



1 **Impacts of passive experimental warming on daytime and**  
2 **night-time respiration in a semi-natural grassland**

3

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7 **Abstract**

8 Soil respiration (SR) is the largest source of CO<sub>2</sub> released from the terrestrial ecosystem. It is  
9 greatly influenced by soil carbon pool, climate warming and daily fluxes i.e., daytime (DT)  
10 and night-time (NT) temperatures. However, there are hardly any studies relating to the  
11 effects of passive experimental warming on Ecosystem respiration (ER) and SR during DT  
12 and NT. We conducted a simulated warming experiment using passive Open Top Chamber  
13 (OTC) in a semi-natural grassland of Doon Valley, in the state of Uttarakhand, India. OTCs  
14 showed an increase in DT and NT soil temperatures. SR and ER were measured within OTC  
15 as well as outside using LI-8100A Automated Soil CO<sub>2</sub> Flux System. We found that SR and  
16 ER increased under passive experimental warming by 38.66% and 20.35% during DT, and  
17 38.8% and 12.41% during NT respectively. SR/ER ratio increased under passive warming  
18 treatment during DT and NT, indicating SR as the major contributor to ER. Temperature-  
19 respiration showed a positive relationship under ambient and warming conditions. Q<sub>10</sub>  
20 analyses revealed that respiration rates are sensitive to passive warming, especially during the  
21 NT. This study addresses the crucial gap of monitoring NT respiration in addition to DT  
22 respiration to estimate the CO<sub>2</sub> efflux and its response to passive experimental warming.

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24

25 **Keywords:** Soil respiration (SR), Ecosystem respiration (ER), Open Top Chamber (OTC),  
26 passive experimental warming, Temperature sensitivity, Night-time respiration, Daytime  
27 respiration



## 28 1. Introduction

29 Soil Respiration (SR) is the major source of carbon dioxide (CO<sub>2</sub>) emitted from the terrestrial  
30 ecosystem (Raich and Schlesinger, 1992; Zhao et al., 2017). The global CO<sub>2</sub> efflux accounts  
31 for approximately 50-95 Pg C yr<sup>-1</sup> (Houghton and Woodwell, 1989; Hashimoto et al., 2015).  
32 The emissions from anthropogenic sources are estimated to be approximately ten times lower  
33 than that of natural sources (Schaefer et al., 2009; Hashimoto et al., 2015). Alterations in the  
34 SR can have a significant impact on atmospheric CO<sub>2</sub> concentrations, resulting in either  
35 positive (Lu et al., 2013) or negative (Schlesinger, 1995; Wang et al., 2014) feedback to the  
36 climate warming.

37 SR, a component of ecosystem respiration (ER), is a combination of both heterotrophic and  
38 autotrophic respiration. It may respond differently to alterations in temperatures and  
39 environmental conditions (Fang et al., 2018). It is observed that the rates of soil respiration  
40 (SR) and ecosystem respiration (ER) are more susceptible to fluctuations due to  
41 environmental factors, as compared to photosynthesis (Valentini et al., 2000). Despite its  
42 importance, respiration is less studied. Furthermore, studies primarily concentrate on DT  
43 respiration rates specifically during the growing season. The study of NT respiration rate is  
44 infrequently conducted as a result of challenges in its measuring during the night (Hu et al.,  
45 2016). DT and NT respiration rates are variable, explained by various microclimatic  
46 parameters including soil temperature (Atkin et al., 2003; Xia et al., 2009) and soil water  
47 content (SWC) (Xia et al., 2009; Balogh et al., 2011). ST and SWC influence respiration rates  
48 by affecting functioning of soil microorganisms, as well as the respiratory enzymes present in  
49 microflora and fauna (Fang and Moncrieff, 2001; Atkin et al., 2003; Moyano et al., 2007). To  
50 evaluate the effect of experimental warming on respiration rates, many studies have  
51 determined the temperature sensitivity of SR and/or ER by calculating Q<sub>10</sub> (Raich and  
52 Schlesinger, 1992; Wan et al., 2007; Fang et al., 2017; Fang et al., 2018). The Q<sub>10</sub> function is  
53 considered a better choice for estimating the respiration rates as it includes all the processes  
54 and factors that may impact the respiration rates (Vesterdal et al., 2012).

55 Grasslands, one of the world's major ecosystems, occur in a wide range of eco-climatic  
56 conditions and are governed by anthropogenic as well as climatic factors (Hall and Scurlock,  
57 1991). The tropical and sub-tropical grasslands in India are mostly anthropogenic in origin  
58 derived from clear-felling, livestock grazing and burning. More actively managed grasslands  
59 are often referred to as semi-natural grasslands. Such grasslands are known to be seral in  
60 nature and are rapidly colonized by shrubs and trees in the absence of management (Queiroz



61 et al., 2014). Soil carbon stock in such grasslands corresponds to at least 10% of the world's  
62 total (Eswaran et al., 1993; Zhao et al., 2017). However, several studies suggest that the  
63 estimation could be as high as 30% of the world's soil carbon (Scurlock and Hall, 1998). The  
64 exploration of interactions between climate change and grasslands has been relatively limited  
65 compared to forests (Hall and Scurlock, 1991). This highlights the significance of  
66 understanding the phenomenon taking place within this particular ecosystem. In the present  
67 study, we simulated passive experimental warming in a semi-natural grassland using an Open  
68 Top Chamber (OTC) to study the impacts of passive warming on SR and ER. The objectives  
69 of the study were to (1) examine the impacts of passive experimental warming on  
70 microclimatic parameters, (2) assess the impacts of passive experimental warming on DT and  
71 NT respiration, and their temperature sensitivities, (3) evaluate SR/ER ratio during DT and  
72 NT, and (4) understand temperature-respiration and moisture-respiration relationships.

73

## 74 2. Materials and Methods

### 75 2.1. Study site

76 The study was conducted in a semi-natural grassland patch in Doon Valley, within Wildlife  
77 Institute of India, Dehradun, Uttarakhand, India (30°17'02" N; 77°58'23" E, 598 m asl). The  
78 area experiences a typical sub-tropical climate with an annual mean air temperature of  $21.8 \pm$   
79  $0.1$  °C (from January 2020 to January 2021) and precipitation of about 2073.3 mm (India  
80 Meteorological Department, 2015). The study site is a maintained grassland with dominant  
81 plant species, including *Dicanthium annulatum*, *Medicago polymorpha*, and *Alternanthera*  
82 *sessilis*, surrounded by trees such as *Butea monosperma*, *Shorea robusta*, and *Bombax ceiba*.  
83 Soil pH, bulk density, electrical conductivity and organic carbon are  $7.12 \pm 0.16$ ,  $0.96 \pm 0.08$   
84  $\text{g cm}^{-3}$ ,  $36.48 \pm 4.43 \text{ dS m}^{-1}$  and  $25.26 \pm 1.35 \text{ g Kg}^{-1}$ , respectively.

85

### 86 2.2. Experimental Design

87 A 2×2 m plot was selected for the study based on uniform vegetation, even terrain, and an  
88 equal proportion of sunlight reaching the ground, with careful avoidance of shade from trees.  
89 The plot was divided into two halves and assigned to control and experimental warming. To  
90 achieve passive warming, hexagonal Open Top Chamber was installed (Fig. 1) during the  
91 first week of January 2019. The chamber was made up of transparent polycarbonate sheets of  
92 3 mm thickness with base, height and upper diameter of 175 cm, 70 cm, and 110 cm



93 respectively. An adjacent paired hexagonal control plot was marked and fenced to avoid  
94 disturbances. In each plot, six cylindrical opaque soil collars (diameter 20 cm and height 11  
95 cm) made of polyvinyl chloride were randomly inserted 2 cm in the soil. Among the six soil  
96 collars in each plot, we allocated three for SR measurements and the remaining three for ER  
97 measurements. These collars were left at the site for the entire duration of the experiment to  
98 minimize disturbance.



99  
100 **Figure 1:** Hexagonal Open Top Chamber and adjacent control plot

101

### 102 2.3. ER, SR, and CO<sub>2</sub> measurements

103 Aboveground vegetation was clipped from three collars 24 hours prior to the measurements  
104 for SR measurement. ER and SR were measured on a clear day using LI-8100A Automated  
105 Soil CO<sub>2</sub> Flux System equipped with the LI-8100-103 opaque chamber (LICOR Inc.,  
106 Lincoln, NE, USA). Respiration readings were taken by gently mounting the LI-chamber on  
107 each collar and observing for 120 seconds with a dead band of 15 seconds. Two types of  
108 measurements were taken (i) 2 times during 0600-1800 hours, twice a week, depending upon  
109 the temperature peak and environmental feasibility and (ii) continuous hourly measurements  
110 for 24 hrs, twice a month. The measurement period was from January 2020 to March 2020.  
111 Data between 0600-1800 hours and 1800-0600 were pooled into DT and NT, respectively.  
112 The LI-8100 instrument recorded surface CO<sub>2</sub> concentration before mounting the chamber  
113 for respiration measurements.

114

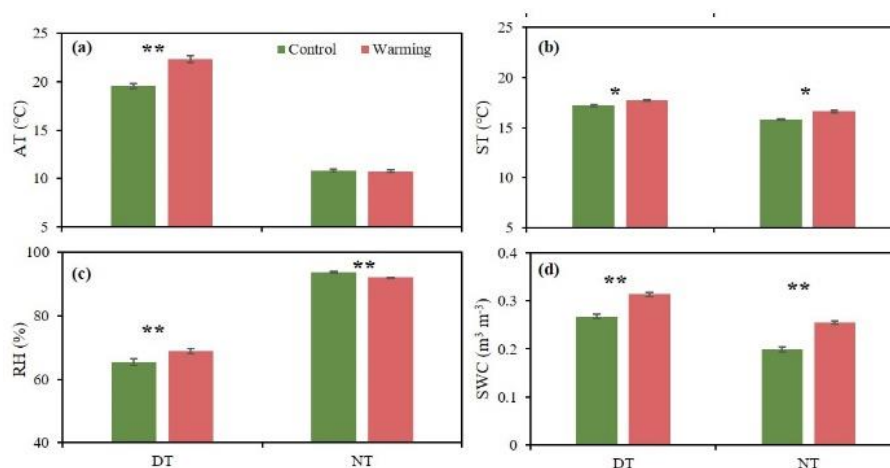




147 **3. Results**

148 **3.1. Microclimate under passive experimental warming**

149 OTC increased only DT AT by 2.78 °C ( $p < 0.001$ ), while increased both DT and NT ST by  
150 0.53 °C ( $p < 0.05$ ) and 0.79 °C ( $p < 0.05$ ), respectively (Fig. 2). We observed higher SWC  
151 under passive experimental warming by 32.39% ( $0.046 \text{ m}^3 \text{ m}^{-3}$ ,  $p < 0.001$ ) during DT and  
152 49.46% ( $0.056 \text{ m}^3 \text{ m}^{-3}$ ,  $p < 0.001$ ) during NT. Relative humidity (RH) increased during DT  
153 by 3.42% ( $p < 0.001$ ) and reduced during NT by 1.72% ( $p < 0.001$ ) under warming.

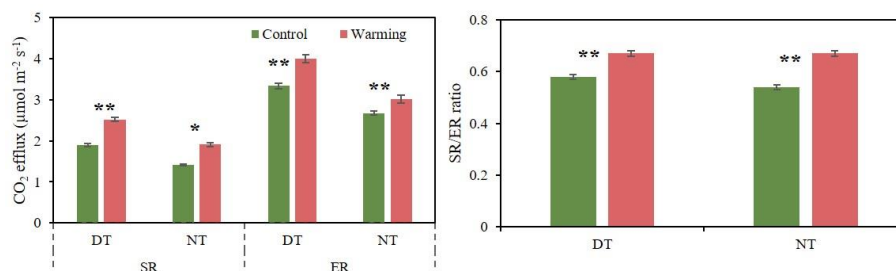


154  
155 **Figure 2.** Daytime (DT), night-time (NT) mean (a) air temperature (AT), (b) soil  
156 temperature (ST), (c) relative humidity (RH), and (d) volumetric soil water content (SWC),  
157 under control and warming condition, significant at \* $<0.05$  and \*\* $<0.001$

158

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159 **3.2. Impacts of passive experimental warming on Respiration rates and SR/ER ratio**  
160 SR and ER ranged from  $0.53\text{-}3.29 \mu\text{mol m}^{-2} \text{ s}^{-1}$  and  $1.58\text{-}7.13 \mu\text{mol m}^{-2} \text{ s}^{-1}$  during DT and  
161 from  $0.71\text{-}2.03 \mu\text{mol m}^{-2} \text{ s}^{-1}$  and  $1.49\text{-}3.85 \mu\text{mol m}^{-2} \text{ s}^{-1}$  during NT respectively. Passive  
162 experimental warming increased both SR and ER by 38.66% and 20.35% during DT and by  
163 38.8% and 12.41% during NT, respectively (Fig. 3). SR/ER ratio also increased under passive  
164 experimental warming from 0.58 to 0.66 and from 0.53 to 0.67 during DT and NT.

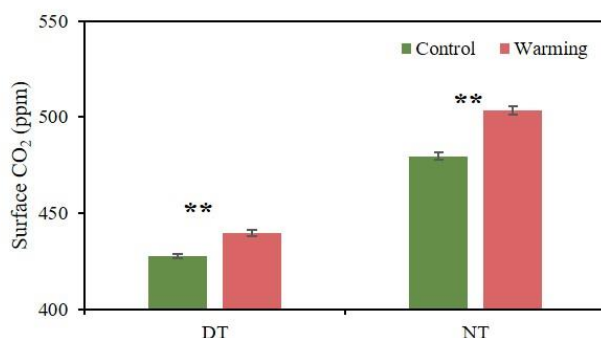


165  
166 **Figure 3.** Daytime (DT), night-time (NT) mean Soil respiration (SR), Ecosystem respiration  
167 (ER), and SR/ER ratio, under control and warming condition, significant at \* < 0.05 and  
168 \*\* < 0.001

169

### 170 3.3. Surface CO<sub>2</sub> levels

171 Surface CO<sub>2</sub> ranged from 396–531 ppm and 422–594 ppm during DT and NT, respectively, in  
172 our study. Passive experimental warming increased the mean surface CO<sub>2</sub> during DT and NT  
173 by 12 ppm (p < 0.001) and 24 ppm (p < 0.001), respectively (Fig. 4.).



174

175 **Figure 4.** Daytime (DT), night-time (NT) mean surface CO<sub>2</sub> levels, under control and  
176 warming condition, significant at \*\* < 0.001

177

### 178 3.4. Temperature-respiration and moisture-respiration relationships

179 In our study, soil temperature and respiration rate were positively correlated. Under passive  
180 experimental warming, the exponential relationships between temperature and respiration  
181 (both SR and ER) increased (Table 1 & 2). ST explained variability in SR significantly (p <  
182 0.001) by 24% and 13% under ambient conditions, which increased to 57% and 78% under  
183 passive experimental warming during DT and NT, respectively. Similarly, ST explained





184 variabilities in ER significantly ( $p < 0.001$ ) by 36% and 48% under ambient conditions,  
 185 which increased to 66% and 87% under passive experimental warming during DT and NT,  
 186 respectively.

187 Soil moisture only showed a significant yet weak relationship with SR during NT. It  
 188 explained 5.6% and 5.3% of its variabilities in control and warming, respectively, while ER  
 189 related to moisture only during DT in control ( $r^2 = 0.056$ ,  $p < 0.01$ ) (Table 1 & 2).

190

191 **Table 1.** Regression models to show exponential relationships between ST and SR ( $R = ae^{bt}$ )  
 192 and linear relationship between SWC and SR ( $R = aM + b$ ). values in the parenthesis represent  
 193 S.E. of the estimate.

			Soil respiration			
			a	b	r <sup>2</sup>	p
<b>R = ae<sup>bt</sup></b>	<b>Day-time</b>	Control	0.776 (0.084)	0.050 (0.006)	0.240	<0.001
		Warming	0.360 (0.042)	0.109 (0.007)	0.571	<0.001
	<b>Night-time</b>	Control	0.915 (0.096)	0.026 (0.006)	0.132	<0.001
		Warming	0.275 (0.027)	0.113 (0.006)	0.779	<0.001
<b>R = aM + b</b>	<b>Day-time</b>	Control	0.781 (0.397)	1.677 (0.114)	0.019	0.050
		Warming	0.255 (0.767)	2.441 (0.246)	0.001	0.740
	<b>Night-time</b>	Control	0.849 (0.338)	1.244 (0.071)	0.056	0.014
		Warming	2.708 (1.109)	1.218 (0.287)	0.053	0.016

194

195 **Table 2.** Regression models to show exponential relationship between ST and ER ( $R = ae^{bt}$ )  
 196 and linear relationship between SWC and ER ( $R = aM + b$ ). values in the parenthesis represent  
 197 S.E. of the estimate.

			Ecosystem respiration			
			a	b	r <sup>2</sup>	p
<b>R = ae<sup>bt</sup></b>	<b>Day-time</b>	Control	1.188 (0.112)	0.059 (0.006)	0.363	<0.001
		Warming	0.312 (0.039)	0.142 (0.007)	0.663	<0.001
	<b>Night-time</b>	Control	1.208 (0.097)	0.048 (0.005)	0.476	<0.001
		Warming	0.165 (0.018)	0.168 (0.006)	0.870	<0.001



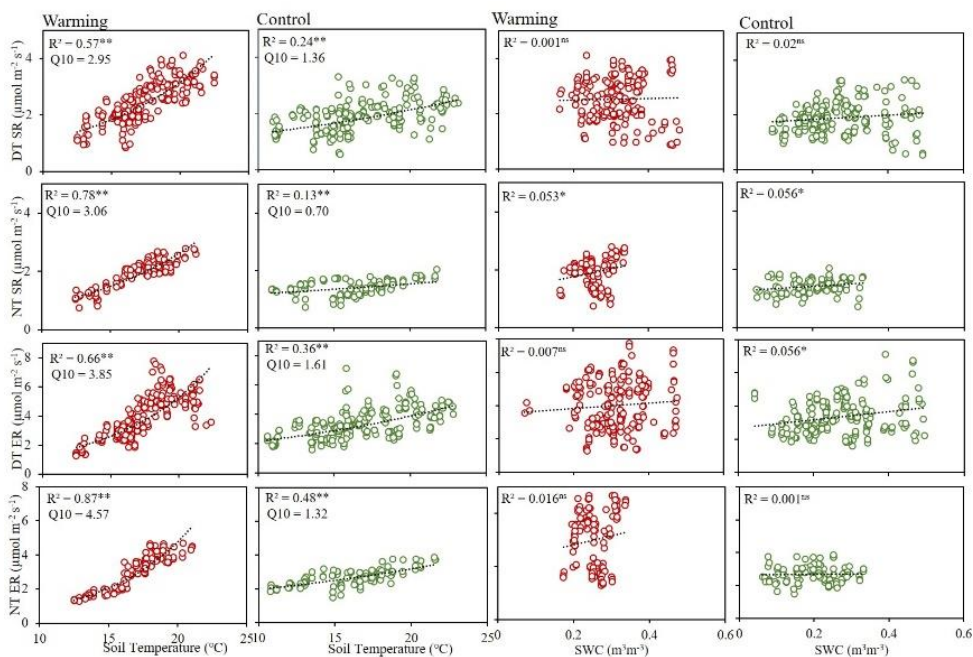
<b>R = aM+b</b>	<b>Day-time</b>	Control	2.427 (0.707)	2.694 (0.201)	0.056	0.001
		Warming	1.666 (1.363)	3.481 (0.440)	0.007	0.223
	<b>Night-time</b>	Control	0.182 (0.679)	2.633 (0.144)	0.001	0.789
		Warming	3.387 (2.604)	2.149 (0.668)	0.016	0.196

198

199 **3.5. Temperature sensitivity of SR and ER**

200 Temperature sensitivity of SR and ER was assessed based on the Q10 values of the  
 201 respiration calculated from the beta value of the temperature-respiration relationship. Q10  
 202 values ranged from 0.70-1.61 and 2.95-4.57 under ambient and warming condition  
 203 respectively, as shown in Fig. 4. We observed that passive experimental warming increased  
 204 Q10 of both SR and ER by 117% and 139% during DT and 337% and 246% during NT,  
 205 respectively.

206



207

208 **Figure 5.** Regression models with Q10 values to show relationships between temperature and  
 209 respiration ( $R=aebt$ ) and SWC and respiration ( $R = aM+b$ ), significant at  $^* < 0.05$  and  
 210  $^{**} < 0.001$ , ns represents non-significant, under control and warming conditions.



## 211 4. Discussion

### 212 4.1. Effect of passive experimental warming on microclimate

213 OTC increased DT AT by 2.78°C, consistent with several other studies in the grassland  
214 ecosystem (Flanagan et al., 2013; Tiruvaimozhi and Sankaran, 2019). ST at 5 cm depth  
215 increased significantly under passive experimental warming both during DT and NT, as  
216 found in other studies (Defrenne et al., 2010; Fang et al., 2017; Tiruvaimozhi and Sankaran,  
217 2019). The possible reason for this may be the reduction in wind speed inside the OTC and,  
218 thus, the reduced diffusion rates (Molau, 1997; Flanagan et al., 2013).

219 Temperature and RH showed an inverse relationship. Passive experimental warming  
220 increased SWC throughout the day, as also reported by Defrenne et al. (2010). This may be  
221 due to the condensation of water droplets on the inner side of the polycarbonate sheets.

222

### 223 4.2. Effects of passive experimental warming on SR, ER and SR/ER ratio

224 Several experimental warming studies have reported either increased (Fang et al, 2017;  
225 Flanagan et al., 2013; Tiwari et al., 2021), decreased (Sharkhuu et al., 2016), or no change  
226 (Sharkhuu et al., 2013; Wan et al., 2007) in respiration rates. In our study, SR and ER  
227 increased under passive experimental warming during DT and NT. This can be due to: (i)  
228 enhanced microbial activity due to the direct effect of increased temperatures, contributing  
229 more to the respiration rates in the warming plot (Flanagan et al., 2013, Fang et al., 2017) (ii)  
230 increase in the amount of carbon substrate available in grassland ecosystem due to the as a  
231 result of increased carbon allocation to the roots, microbes and exudates (Shaver et al., 2000;  
232 Flanagan et al., 2013).

233 The mean SR/ER ratio increased under passive experimental warming from 0.58 to 0.66  
234 during DT and from 0.53 to 0.67 during NT, suggesting that SR is the major contributor to  
235 total ER in our study area.

236

### 237 4.3. Temperature-respiration relationship and Q10 values

238 In our study, ST was the best predictor of SR and ER, consistent with other studies (Fang et  
239 al., 2017; Wan et al., 2007). We observed an exponential relationship between respiration  
240 rates and ST, and a linear relationship between SWC and respiration rates, similar to other  
241 studies (Rey et al., 2011; Thomas, 2012; Escolar et al., 2015). Our study showed a positive  
242 relationship between ST and respiration rates, indicating that respiration increases with the  
243 increase in ST (Fang et al., 2017) in semi-natural grassland. The temperature-respiration



244 relationship became stronger under passive experimental warming, indicating more CO<sub>2</sub>  
245 emissions in the future. This was also supported by the increase in the temperature sensitivity  
246 of respiration (Q<sub>10</sub>) under warming in our study (Escobar et al., 2015). An increase in Q<sub>10</sub>  
247 values was more in the NT than DT, indicating that NT respiration rates are more sensitive to  
248 climate warming than DT.

249

## 250 5. Conclusion

251 In conclusion, passive experimental warming resulted in significant increase in air and soil  
252 temperatures. This study indicates that passive warming is likely to enhance the respiration  
253 rates in sub-tropical grasslands. Furthermore, NT respiration rates are more sensitive to  
254 warming than DT as indicated by increase in Q<sub>10</sub> values in our study. This addresses a  
255 crucial gap in monitoring NT respiration responses along with DT to estimate the CO<sub>2</sub> efflux  
256 and its impact on future climate warming in similar ecosystem.

257

## 258 6. Future works

259 This novel study has underlined the importance of both daytime and night-time respiration in  
260 understanding the respiratory dynamics of an ecosystem. To strengthen the statistical  
261 robustness, future research endeavours should incorporate multiple replicates throughout  
262 various seasons, thereby contributing to a more comprehensive understanding. The  
263 investigation of microclimate impacts remains of utmost importance. Furthermore, an  
264 examination of the biotic regulatory factors, including the impact of plant species  
265 composition, microbial populations, and nutrient availability on the dynamics of respiration,  
266 would contribute to a more thorough understanding of carbon cycling in these ecosystems.

267

### 268 Data availability

269 Data available upon request from the corresponding author.

270

### 271 Author contributions

272 **1. Fund acquisition:** GSR & GT (funds for the equipment and conducting the  
273 experiments)





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