and Late Pleistocene horses (*Equus ferus*) were primarily grazers, although some rare cases have been reported as mixed feeders or browsers, such as at Igue des Rameaux and Schöningen (Kuitems et al., 2015; Rivals et al., 2009, 2015; Uzunidis, 2020). Horse populations from northern and eastern Europe were found to be browsers or mixed feeders, while those from the Mediterranean region tend to be grazers (Rivals et al., 2022).

Finally, the few cervids included in this study exhibit $\delta^{13}C_{diet}$ values that frequently overlap with horses, indicating a mixed feeding behaviour that varies from more closed environments in El Castillo to more open habitats in El Otero. During the Pleistocene, the red deer (*Cervus elaphus*) exhibit a flexible, mixed-feeding behaviour, consuming leaves, shrubs, forbs, grass, and sedges, similar to their present-day counterparts (Merceron et al., 2021; Rivals et al., 2022). Today, this species inhabits diverse habitats ranging from steppes to closed temperate forests.

5.2 Seasonality, mobility and water acquisition: oxygen ratios and intratooth profiles

Average values of $\delta^{18}O_{carb}$ in Vasco-Cantabrian individuals extend between -7.2‰ and -3.3‰ (Table 3). Even if no clear species patterns in $\delta^{18}O_{carb}$ are observed, in general, bovines present slightly lower $\delta^{18}O_{carb}$ values from -7.2‰- to 4.8‰ than other species; horses have a significant variation from -6.6‰ to -3.3‰ and red deer from -6.8‰ to -4.4‰. In Canyars, both species have relatively high $\delta^{18}O_{carb}$ values that fall inside the variation range observed in the Vasco-Cantabrian region, between -5.5‰ and -3.6‰ in bovines and between -4.8‰ and -4.4‰ in horses. Each species shows different $\delta^{18}O_{carb}$ intratooth ranges, with bovines between 1‰ and 3‰, horses mostly around 1.5%, and red deer from 1‰ to 6‰ presenting the higher ranges (Table 3; Appendix D). After applying inverse modelling to correct the dampening effect (Passey et al., 2005b), the majority of teeth increase the $\delta^{18}O_{carb}$ intratooth range, between 3‰ and 8‰ for bovines and 2‰ and 7‰ for horses (Appendix E). Most bovines from Axlor and Labeko Koba and horses from El Castillo and El Otero exhibit well-defined sinusoidal profiles in their $\delta^{18}O_{carb}$ and large intratooth individual ranges, related to the predominant consumption of water sources that reflect seasonal fluctuations between summer and winter. Although not all samples consistently follow this pattern, specific intratooth profiles, particularly those from bovines in El Castillo and Canyars, exhibit sharp profiles with narrow ranges (<1.5‰). This phenomenon was previously reported in the region in preliminary studies conducted at the sites of El Castillo (Jones et al., 2019) and in the Magdalenian levels of El Mirón cave (Geiling, 2020).

Non-sinusoidal profiles observed in the data can be attributed to various factors, including sample techniques and preservation issues and the inherent variability in the original isotopic signal. Factors related to sampling and methods can be connected to 1) the sampling process (e.g. too deep or too distant sampling grooves); 2) the imprecision of the mass spectrometer measurements; 3) uncontrolled effects of samples pretreatments; 4) diagenetic alterations affecting the carbonate fraction. However, it must be noted that technical reasons, whether related to sampling or pretreatment, do not appear to impact the obtained results significantly. First, this study reproduces the same intratooth sampling methods that previously yielded reliable results in similar research (e.g., Pederzani et al., 2023, 2021a). Second, non-significant alterations in intratooth profiles of pretreated horse samples (El Castillo, Labeko Koba, Otero) are noticed in comparison to untreated bovid samples (Appendix D). Some bovid samples show these non-sinusoidal profiles equally. In sites where both species are analysed, no correlation is observed between $\delta^{18}O_{carb}$ and $\delta^{13}C_{carb}$. In tooth enamel, diagenetic alterations are generally less pronounced than in bone due to its higher mineral content. However, carbonates within tooth enamel can be more susceptible to diagenesis and recrystallisation compared to the phosphate fraction, which contains a more extensive reservoir of oxygen and stronger oxygen bonds (Zazzo et al., 2004; Chenery et al., 2012; Bryant et al., 1996). The carbonate content in our samples, ranging from 3.9% to 8.9%, is similar to the proportion found in modern tooth enamel, suggesting no immediate indication of diagenetic alteration. Diagenesis can also be evaluated by comparing the isotopic values of the carbonate and phosphate fractions in a sample, as there is a predictable difference between them. However, phosphate fraction measurements were still unavailable in our study, except at Axlor (Pederzani et al., 2023) where good preservation was attested. Additionally, in the case of diagenetic alteration, we would expect specimens from the same archaeological levels to be affected similarly, which is not the case.

Based on these arguments, it is suggested that the non-sinusoidal $\delta^{18}O_{carb}$ signal observed in some individuals may not be attributed to poor preservation; instead, it likely reflects the original isotopic signature from water input, which appears to be non-seasonal. Several factors can explain why some teeth do not reflect an evident seasonal fluctuation, which could be related to animals' mobility, the isotopic composition of the water sources, and seasonal buffering within those water sources (Pederzani and Britton, 2019). The main factors considered in our study are 1) the high mobility of the animals analysed among ecosystems with different isotopic baselines due to large migrations; 2) the inland-coastal or short altitudinal movements through the region, which lead to the acquisition of water from sources with different isotopic signal; and 3) the acquisition of water from sources with no clear seasonal signal, such as large bodies of water, rivers,

groundwaters, or meltwaters. At mid-latitudes, the temperature effect is currently the dominant factor. However, it is crucial to note that past changes in rainfall density (as the "amount effect"; Dansgaard, 1964) cannot be dismissed from having a more significant role then, particularly during glacial and arid periods. These effects, with their potential to mask temperature oscillations, underscore the urgency and importance of our research in understanding and predicting climate patterns. Furthermore, variability between species and within the same species, even within populations living in the same habitat, is also possible. This can be attributed to multiple factors, from minor differences in foraging and drinking behaviour to slight metabolic and physiological variations, including body size, metabolic rate, breathing rate, moisture content of food, and faeces, among others (Hoppe et al., 2004; Kohn, 1996; Magozzi et al., 2019).

Analyses of nitrogen and sulphur stable isotopes on ungulate bone collagen from Axlor, El Castillo and Labeko Koba (Jones et al., 2018, 2019; Pederzani et al., 2023) have already revealed large variation ranges linked to the existence of several microenvironments just in a few kilometres within the Vasco-Cantabria region. Long migrations and long hunting distances cannot solely explain these diverse values because of the range of species involved and their likely small-scale movements. In our study, the minimal $\delta^{13}C_{carb}$ intratooth variation within individuals (<1‰) indicates limited seasonal changes in their feeding behaviour that influenced the carbon isotopic composition (Appendix D). Therefore, considering the diverse topography of the Vasco-Cantabrian, characterized by steep valleys connecting the Cantabrian Cordillera with the Atlantic Ocean through rivers over short distances (30-50 km), the availability in the past of a wide range of water sources in small areas seems highly likely. Certain drinking behaviours can influence $\delta^{18}O$, as animals may acquire water from various sources, with small streams better reflecting seasonal isotopic oscillations than large lakes or evaporating ponds (see synthesis in Pederzani and Britton, 2019). Systematic consumption of highly buffered water sources can significantly attenuate the final recorded signal. Furthermore, rivers in the region frequently contain meltwater from snow during the winter-spring months and water springs.

5.3 Regional trends and ecological niches

This study provides valuable insights despite the limited sample size at each archaeological level. It establishes a baseline of isotopic values for northern Iberia, allowing for the evaluation of regional trends. In the northwest, in the Vasco-Cantabrian region, the $\delta^{13}C_{carb}$ values obtained oscillated between -13‰ and -8.9‰ and between -7.2‰ and -3.3‰ in the case of $\delta^{18}O_{carb}$ values. These values are within the range expected, considering previous regional studies in ungulates (Carvalho et al., 2022; Jones et al., 2019; Lécuyer et al., 2021; Pederzani et al., 2023). Although oxygen variability trends are less precise, the main factor distinguishing the observed changes over time is the variation of carbon isotopic composition among species and regions. The combination of mean $\delta^{13}C_{diet}$ and $\delta^{18}O_{mw}$ values (Fig. 4; 5) accentuates disparities in ecological niche overlap between horses and bovines, whereas cervids and horses frequently exhibit shared ecological niches. The dissimilarities between bovines and horses could be attributed to shifts in feeding behaviour, which may be accompanied by ecological and environmental changes, either independently or in parallel.

Comparing the entire dataset and across all sites, the consistently lower $\delta^{13}C_{diet}$ values in horses compared to bovids throughout time suggest both animals inhabited open landscapes, with bovines exhibiting a grazer preference while horses show a mix-feeding diet. Only in the Middle-to-Upper Paleolithic transition 18B and 18C levels of El Castillo, an exception is observed with lower $\delta^{13}C_{diet}$ values in bovines, linked to a higher browser input due to a higher habitat in closer environments, such as open forests, similar to those inhabited by the horses. This generates a niche overlapping between horses and bovines, most likely reflecting stable conditions that could support both species in similar ecosystems. Contrarily, in the Châtelperronian and early Aurignacian levels from Labeko Koba, a clear differentiation between horses and bovines is observed, mainly in $\delta^{13}C_{diet}$ values, highlighting the occupation of different parts of the landscape by both species. This

spatially-driven niche separation between species could result from resource competition derived from an unstable climatic period, where species needed to specialise to adapt to the changing conditions. Notable changes are also observed in the $\delta^{18}O_{carb}$ values from Labeko Koba compared to the older El Castillo and Axlor sites, with bovines exhibiting a higher fluctuation range and the lowest values in the region. These trends are consistent with values observed on bone collagen from previous studies in these sites. During the Middle-to-Upper Paleolithic transition in the region, by comparing horses and red deer, a decrease in mean $\delta^{13}C$ (from -21‰ to -20‰) and $\delta^{15}N$ values (from 2.5‰ to 6‰) in bone collagen was observed in contrast to stable red deer mean $\delta^{13}C$ (Fernández-García et al., 2023; Jones et al., 2018, 2019). This decrease was previously interpreted as niche fractionation, derived from an opening landscape, that drove equids into low-quality pastures compared to cervids. Pollen evidence in the region suggests a prevalence of steppe vegetation and low tree cover for the Châtelperronian and Aurignacian (Iriarte-Chiapusso, 2000).

In the same period, Canyars in the northeastern area, higher mean $\delta^{13}C_{diet}$ are observed in both species (between -23.6‰ and -24.4‰), indicating a preference for more open landscapes by bovines and equids. The indication of open areas could be linked to the arid climatic conditions associated with the Heinrich ~~Event~~ Stadial 4, which coincides with the formation of the studied level. This predominance of open areas coincides with the presence of typical steppe herbivore species, such as *Equus hydruntinus* and *Coelodonta antiquitatis*, the microfauna and pollen taxa, and the data offered by the use-wear analysis on ungulate remains identified at the site (Daura et al., 2013; López-García et al., 2022; Rivals et al., 2017).

Aridity is a plausible explanation for the higher niche partitioning observed in Labeko Koba and the higher $\delta^{13}C_{diet}$ values found in Canyars for both species during the Aurignacian. The $\delta^{13}C_{diet}$ results of bovines from Aitzbitarte III interior during the Gravettian are consistent with the trend observed in Labeko Koba, where previous studies have already suggested this time to be notably arid and cold (Arrizabalaga et al., 2010). Finally, in the Magdalenian level of El Otero, higher $\delta^{13}C_{diet}$ values resemble those observed in Canyars. However, this time, carbon values are related to niche partitioning between horses and red deer. In contrast, higher $\delta^{18}O_{mw}$ values might indicate warmer conditions but are still associated with open landscapes in the Vasco-Cantabrian area.

5.4 Late Pleistocene climatic evolution in Northern Iberia

Carbon and oxygen isotopes were used to estimate quantitative parameters related to past temperatures and precipitation. In the case of oxygen isotopic compositions, an evaluation of environmental water composition can be addressed before approaching temperature estimations. When transformed to $\delta^{18}O_{mw}$ using species-adapted correlations and correcting bias in sea water $\delta^{18}O_{mw}$, the summer $\delta^{18}O_{mw}$ values obtained from the modelled teeth range from -8.9‰ to -2.2‰, while the winter values range from -17.1‰ to -8.9‰. These values can be tentatively compared with the current trends observed in $\delta^{18}O_{mw}$ range recorded by the IAEA station (IAEA/ WMO, 2022) in Santander (from -3.5‰ in summer to -6.6‰ in winter) and in Barcelona (from -2.2‰ in summer to -6.3‰ in winter) and the OIPC (Bowen, 2022) estimations for studied locations (from -1‰ to -9‰) (Appendix B). As observed in the present, Canyars exhibit mean annual $\delta^{18}O_{mw}$ values around -8.2‰, which is lower than the current $\delta^{18}O_{mw}$ estimated for this location (-5.4‰) but higher than Labeko Koba mean annual $\delta^{18}O_{mw}$ (-9.5‰). This raises the question of whether the baseline $\delta^{18}O_{mw}$ differences between Canyars and the other sites can be attributed to Mediterranean influence rather than the Atlantic, assuming equivalent air circulation patterns and moisture sources experienced in the past as in the present (Araguas-Araguas and Diaz Teijeiro, 2005; García-Alix et al., 2021; Moreno et al., 2021). However, it's important to note that these comparisons must be approached thoughtfully, considering that moisture fluxes and precipitation trends may have varied significantly during the Pleistocene and the Holocene (Dansgaard, 1964; Shackleton, 1987).

As indicated by the climate reconstructed here, temperatures were colder, and precipitation levels were notably lower in the Late Pleistocene period in this region than they are nowadays (Table 4; Appendix B). From 80 to 50 ka BP, in the Mousterian levels of Axlor, temperatures were slightly colder than today, but older levels showed higher differences between summer and winter temperatures. Rainfall estimations exhibit an unusual arid pattern, possibly affected by bovines predominantly feeding in open areas at that time. This aligns with the impact of basal feeding behaviour on rainfall estimations, as previously advised by Lécuyer et al. (2021). In this case, it is not possible to isolate the effect of diet from environmental interference, but previous studies have highlighted stable climatic conditions at the site (Pederzani et al., 2023). Climatic reconstruction, relying on a compilation of lake sediments from northern Iberia (Moreno et al., 2012) suggests that from late MIS4 to 60 ka cal BP, cold but relatively humid conditions predominated, with drier conditions emerging later. Additionally, stalagmites from the Ejulve cave in the Iberian range indicate a dry climate until 65.5 ka BP, preceding HE6, followed by more humid conditions afterwards (Pérez-Mejías et al., 2019).

During the late Middle Paleolithic and early Aurignacian occupations, the observed shift in the niche configuration of species suggests potential climatic perturbations. There is a decreasing trend in temperatures from the Transitional Aurignacian levels in El Castillo (18C and 18B; ca. 47-46 ka cal BP) to the Châtelperronian (Xinf; 45.1 ka cal BP) and Early Aurignacian (VII-V; from 40.7 to 36.3 ka cal BP) levels in Labeko Koba. Lower mean annual and winter temperatures are particularly notable at El Castillo and Labeko Koba. Labeko Koba levels exhibit high seasonal amplitude, especially at level VII. Additionally, there is a slight decrease in rainfall and increased fluctuations from the Transitional Aurignacian levels from El Castillo (18B-18C) to the Aurignacian levels in Labeko Koba (VII-V). Previous studies in the northern Iberian region underlined an environmental and ecological shift after GS13/HE5, from 48 to 44 ka cal BP, based on a progressive trend to colder temperatures, aridity increase, and open environmental conditions, matching with the late Neanderthal occupations, followed by a population hiatus before the arrival of Anatomically Modern Humans (Fernández-García et al., 2023; Vidal-Cordasco et al., 2022). This episode coincides with the maximum extent of glaciers in this region, as recorded in Lake Enol and Vega Comeya and an significant decrease in plant biomass and herbivore abundance around 44 to 38 ka BP (Ballesteros et al., 2020; Jiménez-Sánchez et al., 2013; Ruiz-Fernández et al., 2022). Moreover, previous isotopic analyses in the region pointed to some ecological alterations considering perturbations observed in the $\delta^{13}C$ and $\delta^{15}N$ of bone collagen (Jones et al., 2018, 2019). This tendency of increased aridity aligns with observations made in regional lake sediments from northern Iberia between 60 and 23.5 ka cal BP, marked by abrupt climate changes associated with HE (Moreno et al., 2012). Supporting this, the marine core MD04-2845 in the northern margin of Iberia reveals a decline in the Atlantic forest and an expansion of steppe and cold grasses from 47 to 40 ka BP (Fourcade et al., 2022).

When comparing the environmental reconstruction of the Aurignacian period between the Vasco-Cantabrian (levels V-IV from Labeko Koba) and the northeastern region (Layer I from Canyars), which are synchronous to HE4 (39 ka BP), this study reveals notably lower rainfall levels for the latter. This is due to the feeding behaviour observed in animals, mainly in open areas. However, these drier conditions align with the specific climatic conditions expected for this period and support previous findings revealing aridity and the predominance of open landscapes (Daura et al., 2013; Rivals et al., 2017). The temperature data indicates that, at Canyars, colder conditions were experienced, especially during the winter season, compared to the present. However, in comparison to Labeko Koba, Canyars experienced warmer conditions. As explained earlier, the Mediterranean basin had consistently higher temperatures, even during colder periods. This is consistent with the persistence of Mediterranean open forests in the surroundings, as indicated by other studies (López-García et al., 2013; Rivals et al., 2017). Continuous natural records are lacking in the northeastern Iberian margin. However, the inland stalagmite record from Ejulve Cave (Pérez-Mejías et al., 2019) and the sedimentary lacustrine sequence of Cañizar de Villarquemado (González-Sampériz et al.,

2020) have identified the most arid intervals during HE5 and HE4. These periods were characterized by steppe vegetation expansions, followed by deciduous woodland expansion. To the south, the Padul sequence agrees with cold and dry conditions alternating with forest recovery (Camuera et al., 2019), as documented in the Alborean Sea (Martrat et al., 2004).

Finally, the sites Aitzbitarte III interior (27.9 ka cal BP) and El Otero (17.3 ka cal BP) provided valuable climatic insights into the Vasco-Cantabrian region during the Upper Paleolithic, specifically during the Gravettian and Magdalenian, respectively. Considering previous research in the region, the climatic trend reported for the Aurignacian, characterised by colder and more arid conditions, was expected to continue or even intensify during the Gravettian (Fernández-García et al., 2023; Garcia-Ibaibarriaga et al., 2019b; Lécuyer et al., 2021). Both sites indicate lower precipitation than today in this area, indicating significant aridity, with ungulates feeding predominantly in open landscapes. However, El Otero's higher mean annual temperatures recorded in the Magdalenian horses respect to other sites within the Vasco-Cantabrian, are consistent with a climatic amelioration following the Last Glacial Maximum (Jones et al., 2021). MIS 2 is marked by the most extreme glacial conditions, as indicated by NGRIP and marine cores in Iberian margins (Martrat et al., 2004; Sánchez Goñi et al., 2002). However, other regional proxies, such as lake sediment and the stalagmite sequence in Pindal Cave (Moreno et al., 2010), suggest a complex and highly variable climate during MIS 2. These proxies identify the coldest and most arid period within MIS 2 as the interval from 18 to 14 ka cal BP rather than the global Last Glacial Maximum (23 to 19 ka cal BP).

5. Conclusions

This study provides a detailed analysis of the temporal evolution of the environment and climatic conditions in northern Iberia, spanning from the Middle Paleolithic to the late Upper Paleolithic, this is from the GS21 to the GS2, ranging from 80 ka BP to 17 ka cal BP. In the Vasco-Cantabrian region, the results reveal a heterogeneous open mosaic landscape, ranging from light forest to meadows and grasslands. This landscape reconstruction is primarily inferred by the feeding locations of the studied animals and, consequently, related to the ecosystems where hominins captured them. Despite shifts in niche configuration observed between equids and bovines, both species typically foraging in open areas, with bovines showing a higher preference for grazing. Only in El Castillo, during the late Mousterian and the Transitional Aurignacian levels, bovines show unusually low $\delta^{13}C_{diet}$ related to higher browsing and overlapping with horse isotopic niche. This might indicate a slightly closed mosaic landscape that could sustain both species. In contrast, only horses from Canyars exhibit a preference for grazing behaviour.

Stable climatic conditions are described for Mousterian in Axlor and El Castillo levels from 80 to 50 ka cal BP. However, some elements indicate environmental perturbations initiated during the Transitional Aurignacian levels of El Castillo, around 48-45 ka BP and after HE5/GS13. After GS12 (44.2-43.3 ka BP), horses and bovines are potentially occupying different ecological niches during the Châtelperronian and early Aurignacian levels of Labeko Koba, pointing to a species' environmental specialisation, which can be a consequence of competition for food resources during an unstable ecological period. The climatic estimations indicate a temperature shift during this period, with a slight decrease in temperatures and evidence of fluctuations in rainfall. Previous environmental studies on the region have underlined ecological stress and increasing aridity from around 42.5 ka cal BP, which may relate to a broader ecosystem decline. When comparing the environmental conditions during the Aurignacian period in the northeast (Canyars) and the northwest (Labeko Koba), the first had higher baseline temperatures but also experienced higher aridity. Animals continued to feed on open landscapes during the Gravettian and Magdalenian levels in the Vasco-Cantabrian region, represented by Aitzbitarte III interior and El Otero. However, there is evidence of a temperature recovery after the LGM at the El Otero.

The results presented here, derived from the first extensive sampling in the Vasco-Cantabrian, establish the basis of future stable isotopic studies on faunal tooth enamel in Iberia. Despite the uncertainties inherent in this work, both $\delta^{18}O$ and $\delta^{13}C$ contributed to the regional climatic characterisation, including the estimation of temperatures and precipitations, as well as the seasonality range between summer and winter. The potential influence of pretreatment effects and uncontrolled diagenetic alterations on the enamel carbonate fraction has been assessed. However, complementary diagenetic tests, using new techniques like $\delta^{18}O_{phos}$ and FTIR analyses are advised in further works to gain more insights into sample preservation. Ongoing sulphur, hydrogen and strontium studies will provide additional information on the mobility patterns of animals hunted by Late Pleistocene hominins and, therefore, will help better understand the ecological and environmental context occupied by Neanderthal and modern humans and their landscape use in this particular region. Finally, a more comprehensive characterisation of the baseline oxygen values would also enhance the environmental interpretation of the existing data.

Appendices

Appendices A, C, D and E are presented after bibliography. Raw data is presented in Appendix B, available at https://github.com/ERC-Subsilience/Ungulate_enamel-carbonate

Field Code Changed

Code availability

R code used to perform plots, temperature and error calculations, Bayesian models code and inverse models in this manuscript can be accessed at GitHub (https://github.com/ERC-Subsilience/Ungulate_enamel-carbonate).

Field Code Changed

Data availability

The available datasets used for this article are provided in the supplementary materials (Appendix A-E).

Author contribution

A.B.M.-A. got the funding and designed the research. A.B.M.-A and M.F.-G. get the permissions for sampling in the regional museums. M.F.-G., K.B, and S.P. defined the analysis strategy. M.F.-G. analysed the data and wrote the manuscript with critical inputs from A.B.M.-A., K.B, and S.P. J.M.G., L.A., M.F.-G., and A.C. M.F.-G., L.A., J.M.G., and A.C. achieved the teeth sampling and lab sample preparation. J.D. and M.S. are responsible for the excavations in Canyars and contribute to the discussion. All the authors revised and commented on the manuscript.

Competing interests

The contact author has declared that none of the authors has any competing interests.

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Appendix A. Sites description

A1. Vasco-Cantabrian sites

Axlor (Dima, Vizcaya, País Vasco)

Axlor is a rock-shelter located in Dima (43.2706; -1.8905), with a continuous Middle Paleolithic sequence from the MIS5 to the MIS3 (DeMuro et al., 2023; Pederzani et al., 2023; Marín-Arroyo et al., 2018). It is placed on the southwestern slope of the Dima Valley, with an elevation of approximately 320 m above sea level (a.s.l.), at 33 km straight from the present-day coastline, next to one of the lowest mountain passes linking the Cantabrian basins and the Alavese Plateau. The site was discovered in 1932 and initial excavations were performed by Barandiarán (1967-1974). J. M. Barandiarán undertook the excavations between 1967 and 1974, identifying eight Mousterian levels (I-VIII) (Barandiarán, 1980).

From 2000 to 2008, new excavations by González-Urquijo, Ibáñez-Estévez and Rios-Garaizar were achieved and, since 2019, these are ongoing by González-Urquijo and Lazuén. Due to the lack of chronology during Barandiarán excavations, among other aspects, work was focused on obtaining a detailed stratigraphy on the new excavation areas to correlate it with Barandiarán's levels (González-Urquijo & Ibáñez-Estévez, 2021; González Urquijo et al., 2005). The new stratigraphic sequence is roughly equivalent to the previous one, but with additional levels not previously identified or excavated by Barandiarán. Some of these levels were deposited before Level VIII (Gómez-Olivencia et al., 2018; 2020). The Middle Paleolithic sequence extends from layers VIII to III (or from N to B-C). Levallois production is predominant in the lower levels (VI to VIII), while Quina Mousterian technocomplex does in the upper ones (from III to V) (Rios-Garaizar, 2012, 2017). Recent chronological data by radiocarbon (Pederzani et al., 2023; Marín-Arroyo et al., 2018) and OSL (Demuro et al., 2023) methods confirm that a sequence Axlor levels VI, VIII, and VIII probably accumulated during MIS5d–a (109–82 ka), while levels D to B probably were formed during the period encompassing the start of MIS 4 (71–57 ka) through to the beginning or middle of MIS 3 (57–29 ka) and upper Level III to 46,200 ±3,000 BP, which calibrates between 45,350 cal BP and beyond the calibration curve at > 55,000 cal BP.

The archaeozoological study indicates an anthropic origin of the faunal assemblage with scarce carnivore activity documented (Altuna, 1989; Castaños, 2005; Gómez-Olivencia et al., 2018). In lower layers, the most abundant taxa are *Cervus elaphus* (VIII) and *Capra pyrenaica* (VII), while in upper layers III-V, C*ervus elaphus* is substituted by *Bos primigenious/Bison priscus* and *Equus sp*. The material included in this work comes from the faunal collection of the Barandiarán excavation currently curated at the Bizkaia Museum of Archaeology (Bilbao), where teeth were sampled, and the stable isotope analyses on enamel phosphate were included in Pederzani et al. (2023).

El Castillo (Puente Viesgo, Cantabria)

El Castillo cave is located in Puente Viesgo (43.2924; -3.9656), with an elevation of approximately 195m a.s.l., at 17 km straight from the present-day coastline. The cave belongs to the karstic system that was formed in the Monte Castillo, which dominates the Pas Valley. The site was discovered in 1903 by H. Alcalde del Río. H. Obermaier carried out the first excavation seasons between 1910 and 1914 when many of the archaeological remains were recovered, mainly from the cave hall. These interventions were done under the supervision of the "Institut de Paléontologie Humaine" (IPH) and Prince Albert I of Monaco. From 1980 to 2011, V. Cabrera and F. Bernaldo de Quirós underwent new excavations focusing on the cave entrance, on the Middle to Upper Paleolithic transitional levels, mainly 16, 18 and 20 (Cabrera-Valdes, 1984). The site has yielded an important stratigraphic sequence, composed by 26 sedimentological units (1-26) related to

different anthropic occupational units, often separated by archaeologically sterile units: Eneolithic (2), Azilian (4), Magdalenian (6 and 8), Solutrean (10), Aurignacian (12, 14, 16 and 18), Mousterian (20, 21 and 22) and Acheulean (24) (Cabrera-Valdés, 1984).

Unit 21 is mostly sterile (Cabrera Valdés, 1984; Martín-Perea et al., 2023), and ESR dated it, yielding a mean date of 69,000 ± 9,200 years BP (Rink et al., 1997). However, Martín-Perea et al. (2023) suggested some dating uncertainty from interpreting the initial stratigraphic nomenclature. They suggest that the ESR dates provided for level 21 by Rink et al. (1997) were erroneously attributed to this unit and it might correspond to 20E, indicating that below that subunit, the chronology is older than 70,000 years BP (Martín-Perea et al., 2023). The Mousterian Unit 20 cave is divided into several subunits (Martín-Perea et al., 2023). In Unit 20, a cave roof collapse took place, transforming the cave system into an open rock shelter. This unit contains abundant archaeological and paleontological remains. Lithic industry consists of sidescrapers, denticulates, notches and cleavers, the majority on quartzite and presents both unifacial, bifacial discoid debitage and Levallois debitage. Unit 20E was attributed to Quina Mousterian by Sánchez-Fernández and Bernaldo De Quiros (2009) and contains a Neanderthal tooth (Garralda, 2005). Considering the geochronological uncertainties for dates on 20E related to Rink et al. (1997), we have decided to rely solely on ESR date of 47,000 ± 9400 BP provided by Liberda et al. (2010) for this level. Unit 20C presents clear evidence of the Mousterian lithic industry and radiocarbon dates of 48,700±3,400 uncal BP (OxA-22204) and 49,400±3,700 uncal BP (OxA-22205) (Wood et al., 2018) and mean ESR date of 42,700 ±9900 BP (Liberda et al., 2010). Level 19 is archaeologically sterile and separates Unit 20 from Unit 18 (Wood et al., 2018).

Unit 18 is divided into 18A (archaeologically sterile), 18B, and 18C. Levels 18B and 18C were classified as Transitional Aurignacian, representing a gradual transformation from the Mousterian to the Aurignacian, which is unique to El Castillo cave (Cabrera et al., 2001; Maíllo and Bernaldo de Quirós, 2010; Wood et al., 2018). These levels' dates and cultural attribution have been the subject of much debate (e.g. Zilhao and D'Errico, 2003; Wood et al., 2018). According to Wood et al. (2018), the last dates of these levels range between 42,000±1,500 uncal BP (OxA-22203) and 46,000±2,400 uncal BP (OxA-21973), which is much earlier than the start of the Aurignacian period in the Cantabrian region (Marín-Arroyo et al., 2018; Vidal-Cordasco et al., 2022). The lithic assemblage of Unit 18 appears to be dominated by Discoid/Levallois technology (Bernaldo de Quirós and Maíllo-Fernández, 2009) but with a high percentage of "Upper Paleolithic" pieces. Additionally, punctual bone industry and pieces with incisions and engravings were discovered in Unit 18 (Cabrera-Valdés et al., 2001). Three deciduous tooth crowns attributed to Neanderthals were found in Unit 18B (Garralda et al., 2022). Above, Unit 17 is sterile but contains scarce lithic and faunal materials, while Level 16 was attributed to the Proto-Aurignacian, with dates of 38,600±1,000 uncal BP (OxA-22200) (Wood et al., 2018).

According to Luret et al. (2020), there was a shift in hunting practices between the Late Mousterian (unit 20) and the Transitional Aurignacian (unit 18). During the Late Mousterian, hunting strategies were less specialized, and the species hunted included red deer, horses, and bovines. However, in Unit 18, a specialization in red deer hunting is observed. However, the explanation of this shift has been proposed as a response to a cultural choice or induced by climatic changes. However, recent taphonomic studies by Sanz-Royo et al. (2023) on the old collections of Aurignacian Delta level reveal a more significant role of carnivores than shown by Luret et al. (2020). The material included in this work comes from the faunal collection recovered during the Cabrera-Valdés and Bernaldo de Quirós excavations curated at Museo de Prehistoria y Arqueología de Cantabria (MUPAC, Santander).

Labeko Koba (Arrastre, Guipúzcoa, País Vasco)

Labeko Koba is a cave in the Kurtzetxiki Hill (43.0619; -2.4833), at 246 m a.s.l. and 29 km straight from the present-day Atlantic coast. In 1987 and 1988, the site was discovered due to the construction of the Arrasate ring road, and a savage excavation was carried out (Arrizabalaga, 2000a). Unfortunately, the site was destroyed after that. The stratigraphic sequence identified nine different levels. The lower Level IX was attributed to the Châtelperronian, based on the presence of three Châtelperron points. Although there is a lack of human remains in few Cantabrian Châtelperronian sites, recent research has suggested that this techno-complex was produced by Neanderthals (Maroto et al., 2012; Rios-Garaizar et al., 2022). Level VII marks the beginning of the Aurignacian sequence, likely Proto-Aurignacian, with a lithic assemblage dominated by Dufour bladelets (Arrizabalaga, 2000a). Levels VI, V, and IV contain lithic assemblages that suggested an Early Aurignacian attribution (Arrizabalaga, 2000b; Arrizabalaga et al., 2009). This site is significant because it is one of the few sites with Châtelperronian assemblages and with both Proto-Aurignacian and Early Aurignacian separated (Arrizabalaga et al., 2009).

Initial radiocarbon dates were inconsistent with the stratigraphy of the site and much more recent than expected for the Early Upper Paleolithic (Arrizabalaga, 2000a). This incoherence was determined to be affected by taphonomic alterations (Wood et al., 2014). Later radiocarbon dates undertaken with an ultrafiltration pre-treatment provided a new regional framework for the regional Early Upper Paleolithic (Wood et al., 2014). The Châtelperronian layer IX inf is dated to 38,100±900 uncal BP (OxA-22562) and 37,400±800 uncal BP (OxA-22560). The Proto-Aurignacian levels cover a period from 36,850±800 uncal BP (OxA-21766) to 35,250±650 uncal BP (OxA-21793). The three Early Aurignacian levels are dated to 35,100±600 uncal BP (OxA-21778) for level VI, ~ 34,000 uncal BP (OxA-21767 and OxA-21779) for level V, and ~ 33,000 BP (OxA-21768 and OxA-21780) for level IV (Arrizabalaga et al., 2009).

Taphonomic studies indicate an alternation in the use of the cave between carnivores and humans, the latter during short occupation periods (Villaluenda et al., 2012; Ríos-Garaizar et al., 2012; Arrizabalaga et al., 2010). Labeko Koba is considered to have functioned as a natural trap where carnivores, mainly hyenas, access animal carcasses. At least in the base of Labeko Koba IX, carnivore activity was higher, and they would have consumed the same prey as humans (Villaluenga et al., 2012). The presence of humans is linked to strategic use as a campsite associated with a small assemblage of lithic artifacts. The most consumed species by Châtelperronian groups were red deer, followed by the consumption of large bovids, equids, and woolly rhinoceros. During the Aurignacian period, there was some stability in human occupations, although they still alternated with carnivore occupations (Arrizabalaga et al., 2010). Cold-adapted fauna such as reindeer and woolly rhinoceros were identified in association with the Châtelperronian. Reindeer and the woolly mammoth and arctic fox were still present during the Aurignacian levels. The original sampling of the teeth studied by this work was performed in the San Sebastian Heritage Collection headquarters, where the Guipuzcoa archaeological materials were deposited at that time.

Aitzbitarte III interior (Rentería, Guipúzcoa, País Vasco)

Aitzbitarte III is an archaeological site located within the Landarbaso karstic system comprising nine caves (43.270; -1.8905). The cave is situated 220 m.a.s.l. and is 10 km away from the present-day coastline. Initial archaeological interventions were carried out at the end of the 19th century by P.M. de Soraluce (Altuna, 2011). Recent excavations were initially conducted in the deep zone inside the cave between 1986 and 1993, where the studied tooth was recovered, and later focused on the cave entrance between 1994 and 2002, by J. Altuna, K. Mariezkurrena, and J. Ríos-Garaizar (Altuna et al., 2011; 2017).

While the cave's entrance area contains a sequence comprising possible Mousterian and Evolved Aurignacian and Gravettian levels (Altuna et al., 2011; 2013), the stratigraphy in the inner cave presents eight levels: level VIII (some tools with Mousterian features), VII (sterile), VIb, VIa and V (Middle Gravettian

technocomplex with abundance of Noailles burins), IV-II (disturbed archaeological levels) and I (surface) (Altuna et al., 2017). Levels V have dates of 24,910 uncal BP (I-15208) and 23,230 uncal BP (Ua-2243); whereas level VI extends from 23,830 ± 345 uncal BP (Ua-2628) and 25,380± 430 uncal BP (Ua-2244) (Altuna, 1992; Altuna et al., 2017), with a possible outlier dated at 21,130 uncal BP (Ua-1917).

The Gravettian occupation in the inner part of the cave was initially thought to be more recent than the one in the cave entrance. However, it was not easy to correlate the two excavation areas due to different sedimentation rates. The abundant human occupations took place during a singular cold phase in the Middle Gravettian with a specialized paleoeconomy focused on the hunting of *Bos primigenius* and *Bison priscus* (85% in level VI and 68% in level V), which is unusual in the Cantabrian region mostly focused on red deer and ibex. Other ungulates present are *Cervus elaphus* and *Rupicapra rupicapra*, and to a lesser extent *Capra pyrenaica, Capreolus capreolus, Rangifer tarandus*, and *Equus ferus* (Altuna et al., 2017; Altuna & Mariezkurrena, 2020). There is a scarce representation of carnivores. The tooth studied was sampled at the Gordailua Center for Heritage Collections of the Provincial Council of Gipuzkoa.

El Otero (Secadura, Voto, Cantabria)

El Otero cave is located in Secadura (Voto) (43.3565; -3.5360), at 129 m.s.a.l and 12 km from the present-day coastline, near the Matienzo valley in a coastal plain environment covered by meadows and gentle hills. The discovery was made in 1908 by Lorenzo Sierra. The site was excavated in 1963 by J. Gonzalez Echegaray and M.A. García Guinea, in two different sectors (Sala I and Sala II) with an equivalent stratigraphic sequence (González Echegaray, 1966). Nine levels were identified in Sala I, from level IX to level I. Levels IX and VIII were initially related to the "Aurignacian-Mousterian, based on lithics assemblages with a combination of both technocomplex features. The overlying levels VI-IV were separated by a speleothem crust (level VII) and were initially related to Aurignacian, due to the presence of end-scrappers, bone points, blades, or burins on truncation (Freeman, 1964; Rios-Garaizar, 2013). Also, perforated deer, ibex, and fox teeth were found in levels V and IV. This site lacked chronological dating methods, until a selection of material from levels VI, V and IV revealed a difference in chrono-cultural attribution (Marín-Arroyo et al., 2018). Radiocarbon results yielded younger dates for such a cultural attribution and showed significant stratigraphic inconsistency. Level VI gave a result of 12,415±55 uncal BP (OxA-32585), two dates in Level V are 12,340±55 (OxA-32509) and 10,585±50 uncal BP (OxA-32510), and a date in Level IV is 15,990±80 uncal BP (OxA-32508). All these results fall into the range of the Late Upper Paleolithic (Magdalenian-Azilian initially identified in levels III-I), eliminating attribution of these levels to the Aurignacian despite the presence of apparently characteristic artefacts. Further assessments of archaeological materials will be needed.

Red deer dominate the assemblage, except for level IV where horses are more abundant. Wild boar, roe deer, and ibex are also present, but large bovids are relatively rare (González Echegaray, 1966). Level IV is the richest and most anthropogenic level, with evidence of butchering in red deer (captured in winter and early summer) and chamois (in autumn). The formation of this level involved humans and carnivores, and although certain data may suggest an anthropogenic predominance, the limited sample analyzed taphonomically and the pre-selection of preserved pieces do not allow for a definitive conclusion (Yravedra & Gómez-Castanedo, 2010). The material included in this work is curated at the Museo de Prehistoria y Arqueología de Cantabria (MUPAC, Santander).

A2. Northeastern Iberia sites

Terrasses de la Riera dels Canyars (Gavà, Barcelona, Cataluña)

Terrasses de la Riera dels Canyars (henceforth, Canyars) is an open-air site located near Gavà (Barcelona) (41.2961;1.9797), at 28 m.s.a.l and 3 km straight from the present-day coastline. The site lies on a fluvial terrace at the confluence of Riera dels Canyars, a torrential stream between Garraf Massif, Llobregat delta and Riera de Can Llong (Daura et al., 2013). Archaeo-paleontological remains were discovered during quarries activities in 2005 and was complete excavated on 2007 by the *Grup de Recerca del Quaternari* (Daura and Sanz, 2006; Daura et al., 2013). This intervention determined nine lithological units. The paleontological and archaeological remains come exclusively from one unit, the middle luthitic unit (MLU), and specifically from layer I. The MLU is composed of coarse sandy clays and gravels, filling a paleochannel network named lower detrital unit (LDU) (Daura et al., 2013). Five radiocarbon dates were obtained on charcoals from layer I, which yield statistically consistent ages from 33,800 ±350 uncal BP to 34,900 ±340 uncal BP, which results in mean age of 39,710 cal BP (from 40,890 to 38,530 cal BP) (Daura et al., 2013; this work).

The layer I of the site has yielded a rich faunal assemblage, consisting of over 5,000 remains. Among the herbivores, the most common species found are *Equus ferus*, *Bos primigenius, Equus hydruntinus*, and *Cervus elaphus* (Daura et al., 2013; Sanz-Royo et al., 2020). *Capra* sp. and *Sus scrofa are* also present, although in lower frequencies. The carnivores found at the site are also noteworthy, with *Crocuta crocuta* and *Lynx pardinus* being the most frequent. Presence of cold-adapted fauna associated to stepped environments is recorded, such as cf. *Mammuthus* sp., *Coelodonta antiquitatis*, and *Equus hydruntinus*. Small mammal analysis, pollen, and use-wear analysis have provided further evidence that a steppe-dominated landscape surrounded the Canyars site, supporting a correlation with the Heinrich Stadial~~Event~~ 4, in coherence with the chronology obtained for the layer (López-García et al. 2013; 2023; Rivals et al., 2017). However, the presence of woodland is also attested by forest taxa within charcoal and pollen assemblages (Daura et al., 2013).

Taphonomic study is ongoing. But several evidences point that hyenas have played an important role in the accumulation of the faunal assemblage (Daura et al., 2013; Jimenez et al. 2019). However, sporadic human presence is documented by few human modifications found in faunal remains (cutmarks and fire alterations). Although the paucity of the lithic assemblage in the site, it shows a clear attribution to Upper Palaeolithic technocomplex, most likely the Early Aurignacian (Daura et al., 2013). Recently, it was documented a perforated bone fragment, which has been identified as a perforated board for leather production (Doyon et al., 2023). All teeth included in this work were sampled in *Laboratori de la Guixera* (Ajuntament de Casteldefels) where the material is stored.

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Appendix C - Individual Bayesian Models

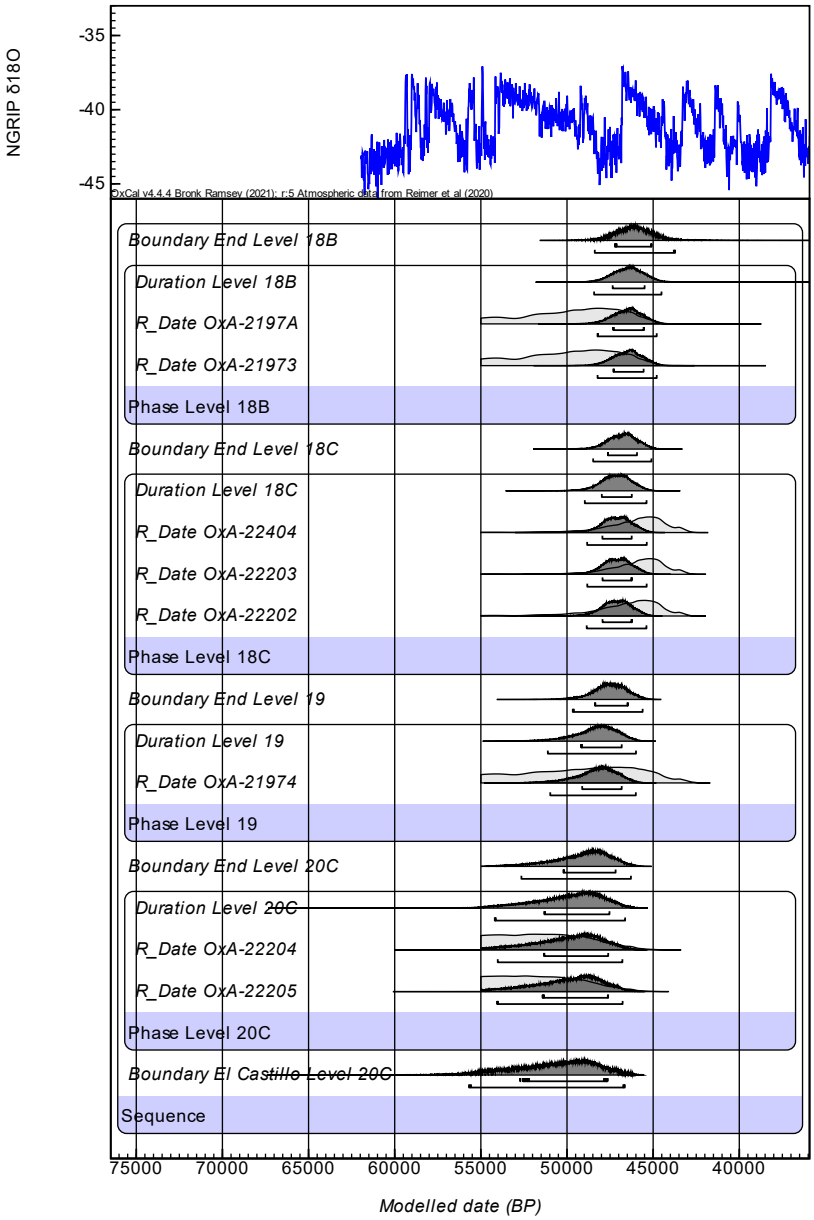

Figure C1. Radiocarbon dates from El Castillo modelled in OxCal4.4 against INTCAL20.

Field Code Changed

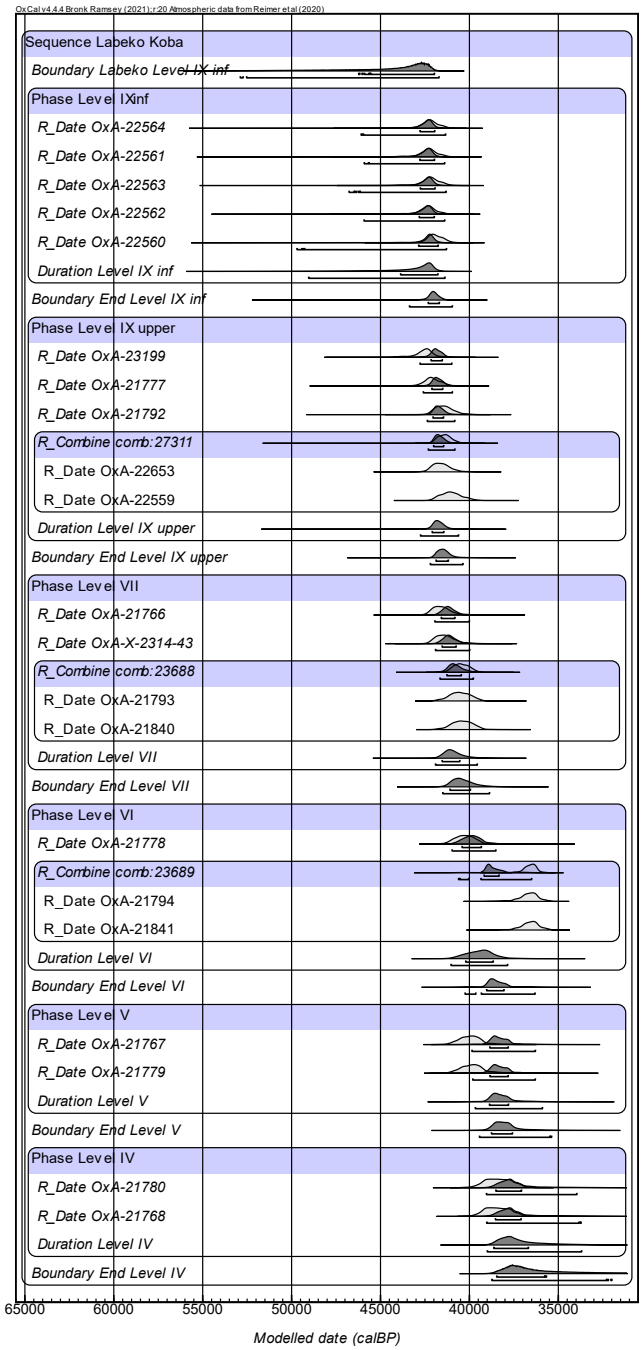

Figure C2. Radiocarbon dates from Labeko Koba modelled in OxCal4.4 against INTCAL20.

Field Code Changed

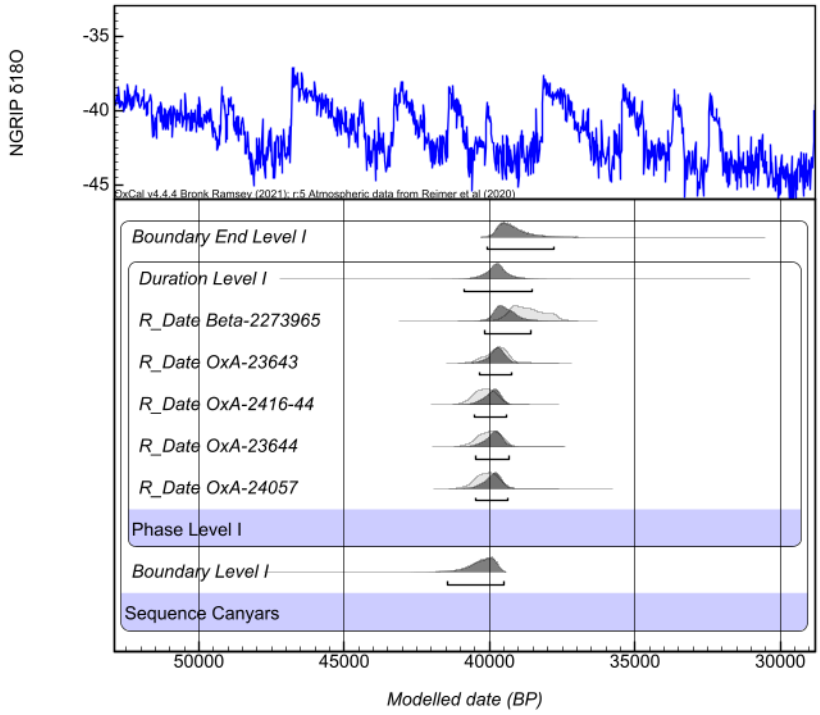

Figure C3. Radiocarbon dates from Canyars modelled in OxCal4.4 against INTCAL20.

Field Code Changed

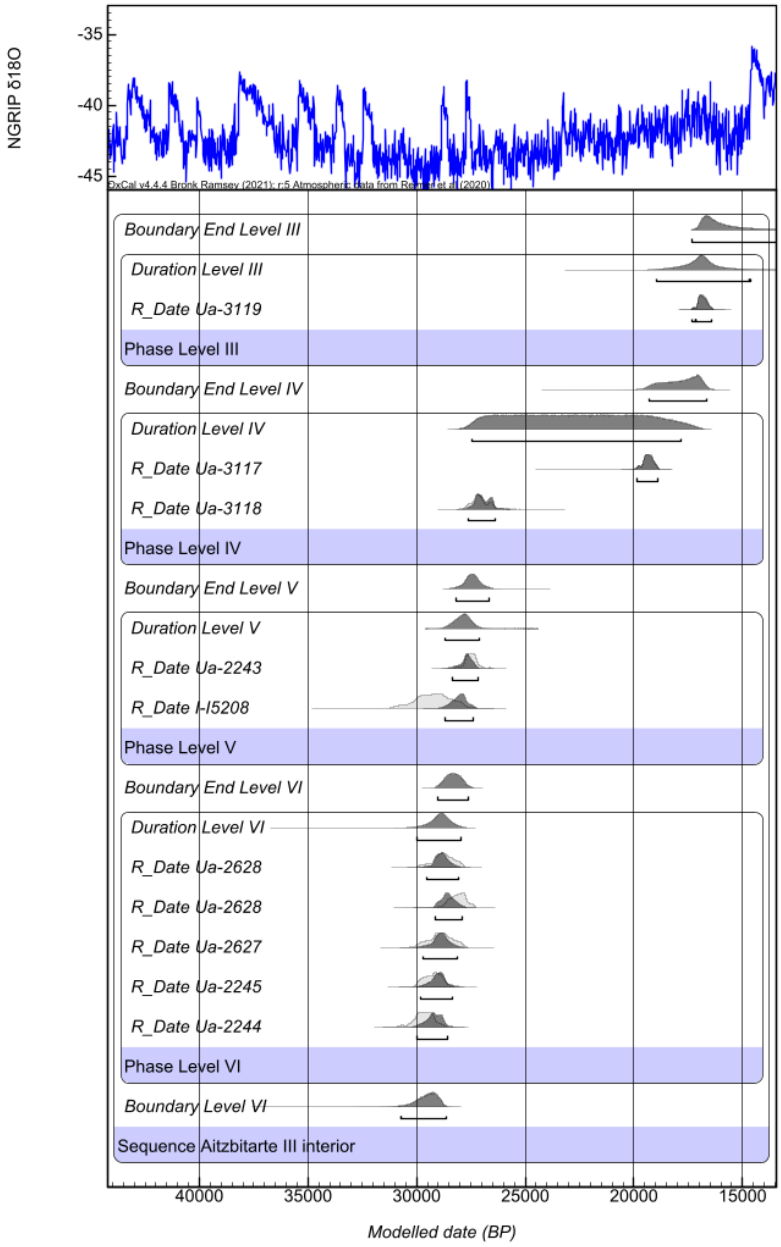

Figure C4. Radiocarbon dates from Aitzbitarte III-interior modelled in OxCal4.4 against INTCAL20.

Field Code Changed

Results of Bayesian Models

El Castillo	Unmodelled (BP)			Modelled (BP)			Indices Amodel 78.8, Aoverall 82.4			
	from	to	%	from	to	%	A	L	P	C
Boundary End Level 18B				48383	43733	95.449.974				97.1
Duration Level 18B				**48438**	**44536**	**95.449.974**				**99.8**
R_Date OxA-2197A	...	45427	95.449.973	48235	44793	95.449.974	98.1		95.2	99.8
R_Date OxA-21973	...	45655	95.449.973	48240	44793	95.449.974	91.9		95.2	99.8
Phase Level 18B										
Boundary End Level 18C				48470	45117	95.449.974				99.8
Duration Level 18C				**48977**	**45382**	**95.449.974**				**99.9**
R_Date OxA-22404	49976	42918	95.449.974	48833	45383	95.449.974	82.2		95.3	99.8
R_Date OxA-22203	49451	42999	95.449.974	48819	45381	95.449.974	76.1		95.2	99.8
R_Date OxA-22202	51146	43039	95.449.974	48861	45386	95.449.974	101.2		95.4	99.8
Phase Level 18C										
Boundary End Level 19				49629	45623	95.449.974				99.7
Duration Level 19				51060	45997	95.449.974				99.7
R_Date OxA-21974	...	44367	95.449.974	50965	45998	95.449.974	120.2		95.3	99.8
Phase Level 19										
Boundary End Level 20C				52583	46286	95.449.974				99.5
Duration Level 20C				**54134**	**46593**	**95.449.974**				**99.3**
R_Date OxA-22204	...	47048	95.449.974	53958	46713	95.449.974	94		95.3	99.3
R_Date OxA-22205	...	47348	95.449.974	53965	46715	95.449.974	86.9		95.3	99.3
Phase Level 20C										
Boundary El Castillo Level 20C				55552	46609	95.449.974				95.3
Sequence										
U(0	68.268.949	3.99E-17	4	68.268.949	5.38E-17	3.776		100		
T(5)	-2.65	2.65	95.449.974							99.9
Outlier_Model General				-2684	2502	95.449.974				100

Table C1. Radiocarbon dates from El Castillo modelled in OxCal4.4 against INTCAL20.

Field Code Changed

Aitzbitarte III Interior	Unmodelled (BP)			Modelled (BP)			Indices Amodel 78.8, Aoverall 82.4			
	from	to	%	from	to	%	A	L	P	C
Boundary End Level III				17300	12910	9.544.997				98
Duration Level III				18960	14630	9.544.997				99.6
R_Date Ua-3119	17270	16390	9.544.997	17300	16430	9.544.997	100.8		95.8	99.8
Phase Level III										
Boundary End Level IV				19320	16640	9.544.997				99.3
Duration Level IV				27430	17820	9.544.997				98.9
R_Date Ua-3117	19830	18900	9.544.997	19840	18910	9.544.997	99.9		95.3	99.6
R_Date Ua-3118	27700	26430	9.544.997	27600	26360	9.544.997	98.1		95.2	99.5
Phase Level IV										
Boundary End Level V				28210	26680	9.544.997				99.7
Duration Level V				**28680**	**27130**	**9.544.997**				**99.9**
R_Date Ua-2243	28260	26610	9.544.997	28370	27190	9.544.997	88.8		95.4	99.8
R_Date I-I5208	30830	27760	9.544.997	28710	27370	9.544.997	57.7		94.8	99.8
Phase Level V										
Boundary End Level VI				29010	27630	9.544.997				99.7
Duration Level VI				29990	27930	9.544.997				99.8
R_Date Ua-2628	29760	27840	9.544.997	29570	28080	9.544.997	118.2		96	99.8
R_Date Ua-2628	28760	27360	9.544.997	29150	27920	9.544.997	67		94.3	99.8
R_Date Ua-2627	29920	27870	9.544.997	29680	28110	9.544.997	120.5		96	99.8
R_Date Ua-2245	30070	28280	9.544.997	29820	28360	9.544.997	108		95.9	99.8
R_Date Ua-2244	30720	28760	9.544.997	30010	28570	9.544.997	77.7		94.9	99.7
Phase Level VI										
Boundary Level VI				30730	28650	9.544.997				96
Sequence										
U(0,4)	3.99E-17	4	9.544.997	5.38E-17	3.772	9.544.997	100			99
T(5)	-2.65	2.65	9.544.997							95.5
Outlier_Model General				-1420	1280	9.544.997				99.9

Table C2. Radiocarbon dates from Labeko Koba modelled in OxCal4.4 against INTCAL20.

Field Code Changed

Canyars	Unmodelled (BP)			Modelled (BP)			Indices Amodel 78.8, Aoverall 82.4			
Boundary End Level I				40090	37770	95.45				95.3
Duration Level I				**40890**	**38530**	**95.45**				**99.7**
R_Date Beta-2273965	39630	37570	9.544.997	40190	38560	95.45	63.2		93.4	99.6
R_Date OxA-23643	40520	39140	9.544.997	40330	39240	95.45	114.2		96.1	99.8
R_Date OxA-2416-44	40880	39450	9.544.997	40540	39400	95.45	99.2		96	99.8
R_Date OxA-23644	40740	39300	9.544.997	40470	39340	95.45	110.5		96	99.8
R_Date OxA-24057	40790	39390	9.544.997	40490	39380	95.45	104.3		96	99.8
Phase Level I										
Boundary Level I				41450	39500	95.45				96.6
Sequence Canyars										
U(0,4)	3.99E-17	4	9.544.997	5.38E-17	3.82	95.45	100			100
T(5)	-2.65	2.65	9.544.997							99.4
Outlier_Model General				-800	1480	95.45				99.9

Table C3. Radiocarbon dates from Canyars modelled in OxCal4.4 against INTCAL20.

Field Code Changed

Labeko Koba	Unmodelled (BP)			Modelled (BP)			Indices Amodel 78.8, Aoverall 82.4			
	from	to	%	from	to	%	A	L	P	C
Boundary End Level IV				38710	32030	9.544.997				98.4
Duration Level IV				**39000**	**33710**	**9.544.997**				**99.8**
R_Date OxA-21768	39700	37030	9.544.997	39050	33820	9.544.997	75.5		80	99.8
R_Date OxA-21780	39780	36910	9.544.997	39050	33960	9.544.997	81.3		82.3	99.8
Phase Level IV										
Boundary End Level V				39470	35440	9.544.997				99.8
Duration Level V				**39730**	**35950**	**9.544.997**				**99.8**
R_Date OxA-21779	41170	38260	9.544.997	39830	36330	9.544.997	21		87.2	99.8
R_Date OxA-21767	41230	38500	9.544.997	39860	36340	9.544.997	15.5		85.5	99.8
Phase Level V										
Boundary End Level VI				40240	36360	9.544.997				99.8
Duration Level VI				**41030**	**37860**	**9.544.997**				**99.9**
R_Date OxA-21841	37710	35420	9.544.997							
R_Date OxA-21794	38040	35460	9.544.997							
R_Combine comb:23689	37350	35900	9.544.997	40620	36500	9.544.997	4.3			99.8
R_Date OxA-21778	41390	39190	9.544.997	40970	38550	9.544.997	90		94.4	99.9
Phase Level VI										
Boundary End Level VII				41490	38890	9.544.997				99.9
Duration Level VII				**41910**	**39570**	**9.544.997**				**99.9**
R_Date OxA-21840	41610	39250	9.544.997							
R_Date OxA-21793	41720	39390	9.544.997							
R_Combine comb:23688	41290	39570	9.544.997	41650	39780	9.544.997	87.3			99.9
R_Date OxA-X-2314-43	42350	40260	9.544.997	41900	40000	9.544.997	96.5		95.4	99.9
R_Date OxA-21766	42520	40530	9.544.997	41950	40020	9.544.997	80.3		94.6	99.9
Phase Level VII										
Boundary End Level IX upper				42190	40360	9.544.997				99.9
Duration Level IX upper				**42750**	**40580**	**9.544.997**				**99.9**
R_Date OxA-22559	42090	39850	9.544.997							
R_Date OxA-22653	42520	40530	9.544.997							
R_Combine comb:27311	42120	40600	9.544.997	42330	40800	9.544.997	95			99.9
R_Date OxA-21792	42370	40330	9.544.997	42380	40820	9.544.997	113.4		95.7	99.9
R_Date OxA-21777	43160	40960	9.544.997	42600	40950	9.544.997	99.5		95.6	99.9
R_Date OxA-23199	43980	41490	9.544.997	42800	40990	9.544.997	52.4		92.8	99.9
Phase Level IX upper										
Boundary End Level IX inf				43420	40970	9.544.997				99.9
Duration Level IX inf				**48940**	**41340**	**9.544.997**				**99.8**
R_Date OxA-22560	42780	40980	9.544.997	49670	41300	9.544.997	75.3		76	99.8
R_Date OxA-22562	43830	41220	9.544.997	45860	41380	9.544.997	102.8		90.9	99.8
R_Date OxA-22563	43250	41010	9.544.997	46280	41300	9.544.997	99.1		89.7	99.8
R_Date OxA-22561	43790	41130	9.544.997	45920	41340	9.544.997	102.3		90.7	99.8
R_Date OxA-22564	43370	41050	9.544.997	46060	41320	9.544.997	101		90.2	99.8
Phase Level IXinf										
Boundary Labeko Level IX inf				52660	41740	9.544.997				96.6
Sequence Labeko Koba										
N(0,2)	-4	4	9.544.997							99.4
Outlier_Model SSimple				...	840	9.544.997				97.5
U(0,4)	3.99E-17	4	9.544.997	5.38E-17	3.932	9.544.997	100			98.3
T(5)	-2.65	2.65	9.544.997							97.5
Outlier_Model General				-6130	9280	9.544.997				99.4

Table C4. Radiocarbon dates from Aitzbitarte III-interior modelled in OxCal4.4 against INTCAL20.

Field Code Changed

Appendix D. Intratooth curve plots

Original curves derived from enamel intratooth sampling on enamel carbonate. Provided by sites. In blue, oxygen stable isotope composition ($\delta^{18}O$), and, in brown, carbon stable isotope composition ($\delta^{13}C$). In the x-axis, the distance from Enamel Rooth Junction (ERJ). Notice that the y-axis can experience some variations between sites.

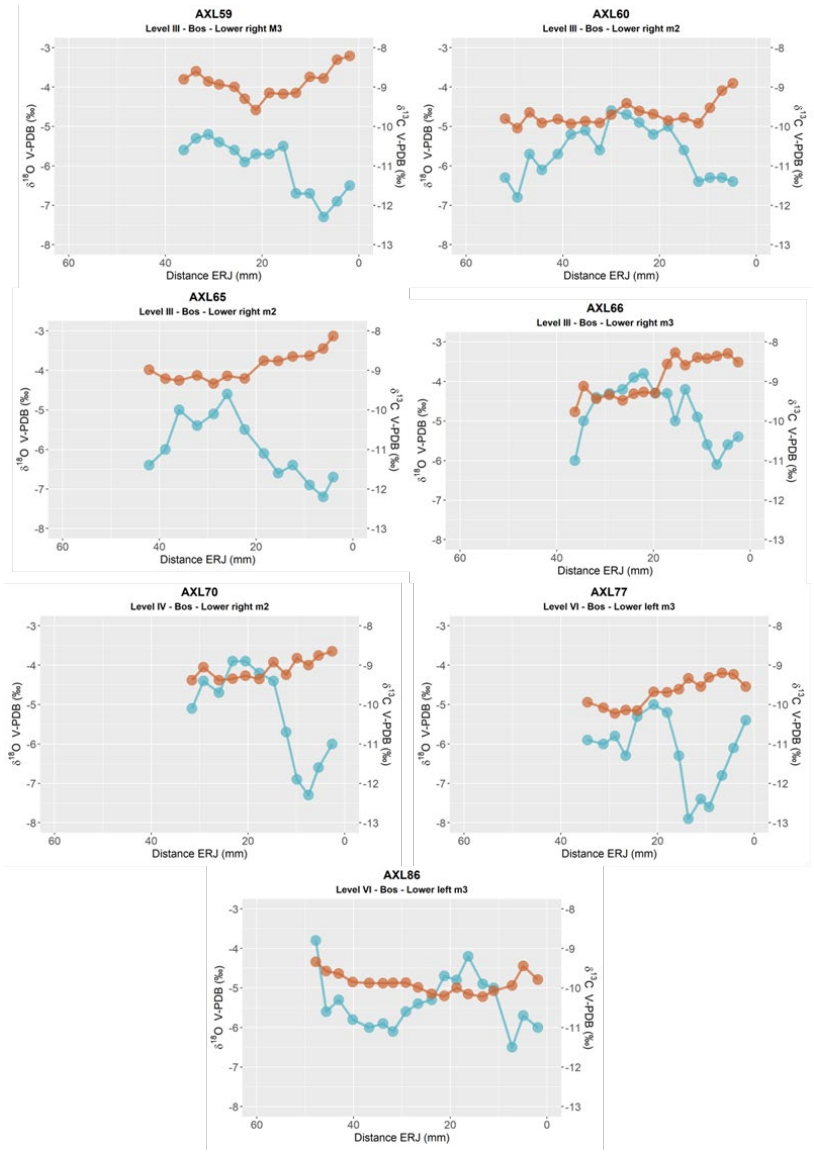

Figure D1. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from Axlor, considering distance from enamel root junction (ERC).

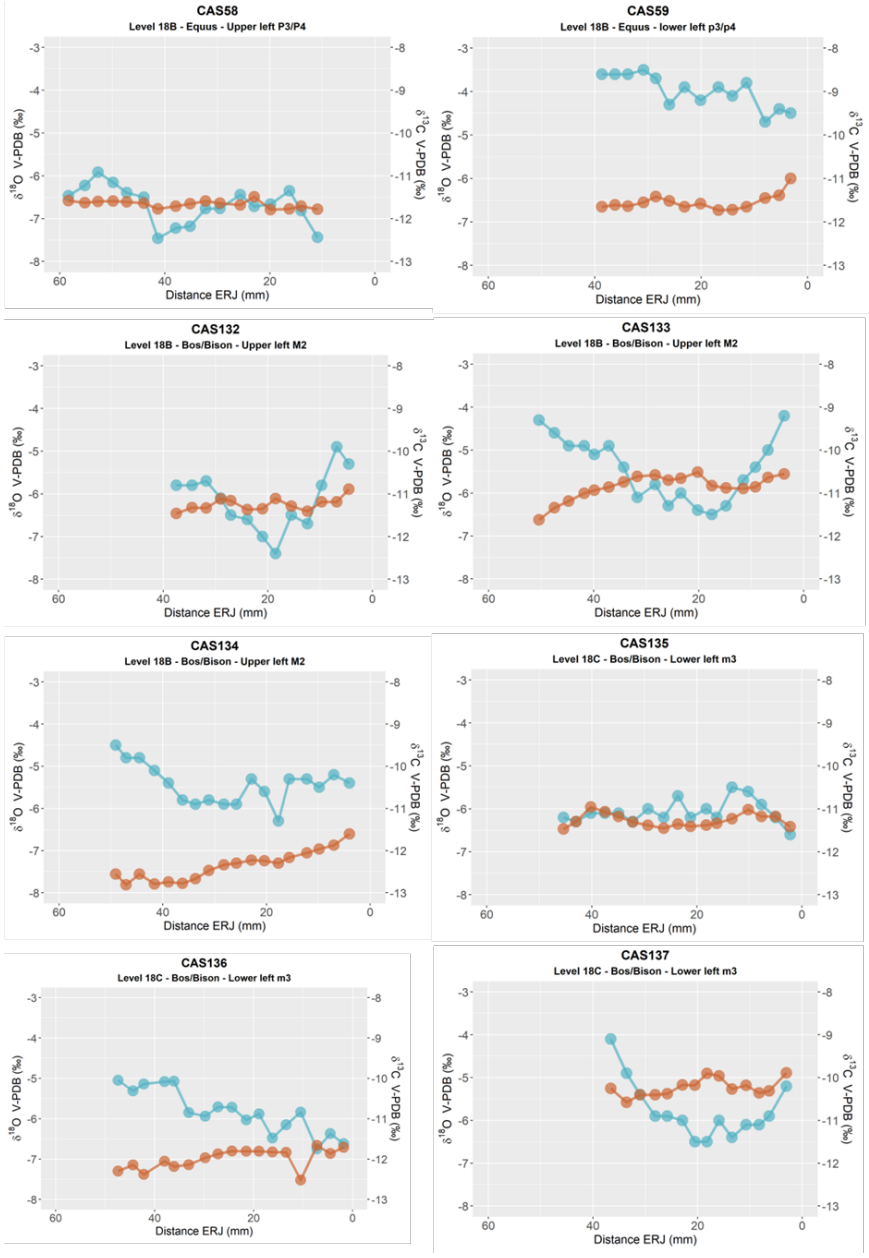

Figure D2. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from El Castillo, considering the sample's distance from the enamel root junction (ERC).

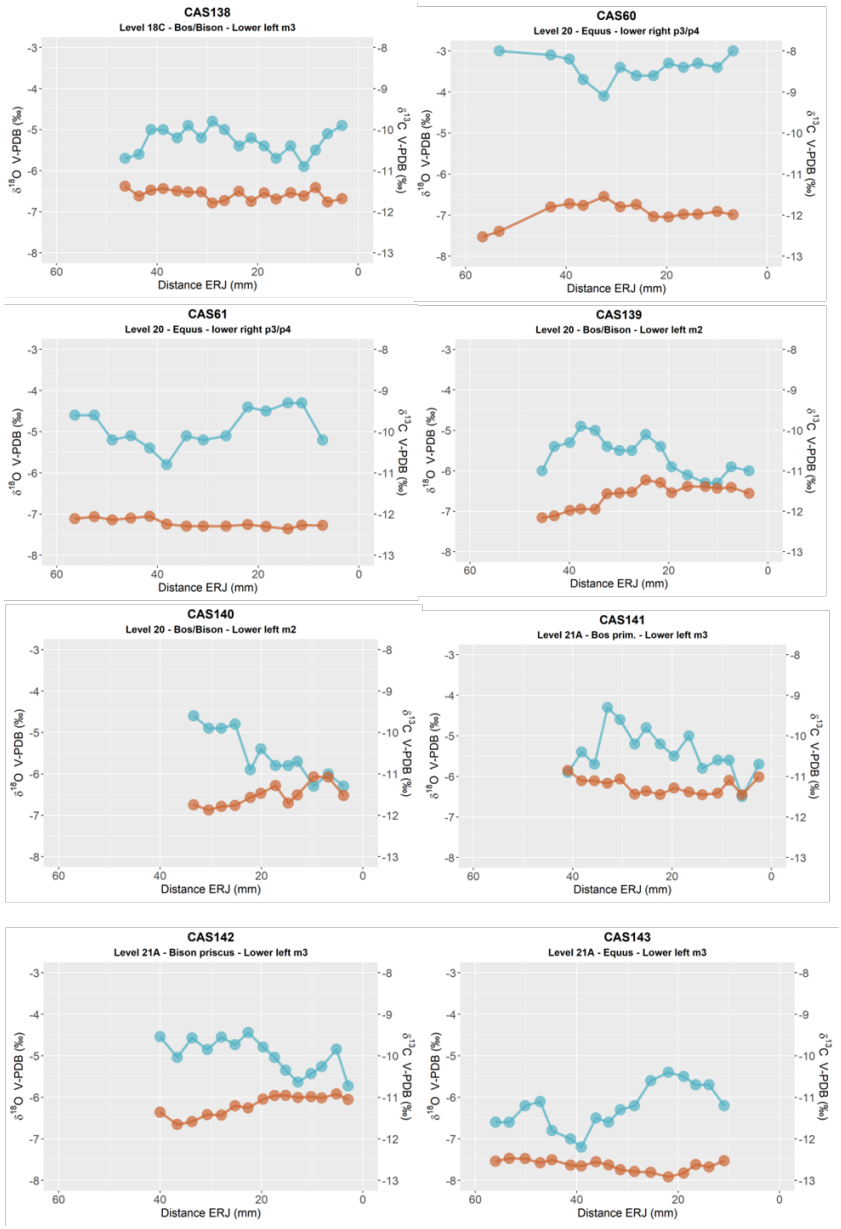

Figure D3. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from El Castillo, considering the sample's distance from the enamel root junction (ERC).

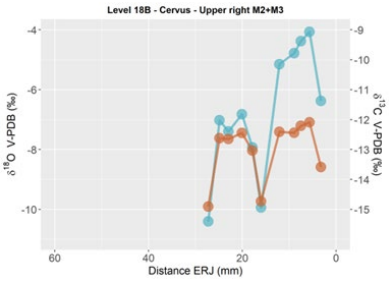

Figure D4. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from El Castillo, considering the sample's distance from the enamel root junction (ERC).

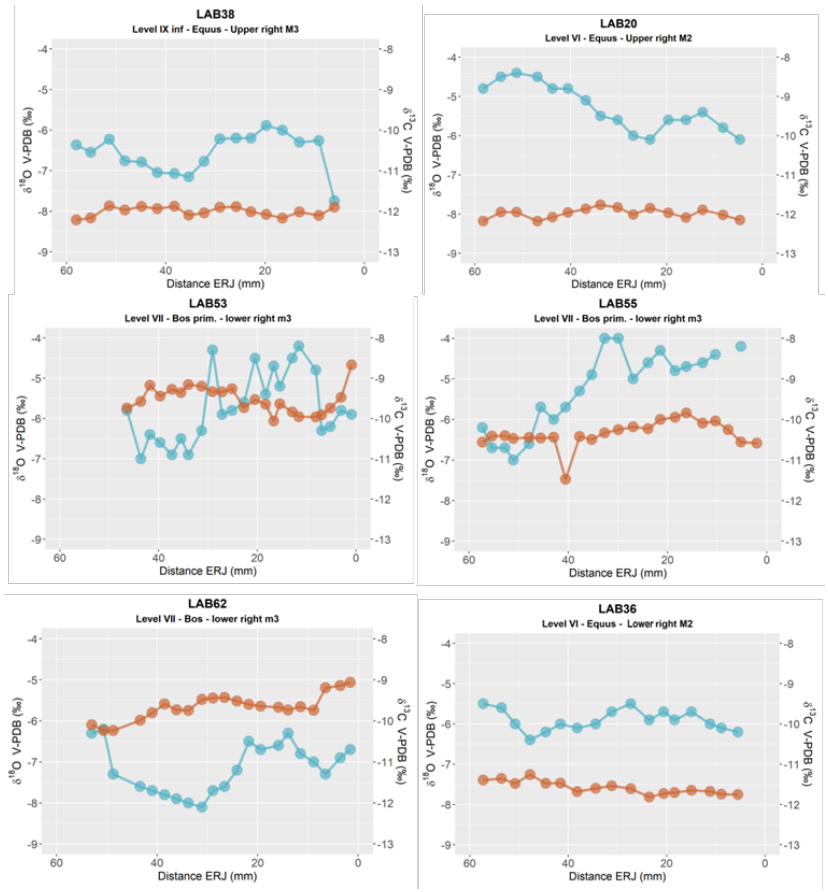

Figure D5. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from Labeko Koba, considering the sample's distance from the enamel root junction (ERC).

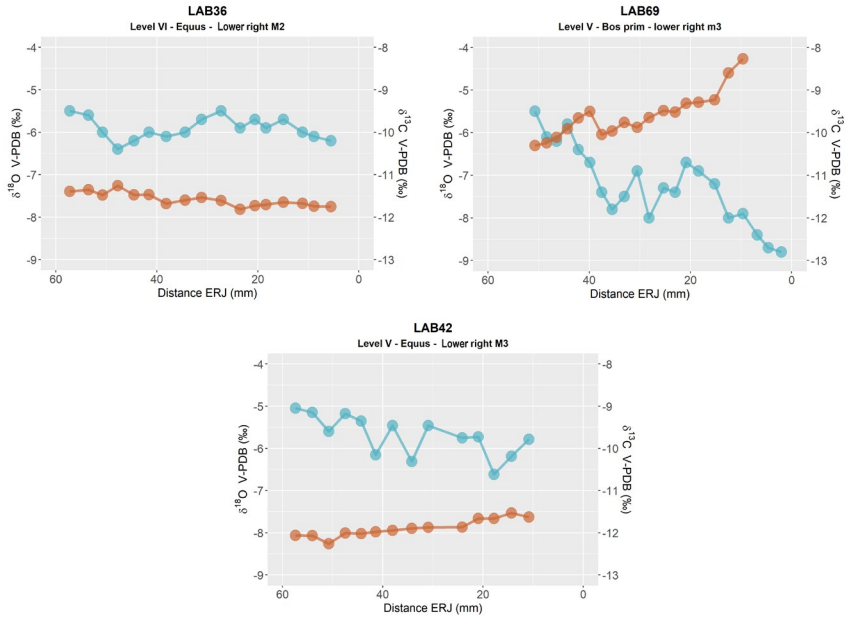

Figure D6. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from Labeko Koba, considering the sample's distance from the enamel root junction (ERC).

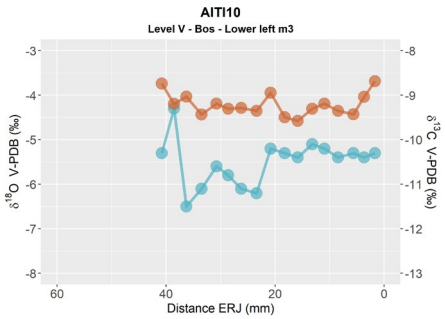

Figure D7. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from Aitzbitarte III interior, considering the sample's distance from the enamel root junction (ERC).

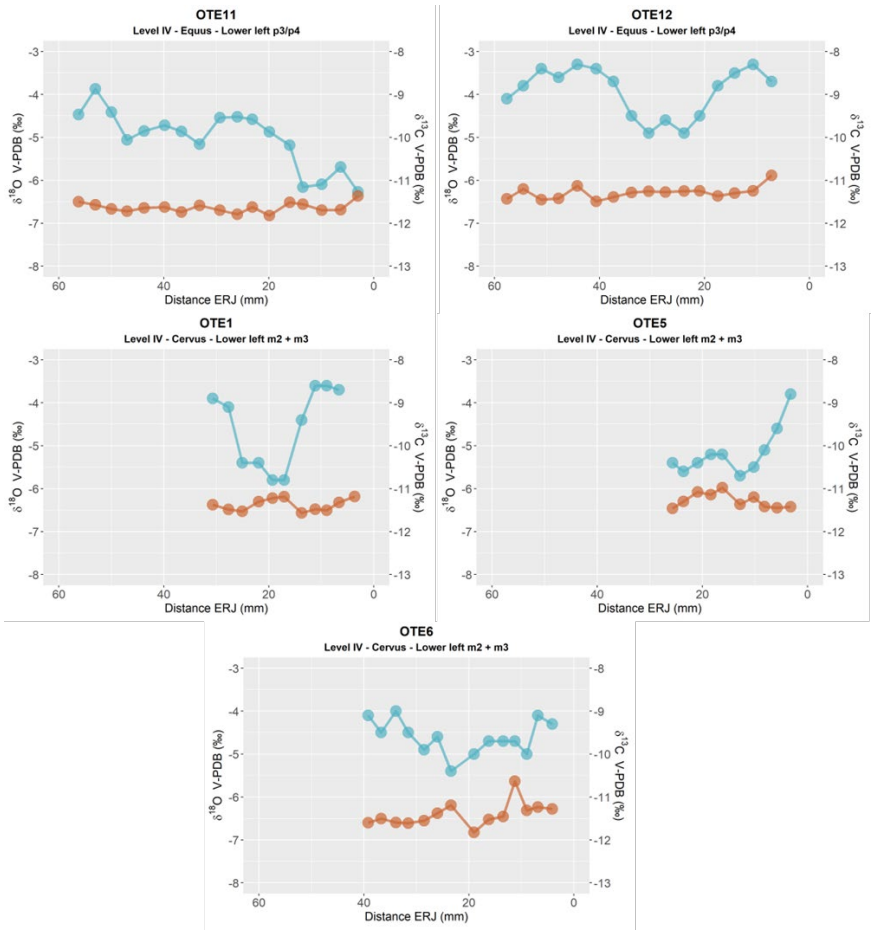

Figure D8. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from El Otero, considering the sample's distance from the enamel root junction (ERC).

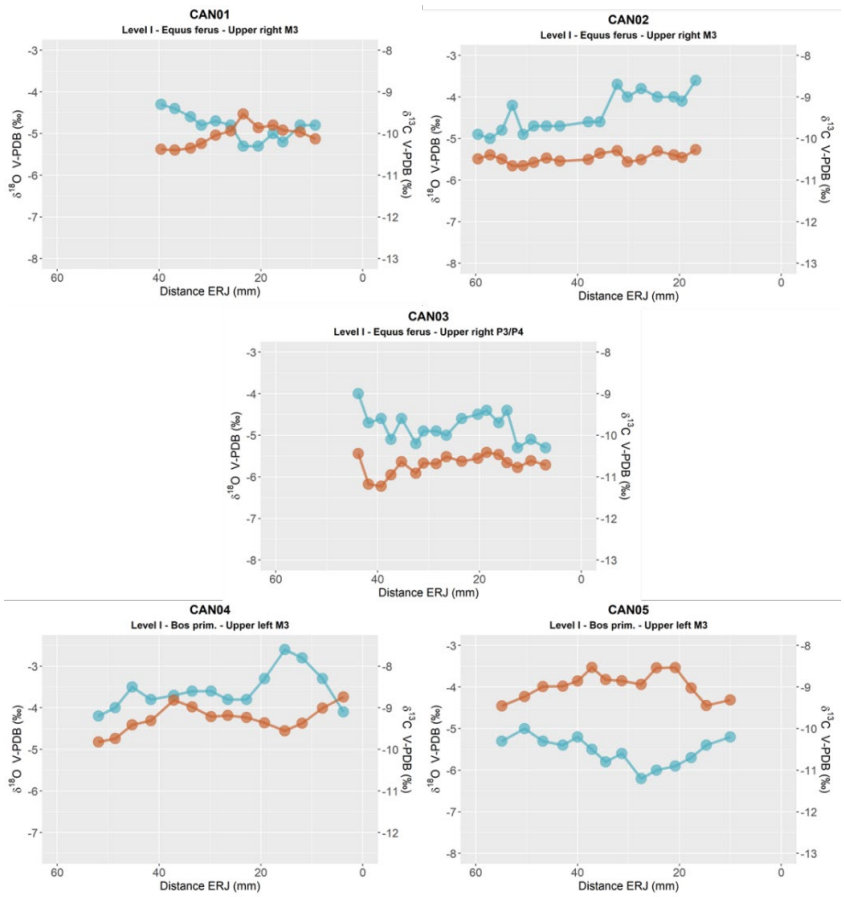

Figure D9. Intratooth plots of oxygen ($\delta^{18}O$) and carbon ($\delta^{13}C$) isotope composition from teeth from Canyars considering the sample's distance from the enamel root junction (ERC).

Appendix E. Inverse Modelling: Methodological Details and Models

The intratooth $\delta^{18}O$ profiles presented in this study were obtained through the application of inverse modelling, using an adapted version of the code published in reference (Passey et al., 2005b). This modeling approach allowed for the correction of the damping effect and the reconstruction of the original $\delta^{18}O$ input time series. The model reproduces the temporal delay between $\delta^{18}O$ changes in the animal's input and their manifestation in tooth enamel, exhibiting a consistent x-direction delay in the modelled $\delta^{18}O$ curve relative to the enamel $\delta^{18}O$ input time series. The model utilizes different species-specific parameters related to enamel formation, which vary between bovines and equids. These parameters have been established based on previous studies (Bendrey et al., 2015; Zazzo et al., 2012; Passey and Cerling, 2002; Kohn, 2004; Blumenthal et al., 2014). For *Bos/Bison* sp., the initial mineral content of enamel is fixed at 25%, the enamel appositional length is set at 1.5 mm, and the maturation length is 25 mm. For *Equus* sp., the initial mineral content of enamel is fixed at 22%, the enamel appositional length is set at 6 mm, and the maturation length is 28 mm.

In addition, the model requires other variables related to sampling geometry, as well as error estimates derived from mass spectrometer measurements. The distance between samples varies for each tooth, but as a general trend, the sampling depth on the tooth enamel surface in the samples of this study represents approximately 70% of the total enamel depth. The standard deviation of the measurements obtained from the mass spectrometer was typically set at 0.12%, taking into account the uncertainty associated with the standards. Finally, the models require a damping factor that determines the cumulative damping along the isotopic profile by adjusting the measured error (Emeas) to the prediction error (Epred). In the teeth analysed in this study, the damping factor ranged from 0.001 to 0.1.

The most likely model solutions were selected, and summer and winter values were extracted from the $\delta^{18}O$ profiles, considering the original peaks and troughs identified in the unmodelled $\delta^{18}O$ profile. This approach was adopted to prevent the introduction of artificial peaks that the model may produce, particularly in teeth without a distinct sinusoidal shape. Flat and less sinusoidal profile are less suitable for the application of the model, given its inherent assumption of an approximately sinusoidal form. Non-sinusoidal curves can lead to complex interpretations in the model outcomes. Consequently, this methodology was not applied to analysed intratooth $\delta^{13}C$ profiles, as the examined individuals did not exhibit appreciable seasonal change.

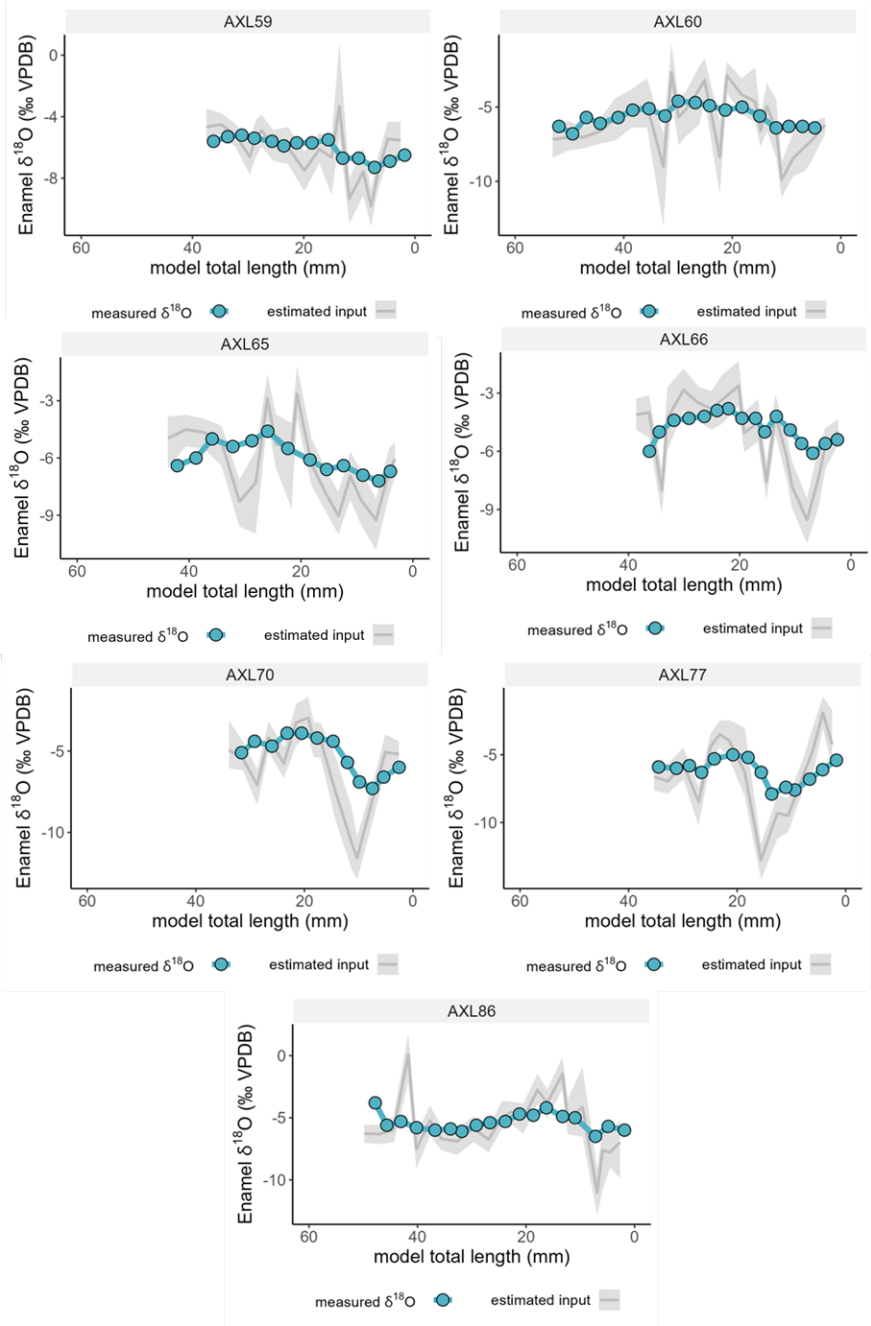

Figure E1. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from Axlor, considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

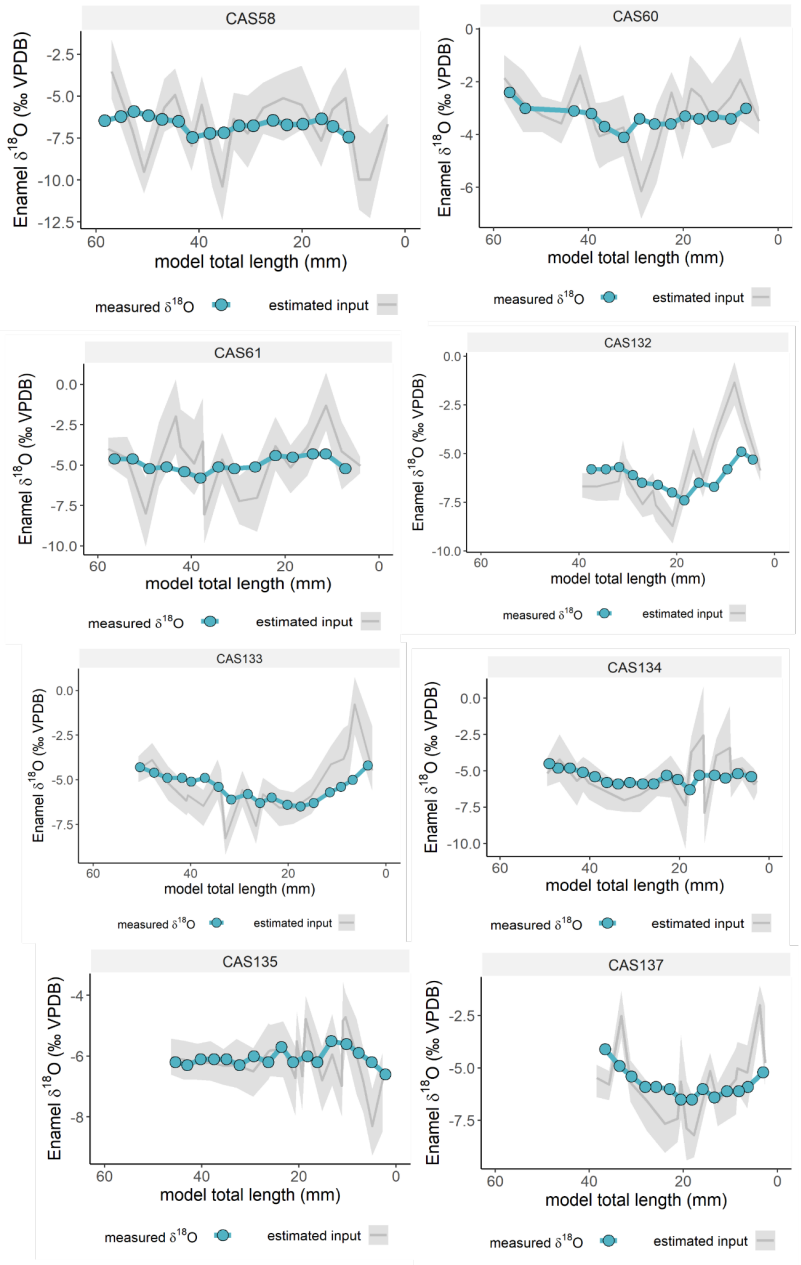

Figure E2. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from El Castillo, considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

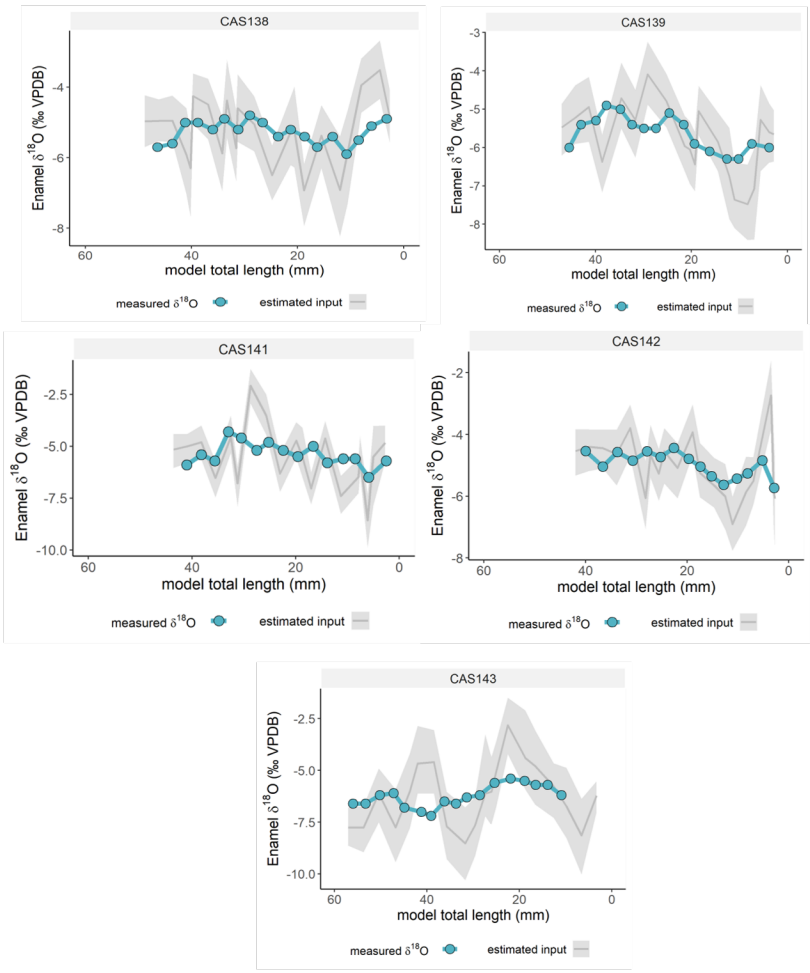

Figure E3. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from El Castillo, considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

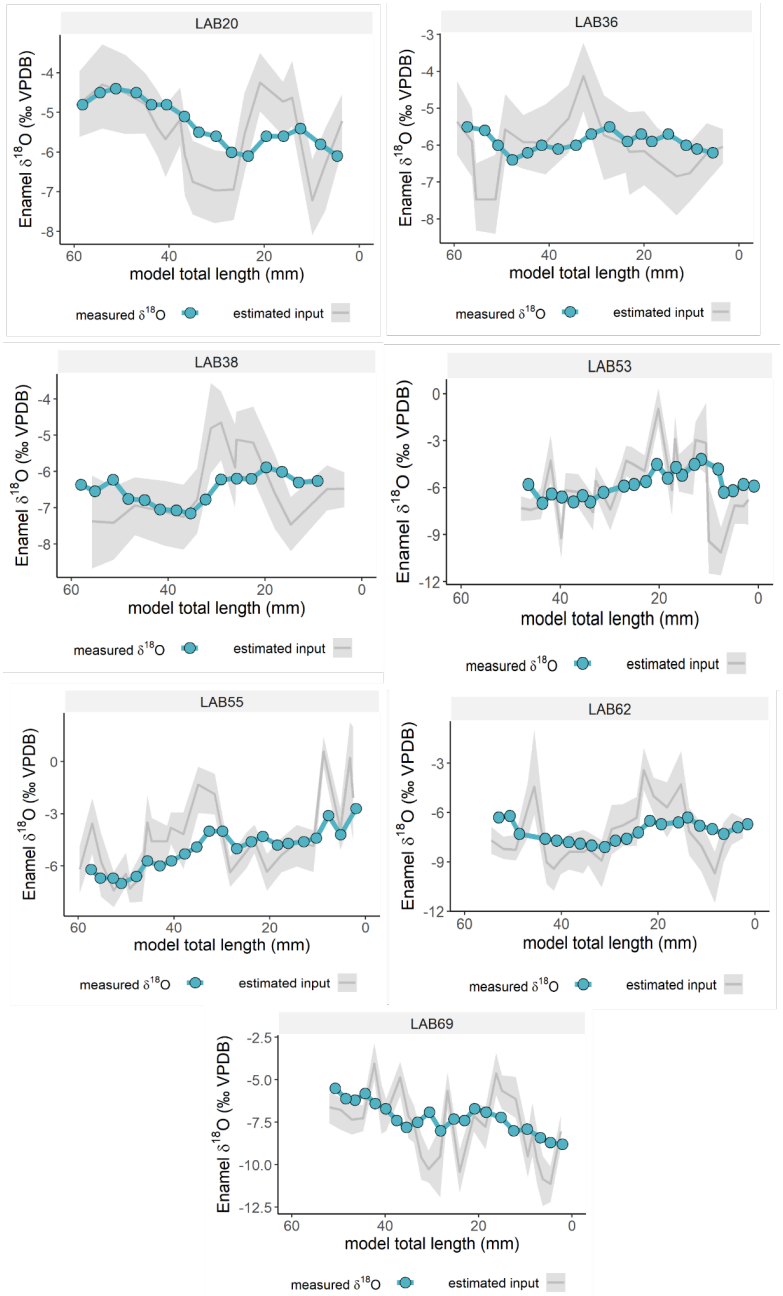

Figure E4. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from Labeko Koba, considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

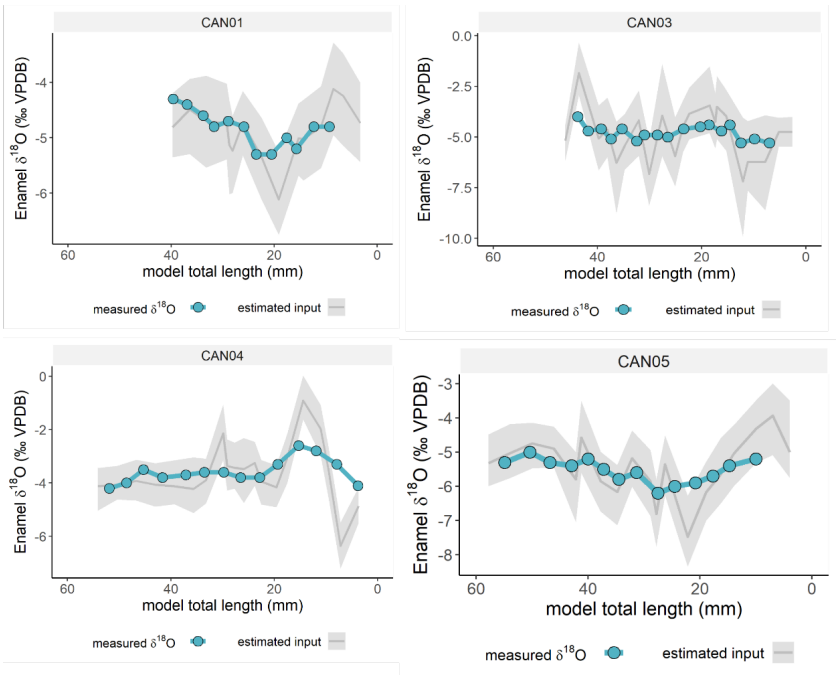

Figure E5. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from Canyars considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

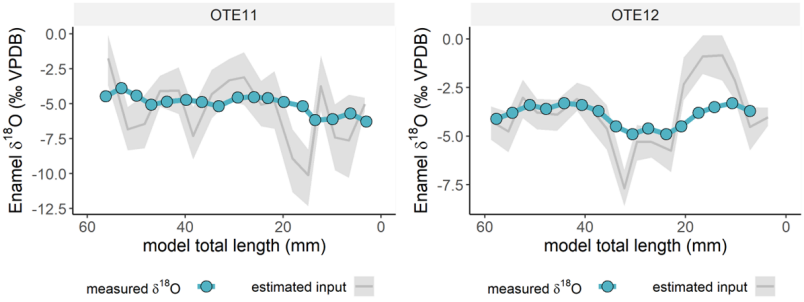

Figure E6. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from El Otero, considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

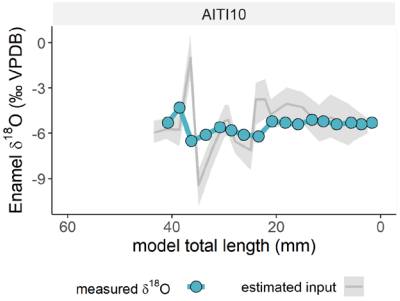

Figure E7. Inverse models for oxygen isotope composition ($\delta^{18}O$) from teeth from Aitzbitarte III interior, considering distance from enamel root junction. The blue line and points correspond to original data and grey line the most likely model solution, with the 95% confidence interval shown in shaded areas.

References Appendix E

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